

SEABROOK UPDATED FSAR

APPENDIX 2C

GEOLOGIC INVESTIGATIONS OF THE SCOTLAND ROAD FAULT
(CLINTON - NEWBURY FAULT), NEWBURY, MASSACHUSETTS, AND
PORTSMOUTH FAULT INVESTIGATIONS

The information contained in this appendix was not revised, but has been extracted from the original FSAR and is provided for historical information.

SCOTLAND ROAD FAULT INVESTIGATIONS

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PUBLIC SERVICE COMPANY OF NEW HAMPSHIRE

SEABROOK STATION

SCOTLAND ROAD FAULT INVESTIGATIONS

Investigations have been conducted over a portion of the Scotland Road fault in Newbury, Massachusetts, to determine the presence, location, orientation and physical characteristics of the fault, and to examine the nature and structure of the unconsolidated Pleistocene deposits which overlie the fault trace. The investigations have indicated that the fault structure is of Permian age, and that Pleistocene deposits overlying the fault zone show no evidence of movement on the fault subsequent to their deposition.

I. LOCATION OF FAULT INVESTIGATIONS

The Scotland Road fault was inferred by A. F. Shride of the U. S. Geological Survey (Shride; 1971) to trend easterly through the towns of West Newbury, Newbury and Newburyport, Massachusetts, about 7 miles to the south of the proposed Seabrook Station (see Figure 1). Shride has interpreted the Scotland Road fault to represent the eastern portion of the Clinton-Newbury fault, which is inferred to trend northeasterly for about 60 miles from the area of Worcester, Massachusetts, to project offshore at Plum Island, Newbury.

Detailed investigations to locate and examine the fault and its overlying Pleistocene deposits have been carried out just to the north of Scotland Road near the north corner of Newbury, Massachusetts, in an open field owned by the Marion H. Marshall Estate (see Figure 2). In this area, the fault forms the boundary between Newburyport granodiorite of presumed Devonian age on the north, and an unnamed complex of diorite and schist of unknown geologic age on the south. Diabase dikes of probable Triassic age intrude both the Newburyport and the unnamed diorite/schist on both sides of the fault.

II. INVESTIGATION PROCEDURES

A. Preliminary - General Area

As preliminary investigation of the Scotland Road fault zone, J. R. Rand walked portions of the fault trace, and inquired of A. F. Shride by telephone as to his studies of the fault zone in the area. R. J. Holt of Weston Geophysical Research, Inc., and J. R. Rand together viewed the inferred trace of the fault zone between Plum Island and Groveland, Massachusetts, by helicopter flying at various altitudes. Diorite ridges aligned parallel to, and about 1000 feet to the south of the trace of the Scotland Road fault in West Newbury are readily seen from the air, but no anomalous physiographic features were noted along the trace of the fault itself. Backhoe trenching investigations over the inferred trace of the fault were attempted on the farm of Miss Alice Elwell, adjacent to Holman Lane,

West Newbury. This exploration, ultimately involving a 232-foot trench excavation in boulder till, sand-cobble till and clay till, was terminated because these glacial materials did not appear suitable for demonstrating the presence or absence of tectonic fault deformation.

B. Final - Property of Marion H. Marshall Estate

As geographic control for all investigations at the final study area on property owned by the Marion H. Marshall Estate in Newbury, a stadia survey of the area and a base map showing all pertinent features were provided by McKenna Associates, Engineers, Portsmouth, New Hampshire (see Plate 1). Technical investigations in the study area have included a seismic refraction survey; the excavation of four backhoe trenches; and the drilling of nine core borings. Laboratory investigations conducted on drill core samples from the study area have included petrographic examinations and radiometric age dating.

III. TECHNICAL INVESTIGATIONS

A. Seismic Refraction Survey

A seismic refraction survey was conducted across the study area during the period November 5-19, 1973, by Weston Geophysical Engineers, Inc., Weston, Massachusetts, to determine thicknesses of unconsolidated overburden and weathered rock materials, as well as velocities of the various geologic materials in the study area. Technical details of this survey are presented in a report by Weston Geophysical Engineers, Inc., attached herewith.

This seismic survey report concludes:

"The bedrock surface, as interpreted from seismic data, does not have any sharp breaks indicating faulting. The seismic velocities of the bedrock do not change sufficiently along the 1000-foot line of investigation to indicate the presence of any significant bedrock anomaly. The fault zone does not exhibit significant velocity differences from adjacent bedrock."

B. Borings Investigations

During the period December 4, 1973, to February 13, 1974, nine borings were put down along the centerline of the seismic refraction survey (Seismic Line "A") to locate, define and sample the Scotland Road fault zone (see Plates 2 and 3). These borings, designated SRF-1 through SRF-9, were drilled by American Drilling and Boring Co., Inc., East Providence, Rhode Island, under the supervision of Geotechnical Engineers, Inc., Winchester, Massachusetts. Geotechnical Engineers' personnel logged the unconsolidated soils materials in these borings, and J. R. Rand logged the bedrock cores. Detailed logs of these borings are attached herewith.

1. Soils

The unconsolidated soils materials encountered in 7 of the study area borings include a blanket of silty clay ranging to 40 feet in thickness,

overlying sandy deposits of varying grain sizes which range to 55 feet in thickness. Locally, a basal section of boulders of a few feet in thickness underlies the sand deposits immediately upon the bedrock surface. Soils materials were not sampled in the two angle borings, SRF-5 and SRF-7.

Plates 3 and 4 describe J. R. Rand's interpretation of the stratigraphy of the soils materials along the line of borings. The geologic interpretation is that of a blanket of glacial-marine clay of late Pleistocene age overlying glacial outwash and marine sands, all underlain by a smooth bedrock surface on which were deposited discontinuous thin sheets of glacial till or ground moraine. The sands in borings SRF-1 and SRF-4, on the southeastern end of the line of borings, are largely yellow-brown, medium- to coarse-grained, and resemble glacial outwash. The sands in SRF-6, SRF-9, SRF-2, SRF-8 and SRF-3 are commonly finer-grained and gray in color, and contain occasional thin interbeds of gray clay. These sands underlying the northern part of the line of borings are interpreted as having been derived from erosion of the outwash, with redeposition in the near-shore marine environment prior to, but historically essentially contemporaneously with deposition of the marine clays. The boundary between the two types of sandy deposits is in the area of SRF-6, where the elevation of the top of the sandy material is low, and the overlying clay blanket is thick.

2. Bedrock

The bedrock in the study area has been defined by outcrops of Newburyport granodiorite at the north end of Seismic Line "S", and by the nine borings which extend intermittently from the outcrop area on the north to Scotland Road on the south. The Newburyport outcrops at the north end of the line consist of massive, mottled pink and green, medium-grained granodiorite which exhibits saussurite alteration of feldspars and chloritization of biotite. The rock does not show evidence of shearing on the outcrop surfaces.

Proceeding southeasterly along the line of borings, the bedrock is seen in cores from SRF-5, SRF-7 and SRF-3 to become progressively more altered chemically and more deformed mechanically, becoming light tannish-green in color, and medium-fine grained and foliated in texture and fabric. With continued distance to the southeast, the bedrock in the hangingwall of the fault is seen in SRF-7, -8, -2, -9 and -6 to be an intensely deformed, light yellow-green welded breccia or cataclastic rock. All of the rock in the fault zone is compact and well consolidated, and no zones of clay gouge or other unconsolidated crushed or sheared materials were encountered in borings in the study area.

Borings SRF-7, SRF-8 and SRF-9 all progressed through the intensely deformed portion of the Scotland Road fault zone into unaltered, dark gray diorite and schist of the unnamed complex which lies to the south of the fault.

In each of these borings, a thin (1" to 2"), tan aphanitic rock layer was cored about 5 feet stratigraphically above the horizon where alteration and cataclastic deformation ceased, and this thin marker has been termed "mylonite" on Plates 3 and 4. Borings SRF-4 and SRF-2 drilled only unaltered bedrock of the diorite/schist complex.

Core in borings SRF-2, -3, -7, -8 and -9 was taken with an orienting barrel. Orientation measurements made by Geotechnical Engineers consistently show schistosity or foliation fabric of cores of the fault zone in these borings to dip in the range 35° to 60° toward the north or $N10^{\circ}W$. On Plate 2, the subcrop of the footwall of the fault is interpreted to strike $N80^{\circ}E$ and to dip to the north at an average of about 44° . The trace of the footwall lies within only about 150 feet of the location inferred by A. F. Shride from his regional mapping studies. The true thickness of the rock section subject to mechanical deformation in the fault zone approaches 300 feet, indicating that the Scotland Road fault is a regional tectonic feature of major geologic significance.

C. Trenching Investigations

At various times during the period November 20, 1973, to March 4, 1974, four backhoe trenches were excavated in the study area to expose and examine the glacial-marine clay which overlies the Scotland Road fault zone (see Plate 2). In all trenches, the organic topsoil zone was about 6 inches to 8 inches thick overlying weathered clay, and was continuous and lay parallel with the nearly planar surface of the study area field.

1. Trench 1

Trench 1, near the north edge of the fault zone, was excavated on November 20, 1973, in massive olive-gray clay to a depth of about 12 feet at the north end of the trench, and was carried for about 150 feet toward the southeast with a depth of 4 feet to 5 feet. A 2-inch to 3-inch layer of fine laminated silty sand occurred in the clay at a depth of 3 feet to $3\frac{1}{2}$ feet below ground surface, sloping gradually to the south. This laminated sand-silt layer was continuous and not disrupted in the southern 100 feet of the trench. At the northern end of the trench, the sand-silt layer merged upward into the weathered portion of the soil zone and became unidentifiable.

2. Trench 2

Trench 2, to the south of the fault trace, was excavated on December 12, 1973, to a depth of 7 feet to 8 feet in clay, and was carried northwesterly for about 50 feet until collapse of the trench walls terminated the work. This trench exposed a thin, flat-lying laminated sand-silt layer in the clay at a depth of about 6 feet. This sand-silt layer generally resembled that found in Trench 1, although the layer was saturated in Trench 2, and small springs issued from it locally when cut by the backhoe bucket.

3. Trench 3

Trench 3 was excavated across the fault zone from south to north on February 26-27, 1974, for a total length of 435 feet and to an average depth of about 7 feet. The trench was cut in olive-gray clay which was internally

massive, but which had a thick-bedded characteristic which permitted measuring the gentle undulating layering structure in the clay. Strike-and-dip plots of these layering features are shown in plan on Plate 2, and the projected layering of the clay is shown schematically in profile on Plate 4.

In addition to gross layering structure seen in the clay throughout the length of the trench, a 2-inch to 4-inch laminated fine sand and silt layer was identified within the clay overlying the footwall trace of the underlying fault zone. This sand-silt marker layer dipped northerly out of the weathered soil zone at about 100 feet north of the south end of the trench, and sloped northerly into a synclinal sag at 135 feet north of the south end of the trench, to rise back into the weathered soil zone and be lost about 170 feet north of the south end of the trench.

The structure of layering in the clay throughout Trench 2 forms gently undulating, open folds which appear generally to parallel the upper surface of the underlying outwash and marine sand deposits. No tight or abrupt folds were seen to disrupt the continuity of layering in the clay, and close examination throughout the length of the trench failed to detect any drag folding within the clay beds. The clay is jointed throughout the trench area, with joints tending to change orientations to conform to changing attitudes of the broad undulations in clay layering. No slickensides or other evidence of displacement were detected on any joints in the trench. No

sand dikes cutting across clay layering or filling joints were found.

No offsets were found in the thin, sagged sand-silt marker horizon which was interbedded in the clay between Stations 100 and 170 in Trench 3.

Between 55 feet to 65 feet north of the south end of Trench 3, the backhoe excavated a pocket into the floor of the trench to a depth of about 14 feet, to determine whether there were any stratigraphic changes to that depth which might be useful to examine while proceeding northerly with the excavation across the fault zone. To the 14-foot depth tested, no sand layers were seen in the clay, and the pocket was backfilled to restore the trench floor to the normal 7-foot depth. Within a few moments of completing and tamping the backfill, several springs erupted from the trench floor within the backfill area, with artesian flows rising 1 inch to 2 inches above the floor of the trench. Fine gray sand suspended in the flowing waters of the several springs rapidly built sand cones several inches thick around the springs. A dam was built across the trench to the north of the springs, to protect the proposed excavation to the north from flooding, and thereafter the southern 80 feet of the trench filled to within 2 feet of ground surface, with the highly mobile fine gray sand continuing to be deposited from the springs onto the floor of the flooding trench.

4. Trench 4

Trench 4 was excavated on March 4, 1974, in an attempt to locate the westerly projection of the laminated sand-silt marker horizon found between

Stations 100 and 170 in Trench 3. A similar layer was found in Trench 4, taking the form of an open synclinal sag which plunged gently to the north-east toward Trench 3. Spoon sampling of the soils in Boring SRF-6, between the two trenches, also had detected a sand-silt layer in the clay at an elevation corresponding with that which projected between the two trenches.

Various points on the sand-silt horizon in each of the two trenches were then surveyed in by McKenna Associates in order to provide locations and elevations with which to define the structure of the horizon as it passed over the footwall and portions of the intensely deformed base of the Scotland Road fault. These surveyed points are designated points "A" through "J" on Plate 2. The structure of the horizon is defined in plan in an insert on Plate 2, and in profiles showing the east wall of Trench 3 and the east and west walls of Trench 4 on Plate 4.

As shown on Plate 2, the structure of the sand-silt marker horizon takes the form of an open, doubly-plunging syncline which strikes south-westerly across the footwall of the fault. No offsets of the sand-silt layer were detected in either trench, and no abrupt folding or drag folds were detected in this layer or in the clay beds in either trench. The sand-silt layer in both trenches does not apparently thicken or show increased grain sizes toward the trough of the syncline. No sand dikes were found in Trench 4, nor were joints slickensided.

No evidence was found to suggest that the synclinal structure of the sand-silt layer crossing the fault in the area of Trenches 3 and 4 was formed by other than passive deformation due to differential settlement of the underlying clay. The relatively non-compressible outwash and marine sands underlying the clay in the study area are at a low elevation beneath the area of this synclinal sag, and the relatively compressible clay section is thick. Conversely, the sand elevation is high and the clay is thin as seen in borings put down to the north and south of the sag. With the gradual post-depositional compaction of the clay materials through time, the thicker clay sections settled more deeply than the thin clay sections, passively producing sags in the originally horizontal layering of the fine-grained clay deposits.

There is no detectable sag in the topsoil zone which overlies the synclinal sag in the sand-silt marker horizon in Trenches 3 and 4, and there is no noticeable variation in thickness of the topsoil zone in these trenches. Since the sand-silt layer does not thicken or show coarser grain sizes toward the trough of the synclinal sag, the sand-silt layer appears to have been deposited on an originally horizontal surface which lay stratigraphically above the present ground surface. Differential settlement and sagging of the sand-silt horizon must have been completed prior to the last erosional beveling of the present ground surface, presumably upon retreat of the last post-glacial marine transgression, since the topsoil zone built

upon this beveled horizon shows no evidence of having sagged over the sand-silt sag or over any other of the gently undulations seen in the clay layering throughout the length of Trench 3. There is no evidence of disruption of any of the sedimentary layers overlying the fault zone in any of the trenches, to suggest movement on the Scotland Road fault subsequent to deposition of the overlying Pleistocene deposits.

D. Age of Pleistocene Deposits

No shells or other organic materials were found in the clay in the study area with which to establish an age of deposition of the clay. The clay deposit is, however, considered correlative with similar glacial-marine clays which blanket portions of the seaboard lowland throughout eastern New England.

Borns (1973) reports that "a major amelioration of climate began prior to 14,200 years ago which resulted in a rapid dissipation of the ice sheet in New England at least by 12,500 years ago". The recession of the ice sheet was accompanied by a marine invasion of the seaboard lowland, with deposition of glacial-marine clay sediments. Borns brackets the time of deposition of the glacial-marine clay in the region between 13,500 and 12,500 years ago.

Schafer and Hartshorn (1965) report that radiocarbon dates of shells from glacial-marine sediments on the seaboard lowland in Maine range from 11,800 to 12,800 years old. Kaye and Barghoorn (1964) have constructed

a curve of sea-level fluctuations for the Boston, Massachusetts, area which describes the last marine submergence as having ended about 12,500 years ago in that area.

It appears, therefore, that the glacial-marine clays of the Newbury study area are at least older than 11,800 years, and are probably in the range of 12,500 to 13,500 years old.

E. Petrographic Examinations

The petrography of eight samples of drill core from borings in and adjacent to the Scotland Road fault has been described by Professor Gene Simmons and Dorothy Richter of Massachusetts Institute of Technology.

Sample	Boring	Depth (feet)	Description
SRF-1A	SRF-1	74.0 to 74.4	Amphibolite breccia
SRF-2A	SRF-2	60.0 to 60.4	Mylonized quartz-muscovite schist
SRF-2B	SRF-2	72.9 to 73.4	Brecciated quartz-muscovite schist
SRF-3A	SRF-3	67.0 to 67.5	Muscovite mylonite
SRF-4A	SRF-4	92.9 to 93.3	Chlorite augen gneiss
SRF-5A	SRF-5	42.1 to 42.6	Sheared granodiorite
SRF-5B	SRF-5	175.1 to 175.6	Altered olivine basalt
SRF-7A	SRF-7	115.9 to 116.4	Ultramylonite

Simmons and Richter conclude from their studies that "the samples (with the exception of sample SRF-5B) all show evidence of dynamic deformation; that is, cataclasis, brecciation and intense crushing--all probably due to motion along the fault. The deformation clearly took place after the regional metamorphism of the rocks (which was probably associated with the Devonian Acadian orogeny). The microcracks produced in the deformational

events appear in thin section to have either annealed, or have been filled by secondary minerals. There is no firm petrographic evidence of recent deformation of these samples". The complete text of the Simmons and Richter report is attached herewith.

A further indication of the old age of deformation of the fault zone is evidenced by sample SRF-5B, from a diabase dike which is enclosed within deformed rocks of the fault zone. Petrographically the dike is seen to be completely undeformed. The dike has been dated radiometrically (K-Ar) at 199 ± 9 million years.

F. Radiometric Age Dating

K-Ar age determination have been obtained on six samples of drill core from borings in and adjacent to the Scotland Road fault by Geochron Laboratories, Division of Krueger Enterprises, Inc., Cambridge, Massachusetts.

Sample	Boring	Depth (feet)	Material	Age
SRF-5A	SRF-5	42.1 to 42.6	whole rock	272 10 M.Y.
SRF-3A	SRF-3	67.0 to 67.5	whole rock	269 10 M.Y.
SRF-2A	SRF-2	60.0 to 60.4	whole rock	256 10 M.Y.
SRF-8A	SRF-8	155.6 to 156.0	sericite/ feldspar	248 9 M.Y.
SRF-1A	SRF-1	74.0 to 74.4	amphibole	324 14 M.Y.
SRF-5B	SRF-5	175.1 to 175.6	whole rock	199 9 M.Y.

Samples SRF-5A, -3A, -2A, and -8A are from within the fault zone; SRF-1A is from the diorite/schist complex which lies to the south of the fault zone; SRF-5B is from an undeformed diabase dike which is enclosed within deformed rocks of the fault zone (see Plate 3). Of apparent geologic

interest is the fact that radiometric ages increase progressively with distance from the footwall of the fault zone. SRF-8A is from about 5 feet above the mylonite band near the footwall of the fault, while SRF-5A is in relatively undeformed granodiorite about 250 feet stratigraphically above the footwall. Radiometric dating of rocks within the Scotland Road fault zone indicates that the fault is of Permian age, and suggests that deformation in the zone may have been active through a period of as much as 20 million years. The dike (SRF-5B) which intruded the fault zone is completely undeformed, indicating that movement on the fault had ceased by Triassic time.

IV. CONCLUSIONS

The Scotland Road fault has been located within 150 feet of the location inferred by A. F. Shride on the basis of his regional field studies. Nine core borings have defined the fault zone as being about 300 feet thick and dipping at about 44° to the north adjacent to Scotland Road in Newbury, Massachusetts. Chemical alteration and mechanical deformation in the fault zone increases progressively from north to south across the fault zone, and alteration effects of faulting terminate abruptly at the footwall of the fault zone, about 5 feet stratigraphically below a thin mylonite band. The fault is a feature of major geological significance in the region.

The fault is geologically very old, of early to middle Permian age, and the altered and deformed bedrock materials in the fault zone are annealed and compact. No unconsolidated gouge, shear zones or polished joint surfaces

were detected in cores from borings drilled across the width of the fault zone. The bedrock surface overlying the fault zone slopes gradually up to undeformed bedrock outcrops at the north edge of the fault zone, and appears from refraction seismic surveys and borings data to be smooth and sub-planar, with no detectable topographic anomalies.

Surficial materials overlying the fault zone include glacial till, glacial outwash and marine sands, and glacial-marine clays, all of Pleistocene age. The youngest of these Pleistocene deposits are the glacial-marine clays, estimated from regional studies to be older than 11,800 years. A thin, essentially horizontal layer of post-Pleistocene topsoil covers the glacial-marine clay in the area.

Examination of the glacial-marine clay in four trenches excavated over the area of the fault zone failed to detect any evidence of tectonic fault displacement in the clay and its interbedded sand-silt layers. Bedding in the clay displayed no abrupt monoclinal or drag folds. Joints were not slickensided. The thin laminated sand-silt horizons interbedded in the clay were not offset. No sand dikes were found in the clay, which directly overlies deposits of highly mobile fine sand..

All evidence observed in the current investigations indicate that Pleistocene deposits overlying the Scotland Road fault have not been subjected to disruption by tectonic faulting.

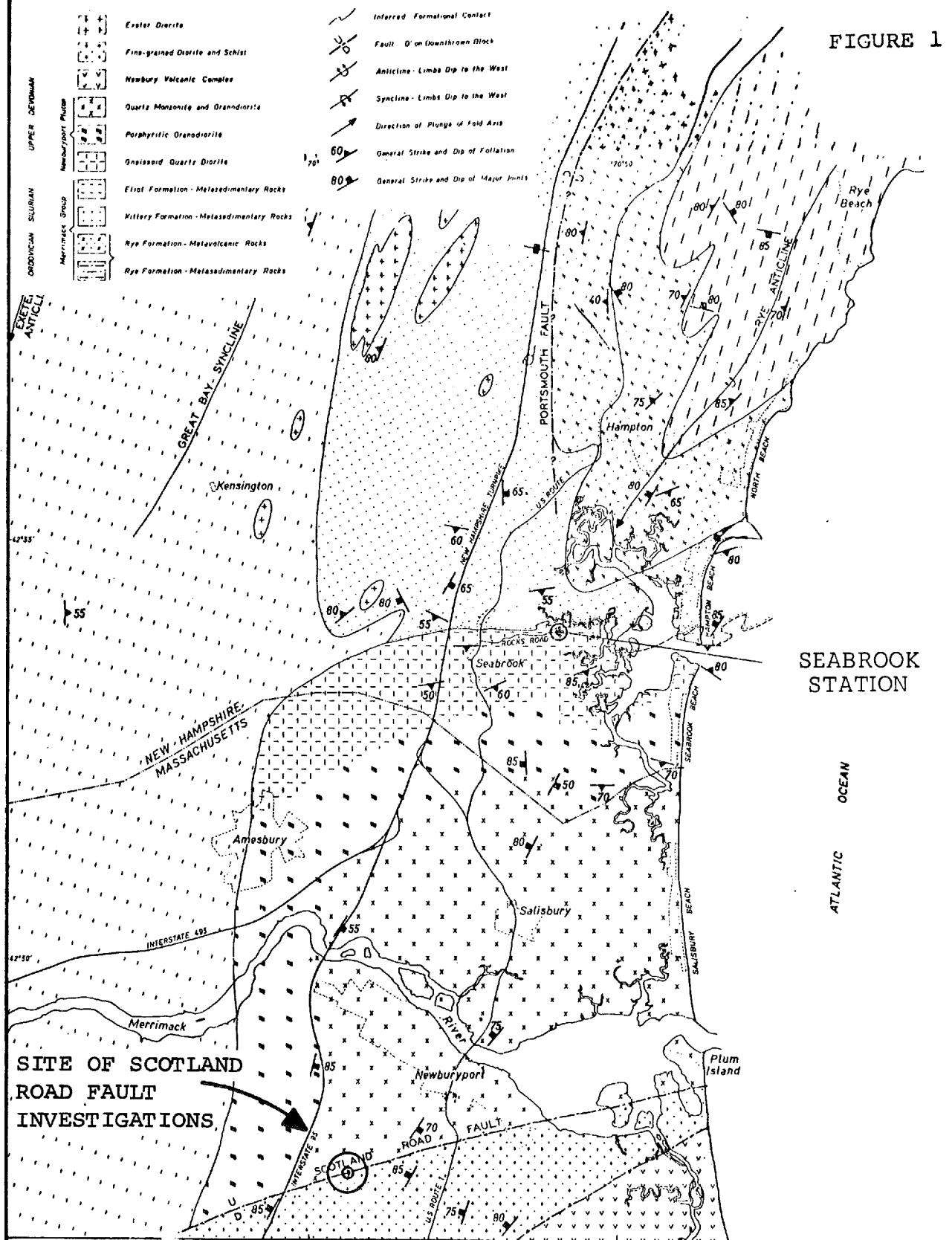
John R. Rand
Consulting Geologist

References:

- Borns, H. W., Jr. (1973) Late Wisconsin Fluctuations of the Laurentide Ice Sheet in Southern and Eastern New England. in The Wisconsinan Stage; Geological Society of America, Memoir 136; Boulder, Colorado.
- Kaye, C. A. and E. S. Barghoorn (1964) Late Quaternary Sea-Level Change and Crustal Rise at Boston, Massachusetts, with Notes on the Autocompaction of Peat. Geological Society of America, Bulletin Vol. 75, 63-80.
- Schafer, J. P. and J. H. Hartshorn (1965) The Quaternary of New England. in The Quaternary of the United States; Princeton University Press; Princeton, New Jersey.
- Shride, A. F. (1971) Igneous Rocks of the Seabrook, New Hampshire-Newbury, Massachusetts, Area. in Guidebook for Field Trips in Central New Hampshire and Contiguous Areas. New England Inter-collegiate Geological Conference - 1971.

NOTE: The study area was visited on March 13, 1974, by M. H. Pease, Jr. and P. J. Barosh, U. S. Geological Survey, Boston. Trenches 3 and 4 were inspected. The trenches were thereupon filled in.

FIGURE 1



BEDROCK GEOLOGY OF THE SEABROOK REGION, NEW HAMPSHIRE - MASSACHUSETTS

0 1 2 3 4
SCALE OF MILES

LOCATION MAP
REGIONAL FAULT INVESTIGATIONS
PUBLIC SERVICE CO. OF NEW HAMPSHIRE
SEABROOK STATION
J. R. Rand, Consulting Geologist

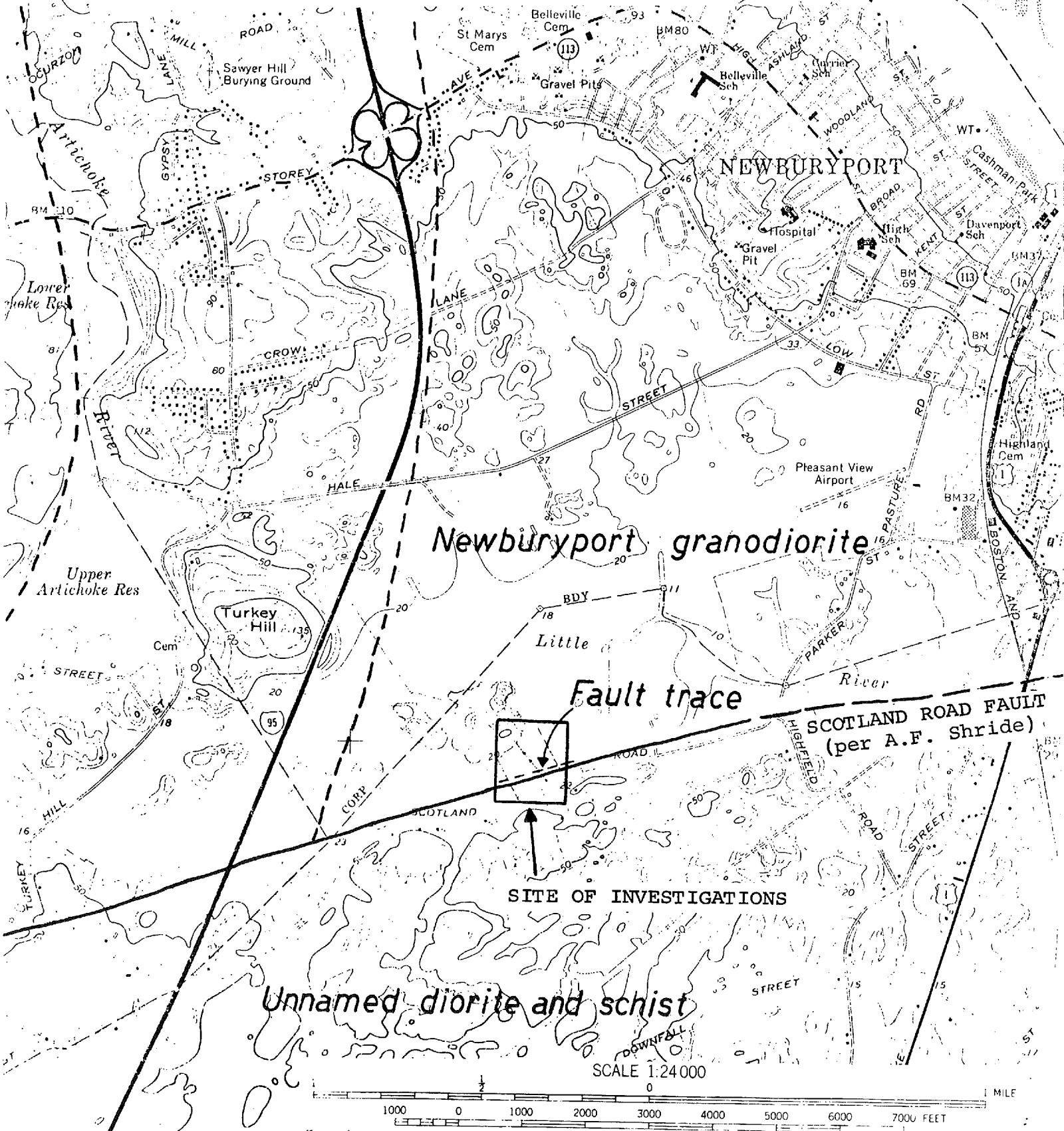
PUBLIC SERVICE CO. OF NEW HAMPSHIRE
SEABROOK STATION
LOCATION MAP - SCOTLAND ROAD FAULT INVESTIGATIONS

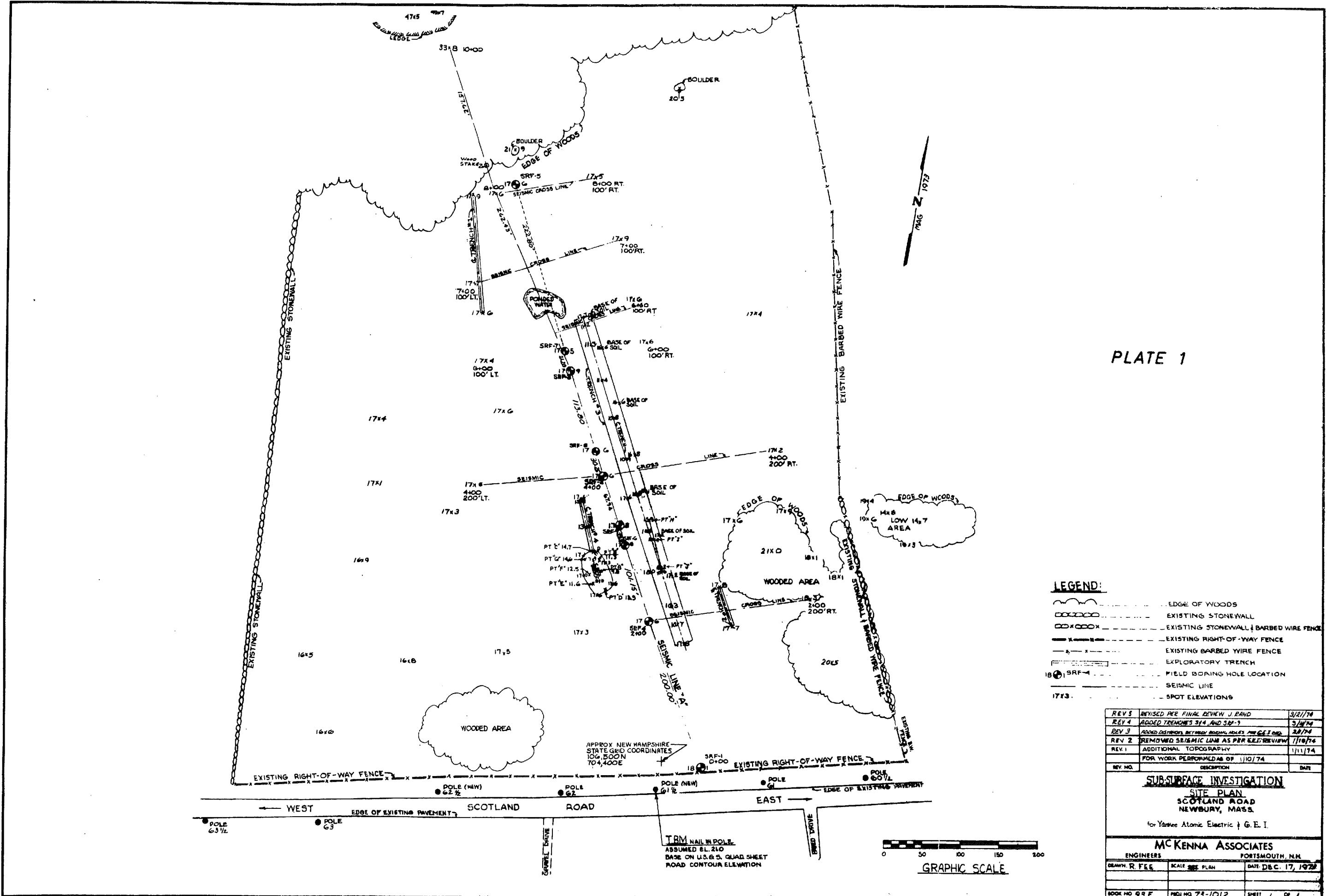
FIGURE 2

on property of

ESTATE of MARION H. MARSHALL
SCOTLAND ROAD, NEWBURY, MASSACHUSETTS

J. R. Rand Consulting Geologist





GEOLOGIC MAP - SCOTLAND ROAD FAULT

on property of

MARION H. MARSHALL ESTATE
NEWBURY, MASSACHUSETTS

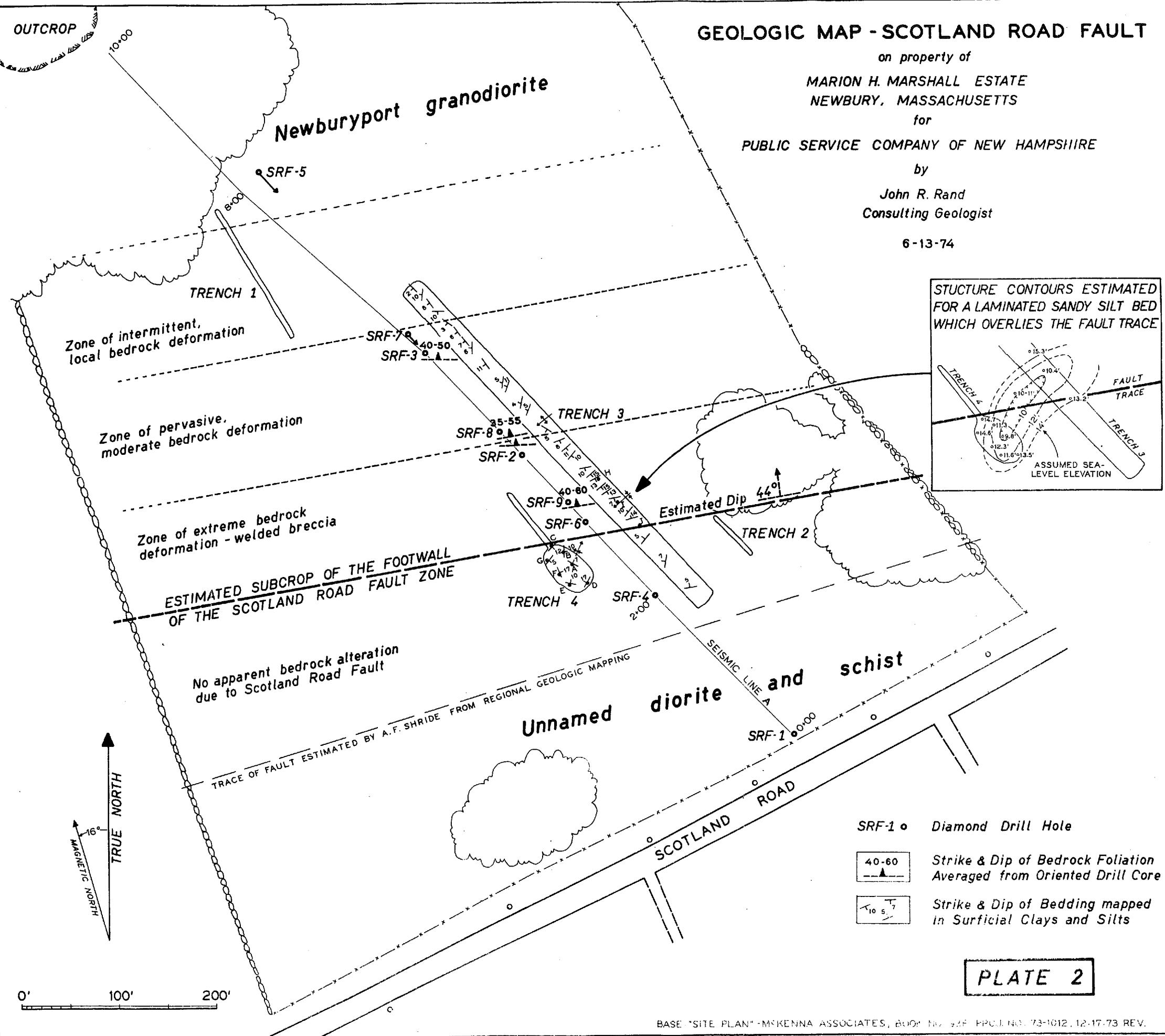
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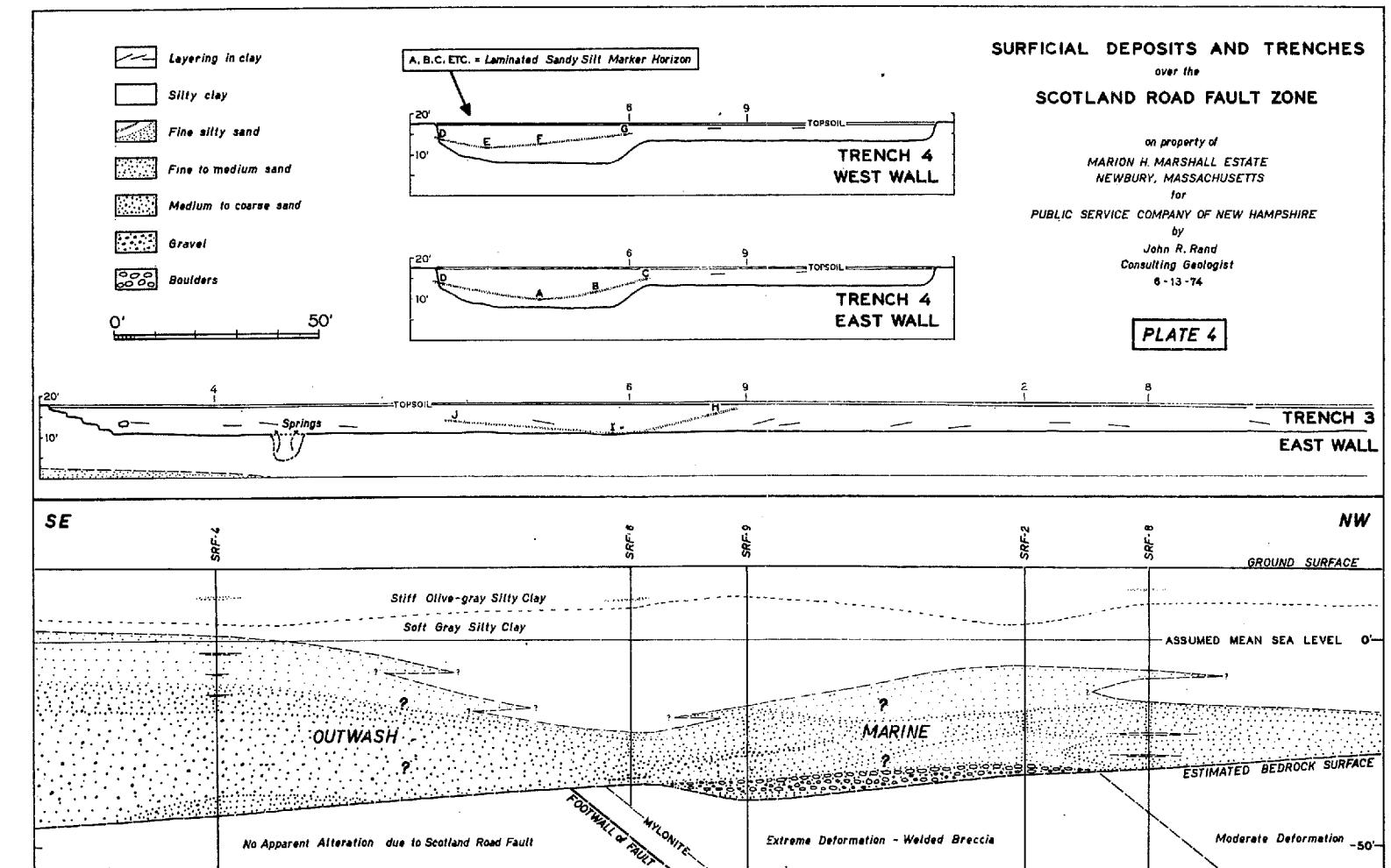
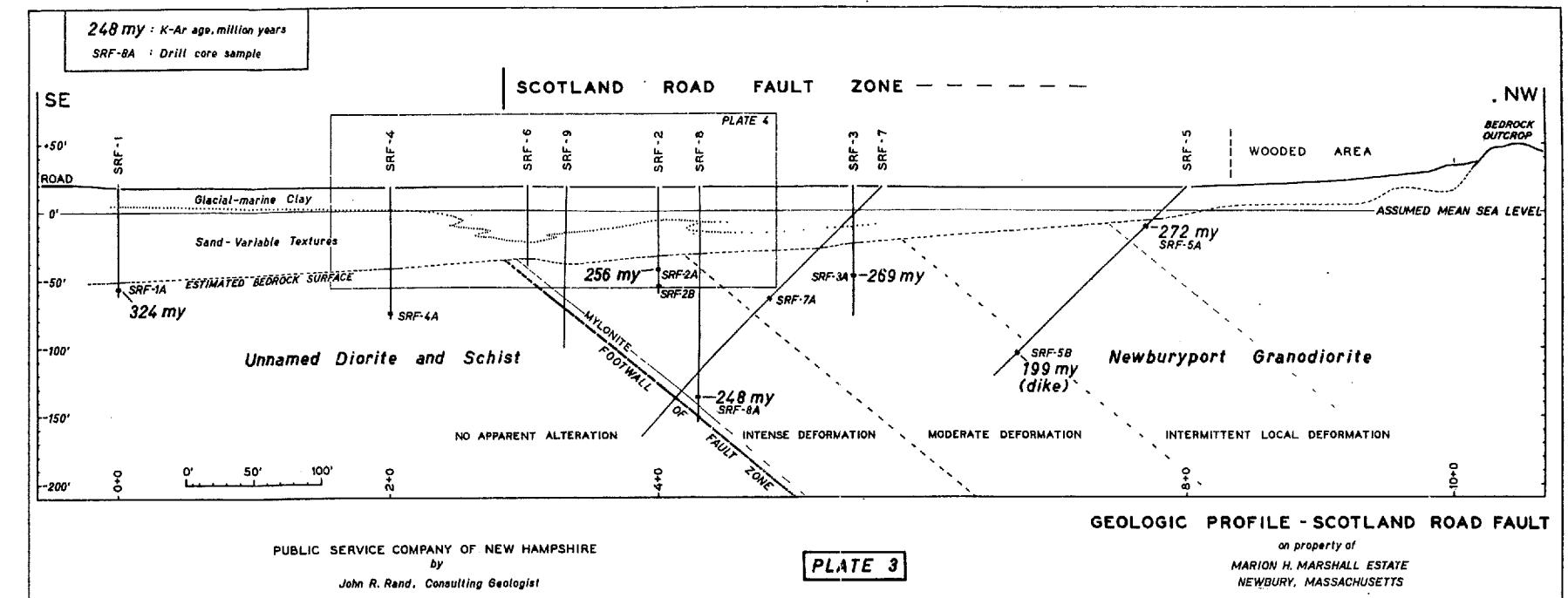
PUBLIC SERVICE COMPANY OF NEW HAMPSHIRE

by

John R. Rand
Consulting Geologist

6-13-74





ATTACHMENT No. 1

SEISMIC REFRACTION SURVEY
SCOTLAND ROAD FAULT ZONE
NEWBURY, MASSACHUSETTS

WESTON GEOPHYSICAL ENGINEERS, INC.
for
PUBLIC SERVICE COMPANY OF NEW HAMPSHIRE

SEISMIC REFRACTION SURVEY

SCOTLAND ROAD FAULT ZONE

NEWBURY, MASSACHUSETTS

for

PUBLIC SERVICE COMPANY

OF NEW HAMPSHIRE



WESTON GEOPHYSICAL ENGINEERS, INC.
WESTON, MASSACHUSETTS

SEISMIC REFRACTION SURVEY
SCOTLAND ROAD FAULT ZONE
NEWBURY, MASSACHUSETTS

INTRODUCTION

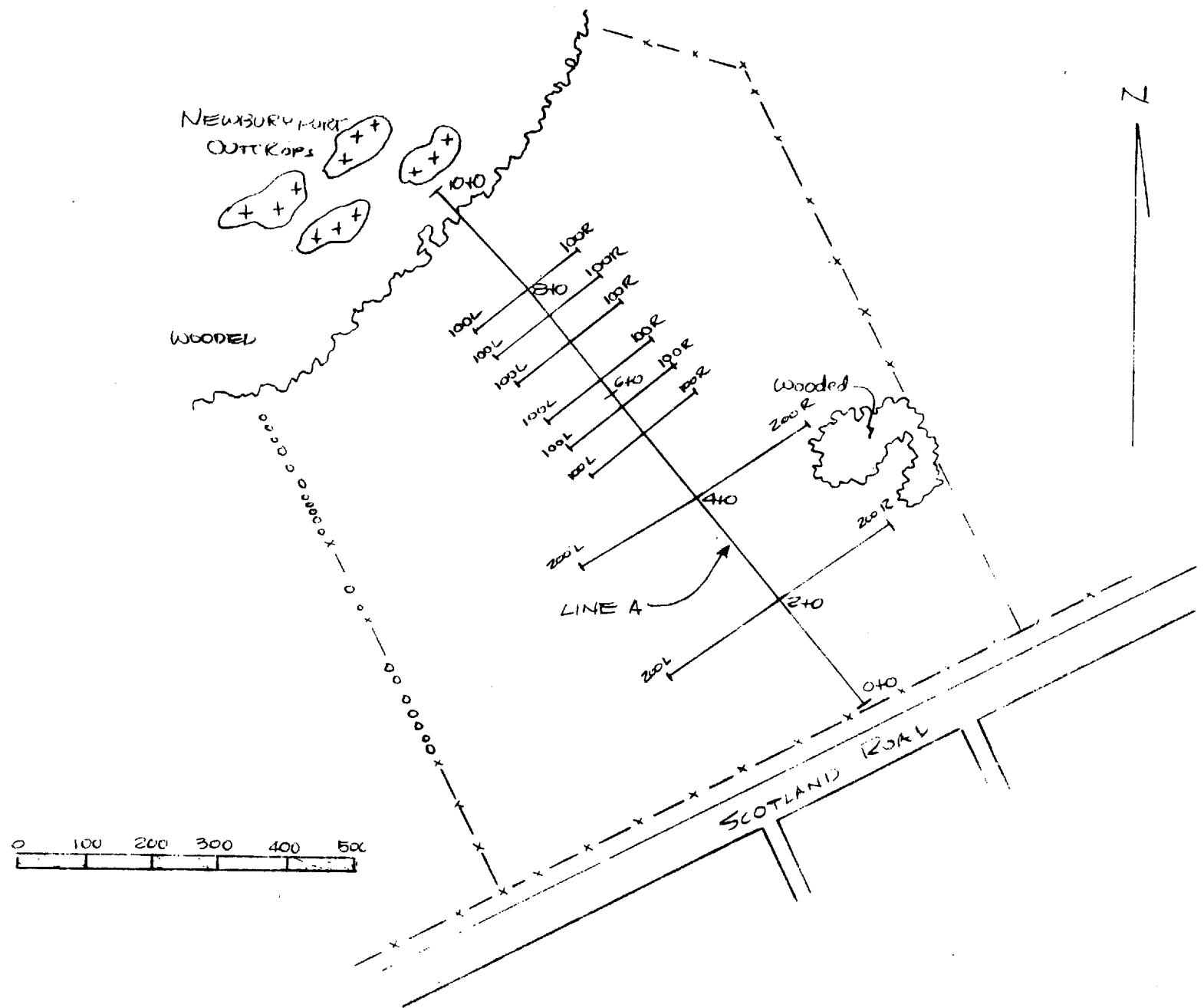
A seismic refraction survey was conducted across the mapped location of the Scotland Road fault, as originally mapped by A. F. Shride (1971) and shown on Figure 1 and Plate 2 of the report. Seismic field work took place during the period of November 5 through 19, 1973. The location of this survey is shown on Figure 1 of this attachment.

The general purpose of this work was to determine thicknesses of overburden and weathered rock materials as well as the velocities of the various geologic materials existing at this location.

RESULTS

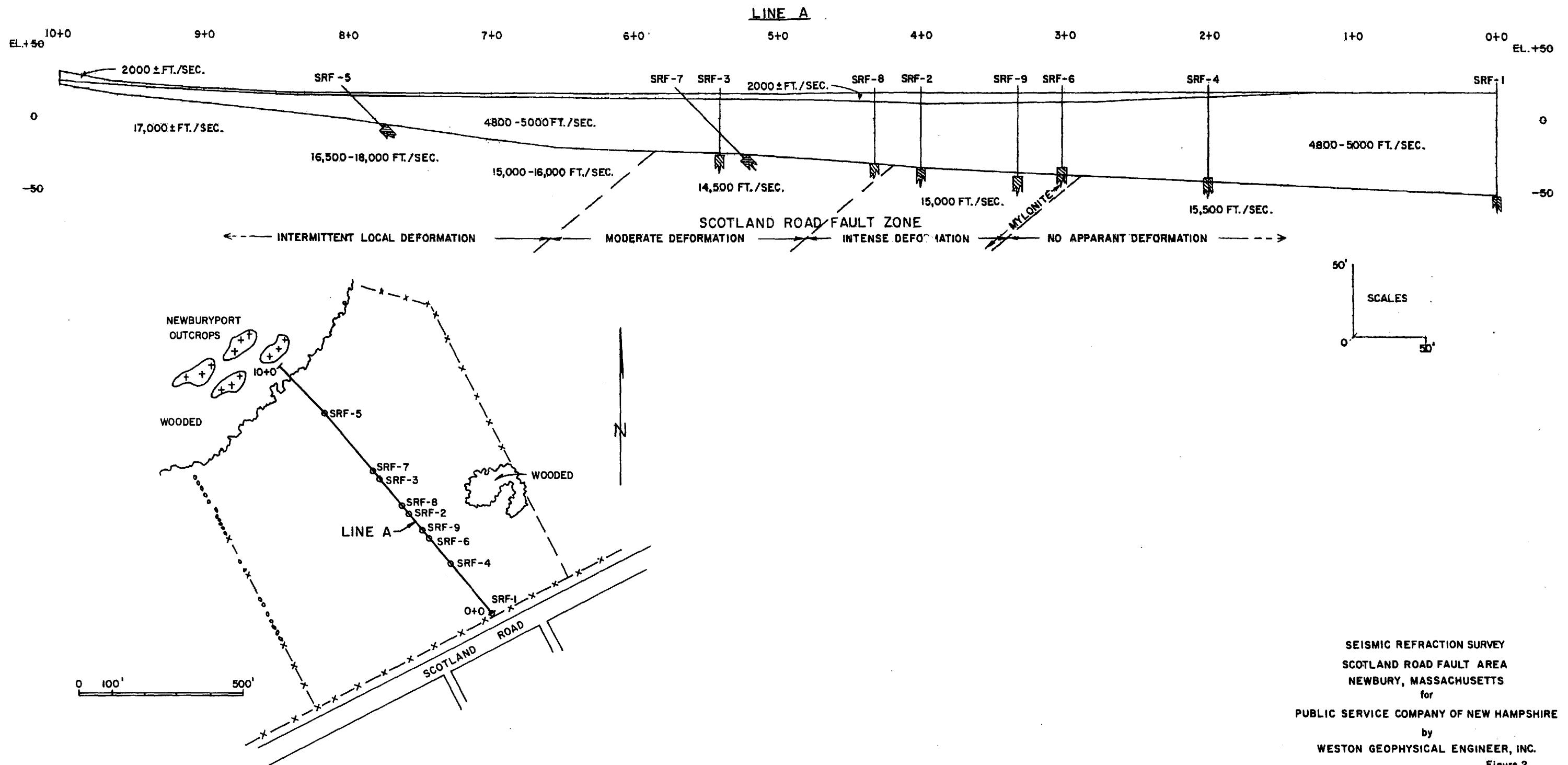
The results of this refraction survey are shown on a profile of the bedrock surface (Figure 2). Also shown on this profile are overburden and bedrock seismic velocities, boring locations, and bedrock depths as found from borings as well as the fault zone, as indicated by J. R. Rand.

The bedrock surface, as interpreted from seismic data, does not have any sharp breaks indicating faulting. The seismic velocities of the bedrock do not change sufficiently along the 1,000-foot line of investigation to indicate the presence of any significant bedrock anomaly. The fault zone does not exhibit significant velocity differences from the adjacent bedrock.



SEISMIC LINE LOCATIONS
SCOTLAND ROAD FAULT
PUBLIC SERVICE CO. of NEW HAMPSHIRE
SEABROOK NUCLEAR STATION

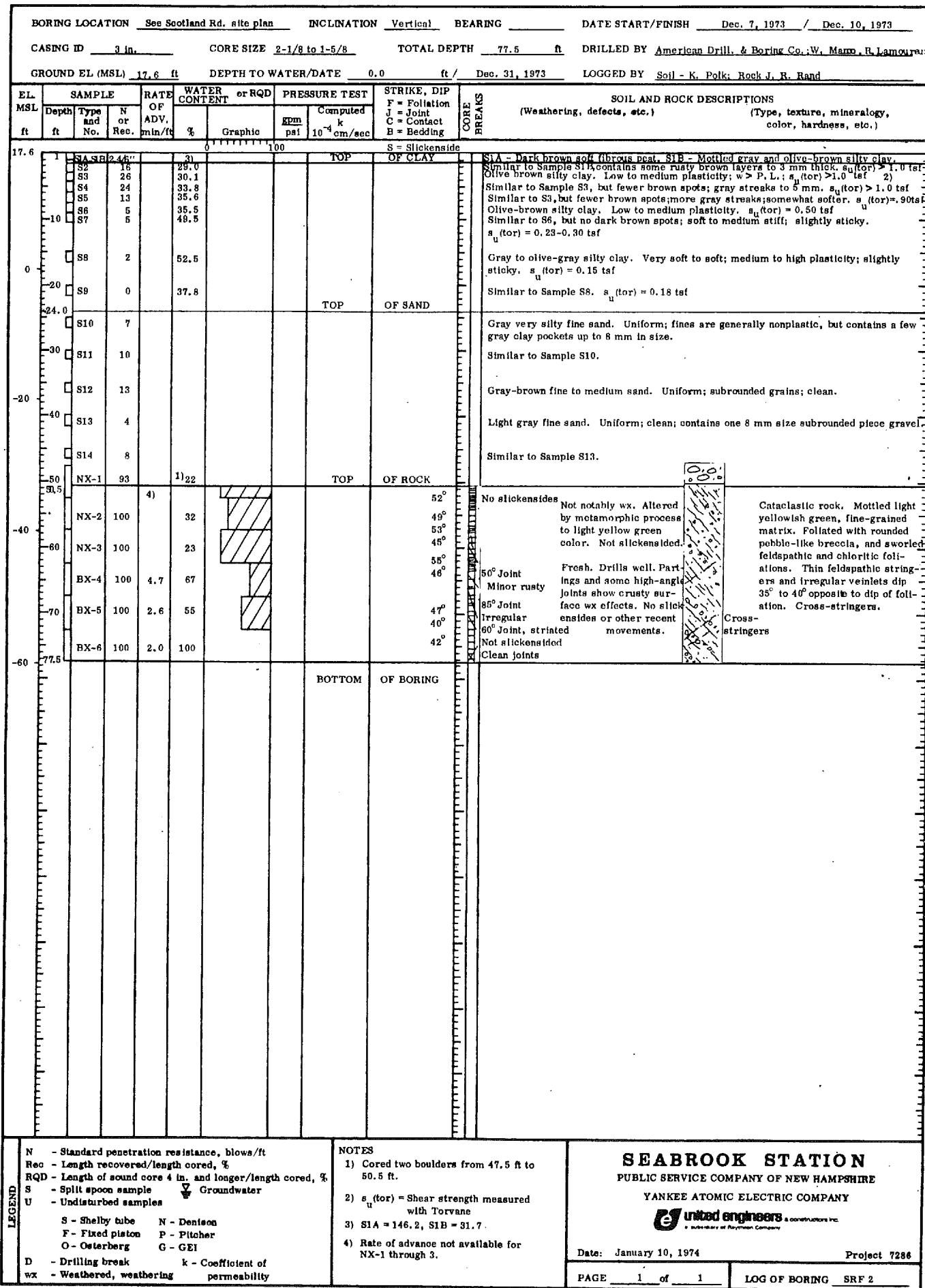
Figure 1

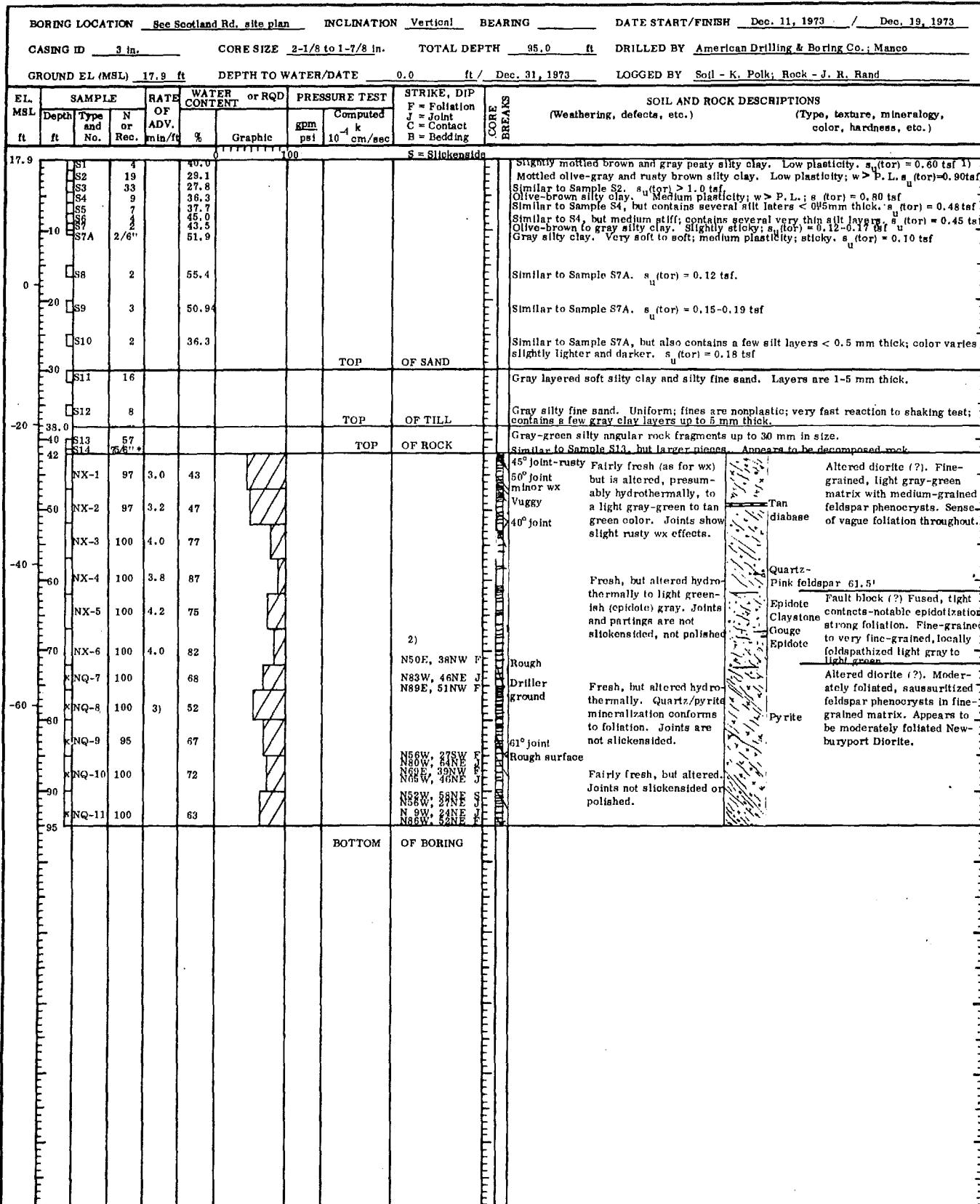


ATTACHMENT No. 2

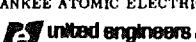
GEOLOGIC AND SOILS LOGS OF BORINGS
SRF-1 THROUGH SRF-9

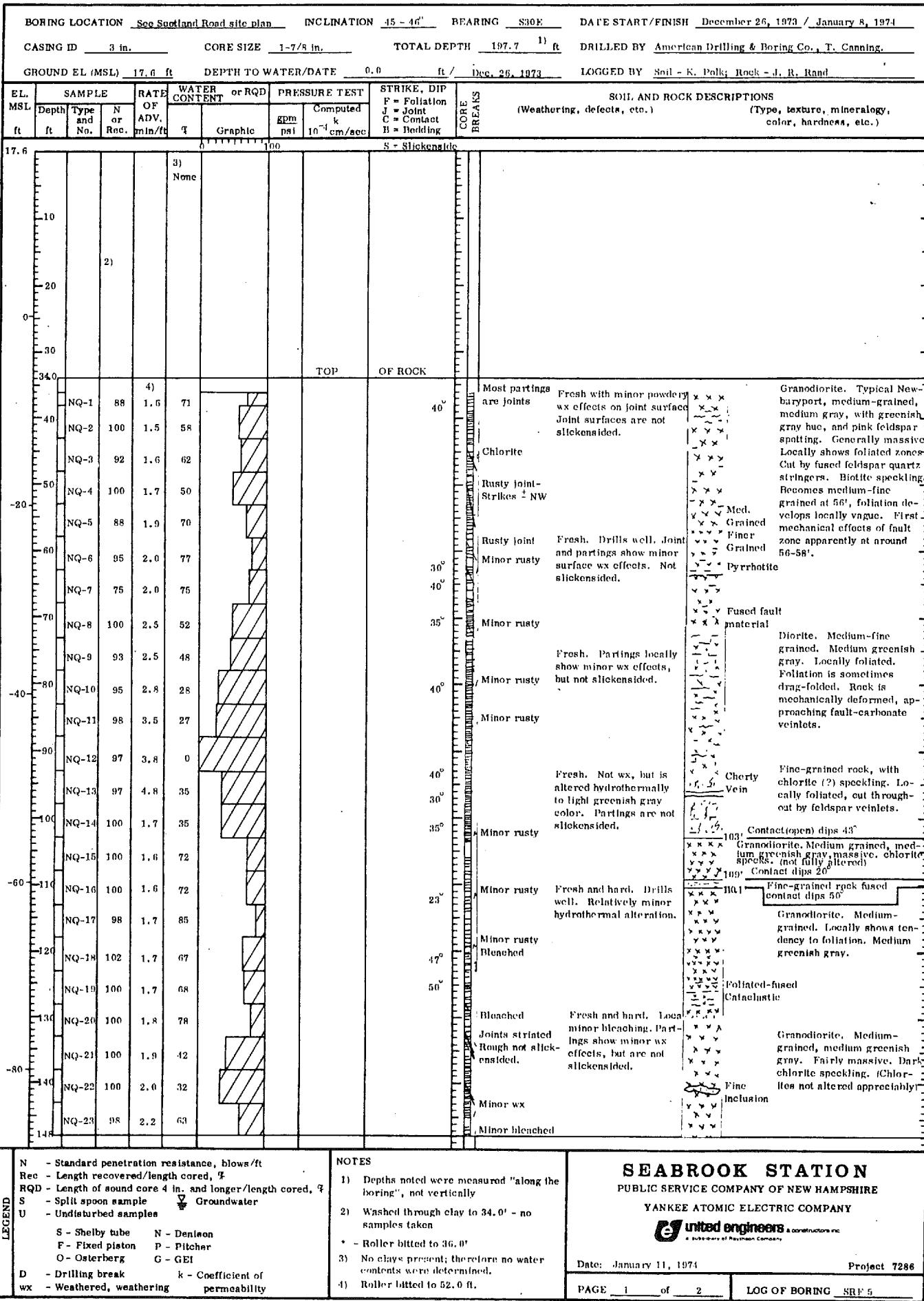
BORING LOCATION See Scotland Rd. site plan				INCLINATION Vertical	BEARING	DATE START/FINISH Dec. 4, 1973 / Dec. 6, 1973		
CASING ID 3 in.				TOTAL DEPTH 69.0 ft	DRILLED BY American Drilling & Boring Co.; W. Manco			
GROUND EL (MSL) 18.1 ft DEPTH TO WATER/DATE 0.5 ft / Dec. 28, 1973				LOGGED BY Soil - K. Polk; Rock - J. R. Rand				
EL. MSL ft	SAMPLE Depth ft	RATE OF ADV. min/ft	WATER CONTENT % Graphic	RQD	PRESSURE TEST gpm psi 10^{-4} cm/sec	STRIKE, DIP F = Foliation J = Joint C = Contact B = Bedding CORE BREAKS S = Slickenslide		
18.1	S1 S2 S3 S4 S5 S6 -10 S7	7 7 21 29 30 31 37.8	33.5 29.4 32.3 31.9 32.4 34.9	100	TOP	OF CLAY	Mottled gray, olive-gray, and brown silty clay. Low plasticity; $w > P.L.$; $s_u(\text{tor}) > 60 \text{ tsf}$. Slightly mottled gray & olive brown silty clay. Low to med. plasticity, $s_u(\text{tor}) > 1.0 \text{ tsf}$. Olive-brown silty clay. Low to medium plasticity; $w > P.L.$ $s_u(\text{tor}) > u 1.0 \text{ tsf}$. Similar to Sample S3. $s_u(\text{tor}) > 1.0 \text{ tsf}$. Similar to S3, but somewhat softer; contains few gray spots to 8 mm. $s_u(\text{tor}) = 0.95 \text{ tsf}$. Similar to S3, but softer; some gray spots. $s_u(\text{tor}) = 0.65 \text{ tsf}$. Similar to Sample S3, but medium stiff; contains a gray silt layer < 0.5 mm thick; color varies slightly olive-brown to olive-gray. $s_u(\text{tor}) = 0.34 \text{ tsf}$.	
13.0	S8 S9 S10 S11 S12 S13 S14 S15 S16A S16B S17A S17B S18 NX-1 NX-2	0 9 10 9 1/2" 19 24 26 31 17 59 15 6" 79 28 6" 92 6" 41 100 3.0			TOP	OF SAND	Gray layered silty clay and clayey fine sand. Silty clay is soft; medium to high plasticity; slightly sticky; very soft when remolded. Layers vary 0.5-10 mm. $s_u(\text{tor}) = 0.22 \text{ tsf}$. Gray silty fine sand. Uniform; fines are nonplastic; very fast reaction to shaking test. Similar to Sample S9, but also contains a few gray clay layers. Brown silty fine sand. Uniform; fines are nonplastic; contains a few rusty-brown fine sand layers. Brown slightly silty fine to medium sand. Uniform; fines are nonplastic; contains a layer of gray clayey gravelly sand with subrounded gravel up to 20 mm in size. Brown very slightly silty uniform fine to medium sand. Light brown silty fine sand. Uniform; fines are nonplastic; contains a few subrounded coarse sand grains and some rusty-brown medium sand layers. Similar to Sample S14. Similar to Sample S16A.	
-40	S16B S17A S17B S18 NX-1 NX-2	15 6" 79 28 6" 92 6" 41 100 3.0			TOP	OF TILL	Gray-brown silty sandy gravel. Widely graded; angular grains; contains gravel pieces up to 30 mm in size; fines are nonplastic. Light gray fine to medium sand. Uniform; angular to subrounded grains; clean. Light gray silty sandy gravel. Angular grains; appears to decomposed rock and rock fragments up to 30 mm in size. Gray silty gravelly fine sand. Uniform; fines are nonplastic; contains angular gravel pieces up to 15 mm in size. Cored boulders.	
-60	NX-3 NX-4	93 98	4.2 3.6	43 83	/	TOP	OF ROCK	75° Joint Clean Fresh and hard. Drills well. Only very slight surface wx effects on joints and partings. 75° Joint Minor wx Diorite. Dark gray with large (?) hornblend crystals (1/2") in fine-grained quartz diorite matrix. 76.3° Gradational contact - fused. Diorite. Massive, fine-grained, dk. gray
-80						BOTTOM	OF BORING	
LEGEND	N - Standard penetration resistance, blows/ft Reo - Length recovered/length cored, % RQD - Length of sound core 4 in. and longer/length cored, % S - Split spoon sample U - Undisturbed samples G - Shelby tube F - Fixed piston O - Osterberg D - Drilling break wx - Weathered, weathering	V - Groundwater N - Denison P - Pitcher G - GEI k - Coefficient of permeability	NOTES 1) - $s_u(\text{tor})$ = Shear strength measured with Torvane.			SEABROOK STATION PUBLIC SERVICE COMPANY OF NEW HAMPSHIRE YANKEE ATOMIC ELECTRIC COMPANY united engineers	Date: January 10, 1974	Project 7286
						PAGE 1 of 1	LOG OF BORING SRF 1	





LEGEND	N - Standard penetration resistance, blows/ft	NOTES		SEABROOK STATION	
	Rec - Length recovered/length cored, %	1) s_u (tor) = Shear strength measured with Torvane.		PUBLIC SERVICE COMPANY OF NEW HAMPSHIRE	
	RQD - Length of sound core 4 in. and longer/length cored, %	2) This is only a partial list of dip and strike data.		YANKEE ATOMIC ELECTRIC COMPANY	
	S - Split spoon sample Groundwater	3) Rate of advance not available for NQ-8 through NQ-11.		A subsidiary of Raytheon Company	
	U - Undisturbed samples			Date: February 13, 1974 Project 7286	
	S - Shelby tube N - Denison				
	F - Fixed piston P - Pitcher				
	O - Osterberg G - GEI				
D - Drilling break k - Coefficient of permeability			PAGE 1 of 1	LOG OF BORING SRF 3	
wx - Weathered, weathering	* Used 300 lb hammer.				

BORING LOCATION			See Scotland Rd. site plan		INCLINATION	Vertical	BEARING	DATE START/FINISH		Dec. 20, 1973 / Jan. 3, 1974	
CASING ID		3 in.	CORE SIZE		2-1/8 to 1-7/8 in.	TOTAL DEPTH	96.0 ft	DRILLED BY	American Drilling & Boring Co.; W. Manco		
GROUND EL (MSL)		17.6 ft	DEPTH TO WATER/DATE		0.0 ft / Jan. 2, 1974			LOGGED BY	Soil - K. Polk; Rock - J. R. Rand		
EL. ft	SAMPLE Type and No.	N or Rec.	RATE OF ADV. m/in/ft	WATER CONTENT %	or RQD Graphic	PRESSURE TEST K/cm psi	Computed k 10^{-4} cm/sec	STRIKE, DIP F = Foliation J = Joint C = Contact B = Bedding	CORE BREAKS	SOIL AND ROCK DESCRIPTIONS (Weathering, defects, etc.)	
ft										(Type, texture, mineralogy, color, hardness, etc.)	
17.6	S1-S1A	1-6		27.5						S = Slickenside	
	S2	19		28.0						SI-Dark brown peaty loam. SIA-Mottled brown, olive-brown, and gray silty clay. Med. stiff to stiff; low plasticity; contains roots > 10 mm in dia., w/ above PL. s_u = 0.75 tsf. SIB-Similar to SIA, but somewhat rockier, some layering; s_u slightly above PL. SIC-Similar to SIA, but more stiff; medium plasticity; somewhat blocky. s_u (tor) > 1.0 tsf. S4-Similar to Sample S3, with some dark brown spots up to 1 mm thick. s_u (tor) > 1.0 tsf. S5-Olive-brown silty clay. Medium stiff to stiff; medium plasticity; contains several dark brown spots. s_u (tor) = 0.52 tsf. S6-Similar to Sample S5. s_u (tor) = 0.45 tsf. S7-Similar to Sample S5. s_u (tor) = 0.35 tsf	
0	S8	13		33.0				TOP	OF SAND	Gray silty clay. Medium stiff; medium to high plasticity; contains some silty fine sand layers up to 20 mm thick near bottom. s_u (tor) = 0.30 tsf	
20	S9	10								Light gray silty fine sand. Uniform; fines are nonplastic; very fast reaction to shaking test; contains a few clay layers up to 1 mm thick.	
30	S10	30								Similar to Sample S9, but contains clay layers up to 5 mm thick.	
34.0	S11	34								Brown medium to coarse sand. Uniform; subrounded grains; contains a few olive-brown clay layers up to 8 mm thick and a few gravel pieces up to 15 mm in size.	
-20	S12	16								Gray-brown silty gravelly sand. Widely graded; subangular to subrounded grains; fines are nonplastic; contains a few gravel pieces up to 20 mm in size.	
-40	S13	17								Brown gravelly silty sand. Widely graded; subrounded grains; fines are nonplastic; contains a few gravel pieces up to 20 mm in size.	
-60	S14	18								Similar to Sample S13, but also contains subangular grains.	
-80	S15	9								Similar to Sample S13, but clean.	
-40	S16	15*								Brown medium sand. Uniform; subrounded grains; clean.	
60	NX-1	100	3.6	0						Rusty	
70	NX-2	77	5.7	0						Soft-fissile Fresh internally. Locally subject to severe wx, softening. Parts on foliation. Not slickensided.	
70	NX-3	100	3.1	14						Striated Chlorite	
80	NX-4	83	4.0	0						Fault-narrow Fresh. Closely broken on high-angle joints or on partings on high-angle foliation. Fresh, not wx.	
80	NX-5	100	3.0	0						82.5 Faulted	
85	NX-6	100	2.8	33						85.0	
90	XNQ-7	100		2)						N20W, 44NW N 2E, 44SE N80W, 60NE N46W, 55NE	
95	XNQ-8	95		45						Feldspar veining apparent faulted	
96	XNQ-9	100		0						Sense of drag-folding at 90.6 ft is up. zone 82.5-85'	
								BOTTOM	OF BORING		
LEGEND					NOTES			SEABROOK STATION			
N	- Standard penetration resistance, blows/ft				1) s_u (tor) = Shear strength measured with Tovane			PUBLIC SERVICE COMPANY OF NEW HAMPSHIRE			
Rec	- Length recovered/length cored, %				2) Rate of advance not available for NQ-7 to 9.			YANKEE ATOMIC ELECTRIC COMPANY			
RQD	- Length of sound core 4 in. and longer/length cored, %				* - Used 300 lb hammer			 A subsidiary of Parsons Company			
S	- Split spoon sample				x - Oriented core			Date: January 10, 1974			
U	- Undisturbed samples							Project 7286			
S	S - Shelby tube							PAGE 1 of 1			
F	F - Fixed piston							LOG OF BORING SRF 4			
O	O - Osterberg										
D	D - Drilling break										
wx	wx - Weathered, weathering										
	N - Denison										
	P - Pitcher										
	G - GEI										
	k - Coefficient of permeability										



LEGEND

N - Standard penetration resistance, blows/ft
 Rec - Length recovered/length cored, %
 RQD - Length of sound core 4 in. and longer/length cored, %
 S - Split spoon sample Groundwater
 U - Undisturbed samples
 S - Shelby tube N - Denison
 F - Fixed piston P - Pitcher
 O - Osterberg G - GEI
 D - Drilling break k - Coefficient of permeability
 wx - Weathered, weathering

ASCEND

N - Standard penetration resistance, blows/ft
 Rec - Length recovered/length cored, %
 RQD - Length of sound core 4 in. and longer/length cored, %
 S - Split spoon sample ∇ Groundwater
 U - Undisturbed samples
 S - Shelby tube N - Denison
 F - Fixed piston P - Pitcher
 O - Osterberg G - GEI
 D - Drilling break k - Coefficient of
 w - Weathered, weathering permeability

NOEB

SEABROOK STATION
PUBLIC SERVICE COMPANY OF NEW HAMPSHIRE

P-E united engineers & Constructors Inc.

Date: January 11, 1974

Project name

PAGE - 2 of 2

LOG OF BORING SRF 5



BORING LOCATION See Scotland Rd. site plan			INCLINATION Vertical	BEARING	DATE START/FINISH Jan. 4, 1974 / Jan. 8, 1974
CASING ID 3 in.			CORE SIZE 1-5/8 in. BX	TOTAL DEPTH 58.0 ft	DRILLED BY American Drilling & Boring, Manco
GROUND EL (MSL) 17.8 ft			DEPTH TO WATER/DATE 0.0 ft / Jan. 30, 1974	LOGGED BY Soil - K. L. Polk; Rock - J. R. Rand	
EL. MSL ft	SAMPLE No.	RATE OF ADV. min/ft	WATER CONTENT % Graphic	PRESSURE TEST K psi 10 cm/sec	STRIKE, DIP F = Foliation J = Joint C = Contact B = Bedding CORE BREAKS S = Slickenside
17.8	S1, S1A	16'3		100	
	S2	15			
	S3	15			
	S4	18			
	S5	13			
	S6	8			
	S7	5			
0	S8	3			Gray silty clay. Soft; medium plasticity; slightly sticky. s_u (tor) = 0.26 tsf
20	S9	4			Similar to Sample S8, but more sticky. s_u (tor) = 0.15 tsf
	S10	5			Similar to Sample S8, but more sticky. s_u (tor) = 0.14 tsf
-30	S11	5			Similar to Sample S8, but more sticky. s_u (tor) = 0.20 tsf
	S12	5			Similar to Sample S8. s_u (tor) = 0.25-0.30 tsf
-40	S13	4		TOP OF SAND	Similar to Sample S8, few silty fine sand layers to 1 mm thick. s_u (tor) = 0.30 tsf
	S14	13'6"			Layered gray silty clay and silty fine sand. Clay is soft; low to medium plasticity; in layers up to 30 mm thick. Sand is uniform; in layers up to 10 mm thick.
	S15	12		TOP OF TILL	
	S16	9'6"			
-46.5	S17	70		TOP OF ROCK	Gray-brown silty medium to coarse sand. Widely graded; fines are nonplastic; sub-angular to subrounded grains; contains a few gravel pieces up to 8 mm in size.
	S18	65'6"			
-53	BX-1	92	3.0	57	Minor rusty Not wx. Altered by hydro- Minor rusty thermal bleaching.
-58				BOTTOM OF BORING	Cataclastic, foliated. Fused breccia, medium-light greenish-gray.

LEGEND

N - Standard penetration resistance, blows/ft
 Rec - Length recovered/length cored, %
 RQD - Length of sound core 4 in. and longer/length cored, %
 S - Split spoon sample Groundwater
 U - Undisturbed samples
 S - Shelby tube N - Denison
 F - Fixed piston P - Pitcher
 O - Osterberg G - GEI
 D - Drilling break k - Coefficient of
 wx - Weathered, weathering permeability

NOTES

1) s_u (tor) = Shear strength measured with Torvane

SEABROOK STATION

PUBLIC SERVICE COMPANY OF NEW HAMPSHIRE

YANKEE ATOMIC ELECTRIC COMPANY

united engineers a subsidiary of Parsons Company

Date: March 9, 1974

Project 7286

PAGE 1 of 1

LOG OF BORING SRF 6



GEOTECHNICAL ENGINEERS INC.

BORING LOCATION See Scotland Road site plan **INCLINATION** 45° **BEARING** S44E or 126 $^\circ$ **T DATE START/FINISH** Jan. 8, 1974 / Jan. 18, 1974

CASING ID	3 in.	CORE SIZE	1-7/8 in.	TOTAL DEPTH	255.0 ft	DRILLED BY	American Drilling & Boring; T. Canning,	
GROUND EL (MSL)	17.5 ft	DEPTH TO WATER	DATE	0.3 ft	Jan. 18, 1974	LOGGED BY	Soil - K. Polk; Rock - J. R. Rand	
EL. MSL	SAMPLE		RATE OF ADV. min/ft	WATER CONTENT	PRESSURE TEST	STRIKE, DIP	SOIL AND ROCK DESCRIPTIONS	
	Depth ft	Type and No. Rec.					Graphic	gpm

17.5

0 100 S = Slickenside

10

20

0

30

40

50

-20

60

65.5

TOP OF ROCK

70

80

90

100

110

120

130

140

148

NQ-1 92 1.5 42

NQ-2 83 1.3 47

NQ-3 100 1.3 13

NQ-4 77 1.3 23

NQ-5 100 1.4 8

NQ-6 100 1.4 0

NQ-7 93 2.0 25

NQ-8 100 2.1 55

NQ-9 97 1.9 47

NQ-10 85 2.1 12

NQ-11 80 2.1 7

NQ-12 100 2.2 55

NQ-13 98 2.1 73

NQ-14 93 1.6 43

NQ-15 97 2.6 63

NQ-16 83 2.3 33

NQ-17 92 2.0 23

5) N44E, 83NW N46W, 33NE N44E, 33NW N44E, 68NW N46E, 81SE N26W, 53NE

Minor rusty Fairly fresh and Cataclastic. Foliated, saus-
Minor rusty hard throughout but suritized, deformed diorite.
1' core lost altered by bleaching. Fine to medium grained, light
Partings are not greenish to tanish gray.

Minor rusty Aphanitic-yellow-green

Driller broken Not wx. Core broken Cataclastic. Fairly fine-
at close intervals throughout into 1' to grained, light greenish gray.
4' pieces by poor drilling.

Discontinuous Not wx. Bleached by Foliated. Apparent dark
vertical joint hydrothermal alter- (chlorite?) minerals scattered
strikes NW ation to greenish locally.

Striated gray.

2' core lost 112' to 114' core Fused intrusive contact, NE strike
lost in soft zone. Minor rough slick- Diabase. Dark gray line graines
enches. Calcite younger than alteration in
vein country rock, Calcite veinlets.

Soft, some talc

Not wx. Hydrothermally altered through- 95.2 Fused intrusive contact, NE strike
out. Fairly hard. Diabase. Dark gray line graines
Drills well. Partings Calcite. Medium grained
tend to parallel foliation. foliated diorite to 103'. Fine-
grained, foliated with small quartz eyes below. Becomes
very fine-grained.

Slickensided-minor polish Internal slick- 114.94 Greenish tan (?) Diabase (?) Bleached
enches parallel to foliation to greenish tan, aphanitic. May be
Fused mylonite. Hairline foliation.
Welded breccia

Misatch 1' core lost Welded breccia Cataclastic. Fine-
1' core lost Not wx. Hydrothermally altered. Light greenish gray.
Somewhat softened by alteration tan bleach. Joints Locally foliated.

Light greenish tan

Diorite-granodiorite. Medium granodiorite. Light tanish gray (extensive bleaching), foliated, medium grained.

NOTES

- Angle hole
- Washed through soil from 0-65.5' - no sample taken.
- Roller hit to 66.0'
- No clays present; therefore, no water contents were determined.
- This is only a partial list of dip and strike data.

SEABROOK STATION
PUBLIC SERVICE COMPANY OF NEW HAMPSHIRE
YANKEE ATOMIC ELECTRIC COMPANY

united engineers & Constructors Inc
Subsidiary of Parsons Company

Date: February 13, 1974

Project 7286

LEGEND

- N - Standard penetration resistance, blows/ft
- Rec - Length recovered/length cored, %
- RQD - Length of sound core 4 in. and longer/length cored, %
- S - Split spoon sample Groundwater
- U - Undisturbed samples
- S - Shelby tube N - Denison
- F - Fixed platen P - Pitcher
- O - Osterberg G - GEI
- D - Drilling break k - Coefficient of permeability
- wx - Weathered, weathering

PAGE 1 of 2

LOG OF BORING SRF 7

REC	- Length recovered/length cored, %
RQD	- Length of sound core 4 in. and longer/length cored, %
S	- Split spoon sample Groundwater
U	- Undisturbed samples
	S - Shelby tube N - Denison
	F - Fixed piston P - Pitcher
	O - Osterberg G - GEI
D	- Drilling break k - Coefficient of
WX	- Weathered, weathering permeability

NOTES

- 1) Angle hole
 - 2) Washed through soil from 0-65.5' - no sample taken.
 - 3) Roller hit to 66.0'
 - 4) No clays present; therefore, no water contents were determined.
 - 5) This is only a partial list of dip and strike data.

SEABROOK STATION

PUBLIC SERVICE COMPANY OF NEW HAMPSHIRE

YANKEE ATOMIC ELECTRIC COMPANY



Date: February 13, 1974

Project 7288

PAGE 1 of 2 LOG OF BORING SBE 2

BORING LOCATION			See Scotland Road site plan		INCLINATION	45°	BEARING	S44E or 136°	DATE START/FINISH	Jan. 8, 1974 / Jan. 18, 1974
CASING ID		3 in.	CORE SIZE		1-7/8 in.	TOTAL DEPTH	255.0 ft	DRILLED BY	American Drilling & Boring; T. Canning,	
GROUND EL (MSL)			17.5 ft	DEPTH TO WATER/DATE		0.3 ft / Jan. 18, 1974	LOGGED BY	Soil - K. L. Peck; Rock - J. R. Hand		
EL. ft	MSL ft	SAMPLE Type and No.	RATE OF ADV. min./ft	WATER CONTENT %	or RQD	PRESSURE TEST GFM psi	Computed k 10 ⁻⁴ cm/sec	STRIKE, DIP F = Foliation J = Joint C = Contact B = Bedding	CORE BREAKS	SOIL AND ROCK DESCRIPTIONS (Weathering, defects, etc.) (Type, texture, mineralogy, color, hardness, etc.)
S - Slickenside CONTINUED FROM PREVIOUS PAGE										
-150		NQ-18	92	2.0	35					149.7, Broken contact. No visible attitude.
		NQ-19	100	2.4	48					Less medium gray fused breccia.
-160		NQ-20	100	2.2	32			Breccia-fused		Hairline veinlets. Medium dark green (epidote) beginning 155'
-100		NQ-21	100	1.9	45			Driller ground		Epidote
-170		NQ-22	97	1.9	28					Not so bleached as above. Not slickensided.
		NQ-23	100	2.0	51					Offset veinlets
-180		NQ-24	100	2.1	55					Fused
		NQ-25	100	1.7	28					Cataclastic. Fine-grained, locally foliated and brecciated (fused). Medium-dark greenish gray. Epidotized.
-190		NQ-26	100	1.6	72					
-120		NQ-27	100	1.7	25					
		NQ-28	100	1.8	40					
-200		NQ-29	100	1.9	87	N53E, 24NW	J	Moderate bleaching alteration		
		NQ-30	100	1.7	80	N78E, 38NW	J			
-210		NQ-31	100	1.9	52	N80E, 32NW	J			
		NQ-32	100	2.3	55	N80E, 43SE	F			
-220		NQ-33	100	2.1	53	N86W, 42NE	F			
		NQ-34	100	2.1	90	N84W, 90	S			
-230		NQ-35	100	2.5	54	N86E, 70NE	S			
		NQ-36	100	2.6	65	N66E, 17NW	S			
-240		NQ-37	100	2.3	50	N66E, Horizontal	S			
		NQ-38	100	2.6	43	N80E, 73NW	S			
-250		NQ-39	100	2.6	46	N78E, 77SE	S			
-265		NQ-40	64	2.5	18					
						Bottom	OF BORING			

LEGEND	N	- Standard penetration resistance, blows/ft	NOTES x - Oriented core
	Rec	- Length recovered/length cored, %	
	RQD	- Length of sound core 4 in. and longer/length cored, %	
	S	- Split spoon sample	
	U	- Undisturbed samples	
S - Shelby tube		N - Denison	
F - Fixed piston		P - Pitcher	
O - Osterberg		G - GEI	
D - Drilling break		k - Coefficient of permeability	
wx - Weathered, weathering			

SEABROOK STATION
PUBLIC SERVICE COMPANY OF NEW HAMPSHIRE

YANKEE ATOMIC ELECTRIC COMPANY

 **united engineers** contractors inc
A subsidiary of Parsons Company

Date: April 18, 1974

Project 7286

PAGE 2 of 2

LOG OF BORING SRF 7

N	- Standard penetration resistance, blows/ft
Rec	- Length recovered/length cored, %
RQD	- Length of sound core 4 in. and longer/length cored, %
S	- Split spoon sample Groundwater
U	- Undisturbed samples
	S - Shelby tube N - Denison
	F - Fixed piston P - Pitcher
	O - Osterberg G - GEI
D	- Drilling break k - Coefficient of
wx	- Weathered, weathering permeability

NOTES

- 1) Roller bitted to 53 ft.
- 2) s_{u} (tor) = Shear strength measured with Torvane
- 3) This is only a partial list of dip and strike data.

SEABROOK STATION
PUBLIC SERVICE COMPANY OF NEW HAMPSHIRE

YANKEE ATOMIC ELECTRIC COMPANY

United engineers construction
Montreal, Quebec

— 1 —

Date: March 9, 1974

Project 7285

PAGE 1 of 1

LOG OF BORING SRF

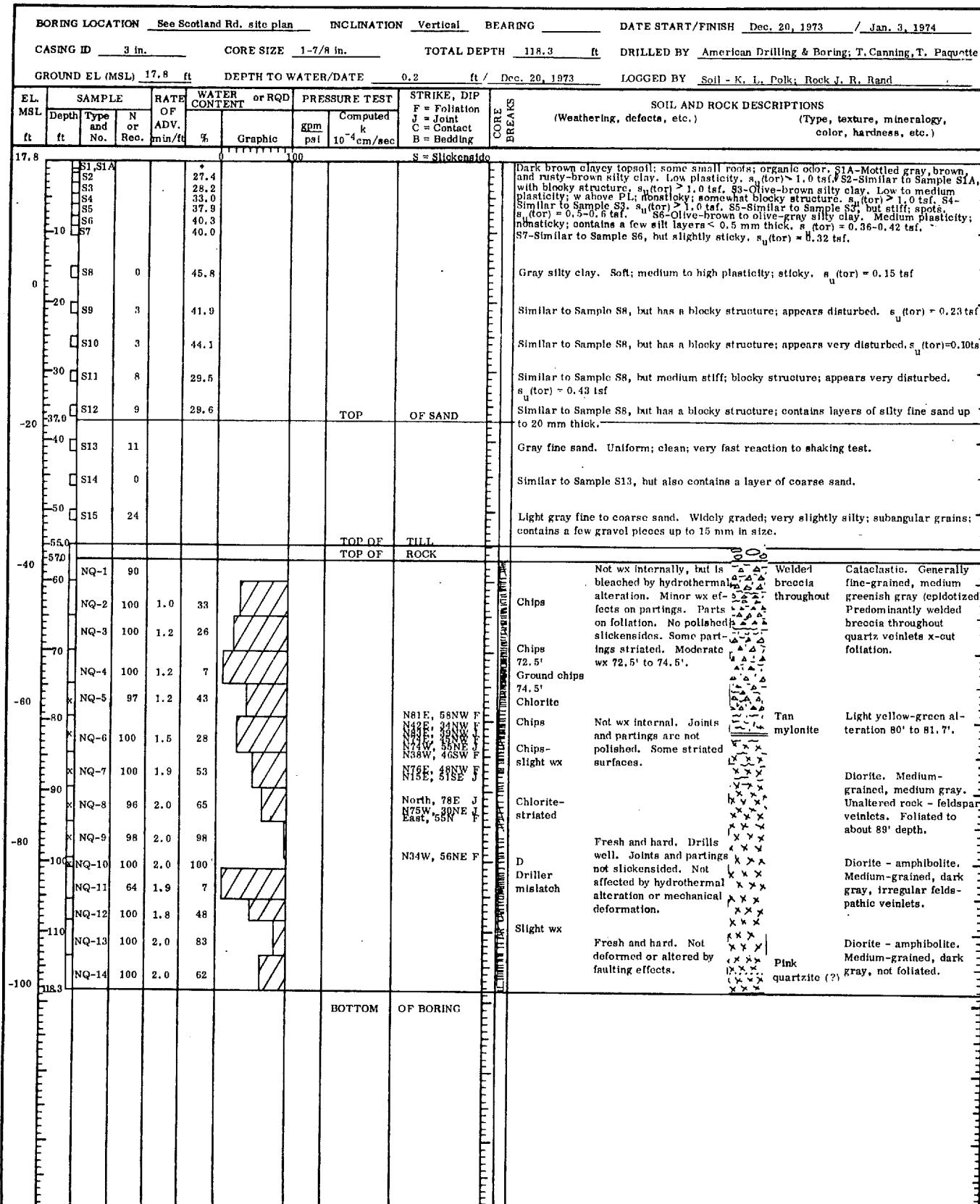


GEOTECHNICAL ENGINEERS INC.

BORING LOCATION See Scotland Rd. site plan				INCLINATION Vertical	BEARING _____	DATE START/FINISH Jan. 25, 1974 / Feb. 19, 1974					
CASING ID 3 in.		CORE SIZE 1-7/8 in.		TOTAL DEPTH 172.0 ft	DRILLED BY American Drilling & Boring; T. Canning						
GROUND EL (MSL) 17.6 ft		DEPTH TO WATER/DATE		Tidal ft / -	LOGGED BY Soil - K. L. Polk; Rock - J. R. Rand						
EL. MSL ft	SAMPLE Type and No. ft	RATE OF ADV. min/ft	WATER CONTENT % Graphic	PRESSURE TEST gpm psi 10 ⁻⁴ k cm/sec	STRIKE, DIP F = Foliation J = Joint C = Contact B = Bedding	CORE BREAKS					
SOIL AND ROCK DESCRIPTIONS (Weathering, defects, etc.)						(Type, texture, mineralogy, color, hardness, etc.)					
S = Slickenside CONTINUED FROM PREVIOUS PAGE											
145	X NQ-23	100	1.1	57							
150	X NQ-24	100	1.2	78	N72E, 68NW J,F N85W, 81SW J N85W, 20NE S	Not polished.					
-140	X NQ-25	100	1.2	77	N86W, 60NE F N82E, 28NW J N59W, 40SW J N85E, 56NW F N34W, 19NE J,F N75E, 51NW F	Not wx. hydrothermally altered to 163. 6'. Fresh, essentially un-altered below partings generally parallel!					
160	X NQ-26	98	1.2	47	Chips	160. 2' to 160. 4'					
170	X NQ-27	98	1.5	57	Chips	Mylonite					
172	X NQ-28	100	1.2	71	Not slickensided	Fault zone-transitional-not slick.					
					Smooth joint	Diorite. Slight alteration and foliated to about 168 ft. Medium-fine grained, medium gray.					
BOTTOM OF BORING											
NOTES											
LEGEND	N	- Standard penetration resistance, blows/ft									
	Rec	- Length recovered/length cored, %									
	RQD	- Length of sound core 4 in. and longer/length cored, %									
	S	- Split spoon sample									
	U	- Undisturbed samples									
	S	- Shelby tube									
	F	- Fixed piston									
	O	- Osterberg									
	D	- Drilling break									
	WX	- Weathered, weathering									
N - Denison P - Pitcher G - GEI k - Coefficient of permeability											
SEABROOK STATION PUBLIC SERVICE COMPANY OF NEW HAMPSHIRE YANKEE ATOMIC ELECTRIC COMPANY											
 A subsidiary of Kiewit Company											
Date: March 9, 1974 Project 7286											
PAGE 2 of 2 LOG OF BORING SRF 8											



GEOTECHNICAL ENGINEERS INC.



ATTACHMENT No. 3

PETROLOGY AND PRELIMINARY INTERPRETATION
OF EIGHT SAMPLES OF DRILL CORE
FROM THE SCOTLANT ROAD FAULT
NEWBURY, MASSACHUSETTS

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PETROGRAPHY AND PRELIMINARY INTERPRETATION
OF EIGHT SAMPLES OF DRILL CORE
FROM THE SCOTLAND ROAD FAULT,
NEWBURYPORT, MASSACHUSETTS

Weston Geophysical Research, Inc.
Post Office Box 306
Weston, Massachusetts 02193

Gene Simmons
Dorothy Richter
15 June 1974

SUMMARY

The petrography of eight samples of drill core from the vicinity of the Scotland Road fault, Newburyport, Massachusetts is described in this report. The samples (with the important exception of sample SRF-5B) all show evidence of dynamic deformation; That is, cataclasis, brecciation, and intense crushing--all probably due to motion along the fault. The deformation clearly took place after the regional metamorphism of the rocks (which was probably associated with the Devonian Acadian orogeny). The microcracks produced in the deformational events appear in thin section to have either annealed, or have been filled by secondary minerals. There is no firm petrographic evidence of recent deformation of these samples.

Sample SRF-5B may be a very important clue to the history of movement on the Scotland Road Fault. It is an altered olivine basalt that seems to be completely free of deformation. If the thin section is representative of a significant volume of this rock, then it may show that no deformation has occurred on the Scotland Road Fault since this rock last cooled below about 500°C. An even stronger statement can be made with respect to movement on the fault after the alteration of the rock was completed: Because the strength of diabase decreases with alteration and because of the absence of deformational features in Sample SRF-5B, we are quite sure that no movement occurred on the fault after the alteration was completed.

Table 1 is a summary of the rock types in the Scotland Road fault suite. Detailed petrographic descriptions and photomicrographs of textural features are given on the following pages.

Table 1. Summary of Samples

<u>Sample #</u>		<u>Rock Type</u>
SRF-1A		Amphibolite breccia
SRF-2A		Mylonized quartz-muscovite schist
SRF-2B		Brecciated quartz-muscovite schist
SRF-3A	67'	Muscovite mylonite
SRF-4A	43'	Chlorite augen gneiss
SRF-5A	42'	Sheared granodiorite
SRF-5B	175'	Altered olivine basalt
SRF-7A	116'	Ultramylonite
SRF-8A	155'	Brecciated quartz-muscovite schist
SRF-8B	146.5'	Brecciated quartz-muscovite schist
SPF-9A	80'	Brecciated quartz-muscovite schist

PETROGRAPHY OF SAMPLE SRF-1A

Name: Amphibolite breccia

Macroscopic Description

This sample is a coarse-grained dark green breccia.

Large (to 1.5cm) angular fragments of dark green amphibole appear set in a finer matrix of crushed amphibole and finer-grained white minerals. Zones of continuous mylonized and sheared materials cut across the sample.

Microscopic Description

Texture

The texture of the thin section is very complex.

Large single crystals can be seen to be split, sheared, rotated, and crushed. The original foliation of the schist is totally disrupted and the crystals are now randomly oriented. Multiple sets of fine parallel cracks and/or inclusion trains can be traced from an amphibole crystal into an adjacent feldspar grain. Coherent fragments of crystals can be "fitted" back together by eye, but they are separated by fibrous chlorite. Large crystals have cataclastic material along grain edges. Calcite veins which crosscut the breccia are themselves deformed, and crosscut by thin veins of undeformed plagioclase.

Mineralogy

Hornblende is the dominant constituent of the rock. It is optically negative with a large axial angle, and pleochroic from pale green to dark greenish brown.

Crystal fragments range in size from 0.01mm- 1.5cm.

The crystals contain abundant inclusion trains and cataclastic material occurs within crystals and bevels grain boundaries. The hornblende appears to be unaltered except for a few overgrowths of blue-green amphibole.

Plagioclase is the second most abundant mineral in the rock. It occurs as untwinned crystals which were probably a part of the original amphibolite schist.

Plagioclase forms large (0.5 - 1.0mm) crystals which are completely covered with linear sets of dusty inclusions. Most crystals appear strained and broken; healed fractures are marked by strings of quartz, calcite, and fresh plagioclase.

Chlorite forms pale green, fibrous, slightly pleochroic aggregates. All crystals display a consistent anomalous "tiger eye" brown interference color. Some of the chlorite seems to be post-brecciation recrystallized mylonitic material which appears to be stretched between crystals. A lesser amount of chlorite appears to be retrograded biotite which is recognized by small amounts of relict biotite and remnant pleochroic haloes.

Calcite appears in veins and fills interstices in the matrix of the rock. Most of the calcite in the veins

is highly distorted and elongated; but there are also minor amounts of undistorted calcite in thin younger veins.

Sphene occurs in accessory amounts as small nodular crystals associated with fuzzy aggregates of leucoxene.

Opaque minerals form stringy aggregates in the mica flakes and more rarely occur as roundish single crystals in the matrix of the rock.

Apatite and Cordierite occur as small euhedral crystals in the matrix of the sample.

Estimated modal composition

amphibole	45%
plagioclase	30%
calcite	15%
opaque	5%
accessories	5%

PETROGRAPHY OF SAMPLE SRF-2A

Name: Mylonized quartz muscovite schist

Macroscopic Description

Sample SRF-2A is a light greenish-grey rock. It appears in hand specimen to be a brecciated cataclasite; in other words, it has a very complex texture which may be the result of multiple deformations. The sample can be separated into different domains of fragments of coarser and finer grained material. The fragments are separated by fine-grained, lighter colored material.

Microscopic Description

Texture

The domains mentioned above appear in thin section as very fine mosaics of granular quartz grains and scaly muscovite. The average grain size is about 0.0mm.

The coherent fragments are separated by shear zones of chlorite, calcite, sphene and ultrafine material which is unresolvable with high magnification.

Mineralogy

Quartz is abundant in the rock fragments and occurs as small (0.2mm) roundish grains. Many grains appear to be crushed and granulated. Most grains have undulose extinction. The quartz crystals are almost always separated from each other by a film of minute mica flakes, except in the coarser grained fragments where they are in direct contact along sutured grain boundaries.

Muscovite occurs as small scaly clusters of crystals.

Muscovite is a major constituent of the rock and has three modes of occurrence -- 1) as minute aggregates completely replacing what was probably feldspar, 2) as thin films around individual quartz crystals, and 3) as part of the shear zones between the rock fragments.

Calcite forms small aggregates in the shear zones and small veins which cut the rock.

Chlorite occurs in the shear zones between the fragments as irregular stringers.

Estimated modal composition

quartz	40%
muscovite	35%
calcite	15%
chlorite	5%
unresolvable material	5%

PETROGRAPHY OF SAMPLE SRF-2B

Name: Brecciated quartz-muscovite schist

Macroscopic Description

This sample is a medium greyish green brecciated rock which is very similar to sample SRF-2A in hand specimen. It is slightly coarser grained than the latter sample but has a similar texture of sheared and brecciated metamorphic rock fragments up to 2cm in size.

Microscopic Description

Texture

The thin section shows a complex texture of brecciated quartz-muscovite rock. The fragments are of various sizes but have an internal uniform grain size of 0.1mm or less. The fragments are separated by zones of unresolvably fine minerals mixed with calcite.

Mineralogy

Quartz is one of the most abundant minerals in this rock. It occurs as irregular but generally ovoid grains which appear to be highly strained and are 0.1mm in size. Most of the quartz grains are not in contact with other quartz grains, and contain relatively few inclusions and bubble trains.

Muscovite forms small scaly masses which thinly separate quartz grains. The muscovite contains many small inclusions of opaques. Muscovite is a common mineral in the shear zones where it has a weblike pattern.

Chlorite is not very abundant in the main body of the rock but it is quite common in the sheared zones between the rock fragments. It is generally very pale green, only slightly pleochroic, and very weakly birefringent.

Biotite occurs as a few relict grains associated with some of the chlorite.

Calcite, clouded with fluid inclusions, fills the shear zones and younger veins. It is also present in the matrix of the fragments as small subhedral crystals.

Opaque grains are widely dispersed throughout the thin sections as minute single crystals and aggregates.

Garnet crystals are present in the sample but are very rare. Crystals <0.1mm in size appear brownish at the core because of tiny opaque inclusions.

Estimated Modal Composition

quartz	35%
muscovite	40%
calcite	15%
chlorite	5%
opaques & accessories	5%

PETROGRAPHY OF SAMPLE SRF-3A 67'

Name: Muscovite Mylonite

Macroscopic Description

This sample is a massive rock, mottled light and dark grey, and almost gneissic in texture. Most grains are too fine-grained to be recognized although enough larger quartz grains are visible to give the sample its banded appearance.

Microscopic Description

Texture

The sample is very fine-grained ($\sim 0.01\text{mm}$) and vaguely schistose in thin section. Very faint outlines of lenticular shapes seem to mark former brecciated fragments. These fragments are obscured by a fine network of stringy mica which have a preferred orientation in another direction. The complex texture of this sample suggests multiple periods of deformation.

Mineralogy

Muscovite is abundant in this sample as ultrafine crystals which are often optically aligned to give a weblike appearance of the mineral. Muscovite is very finely mixed with quartz in the matrix of the rock. It is the major mineral in the sample, although one cannot see it in hand specimen.

Quartz occurs as isolated fragmental crystals in the sample. It generally has indistinct grain boundaries.

Quartz also appears to be mixed with the muscovite at a very fine scale.

Calcite occurs commonly as 0.5mm roundish crystals in the matrix and as thin aggregates following the schistosity.

Opaque grains occur in small knots with streamlined outlines, and small crystals following schistosity.

Estimated Modal Composition

muscovite	70%
quartz	15%
calcite	10%
opaques	5%

Note: Another thin section from this core exhibits similar textures but contains small domains which are calcite rich.

PETROGRAPHY OF SAMPLE SRF-4A 43'

Name: Chlorite augen gneiss

Macroscopic Description

This sample is a fine-grained augen gneiss. It has a dark green matrix of indistinguishable minerals and 0.5mm "eyes" of white crystals. The sample shows strong directional foliation which is crosscut by younger veins of light colored minerals.

Microscopic Description

Texture

In thin section, the sample shows a complex, almost chaotic texture. It is basically a mosaic of fragmental quartz and feldspar crystals and aggregates with lenticular shapes sandwiched by shear zones of chlorite, calcite, and opaques. Thin veins of calcite cut the foliation.

Mineralogy

Chlorite is the most abundant mineral in the rock.

It is pale green, pleochroic, and exhibits anomalous brown interference colors. Very fine, scaly aggregates of chlorite are commonly finely mixed with quartz and opaque grains. Larger crystals of chlorite show small amounts of relict biotite.

Plagioclase occurs as intensely sericitized, poorly twinned, fragmented crystals in the augen.

Quartz has three modes of occurrence in this sample: 1) large broken crystals in the augen, 2) very finely

mixed in the matrix, and 3) fresh crystals in thin veinlets.

Calcite is a very common mineral in the matrix, shear zones, and in veins. It commonly has deformed twin planes.

Orthoclase occurs in accessory amounts as anomalously fresh appearing fragmental crystals in the augen.

Opaque grains are widely dispersed throughout the thin section as minute crystals.

Estimated Modal Composition

chlorite	35%
plagioclase	20%
quartz	15%
calcite	20%
orthoclase	5%
opaque	5%

Note -- the bulk mineral composition of this sample suggests that its protolith was a mafic igneous rock.

PETROGRAPHY OF SAMPLE SRF-5A 42'

Name: Sheared granodiorite

Macroscopic Description

This sample appears in hand specimen to be a massive, coarse-grained igneous rock with no evidence of deformation. The average grain size is approximately 1mm. Visible in hand specimen are pink feldspar, white quartz, and an unknown green mineral.

Microscopic Description

Texture

The thin section has the hypidiomorphic granular texture typical of plutonic rocks. Equidimensional crystals showing varying degrees of alteration are crosscut by thin veinlets. The major deformational features in the thin section are: healed cracks, undulose extinction of the minerals, and a narrow shear zone.

Mineralogy

"Plagioclase", once a major component of this sample, has been completely kaolinized with only a few rare traces of the original twinning or textures left. The kaolinization reaction produces excess SiO_2 which can be seen in the thin section as a thin rim around each kaolinized grain. These peculiar rims are optically uniform around each crystal. The rims only occur along feldspar-feldspar contacts but do not occur along feldspar-quartz contacts.

Quartz occurs as 1 mm blocky crystals with undulose extinction and numerous inclusion trains. Quartz-feldspar boundaries are generally smooth whereas quartz-quartz boundaries are sutured, a sign of partial recrystallization. Quartz also occurs in the rims around kaolinized feldspar grains as mentioned above.

Microcline occurs as slightly altered crystals with a microperthitic texture.

Chlorite forms pseudomorphs after biotite and amphibole. It is medium green, weakly pleochroic, and contains abundant needles of opaques.

Calcite occurs as small clusters of crystals finely mixed with kaolinite alteration products, as thin veinlets, and as aggregates in the matrix. Calcite also fills the one shear zone in the thin section.

Accessory minerals in this rock are opaques, apatite, and sphene.

Estimated Modal Composition

"plagioclase"	40%
microcline	20%
quartz	25%
chlorite	12%
opaque & accessories	3%

PETROGRAPHY OF SAMPLE SRF-5B 175'

Name: Altered olivine basalt

Macroscopic Description

This is a massive, dark greenish grey aphanitic rock.

Small dark phenocrysts (0.5 - 1.0mm) and 0.5mm white amygdules are visible in the black groundmass. There are no signs of deformation such as shear zones or even veins.

Microscopic Description

Texture

The sample has a very fine-grained (<0.1mm) intersertal texture. The matrix texture is somewhat obscured by partial alteration of the minerals. The vesicles are rimmed with fibrous minerals. The phenocrysts are completely replaced by alteration minerals.

Mineralogy

Plagioclase occurs as small (0.1mm or less) laths in the matrix of the rock. It does not form any phenocrysts. The plagioclase is generally poorly twinned and partially altered to a sericitic product.

Pyroxene crystals occur as small roundish grains with small scale intergrowths with opaque rods. It is pinkish brown in color and is probably augite.

Serpentine completely replaces roundish 1.0mm phenocrysts of olivine. Serpentine also occurs as fibers in the matrix of the rock, and as the lining of the amygdules.

Calcite forms twinned single crystals in the amygdules
and is otherwise rare in the matrix.

Estimated Modal Composition

plagioclase	35%
pyroxene	35%
serpentine	10%
calcite	10%
sericitic alteration	10%

Note -- This sample is probably from a dike which post-dates
movement along the Scotland Road fault since it is
completely undeformed.

PETROGRAPHY OF SAMPLE SRF-7A 116'

Name: Ultramylonite

Macroscopic Description

This is a compact, extremely fine-grained, mustard colored rock. A few small whitish augen (0.5 - 1.0mm) are visible in the hand specimen. The matrix is buff colored, highly sheared looking material.

Microscopic Description

Texture

This is an ultrafine-grained crush breccia. The original texture of the rock is totally obliterated. The apparent mineral layering is due to 'smearing' of the grain in local shear zones.

Mineralogy

The rock is so fine-grained that individual crystals are difficult to discriminate, except in the few augen of quartz, calcite, and opaque minerals. The matrix is extremely finely-ground quartz, mica, calcite, sphene, apatite, and opaque minerals. Calcite occurs in small nodules which show some signs of recrystallization.

Note -- the fine-grained nature of this rock precludes any further discussion of its mineralogy or texture.

PETROGRAPHY OF SAMPLE SRF-8A 155

Name: Brecciated quartz-muscovite schist

Macroscopic Description

This sample is a dark greenish grey rock. On a fresh surface it appears to be a fine grained quartzite cut by narrow black shear zones and mottled tan zones. The wet sawed surface shows the texture of a breccia with distinct fragments ranging in size from 1mm to 1cm. The fragments are separated by the tan material; both are cut by the black shear zone.

Microscopic Description

Texture

The texture in thin section is similar to other samples in the suite. Lenticular fragments of various sizes of quartz muscovite rock are separated by ultrafine-grained shear zones. Average grain size is 0.1mm. The relative proportions of quartz and muscovite varies from fragment to fragment.

Mineralogy

Quartz occurs as roundish grains which are almost always isolated from each other by varying amounts of muscovite. Some of the crystals appear to be broken.

Muscovite forms scaly masses which are vaguely schistose.

Muscovite is a major component of the rock, filling interstices, between quartz grains, shear zones. It forms the bulk of several lithic fragments.

Chlorite is a major constituent of the sheared zones between lithic fragments although it is not abundant in the fragments themselves. It is pale green, slightly pleochroic, and exhibits anomalous blue interference colors.

Opaque grains, finely mixed with leucoxene, form intricate intergrowths pseudomorphous after tabular biotite plates and occur as euhedral crystals in the lithic fragments. Calcite is common in the shear zones as elongate crystals. It also occurs as minute single crystals in the lithic fragments, and in a few thin, undeformed veins.

Sphene forms fine granular aggregates in the matrix of the fragments and occurs as stringers in the shear zones.

Estimated modal composition

Quartz	45%
Muscovite	30%
Chlorite	10%
Opaque	5%
Calcite	5%
Sphene	5%

PETROGRAPHY OF SAMPLE SRF-8B 146.5'

Name: Brecciated quartz-muscovite schist

Macroscopic Description

This sample is strikingly similar to SRF-8A in hand specimen. It is dark greenish-grey in color. On a fresh broken surface, it appears fine grained and structureless. On the sawed surface, one can see lenticular fragments of various sizes, thinly outlined by lighter colored material. The core is broken along a major fracture surface.

Microscopic Description

Texture

The texture of the sample is variable and complex. The rock fragments consist of roundish quartz grains and scaly mica; the grain size and composition of the fragments vary. The lithic fragments are separated by mylonite which consists of ground quartz, mica, chlorite, and calcite.

Mineralogy

Quartz is the most abundant and most coarsely grained mineral in the rock. It occurs as roundish grains which vary in size (0.1-0.3mm) and abundance (60%-40%) in the different lithic fragments. The crystals commonly contain inclusions. Quartz crystals are rare in contact with each other. A minor amount of quartz occurs in thin veins which cut the rock and probably post-date the brecciation.

Muscovite occurs as scaly aggregates whose crystals are

much less than 0.1mm in size. The aggregates form most of the matrix of the lithic fragments. Submicroscopic muscovite appears to occur in the mylonized zones. Chlorite forms pale green 0.1mm crystals in the shear zones.

Chlorite less commonly occurs in the matrix of the lithic fragments.

Opaque grains occur in the shear zones, in the matrix and in a few rare veins.

Carbonate forms irregular clusters of crystals in the shear zones but does not occur in the lithic fragments.

Sphene occurs in minor amounts as grainy aggregates in the matrix of fragments and in the mylonized zones.

Estimated modal composition

Quartz	40%
Muscovite	40%
Chlorite	10%
Calcite	5%
Opaques	5%

PETROGRAPHY OF SAMPLE SRF-9A 80'

Name: Brecciated quartz-muscovite schist

Macroscopic Description

The texture of this sample is similar to that of samples 8A and 8B, although the rock is light tannish-grey in color. Lenticular and irregularly shaped fragments 0.1-1cm in size are recognizable in a highly sheared matrix. Individual minerals are too fine-grained to recognize in hand specimen. Thin veins of light-colored minerals and, more rarely, opaques are present.

Microscopic Description

Texture

The thin section exhibits the chaotic texture of the rock. Lenticular quartz-muscovite lithic fragments are elongate parallel to foliation. Mylonized zones appear to be structureless. Irregular semi-parallel veinlets cut the foliation.

Mineralogy

Quartz occurs as roundish grains in the lithic fragments.

The grains appear to be highly strained and in places broken. They commonly contain linear arrays of inclusions. Very finely ground quartz is apparently a constituent in the mylonite zones. Several thin veins of quartz cut the rock. The margins of the veins are commonly sutured and show signs of recrystallization; in some places the vein quartz is optically continuous with quartz grains which it cuts.

Muscovite forms scaly masses between quartz grains in the lithic fragments. The individual crystals are minute but seem to show a general preferred orientation parallel to the foliation. Muscovite appears to be relatively more abundant in the finer-grained lithic fragments than in the coarser-grained fragments.

Calcite is prominent in the mylonized zones and in a few veins. It occurs less commonly in the matrix of the rock fragments.

Sphene aggregates are also common in the shear zones but sparsely distributed in the rest of the rock.

Opaques seem to be concentrated in the shear zones between lithic fragments in clusters of 0.1mm crystals. They also occur in a few veins and as euhedral crystals in the fragments.

Estimated modal composition

Quartz	35%
Muscovite	30%
Calcite	20%
Sphene	5%
Opaques	10%

Note: The light color of this sample is apparently due to the virtual absence of chlorite in the shear-zones coupled with the relative abundance of calcite.



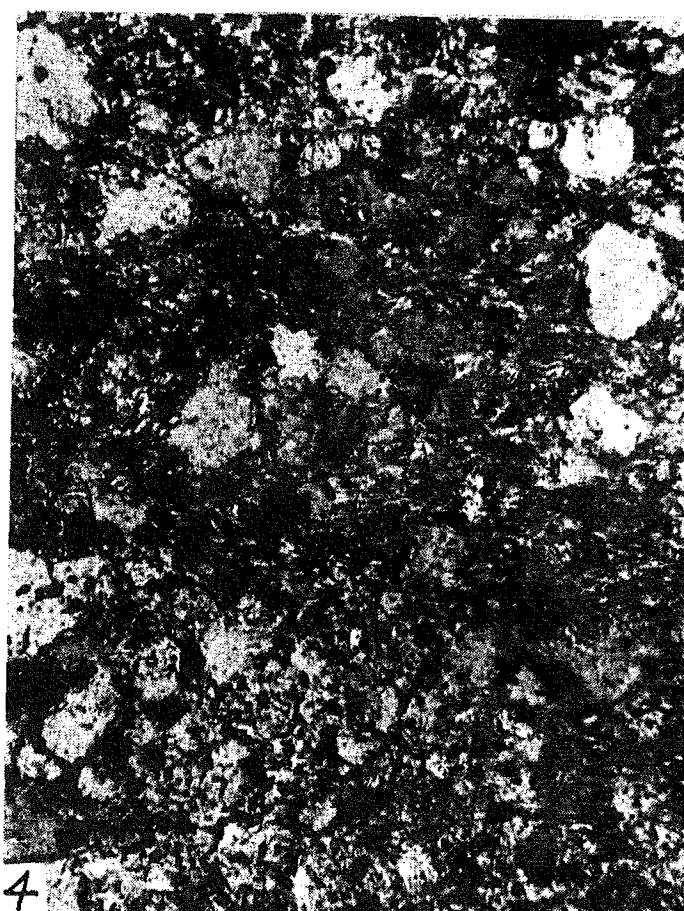
Photo 1. Sample SRF-1A.
Amphibolite breccia. Plane
polarized light. Width of
field 1.5mm. This photo-
micrograph shows a typical
field of view of this sample.
Note that the large dark-
ish hornblende crystals
are sheared. The lighter
grey crystals are plagio-
clase. See also Photo 2.



Photo 2. Sample SRF-1A. Amphi-
bolite breccia. Crossed polarized
light. Width of field 1.5mm.
This photomicrograph shows a
major shear zone in the rock.
The elongate crystals are de-
formed calcite. See also Photo 1.



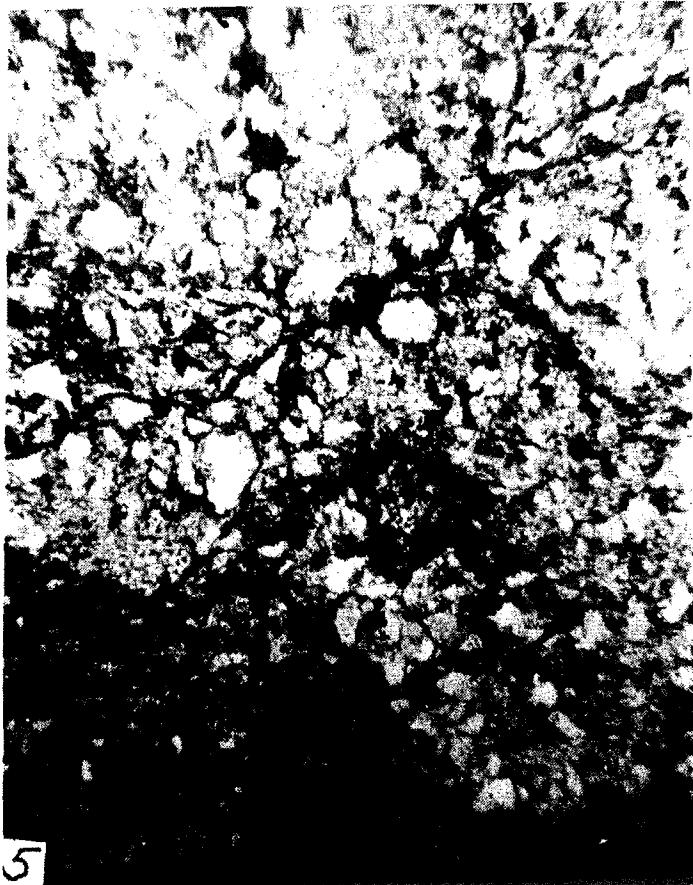
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4

Photo 3. Sample SRF-2A. Mylonized quartz-muscovite schist. Cross polarized light. Width of field 1.5mm. This photomicrograph shows one large lithic fragment covering 3/4 of the photograph and consisting of roundish quartz grains and fuzzy muscovite. The dark zones around the fragment are shear zones of chlorite and other unresolvable minerals. See Photo 4 for an enlargement of the lithic fragment.

Photo 4. Sample SRF-2A. Mylonized quartz-muscovite schist. Plane polarized light. Width of field 0.5mm. This photo is an enlargement of the large lithic fragment shown in Photo 3. The roundish grains are quartz, and the matrix is scaly muscovite, opaques, sphene, and tiny euhedral crystals of calcite as in the left center of the photo.



5



6

Photo 5. Sample SRF-2B.
Brecciated quartz-muscovite
schist. Plane polarized
light. Width of field 1.5mm.
This photomicrograph shows
the chaotic texture typical
of this rock. Note the lenticular
fragments of varying
grain sizes. The round
white crystals are quartz;
the darker minerals are
scaly muscovite, sphene,
calcite and opaques. See
also Photo 6, an enlargement
of a part of this field
magnified. Note the similarity
of this sample with
SRF-2A.

Photo 6. An enlargement of
a portion of Photo 5. Sample
SRF-2B. Brecciated quartz-muscovite
schist. Plane polarized
light. Width of field 0.5mm.
This photomicrograph shows the
chaotic texture typical of this
rock. The round white crystals
are quartz; the darker minerals
are scaly muscovite, sphene, cal-
cite and opaques. Note the
similarity of this sample with
SRF-2A.



7

Photo 7. Sample SRF-3A 67°.
Muscovite mylonite. Crossed
polarized light. Width of
field 1.5mm. This photo-
micrograph shows the
typical texture of this
very fine-grained sample.
The few larger grains are
fragmented quartz crystals.
They are set in a finely
ground matrix of quartz,
muscovite and lesser amounts
of calcite, sphene, and
opaques.



8

Photo 8. Sample SRF-4A 43°.
Chlorite augen gneiss. Plane
polarized light. Width of field
1.5mm. This photomicrograph shows
a polycrystalline 'eye' (lower
half of photo) in a crushed and
sheared matrix. The light grains
in the photo are mostly plagio-
clase and quartz. The large
darker grey crystals are chlorite.
Note the concentration of
opaques in the shear zone in the
upper right corner.



9



10

Photo 9. Sample SRF-5A 42°.
Sheared granodiorite.
Crossed polarized light.
This photomicrograph shows
a typical field of view
of this sample. Note
the large fuzzy grains.
They are kaolinized plagi-
oclase crystals which have
narrow rims of optically
continuous quartz. These
rims were probably pro-
duced as a result of the
kaolinization. Note that
the rims do not continue
along a quartz-plagioclase
grain boundary at the left.
The medium grey grains are
micropertite, and the
light grey grains are quartz.

Photo 10. Sample SRF-5B 175°.
Altered olivine basalt. Plane
polarized light. Width of field
1.5mm. This photomicrograph is a
good example of the texture of
this sample. In the upper left
is an amygdale filled with twinned
calcite and lined with fibrous
serpentine. At the right is a
phenocryst of olivine which has
been completely replaced by serpen-
tine. The matrix consists of
laths of plagioclase (white) and
darker crystals of pyroxene and
black opaques. See also photo
11, an enlargement of the matrix.



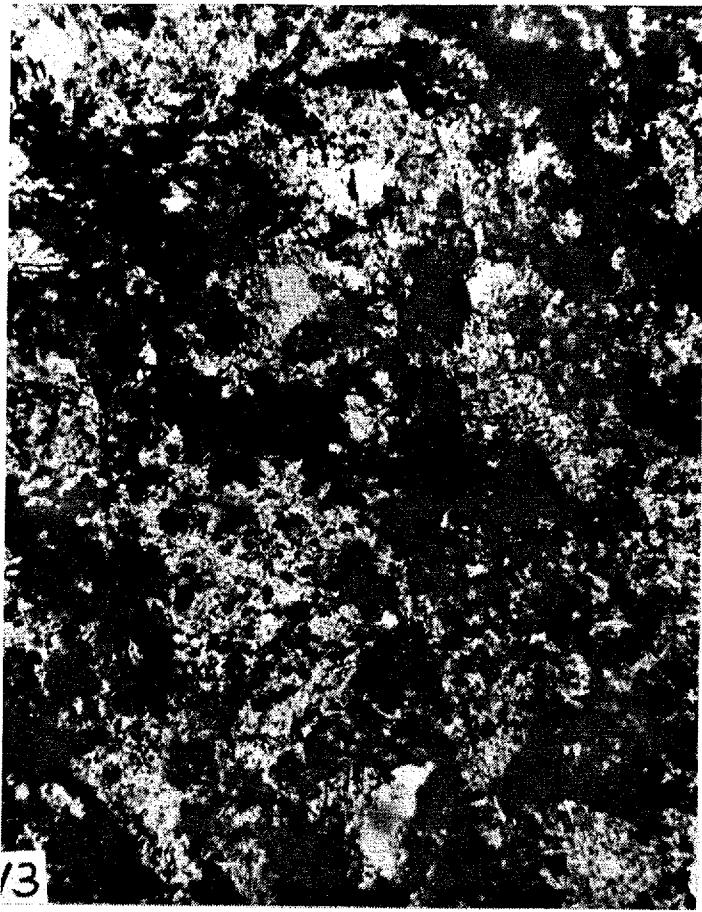
11



12

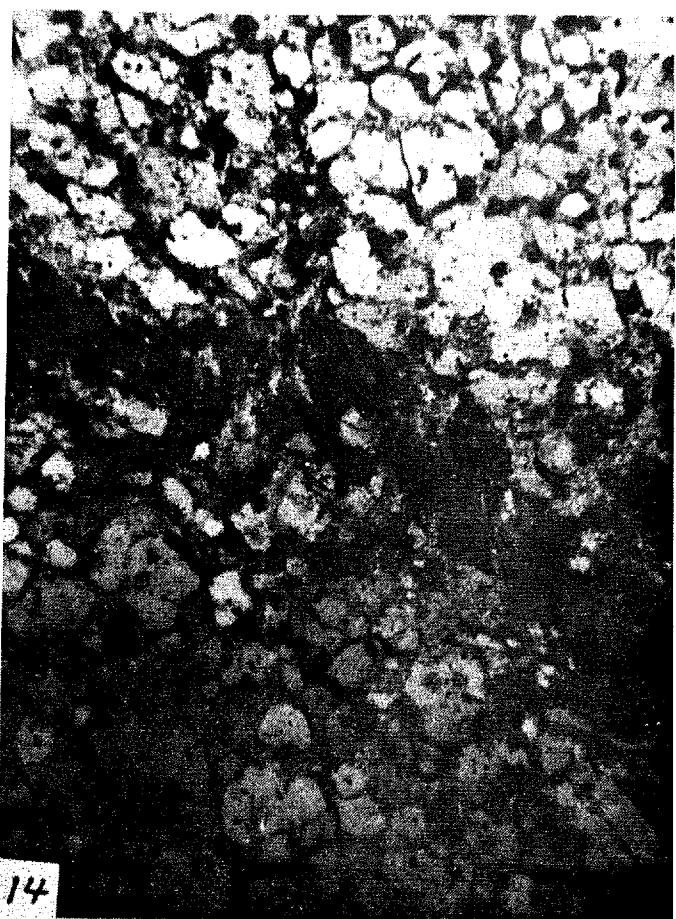
Photo 11, an enlargement of a portion of photo 10. Sample SRF-5B 175'. Altered olivine basalt. Plane polarized light. Width of field 0.5mm. This photomicrograph is an enlargement of the matrix.

Photo 12. Sample SRF-7A 116'. Ultramylonite. Plane polarized light. Width of field 1.5mm. This photomicrograph shows typical texture of this rock. Dark shear zones can be distinguished against the background of highly crushed minerals. See also photo 13.



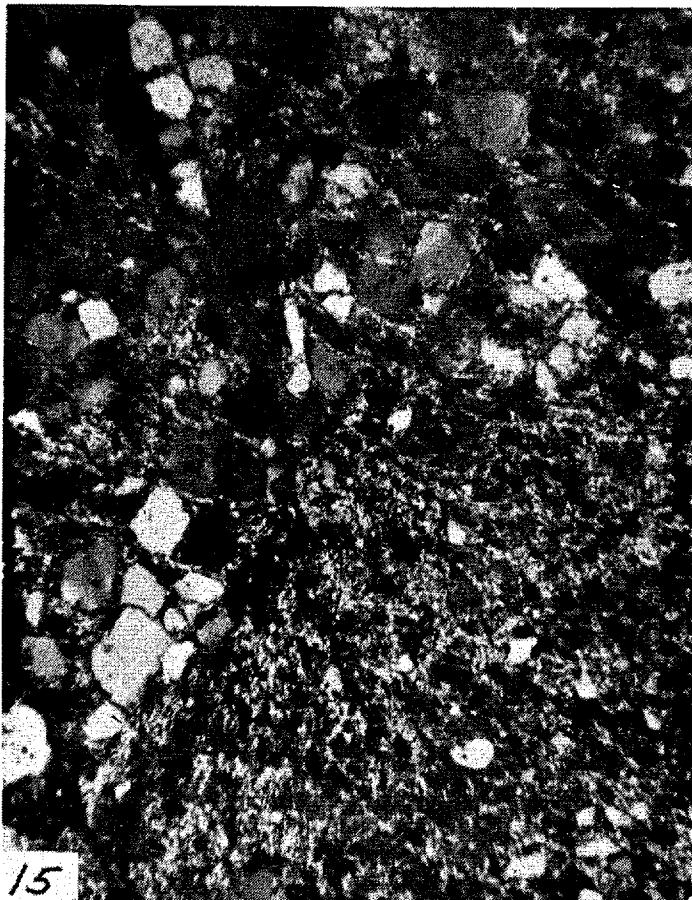
13

Photo 13. an enlargement of a portion of photo 12. Sample SRF-7A 116'. Ultramylonite. Plane polarized light. Width of field 0.5mm. The rock is so pulverized that only a few grains can be identified with certainty--some dark nodular sphene, a few quartz grains and a few aggregates of calcite.



14

Photo 14. Sample SRF-8A 155'. Brecciated quartz-muscovite schist. Plane polarized light. Width of field 1.5mm. This photomicrograph shows a typical field of view. Two large lithic fragments are separated by a dark grey shear zone consisting of chlorite, calcite, and finely ground quartz and muscovite. The white grains in the rock fragments are quartz which are surrounded by darker muscovite, calcite, sphene, and opaque grains.



15



16

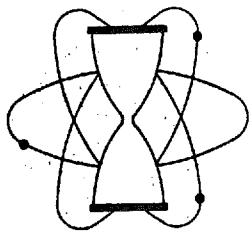
Photo 15. Sample SRF-8B 146.5°. Brecciated quartz muscovite schist. Cross polarized light. Width of field 1.5mm. This photomicrograph shows parts of three lithic fragments. Two of the fragments are coarser-grained than the fragment in the lower right. The larger roundish grains are quartz and the fuzzy material is fine grained masses of muscovite. A thin black line of chlorite and opaques separate the three fragments. Note the similarity of this sample to SRF-8A.

Photo 16. Sample SRF-9A 80°. Brecciated quartz-muscovite schist. Cross polarized light. Width of field 1.5mm. The left hand side of the photomicrograph shows a lithic fragment of roundish quartz grains surrounded by scaly masses of muscovite. At the right is a stringy mylonite zone consisting of pulverized quartz and muscovite with carbonate and opaques. This sample is similar to samples SRF-8A and 8B except for the absence of chlorite.

ATTACHMENT No. 4

K-Ar AGE DETERMINATIONS OF SIX SAMPLES
FROM THE SCOTLAND ROAD FAULT ZONE

GEOCHRON LABORATORIES DIVISION,
KRUEGER ENTERPRISES, INC.
CAMBRIDGE, MASSACHUSETTS 02139
for
WESTON GEOPHYSICAL RESEARCH, INC.
WESTBORO, MASSACHUSETTS 01581



KRUEGER ENTERPRISES, INC.
GEOCHRON LABORATORIES DIVISION

24 BLACKSTONE STREET • CAMBRIDGE, MASSACHUSETTS 02139 • (617) 876-3691

16 May 1974

Richard J. Holt
Weston Geophysical
P.O. Box 364
Weston, MA 02193

Dear Mr. Holt:

Enclosed are the analytical reports of the K-Ar age determinations on two (2) of the six (6) samples sent to us by Gene Simmons at M.I.T. I have already given these results to you by telephone.

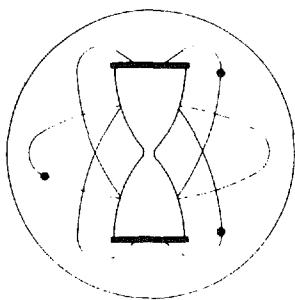
We analyzed sample 5B as a whole rock and obtained an age of about 199 m.y., and we analyzed a sericite concentrate from 8A and obtained an age of about 248 m.y. This latter concentrate contained a significant amount of feldspar, but with a sample of this sort it is often not possible to obtain a high quality mica concentrate. The measured age of sample 8A should be a reasonably good metamorphic age for the rock.

If you have any questions, please do not hesitate to contact me. In the meantime, I am enclosing our invoice for this work. I will contact you as soon as the remaining samples have been analyzed.

Sincerely,

Richard H. Reesman
General Manager

RHR/dm
encl: 2 reports & invoice #4401



KRUEGER ENTERPRISES, INC.

GEOCHRON LABORATORIES DIVISION

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POTASSIUM-ARGON AGE DETERMINATION

REPORT OF ANALYTICAL WORK

Our Sample No. R-2813

Date Received: 22 April 1974

Your Reference: SRF 5B (175.1')

Date Reported: 14 May 1974

Submitted by:

Richard Holt
Weston Geophysical Res. Inc.
P.O. Box 364
Weston, MA 02193

Sample Description & Locality: Dark basalt drill core, SRF 5B (175.1')

Material Analyzed: Whole rock, crushed to -40/+100 mesh.

$\text{Ar}^{40*}/\text{K}^{40}$ = .01230

AGE = 199 ± 9 M.Y.

Argon Analyses:

Ar^{40*} , ppm.	$\text{Ar}^{40*}/\text{Total Ar}^{40}$	Ave. Ar^{40*} , ppm.
.01647	.686	.01638
.01628	.645	

Potassium Analyses:

% K	Ave. %K	K^{40} , ppm
1.095	1.091	1.331
1.087		

Constants Used:

$$\lambda_\beta = 4.72 \times 10^{-10} / \text{year}$$

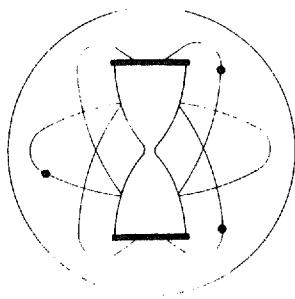
$$\text{AGE} = \frac{1}{\lambda_e + \lambda_\beta} \cdot \ln \left[\frac{\lambda_\beta + \lambda_e}{\lambda_e} \times \frac{\text{Ar}^{40*}}{\text{K}^{40}} + 1 \right]$$

$$\lambda_e = 0.585 \times 10^{-10} / \text{year}$$

$$\text{K}^{40}/\text{K} = 1.22 \times 10^{-4} \text{ g./g.}$$

Note: Ar^{40*} refers to radiogenic Ar^{40} .

M.Y. refers to millions of years.



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GEOCHRON LABORATORIES DIVISION

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POTASSIUM-ARGON AGE DETERMINATION

REPORT OF ANALYTICAL WORK

Our Sample No. M-2820

Date Received: 26 April 1974

Your Reference: SRF 8A

Date Reported: 15 May 1974

Submitted by: Richard J. Holt
Weston Geophysical
P.O. Box 364
Weston, MA 02193

Sample Description & Locality: Sericitized meta-sediment, drill core #SRF 8A.

Material Analyzed: Sericite concentrate with substantial feldspar remaining.

$\text{Ar}^{40*}/\text{K}^{40}$ = .01550

AGE = 248 ± 9 M.Y.

Argon Analyses:

Ar^{40*} , ppm.	$\text{Ar}^{40*}/\text{Total Ar}^{40}$	Ave. Ar^{40*} , ppm.
.09410	.891	.09629
.09848	.791	

Potassium Analyses:

% K	Ave. %K	K^{40} , ppm
5.086	5.092	6.212
5.099		

Constants Used:

$$\lambda_\beta = 4.72 \times 10^{-10} / \text{year}$$

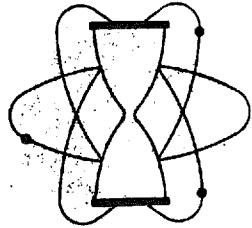
$$\text{AGE} = \frac{1}{\lambda_e + \lambda_\beta} \ln \left[\frac{\lambda_\beta + \lambda_e}{\lambda_e} \times \frac{\text{Ar}^{40*}}{\text{K}^{40}} + 1 \right]$$

$$\lambda_e = 0.585 \times 10^{-10} / \text{year}$$

$$\text{K}^{40}/\text{K} = 1.22 \times 10^{-4} \text{ g./g.}$$

Note: Ar^{40*} refers to radiogenic Ar^{40} .

M.Y. refers to millions of years.



KRUEGER ENTERPRISES, INC.
GEOCHRON LABORATORIES DIVISION

24 BLACKSTONE STREET • CAMBRIDGE, MASSACHUSETTS 02139 • (617) 876-3691

31 May 1974

Richard Holt
Weston Geophysical Research Inc.
P.O. Box 364
Weston, MA 02193

Dear Mr. Holt:

Enclosed are the analytical reports of the K-Ar age determinations on the remaining four (4) samples of the six (6) we received from Gene Simmions last month.

The amphibole in SRF 1A gave an age of 324 m.y. Samples SRF 2A, SRF 3A, and SRF 5A 42' were analyzed as whole rocks and gave indistinguishable ages of 256 m.y., 269 m.y., and 272 m.y. respectively.

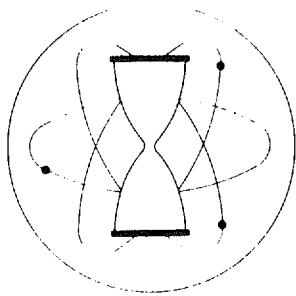
Judging from past analyses we have done for you I suspect these numbers are about what you expected.

If you have any questions, please do not hesitate to contact me. In the meantime, I am enclosing our invoice for this work. We look forward to serving you again in the near future.

Sincerely,

Richard H. Reesman
Richard H. Reesman
General Manager

RHR/dm
encl: 4 reports & invoice # 4414



KRUEGER ENTERPRISES, INC.

GEOCHRON LABORATORIES DIVISION

24 BLACKSTONE STREET • CAMBRIDGE, MA. 02139 • (617)-876-3691

POTASSIUM-ARGON AGE DETERMINATION

REPORT OF ANALYTICAL WORK

Our Sample No. A-2814

Date Received: 22 April 1974

Your Reference: SRF 1A up

Date Reported: 31 May 1974

Submitted by:

Richard Holt
Weston Geophysical Res. Inc.
P.O. Box 364
Weston, MA 02193

Sample Description & Locality: Coarse-grained amphibolite

Material Analyzed: Amphibole concentrate, -40/+100 mesh.

Ar^{40*}/K⁴⁰ = .02069

AGE = 324 ± 14 M.Y.

Argon Analyses:

Ar ^{40*} , ppm.	Ar ^{40*} / Total Ar ⁴⁰	Ave. Ar ^{40*} , ppm.
.01967	.679	.01974
.01981	.704	

Potassium Analyses:

% K	Ave. %K	K ⁴⁰ , ppm
.786	.782	.954
.778		

Constants Used:

$$\lambda_\beta = 4.72 \times 10^{-10} / \text{year}$$

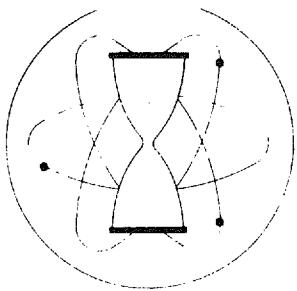
$$\text{AGE} = \frac{1}{\lambda_e + \lambda_\beta} \ln \left[\frac{\lambda_\beta + \lambda_e}{\lambda_e} \times \frac{\text{Ar}^{40*}}{\text{K}^{40}} + 1 \right]$$

$$\lambda_e = 0.585 \times 10^{-10} / \text{year}$$

$$\text{K}^{40}/\text{K} = 1.22 \times 10^{-4} \text{ g./g.}$$

Note: Ar^{40*} refers to radiogenic Ar⁴⁰.

M.Y. refers to millions of years.



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GEOCHRON LABORATORIES DIVISION

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POTASSIUM-ARGON AGE DETERMINATION

REPORT OF ANALYTICAL WORK

Our Sample No. R-2817

Date Received: 26 April 1974

Your Reference: SRF 2A

Date Reported: 31 May 1974

Submitted by: Richard J. Holt
Weston Geophysical
P.O. Box 364
Weston, MA 02193

Sample Description & Locality: Sericite schist

Material Analyzed: Whole rock, crushed to -60/+100 mesh.

$\text{Ar}^{40*}/\text{K}^{40} = .01604$

AGE = 256 \pm 10 M.Y.

Argon Analyses:

Ar^{40*} , ppm.	$\text{Ar}^{40*}/\text{Total Ar}^{40}$	Ave. Ar^{40*} , ppm.
.03235	.676	.03307
.03378	.807	

Potassium Analyses:

% K	Ave. %K	K^{40} , ppm
1.699	1.689	2.061
1.680		

Constants Used:

$$\lambda_\beta = 4.72 \times 10^{-10} / \text{year}$$

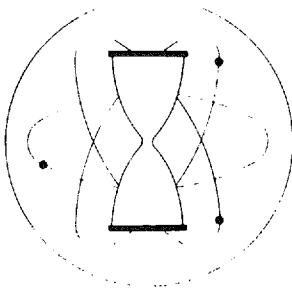
$$\text{AGE} = \frac{1}{\lambda_e + \lambda_\beta} \ln \left[\frac{\lambda_\beta + \lambda_e}{\lambda_e} \times \frac{\text{Ar}^{40*}}{\text{K}^{40}} + 1 \right]$$

$$\lambda_e = 0.585 \times 10^{-10} / \text{year}$$

$$\text{K}^{40}/\text{K} = 1.22 \times 10^{-4} \text{ g./g.}$$

Note: Ar^{40*} refers to radiogenic Ar^{40} .

M.Y. refers to millions of years.



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POTASSIUM-ARGON AGE DETERMINATION

REPORT OF ANALYTICAL WORK

Our Sample No. R-2818

Date Received: 26 April 1974

Your Reference: SRF 3A

Date Reported: 31 May 1974

Submitted by: Richard J. Holt
Weston Geophysical
P.O. Box 364
Weston, MA 02193

Sample Description & Locality: Sericite schist

Material Analyzed: Whole rock, crushed to -60/+100 mesh.

Ar^{40*}/K⁴⁰ = .01690

AGE = 269 ± 10 M.Y.

Argon Analyses:

Ar ^{40*} , ppm.	Ar ^{40*} / Total Ar ⁴⁰	Ave. Ar ^{40*} , ppm.
.07748	.913	.07756
.07763	.787	

Potassium Analyses:

% K	Ave. %K	K ⁴⁰ , ppm
3.782	3.761	4.589
3.741		

Constants Used:

$$\lambda_\beta = 4.72 \times 10^{-10} / \text{year}$$

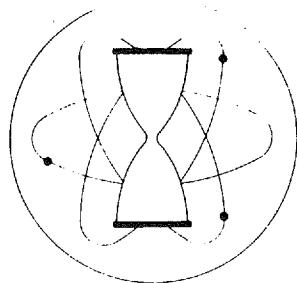
$$\text{AGE} = \frac{1}{\lambda_e + \lambda_\beta} \ln \left[\frac{\lambda_\beta + \lambda_e}{\lambda_e} \times \frac{\text{Ar}^{40*}}{\text{K}^{40}} + 1 \right]$$

$$\lambda_e = 0.585 \times 10^{-10} / \text{year}$$

$$\text{K}^{40}/\text{K} = 1.22 \times 10^{-4} \text{ g./g.}$$

Note: Ar^{40*} refers to radiogenic Ar⁴⁰.

M.Y. refers to millions of years.



KRUEGER ENTERPRISES, INC. GEOCHRON LABORATORIES DIVISION

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POTASSIUM-ARGON AGE DETERMINATION

REPORT OF ANALYTICAL WORK

Our Sample No. R-2819

Date Received: 26 April 1974

Your Reference: SRF 5A 42'

Date Reported: 31 May 1974

Submitted by: Richard J. Holt
Weston Geophysical
P.O. Box 364
Weston, MA 02193

Sample Description & Locality: Altered granodiorite

Material Analyzed: Whole rock, crushed to -60/+100 mesh.

$\text{Ar}^{40*}/\text{K}^{40} = -0.1710$ AGE = 272 + 10 M.Y.

Argon Analyses:

Ar ⁴⁰ *, ppm.	Ar ⁴⁰ * / Total Ar ⁴⁰	Ave. Ar ⁴⁰ *, ppm.
.06782	.879	.06893
.07003	.872	

Potassium Analyses:

% K	Ave. %K	K ⁴⁰ , ppm
3.341	3.304	4.030
3.267		

Constants Used:

$$\lambda_\beta = 4.72 \times 10^{-10} / \text{year}$$

$$\lambda_e = 0.585 \times 10^{-10} / \text{year}$$

$$K^{40}/K = 1.22 \times 10^{-4} \text{ g./g.}$$

$$\text{AGE} = \frac{1}{\lambda_e + \lambda_\beta} \ln \left[\frac{\lambda_\beta + \lambda_e}{\lambda_e} \times \frac{Ar^{40}*}{K^{40}} + 1 \right]$$

Note: Ar⁴⁰* refers to radiogenic Ar⁴⁰.
M.Y. refers to millions of years.

GEOLOGICAL INVESTIGATIONS
of the
PORTSMOUTH FAULT
(Novotny - 1963)
PORTSMOUTH-HAMPTON, NEW HAMPSHIRE

PUBLIC SERVICE COMPANY OF NEW HAMPSHIRE
SEABROOK STATION

SEPTEMBER 1974

PORPSMOUTH FAULT INVESTIGATIONS

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PUBLIC SERVICE COMPANY OF NEW HAMPSHIRE

SEABROOK STATION

PORTSMOUTH FAULT INVESTIGATIONS

Investigations have been conducted along the general path of the inferred Portsmouth fault between Portsmouth and Hampton, New Hampshire, in an attempt to locate and define the inferred feature, and to examine the nature and structure of unconsolidated Pleistocene deposits which overlie bedrock in the area. (Figure 1)

All investigations have failed to locate or to suggest the existence of the Portsmouth fault. Well-stratified outwash sand deposits of Pleistocene age, as exposed in the walls of a number of gravel pits at scattered localities along the general trend of the inferred fault, show no evidence of tectonic faulting subsequent to their deposition.

Consideration of all available facts leads to the conclusion that the Portsmouth fault does not exist.

I. DEFINITION OF THE PORTSMOUTH FAULT

The Portsmouth fault was originally postulated by R. F. Novotny to trend southerly in an arcuate path for a total length of $12\frac{1}{2}$ miles from Pierce's Island, Portsmouth, to the Taylor River, Hampton, New Hampshire (Novotny; 1963). Novotny's bases for postulating the fault include: 1) brecciated and faulted rocks in the Kittery formation in an exposure on Route 1 By-pass, Portsmouth; 2) brecciated and partly silicified Kittery formation rocks exposed on the southeastern shore of Goat Island, New Castle; 3) brecciated and partly silicified Kittery formation rocks exposed near the east end of Brumley Hill, North Hampton; 4) the presence of granitic intrusives in the Rye formation near the Kittery contact; 5) an apparently unconformable stratigraphic relationship between the Rye and Kittery formations along the trend of their contact zone.

Novotny further interpreted the Portsmouth fault to form the steeply west-dipping contact between the Rye and Kittery formations. Displacement was inferred to be down on the west, suggesting a normal fault. Outcrops were reported to be too few and too poor to attempt calculation of fault displacement.

II. INVESTIGATION PROCEDURES

A. Preliminary - General Area

As a preliminary investigation of the Portsmouth fault, J. R. Rand walked portions of the fault trace as it was defined by Novotny, and examined gravel pits and highway road cuts and construction excavations in a strip about 2 miles wide overlapping the postulated trace of the fault from Portsmouth to Hampton. Each specific outcrop cited by Novotny as proof of faulting was also examined. R. J. Holt of Weston Geophysical Research, Inc. and J. R. Rand together viewed by helicopter the inferred trace of the fault between the Seabrook site and Gerrish Island, Maine. J. R. Rand also studied commercial aerial photographs covering the zone from the site to North Hampton, and his own color photographs taken along the path of the zone during the helicopter inspection. Backhoe trenching and ground magnetometer surveys have also been conducted in Greenland, New Hampshire, 8 to 9 miles north of the Seabrook site, in an effort to locate the fault (Point "A" on Figures 1 and 2). Several bedrock samples were taken along the zone for radiometric age dating.

B. Detailed - Breakfast Hill Road, Greenland

Just to the northeast of the intersection of the New Hampshire Turnpike and Breakfast Hill Road, Greenland, a wide area of outwash sands, ice-contact gravels and cobble till deposits was excavated for Turnpike construction subsequent to Novotny's field studies in the area (Point "A" on Figures 1 and 2; Figure 3; Figure 4). Within this large area, numerous low, glacially striated surfaces of Rye formation bedrock are now exposed in the floor of the reclaimed borrow area, in contradiction to Novotny's interpretation of Kittery formation terrane in this area. Survey control for investigations was provided by McKenna Associates, Portsmouth (map attached).

1. Coakley Sand Pit

As shown on Figure 3, backhoe trenching in an operating sand pit at the northwest corner of the area exposed additional outcrop of the Rye formation bedrock. Boring PF-1 was drilled on a N50W (True) bearing at an inclination of about 43° to a depth of 276', taking oriented core samples, in a search for a possible Rye/Kittery contact in an apparent folded structure which underlay well-stratified and undisturbed outwash sands exposed in the north wall of the pit.

Boring PF-1 encountered only interbedded gneiss, fine-grained schist and thin interbedded quartzites of the Rye formation, and was terminated as it passed to the west of the edge of the sand pit. The structure of the Rye formation in the boring, as indicated by orientation measurements of bedrock foliation, is that of a tight syncline which dips steeply to the west. Five zones of welded breccia were encountered in the boring, the thickest of which included 7.5' of welded quartzite breccia at 249.5' to 257' depths in the hole. The brecciated rock in PF-1 was fresh, compact, thoroughly welded or annealed, and did not show polished or slickensided surfaces on partings.

No mineralization, hydrothermal alteration, shear zones, or other evidence of major faulting was encountered in the boring. The welded brecciation is of the type found frequently in borings in metamorphic rocks in the region, and is interpreted to be associated with strains developed at the time of folding and metamorphism of the region during the Acadian orogeny. Two diabase dikes encountered in the boring were fresh, unaltered, and showed normal intrusive contacts.

2. Loch-Coombs Reclaimed Borrow Area

As shown on Figure 4, three core borings (PF-2, PF-3, PF-3A) were drilled across the property line between lands of Anthony Loch and Richard Coombs, at the north edge of a reclaimed borrow area to the north of Breakfast Hill Road, to investigate the western boundary of a local magnetic anomaly.

a. Ground Magnetometer Survey

Because the bedrock exposed throughout the Breakfast Hill study area is represented only by Rye formation metavolcanic rocks for as much as one-half mile to the west of Novotny's fault trace, and comprises no outcrops of Kittery formation quartzites as had been interpreted by Novotny, the presence of a fault contact between these two formations in this area cannot, by definition, exist. Having no formal contact to investigate for these current studies, Weston Geophysical Engineers, Inc. undertook a ground magnetometer survey to determine whether any anomalous magnetic features might occur which could suggest faulting within the Rye formation itself. Technical details of this survey are presented in a report by Weston Geophysical Engineers, Inc., attached herewith.

The magnetometer surveys show no anomalous magnetic intensities in the zone of Novotny's fault trace in five profiles which were conducted across the inferred trace at intervals influencing a zone of almost 4,000' along the trace from north to south. In the area of Coombs Pond (Figure 4), a local magnetic anomaly high was detected on 3 survey lines (Lines 6, 2NR and 2R). The apparent alignment of this anomaly is about N10E, parallel to the strike of bedrock foliation in the area. Novotny's inferred fault trace in the same general area strikes about N40E, transverse to foliation.

b. Borings Investigations

Two borings, PF-2 and PF-3A, were drilled at approximately 40° inclination to the southeast to investigate bedrock conditions at the western boundary of the local magnetic anomaly. A third boring, PF-3, was drilled vertically to determine bedrock depth prior to drilling PF-3A. The results of these borings are generalized in cross section on Figure 4, on which also is projected the magnetic profile of Mag. Line 6.

Overburden, which was not specifically sampled in these three borings, is comprised of outwash sands overlying a sandy boulder till. Boring PF-2 was drilled to a depth of 271' (about 201' in bedrock), in light gray banded gneiss and dark green amphibolite, intruded locally by weakly magnetic diabase dikes. PF-3 was drilled to a depth of 50' (10' in bedrock) in gray and greenish gneiss. PF-3A was drilled to a depth of 204.3' (124' in bedrock) in gray banded gneiss, dark green amphibolite and, at the bottom 5' of the boring, notably magnetic, salmon-feldspar gneiss, with a single diabase dike. The location of the basal magnetic gneiss in PF-3A conforms reasonably with the downward projection on the local bedrock structure of the magnetic anomaly found by surface surveys. The weakly magnetic dikes in the borings conform with a slight increase in magnetic intensity found by surface surveys.

The condition of bedrock in PF-2, PF-3 and PF-3A was weakened by weathering effects on moderately closely-spaced jointing to about -70' Elevation. In no boring, however, were there slickensided or polished joint surfaces, gouge zones, hydrothermal alteration or any other visible evidence of bedrock faulting.

C. Petrographic Examinations

The petrography of three samples of drill core from Boring PF-2 has been described by Professor Gene Simmons and Dorothy A. Richter of Massachusetts Institute of Technology.

<u>Sample</u>	<u>Depth</u>	<u>Field Description</u>	<u>Petrographic Description</u>
PF-2A	99.5 - 99.9'	Gneiss	Felsic Metatuff
PF-2B	136.0 - 136.5'	Diabase	Metabasalt
PF-2C	262.0 - 262.4'	Amphibolite	Fine Grained Amphibolite

Simmons and Richter conclude from their studies that "Evidence for dynamic structural deformation, either recent or ancient, is entirely absent. In summary, we find no petrographic evidence that these three samples are associated with a fault. If a fault does exist in the region from which these samples were obtained, then either its deformation was not so pervasive as to effect these three samples, or else the deformation occurred before metamorphism and all petrographic evidence has been erased by the last metamorphic event".

The full report by Simmons and Richter is attached herewith.

III. RESULTS OF INVESTIGATIONS ALONG THE INFERRRED FAULT

None of the current investigations along the path of the inferred Portsmouth fault has detected or suggested the presence of a through-going fault structure along the zone of the Rye/Kittery contact between Portsmouth and Hampton. No exposure of Pleistocene deposits seen along this zone has shown internal structures suggestive of tectonic fault displacement.

A. Novotny's "Faulted" Outcrop Exposures

1. Route 1 By-pass, Portsmouth (Point "B" on Figures 1 and 2)

Novotny cites a road cut on the north side of the Route 1 By-pass in Portsmouth as suggesting the presence of the Portsmouth fault nearby, but not within, the road cut exposure. This exposure shows two steeply west-dipping zones of weathered and rusty rock material interlayered in gneiss and quartzite. In one of these weathered zones, an open drag fold was interpreted by Novotny to represent differential movement, down on the west. This folding could also represent simple folding of the beds, signifying an anticline to the west.

The rock materials within these two weathered zones are not slicken-sided or mineralized, and the rock adjacent to the weathered zones shows no hydrothermal alteration. Very similar open folding can be seen in an unweathered exposure of quartzitic rock near the Rye/Kittery contact, 3.54 miles S52W of this locality, on the west right-of-way of the New Hampshire Turnpike, where there is no evidence of fracturing. Fold structures of the type

seen at the Route 1 By-pass and on the New Hampshire Turnpike right-of-way are most logically explained as simple small-scale drag folding formed during the regional folding of the Rye anticline. The exposure on the Route 1 By-pass is suggestive of faulting only because it is weathered. Rye formation rocks occur on both sides of the weathered zone at the Route 1 exposure.

2. Goat Island, New Castle (Point "C" on Figures 1 and 2)

Bedrock structure on the southeast shore of Goat Island is a complex jumble of brecciated Rye formation metavolcanics and quartzite. The breccia is welded, and is intruded by diabase dikes. No "trend" of faulting is apparent at this locality to suggest a through-going fault plane which might connect this exposure with the exposure cited on the Route 1 By-pass, 2.1 miles to the southwest. The apparently random distribution of metavolcanics and quartzite breccia blocks suggests that fault structure in this area may represent explosion breccia, which Hussey (1962) has also found as discontinuous masses 3 miles to the east on Gerrish Island, Maine. Hussey suggests that the breccia at Gerrish Island may relate to volcanic activity associated with the Cape Neddick and Tatnic volcanic complexes, southwestern Maine.

3. Brumley Hill, North Hampton (Point "F" on Figure 1)

The brecciated quartzite cited by Novotny for the east end of Brumley Hill showed some healed fracturing and rusty staining in a dark, fine-grained quartzite. Billings (1956) interpreted this area to lie in a broad fold zone in the Rye formation. No through-going shears were apparent in the exposure to suggest the presence of faulting. The exposure no longer exists, having been removed during construction of a new north-bound lane of the New Hampshire Turnpike.

B. Granite in the Rye Formation

Novotny states (1963; p. 147): "Although metamorphic zones are apparently not displaced because of the fault, the presence of concordant foliated and granulated Breakfast Hill granite only in the Rye formation and near the Kittery formation contact supports the hypothesis of a fault developed during the Acadian period of orogeny, along which deeply buried and intruded portions of the Rye formation were elevated". (Point "D" on Figures 1 and 2)

Foliated granite, seen in a number of places in the Rye formation, appears to be a primary metamorphic constituent of that formation, having formed by recrystallization ("granitization") of the inherently feldspathic Rye formation rocks. These granitic masses appear genetically related to a process of metamorphism within the Rye, rather than to plutonic intrusions from a separate deep-seated source. Because of the fundamental lack of feldspar in the Kittery formation, furthermore, no comparable granitization of the Kittery could have occurred at the time the Rye was being recrystallized and granitized.

Whereas the granites of the Rye formation to the east of the Rye/Kittery contact do not in themselves offer any proof that the Rye has been elevated relative to the Kittery, plutonic intrusives of the Exeter diorite are found in the Kittery formation to the west of the Rye/Kittery contact, tending to negate an hypothesis of fault displacement based on the presence or absence of igneous rocks in the metamorphic terrane. (Point "E" on Figures 1 and 2)

C. Unconformable Rye/Kittery Stratigraphy

Whereas Novotny interpreted an unconformable stratigraphic relationship between the Rye and Kittery formations in the area between Portsmouth and Hampton, outcrops of the two formations are widely scattered, and the contact between these formations is nowhere exposed along the 12½ mile path of the inferred Portsmouth fault. On Gerrish Island, Maine, about 5 miles east of Portsmouth, Hussey (1962) interprets the Rye/Kittery contact to be conformable, grading upward through progressively less feldspathic gneisses of the Rye formation into biotite quartzites typical of the Kittery.

Novotny, Hussey and Billings (1956) all define the Rye formation as metavolcanic and the Kittery as metasedimentary, predominantly quartzite. Novotny interprets the contact between these two formations to be defined by a major fault structure, while Hussey and Billings do not. Novotny, furthermore, defines the geographic location of the Rye/Kittery contact as much as three-quarters of a mile to the east of the contact trace defined by Hussey and Billings. Figure 1 shows by a dotted line the contact between the Rye metavolcanic member and the Kittery formation as defined by Billings to the southwest and by Hussey to the northeast.

Current investigations have indicated that Novotny's contact trace trends from Portsmouth to North Hampton through a terrane characterized only by bedrock exposures of the Rye formation metavolcanic member. Since the metavolcanic member of the Rye is made up of an original sequence of different types of volcanic rocks and interbedded sedimentary units, unconformable stratigraphic relationships might be expected in the zone where Novotny has defined the Rye/Kittery contact. Such relationships would not, however, signify the presence of a major fault zone. Furthermore, foliation structure symbols shown on Figure 1 (after Novotny and Hussey; and J. R. Rand reconnaissance) indicate a reasonable parallelism of bedrock structure along Novotny's inferred fault trace in this area, with no suggestion of the alledged formational unconformity.

D. Radiometric Age Dating

Four outcrop samples (PF-S1, -S2, -S3, -S4) were taken at intervals along the path of the inferred fault for radiometric age dating (K-Ar). The locations and K-Ar ages of these samples, along with three other samples taken from Borings B2, B4 and B9 at the site area in 1969, are defined on Figure 1. Age determinations were obtained by Geochron Laboratories, Division of Krueger Enterprises, Inc., Cambridge, Massachusetts.

<u>Sample</u>	<u>Location</u>	<u>Rock Type</u>	<u>Material</u>	<u>K-Ar Age</u>
PF-S1	Towle Road, Hampton	Quartzite	Biotite	268±10 M.Y.
PF-S2	Rte. 151, North Hampton	Quartzite	Amphibole	308±14 M.Y.
PF-S3	Rte. 1, Portsmouth	Gneiss	Muscovite	294±10 M.Y.
PF-S4	Rte. 1, Portsmouth	Quartzite	Mica-Quartz	262±11 M.Y.
B2	129.5' - Boring B2	Qtz. Diorite	Biotite	294± 9 M.Y.
B4	93.0' - Boring B4	Schist	Biotite	254± 9 M.Y.
B9	12.3' - Boring B9	Bio. Diorite	Biotite	284± 9 M.Y.

No anomalously young ages were found in this dating program. All ages found conform to previously reported regional data which indicates a Permian thermal event for the area (Zartman et al, 1970). The lower ages obtained in this investigation (PF-S1, PF-S4 and B4) are mineral dependent, with argon loss associated with the fine-grained materials analyzed.

IV. CONCLUSIONS

Field investigations have shown that

1. The graphic trace of the alleged Portsmouth fault bears no meaningful spatial relationship to the contact between the Rye and Kittery formations, along which the fault was postulated by Novotny to trend.
2. There is no evidence of the alleged unconformable relationship between the Rye and Kittery formations.
3. There is no evidence of anomalous magnetic intensities on the inferred fault trace in Greenland, New Hampshire.
4. Examination of drill cores in the area of the alleged fault trace in Greenland, complimented by petrographic studies of core samples, indicate no evidence of faulting in that area.
5. There is no evidence of a through-going fault structure associated with the specific bedrock exposures cited by Novotny as indicating the presence of the Portsmouth fault.
6. There is no justification for ascribing the presence of granitic rocks at ground surface in the Rye formation terrane to the differential uplift of these rocks along a nearby fault.
7. There are no meaningful variations in radiometric ages of rocks along the alleged fault trace.
8. Ground and aerial examinations have failed to detect any anomalous landforms or stream patterns along the trace of the alleged fault.
9. Pleistocene deposits exposed in road cuts and gravel pits along the alleged fault trace show no features which might imply tectonic faulting in the area.

The current investigations have concluded that the Portsmouth fault does not exist.

John R. Rand
Consulting Geologist

September 1974

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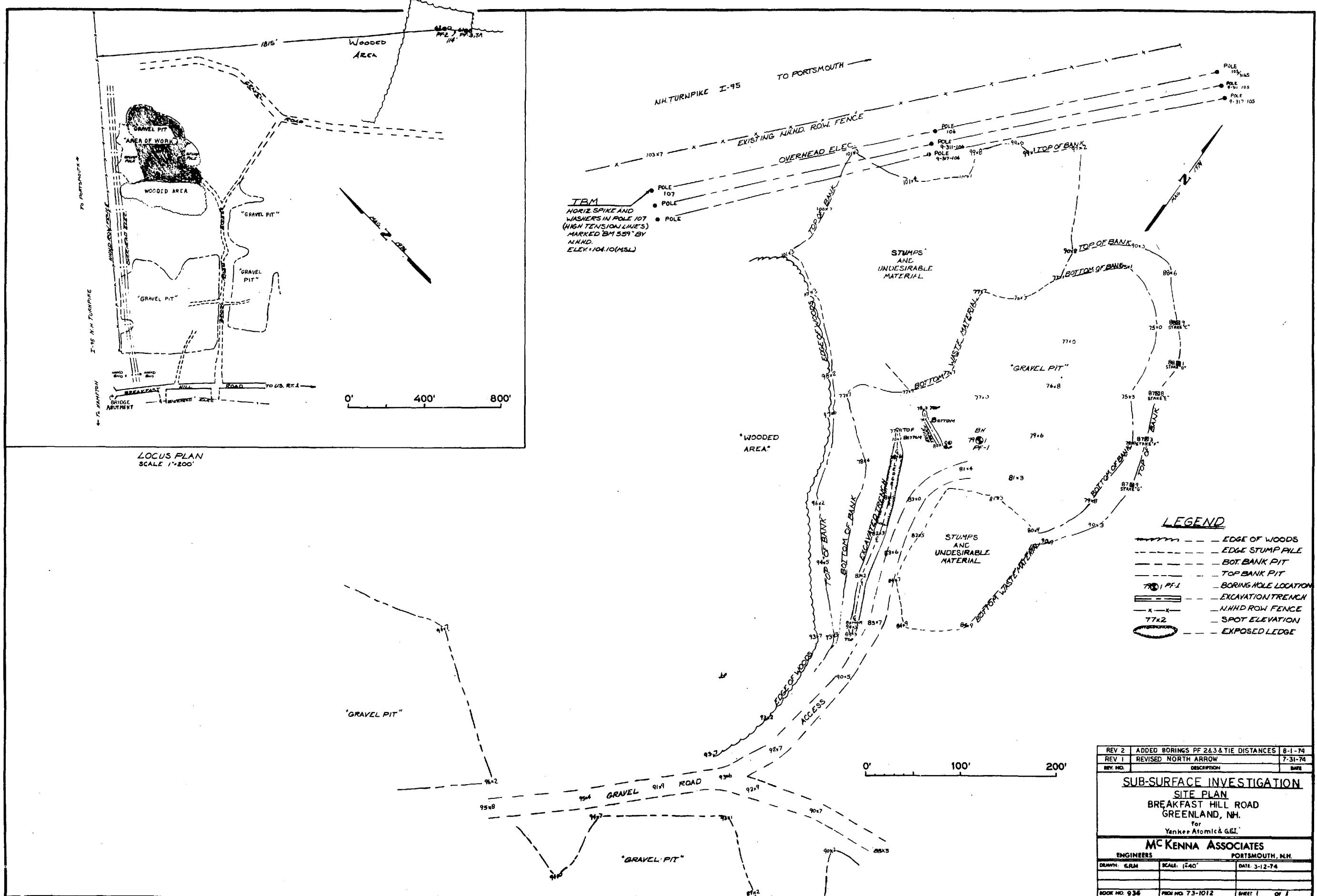
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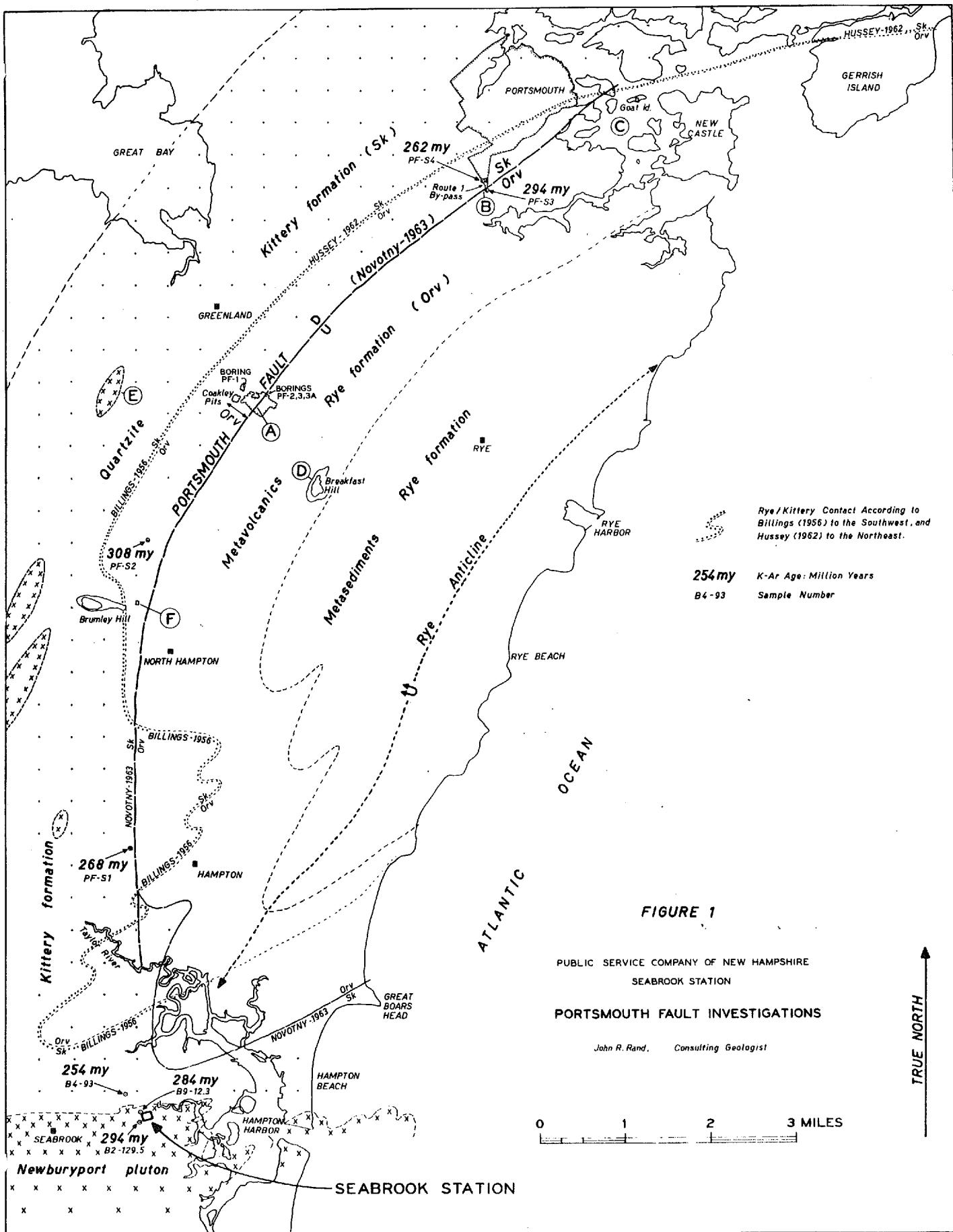
Hussey, A. J. II (1962) The Geology of Southern York County, Maine. Special Geologic Studies, No. 4, Maine Geological Survey; Augusta, Maine.

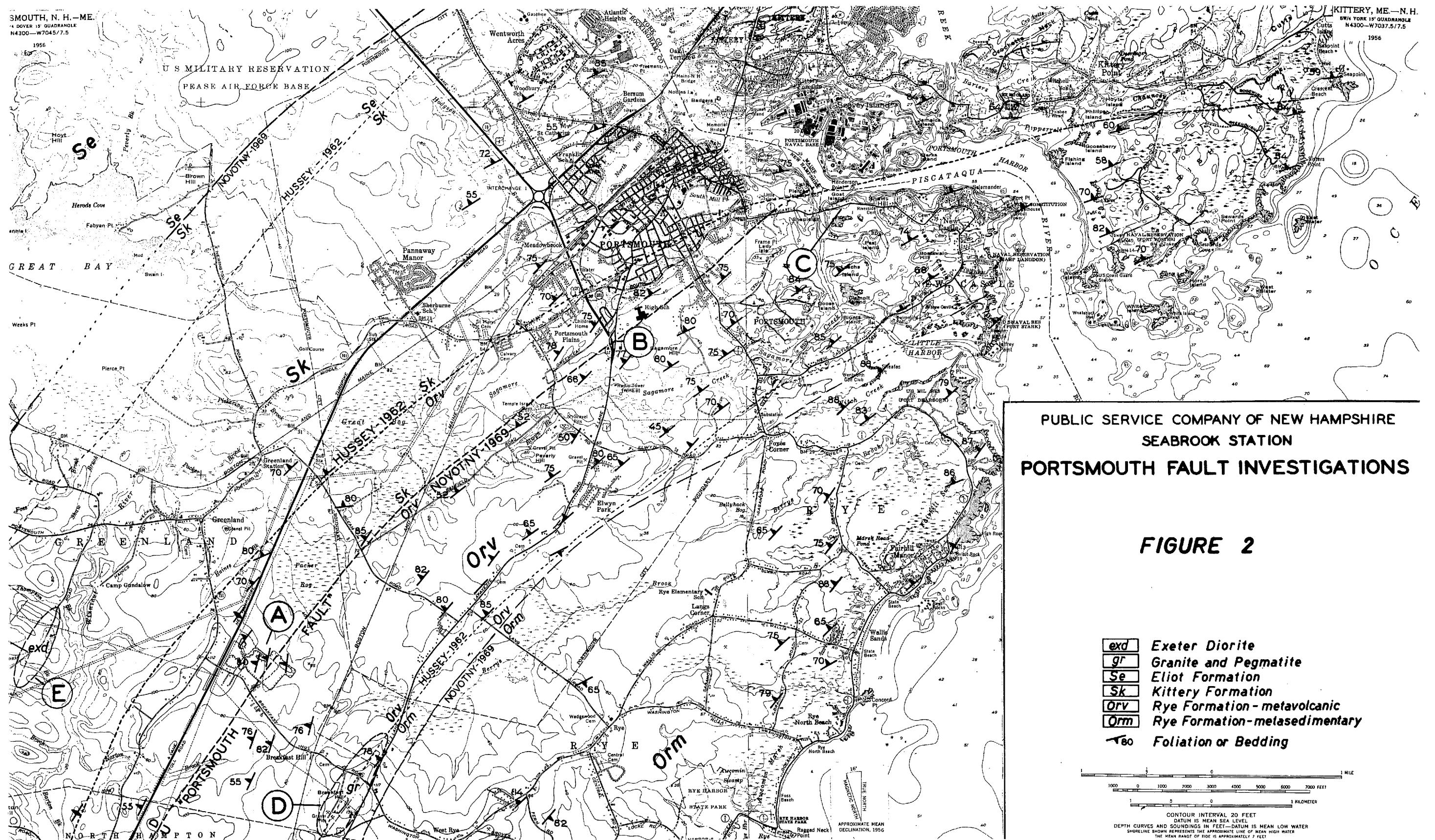
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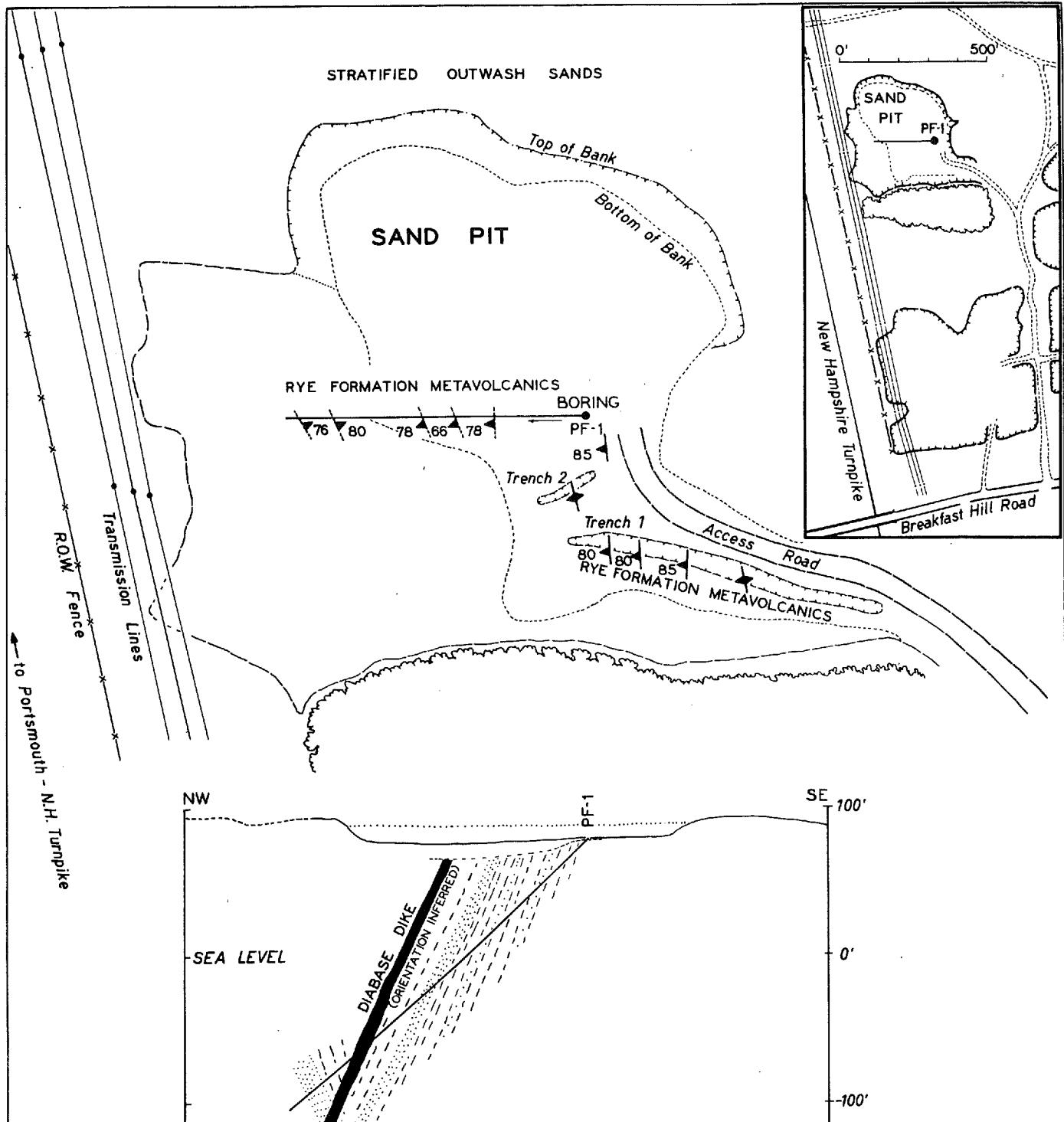


FIGURE 3

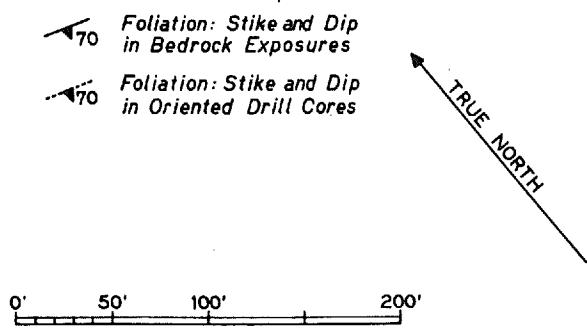
PUBLIC SERVICE COMPANY OF NEW HAMPSHIRE
SEABROOK STATION

PORTSMOUTH FAULT INVESTIGATIONS

COAKLEY SAND PIT

BREAKFAST HILL ROAD GREENLAND, NEW HAMPSHIRE

John R. Rand, Consulting Geologist



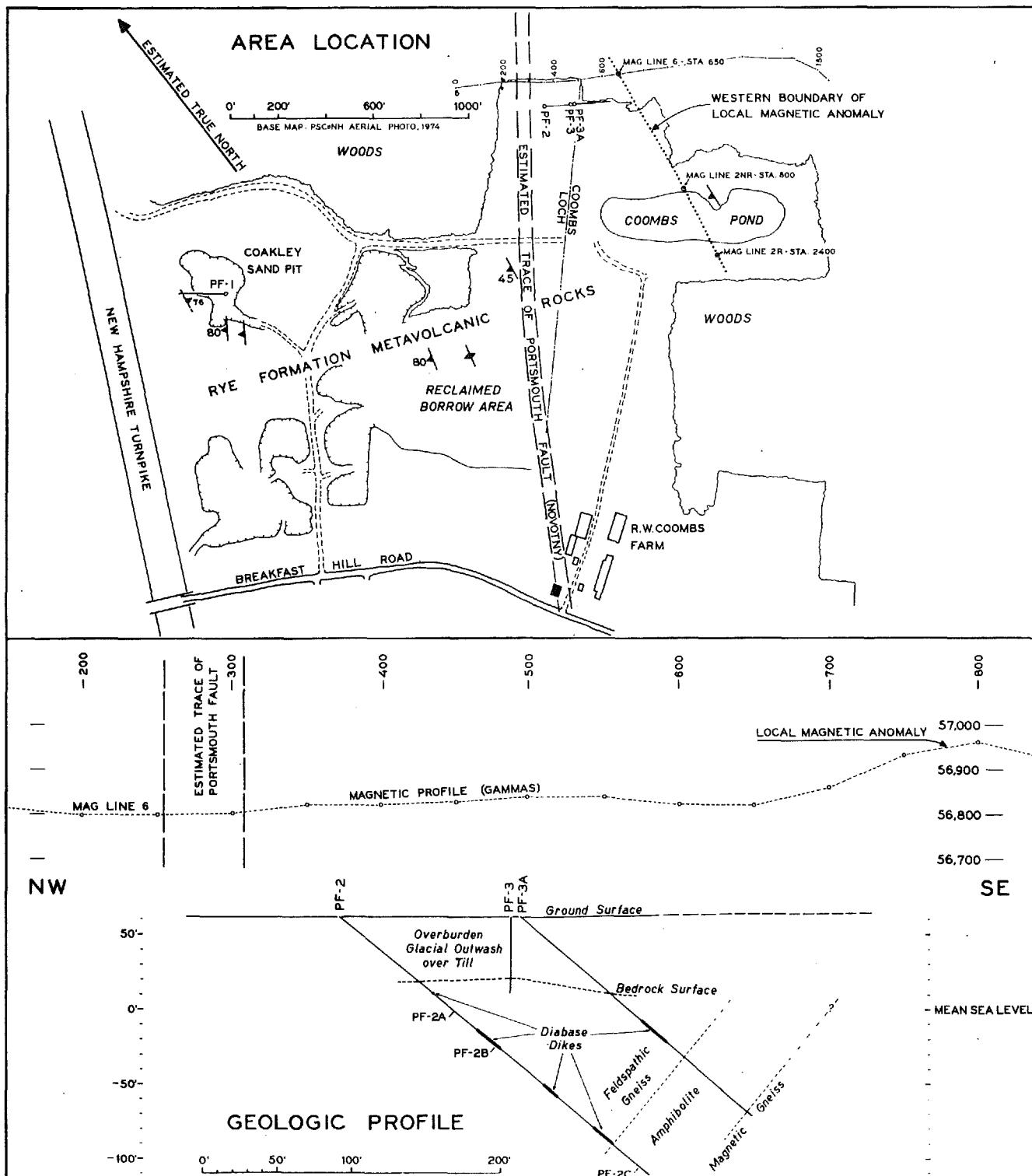


FIGURE 4

PUBLIC SERVICE COMPANY OF NEW HAMPSHIRE
SEABROOK STATION

PORTSMOUTH FAULT INVESTIGATIONS

COOMBS POND AREA

BREAKFAST HILL ROAD - GREENLAND, NEW HAMPSHIRE

John R. Rand, Consulting Geologist

ATTACHMENT NO. 1

GROUND MAGNETOMETER SURVEY
BREAKFAST HILL ROAD AREA
GREENLAND, NEW HAMPSHIRE

WESTON GEOPHYSICAL RESEARCH, INC.
WESTBORO, MASSACHUSETTS

GROUND MAGNETOMETER SURVEY

BREAKFAST HILL ROAD AREA

GREENLAND, NEW HAMPSHIRE

SUMMARY

This report details a ground magnetometer survey conducted by Weston Geophysical Research, Inc. in the vicinity of Breakfast Hill Road, Greenland, New Hampshire. This study was completed in conjunction with a general geologic investigation of the inferred Portsmouth fault, as proposed by Novotny (1963).

Five separate magnetic lines were run across the trace of the inferred fault. No magnetic evidence for faulting was found on any of the profiles.

INSTRUMENTATION

The survey was begun with a vertical field, torsion magnetometer (Askania, Model Gfz), which is tripod mounted and must be leveled prior to each reading. Because this procedure is difficult in soft or swampy ground, which is extensive in the investigation area, the vertical field magnetometer was replaced with a total field, proton precession magnetometer (Geometrics, Model G-816), which requires neither tripod nor leveling.

METHOD

The survey method consisted of making total magnetic field intensity measurements at paced intervals along a predetermined line. The interval used varied from 50 to 100 feet. The magnetic sensor was oriented north (magnetic) for each reading, and readings were repeated to insure precision. A base station was established, and base station readings were taken regularly to determine the diurnal variation of the earth's magnetic field during a given portion of the survey. The diurnal variation has been removed from the final profiles. Careful notes were taken during the survey so that the presence of magnetic interference sources (i. e., power lines, buried metal, houses, parked vehicles, etc.) could be considered in the final analysis.

RESULTS

Total field intensity magnetic profiles were made from data for five traverses in the area of investigation. As shown in Figure Ala, Profiles 2R, 2NR and 6 are located at distances extending up to approximately 2,500 feet northeast of Breakfast Hill Road, near Coombs Pond. Profiles 4 and 5 are located at distances extending up to approximately 1,500 feet southwest of Breakfast Hill Road. All five magnetic profiles crossed Novotny's inferred fault trace at nearly perpendicular angles. Any magnetic expression of Novotny's inferred fault (within the Rye formation) should, therefore, have been readily apparent.

Figure A1a locates the inferred fault trace relative to magnetic profiles reported in Figure A1b at or near the following profile points: 3+0 on Line 6, 1+5 on Line 2NR, 16+0 on Line 2R, 15+0 on Line 4, and 13+0 on Line 5.

No evidence of the postulated fault was found. Further examination of the profiles indicates that localized anomalies, probably due to local variations in magnetic mineral concentrations known to be present in the Rye formation, appear on each of the traverses near Coombs Pond. Profiles 6, 2NR and 2R show such an anomalous condition, which appears to trend N10E in the vicinity of the three lines. It should be noted that this strike is parallel to the bedrock foliation of the area.

MAGNETIC PROFILES
FIGURE A1b

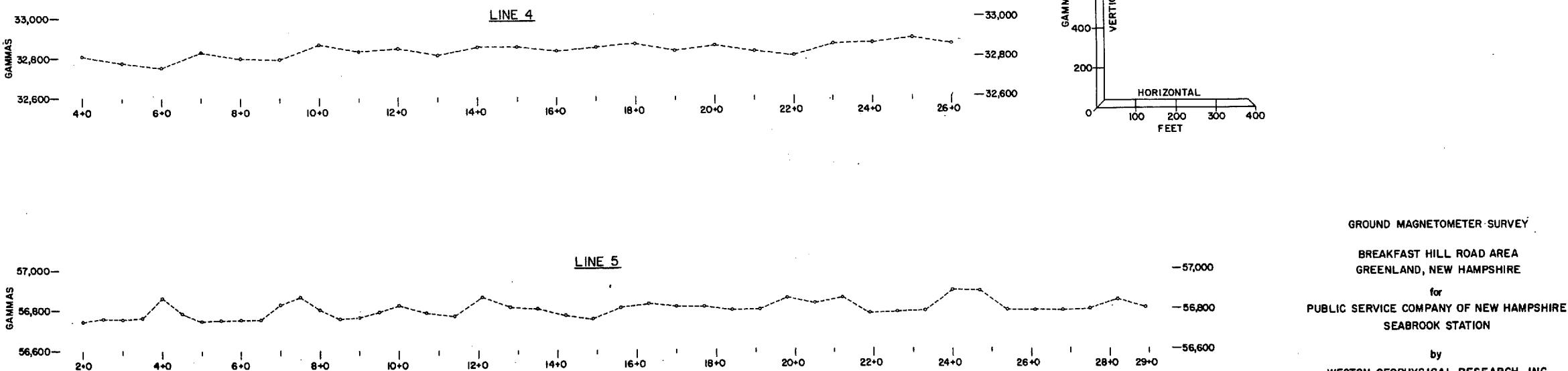
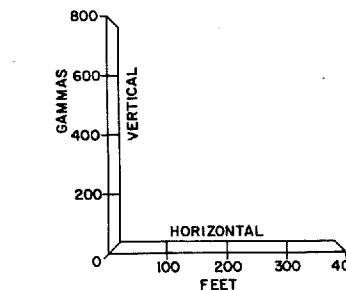
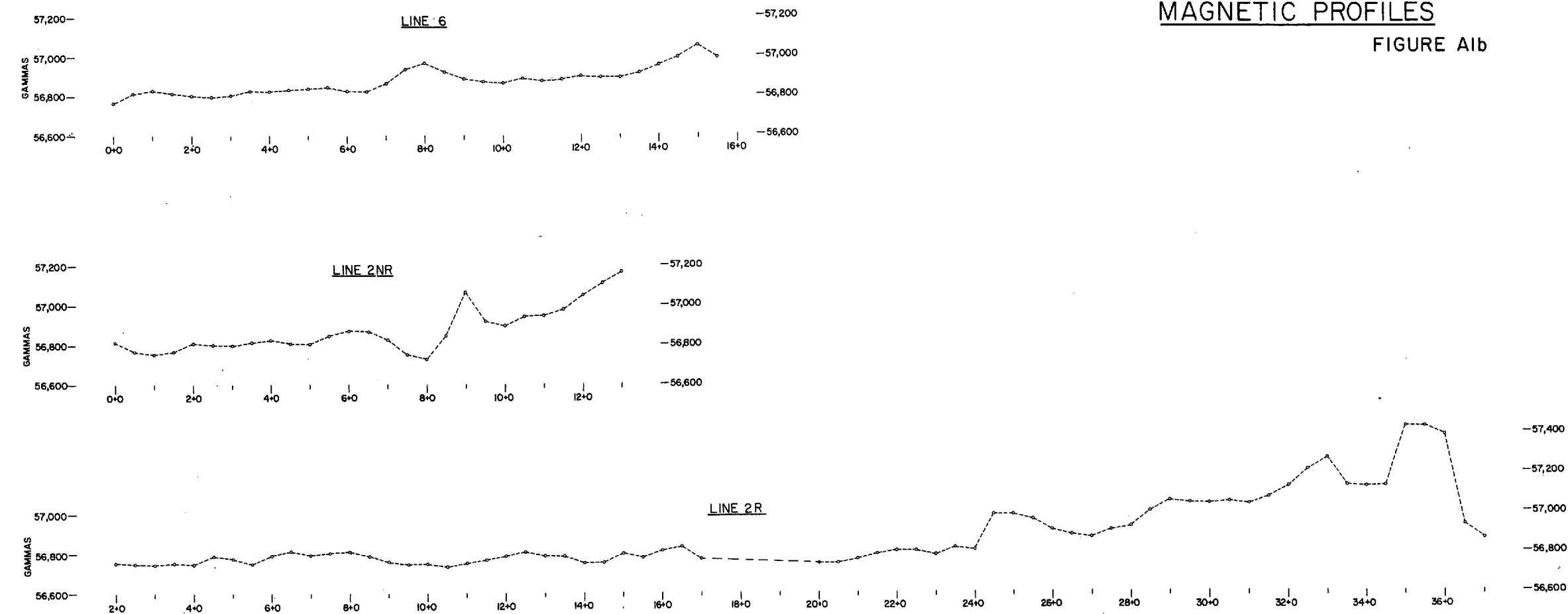
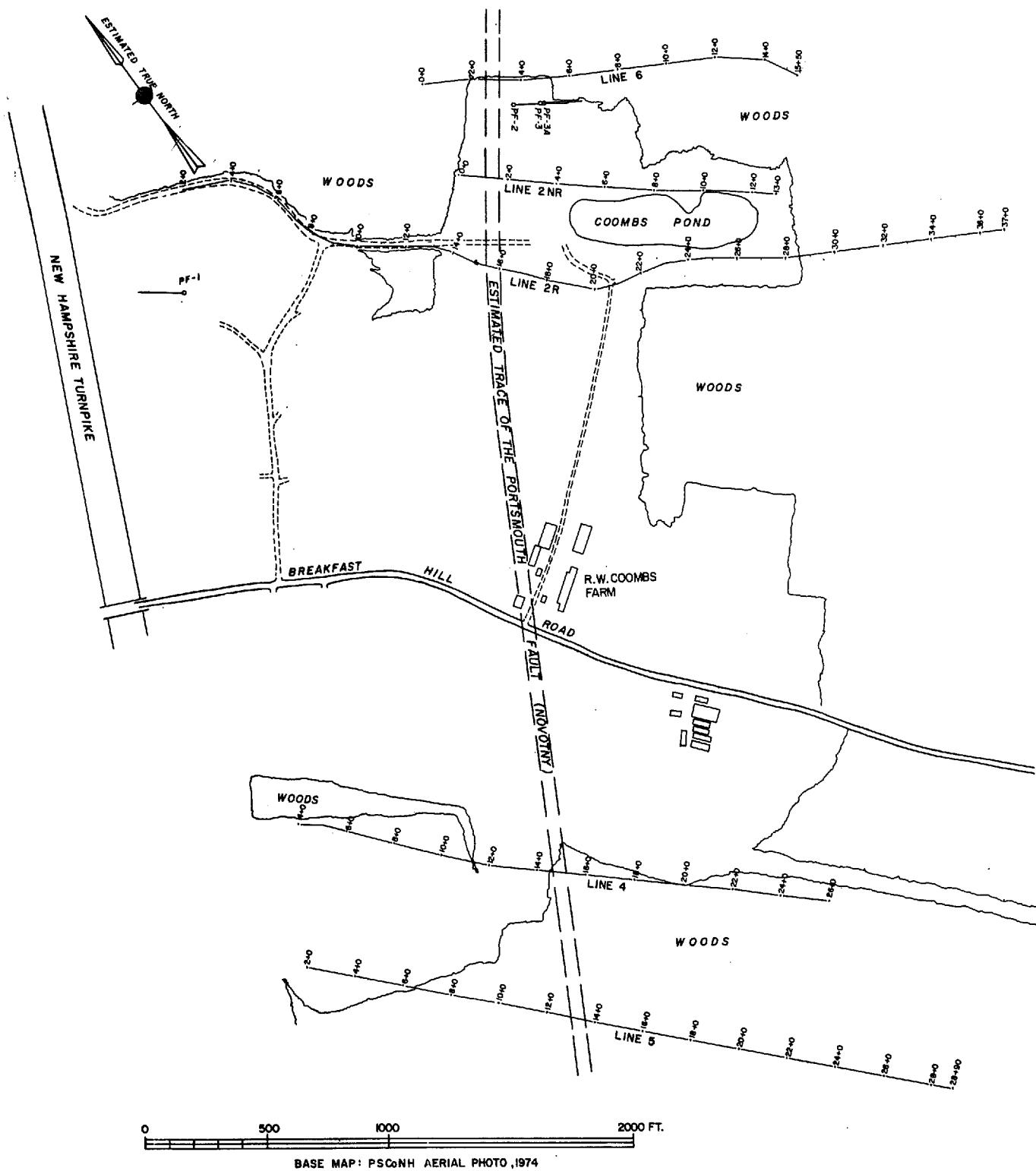


FIGURE A1



AREA LOCATION
FIGURE A1a

ATTACHMENT NO. 2

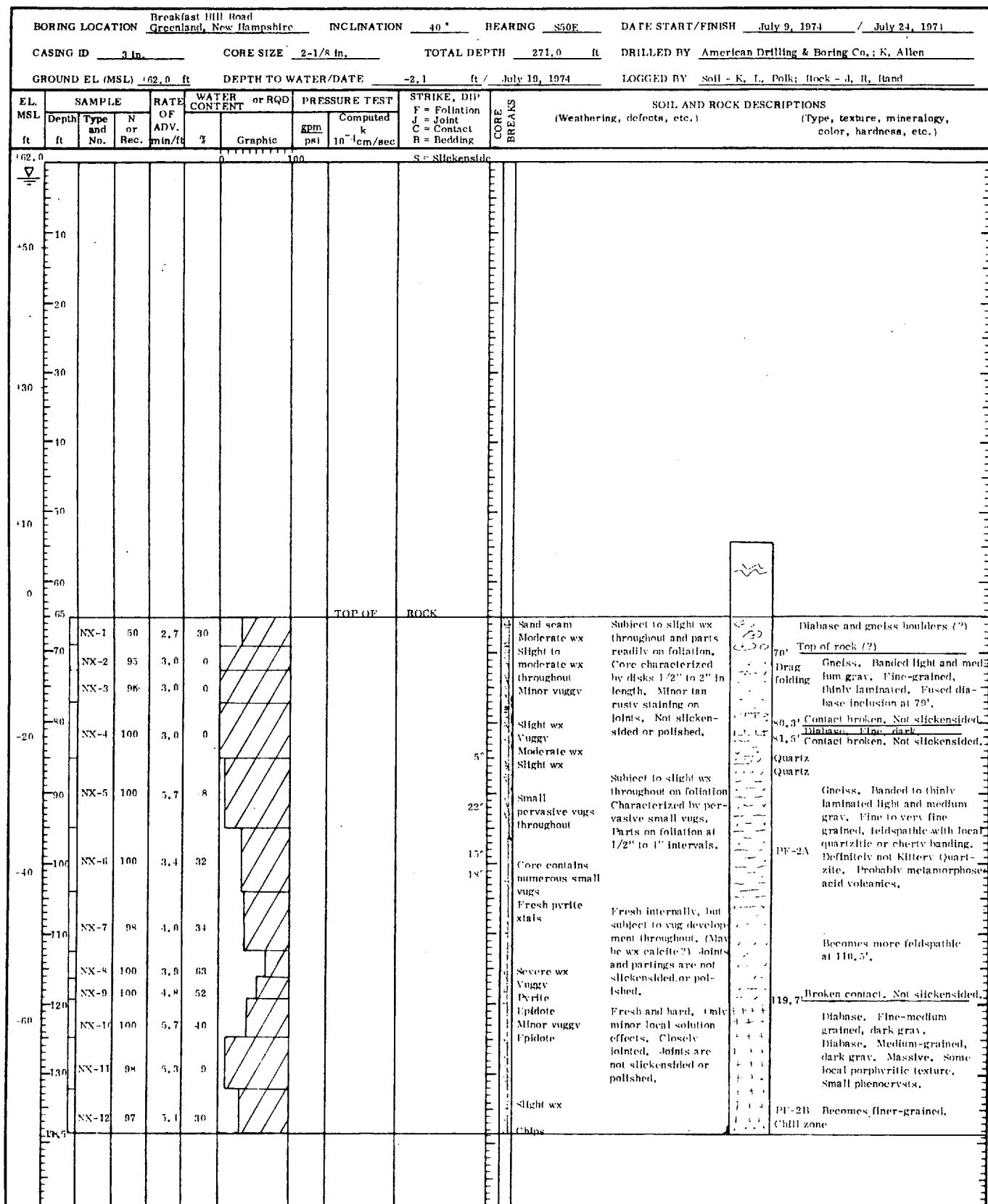
GEOLOGIC BORINGS LOGS

BORINGS PF-1, PF-2, PF-3, PF-3A

LEGEND	N - Standard penetration resistance, blows/ft Rec - Length recovered/length cored, % RQD - Length of sound core 4 in. and longer/length cored, % S - Split spoon sample ▽ Groundwater U - Undisturbed samples	NOTES	SEABROOK STATION PUBLIC SERVICE COMPANY OF NEW HAMPSHIRE YANKEE ATOMIC ELECTRIC COMPANY  United engineers Constructors Inc. A subsidiary of Bechtel Company
	S - Shelby tube N - Denison F - Fixed piston P - Pitcher O - Osterberg G - GEI D - Drilling break k - Coefficient of wx - Weathered, weathering permeability		Date: May 14, 1974. Project 7286 PAGE <u>2</u> of <u>2</u> LOG OF BORING <u>PF 1</u>



ORTHOCLINICAL ENGINEERS INC





N	- Standard penetration resistance, blows/ft
Rec	- Length recovered/length cored, %
RQD	- Length of sound core 4 in. and longer/length cored, %
S	- Split spoon sample ∇ Groundwater
U	- Undisturbed samples
	S - Shelby tube N - Denison
	F - Fixed piston P - Pitcher
	O - Osterberg G - GEI
D	- Drilling break k - Coefficient of
wx	- Weathered, weathering permeability

NOTES

- 1) Washed through soil 0-80 ft, no samples taken.
- 2) No drill times available.

SEABROOK STATION
PUBLIC SERVICE COMPANY OF NEW HAMPSHIRE
YANKEE ATOMIC ELECTRIC COMPANY

UNITED ENGINEERS & CONSTRUCTORS INC.

Date: August 15, 1974

Project 7286

PAGE 1 of 2

LOG OF BORING PE-3A

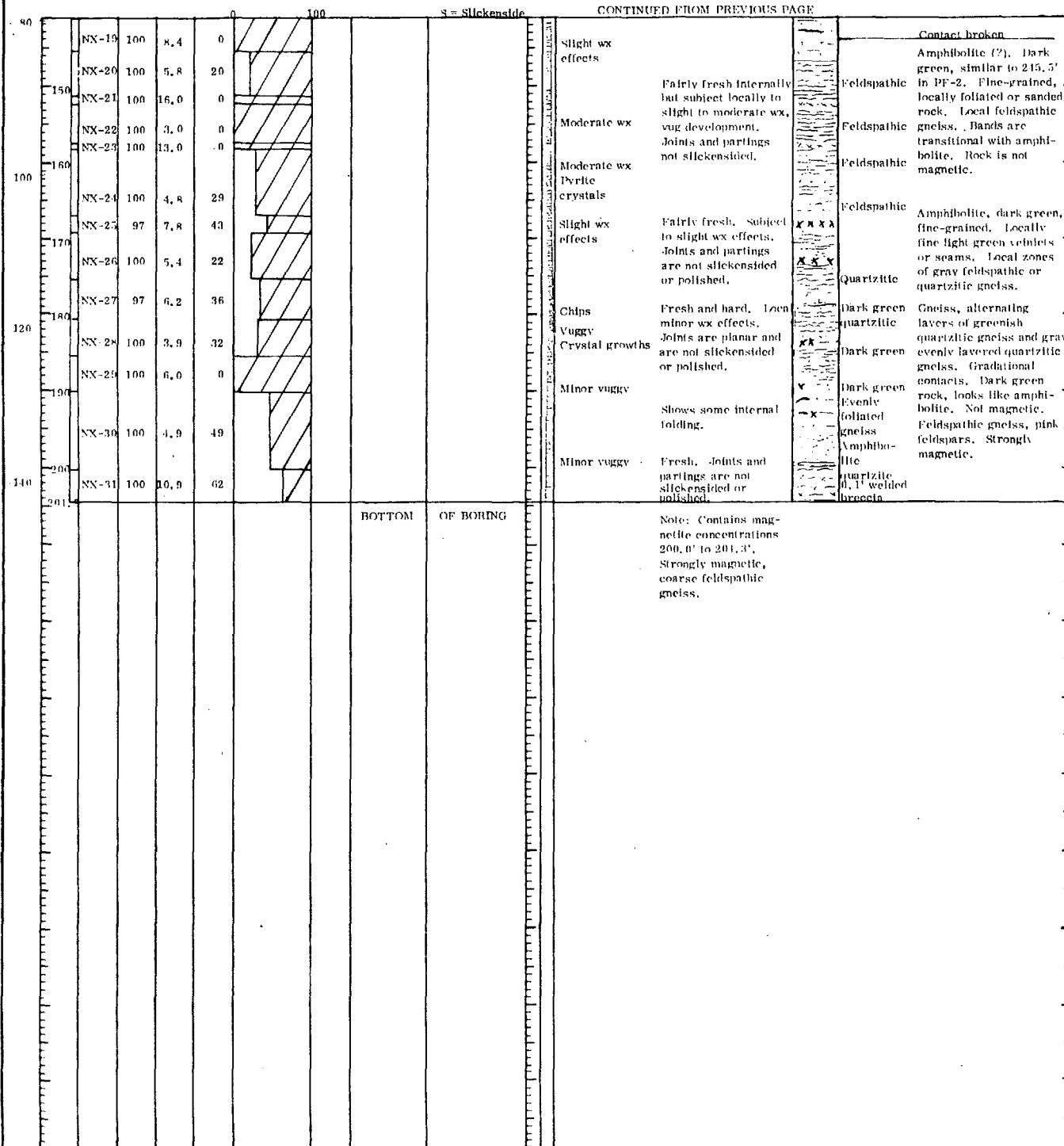


BORING LOCATION Breakfast Hill Road
Greenland, New Hampshire INCLINATION 41° BEARING 550° DATE START/FINISH July 30, 1974 / August 8, 1974

CASING ID 3 in. CORE SIZE 2-1/8 in. TOTAL DEPTH 204.3 ft DRILLED BY American Drilling and Boring Co.; K. Allen

GROUND EL (MSL) 101.8 ft DEPTH TO WATER/DATE -2.5 ft / July 30, 1974 LOGGED BY Soil - K. L. Polk; Rock - J. R. Rand

EL. ft MSL	SAMPLE Depth ft	Type and No. Rec.	N or R	WATER CONTENT OF ADV. min/ft	RATE OF ADV. gpm psi	PRESSURE TEST Graphic	COMPUTED K 10 ⁻⁴ cm/sec	STRIKE, DIP F = Foliation J = Joint C = Contact B = Bedding	CORE BREAKS	SOIL AND ROCK DESCRIPTIONS (Weathering, defects, etc.)		(Type, texture, mineralogy, color, hardness, etc.)
										(Weathering, defects, etc.)	(Type, texture, mineralogy, color, hardness, etc.)	



LEGEND	N - Standard penetration resistance, blows/ft	NOTES	SEABROOK STATION	
	RRC - Length recovered/length cored, %		PUBLIC SERVICE COMPANY OF NEW HAMPSHIRE	
RQD - Length of sound core 4 in. and longer/length cored, %	S - Split spoon sample		YANKEE ATOMIC ELECTRIC COMPANY	
U - Undisturbed samples	W - Groundwater		 united engineers	
S - Shelby tube	N - Denison		Date: August 15, 1974	Project 7286
F - Fixed piston	P - Pitcher			
O - Osterberg	G - GEI			
D - Drilling break	k - Coefficient of			
wx - Weathering	permeability			
		PAGE 2 of 2	LOG OF BORING PF-3A	

ATTACHMENT NO. 3

PETROGRAPHY AND PRELIMINARY INTERPRETATION
OF THREE SAMPLES OF DRILL CORE
FROM THE PORTSMOUTH FAULT
GREENLAND, NEW HAMPSHIRE

Gene Simmons
Dorothy A. Richter

MASSACHUSETTS INSTITUTE OF TECHNOLOGY
CAMBRIDGE, MASSACHUSETTS

for

WESTON GEOPHYSICAL RESEARCH, INC.
WESTBORO, MASSACHUSETTS

PETROGRAPHY AND PRELIMINARY INTERPRETATION
OF THREE SAMPLES OF DRILL CORE
FROM THE PORTSMOUTH FAULT,
GREENLAND, NEW HAMPSHIRE

Weston Geophysical Research, Inc.
Post Office Box 306
Weston, Massachusetts 02193

Gene Simmons
Dorothy Richter
26 August 1974

SUMMARY

In this report we describe three samples of drill core from the vicinity of the alleged Portsmouth fault near Greenland, New Hampshire. The three samples are metamorphic rocks. The pronounced laminations in sample PF-2A appear to be of primary depositional origin rather than of tectonic origin. The peculiar arcuate structures common to both samples PF-2A and PF-2C are reminiscent of glass shards, which suggests that the rocks are totally recrystallized meta-tuffs or reworked volcanic detritus of different compositions. Sample PF-2B is a partially recrystallized basalt which is probably younger than the other two samples. All three samples lack substantial preferred orientation of their minerals. Evidence for dynamic structural deformation, either recent or ancient, is entirely absent. In summary, we find no petrographic evidence that these three samples are associated with a fault. If a fault does exist in the region from which these samples were obtained, then either its deformation was not so pervasive as to affect these three samples, or else the deformation occurred before metamorphism and all petrographic evidence has been erased by the last metamorphic event.

PETROGRAPHY OF SAMPLE PF-2A 99.5-99.9'

Name: Felsic metatuff

Macroscopic Description

This sample is a finely laminated schist. It is light grey in color and fine grained. Layers of alternating light and dark colors are probably due to segregation of mineral phases. Euhedral crystals of pyrite (~ 1/2 mm) are abundant. The texture is punctuated by light colored augen and irregular 0.5 mm pores. This 5 inch core shows no veins, folds, and only a few large cracks.

Microscopic Description

Texture

The average grain size is less than 0.05 mm. Laminations are the product of the effect of variations in grain size, in the proportions of quartz to mica, and the abundance of opaques. Micas tend to show a preferred orientation of flakes at an angle of about 60° to the laminae. Most of the veinlet-like seams of quartz follow the foliation; although a few seams cross-cut the foliation they are not common and their margins have recrystallized to blend with the rest of the rock. The augen are pods of fine grained quartz. Some of the pores have minor amounts of weathering products around the rims.

A few large microcracks that are now completely healed were observed in the thin section. They are marked by

chlorite, quartz, and trains of discrete grains of opaques (probably pyrite). However, there is no other textural evidence of penetrative deformation. The thin laminations and indications of flow structures imply that this sample is a recrystallized silicious tuff or reworked volcanic detritus.

Mineralogy

Quartz is the most abundant mineral in the thin section.

It occurs in very fine (0.01-0.1 mm) anhedral aggregates. The individual crystals appear strained and have sutured grain boundaries. Coarser grained quartz occurs in thin seams and pods which are generally parallel to the layering.

Muscovite occurs as small flakes between quartz grains.

It commonly shows a preferred orientation at about 60° to the layering. Muscovite rarely occurs in multigranular aggregates.

Chlorite occurs scattered through the matrix, in minor amounts in thin seams both with and without quartz, and in a few of the darker laminae in the sample. It is pale green, fibrous, and exhibits blue and brown interference colors.

Opaque grains occur in thin, discontinuous layers in the sample. Some seem to be dendrites parallel to the layers, and others are small nodules. Many crystals can be seen in hand specimens to occur as well formed

cubes.

Calcite and sphene occur in accessory amounts in some
of the layers.

Estimated Modal Composition

quartz	60%
muscovite	15%
chlorite	15%
others	10%
	100%

PETROGRAPHY OF SAMPLE PF-2B 136-136.5'

Name: Metabasalt

Macroscopic Description

This massive dark grey sample has a fine grained, uniform, phaneritic texture. Feltly plagioclase crystals (1-2 mm size) set in a dark groundmass are easily recognized with a hand lens. The plagioclase (Hardness 6) is evidently quite altered since it is readily pulverized by probing with a knife point (Hardness 5.5). In the black groundmass biotite flakes are large enough to be seen. Pyrite is present as widely dispersed anhedral grains. There are no veins or major cracks visible in the core. A few open pores are present.

Microscopic Description

Texture

The thin section displays a primary intersertal texture which is partially masked by secondary minerals. Plagioclase laths (0.5 mm) form a mat with ferromagnesian and secondary minerals filling the interstices. Cleavage cracks are not abundant. There is no evidence of healed cracks, no veinlets, and no other signs of structural disruption. Even the larger feldspar crystals are remarkably free of all types of micro-cracks.

The absence of deformation structures in this rock indicates that no significant non-hydrostatic stress has

existed after the last metamorphic event. Hence, if a fault is present in the vicinity of this rock, stresses, if any, have been small since the time of last metamorphism of the rock.

Mineralogy

Plagioclase originally composed about 40% of the rock. It is now very highly altered to sericitic products. Most of the lath-like crystals have a turbid appearance, and are uniform 0.5 mm. There are a very few larger crystals which are now sericitized.

Clinopyroxene (probably augite) occurs as abundant roundish grains 0.1 - 0.3 mm in diameter. The crystals have poor cleavage and weak zonation. The clinopyroxene is interpreted to be relict in this biotite grade metamorphic assemblage.

Opaque grains are relatively abundant in thin section. They commonly have square outlines, and occur in clumps with pyroxene and biotite.

Biotite occurs as subhedral crystals in the matrix. Basal sections are reddish brown while other orientations are pleochroic from yellowish brown to dark brown. The biotite is probably metamorphic in origin.

Chlorite is a common mineral in the matrix of this rock. It is pale green and fibrous. There are a few ovoid mats of chlorite about 1 mm in diameter which may represent replaced olivine crystals.

Apatite is an accessory mineral in this sample. Euhedral crystals are minute but common.

Actinolite needles are dispersed through the section.

Incipient blue green actinolite also seems to be present in some chlorite mats.

Minor amounts of sphene and hematite are also present in the rock.

Estimated Modal Composition

plagioclase 40%
 (plagioclase
 + sericite)

clinopyroxene 15%

opaque 10%

biotite 15%

chlorite 15%

apatite and
accessories 5%

 100%

PETROGRAPHY OF SAMPLE PF-2C 262.0-262.4'

Name: Fine Grained Amphibolite

Macroscopic Description

This specimen is a very fine grained dark green rock. The individual minerals are too small to identify with a hand lens. The rock is massive and non-foliated. It is cut by a weblike network of calcite and quartz veinlets. Small clots of pyrite are visible.

Microscopic Description

Texture

The sample displays a complex texture in thin section. The average grain size is about 50 microns. There is no preferred orientation or systematic foliation although the constituent minerals are metamorphic. There is a vague layering to the rock marked by arcuate clumps and thin layers of epidote and calcite. Calcite-quartz veins which randomly crosscut the rock are partially recrystallized.

Mineralogy

Amphibole (probably hornblende) and chlorite, in about equal proportions, are in the sample. The amphibole occurs as brownish green stubby, poorly formed crystals finely mixed with chlorite. The crystals are pleochroic from pale green to brownish green. There is no apparent preferred orientation of the grains.

Chlorite is also a major phase in the rock. It is generally

pale bluish green and forms both platy mats and stringy aggregates.

Epidote occurs as minute granular crystals clustered in veins, in arcuate clumps, and scattered through the matrix.

Quartz forms spongy crystals in the matrix barely resolvable at high magnification, and clear 0.1 mm crystals in veins.

Sphene is widely distributed as nodular aggregates and a few 0.1 mm subhedral crystals.

Apatite is present as accessory crystals.

Calcite is common in fine grained veins and in lesser amounts in the matrix.

Opaque grains are usually associated with veins and are not common in the matrix.

Estimated Modal Composition

amphibole	25%
chlorite	25%
epidote	20%
quartz	20%
calcite	5%
opaque	
sphene &	5%
apatite	
	100%

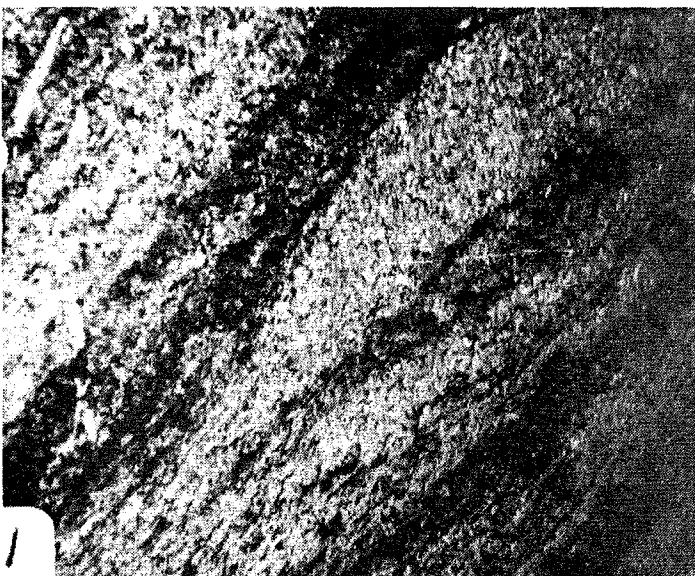


Photo 1. Sample PF-2A 99.5-99.9'. Felsic metatuff. Plane polarized light. Width of field 1.5 mm. The photomicrograph shows the fine grained nature of the sample. Roundish white spots are quartz which are obscured by muscovite and chlorite. The thin discontinuous laminae are composed of sphene, calcite, iron oxides, and chlorite. (The black circles are bubbles in the epoxy.)



Photo 2. Sample PF-2A 99.5-99.9'. Felsic metatuff. Plane polarized light. Width of field 0.5 mm. This photomicrograph is an enlarged view of the matrix and shows one of the few quartz veinlets which crosscuts the laminae. The thin, discontinuous laminae are composed of sphene, calcite, iron oxides, and chlorite. In this view, the dark laminae are almost opaque because the individual grains are only about $1-2\mu$.

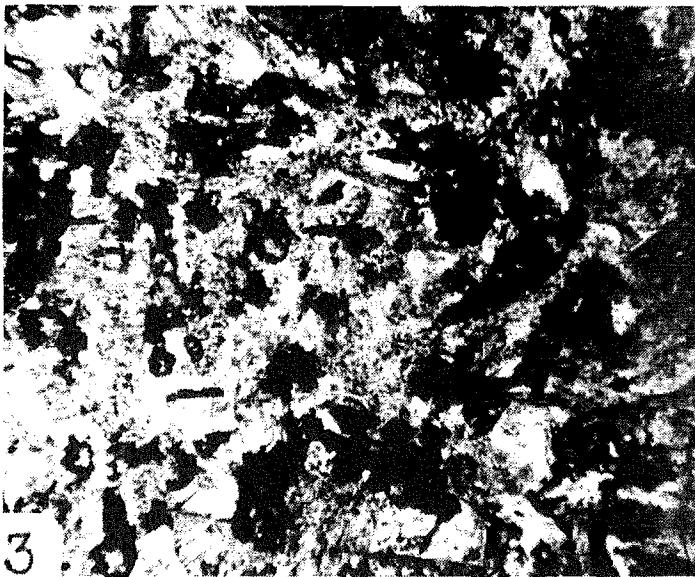


Photo 3. Sample PF-2B 136-136.5'. Metabasalt. Plane polarized light. Width of field 1.5 mm. This photomicrograph shows the typical textures observed in this sample. The light grey dusty looking background is altered plagioclase. Ovoid darker grains are relict clinopyroxene. Note the abundance of black grains; they are both opaque minerals and iron-rich biotite. See photo 4 for the details of the fabric.



4

Photo 4. Sample PF-2B 136-136.5'. Metabasalt. Plane polarized light. Width of field 0.5 mm. This photomicrograph shows the typical details of the fabric. Note how pervasively altered the plagioclase is. Note also the hexagonal biotite plates; the euhedral form implies that the biotite is metamorphic.



5

Photo 5. Sample PF-2C 262-262.4'. Fine grained amphibolite. Plane polarized light. Width of field 1.5 mm. This photomicrograph shows a typical view of this sample. The fine light and medium grey crystals are intergrown amphibole and chlorite; the white grains are quartz; and the darkest aggregates are clusters of epidote-calcite-sphene. Note the abundant arcuate quartz and epidote shapes; these are all polygranular.



6

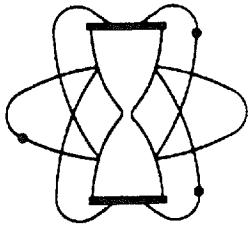
Photo 6. Sample PF-2C 262-262.4'. Fine grained amphibolite. Plane polarized light. Width of field 0.5 mm. This photomicrograph shows the intimate amphibole-chlorite intergrowths, and a granular epidote-sphene seam which arches across the field of view.

ATTACHMENT NO. 4

K-Ar AGE DETERMINATIONS OF SEVEN
SAMPLES RELATED TO THE INFERRED PORTSMOUTH FAULT

GEOCHRON LABORATORIES DIVISION
KRUEGER ENTERPRISES, INC.
CAMBRIDGE, MASSACHUSETTS

for
WESTON GEOPHYSICAL RESEARCH, INC.
WESTBORO, MASSACHUSETTS



KRUEGER ENTERPRISES, INC.
GEOCHRON LABORATORIES DIVISION

24 BLACKSTONE STREET • CAMBRIDGE, MASSACHUSETTS 02139 • (617) 876-3691

20 August 1974

Richard J. Holt
Weston Geophysical Res. Inc.
P.O. Box 550
Westboro, MA 01581

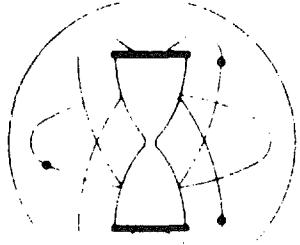
Dear Mr. Holt:

Enclosed are the analytical reports Mr. Rand requested. They are B-1236, B-1237 and B-1238 which were submitted for analyses on 20 January 1969.

Please forward these reports to Mr. Rand and if we can be of any further assistance, please do not hesitate to contact us.

Sincerely,

Derreth McStowe
Office Manager



KRUEGER ENTERPRISES, INC.

GEOCHRON LABORATORIES DIVISION

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POTASSIUM-ARGON AGE DETERMINATION

REPORT OF ANALYTICAL WORK

Our Sample No. B-1236 **Date Received:** 20 January 1969

Your Reference: B2 129.5 **Date Reported:** 31 January 1969

Submitted by: Ed Levine
Weston Geophysical Research Inc.
P.O. Box 364
Weston, MA

Sample Description & Locality:

Newburyport quartz diorite, biotite-bearing phase, drill core B2,
Seabrook, N.H.

Material Analyzed:
Biotite concentrate, -20/+100 mesh

$\text{Ar}^{40*}/\text{K}^{40}$ = 0.0186 AGE = $294 (\pm 9) \times 10^6$ yrs.

Argon Analyses:

Ar^{40*} , ppm.	$\text{Ar}^{40*}/\text{Total Ar}^{40}$	Ave. Ar^{40*} , ppm.
0.1431	0.950	0.1432
0.1432	0.953	

Potassium Analyses:

% K	Ave. %K	K^{40} , ppm
6.295	6.306	7.693
6.316		

Constants Used:

$$\lambda_\beta = 4.72 \times 10^{-10} / \text{year}$$

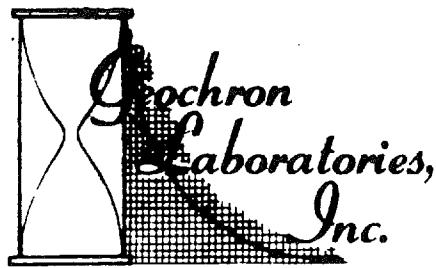
$$\lambda_e = 0.585 \times 10^{-10} / \text{year}$$

$$\text{K}^{40}/\text{K} = 1.22 \times 10^{-4} \text{ g./g.}$$

$$\text{AGE} = \frac{1}{\lambda_\beta + \lambda_e} \ln \left[\frac{\lambda_\beta + \lambda_e}{\lambda_e} \times \frac{\text{Ar}^{40*}}{\text{K}^{40}} + 1 \right]$$

Note: Ar^{40*} refers to radiogenic Ar^{40} .

M.Y. refers to millions of years.



24 Blackstone Street, Cambridge, Mass. 02139
Telephone: TRawbridge 8-8681

REPORT OF ANALYTICAL WORK

POTASSIUM-ARGON AGE DETERMINATION

Our Sample No. B-1237

Your Reference: ME #4 93'

Submitted by: Mr. Ed. Levine
Norton Geological Research, Inc.
P. O. Box 204
Norton, Mass.

Date Received: 29 January 1969

Date Reported: 31 January 1969

Sample Description & Locality:

Biotite-rich metasediment of the Merrimack Group,
Drill Core ME #4, 93', Seabrook, N. H.

Material Analyzed:

Biotite concentrate, -60/+200 mesh. The biotite was too fine grained to be completely free grains, therefore, a concentrate of the most biotite-rich grains was used. Estimated 70-80% biotite.
Ar^{40*}/K⁴⁰ = 0.0159 AGE = 254 (±9) × 10⁶ years.

Argon Analyses:

Ar ^{40*} , ppm.	Ar ^{40*} / Total Ar ⁴⁰	Ave. Ar ^{40*} , ppm.
0.0483	0.892	0.0483
0.0483	0.897	

Potassium Analyses:

% K	Ave. %K	K ⁴⁰ , ppm
2.430	2.486	3.033
2.542		

Constants Used:

$$\lambda_\beta = 4.72 \times 10^{-10} / \text{year}$$

$$\lambda_e = 0.585 \times 10^{-10} / \text{year}$$

$$K^{40}/K = 1.22 \times 10^{-4} \text{ g./g.}$$

$$\text{AGE} = \frac{1}{\lambda_e + \lambda_\beta} \ln \left[\frac{\lambda_\beta + \lambda_e}{\lambda_e} \times \frac{\text{Ar}^{40*}}{K^{40}} + 1 \right]$$

Note: Ar^{40*} refers to radiogenic Ar⁴⁰.



24 Blackstone Street, Cambridge, Mass. 02139

Telephone TRowbridge 6-3691

REPORT OF ANALYTICAL WORK

POTASSIUM-ARGON AGE DETERMINATION

Our Sample No. B-1238

Your Reference: B-1238

Submitted by:

Mr. M. Levine
Western Geophysical Inc., Lab.
P. O. Box 364
Norton, Mass.

Date Received:

29 January 1969

Date Reported: 31 January 1969

Sample Description & Locality:

Motite phase of Newburyport Quartz diorite, Drill core
B9, Seabrook, N.H. Coarse-grained diorite in igneous contact
with dark, fine-grained rock.

Material Analyzed: Motite concentrate. -40/+100 mesh, from coarse igneous phase.
Fresh motite, 75%, Chlorite, 15%; Amphibole, 10%.

$\text{Ar}^{40*}/\text{K}^{40}$ = 0.0179

AGE = 254 (±9) $\times 10^6$ years.

Argon Analyses:

Ar^{40*} , ppm.

0.0854
0.0904*
0.0860

$\text{Ar}^{40*}/\text{Total Ar}^{40}$

0.935
0.917

Ave. Ar^{40*} , ppm.

0.0857

(*Poor gas sample - not used in age calculation).

Potassium Analyses:

% K

3.998

Ave. %K

3.933

K^{40} , ppm

4.798

3.868

Constants Used:

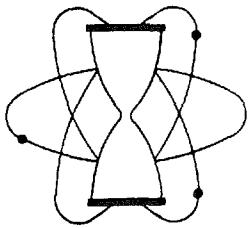
$$\lambda_\beta = 4.72 \times 10^{-10} / \text{year}$$

$$\lambda_e = 0.585 \times 10^{-10} / \text{year}$$

$$\text{K}^{40}/\text{K} = 1.22 \times 10^{-4} \text{ g./g.}$$

$$\text{AGE} = \frac{1}{\lambda_e + \lambda_\beta} \ln \left[\frac{\lambda_\beta + \lambda_e}{\lambda_e} \times \frac{\text{Ar}^{40*}}{\text{K}^{40}} + 1 \right]$$

Note: Ar^{40*} refers to radiogenic Ar^{40} .



KRUEGER ENTERPRISES, INC.
GEOCHRON LABORATORIES DIVISION

24 BLACKSTONE STREET • CAMBRIDGE, MASSACHUSETTS 02139 • (617) 876-3691

19 August 1974

Richard J. Holt
Weston Geophysical Res. Inc.
P.O. Box 550
Westboro, MA 01581

Dear Mr. Holt:

Enclosed are the analytical reports of the K-Ar age determinations on the seven (7) rock samples described in John Rand's letter of 18 July 1974.

These samples were a little difficult to work with because of the type of materials, however we did the best we could with them. The measured K-Ar ages are about what I would expect for these rocks.

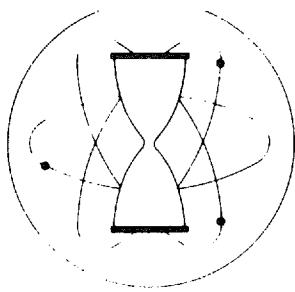
I will be away for a few days, but Hal Krueger will be here. I have discussed these results with him, and he is quite familiar with the geology of the area in question and with the work we did for you in this area several years ago. He will be happy to discuss these results with you in greater detail if you care to give him a call.

In the meantime, I am enclosing our invoice for this work. We look forward to serving you again in the near future.

Sincerely,

Richard H. Reesman
Richard H. Reesman
General Manager *D.M.*

RHR/dm
nelc: 7 reports & invoice #4473
cc: J.R. Rand (letter)



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GEOCHRON LABORATORIES DIVISION

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POTASSIUM-ARGON AGE DETERMINATION

REPORT OF ANALYTICAL WORK

Our Sample No. B-2882

Date Received: 22 July 1974

Your Reference: PF - S1

Date Reported: 16 August 1974

Submitted by: Richard J. Holt
Weston Geophysical Res., Inc.
P.O. Box 550
Westboro, MA 01581

Sample Description & Locality: Kittery quartzite
Towle Road, Hampton-Exeter Expressway
Hampton, New Hampshire

Material Analyzed: Chloritized biotite concentrate, -80/+200 mesh.

$\text{Ar}^{40*}/\text{K}^{40}$ = .01687

AGE = 268 \pm 10 M.Y.

Argon Analyses:

Ar^{40*} , ppm.

$\text{Ar}^{40*}/\text{Total Ar}^{40}$

Ave. Ar^{40*} , ppm.

.06717
.06588

.834
.862

.06653

Potassium Analyses:

% K

Ave. %K

K^{40} , ppm

3.224
3.242

3.233

3.944

Constants Used:

$\lambda_\beta = 4.72 \times 10^{-10}$ / year

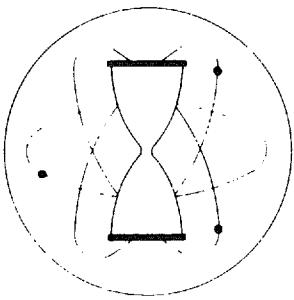
$$\text{AGE} = \frac{1}{\lambda_e + \lambda_\beta} \ln \left[\frac{\lambda_\beta + \lambda_e}{\lambda_e} \times \frac{\text{Ar}^{40*}}{\text{K}^{40}} + 1 \right]$$

$\lambda_e = 0.585 \times 10^{-10}$ / year

$\text{K}^{40}/\text{K} = 1.22 \times 10^{-4}$ g./g.

Note: Ar^{40*} refers to radiogenic Ar^{40} .

M.Y. refers to millions of years.



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POTASSIUM-ARGON AGE DETERMINATION

REPORT OF ANALYTICAL WORK

Our Sample No. A-2883

Date Received: 22 July 1974

Your Reference: PF - S2

Date Reported: 16 August 1974

Submitted by: Richard J. Holt
Weston Geophysical Res., Inc.
P.O. Box 550
Westboro, MA 01581

Sample Description & Locality: Rye fm. feldspathic quartzite
Winnicut Road, Route 151
North Hampton, New Hampshire

Material Analyzed: Amphibole concentrate, -80/+200 mesh. Estimated composition:
95% gray-black amphibole, 5% adhering groundmass.

$\text{Ar}^{40*}/\text{K}^{40}$ = .01960

AGE = 308 \pm 14 M.Y.

Argon Analyses:

Ar^{40*} , ppm.	$\text{Ar}^{40*}/\text{Total Ar}^{40}$	Ave. Ar^{40*} , ppm.
.01794	.674	.01773
.01752	.668	

Potassium Analyses:

% K	Ave. %K	K^{40} , ppm
.752	.741	.904
.731		

Constants Used:

$$\lambda_\beta = 4.72 \times 10^{-10} / \text{year}$$

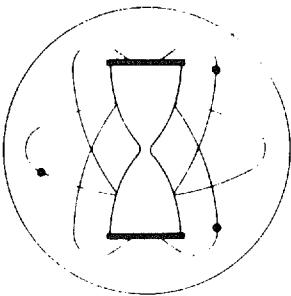
$$\lambda_e = 0.585 \times 10^{-10} / \text{year}$$

$$\text{K}^{40}/\text{K} = 1.22 \times 10^{-4} \text{ g./g.}$$

$$\text{AGE} = \frac{1}{\lambda_e + \lambda_\beta} \ln \left[\frac{\lambda_\beta + \lambda_e}{\lambda_e} \times \frac{\text{Ar}^{40*}}{\text{K}^{40}} + 1 \right]$$

Note: Ar^{40*} refers to radiogenic Ar^{40} .

M.Y. refers to millions of years.



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POTASSIUM-ARGON AGE DETERMINATION

REPORT OF ANALYTICAL WORK

Our Sample No. M-2884

Date Received: 22 July 1974

Your Reference: PF - S3

Date Reported: 16 August 1974

Submitted by: Richard J. Holt
Weston Geophysical Research Inc.
P.O. Box 550
Westboro, MA 01581

Sample Description & Locality: Rye fm. feldspathic gneiss
Route 1 Bypass, Lafayette Road
Portsmouth, New Hampshire

Material Analyzed: Muscovite concentrate, -80/+200 mesh. Estimated composition:
90% muscovite, 5% biotite, 5% quartz and feldspar.

Ar^{40*}/K⁴⁰ = .01864

AGE = 294 ± 10 M.Y.

Argon Analyses:

Ar ^{40*} , ppm.	Ar ^{40*} / Total Ar ⁴⁰	Ave. Ar ^{40*} , ppm.
.1522	.852	.1500
.1478	.782	

Potassium Analyses:

% K	Ave. %K	K ⁴⁰ , ppm
6.563	6.597	8.048
6.631		

Constants Used:

$$\lambda_\beta = 4.72 \times 10^{-10} / \text{year}$$

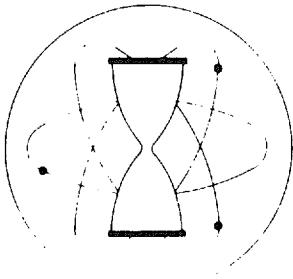
$$\lambda_e = 0.585 \times 10^{-10} / \text{year}$$

$$K^{40}/K = 1.22 \times 10^{-4} \text{ g./g.}$$

$$\text{AGE} = \frac{1}{\lambda_e + \lambda_\beta} \ln \left[\frac{\lambda_\beta + \lambda_e}{\lambda_e} \times \frac{\text{Ar}^{40*}}{K^{40}} + 1 \right]$$

Note: Ar^{40*} refers to radiogenic Ar⁴⁰.

M.Y. refers to millions of years.



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POTASSIUM-ARGON AGE DETERMINATION

REPORT OF ANALYTICAL WORK

Our Sample No. M-2885

Date Received: 22 July 1974

Your Reference: PF - S4

Date Reported: 16 August 1974

Submitted by: Richard J. Holt
Weston Geophysical Res., Inc.
P.O. Box 550
Westboro, MA 01581

Sample Description & Locality: Rye fm. feldspathic quartzite
Route 1 Bypass, Greenleaf Road
Portsmouth, New Hampshire

Material Analyzed: Concentrate of fine-grained mica-quartz aggregates, -80/+200 mesh.

$\text{Ar}^{40*}/\text{K}^{40}$ = .01645

AGE = 262 ± 11 M.Y.

Argon Analyses:

Ar^{40*} , ppm.	$\text{Ar}^{40*}/\text{Total Ar}^{40}$	Ave. Ar^{40*} , ppm.
.02042	.625	.02046
.02049	.645	

Potassium Analyses:

% K	Ave. %K	K^{40} , ppm
1.015	1.019	1.243
1.023		

Constants Used:

$$\lambda_\beta = 4.72 \times 10^{-10} / \text{year}$$

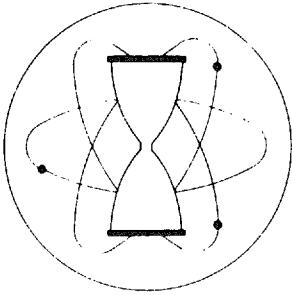
$$\lambda_e = 0.585 \times 10^{-10} / \text{year}$$

$$\text{K}^{40}/\text{K} = 1.22 \times 10^{-4} \text{ g./g.}$$

$$\text{AGE} = \frac{1}{\lambda_e + \lambda_\beta} \ln \left[\frac{\lambda_\beta + \lambda_e}{\lambda_e} \times \frac{\text{Ar}^{40*}}{\text{K}^{40}} + 1 \right]$$

Note: Ar^{40*} refers to radiogenic Ar^{40} .

M.Y. refers to millions of years.



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POTASSIUM-ARGON AGE DETERMINATION

REPORT OF ANALYTICAL WORK

Our Sample No. A-2886

Date Received: 22 July 1974

Your Reference: SRF - S1

Date Reported: 16 August 1974

Submitted by: Richard J. Holt
Weston Geophysical Res., Inc.
P.O. Box 550
Westboro, MA 01581

Sample Description & Locality: Diorite
Scotland Road, Interstate 95
Newbury, Massachusetts

Material Analyzed: Amphibole concentrate, -80/+200 mesh. Estimated composition:
85% amphibole, 10% biotite, 5% chlorite.

Ar^{40*}/K⁴⁰ = .02764

AGE = 422 ± 17 M.Y.

Argon Analyses:

Ar ^{40*} , ppm.	Ar ^{40*} / Total Ar ⁴⁰	Ave. Ar ^{40*} , ppm.
.03714	.807	.03892
.04070	.389	

Potassium Analyses:

% K	Ave. %K	K ⁴⁰ , ppm
1.154	1.154	1.407
1.154		

Constants Used:

$$\lambda_\beta = 4.72 \times 10^{-10} / \text{year}$$

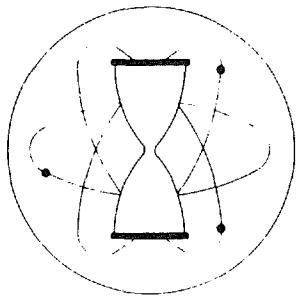
$$\lambda_e = 0.585 \times 10^{-10} / \text{year}$$

$$K^{40}/K = 1.22 \times 10^{-4} \text{ g/g.}$$

$$\text{AGE} = \frac{1}{\lambda_e + \lambda_\beta} \ln \left[\frac{\lambda_\beta + \lambda_e}{\lambda_e} \times \frac{\text{Ar}^{40*}}{K^{40}} + 1 \right]$$

Note: Ar^{40*} refers to radiogenic Ar⁴⁰.

M.Y. refers to millions of years.



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POTASSIUM-ARGON AGE DETERMINATION

REPORT OF ANALYTICAL WORK

Our Sample No. A-2887

Date Received: 22 July 1974

Your Reference: SRF - S2

Date Reported: 16 August 1974

Submitted by: Richard J. Holt
Weston Geophysical Res., Inc.
P.O. Box 550
Westboro, MA 01581

Sample Description & Locality: Schist
Highfield Road, Abandoned RR grade
Newbury, Massachusetts

Material Analyzed: Chlorite - amphibole concentrate, -80/+200 mesh. Estimated composition: 40% amphibole, 60% chlorite.

$\text{Ar}^{40*}/\text{K}^{40}$ = .01932

AGE = 304 \pm 15 M.Y.

Argon Analyses:

Ar^{40*} , ppm.	$\text{Ar}^{40*}/\text{Total Ar}^{40}$	Ave. Ar^{40*} , ppm.
.01162	.381	.01149
.01136	.548	

Potassium Analyses:

% K	Ave. %K	K^{40} , ppm
.492	.487	.594
.483		

Constants Used:

$$\lambda_\beta = 4.72 \times 10^{-10} / \text{year}$$

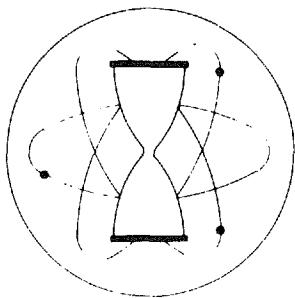
$$\lambda_e = 0.585 \times 10^{-10} / \text{year}$$

$$\text{K}^{40}/\text{K} = 1.22 \times 10^{-4} \text{ g./g.}$$

$$\text{AGE} = \frac{1}{\lambda_e + \lambda_\beta} \ln \left[\frac{\lambda_\beta + \lambda_e}{\lambda_e} \times \frac{\text{Ar}^{40*}}{\text{K}^{40}} + 1 \right]$$

Note: Ar^{40*} refers to radiogenic Ar^{40} .

M.Y. refers to millions of years.



KRUEGER ENTERPRISES, INC.

GEOCHRON LABORATORIES DIVISION

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POTASSIUM-ARGON AGE DETERMINATION

REPORT OF ANALYTICAL WORK

Our Sample No. B-2888

Date Received: 22 July 1974

Your Reference: SRF - S3

Date Reported: 16 August 1974

Submitted by: Richard J. Holt
Weston Geophysical Res., Inc.
P.O. Box 550
Westboro, MA 01581

Sample Description & Locality: Newburyport granodiorite
Parker Street, Little River area
Newburyport, Massachusetts

Material Analyzed: Chlorite-biotite concentrate, -80/+200 mesh. Estimated
composition: 70% chloritized biotite, 30% quartz.

$\text{Ar}^{40*}/\text{K}^{40}$ = .01860

AGE = 294 ± 20 M.Y.

Argon Analyses:

Ar^{40*} , ppm.

$\text{Ar}^{40*}/\text{Total Ar}^{40}$

Ave. Ar^{40*} , ppm.

.005765

.325

.005548

.005330

.370

Potassium Analyses:

% K

Ave. %K

K^{40} , ppm

.245

.244

.298

.244

Constants Used:

$$\lambda_\beta = 4.72 \times 10^{-10} / \text{year}$$

$$\text{AGE} = \frac{1}{\lambda_e + \lambda_\beta} \ln \left[\frac{\lambda_\beta + \lambda_e}{\lambda_e} \times \frac{\text{Ar}^{40*}}{\text{K}^{40}} + 1 \right]$$

$$\lambda_e = 0.585 \times 10^{-10} / \text{year}$$

$$\text{K}^{40}/\text{K} = 1.22 \times 10^{-4} \text{ g./g.}$$

Note: Ar^{40*} refers to radiogenic Ar^{40} .

M.Y. refers to millions of years.