An overal approach linking mobility and tsunami potential used for the Palos Verdes area (California)



Example: The Currituck Slide (165 km³)





Geomorphology and seismic to reconstruct slide sequences

Locat

et al.

2009

108

57

165

Slide Volumes

Prior

et al.

1986

78

46

124

Slide

2

Total





What could be a reasonable geometry for back analysis of the conditions prior to failure ?



Stratigraphic units and failure surfaces used for slide 1 and 2, and the hypothetical deltaic infill for a delta edge located at 0, 5 and 20 km from actual shelf break position assuming a fix position of the base of the foreslope. (vertical scale exaggeration is 3.75).



Gemorphology of the debris flow deposit and yield strength

Yield strength as a function of the critical height in the depositional zone. The colored box is for a range of reported thickness for the various depositional lobes. Black dot is for a height of 30 m and a yield strength of 2.0 kPa.

$$t_v = H_c (g' sinb)$$



Currituck mobility analysis



Initial acceleration may be unrealistic because as it assumes the material is already remoulded

Did the slide took place as a single event?

volume Initial and run-out distance for two values of the yield strength. Volume at (a) is from Prior et al. (1986) and his taken at 128 km³. Slide 1, and volume (b) is from our computation at 165 km³. Slide 1 and slide 2 are from models shown before. Field maximum run out is taken from field at 190 km.



Comments on the Currituck slide

- The Currituck slide took place as a single event.
- It involved a volume of sediment between 150 km³ and 165 km³.
- The mobilized yield strength was of the order of 2000 to 4000 Pa. (still may involved some water intake)
- It was triggered by a catastrophic event that must have required a sudden increase in pore pressure, likely due to an earthquake or a process rapidly generating a failure over a large surface.

Retrogressive failures and tsunamis: special conditions:

- 1. Presence of a weak layer (or *weakable*) ensuring a rapid propagation of the failure plane and the bulk mobilization of the sliding mass over on a remolded layer with a very low shearing resistance. Role of a film of water ?
- 2. Minimum acceleration must be reached before significant postfailure transformation takes place, e.g. desintegration, breakage into lumps, etc...
- **3.** Actual signature of retrogressive failure are mostly concentrated near the final escarpment with little knowledge of the slide dynamics in the lower starting zone.
- 4. Any phase of a retrogressive failure can generate a tsunami, as long as a significant volume of the sliding mass gains enough acceleration (everything else being constant)

Rapid development of a failure plane by the formation of a film of water at the hydraulic interface





Which phase triggered the tsunami?

Grand Banks, 1929





Storegga, 8200 yBP

Trigger: earthquake Storegga: V = 2.3 km³ $M = \sim$ 7.2, Grand Banks: V = ~100 km³

Comments on: Earthquakes and tsunamigenic slides







Photographs of the 1929 tsunami deposit (Tuttle et al., 2004).

Allan Ruffmann

Simulation du Tsunami du glissement des Grands Bancs slide (1929) par B. Bornhold (Uvic) – COSTA-Canada



1929 Grand Banks earthquake and tsunami, 1929

A proposal:

Already as part of the COSTA (2000-2003) project there were discussions to initiate an international study of the Grand Banks slide and tsunami with great interest from Norway, Spain and Canada, and I know of more recent interest. I think that in the light of the Japan disaster, this should consider even more seriously. We must remember that this is the most significant tsunamigenic landslide hitting the west coats in historical time.

Fredericton

An interesting question:

and a second second

How about the remaining strong earthquake potential in glaciated areas ?

Earthquake Risks in Glaciated Areas



Historical Seismicity on the American North Atlantic Seaboard





Ungava, Québec, Earthquake of 1989



Main Facts

8.5 km of surface faulting, Maximum throw I.8 m on a thrust fault Magnitude 6.3 Ms Hypocentre in Canadian Shield



Source: John Adams, GSC

Earthquake Risks

Will earthquake activities (frequency and magnitude) increase over the next few thousand years ?

0.00

0.00

0.09

0.06

0.03

0.00

0.03

44

12 Mary

0.06

0.03

Uniform h at 1E21

283

0.09

0.06

0.12



Impact on risk assessment

that the

amount of resound sitess available to trigger earthquake is decreasing with time. The intensity would depend on mantle viscosity to dissipate the stored strain energy due to glacio-isostasy

ttp://www.geo.ucalgary.ca/~wu/dFSMRate.html



Case of the 1663 Charlevoix Earthquake ($M = 7.8 \pm 0.6$)

Earthquake magnitude, intensity and distance using the model of Bakun and Hopper (2004) for Eastern North America.



Understanding earthquake source, frequency and magnitude is a key element needed to predict submarine slide hazards

Distance from the epicenter (km)

Locat 2011





2 AND CONTRACTOR OF STREET La Malbaie Bathymetry (>30m) of the Mv cottage! St. Lawrence Estuary between Tadoussac and Baie St. Paul HOMOTOTOTS:

00.00000

Saguenay Fjord

10 km

00.00001



Mapping slide prone areas and the potential type of failures is necessary to implement any regional risk assessment strategy. Potential trigger may vary according to the time laps (e.g. A new ice age in 10 000 years ?) and quantification mya depend on processes (not statistical)

Concluding remarks

Critical geomechanical properties of tsunamigenic failures are:
Strength of material (drained or undrained): in situ measurements may be essential
Sensitivity of a weak layer
Deformability and initial rate process (desintegration)

•Structure (rocks)

Limits:

Development of spread failure criteria still ungoing (A. Locat)
Acceleration still is difficult to predict correctly: Newer approaches using deformation models may help understand the transition between failure and post-failure
Mapping areas prone to failure and potential triggers
How about easy access to 3D seismics

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From deposition to failure: a simplified approach

