

**Attachment 1**

**Bell Bend Nuclear Power Plant Supplemental Information for FSAR Section 2.5**

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## 2.5 GEOLOGY, SEISMOLOGY, AND GEOTECHNICAL ENGINEERING

{This section of the U.S. EPR FSAR is incorporated by reference with the following supplements.

The summary includes a synopsis of FSAR Sections 2.5.1 through 2.5.5, including a brief description of the Site, investigations performed, results of investigations, conclusions, and identification of the organization that performed the work.

This section is intended to demonstrate compliance with the requirements of paragraph (c) of 10 CFR 100.23, "Geologic and Seismic Siting Criteria" (CFR, 2007a). Regulatory Guide (RG) 1.208, "A Performance-Based Approach to Define the Site-Specific Earthquake Ground Motion", (NRC, 2007a) is the primary guidance document for the development of the Site Probabilistic Seismic Hazard Analysis (PSHA) and Ground Motion Response Spectrum (GMRS).

The terms site region, site vicinity, site area, and site are used in this section to describe the specific areas of investigation. These terms correspond to the following areas:

1. The site region is that area within 200 miles (320 km) of the site location.
2. The site vicinity is that area within 25 miles (40 km) of the site location.
3. The site area is that area within 5 miles (8 km) of the site location.
4. The site is that area within 0.6 mile (1 km) of the site location.

### SUMMARY

The proposed Bell Bend Nuclear Power Plant (BBNPP) site is located south and west of the existing Susquehanna Steam Electric Station (SSES) in Salem Township, Luzerne County, Pennsylvania. The BBNPP site is found approximately 5 mi (8 km) northeast of the Borough of Berwick, Pennsylvania, and 1.5 mi (2.4 km) to the north and west of the north branch of the Susquehanna River. The major metropolitan centers closest to the site include: Wilkes-Barre, located 19 mi (31 km) to the northeast; Allentown, PA, approximately 50 mi (80 km) to the southeast; and Harrisburg, PA, which is approximately 70 mi (100 km) to the southwest.

The BBNPP Owner Controlled Area (OCA) is 882 ac (357 ha). The BBNPP site occupies an area of 424 acres (172 hectares) within the OCA. The BBNPP is not within the Exclusion Area Boundary for SSES Units 1 and 2.

FSAR Section 2.5 provides information on the seismic, geologic, and geotechnical characteristics of the site and the region surrounding the site. The purpose of this information is to permit an adequate evaluation of the proposed site, to support evaluations performed to estimate the site-specific ground motion response spectrum (GMRS), and to permit adequate engineering solutions to actual or potential geologic and seismic hazards at the proposed site. Details of the studies and investigations performed as well as the findings and conclusions are presented in sections 2.5.1 through 2.5.5.

The primary conclusions of Section 2.5 are as follows:

1. **Section 2.5.1 Basic Geologic & Seismic Information** – The site lies in a stable geologic region, and no geologic or man-made hazards have been identified within the



site area. The region has also experienced only minor earthquake activity, with no measured historical epicenter located within 50 miles of the Site.

2. **Section 2.5.2 Vibratory Ground Motion** – A PSHA was developed in accordance with Regulatory Guide 1.208, using: (1) the USGS 2002 catalog data update of the United States National Seismic Hazard Maps, (2) the EPRI-SOG 1986 seismic source zones, (3) updated ground motion models (EPRI, 2006), (3) deaggregation and site response analysis according to NUREG 6728-2B; and (4) a performance based Ground Motion Response Spectra (RG 1.208). Sensitivity studies which were performed on the 2002-2007 seismic events for 500 mi (805 km) and the Charlevoix EPRI source zone defined prior to the 1988 Sagueny Earthquake. These studies confirmed the results of the PSHA.
3. **Section 2.5.3 Surface Faulting** – The site vicinity exhibits little evidence of faulting, and all of it is non-capable. In addition, a targeted field investigation for the site vicinity discovered no evidence of paleoliquefaction features.
4. **Section 2.5.4 Stability of Subsurface Materials and Foundations** – Subsurface investigations concluded that the site analysis was bounded by the U.S. EPR FSAR for settlement, bearing capacity, hydraulic gradient and that the subsurface bedrock materials are stable with no potential for soil liquefaction. The Nuclear Island will be founded on the Mahantango Formation, which will be augmented with engineered fill to support the surface founded structures.
5. **Section 2.5.5 Stability of Slopes** – The site exhibits limited natural slopes and the only permanent slopes are those for the ESWEMS Retention Pond. All permanent and temporary slopes are shown to have adequate stability safety factors.

Additional details are summarized below.

## **Basic Geologic and Seismic Information**

### Regional Geology

The BBNPP site lies within the Appalachian Mountain Section of the Ridge and Valley Province that consists of long, narrow ridges and broad to narrow valleys exhibiting moderate to very high relief. These ridges and valleys are a direct result of lithologic disparities in erosional resistance and the folded and faulted structures developed in the geologic past, when the mountains were built, during the Alleghanian Orogeny.

This Province is primarily a zone containing Cambrian to Pennsylvanian rocks that were folded and faulted during the Alleghanian Orogeny that occurred during late Pennsylvanian through Permian times, nearly 300 million years ago. In addition to the geologic events that affected the entire Ridge and Valley Physiographic Province, three glacial advances affected the site vicinity during the Pleistocene Epoch.

The BBNPP site region is located in a stable continental region (SCR) characterized by low rates of crustal deformation with no active plate boundary conditions. There is no evidence for late Cenozoic seismogenic activity of any tectonic feature or structure within the site region (within 200 mi, 322 km).

Seismic activity associated with some of the larger tectonic features located beyond the site region has persisted to the present time. The nearest active tectonic feature is associated with the Newbury Liquefaction Features, located approximately 300 miles (480 km) to the northeast of the site. The Newbury feature in northeastern Massachusetts has been assigned to Class "A" status because of eyewitness reports of liquefaction during an earthquake in 1727, but the causative fault responsible for the ground motion and liquefaction remains unidentified. Three other features located beyond the 200-mile (322-km) site region have been included in the calculation of the BBNPP seismic hazard at low frequencies, based on updated models with new paleoseismological data. These features are: the New Madrid Seismic Zone, the Charleston Seismic Zone, and the Charlevoix-La Malbaie Seismic Zone in Canada.

#### Site Geology

Significant surface and subsurface investigations were conducted for the BBNPP site area foundation investigation, including field surveys, drilling, and geophysical exploration. Additional information was obtained from the Susquehanna Steam Electric Station Units 1 and 2 (SSES) FSAR.

The site is adjacent to the Susquehanna River near the southern edge of glaciation in Pennsylvania. The BBNPP site area is located within the Susquehanna Lowland Section of the Ridge and Valley Physiographic Province and is bordered by the Anthracite Valley Section to the north, and the Anthracite Upland Section to the south. The site area is underlain by folded Late Silurian, Devonian, and Lower Mississippian formations, with the Upper Devonian Mahantango Formation forming the bedrock directly beneath the site at the approximate elevation of 660 ft (20 m) msl. The most recent geologic influence on the site is the Wisconsinan glaciation that deposited glacial materials (including kame terrace, moraines and outwash) on the bedrock surface in a layer approximately 40 feet thick.

The topography within 5 mi (8 km) of the site consists of low to moderately high, linear ridges and valleys that primarily follow structural trends of the local geologic formations. Local elevations ranging from about 260 ft (79 m) to nearly 2,368 ft (722 m) msl. The BBNPP is planned to be constructed at a final grade elevation of 674 ft (205 m) msl, approximately 200 ft (61 m) above the Susquehanna River, and will be set back approximately 7,500 ft (2,286 m) from the River bank.

The local geologic formations have been subjected to a series of mountain-building episodes, including the Grenville, Taconic and Alleghanian orogenies. The local structure of the Ridge and Valley Province was imparted to the area during the Alleghanian Orogeny at the end of the Permian Period, nearly 250 million years ago. The Site geologic history has been quiet since the end of the Permian. At that time the local portion of the crust became more stable and tensional stresses predominated through the Cretaceous Period. The only disturbance of this quiet state was the advance of several ice sheets in the Pleistocene. However, since the site is located at the extreme southern limit of the glaciated area, the ice sheets were at their thinnest and any crustal depression or subsequent rebound from the ice load has been minimal.

Geologic studies to determine the site structural characteristics have been performed utilizing data obtained from site borings and geophysical surveys. In addition, bedrock exposures were mapped throughout the site area. A thorough search for faulting and detailed mapping of excavations was also performed. The site occupies a position on the northern limb of the northeast end of the Berwick Anticline. A pair of faults has been mapped near the axis of the anticline, but they have been mapped as being related to the Alleghanian Orogeny, and are not active today. No geologic hazards have been identified during either investigation.

The site is located in a region that has experienced only infrequent minor earthquake activity, with no earthquake-related epicentral locations detected within a 50-mile (80 km) radius of the site.

Investigations at the Site have not revealed any adverse geologic conditions that can be attributed to man's activity. The addition or withdrawal of subsurface fluids, including ground water, at the Site has not been significant. Material extraction in the Site vicinity has consisted of minor amounts of sand and gravel for roadbuilding. At present, there are no active mining operations within 11 miles (18 km) of the Site. There has been no mining or petroleum production in the Site area that would cause any surface or subsurface subsidence.

### **Vibratory Ground Motion**

Section 2.5.2 provides a detailed description of the vibratory ground motion assessment that was carried out to develop the BBNPP Ground Motion Response Spectra (GMRS). The GMRS is the first step in defining the Site Safe Shutdown Earthquake (SSE) response spectra. The Site SSE for BBNPP is further defined in Section 3.7.1.1.1 after reconciliation with the Certified Seismic Design Spectra (CSDRS) curves. The vibratory ground motion assessment was performed using a Probabilistic Seismic Hazard Analysis (PSHA). The PSHA was conducted in accordance with RG 1.208.

The PSHA process began with the creation of an updated seismic catalog. The United States Geological Service (USGS) "documentation for the 2002 Update of the National Seismic Hazard Maps"(USGS, 2002) was used as a starting point for earthquake catalog selection. The earthquake catalog was extended to include events up to the end of 2007.

The Electric Power Research Institute-Seismicity Owners Group (EPRI-SOG) seismic source zones defined in the 1989 EPRI/SOG study (EPRI, 1989a) have been adopted for updating the BBNPP site PSHA. Adjustments were required to include characteristic earthquake models that must be used to properly account for more recent information on the seismic activity in the New Madrid and Charleston seismic zones. The hazard contribution of the New Madrid Seismic Zone (NMSZ) was incorporated through a characteristic earthquake model of the New Madrid Fault System (NMFS). Results of several post-EPRI studies have demonstrated that the parameters of the Charleston seismic source need to be updated. The present PSHA for the BBNPP has adopted the Updated Charleston Seismic Source (UCSS) model that was also used in the seismic hazard studies that support the recent FSAR for the CCNPP Unit 3 (UniStar Nuclear, 2007). The Early Site Permit (ESP) Application for the Clinton NPP (EGC, 2006) submitted to the NRC on April 16, 2006 by Exelon Generation Company (EGC), and USGS new interpretations were included in the development of the NMFS characteristic earthquake model.

The following sources are considered for the PSHA at the BBNPP Site:

1. EPRI general area source zones extracted from the 1989 EPRI-SOG study.
2. Updated New Madrid Seismic Zone (NMSZ) as a characteristic model earthquake,
3. The Updated Charleston Seismic Source (UCSS) model

Site response analyses were conducted to evaluate the effect of the site geologic conditions on the generic Central and Eastern United States (CEUS) hard rock ground motions. The intent of the analyses is to develop ground motions at the surface that are consistent with the hazard levels defined for the generic rock conditions.

Uniform Hazard Response Spectra (UHRS) at hard rock are derived as well as their corresponding de-aggregation results for 1E-4, 1E-5, and 1E-6 hazard levels. The low frequency controlling event, 1 to 2.5 hertz, and the high frequency controlling event, 5 to 10 hertz, are identified from the de-aggregation results. These controlling events are prescribed in terms of magnitude (M) and distance (R) pairs. From the M-R pair defining each controlling event, the rock motion time histories whose epicenter distance and magnitude are close to the ones of the controlling event are selected.

For each controlling event, a spectral shape is adopted from the shapes for the Central and Eastern United States (CEUS) site according to the corresponding magnitude and distance pair (NUREG 6728). Then the response spectrum is scaled to match the rock UHRS at the spectral frequencies of 1.75 hertz (low frequency controlling event) and 7.5 hertz (high frequency controlling event). Next, the rock motion time histories are scaled to match the corresponding scaled controlling response spectrum. These scaled time histories are utilized as input to the Site Response Analyses.

A best estimate soil profile, with best estimate shear-wave velocity and material stiffness, density, and damping curves, is developed based on the site specific BBNPP Plant subsurface geophysical and geotechnical investigation. The site analysis uses randomized material properties and layer thicknesses based on the best estimate values and their variability. The results are expressed as mean site amplification functions corresponding to each controlling event. These factors are used to calculate the mean soil UHRS. In turn, the mean site amplification functions and the mean soil UHRS are used to derive the GMRS, with the application of the performance approach defined by RG 1.208. Vertical to horizontal ratios are subsequently used to calculate the vertical GMRS.

Two sensitivity analyses were performed as part of the BBNPP PSHA: (1) to assess the contribution to the hazard from the 2002-2007 seismicity with a 500 mile radius around the site, and (2) to assess the site specific significance of a modified Charlevoix seismic zone. In addition to the sensitivity analyses, recent research and interpretations, related to the St. Lawrence valley, the New England zone, and the Ramapo fault zone, have been evaluated in the context of the BBNPP PSHA. The first sensitivity study concluded that the 2002-2007 data results in a non-conservative prediction, (i.e., would result in additional margin) of seismic occurrence rates when compared to the USGS 2002 catalog. This would be a non-conservative reduction in the PSHA. The second sensitivity analysis concludes that the modification of the Charlevoix seismic zone to properly incorporate post-EPRI 1986 earthquakes does not have an effect in the BBNPP PSHA, mainly due to the distance to the site.

#### Surface Faulting

In order to assess the potential or lack of potential for surface rupture in the site region, a variety of detailed subsurface and surface investigation tools were employed. A detailed review of existing information was performed, including geologic maps, seismologic survey data, the USGS earthquake catalog, aerial and satellite imagery, local knowledge from local researchers, and published references. Review of the geotechnical drilling information from the adjacent SSES Site was undertaken to verify the lateral continuity of strata across the BBNPP site. In 2007 a surface outcrop survey was carried out within the site area. In addition, in October 2008, additional field investigations were performed by geologists and engineers to assess the presence of paleoliquefaction features along waterways within the site vicinity, and also to ground-truth the presence and surface expression of nearby Paleozoic-age faults.

Seismic refraction studies interpreted by Weston in 2008 show no indication of offset of the top of the Mahantango Shale, which is the local bearing stratum in the site area.

There are no documented zones of Quaternary deformation within the BBNPP site vicinity. No evidence of seismic-related disturbance has been found within the Mahantango Shale or the overlying glacial and post-glacial deposits.

It has been determined that there is no potential for tectonic rupture within the site area, and there are no capable tectonic sources within the site vicinity.

#### Stability of Subsurface Materials and Foundations

The natural topography at the BBNPP site is gently sloping. The maximum variation in relief is about 144.5 ft (44 m) across the site. Average elevation across the Site is about 680 ft (207 m).

The upper 400 ft (122 m) of the BBNPP soils was the subject of subsurface investigation. The site geology is comprised of glacial and postglacial soil deposits underlain by bedrock, which is on average 38.9 ft (11.9 m) below the ground surface. The subsurface is divided into the following stratigraphic units:

- Overburden Soils: Glacial tills that grade from fine silty and clayey soils at the surface to cobbly sand with boulders at the bottom. These soils cannot be used for foundations, because there is a potential for liquefaction.
- Bedrock Formation: the Mahantango Formation that is a dark gray to black shale, and will provide the foundation for the majority of the principal plant structures.

The field investigation for the site was performed in accordance with NRC Regulatory Guide 1.132, "Site Investigations for Foundations of Nuclear Power Plants" (NRC, 2003a). A thorough field investigation program was planned and implemented, and included the following:

- Boring Program
  - Wash Rotary Drilling and Standard Penetration Test
  - Rock Coring (NQ Wireline)
- In-Situ Pressure meter Testing
- Geophysical Exploration
  - Downhole Tests
  - PS Suspension Logging Tests
  - Deviation Surveys
  - Seismic Refraction Surveys

In total, 48 boreholes were completed for the BBNPP site, of which 27 boreholes were located in the vicinity of the proposed Category 1 structures. Three borings were extended to 400 feet (122 m) for detailed core logging and geophysical testing at the location of the reactor building.

A Hydrogeologic Field Investigation collected site-specific data to support a comprehensive hydrogeological evaluation of the plant site and surrounding areas. The data collected were utilized to support the surface hydrology analysis, hydrogeological characterization, and the development of a groundwater flow model.

A comprehensive laboratory testing program was performed on disturbed and undisturbed soil and rock samples, including the following:

1. Index and engineering classification
2. Strength
3. Consolidation
4. Permeability
5. Chemical Tests
6. Resonant Column Torsional Shear
7. Free-Free Resonant Column

The number and type of tests performed were consistent with the field investigation findings and the overall uniform conditions found at the site, and were performed in accordance with the guidance of the NRC Regulatory Guide 1.138, "Laboratory Investigations of Soils and Rocks for Engineering Analysis and Design of Nuclear Power Plants" (NRC, 2003b).

Category 1 Granular Structural Fill and Backfill for the eventual construction of the plant was identified from local sources and tested to determine the relevant engineering properties to be used in design analyses.

Recommendations of Soil, Fill and Rock properties were developed for all materials based on a combination of field measurements, laboratory testing, engineering analysis and judgment, and reference materials. Recommended properties are summarized in Section 2.5.4.2.5.

Foundation interfaces between the planned structures and site foundation materials were evaluated, and design parameters selected, in accordance with the requirements for COL applicants referencing the U.S. EPR FSAR.

#### Stability of Slopes

The BBNPP site is comprised of generally flat topography in the vicinity of the primary structures and components. The Site is planned to be graded to establish the final grade for the project, resulting in minor cuts and fills, as well as slopes.

The stability of temporary and permanent slopes was evaluated using limit equilibrium methods, resulting in a Factor of Safety for the slope section analyzed.

The ESWEMS Retention Pond slopes are the only permanent slopes planned for the BBNPP. Slope stability analysis results indicate that the pond side slopes have Factor of Safety values ranging from 2.0 to 9.2, indicating that the proposed slope design and configuration is stable under all considered loading conditions.

Temporary cut and fill slopes will exist in dry conditions during construction. Slope stability analysis results indicate the temporary slopes have Factor of Safety value of 1.3, indicating that the proposed slope design is stable under all considered loading conditions.

## **2.5.1 BASIC GEOLOGIC AND SEISMIC INFORMATION**

The U.S. EPR FSAR includes the following COL Item in Section 2.5.1:

A COL applicant that references the U.S. EPR design certification will use site-specific information to investigate and provide data concerning geological, seismic, geophysical, and geotechnical information.

The COL Item is addressed as follows:

This section presents information on the geological and seismological characteristics of the site region (200 mi (322 km) radius), site vicinity (25 mi (40 km) radius), site area (5 mi (8 km) radius) and site (0.6 mi (1 km) radius). Section 2.5.1.1 describes the geologic and tectonic characteristics of the site region. Section 2.5.1.2 describes the geologic and tectonic characteristics of the site vicinity and location. The geological and seismological information was developed in accordance with the following NRC guidance documents:

- Regulatory Guide 1.70, Section 2.5.1, "Basic Geologic and Seismic Information," (NRC, 1978)
- Regulatory Guide 1.206, Section 2.5.1, "Basic Geologic and Seismic Information," (NRC, 2007) and
- Regulatory Guide 1.165, "Identification and Characterization of Seismic Sources and Determination of Safe Shutdown Earthquake Ground Motion," (NRC, 1997).

### **2.5.1.1 Regional Geology (200 mi (322 km) radius)**

The U.S. EPR FSAR includes the following COL Item in Section 2.5.1.1:

Regional geology is site specific and will be addressed by the COL applicant.

This COL Item is addressed as follows:

This section discusses the physiography, geologic history, stratigraphy, and tectonic setting within a 200 mi (322 km) radius of the site. The regional geologic map and regional physiographic map, as shown in Figure 2.5-5 and Figures 2.5-7 (USGS, 2002), respectively contain information on the geology, stratigraphy, and tectonic setting of the region surrounding the {BBNPP site}. Summaries of these aspects of regional geology are presented to provide the framework for evaluation of the geologic and seismologic hazards presented in the succeeding sections.

{Section 2.5.1.1.1 through Section 2.5.1.1.4.4.13 are added as a supplement to the U.S. EPR FSAR.

#### **2.5.1.1.1 Regional Physiography and Geomorphology**

The BBNPP site lies within the Ridge and Valley Physiographic Province as shown in Figure 2.5-7. The area within a 200 mi (322 km) radius of the site encompasses parts of seven other physiographic provinces. These are the Appalachian Plateaus Province, the Piedmont Province, the New England Province, the Atlantic Coastal Plain Province, the Blue Ridge Province, the Central Lowlands Province, and the Adirondack Province (Barnes, 2002). The



physiographic provinces in the site region are shown on Figure 2.5-7. Each of these physiographic provinces is briefly described in the following sections. A map showing the different sections and subsections within the physiographic provinces of Pennsylvania, as depicted by the Pennsylvania Geological Survey (PGS), is shown on Figure 2.5-8.

#### **2.5.1.1.1.1 Physiography and Geomorphology of the Ridge and Valley Physiographic Province**

The Ridge and Valley Physiographic Province occupies most of central Pennsylvania, extending from the West Virginia and Maryland borders to northeastern Pennsylvania including the site region. The Ridge and Valley Province is bordered to the west and north by the Appalachian Plateaus Province and to the southeast by the Piedmont Province as shown in Figures 2.5-6 (DCNR, 2000), 2.5-7 (USGS, 2002), and 2.5-8 (Sevon, 2002). On a regional scale, the Ridge and Valley Province extends in a northeast striking zone of varying width from Alabama to New Jersey. The Ridge and Valley Province is primarily a zone containing Cambrian to Pennsylvanian rocks folded and faulted during Alleghenian orogenic events that affected eastern North America during late Mississippian through Pennsylvanian times (Hatcher, 1987). Folding resulted in complex geologic structure including synclinal, anticlinal, and nappe structures while faulting resulted in thrust faults, and lateral ramp thrusts. In addition, significant strike slip components are recognized in many Ridge and Valley Faults (Faill and Nickelsen, 1999).

The complex geologic structure and varying lithologies, including repetition of some parts of the stratigraphic section, directly affect the geomorphology of the region and site vicinity. The Appalachian Mountain Section of the Ridge and Valley Province consists of long, narrow ridges and broad to narrow valleys exhibiting moderate to very high relief that exist as a direct result of lithologic disparities in resistance to erosion and location of these resistant lithologies directly related to structure. These ridges typically are the remnant flanks of breached anticlines, typically capped by Cambrian sandstone and quartzite, and synclines underlain by resistant cherty limestones and sandstones of the Upper Silurian and Lower Devonian Keyser and Oriskany formations (DCNR, 2008a).

In addition to the importance of lithology in defining resistance to erosion and the formation of ridges versus valleys, the susceptibility of limestone and dolomite within the Ridge and Valley Province has resulted in significant formation of karst features such as caverns, surface subsidence and collapse, springs, and disappearing streams. Limestone and/or dolomite formations occur in various thicknesses in Cambrian through Pennsylvanian rocks in the Ridge and Valley Province. While not as prevalent as in the Appalachian Plateaus Province, karst is significant in the Ridge and Valley Province. In addition, the abundant fractures, folds, and faults of the Ridge and Valley Province serve as zones where dissolution can initiate and concentrate. However significant the presence of karst throughout some of the Ridge and Valley Province, there is only a small area of potential karst development in the extreme western extent of the site area (5 mi (8km) radius) and based on the lithology of the Mahantango shale and of the Timmers Rock Sandstone, bedrock at the site exhibits no potential for karst development within the site location (0.6 mi (1km) radius).

The Ridge and Valley Province has a sub-section known as the Great Valley Section, as depicted on Figure 2.5-6. The Great Valley Section of the Ridge and Valley Province consists of a very broad lowland that lies south of the Blue Mountain Sub-Section as depicted in Figure 2.5.8 in southeastern Pennsylvania. The lowland has gently undulating hills eroded into shales and siltstones on the north side of the valley and a lower elevation, flatter landscape developed

on limestones and dolomites on the south side (DCNR, 2008b). Elevations in the Appalachian Mountain Section range from 440-2,775 ft (134-846 m) msl while elevations in the Great Valley Section range from 140-1,100 feet (43-335 m) msl (DCNR, 2008b). In addition to the geologic events that affected the entire Ridge and Valley Physiographic Province, four glacial events affected the site region (200 mi (322 km) radius).

Three of four main periods of continental glaciation occurring in Pennsylvania directly affected the site vicinity (25 mi (40 km) radius) as shown on Figure 2.5-9 (Sevon, 2000c). These glacial events occurred in the following order from oldest to youngest; Early Pleistocene, Early middle Pleistocene, Middle Pleistocene, and Late Pleistocene (Braun, 2004). The oldest glaciation extended the farthest south, with each subsequent glacial event never advancing past the previous one, as shown in Figure 2.5-9. These older glacial advances are more difficult to identify due to the eroding attributes of more recent glaciers. The area south of the late Pleistocene glacial limit is characterized by extensive colluvial deposits and other features of periglacial origin (Braun, 2004) including frost riving and congelifluction (Sevon, 1999). The limit of the late Pleistocene glacial event, also known as the Late Wisconsinan (17,000-22,000 yrs), is marked by heads-of-outwash in the valleys with an indistinct moraine on adjacent hillsides (Braun, 2004) and is labeled as Olean Till as shown in Figure 2.5-9. The overall trend of the late Wisconsinan margin across northeastern Pennsylvania is approximately west-northwest, and hilltop striae on the Appalachian and Pocono plateaus within 30 mi (48 km) of the margin indicate a regional ice flow direction of north-south to south-southwest (Braun, 1988). The Late Illinoian (132,000-198,000 yrs) glacial event advanced only a few miles from the more recent Late Wisconsinan event, as shown in Figure 2.5-9, and is identified by heads-of-outwash in the valleys and discontinuous patches of till or colluvium derived from till (Braun, 1988). Pre-Illinoian glaciations advanced approximately 20-40 mi (32-64 km) beyond the Late Illinoian limit, as shown in Figure 2.5-9. Glacial lake sediments and two belts of "markedly thicker glacial deposits" suggest that Pre-Illinoian stage (>770,000 yrs) northeastern Pennsylvania was subjected to two glacial events (Braun, 2004). The first of which extended to the maximum glacial limit as shown in Figure 2.5-9, and the second extended only several miles northeast of the maximum glacial limit (Braun, 2004). In addition to the deposition of glacial sediments, surface hydrology has been affected by glacial melt and outwash. During glacial retreats, large volumes of glacial melt-waters formed broad, high energy streams including the Susquehanna, and other neighboring rivers such as the Delaware and Potomac Rivers.

The geologic structure, lithologic makeup, and glacial history of the Ridge and Valley Province define the physiography and geomorphology of the majority of the site vicinity (25 mi (40km) radius and the site area (5 mi (8 km) radius). The site area (5 mi (8 km) radius) physiography is discussed in further detail in Section 2.5.1.2.1.

#### **2.5.1.1.2 Physiography and Geomorphology of the Appalachian Plateaus Physiographic Province**

Located west of the Ridge and Valley Province, the Appalachian Plateaus Physiographic Province includes the western part of the Appalachian Mountains, and stretches from New York to Alabama. The mountains within the Appalachian Plateaus Province are generally long, narrow, and even crested and valleys are highly variable in width and elevation (Way, 1999). Much of the current day landscape of this Province in Pennsylvania developed during multiple periods of glaciation within the Pleistocene period (Way, 1999). The Allegheny Front is the topographic and structural boundary between the Appalachian Plateaus and the Ridge and Valley Province (Clark, 1992). It is a bold, high escarpment, underlain primarily by clastic sedimentary rocks capped by sandstone. In eastern West Virginia. Elevations along this escarpment reach 4,790 ft (1,460 m) msl (Hack, 1989) while in Pennsylvania, its highest point is

3,210 ft (978 m) msl. West of the Allegheny Front, the Appalachian Plateaus topographic surface slopes gently down to the northwest. A large portion of the Appalachian Plateaus Province lies within 200 mi (322 km) of the BBNPP site as shown in Figure 2.5-7, and the northern portion of the site vicinity (25 mi (40 km) radius) encompasses the Appalachian Plateaus Physiographic Province as depicted on Figure 2.5-6. Note that no other physiographic province lies within the BBNPP site vicinity (25 mi (40 km) radius).

#### **2.5.1.1.1.3 Physiography and Geomorphology of the Piedmont Physiographic Province**

The Piedmont Physiographic Province extends southwest from New York, through southeast Pennsylvania, to Alabama and lies southeast of and adjacent to, the Ridge and Valley Physiographic Province as shown in Figure 2.5-7. The Piedmont Province is about 60 mi (97 km) wide in southeastern Pennsylvania and narrows northward to about 10 mi (16 km) wide in southeastern New York. Elevation in the Piedmont Province ranges from 20-1,355 feet (6-413 m) mean sea level (msl) (DCNR, 2007b; DCNR, 2007c; and DCNR, 2007d).

In Pennsylvania, the Piedmont Province is divided into the Piedmont Lowland Section, the Gettysburg-Newark Lowland Section, and the Piedmont Upland Section. With the exception of the Piedmont Lowland Section, the majority of the Piedmont Province consists mainly of rolling low hills and valleys (DCNR, 2007a). The Piedmont Lowland Section consists of broad, moderately dissected valleys separated by broad low hills and is developed primarily on limestone and dolomite rock highly susceptible to karst topography (DCNR, 2007b). The Gettysburg-Newark Lowland Section runs adjacent to the Great Valley Section of the Ridge and Valley Province as shown in Figure 2.5-8. The Gettysburg-Newark Section consists of rolling low hills and valleys developed on sedimentary fluvial and lacustrine clastic rock deposits that represent a series of exposed faulted rift basins (Root, 1999). Metamorphic rocks of varying affinity comprise the surface rock outside of the rift basins within the Gettysburg-Newark Lowland Section. The Piedmont Upland Section exhibits gently rolling hills and valleys. Drainage in the Piedmont Upland Section is often controlled by a well developed foliation in predominant schists with drainage developing along foliation or normal to foliation (DCNR, 2007).

#### **2.5.1.1.1.4 Physiography and Geomorphology of the New England Physiographic Province**

The New England Physiographic Province is bounded on the north by the Ridge and Valley Province and on the south by the Piedmont Province as shown in Figure 2.5-7. The New England Province, aligned in a northeast-southwest direction, extends from the eastern border of Pennsylvania to mid-southeastern Pennsylvania occupying only a small amount of area as compared to the surrounding provinces. The province has an average width of about 5 mi (8 km) within Pennsylvania, and consists of circular to linear, rounded low hills or ridges that project upward in significant contrast to surrounding lowlands (DCNR, 2007e). The hills and ridges are made up of granitic gneiss, granodiorite, and quartzite thus making them very resistant to erosion (DCNR, 2007e). This province has a local relief ranging from 300-600 ft (91-183 m) msl with elevations ranging from 140-1,364 ft (43-416 m) msl (DCNR, 2007e).

#### **2.5.1.1.1.5 Physiography and Geomorphology of the Atlantic Coastal Plain Physiographic Province**

The Atlantic Coastal Plain Physiographic Province lies east of, and adjacent to, the Piedmont Province and occupies much of the eastern seaboard, as shown in Figure 2.5-7. In Pennsylvania, this area is designated as the Lowland and Intermediate Upland Section of the

Atlantic Coastal Plain Province as shown in Figure 2.5-8. This section consists of a flat upper terrace surface that is cut by numerous short streams, which are typically narrow and steep sided (DCNR, 2008c). The province is aligned in a northeast-southwest direction and is, on average, 6 mi (10 km) wide in Pennsylvania but attains a width of up to 50 mi (80 km) in New Jersey. The unconsolidated to poorly consolidated sand and gravel deposits of the Coastal Plain, dip gently to the southeast (NJGS, 2003). These sediments rest on various metamorphic rocks (DCNR, 2008c). Local relief is very low in the Lowland and Intermediate Upland Section of Pennsylvania, and elevations range from sea level to 200 ft (61 m) msl (DCNR, 2008c). The highest elevation of the Atlantic Coastal Plain Province in New Jersey is 391 ft (119 m) msl (NJGS, 2003).

#### **2.5.1.1.1.6 Physiography and Geomorphology of the Blue Ridge Physiographic Province**

The Blue Ridge Province extends less than 50 miles (80 km) into Pennsylvania from the south and is approximately 70 miles (113 km) from the BBNPP site. The Blue Ridge Physiographic Province is bounded on the east by the Piedmont Province and on the west by the Valley and Ridge Province as shown in Figure 2.5-7 (USGS, 2002). Figures 2.5-6 (DCNR 2000c) and 2.5-8 (Sevon, 2002) do not include the Blue Ridge in the statewide designation of physiographic provinces. The Blue Ridge Province extends from Pennsylvania to Georgia in a northeast-southwest direction and is underlain primarily by metamorphosed Precambrian and Early Paleozoic igneous and sedimentary rock (VADOT, 2008). The Blue Ridge Province is recognized as the core of the Appalachian Mountains, emplaced during Alleghanian tectonism along a regional detachment structure (Hatcher 2004). Soils of the Blue Ridge are predominantly colluvium with small amounts of alluvium along the rivers and streams (VADOT, 2008). Residuum occurs locally but is limited in thickness and lateral extent due to aggressive erosive forces caused by steep topography. The Blue Ridge is a long, linear province which ranges in width from about 5 mi (8 km) in Maryland to over 50 mi (80 km) in North Carolina. Elevations in the Blue Ridge Province exceed 6,600 (2,012m) feet in North Carolina and Tennessee.

#### **2.5.1.1.1.7 Physiography and Geomorphology of the Central Lowlands Physiographic Province**

The Central Lowlands Physiographic Province, also known as the Ontario Lowlands, has relatively low relief (Komor, 1998) and is located between the Appalachian Plateaus Province to the south and Lake Ontario to the north as seen on Figure 2.5-7. The Central Lowlands were subjected to glaciation and as a result, consist mainly of unconsolidated surficial materials including mostly sands and gravels (DCNR, 2008d). Elevation within the province ranges from 570 ft (174 m) to approximately 1,000 ft (305 m) as erosion processes along the shores of Lake Erie have created a steep lake-land interface along much of the shoreline (DCNR, 2008d).

#### **2.5.1.1.1.8 Physiography and Geomorphology of the Adirondack Physiographic Province**

The Adirondack Physiographic Province has moderate to high relief throughout and its circular shape attains a diameter of over 150 mi (241 km). Elevations within the Adirondack Province range from 1,500 ft. (457 m) to 5,344 ft.(1,629 m) The Adirondack Physiographic Province is located in northern New York and is surrounded by the Ridge and Valley Province to the southeast, the Appalachian Plateaus to the south, the Central Lowlands Province to the west, and the St. Lawrence Valley Province to the north as seen in Figure 2.5-7. The bedrocks of the Adirondacks are primarily Precambrian to early Paleozoic, metamorphic rocks and are part of the great Canadian Shield that has been uplifted to its present day geography (McDonnell,

2008). The bedrock generally supplements long, straight valleys, gently curved ridges, and a radial drainage pattern (Komor, 1998) throughout the province.

#### **2.5.1.1.2 Regional Geologic History**

The BBNPP site is located within the Appalachian orogenic belt, a geologic region marked by a complex history of orogenic events, rift sequences, subsequent depositional sequences, eustatic and local sea level changes, and glacial events. The region's position along an active continental margin at intervals in the Proterozoic and Paleozoic Era has developed the structural and stratigraphic characteristics that define the seismotectonic setting. Episodes of continental collisions have produced a series of terranes separated, in part, by low angle detachment faults (Pohn, 2000). Sources of seismicity may occur in the stratigraphy along structures within the North American basement, along the terranes, and over thrust plates. Tectonic episodes of continental rifting have produced high angle normal and boundary faults that extend to the aforementioned detachment faults and in some cases through the upper crust. These rift related faults do not extend to within the site vicinity (25 mi (40 km) radius) and have minimal to no influence on BBNPP site (0.6 mi (1 km) radius) seismicity. Direct evidence of these deformational events is visible in the Ridge and Valley province, as described in Section 2.5.1.1.4.3, and borehole data as described in Section 2.5.1.1.3 which identifies the Precambrian rock as Grenvillian. The site region is located currently on the passive margin of the North American plate following latest Permian and early to middle Mesozoic Era continental extension and rifting. The events that have affected the site region (200 mi (322 km) are discussed in terms of orogenesis, depositional basins and environments, and their effects on site region (200 mi (322 km) to site vicinity (25 mi (40 km) radius) geology.

#### **2.5.1.1.2.1 Grenville Orogeny**

The Grenville Orogeny occurred as a result of the Mesoproterozoic collision of North and South America (Eriksson, et al, 2003) during the construction of the supercontinent Rodinia (Millot, 2001) between 1.0 and 1.20 Ga (Murphy, et al, 2000). The Grenville Orogeny occurred along the southern and eastern margin of the North American craton and resulted in the accretion and crustal evolution of the eastern and southern margins of the North American Craton during the Proterozoic. Crustal evolution along the Grenville has been studied to determine its role in evolution of the Appalachian Orogen (Mueller, et al, 2008). This Proterozoic orogenic province is comprised of a "laterally continuous region of high grade, polymetamorphic terranes and is also recognized as a multiphase orogenic event with distinct zones connected by ductile shear zones (Streepey and Johnson 2001) The Grenville Province is also described as primarily orthogneisses (Mueller, et al, 2008) Grenville basement rocks outcrop in numerous locations along the Appalachian mountains from Alabama through Nova Scotia and in Texas (Reese and Sharon, 2004) and form the basement upon which Paleozoic orogenic and depositional events and sequences occurred. The most significant outcrop of the Grenville Province within the site region is the Adirondack Mountains. The location of the Adirondack physiographic province, comprised of the Grenville age Adirondack Mountains, is depicted on Figure 2.5-7. The site region (200 mi (322 km) radius), site vicinity (25 mi (40 km) radius), site area (5 mi (8 km) and site location (0.6 mi (1 km) radius) all have Grenville rocks as the Precambrian basement. The depth to Grenville basement at the BBNPP site is approximately 33,000 ft (10,058 m) as shown on Figure 2.5-19.

#### **2.5.1.1.2.2 Late Precambrian to Upper Cambrian Rifting and Late Precambrian to Late Ordovician Deposition**

Following the Grenville orogeny, late Neoproterozoic to early Cambrian rifting initiated the breakup of the supercontinent, Rodinia, and led to the formation of basins in which late Precambrian sediments accumulated in the rift-related Appalachian foreland basin. Rift related structures such as the Central Pennsylvania trough, a northeast trending rift graben that exists to the south approximately 50 miles (80 km) from the BBNPP site, is analogous to a series of basement troughs that formed during the opening of the Iapetus Ocean (Gao, et al, 2000). Age analysis of a group of intrusive to hypabyssal rocks associated with rift related normal faulting in the North Carolina and Virginia Blue Ridge reveal age ranges from 760 to 700 Ma (Bailey and Tollo, 1998) and ages of initial Rodinian breakup have been determined at 780 Ma (Harlan and LeCheminant, 2003), indicating a time range for initial rifting. Rifting continued into Earliest Cambrian until the completion of the opening of the Iapetus Ocean and change to a passive continental margin approximately at the Upper Cambrian (Gao, et al, 2000). The initiation of rifting was also the onset of the deposition of a thick sequence of terrigenous and volcanic detritus, carbonate rocks, and distal more argillaceous sediment in the Appalachian foreland basin that indicates a predominantly transgressive sea with minor local and temporal exceptions (Kauffman, 1991 and Hasson and Haase, 1988). Contribution of detritus to the basin from the Grenville Province is established by a predominance of detrital zircons in Cambrian sandstones (Eriksson, 2003). Sediment load enhanced basin subsidence and faulting, which resulted in localized zones of sequence thickening (Hasson and Haase, 1998) during the Middle Cambrian through the termination of rifting. The Waynesboro, Pleasant Hill, and Warrior Formations represent the site area (5 mi (8 km) radius) stratigraphy that is part of this sequence (Figure 2.5-21). Deposition of terrigenous and carbonate sequences continued along the passive margin of the North American craton from the Upper Cambrian through the Upper Ordovician. Fluctuations in sea level shifted the shoreline and depositional environment throughout this time as indicated by stratigraphic variations from sandstones to shales indicating significant variations in distance of the depositional environment to the sediment source. **The west-northwest extent of Iapetan rifting is significant to the current seismic regime. Faults related to Iapetan rifting and the Iapetan passive margin have been identified as the source of seismic activity in Tennessee, Virginia, and Quebec (Wheeler, 1996). Specific seismogenic sources are discussed in detail in Section 2.5.1.1.4.**

#### **2.5.1.1.2.3 Taconic Orogeny and Clastic Wedge Deposition**

The North American craton became a convergent margin with the onset of the Taconic Orogeny at the end of the Middle Ordovician and continuing through the Middle Silurian (Gao et al, 2000). The Taconic Highlands, an island arc terrane (Trembley, Bedard, and Lauziere, 1997) converged to the east of the North American craton and became the dominant source of detritus to the Appalachian basin. Along the margin of the craton, especially in New England and Canada, suites of ocean floor and island arc were accreted to the continent along with what is regionally called the Taconic Thrust Belt (Hayman and Kidd, 2002). Deformation from the Taconic orogeny was imparted throughout the continental margin, as metamorphic events of Taconian (457 Ma) time are recorded in the Blue Ridge of western North Carolina (Moecher and Miller, 2004). An example of Taconic deformation within the site region (200 mi (322 km) radius) is the development of the Hamburg Nappe (Pohn, 2000), an overthrust fold derived from folding of basinal sediment due to thrusting of the Taconic front onto the continent. Similar structures include the Lebanon Valley, Irish Mountain, Applebutter, Musconetcong and Lon Station-Paulins Kill which form the Musconetcong nappe megasystem. During this time the site region (200 mi (322 km) radius) existed within the Appalachian basin and received significant

amounts of detritus derived from the Taconic Highlands to the east and significantly less detritus from the continent. This led to a clastic wedge thickening to the east toward the Taconic sediment source (Thompson, 1999 and Castle, 2005) and thinning toward the continental margin. The timing of the Taconic Orogeny ranges from the Middle Ordovician into the Middle Silurian. The site vicinity (25 mi (40 km) radius) was positioned on the western side of the Appalachian basin during the Taconic Orogeny. The effect of the Taconic Orogeny and Clastic Wedge deposition on the site vicinity (25 mi (40 km) radius) and site area (5 mi (8 km) radius) is primarily the deposition of the Middle Ordovician to Lower Silurian strata as noted on Figure 2.5-21. The Martinsburg, Bald Eagle and Juniata Formations were derived from the erosion of the tectonic highland. The composition of the aforementioned Formations is provided in Section 2.5.1.1.3.

#### **2.5.1.1.2.4 Middle Silurian through Early Devonian**

During the Middle Silurian through Early Devonian Time the site region (200 mi (322 km) radius) existed as part of an extensive basin of deltaic, reef, alluvial and shallow marine environments (Ver Straeten and Brett, 2000, and Diedrich and Wilkinson, 1999). The Keefer, Bloomsberg, and Mifflintown Formations are the representative strata for this time. The eastern margin of the North American craton was passive while the eroding Taconic Highlands continued to load sediment into the Appalachian Basin.

#### **2.5.1.1.2.5 Acadian Orogeny**

The Acadian Orogeny began at the onset of the Middle Devonian within the site region (200 mi (322 km) radius), the result of the Avalon Terrane and Baltica (Eusden, et al, 2000) converging with the North American craton. The Acadian Orogeny affected the northern Appalachians of New England in the late Silurian (Bradley and Tucker, 2001). The site vicinity (25 mi (40 km) radius) remained within the Appalachian basin area while the Acadian mountain range, to the east of the site area, was subjected to erosional processes. These eroded sediments were deposited in the site area and are represented by the modern day black and gray shales of the Marcellus formation that underlie the site (Ver Straeten and Brett, 1999). The effect of the Acadian Orogeny is more pronounced in terms of metamorphism and magmatism in the northeastern United States (Maine, New Hampshire, Vermont), but is significant to the site area in that the source material for the formation of the Devonian rocks, notably the Marcellus and Mahantango Formations, that outcrop in the site area was primarily derived from the Acadian Highlands. (Ver Straeten and Brett, 1999) identify the progression of a terrigenous clastic wedge from the orogenic belt to the west during multiple phases of orogenic uplift during the Devonian that was the source of Devonian deposition (Ver Straeten and Brett, 1999). Basin flexure and bulging produced local sea level and depositional environment changes that resulted in reef structures that are noted in the Onondaga Formation which outcrops within the site area (5 mi (8 km) radius). The Acadian Orogeny and syn-orogenic deposition within the BBNPP site region (200 mi (322 km) radius) came to a close at the Middle Mississippian period. Upper Devonian and Lower Mississippian strata were deposited in the site vicinity (25 mi. (40 km) radius) but have been eroded by subsequent uplift during the Alleghanian Orogeny. The Middle Mississippian through Middle Pennsylvanian, approximately 350-300 Ma, (Hatcher, 1987) was an orogenic hiatus in which the Mauch Chunk and Pottsville Formations were deposited within the site region and site area. **The composition of the aforementioned Formations is provided in Section 2.5.1.1.3.**

#### 2.5.1.1.2.6 Alleghanian Orogeny

Convergence of Peri Gondwana with Laurentia at margins identified in modern geographic terms as northwest Africa and eastern North America led to the formation of the supercontinent of Pangaea. The continental collision caused the eastern and southern margins of North America to undergo uplift and deformation in what is referred to as the Alleghanian Orogeny. Late Pennsylvanian dextral transpression (a combination of convergent and transform plate boundaries) was the initial interaction of the aforementioned continents (Engelder and Whitaker, 2006). As the convergent margin evolved in the early Permian, intense brittle and ductile deformation in the form of thrusting, folding, and varying degrees of metamorphism took place (Steltenpohl, 1988) and (Schumaker, 2002). The current geologic setting along the eastern, southeastern, and south-central (although much is now buried under Gulf Coast Basin sediment) United States is strongly defined by Alleghanian deformation. In many cases, pre-existing faults related to Grenville, Taconic, or Acadian deformation were reactivated to develop regional detachment structures or decollements along which the Piedmont, Blue Ridge, and Ridge and Valley were transported as much as 180 mi (300) km to the northwest in the middle to late Permian (Engelder and Whitaker, 2006). This crustal shortening and overthrusting developed the structural setting of deep seated regional thrust faults, intense folding, and varying degrees of metamorphism that is prevalent in the Ridge and Valley, Blue Ridge, and Piedmont Physiographic provinces. Throughout the Alleghanian orogenic process the stress regime maintained a dextral or right lateral transpressional component (Ong et al, 2007) with the exception of Latest Permian dextral motion (Steltenpohl, 1988). This predominant dextral transpression is at least partially responsible for the oroclinal structures in the Ridge and Valley and Blue Ridge Provinces. A large potentially oroclinal component of the Ridge and Valley exists within the site region (200 mi (322 km) radius) that can be seen on Figure 2.5-6 as the change in regional structural fabric from a north east strike to a east-northeast strike. Change of structural orientation of 19 degrees to the east of dominant strike is measured in northeast Pennsylvania (Harrison et al, 2004).

Examples of Alleghanian deformation within the site vicinity (25 mi (40 km) radius) and site area (5 mi (8 km) radius) include the Berwick and Light Street Faults, depicted on Figure 2.5-27 and discussed in Sections 2.5.1.2.4.1 and 2.5.1.2.6.4. These faults are recognized as exhibiting reverse, to the northwest vergence and are classified as Alleghanian thrust faults (Inners, 1978). Field studies did not identify offset in terrace gravels overlying the Light Street and Berwick Faults (Inners, 1978). In addition, the Berwick Anticline, an east-northeast striking, gently northeast plunging anticline trends directly through the site area (5 mi (8 km) radius). The Berwick Anticline is an asymmetric structure in the site area (5 mi (8 km) radius) with the north-northwest limb dipping steeply to the north-northwest and the southern limb dipping more gently to the south-southeast. The orientation of this structure is classic Ridge and Valley Alleghanian deformation as presented by Hatcher (Hatcher, 1987). In addition to crustal deformation, the Alleghanian Orogeny had an important effect on the depositional regime in the Appalachian Basin and essentially closed the basin at the end of the Permian. Marked changes in Pennsylvanian rocks are a result of syn-orogenic flysch deposits derived from the uplifted continental margin (Thomas, 2004). The Pottsville, Conemaugh, Allegheny, and Monongahela Formations that lie within the site region (200 mi (322 km) radius) and exhibit lithologies that were influenced by Alleghanian derived detritus exemplify this relationship. Within the site area (5 mi (8 km) radius) the Pottsville and Llewellyn Formations, as shown on Figure 2.5-27 exhibit the Alleghanian sedimentary source. The Ridge and Valley Province within which the site area mostly lies, (5 mi (8 km) radius) is bounded to the northwest by the Allegheny structural front that separates the Appalachian Plateaus province from the Ridge and Valley. Surface expression of Alleghanian deformation within the Appalachian Plateaus Province is less intense than in the



Ridge and Valley. Alleghanian stress was accommodated differently in different lithologies, resulting in less pronounced deformation (more gentle folding and fewer thrust faults) within the surface strata (Schumaker, 2002). Deformation during the Alleghanian is most notable as uplift of the Appalachian Province and brittle deformation. The current structural setting on cross section view of the site region (200 mi (322 km) radius) is depicted on Figure 2.5-10a. The structural and physiographic position of the site area at the end of Pennsylvanian time placed it at elevations above depositional levels, and sub-aerial exposure and erosion ensued.

#### **2.5.1.1.2.7 Early Mesozoic Extensional Episode (Triassic Rifting)**

During the Late Triassic, at the onset of the breakup of Pangea, the eastern North American plate and African plate began to separate to create the Atlantic Ocean. A series of rift basins, such as the Gettysburg-Newark basin developed in southeastern Pennsylvania and along the North American coastline, respectively in what is referred to as the North American Rift System (Schlische, 2002). The rift basins are arranged primarily in northeast-southwest asymmetric trend and are located from South Carolina through Massachusetts (USGS, 1985). Normal faulting under the extensional regime often occurred along pre-existing Paleozoic structures (Olsen and Schlische, 1990). Subsequently, the basins were filled with sediments such as conglomerates, sandstones and shales and exhibit evidence of syn-rift deposition (lessening of offset upward in the basin deposits) (Schlische, 2002). The Culpepper, Gettysburg, and Newark Basins lie within the site region (200 mi (322 km) radius) and are shown on Figure 2.5-15d. As basin subsidence continued through the Triassic, the depositional environment within the basin became increasingly sub-aqueous. Outboard of the basins, carbonate platform deposition along the nascent continental margin occurred. Mantle derived basaltic intrusions occurred within the faulted crust which are evident in diabase dikes and sheets of the Piedmont Province of the eastern United States (Philpotts et al, 1985). During the early Jurassic period, the process of seafloor spreading caused deep-seated magma to approach the surface. Volcanic deposits ranging from 6-9.3 mile (10-15 km) thickness formed along the entire U.S Atlantic margin in the Middle Jurassic (Sheridan et al, 1993). The magma created the basalt located in the Gettysburg-Newark basin of the Piedmont province (Schlische, 2003). **Northwest-southeast-directed post rift activities in the Mesozoic basin caused inversion to many structures present during this time (Withjack, 1998).** Following rifting, subsidence, and volcanism, the Atlantic Margin became a passive margin. The structural and seismotectonic influence of Mesozoic rifting affects the site region (200 mi (322 km) with respect to potential seismogenic structures. These structures, such as the Ramapo Fault have exhibited seismogenic potential (Ratcliffe, 1971) and are discussed in further detail in section 2.5.1.1.4.

#### **2.5.1.1.2.8 Late Mesozoic and Cenozoic History**

During the Cretaceous Period of the Mesozoic Era, the site vicinity (25 mi (40 km) radius) and the majority of the site region (200 mi (322 km) radius) experience sub-aerial exposure and various rates of erosion based on climate. The notable exception is the Atlantic Coastal Plain Physiographic Region (Figure 2.5-7) that was inundated by the transgressive Cretaceous sea within which significant deposition took place. The extent of the Cretaceous transgressive sea was well inland of the current Atlantic Coastal Plain limit, but erosion during the Cenozoic Era moved the limit of the sedimentary Atlantic Coastal Plain to the east and southeast. The Atlantic Coastal Plain remained inundated during much of the Cenozoic Era as exemplified by Paleocene through Pleistocene deposits as shown on Figure 2.5-5. Inland of the shoreline from the Cretaceous Period to early Cenozoic Era, the majority of the site region (200 mi (322 km) radius) was subjected to chemical and physical weathering and erosion at varying degrees of intensity based on climatic conditions. The depositional setting that has existed throughout the

Cenozoic within the site vicinity (25 mi (40 km) radius) and site area (5 mi (8 km) radius) were topographic lows upon which alluvial and lacustrine deposits accumulated. The notable exception was the advance of glacial ice sheets that had significant effects in terms of exhuming and re-depositing surficial material on a regional scale and changing the loading and unloading conditions (which can initiate deformation in the rock mass) on the site vicinity (25 mi (40 km) radius) and site area (5 mi (8 km) radius).

From Pre-Illinoian to late Wisconsinan, three major glacial advancements occurred from ice accumulation in Canada advancing into Pennsylvania, and into the BBNPP site area. The glaciers were located at the northern portion of Pennsylvania and covered most of the Appalachian Plateaus province. The earlier glaciers migrated south approximately 770,000 years ago while the most recent occurred about 17,000-22,000 years ago (Barnes, 2002) and (Braun, 2007). Figure 2.5-9 shows the limit of glacial advance in Pennsylvania with the site vicinity (25 mi (40 km) radius) superposed. A map showing surficial glacial deposits within the BBNPP site (0.6 mi (1 km) radius) is presented on Figure 2.5-24.

The glacial advancements scoured valleys and deposited till, sand and gravel outwash material throughout BBNPP site area while the nearby Susquehanna River deposited sand and gravel outwash, filling the bottoms of valleys. During the period of elevated physical weathering, freezing and thawing at the surface caused the breakup of large quantities of rock at the crests of ridges in the Ridge and Valley province. As a result, the crests of these ridges were lowered by several feet. In addition, loose talus rock accumulated on the slopes of many ridges within central Pennsylvania.

The geologic history of the site region (200 mi (322 km) radius) is complex. The current geologic processes affecting the site are limited to weathering and erosion of existing material, and to the regional stress field that affects the passive Atlantic margin. Figure 2.5-10b shows the current stress fields in the eastern portion of North America. Minimal isostatic uplift that has remained relatively constant throughout the Cenozoic (Matmon et al, 2003) is currently a component of the geologic setting. With respect to seismic stability and geologic hazards relevant to the site vicinity (25 mi (40 km) radius) the site area (5 mi (8 km) radius), and site location (0.6 mi (1 km) radius) are positioned in a stable geologic setting.

#### **2.5.1.1.3 Regional Stratigraphy**

This section contains information and analysis on the regional stratigraphy within the major physiographic provinces in the Northeastern United States that are located within the BBNPP region. The regional geology and generalized stratigraphy within a 200 mi (322 km) radius of the BBNPP site is presented schematically in Figure 2.5-13a.

Reports and scientific journal articles, published primarily by staff of the U.S. Geological Survey (USGS) and the Pennsylvania Geological Survey, have been used to develop the following descriptions of regional and local geology. After a thorough search and evaluation of literature ranging from the 1930s to 2008, the references cited in this document are considered to be the leading and most authoritative references relative to BBNPP regional geology, stratigraphic nomenclature, and stratigraphic relationships. In some cases, the best sources of information (most complete and most detailed) for local geology were published in the 1970s and 1980s. In all cases, when more recent publications provide new information, new data, or reinterpretations of old concepts, the newer information has been used and has been cited.

### **2.5.1.1.3.1 Stratigraphy of the Ridge and Valley Physiographic Province**

#### **2.5.1.1.3.1.1 Pre-Cretaceous Basement Rock**

The crystalline basement rock underlying Pennsylvania is of Precambrian age (Saylor, 1999) and rarely exposed, except in the Piedmont Province of southeastern Pennsylvania (Figure 2.5-6). Estimated depth of this basement rock at the BBNPP site is presently at approximately 33,000 ft (10,058 m) below ground surface (bgs), as shown in Figure 2.5-11 and Figure 2.5-19. Due to the lack of exposure and the relatively great depth to Precambrian rocks in the Ridge and Valley Province, information on Precambrian basement rock is extrapolated from several exploratory wells in western Pennsylvania (Saylor, 1999; Gold, 2008). It is inferred from these deep wells that the Precambrian basement is approximately 1 billion years (Ga) old (Gold, 2008) and it is composed of metamorphosed greenschist or amphibolite. It is also inferred that this Precambrian basement is a regular, gently sloping surface, dipping eastward and forming the western margin of the Appalachian miogeosyncline (Saylor, 1999). Earliest deformation of this basement rock appears to have occurred during the Grenville Orogeny (Saylor, 1999), resulting in multiple folding events and faulting. Due to the heavily metamorphosed state of this Precambrian basement, little is known as to its depositional environment of the original sedimentary deposits.

The closest boreholes to the BBNPP site that penetrate the basement rock are located in Erie and Crawford counties, Pennsylvania, located about 200 mi (322 km) northwest of the BBNPP site (Figure 2.5-11). The borings that penetrate the underlying Precambrian basement in northwestern Pennsylvania, eastern Ohio, and northern West Virginia have encountered metamorphic or igneous rocks (Saylor, 1999). For example, a well labeled Temple No. 1 in Mercer County PA, located approximately 208 mi (335 km) west of the BBNPP site, was drilled into a biotite granite/quartz-biotite gneissic basement rock at 9,810 ft (2,990 m) depth (Saylor, 1999). Another well, labeled Fleck in Mercer County PA, located 205 mi (330 km) west of the BBNPP site, was drilled into basement rock at a depth of 9,136 ft (2,785 m) with rock composition including weathered chloritic schist and granite grading into gneiss (Saylor, 1999). The basement rock was only sampled in the drill cuttings and suggests a gneiss/schist from the mineralogy present (i.e., biotite, chlorite, and clear quartz).

Overlying the Precambrian metamorphic and igneous basement of the Ridge and Valley Province are the clastic sedimentary deposits of the Early Cambrian with a transition to the carbonate rich sediments of the Early Ordovician. These early Cambrian deposits created a wedge of terrigenous sediments, best described today as the Chilhowee Group, which were the result of marine waters of the Iapetus (Proto-Atlantic) Ocean slowly transgressing across the continent shortly after the Grenville Orogeny (Kauffman, 1999). Overlying these sediments is a carbonate platform (Bradley, 1989; Kauffman, 1999) showing signs of uplift and erosion during the Taconic Orogeny during the Ordovician (Bradley, 1989). Above the clastic sediments of the Chilhowee Group is the brown sandstone interbedded with red and green shale beds of the Waynesboro Group (Kauffman, 1999). The Waynesboro Group, according to Kauffman (Kauffman, 1999) is the oldest exposed outcrop in Central Pennsylvania with an Early to Middle Cambrian age (Figure 2.5-13a). Overlying the Waynesboro Group is a limestone formation identified as the Warrior Formation (Ryder, 1992) of Middle to Late Cambrian age. The lithology of the Warrior Formation is further defined by Kauffman (Kauffman, 1999) as a dark, fossiliferous, fine grained limestone interbedded with silty dolomite and has a thickness of up to 1,340 ft (408 m) in the Ridge and Valley Province. Bordering the Cambrian-Ordovician contact and overlying the Warrior Formation is the Gatesburg Formation. The Gatesburg Formation consists of a series of sequential sandstone and dolomite units and can be labeled as Late

Cambrian age through the identification of gastropod fossils in the uppermost member (Ryder, 1992).

Significant deposition of sediment occurred in the Appalachian Basin during the Ordovician Period, and these sediments are present throughout the Ridge and Valley Province from West Virginia up through Maryland eastward to northeastern Pennsylvania and New York (Figure 2.5-13a). The Ridge and Valley rocks of Ordovician age are primarily sedimentary in nature, with evidence of uplifting during the Taconic Orogeny. According to Thompson (Thompson, 1999), the Ordovician sedimentation can be broken down into three major phases with early Ordovician being a depositional environment of a stable carbonate-platform. During Middle Ordovician, there was a submergence of the carbonate-platform, due to the Taconic Orogeny, with marine limestone and siliciclastic sedimentation during the submergence (Thompson, 1999). This submergence resulted in the creation of a basin which was infilled with additional marine limestone and siliciclastic sediments (Thompson, 1999). Stratigraphically, Early Ordovician rocks are generally referred to as part of the Beekmantown Group (Harper, 2003), are composed primarily of dolomite-limestone, and reach a thickness of up to 4,200 ft (1280 m) (Thompson, 1999). The Middle Ordovician shows a transition zone from the dolomite-limestone to rocks of primarily limestone composition deposited in both shallow and deep-water environments (Thompson, 1999). In central-Pennsylvania, the Loysburg Formation best represents this transition from a tidal-zone to a shallow marine zone with a dolomitic and stromatalite rich limestone underlying a coarse grained, fossiliferous limestone (Thompson, 1999). It is also during the Middle Ordovician that the Iapetus Ocean stopped widening and began to close; meaning this formerly passive area of sedimentation became tectonically active, thus giving birth to the Taconic Orogeny (Cotter, 2008). This active margin setting became the depositional environment of the sandstone and greywacke-shales that comprise almost 3,500 ft (1,067 m) of Late Ordovician formations including the Juniata, Bald Eagle, and Reedsville Formations of central Pennsylvania.

During the early Silurian Period, shallow marine conditions returned to central Pennsylvania (Cotter, 2008) as it became a depositional environment for sediments being eroded and transported from the Taconic highlands in the eastern part of the state. The Silurian basement rocks throughout Pennsylvania have a thickness ranging from 3,000 ft (914 m) in central Pennsylvania to 4,000 ft (1,219 m) in northeastern Pennsylvania (Laughrey, 1999). The Silurian represents a transition from a coastal plain in the east to a delta in the west, through the alluvial clastic deposits of the Shawangunk and Tuscarora formations (eastern and central Pennsylvania respectively) to the offshore facies of the Medina Formation of western Pennsylvania (Laughrey, 1999). The Tuscarora Formation, prevalent throughout the Ridge and Valley Province in central Pennsylvania, is composed primarily of quartzose, sublithic, and argillaceous sandstones and shales (Laughrey, 1999) and ranges in thickness from 492 ft (150 m) to 656 ft (200 m). The Rose Hill, Keefer, and Mifflintown formations (in ascending order) best describe the stratigraphic members of the Middle Ordovician. Rose Hill Formation, which overlies the Tuscarora Formation, is defined as predominantly an olive shale with interbedded layers of hematitic sandstone, purplish shale, and fossiliferous limestone (Laughry, 1999). The Keefer Formation is described mainly as a quartzose and hematitic sandstone with some mudstone and the overlying Mifflintown Formation is composed of shallow marine mudrocks and limestones (Laughry, 1999). The Upper Silurian is identified by the Bloomsburg Formation, a grayish-red clay-siltstone with some interbedded sandstone, transitioning to the limestone and thin shale beds of the Tonoloway Formation (Laughrey, 1999).

In Pennsylvania, the Devonian-age rocks represent a "westward-thinning wedge of sediments" that range in thickness from 2,400 ft (732 m) in the western portion of the state, to over 12,000 ft

(3,658 m) in the east (Harper, 1999). These Devonian sediments are generally broken down into two basic groups: the Pre-Acadian Orogeny comprised of stable shelf sedimentary deposits and Post-Acadian Orogeny strata that emphasize the presence of "tectonism, subsidence, and filling of a foreland basin" (Milici, 2006). The Keyser Formation forms the base for the Devonian-age rocks of the Ridge and Valley Province in Pennsylvania. This formation is primarily composed of gray, fossiliferous limestone (Laughrey, 1999). Above the Keyser Formation lie other stages of the Lower Devonian, including (in ascending order) the cherty limestone of the Helderberg Stage, the quartz rich sandstones, shales and siltstones of the Deerpark Stage, and the detrital sediments of the Onesquethawan Stage (Harper, 1999). The Onesquethawan Stage carries into and becomes the basement for the Middle Devonian timeframe which consists of basinal marine shales to nonmarine sandstone. Other stages within the Middle Devonian Ridge and Valley Province include the fossiliferous shale of the Needmore Formation, the argillaceous and silty Selinsgrove Limestone, the volcanic Tioga ash and shales, and the Mahantango Formation (Harper, 1999). The Mahantango Formation, which comprises the bedrock of the BBNPP site, is described by Harper (Harper, 1999) as "a complex series of interbedded shales, siltstones, and sandstones ranging from 1,200 ft (366 m) to 2,200 ft (671 m) thick." Milici (Milici, 2006) also refers to the Mahantango Formation as silty shale. The fossiliferous shaley limestone of the Tully Limestone Formation forms the uppermost portion of the Mahantango Formation (Harper, 1999).

The marine and non-marine rocks of the Upper Devonian Period represent sediment deposition during the progradation of the Catskill deltaic system (Harper, 1999). This system, as it relates to the Ridge and Valley Province in central Pennsylvania, can be "broadly defined" by four main depositional episodes including (in ascending order) the rarely fossiliferous basinal shales of the Harrell Formation, the interbedded shales, siltstones, and sandstones of the Brallier Formation (equivalent to the Trimmers Rock Formation), the shales, thin siltstone, sandstones, and conglomerates of the Scherr and Lock Haven Formations, and the nonmarine sandstones and mudrock that overlap the Devonian-Mississippian boundary (Harper, 1999).

The Mississippian Period of the Ridge and Valley Province is a topic of on-going research but the most commonly accepted 'boundary' between the Mississippian and Upper Devonian is the Spechty Kopf Formation (Berg, 1999). The Spechty Kopf Formation, which ranges up to 1,280 ft (390 m) in thickness, is typically associated with the unconformity lying between the previously discussed Catskill Formation and the fluvial sandstones of the Pocono Formation (Berg, 1999). The Spechty Kopf Formation is predominantly sandstone with some interbedded shale and siltstone. Above it lies the Pocono Formation which, in north-eastern Pennsylvania, consists mainly of non-red medium to coarse-grained sandstones and conglomerates (Brezinski, 1999). In central Pennsylvania, the Pocono Formation is better represented by the Huntley Mountain and Rockwell Formations which are characterized by greenish-gray to tan sandy siltstone and silty shale with some sandstone (Brezinski, 1999). The red shales, sandstones, and conglomerates of the Mauch Chunk Formation (Van Diver, 1993) mark the original uplifting of the Alleghanian Orogeny, as well as the uppermost boundary of the Mississippian strata in the Ridge and Valley Province. The Mauch Chunk Formation varies in thickness throughout the state, but is generally between 3,000 ft (914 m) and 4,000 ft (1,219 m) thick (Brezinski, 1999).

Above the Mauch Chunk Formation in northeastern and central Pennsylvania lies the Lower Pennsylvanian age Pottsville Formation, which ranges in thickness from 100 ft (30 m) to 1,600 ft (488 m) and is composed mainly of conglomerate and conglomeratic sandstone, with some sandstone and coal (Edmunds, 1999). Overlying the Pottsville Formation and marking the boundary between the Pennsylvanian and Permian Periods is the Llewellyn Formation. The Llewellyn Formation has a thickness of up to 3,500 ft (1,067 m) and consists mainly of

conglomerates and sandstones with numerous coal beds and some clayey shale (Edmunds, 1999).

As shown in Figure 2.5-6, there are no post-Pennsylvanian age rock units present in the Ridge and Valley Province of Central and Northeastern Pennsylvania. Subsidence and deposition in the northern end of the Appalachian Basin ended in the Pennsylvanian Period and the beginning of the Alleghanian Orogeny. No deposition of geologic materials occurred in the Ridge and Valley Province between the Pennsylvanian Period and the Pleistocene glacial epochs. During this long time period, the Paleozoic rocks were subjected to weathering and erosion processes.

#### **2.5.1.1.3.1.2 Tertiary and Quaternary Deposits**

No Tertiary or pre-Pleistocene Quaternary deposits have ever been identified in the Ridge and Valley Province.

Pleistocene glaciers advanced over the northeastern end of the Ridge and Valley Province and covered the bedrock with variable thicknesses of glacial deposits, including tills, kame terraces, eskers, and outwash. The thickest glacial deposits include kame terraces and outwash which was deposited along major valleys leading away from retreating glaciers. These deposits can be more than 100 feet (30 m) thick and contain boulders, gravel, and coarse sand.

#### **2.5.1.1.3.2 Stratigraphy of the Piedmont Physiographic Province**

There are three distinct sections that comprise the Piedmont Physiographic Province. The first is the Gettysburg-Newark Lowland, the second is the Piedmont Lowland, and the third is the Piedmont Upland (Figure 2.5-8 shows the locations of these sections).

##### **2.5.1.1.3.2.1 Gettysburg-Newark Lowland Section**

The Gettysburg-Newark Lowland Section forms a 140-mile (225-km) arc across southeastern Pennsylvania with a series of exposed rift basins of Late Triassic to Early Jurassic age that are filled with fluvial and clastic sediments. Sediments filling the basins include conglomerates, shales, siltstones, and sandstones. These basins are underlain by nonmetamorphic Cambrian and Ordovician basement rocks and are bordered "by a continuous, complex system of normal faults" (Root, 1999).

##### **2.5.1.1.3.2.2 Piedmont Lowland Section**

Mesozoic sedimentary rocks of the Piedmont Province occur primarily within the highly folded and faulted region of the Piedmont Lowland section (Figure 2.5-8). The sediments were deposited in a series of northeast-trending basins. Sediments filling the basins include conglomerates, shales, siltstones and sandstones, and basic igneous intrusive dikes, diabase, and lava flows (VADOT, 2008). The Lower Mesozoic sediments deposited in these basins usually are referred to as Triassic basin deposits, although the basins are now known to also contain Lower Jurassic rocks. The folding and faulting of this section, as well as lithologies, are very similar to those found in the Great Valley Section of the Ridge and Valley Province, where Cambrian quartzite and Precambrian gneiss are brought into contact with rocks as young as the lower Ordovician (Gray, 1999).

#### **2.5.1.1.3.2.3 Piedmont Upland Section**

Crystalline rocks primarily occur within the Piedmont Upland section of the Piedmont Province. The crystalline rocks consist of deformed and metamorphosed meta-sedimentary and meta-igneous rocks, with overlying saprolite (VDEQ, 2008). The rocks belong to a number of northeast-trending belts that are defined on the basis of rock type, structure and metamorphic grade and are interpreted to have formed along and offshore of ancestral North America (Pavlides, 1994).

Surficial sediments in the Piedmont Province consist of residual and transported material. The residual soils have developed in place from weathering of the underlying rocks, while the transported material - alluvium and colluvium - has been moved by water or gravity and deposited as unconsolidated deposits of clay, silt, sand, and gravel. Surficial sediments in the Piedmont Upland section are interpreted to be the product of Cenozoic weathering, Quaternary periglacial erosion and deposition, and recent anthropogenic activity (Sevon, 2000).

Residual soil in the Piedmont Province consists of completely decomposed rock and saprolite. Residual soils occur almost everywhere, except where erosion has exposed the bedrock on ridges and in valley bottoms. Saprolite comprises the bulk of residual soil in the Piedmont Province and is defined as an earthy material in which the major rock-forming minerals (other than quartz) have been altered to clay but the material retains most of the textural and structural characteristics of the parent rock. The saprolite forms by chemical weathering, its thickness and mineralogy being dependent on topography, parent rock lithology, and the presence of surface and/or groundwater (Cleaves, 1992).

Relief affects the formation of soils by causing differences in internal drainage, runoff, soil temperatures, and geologic erosion. In steep areas where there is rapid runoff, and little percolation of water through the soil and little movement of clay, erosion is severe and removes soil as rapidly as it forms. Gently sloping areas, on the other hand, are well drained and geologic erosion in these areas is generally slight. The characteristics of the underlying rock strongly influence the kind of changes that take place during weathering. Because of differences in these characteristics, the rate of weathering varies for different rock types. The igneous, metamorphic, and sedimentary rocks of the Piedmont Province are all sources of parent material for the soils.

Colluvium in the Piedmont Province occurs discontinuously on hilltops and side slopes, while thicker colluvium occurs in small valleys lacking perennial streams. Alluvium is present in all valleys with perennial streams (Sevon, 2000).

#### **2.5.1.1.3.3 Stratigraphy of New England Physiographic Province**

The basement rocks of the Reading Prong Section of the New England Physiographic Province are believed to have formed during the Greenville Orogeny and are composed of metamorphosed sedimentary rocks. These rocks were then subjected to the intense thrust faulting and continual folding associated with the Taconic Orogeny, thus creating a complex nappe megasystem (Drake, 1999). Continued folding and faulting during the Alleghanian Orogeny has lead to "extremely complicated geologic relations" (Drake, 1999) within the Reading Prong. The Middle Proterozoic carbonate and crystalline rocks that were transported overtop of the basement rocks (Drake, 1999), were also subjected to folding and faulting and range in sequence depending upon the area of the Reading Prong being studied. Seismic-reflection studies have suggested that the basement of the Reading Prong ranges in thickness

from 15,000 ft (4,572 m) in the easternmost part of the Pennsylvania, to 45,000 ft (13,716 m) in Lebanon and Lancaster Counties (Drake, 1999).

#### **2.5.1.1.3.4 Stratigraphy of Atlantic Coastal Plain Physiographic Province**

The Atlantic Coastal Plain Physiographic Province is one of the flattest of the many physiographic provinces within the site region. The province covers more than 3,200 mi (5,150 km) from Cape Cod to the Yucatan Peninsula, and forms the continental shelf along the Atlantic Ocean (Komor, 1998). The province represents repeated cycles of transgression and regression of the ocean resulting in over 100 million years of sediment accumulation (Komor, 1998). Underlying most of the Atlantic Coastal Plain Physiographic Province are sediments of Cretaceous and Tertiary age with Pleistocene fluvial sediments overlying areas in and around the part of the province within the state of New Jersey. These Cretaceous sediments, in addition to glacial outwash deposits from the Pleistocene continental glaciers, comprise the underlying geology of Long Island and the eastern shores of Staten Island (Komor, 1998). The total sediment accumulation in the Atlantic Coastal Plain Province is nearly 30,000 ft (9,144 m) thick (Komor, 1998).

#### **2.5.1.1.3.5 Stratigraphy of Appalachian Plateau Physiographic Province**

The Appalachian Plateau Physiographic Province is underlain by rocks that are continuous with those of the Ridge and Valley Province but, in the Appalachian Plateau, the layered rocks are nearly flat-lying or gently tilted and warped, rather than being intensely folded and faulted. Rocks of the Allegheny Front along the eastern margin of the province consist of thick sequences of sandstone and conglomerate, interbedded with shale, ranging in age from Devonian to Pennsylvanian. The stratigraphic formations in the Appalachian Plateau are nearly identical to the strata in the Ridge and Valley Province. In southwest Pennsylvania (Figure 2.5-6) and West Virginia, however, the Appalachian Plateau Province includes sedimentary rocks of Permian age (sandstones, shales, and coal seams) (Hack, 1989).

The Appalachian Plateau Province is the only one of the provinces described above that has rock strata and glacial deposits that are similar to the geologic formations found at the BBNPP site. The rocks of these two provinces (Appalachian and the Ridge and Valley) were deposited in the same general Appalachian structural basin (miogeosyncline) and the sedimentary formations are continuous across the two provinces. On the other hand, there are great differences in geologic age, in structural and tectonic features, and sediment provenance between the Ridge and Valley geologic formations and the geologic materials found in the other provinces described above (except for the Appalachian Plateau Province).

#### **2.5.1.1.4 Regional Tectonic Setting**

The seismotectonic framework of a region, which includes the basic understanding of existing tectonic features and their relationship to the contemporary stress regime and seismicity, forms the foundation for assessments of seismic sources. In the probabilistic seismic hazard study performed by EPRI (EPRI, 1986), seismic source models were developed for the Central and Eastern United States (CEUS) based on tectonic setting; the identification and characterization of "feature-specific" source zones; and the occurrence, rates, and distribution of historical seismicity. The EPRI source model development included the independent interpretations performed by six Earth Science Teams. The seismic source models developed by each of the six teams were based on the tectonic setting and the occurrence, rates, and distribution of historical seismicity. The original seismic sources identified by EPRI (EPRI, 1986) are



thoroughly described in the EPRI study reports (EPRI, 1986), and are summarized in Subsection 2.5.1.1.4.3.6.

The EPRI models reflected the general state of knowledge of the geosciences community in the mid-to-late 1980's.

Since the EPRI study, additional geologic, seismologic, and geophysical research has been performed in the site region. This subsection presents a summary of the current state of knowledge in the regional tectonic setting and highlights the more recent information that is relevant to the identification of seismic sources for the BBNPP site. The following subsections describe the region in terms of:

- Contemporary Plate Tectonic Setting of the Atlantic Margin (subsection 2.5.1.1.4.1)
- Origin and Orientation of Contemporary Tectonic Stress (subsection 2.5.1.1.4.2)
- Gravity and Magnetic Data and Features of the Site Region and Site Vicinity (subsection 2.5.1.1.4.3)
- Regional Tectonic Structures (subsection 2.5.1.1.4.4)
- Seismic Sources as Interpreted by EPRI Groups (Subsection 2.5.1.1.4.5)

The BBNPP site region is a stable continental region (SCR) characterized by low rates of crustal deformation and no active plate boundary conditions (EPRI, 1986). Nearly 70% of SCR earthquakes with magnitude 6 or greater occurred in areas of Mesozoic and Cenozoic extended crust. Additional evidence shows an association between Late Proterozoic rifts and modern seismicity in eastern North America. There is no evidence for late Cenozoic seismic activity of any tectonic feature or structure in the site region (Crone, 2000) (Wheeler, 2005). No new structures or features have been identified in the site region since 1986 as described in the EPRI study (EPRI, 1986), but recent studies have provided a better understanding of the previously identified tectonic features and their associated seismicity. Among these features are the Ramapo fault and the Reading Prong (Sykes, 2008). These later studies were reviewed as input to this study, and are referenced throughout the subsections.

#### **2.5.1.1.4.1 Contemporary Plate Tectonic Setting of the Atlantic Margin**

The Late Precambrian to recent geologic history and plate tectonic evolution of the site region is summarized in Section 2.5.1.1.2. Several recent studies have concentrated on the relationship between the stratigraphy and structure during the Paleozoic era as it relates to orogenies and plate tectonics (Pazzaglia, 1994) (Pohn, 2000 and 2001) (Hibbard, 2006) (Cotter, 2008). These three main orogenic phases during the Paleozoic: Taconic, Acadian, and Alleghanian Orogenies, begin with accumulation of marine sediments and volcanic deposits, followed by deformation by structural folding and faulting, and end with tectonic uplift of mountains, and erosion of uplifted land. A particular consequence of orogeny is the production of sediment as uplifted mountains erode. Thus, each phase created a delta, filling shallow seas on the continental side of the orogeny. Clastic fans were deposited in terrestrial, coastal, near-shore, and off-shore settings (Aber, 2001).

During the break up of Pangaea in the Middle Triassic, rift basins developed in eastern North America. The rift basins were typically asymmetrical and trended northwest to southeast as the current Atlantic passive continental margin has evolved since rifting initiated in the Early

Triassic. The progression from active continental rifting to sea-floor spreading and a passive continental margin included: (1) initial rifting and hot-spot plume development, (2) thinning of warm, buoyant crust with northwest-southeast extension, normal faulting and deposition of synrift sedimentary and volcanic rocks, and (3) cooling and subsidence of thinned crust and deposition of postrift sediments on the coastal plain and continental shelf, slope, and rise (Klitgord, 1988) (Klitgord, 1995). The transition between the second (rifting) and third (drifting) phases during the Early Jurassic marked the initiation of a passive margin setting for the site region, in which active spreading migrated east, away from the margin.

The continental margin moved away from the spreading center of the mid-Atlantic and horizontal northwest-southeast tension changed to horizontal compression as gravitational potential energy from the spreading ridge exerted a lateral "ridge push" force on the oceanic crust. Northwest-southeast directed post rift activities in the Mesozoic basin caused inversion to many structures present during this time (Withjack, 1998). The transition from a rift to a drift margin through the remainder of the Mesozoic and into the Cenozoic along with the westward push of the continental divide, dominated the tectonic and geomorphic development of the eastern United States up to the modern time period (Schlische, 2003).

Contractional post-rift deformation is interpreted to record the change in stress regime from horizontal maximum extension during rifting to horizontal maximum compression during passive margin drifting. The hypothesis that the change in stress regime following rifting was known prior to the 1986 EPRI study (e.g., (Sanders, 1963) (Swanson, 1982) (Wentworth, 1983)). However, significant advances in the characterization of the rift to drift transition and post-rift deformation have occurred since then (Withjack, 1998) (Schlische, 2003).

Latest tectonic processes in the Cenozoic Era include vertical tectonics associated with lithospheric flexure (Pazzaglia, 1993). Vertical tectonics are dominated by cooling of the extended continental, transitional, and oceanic crust as the spreading center migrated eastward, and the erosion of the Appalachian Mountains to the Coastal Plain and extension of the Continental Shelf and Slope.

Based on models of the Cenozoic flexural deformation, surface material from the Appalachian Mountains eroded and was deposited on the Coastal Plain and Continental Shelf (Pazzaglia, 1993). The sediment is mainly deposited in the Salisbury Embayment and Baltimore Canyon Trough (Figure 2.5-10a). The Fall Line is the axis for the depositional downward pressure and the uplift from the erosional environment. An elastic model predicts the uplift in the Piedmont province to be as much as 33 ft (10 m) per million years (Pazzaglia, 1994). Figure 2.5-10a illustrates present conditions with location map and composite cross section of the middle U.S. Atlantic passive margin (Pazzaglia, 2004). The Susquehanna River terrace profiles, the Coastal Plain stratigraphic sections, geodynamic model cross-sections, and offshore load volumes are defined along cross section A-A', in the lower part of the same Figure 2.5-10a.

#### **2.5.1.1.4.2 Origin and Orientation of Contemporary Tectonic Stress**

Since 1986, an international effort to collate and evaluate stress indicator data has resulted in the publication of a new World Stress Map in 1989 that is periodically updated (Heidbach, 2008). The World Stress Map (WSM) is the global repository for contemporary tectonic stress data from the Earth's crust. It was originally compiled by a research group headed by Mary Lou Zoback as part of the International Lithosphere Program (ILP). Since 1995 the WSM is a research project of the Heidelberg Academy of Sciences and Humanities. The WSM research team is integrated into the Tectonic Stress Group of the Geophysical Institute at the Karlsruhe

University. The WSM is a task group of the International Association of Seismology and Physics of the Earth's Interior (IASPEI) (Heidbach, 2008).

As observed in Figure 2.5-10b (Heidbach, 2008) throughout the BBNPP site region, the average P axes of the earthquakes are oriented NE-SW and are aligned with the direction of the current plate driving stress that in the BBNPP region is characterized by northeast-southwest-directed horizontal compression. Since the 1986 EPRI evaluation of plate tectonic stress for the CEUS several authors have studied the stress magnitudes in the crust, deformation of intraplate lithosphere, joints and veins in regional tectonic perspective, and contemporary stress variations (Herman, 2005) (Zoback, 2002 and 1991) (Engelder, 2001) (Hancock, 1989) (Evans, 1989) (Hickman, 1985). As indicated in these subsequent studies, the tectonic stress created by the Mid Atlantic ridge forced stress orientation to align northeast to southwest on the North American plate. They also indicated that the lithosphere cannot be deformed more rapidly because of the limited amount of tectonic force available to drive that deformation in the current stress field (Zoback, 2002) and that the stress field for the site region is considered stable. Other potential forces acting on the North American plate are considered minor stress levels of magnitude and orientation.

Study of neotectonic joint attributes by Hancock (Hancock, 1989) and Engelder (Engelder, 1980) that included the Appalachian Plateau and other places in England, France, the Arabian Platform and the Ebro basin in Spain, showed that late-formed joints have potential value for tracking the contemporary stress field in regions where in-situ measurements are not available. Their conclusion showed that these joints have the characteristics of neotectonic joints and approximately parallel to directions of contemporary horizontal maximum stress (SH) known from in-situ stress measurements. The latter shows that although there may be a slight misalignment between joint strike and the direction of the greatest horizontal stress, these joints reflect the contemporary stress field.

Current understanding of the tectonic stress in the CEUS based on published reports since the publication of the EPRI source models in 1986 show only slightly different localized characterization of the northeast-southwest orientation of the maximum compressive principal stress (Hancock, 1989). The current characterization of potential activity of tectonic structures for the BBNPP Site remains valid.

#### **2.5.1.1.4.3 Gravity and Magnetic Data and Features of the Site Region and Site Vicinity**

Gravity and magnetic anomaly datasets of the site region have been published since the 1986 EPRI study. Regional maps of the gravity and magnetic fields are presented for North America by the Geological Society of America (GSA), as part of the Society's Decade of North America Geology (DNAG) project (Tanner, 1987) (Hinze, 1987) as shown in Figure 2.5-17 (Kucks, 1999) and Figure 2.5-18 (Bankey, 2002).

These maps present the potential field data at 1:5,000,000-scale, and show gravity and magnetic anomalies with wavelengths. Regional gravity anomaly maps are based on Bouguer gravity anomalies onshore and free-air gravity anomalies offshore. The primary sources of magnetic data reviewed for this BBNPP study are from aeromagnetic surveys onshore and offshore (Kucks, 1999). Large-scale compilations (1:2,500,000-scale) of the free-air anomalies offshore and Bouguer anomalies onshore were published in 1982 by the Society of Exploration Geophysicists (Lyons, 1982) (Sheridan, 1988). The DNAG magnetic anomaly maps were based on a prior analog map of magnetic anomalies of the U.S. published in the early 1980's (Zietz, 1982) (Behrendt, 1983) (Sheridan, 1988).

In addition, the DNAG Continent-Ocean transect program published a synthesis of gravity and magnetic data with seismic and geologic data (Klitgord, 1995). No gravity and magnetic data published since 1986 reveal new anomalies related to geologic structures. The following sections discuss the gravity and magnetic anomalies.

#### **2.5.1.1.4.3.1 Gravity Data and Features**

Gravity data compiled at 1:5,000,000-scale for the DNAG project provide documentation of previous observations that the gravity field in the site region is characterized by a long-wavelength, east-to-west gradient in the Bouguer gravity anomaly over the continental margin (Kucks, 1999) (Figure 2.5-17). Bouguer gravity values increase eastward from about -80 milligals (mgal) in the Ridge and Valley Province of western Virginia to about +10 mgal in the Coastal Plain Province (Figure 2.5-17). Gravity highs, or positive anomalies, are created by accumulations of dense rock units while gravity lows are from mass deficiencies. The folded and faulted structures, basins, igneous intrusions, lithologic variations, and basement uplifts create variations in mass. Gravity anomalies occur from density contrast in size, depth, and structural depth. Long wavelengths show shallow structures or highly concentrated deep structures. Shorter wavelengths are created by shallower structures (Lavin, 1999). As shown on (Figure 2.5-17), gradient gravity extends from Canada to Alabama and parallels the Appalachian Mountains. The Mesozoic rift basins show gravity lows and northeast-trending border faults.

The gravity map also shows northeast-trending, long wavelengths of gravity highs and lows. The alignments are variations of thickness of the sedimentary rocks and crustal structures (Lavin, 1999). Low gravity dominates the western part of Pennsylvania and eastern Ohio, including areas such as Beaver Falls gravity lows and Somerset gravity high. The Chambersburg anomaly is another low, broad, northeast-trending gravity low which extends the length of the Appalachian Mountain system. In the northeast, the Scranton gravity high is surrounded by the Williamsport and Reading lows. The lows are deep Paleozoic sedimentary basin and/or increased crustal thickness. The Scranton gravity high is related to mafic material during late Precambrian rifting (Lavin, 1999). All anomalies were known at the time of the 1986 EPRI study. In summary, gravity data published since the mid-1980's confirm and provide additional documentation of previous observations (i.e., pre-EPRI) across this region of eastern North America, and do not reveal any new anomalies related to geologic structures previously unknown to EPRI (EPRI, 1986) that would impact the BBNPP site.

#### **2.5.1.1.4.3.2 Magnetic Data and Features**

Magnetic data compiled for the 2002 Magnetic Anomaly Map of North America reveal numerous northeast-southwest-trending magnetic anomalies, generally parallel to the structural features of the Appalachian orogenic belt (Bankey, 2002) (Figure 2.5-18). The magnetic map allows a visualization of the geological structure of the upper crust in the subsurface showing the spatial geometry of bodies of rock and the presence of faults and folds. Prominent north- to northeast-trending magnetic anomalies in the BBNPP site region include the interior New York-Alabama Lineament, the New Bloomfield high, subsurface nappes near Scranton and Allentown, anomalies over largely subsurface Proterozoic rocks at the Reading Prong, Philadelphia and Lancaster, and an inferred basement fault located south of Pittsburgh (King, 1999). The 1,000 mi (1,609 km) long lineament in aeromagnetic maps of the eastern U.S. is referred to as the "New York-Alabama Lineament" (NY-AL) (Figure 2.5-18). The NY-AL primarily is defined by a series of northeast-southwest trending linear magnetic anomalies in the Ridge and Valley province of the Appalachian fold belt. The NY-AL is located about 50 mi (80 km) northwest of the BBNPP site. Based on studies by King (King, 1999), the NY-AL lineament divides the

basement into two magnetically distinct areas (Figure 2.5-18). To the southeast, the few anomalies present are very broad and have gentle gradients consistent with the profound basement depths of the region. To the northwest, numerous anomalies indicate a basement composed of large units of rock with strongly contrasting magnetic properties. King (King, 1999) has interpreted the NY-AL to be a major strike-slip fault in the Precambrian basement beneath the thin-skinned, fold-and-thrust structures of the Ridge and Valley province and created a base model for the Appalachian fold belt.

The Clingman-Ocoee lineament is an approximately 750 mi (1,207 km) long, northeast-trending aeromagnetic lineament that passes through parts of the Blue Ridge and eastern Ridge and Valley provinces from Alabama to Pennsylvania (King, 1999). The Clingman-Ocoee lineament is sub-parallel to and located about 30 to 60 mi (48 to 97 km) east of the NY-AL. These lineaments are located about 50 mi (80 km) southeast of the BBNPP site. The Clingman-Ocoee lineament is interpreted to represent a source or sources in the Precambrian basement beneath the accreted and transported Appalachian terrains (Nelson, 1983). The Clingman-Ocoee block is a Precambrian basement block bounded by the NY-AL and Clingman-Ocoee lineaments (Johnston, 1985b).

The Newark and Gettysburg rift basins consist of clastic rocks. The basins present magnetic anomalies consisting of elongated shaped bodies of diabase. The Mesozoic rocks have been downfaulted against Proterozoic and Paleozoic rocks (King, 1999).

The Buckingham Mountain anomaly is produced by a fault-bound structure creating a northeast trending ridge, and dividing the Newark basin. The faults cut the Mesozoic rocks and bound small diabase sheets on the north, just as the larger sheets are bounded along the northern boundary fault. The Buckingham magnetic high indicates a large subsurface ridge of magnetic Proterozoic rocks extending 15 mi (24 km) southwest (King, 1999).

The magnetic anomalies over the Reading Prong are produced by a complex of magnetite-rich, gneissic Proterozoic rocks at the surface (King, 1999). These rocks are related to the center of a nappe system that is over thrust from the southeast. Small anomalies occur east of Lancaster and are related to gneisses exposed in the Minde Ridge anticline and related structures. The magnetic data indicate similar rocks at shallow depths to the west toward Lancaster and to the east of the Honey Brook Upland, under the Triassic Basin (King, 1999).

In summary, magnetic data published since the mid-1980's confirm and provide additional documentation of previous observations (i.e., pre-EPRI) across this region of eastern North America, and do not reveal any new anomalies related to geologic structures previously unknown to EPRI (EPRI, 1986) that would impact the BBNPP site.

#### **2.5.1.1.4.4 Regional Tectonic Structures**

A tectonic map of the important structures in the BBNPP site region is shown in Figure 2.5-15, while cross sections are shown in Figure 2.5-10a. Since the EPRI study (EPRI, 1986) was completed, new tectonic features have been proposed and described in the site region, and previously described features have been characterized in more detail. New features identified since the EPRI study (EPRI, 1986) in the BBNPP site region include folds and faults (Pohn 2000) (Wheeler 2006). In the sections below, specific tectonic features and the evidence of their activity, or absence of activity are discussed. Specific descriptions are provided for those features within the BBNPP site region that could have an impact or produce a significant change in the EPRI seismic source model.

Table 2.5-62 contains a listing of the Tectonic Events that have affected the site region, the timing of each event, a brief description of the prevailing conditions during that event, the associated tectonic and non-tectonic features, and the section where these features are discussed. Some features are listed on the table multiple times, to acknowledge their prolonged activity or reactivation through geologic time.

Regional tectonic structures within and outside the 200 mi (322 km) BBNPP site region are described in this section, based on the following breakdown:

- Significant Seismic Sources Beyond 200 mi (322 km) radius that could impact the BBNPP site (subsection 2.5.1.1.4.4.1)
- Quaternary Tectonic Features (subsection 2.5.1.1.4.4.2)
- Relevant Tectonic Features with Associated Seismicity (subsection 2.5.1.1.4.4.3)
- Relevant Tectonic Features with No Associated Seismicity (subsection 2.5.1.1.4.4.4)

#### **2.5.1.1.4.4.1 Significant Seismic Sources Beyond 200 mi (322 km) Radius**

Description of the tectonic features begins with seismic sources significant to BBNPP, that are outside the BBNPP site Region, or greater than 200 mi (322 km) from the site. The following features are included in this subsection:

- The New Madrid Seismic Zone (NMSZ) (Figure 2.5-15a),
- The Charleston Seismic Zone (CSZ) (Figure 2.5-15b),
- The Charlevoix-La Malbaie Seismic Zone in Canada (Figure 2.5-15c), and
- The Saint Lawrence Rift Valley

The Madrid Seismic Zone (NMSZ), the Charleston Seismic Zone (CSZ), and the Charlevoix-La Malbaie Seismic Zone in Canada, are known contributors to the BBNPP seismic hazard at low frequencies, based on updated models with new paleoseismological data on characteristic earthquake events from these sources (see subsection 2.5.2 for further details). The Saint Lawrence Rift is not considered a contributor to the BBNPP seismic hazard. However, because it is considered a seismic source in EPRI (EPRI 1986), it is discussed here for completeness.

##### **2.5.1.1.4.4.1.1 New Madrid Seismic Zone (NMSZ)**

The New Madrid Seismic Zone and Reelfoot Scarp are located in the central part of the Mississippi River Valley, approximately 800 miles (1290 km) WSW of the BBNPP Site. This section discusses general details and characteristics of New Madrid Seismic Zone and Reelfoot Scarp because they are considered to be a contributor to the seismic hazard at the BBNPP site.

At present, structural and tectonic information about specific seismogenic faults is limited, in part because the seismogenic faults are not expressed or are poorly expressed at the surface. Furthermore, the entire river valley is covered by late Quaternary sediments, so only the geologically youngest deformation is expressed at the surface. The Reelfoot scarp is a topographic escarpment that extends south-southeastward from near the town of New Madrid, Missouri, along the western margin of Reelfoot Lake, to a point south of the lake. It is the most

prominent geomorphic feature in the entire seismic zone that is clearly known to have a tectonic origin (Figure 2.5-15a) (Csontos, 2008).

In the winter of 1811-1812, at least three major earthquakes occurred in the New Madrid seismic zone, and the area remains the most seismically active area in central and eastern North America. The 1811-1812 earthquakes were among the largest historical earthquakes to occur in North America and were perhaps the largest historical intraplate earthquakes in the world. The New Madrid earthquakes consisted of a series of at least 3 and possibly 4 major ( $M \geq 7.5$ ) events during a period of 2 months in the winter of 1811 and 1812. Strong aftershocks persisted in the region for at least one year. The first major earthquake occurred at 2:15 a.m. (all times are local times) on December 16, 1811. It was followed by another major earthquake at 8:15 a.m. the same day that was the smallest of the 4 major events. Most reports of the New Madrid earthquakes note three principal events, and this event is commonly not cited as one of the principal earthquakes. The next major event occurred at 9 a.m. on January 23, 1812. Historical accounts suggest that this event was intermediate in size between the first and last major shocks. The last and largest earthquake occurred at 3:45 a.m. on February 7, 1812 (Nuttli, 1972; Street and Nuttli, 1990).

The earthquakes produced widespread liquefaction throughout the seismic zone and prominent to subtle surface deformation in several areas, but they did not produce any known surface faulting (Crone, 2000). Other than the pervasive sand blows throughout the seismic zone, the Reelfoot scarp is the most prominent geomorphic feature that has been produced by the modern tectonism in the New Madrid Seismic Zone. Recent studies of the scarp have provided valuable information on the recurrence of deformation on the scarp, calculated uplift rates, and the history of faulting. Modern or historic earthquakes that induced liquefaction in the same region can serve as calibration events for paleoearthquakes. In the NMSZ, for example, comparison of the size and spatial distributions of historic and pre-1811 sand blows suggests source areas and magnitudes for paleoearthquakes similar to the very large to great earthquakes of 1811–1812 (Tuttle, 2001). Many features shown on maps of the New Madrid seismic zone (Figure 2.5-15a) are not discrete faults; the most notable exception is the Reelfoot scarp, which is located along the western shore of Reelfoot Lake in extreme northwestern Tennessee. The three most commonly noted features associated with the contemporary deformation are: (1) the Lake County uplift and Reelfoot scarp, (2) areas of suspected co-seismic subsidence, and (3) areas of abundant liquefaction during the 1811-1812 earthquakes. The locations of these features are derived from the digital data used to generate the suite of seismotectonic maps of the New Madrid, Missouri area. The modern seismicity in the New Madrid Seismic Zone is intimately associated with the Reelfoot rift (Braile, 1997), a northeasterly-trending, 43.5 mile (70 km) wide graben that has as much as 1.2 mi (2 km) of structural relief on magnetic basement. The rift is best defined by magnetic data, which also reveals the presence of major positive magnetic anomalies along the flanks and axis of the rift that are inferred to be mafic plutons (Braile, 1997).

Recent compilation in Crone (Crone, 2000) have shown that the Reelfoot scarp is approximately 19.9 miles (32 km) long and the subjacent Reelfoot fault may be as much as 43.5 miles (70 km) long. However, this fault is the only feature that has conspicuous surface expression and therefore can be studied at the surface. Based on the historical seismicity, there may be other significant but unexposed faults in the seismic zone.

The New Madrid seismic zone is not a single feature but is defined by the region of abundant seismicity in the central Mississippi River Valley. Accordingly, it is not possible to define an average strike, but most of the seismicity is associated with the Reelfoot rift, which has a

northeasterly trend. The Reelfoot scarp is a well-defined feature, but is small in comparison to the general dimensions of the seismic zone. The average azimuth of the Reelfoot scarp is 337°.

In late Quaternary time and probably in earlier episodes, tremendous volumes of glacial melt-water from much of North America flowed down the Mississippi-Ohio Rivers drainage system and through the northern Mississippi Embayment. Braided streams transported the melt-water deposited outwash sand and gravel into the embayment. These braided stream deposits are typically tens of meters thick in the New Madrid region. In early Holocene time, the Mississippi River changed from a braided stream to a meandering regime and began developing the modern meander belt. As a meandering river, fine-grained overbank sediment that was deposited as annual floods spread across wide expanses of the modern river valley.

The contemporary seismicity and current deformation in the New Madrid region is controlled by a regional stress field in which the maximum compressive stress is oriented approximately east-northeast to west-southwest (Van Arsdale, 1995 and 2000). Within this stress field ancient faults, most of which originally formed as extensional features during rifting, have been reactivated mainly as strike-slip faults. The modern seismicity is largely associated with rift-related features and is concentrated into three major trends that form a zigzag pattern that has an overall northeasterly trend. The southwestern-most trend is a narrow, linear, 74.6 miles (120 km) long zone of earthquakes in northeastern Arkansas and extreme southeastern Missouri; this zone of earthquakes roughly coincides with the position of an axial fault zone that is commonly present along the center of most rifts. Based on combined information from seismological, seismic-reflection profiling, geomorphic, and geological studies, the Reelfoot Scarp is interpreted as an east-dipping monocline which is the surface expression of a fault-propagation fold associated with the underlying blind Reelfoot thrust fault (Van Arsdale, 1995 and 2000).

The most widespread expression of recent strong earthquakes in the New Madrid region is the abundant liquefaction features (sand blows and sand-filled fissures), which are concentrated in a 24.9 to 37.3 mile (40 to 60 km) wide belt from near Charleston, Missouri on the northeast to south of Marked Tree, Arkansas (Obermeier, 1988). Geologic conditions in the New Madrid region are near optimum for the development of liquefaction features during strong earthquakes: a thin (6.6 – 26.2 feet (2-8 m) thick, fine-grained "topstratum" deposit overlies water-saturated, unconsolidated "substratum" sand and gravel. Extensive liquefaction occurred during the 1811-1812 earthquakes; locally the ground surface was buried by more than 3.3 feet (1 m) of liquefied sand, and hundreds of square kilometers of the land surface have been mapped as being more than 25 percent covered by liquefied sand (Obermeier, 1988).

Detailed studies of the Reelfoot scarp in northwestern Tennessee have documented evidence of three deformation events within the past 2,400 years and characterized the style of near surface deformation associated with the scarp (Kelson, 1996). Late Holocene fluvial deposits are warped into a 26.2 feet (8 m) high, east-facing monocline. Borehole data and trenches at three sites characterized the style of near-surface deformation associated with the scarp and constrain the timing of three deformation events on the scarp.

Deformation on the scarp associated with the 1811-1812 New Madrid earthquake sequence produced extensive liquefaction, folded the fluvial sediments, and caused minor reactivation of small faults that bound an extensional graben in the uplifted (hanging wall) of the Reelfoot scarp. The penultimate deformation event occurred between A.D. 1260 and 1650 (350-740 yr B.P.), produced about 4.3 feet (1.3 m) of throw in the graben bounding faults, and caused folding and development of the scarp. The oldest documented event associated with the scarp occurred between A.D. 780 and 1000 (1000-1120 yr B.P.), and initially produced the small



graben in the hanging wall of the Reelfoot fault. Liquefaction features of similar age to, but of smaller size than, the Daytona Beach sand blow occur near Blytheville (93 mi (150 km) northeast of Marianna) and Montrose, Arkansas (109 mi (175 km) southeast of Marianna). A very large ( $M > 7.2$ ) earthquake centered near Marianna about 3500 B.C. might account for the liquefaction in all three areas. The large sand blows at the St. Francis site are similar to compound sand blows in the NMSZ, suggesting that a New Madrid-type earthquake sequence was centered near Marianna about 4800 B.C (Tuttle 2006).

Paleoseismic studies have suggested a recurrence interval of about 500-1100 years for earthquakes that are large enough to produce significant surface deformation or liquefaction in various parts of the seismic zone, with most recent studies suggesting that there were about 900 years between the last two New-Madrid-size events (A.D. 900 to A.D. 1811) and that widespread liquefaction occurs every few hundred years (Crone, 2000). However, the record studied thus far is too short to be used for a long-term recurrence rate. Liquefaction data indicate that New Madrid events occurred every  $500 \pm 300$  years during the past 1200 years. Furthermore, this recurrence rate for very large earthquakes is not easily reconciled with the small amount of crustal deformation observed in the region, suggesting that the NMSZ became active during the Quaternary and that New Madrid earthquakes may be temporally clustered in this intraplate region (Tuttle 2002). The detailed investigations of the Reelfoot scarp described above provide information that permitted Kelson (Kelson, 1996) to estimate a recurrence interval of 150-900 years, with a more likely range of about 400-500 years. These rate estimates only apply to the Reelfoot scarp, and not to overall behavior of major events for the entire seismic zone.

Epicentral patterns, correlative geophysical data, and historical seismic energy release indicate the significance of New Madrid area seismicity, both within the Reelfoot segment of the rift structures and in areas outside of this segment, particularly to the north (Braile, 1997). It is also clear that deep structure of the crust, including thickness variations in the upper crust and the presence of a high-density lower crustal layer, is a controlling factor in New Madrid seismicity.

Although there are many unanswered questions about New Madrid area seismicity and the causes of earthquakes in this intraplate setting, much has been learned in the past several decades, particularly from the detailed epicentral patterns that are beginning to be illuminated by the seismograph network data and from the geological and geophysical studies (Braile, 1997) of the shallow and deep (crustal and upper mantle) structures within the NMSZ.

Calculations performed for the PSHA for BBNPP determined that the New Madrid Seismic Zone is a contributor to the seismic hazard at the BBNPP Site in the low-frequency range.

#### **2.5.1.1.4.4.1.2 Charleston Seismic Zone (CSZ)**

The Charleston, South Carolina, Seismic Zone is located 590 miles (950 km) SSW from the BBNPP Site. The Charleston region is considered to have the highest seismic hazard of the east coast of the United States (Jaume, 2006). The CSZ is considered a contributor to the hazard at the BBNPP Site at low frequencies.

The August 31, 1886 Charleston earthquake (magnitude 6.9-7.3) was the largest historical event to strike a well-populated region of the eastern United States (Jaume, 2006). Historical evidence for felt earthquakes in the region before the 1886 event and continuing low-level seismicity in the epicentral area, plus evidence for multiple earthquakes producing liquefaction features in the South Carolina coastal plain, result in this region having the second highest seismic hazard in the continental United States east of the Rocky Mountains (Jaume, 2006).

Geomorphic, geologic, and geophysical data suggest the presence of an approximately 373 mi (600 km) NNE trending buried fault system in the Coastal Plain of the Carolinas and Virginia, named the East Coast Fault System (ECFS), illustrated in Figure 2.5-15b (Marple, 2000). This fault system is expressed by anomalous changes in fluvial geomorphology that locally coincide with one or more than the following features: linear aeromagnetic anomalies, buried faults interpreted from seismic reflection data, surface faults offsetting Pliocene-Pleistocene surficial units, locally brecciated phyllites and argillites, gently up-warped sediments, topographic highs, and seismicity near Summerville, South Carolina.

The suggested fault system traverses the epicentral area of the 1886 Charleston earthquake and lies west of paleoliquefaction sites along the outer South Carolina Coastal Plain. Therefore, this fault system could be the source of the Charleston Earthquake and other large prehistoric earthquakes. Although most features described individually do not provide conclusive evidence of faulting, collectively they strongly suggest the presence of the East Coast Fault System (Marple, 2000).

The shallow subsurface Tertiary stratigraphy of the greater Charleston region reflects the tectonic development and setting of the area over the past 34 Ma. Upper Eocene and Oligocene stratigraphic horizons show a net regional dip toward the southwest or south, whereas Miocene and Pliocene horizons show a shift to net regional dips toward the southeast (Weems, 2002). The tendency for many stratigraphic horizons to pinch out northeastward within the Charleston area indicates that it has been a persistent hinge zone between the Southeast Georgia embayment to the southwest, and the Cape Fear arch to the northeast. Localization of relative strain and motion along this hinge zone probably explains why the Charleston area is much more active seismically than other areas within the Atlantic Coastal Plain (Weems, 2002).

Results from a comprehensive loss assessment for the state of South Carolina (Wong, 2005) indicate that a future repeat of the 1886 earthquake would be catastrophic, resulting in possibly 200 deaths, more than 44,000 injuries, and a total economic loss of \$20 billion in South Carolina alone. By characterizing the nature and scope of potential impacts, the study performed by Wong (Wong, 2005) for South Carolina represents a starting point for renewed and hopefully more effective efforts in earthquake hazards and loss mitigation for this, distant to the BBNPP site, capable source.

Calculations performed for the PSHA for BBNPP determined that the Charleston Seismic Zone is a contributor to the seismic hazard at the BBNPP Site in the low-frequency range.

#### **2.5.1.1.4.4.1.3 Charlevoix-La Malbaie Seismic Zone**

The Charlevoix-La Malbaie Seismic Zone (CSZ) area is located 530 mi (853 km) NNE of the BBNPP Site, and 62 mi (100 km) NE of Quebec City, on the north shore of the St. Lawrence River (Figure 2.5-15c) (Tremblay, 2003). It is considered a contributor to the hazard at the BBNPP Site at low frequencies.

The bedrock of the area mostly consists of metamorphic rocks of the Precambrian crystalline basement (Lemieux, 2000). However, along large segments of the shoreline, as well as in two river valleys ("Du Gouffre" and Malbaie rivers), Paleozoic rocks of the St. Lawrence Lowlands are well exposed and unconformably overlie the Grenvillian basement (Lemieux, 2000).

The Charlevoix area has been affected by several tectonic events of regional significance. In Late Precambrian to Early Cambrian times, the opening of the Iapetus Ocean is recorded in Charlevoix by isolated and discontinuous segments of the St. Lawrence Lowlands platform.

Following the formation of these rift-related faults, the Iapetus ocean is inferred to have covered the Grenvillian basement, which led to the deposition of siliciclastic and calcareous sediments that form the lower part of the St. Lawrence Lowlands in the area (Lemieux, 2000).

Historically, only a few eastern Canadian earthquakes of magnitude greater than 5.5 had observed geological effects such as surface faulting, liquefaction, submarine slumping, rock avalanches, rock falls, landslides, railroad embankment slides, and one tsunami (Lamontagne, 2002). The extent of the Charlevoix Seismic Zone has been defined using instrumentally recorded data (Anglin, 1981). The Charlevoix is the most active region, with five  $M > 6$  events, in 1663, 1791, 1860, 1870, and 1925. Work by Doig (Doig, 1990) has confirmed in part the dates of these earthquakes, based on the relative spacing of the layers and the accelerator-mass spectrometer  $^{14}\text{C}$  dating of a twig in one of the cores taken from Lake Tadoussac, in Charlevoix (Doig, 1990).

According to available geochronological data, a major meteoritic impact occurred in the Late Devonian, and is recorded by abundant shattercone localities in basement and Paleozoic rocks and by circular topographic features that characterize the Charlevoix area.

Finally, the occurrence of a Mesozoic tectonic event has been suggested for the area and consisted of fault reactivation attributed to the opening of the Atlantic Ocean. The geometry and nature of the faults related to impact cratering in the Charlevoix will have to be determined in order to reconstruct and understand the regional tectonic evolution (Lemieux, 2000). Within the boundaries of the Charlevoix impact crater, two types of fault-related breccias exist: coarse-grained cataclastic breccias, and polymictic clastic matrix breccias (Tremblay, 2003).

Most faults of the Charlevoix Fault Zone are brittle faults, usually normal faults characterized by down-dip motion, but strike-slip motion is locally observed; very few faults have been observed with reverse movement (Lemieux, 2000). The St. Lawrence rift system consists of steeply inclined, dip-slip faults characterized by normal-sense movement along longitudinal faults and a variable component of strike-slip faulting along traverse faults (Tremblay, 2003). In La-Malbaie area most faults display normal-sense motion, but reverse faults have also been observed locally. A predominant northwest-trending fault system accounts for more than 30% of fault data (Lemieux, 2000).

The Charlevoix area is a region of significant geologic interest, particularly since the discovery of the impact structure. Previous mapping and analysis have revealed a major part of the regional tectonic history of the area but relationships between the inferred impact cratering with structures left by the Iapetus rifting or the Mesozoic reactivation have still to be determined (Lemieux, 2000).

High resolution marine seismic reflection data acquired in the St. Lawrence River estuary (Tremblay, 2003), identify submarine topographic relief attributed to the St. Lawrence rift system. Northwest-trending seismic profiles suggest that normal faults fringing the St. Lawrence river are associated with a major topographic depression in the estuary, the Laurentian Channel trough, with up to 2,997 ft (700 m) of basement relief (Lemieux, 2000). The analysis of some 124 mi (200 km) of deep seismic reflection lines in the CSZ did not reveal any clear evidence of a surface rupture under the St. Lawrence river including in areas where most  $M > 4.0$  earthquakes occurred in the 20th century (Lamontagne, 2002).

Since 1924, seismicity at Charlevoix has been almost entirely limited to a 44 mi (70 km) zone along the St. Lawrence River (Doig, 1990).

Calculations performed for the PSHA for BBNPP determined that the Charlevoix-La Malbaie is a contributor to the seismic hazard at the BBNPP Site in the low-frequency range.

#### **2.5.1.1.4.4.1.4 Saint Lawrence Rift Valley**

The Saint Lawrence Valley or St. Lawrence rift system is a seismically active zone parallel to the Saint Lawrence River which extends approximately 620 mi (1000 km) from Quebec to Newfoundland, at a minimum distance of 450 mi (724 km) NNE of the BBNPP Site (Feature 36 in Figure 2.5-15). The Charlevoix Seismic Zone, one of the most active seismic regions in eastern Canada, is situated in the Saint Lawrence Valley. The Saint Lawrence Rift System, incorporating the Charlevoix Seismic Zone (described in section 2.5.1.1.4.4.1.3) has the potential for producing moderate to large earthquakes. The March 1, 1925 earthquake with an estimated magnitude between 6.0 to 6.5 Mb is the largest instrumentally recorded earthquake of the Charlevoix Seismic Zone. Several other large earthquakes, however, have been reported in the region since the first settlements (Bent, 1992). Hypocentral data with additional time has shown no migration of seismicity (Anglin, 1984), which is indicative of confinement of high activity to the same area.

Recent research has been performed in the southeastern corner of Lake Ontario and western Lake Ontario in search for evidence of neotectonic faulting (Wallach, 2002). In the Rochester Basin, vertical separations of layers of unconsolidated sediments and the underlying Paleozoic bedrock had been recognized. It may be possible to interpret that the observed displacements of the units are due to the recent tectonic faulting (Wallach 2002). Such interpretation, however, remains uncertain. Finding Quaternary faulting in Eastern North America (ENA) will continue to be difficult. More rigorous study and work is needed in the study area to validate these interpretations.

The EPRI study (EPRI, 1986) excludes the Charlevoix Seismic Zone from the Saint Lawrence Rift System and considers it as a separate seismic zone and that approach was taken for this BBNPP study. The Saint Lawrence Rift System is considered to be a capable tectonic source. Calculations performed for the PSHA for BBNPP determined that the Saint Lawrence Rift Valley (excluding the CSZ) is not a significant contributor to the seismic hazard at the BBNPP Site.

#### **2.5.1.1.4.4.2 Quaternary Tectonic Features**

The U.S. Geological Survey maintains a nationwide database on features that have known or suggested Quaternary tectonic faulting. Geologic information on the Quaternary faults, folds, and earthquake-induced liquefaction in the eastern United States was compiled by Crone and Wheeler (Crone, 2000). An update containing new assessments was published by Wheeler (Wheeler, 2005). Tectonic features, described by Crone and Wheeler (Crone, 2000), and Wheeler (Wheeler, 2005) (Wheeler, 2006) within a 200 mi (322 km) radius (site region) of the BBNPP site were identified, are shown on Figures 2.5-15d and 2.5-15e.

The features within the site region and one outside the region that has demonstrated Quaternary activity, are categorized into one of four Classes, as follows (Crone, 2000). A summary of BBNPP features in each Class is provided as well:

Category	Number of Features	Description
Class A	1	Geologic evidence demonstrates the existence of a Quaternary fault of tectonic origin, whether the fault is exposed by mapping or inferred from liquefaction or other deformational features.
Class B	0	Geologic evidence demonstrates the existence of Quaternary deformation, but either (1) the fault might not extend deeply enough to be a potential source of significant earthquakes, or (2) the currently available geologic evidence is too strong to confidently assign the feature to Class C but not strong enough to assign it to Class A.
Class C	16	Geologic evidence is insufficient to demonstrate (1) the existence of tectonic faulting, or (2) Quaternary slip or deformation associated with the feature.
Class D	0	Geologic evidence demonstrates that the feature is not a tectonic fault or feature. This category includes features such as joints, landslides, erosional or fluvial scarps, or other landforms resembling fault scarps but of demonstrable nontectonic origin.

Note that only one (1) feature was categorized as Class A – demonstrating convincing evidence of Quaternary activity. This is the Newbury Liquefaction Feature. This feature is outside the BBNPP site region (200 mi (322 km) radius). In addition, no features in the site region were identified as having sufficient evidence of Quaternary deformation (Class B).

The sixteen (16) Class C features in the site region are listed in this section and are shown in Figure 2.5-15 and Table 2.5-62. Information on these features was reviewed for this study, and these features are not considered in the seismic hazard assessment as characteristic tectonic sources for the BBNPP site.

The known or suggested Quaternary tectonic features that may be significant for the assessment of the seismic hazard at the BBNPP site are described briefly in the following subsections. More detailed discussion of the seismicity and evidence regarding the magnitude and recurrence of earthquakes within the seismic zones is provided in this subsection, and Subsection 2.5.2.

Following is a list of the seventeen (17) Quaternary features described below.

1. Newbury Liquefaction Feature (Class A)
2. Ramapo Fault System (Class C)
3. Furlong-Flemington Fault System (Class C)
4. Kingston Fault (Class C)
5. New York Bight Fault (Class C)
6. Mosholu Fault (Class C)
7. New Castle County Faults (Class C)
8. Upper Marlboro Faults (Class C)
9. Dobbs Ferry Fault Zone (Class C)
10. Lancaster Seismic Zone (Class C)
11. Cacoosing Valley Earthquake Sequence (Class C)
12. Moodus Seismic Zone (Class C)
13. Clarendon-Linden Fault Zone (Class C)
14. Offset Glaciated Surfaces (Class C)
15. Fall Lines of Weems (Class C)
16. Everona Fault-Mountain Run Fault Zone (Class C)
17. Stafford Fault System (Class C)

#### **2.5.1.1.4.4.2.1 Newbury Liquefaction Features (Class A)**

The Newbury area lies in northeastern Massachusetts 300 miles (480 km) to the northeast of the site, and has been the subject of recent research. Crone and Wheeler (Crone, 2000), in their compilation of faults and tectonic features of central and eastern United States selected the name "Newbury Liquefaction Features" for the part of the structure located near Newbury (Essex County, MA) and assigned it to Class "A". Their reason for assignment to Class "A" was eyewitness reports of liquefaction during an earthquake in 1727 (MMI VII, Magnitude = 4.8) and sand dikes that were attributed to the 1727 earthquake. Tuttle and Seeber (Tuttle, 1991) concluded that the liquefaction was caused by strong ground motion but the causative fault responsible for the ground motion and liquefaction remains unidentified.

The Clinton-Newbury Fault Zones are northeast trending faults that extend from Connecticut to New Brunswick Canada. The fault zones are approximately 262 mi (422 km) northeast of the BBNPP site (Feature 21 in Figure 2.5-15). Strike slip movements have been documented along the fault zone. The trend along this fault system has been the location of several moderate to

large earthquakes (EPRI, 1986). Based on research, the fault zone has been assigned an activity probability of 0.2, since location uncertainties prevented association of the apparent trend of earthquakes with the trend of the tectonic feature. The stress information in the area of this tectonic feature was inconclusive (EPRI, 1986). Information on this feature was reviewed for this study, and this feature is considered to be a capable source, but not a characteristic tectonic source for the BBNPP site.

#### **2.5.1.1.4.4.2.2 Ramapo Fault System (Class C)**

The Ramapo Fault System is located in northern New Jersey and southern New York State, approximately 124 mi (200 km) north-northeast from the BBNPP site (Figure 2.5-15). This fault system consists of northeast-striking, southeast-dipping, normal faults and classified as border faults for the Mesozoic Newark Basin (Jacob, 2004) (Feature 15 in Figure 2.5-15). The Ramapo Fault System has a northeast strike and dips (approximately 70°) to the southeast. The term "Ramapo Seismic Zone" (RSZ) is used for the seismically active 7.5 mi (12 km) wide eastern area of Reading Prong (Sykes, 2008). Since post 1974 earthquakes have been located with higher accuracy, this 7.5 mi (12 km) width cannot be attributed to location errors. Therefore, it is concluded that more than one fault must be involved in generating the earthquakes (Sykes, 2008). Reactivation of the Ramapo fault during the Quaternary period has not been demonstrated. Results of core analyses in six localities of Ramapo and other basin-border faults showed that the most recent slip was extensional at each locality. The extensional tectonic episode did not extend beyond the Mesozoic and there is no evidence of post-Jurassic displacement (Sykes, 2008).

Although earlier tectonic episodes characterize parts of the Fault System evolution, it is best characterized as a normal, Mesozoic basin boundary fault (EPRI, 1986). Evidence of the repeated slip since Proterozoic time (including Mesozoic extensional reactivation) is contained in different faults of the system (Ratcliffe, 1971). Earthquakes have occurred in the general vicinity of the Ramapo Fault System. Many of the earthquakes have not been well located mainly because of poor seismic station distribution prior to 1970's. Therefore, while the association of earthquakes with the Ramapo Fault System is possible, the uncertainty in the locations allows the association with other structures in the area as well. Additional seismographs have been installed in the area of the Ramapo Fault since 1970's and a large amount of micro earthquake data has been recorded. Based on earthquake hypocenters and single event focal mechanisms, Aggarwal and Sykes (Aggarwall, 1978) inferred a reverse slip on a surface dipping 60°-65° southeast from the trace of the Ramapo Fault. The history of repeated slip during the Proterozoic and correlation of fault trend with epicenters, led Aggarwall and Sykes (Aggarwall, 1978) to conclude that the Ramapo Fault System is active. Based on the scattered epicenters, they also concluded that the seismicity was concentrated along a group of northeast-trending faults of which the Ramapo Fault appeared to be the most active. Yang (Yang, 1981) concluded that the Ramapo Fault System is probably the most active fault system in the greater New York City area.

Examination of small earthquakes and re-evaluation of some focal mechanisms in 1980's did not favor the association of the epicenters with the Ramapo Fault. Seborowski (Seborowski, 1982) studied a sequence of micro earthquakes near Annsville, New York, recorded during January, 1980, and derived a composite focal mechanism solution. Their solution indicated east-northeast compression resulting in thrust motion on a north-northwest striking fault plane. This direction is transverse to the northeast trend of the major structures in the epicentral region including the Ramapo Fault (Aggarwall, 1978). The dominant reverse mechanisms (Aggarwall, 1978) imply east-southeast maximum (horizontal) compressive stress. However, the maximum

horizontal compressive stress trend throughout most of the eastern United States including the area of the Ramapo Fault System is east-northeast (Crone, 2000).

An improved 3-D velocity model (Thurber, 1985), which shows 10-15 percent velocity difference across the Ramapo Fault, has changed some of the epicenters and depths of the earthquakes studied by Aggarwal and Sykes (Aggarwall, 1978). Kafka (Kafka, 1985) used earthquake data of the greater New York City area and refined the catalog to eliminate station and detection bias from the network seismicity. This allowed for uniform measurement of magnitudes and earthquake locations. The results showed half of the earthquakes occurred about 6 mi (10 km) from the Ramapo Fault and about half were located about 31 mi (50 km) from this fault -- around the northern part of Newark Basin. Kafka (Kafka, 1985) concludes that "while the Ramapo Fault can by no means be ruled out as a possible source zone for earthquakes in the greater New York City area, the cause of earthquakes in this region is, in the final analysis, still unknown."

In general, even though the epicenters align along the Ramapo Fault, the association is less significant than the one suggested by Aggarwal and Sykes (Aggarwall, 1978). Therefore, the Ramapo Fault system seems to dominate the seismicity (Crone, 2000). Many earthquakes in the area have been attributed to the reactivation of the Ramapo Fault by the present-day compressional stress field. The results of core analyses (in the area of Ramapo Fault in New York and New Jersey) by Ratcliffe and Burton (Ratcliffe, 1984) are not consistent with the reactivation of Ramapo Fault and related faults in the present-day stress field.

There is an apparent discrepancy between the distribution of earthquakes in the Ramapo Seismic zone and the lack of displacement in the last 150 Ma in the localities where the cores have been taken (Sykes, 2008). As a possible explanation for this discrepancy, it is assumed that earthquakes may originate from other preexisting faults which may or may not strike similar to Ramapo fault. It is not clear which faults are active. The seismicity data used and processed cannot be solely used to delineate a single fault or multiple faults. Geologic evidence of Holocene fault movements is also very hard to find in the study area.

It is not possible to conclude that the Ramapo fault is an active feature. Therefore, even though it cannot be ruled out as a possible source for some of the observed earthquakes, the cause of the earthquakes is still unknown. This is a conclusion that has been reached by most previous studies (Kafka, 1985). It is not clear which magnitude will probably represent the maximum magnitude for the area or for the Ramapo fault assuming that it is active. The largest observed earthquake in the area is the 1884 (offshore) New York earthquake with  $m_b L_g = 5.25$ . Based on the observed seismicity a maximum magnitude of 6 or slightly higher is appropriate.

The EPRI seismic source model does not consider the Ramapo fault system as a separate seismic source. For the BBNPP PSHA, the maximum magnitudes of the EPRI seismic source zones (that encompass the Ramapo seismic zone) range from 5.3 to 7.1. These values have been presented in different logic trees by six EPRI teams and adequately characterize the upper level seismicity. Even though a larger source zone (currently used in BBNPP PSHA) tends to diffuse the hazard, this would only have an effect if the RSZ was located closer to the site. Based on the available data and information, the current PSHA results adequately reflect the hazard at the BBNPP from the Ramapo fault zone.

Some branches of the Ramapo fault system (not the Ramapo fault itself) may extend into southeastern Pennsylvania. Those branches have similar distances from the site. Even if it is assumed the Ramapo fault as an active source (with  $M_{max} \sim 6.0$ ), it cannot produce significant ground motions at the site because of the relatively large distance away



Crone (Crone, 2000) also summarized a few reports that indicate some forms of Quaternary deformation near the Ramapo Fault but argue that none of the reports provide convincing evidence for Quaternary faulting or sudden offset which can be used to distinguish prehistoric seismic slip from a seismic creep. The Ramapo Fault or Fault System is probably capable of generating small or rare large earthquakes but it has been assigned to Class C feature since evidence for quaternary faulting has not been identified (Crone, 2000). Information on this feature was reviewed for this study, and this feature is considered to be a capable source, but not a characteristic tectonic source for the BBNPP site.

#### **2.5.1.1.4.4.2.3 Furlong-Flemington Fault System (Class C)**

Furlong Fault is located in the west portion of the Newark Basin, near New Hope in eastern Pennsylvania. This major intrabasinal Mesozoic fault connects to the north, with the Flemington fault in New Jersey (Ratcliffe, 1988). It is located about 75 mi (121 km) southeast of the BBNPP (Feature 24 in Figure 2.5-15). The Furlong and Flemington faults have been considered as one fault system (Root, 1999). Ratcliffe and Burton (Ratcliffe, 1988), using data from coring and surface observations, determined that the Furlong-Flemington fault zone consists of two closely parallel faults that dip at 47° to 50° to the southeast. Structural analysis also indicated a normal fault with some component of strike slip for the Furlong fault. Many of the Mesozoic border faults, such as the Ramapo and Flemington faults, coincide with thrust Paleozoic faults. It seems that the reactivation of the Paleozoic thrusts by Mesozoic border faults controlled the overall structure of the basin (Ratcliffe, 1985). However, there is no indication or evidence of later activities for the fault system. Seismicity has not been associated with either Flemington or Furlong faults or with larger Chalfont Fault, which is intersected by the Furlong Fault. Based on the lack of evidence of recent or Quaternary activity, Furlong-Flemington Fault System is not a considered to be a capable tectonic source.

#### **2.5.1.1.4.4.2.4 Kingston Fault (Class C)**

The Kingston Fault is located in central New Jersey, approximately 87 mi (140 km) east-southeast of the BBNPP (Feature 6 in Figure 2.5-15). The Kingston Fault is 8 mi (13 km) long, north to northeast strike with a nearly vertical dip (Owens, 1995). The fault is located in the Mesozoic formation of the Newark Basin (Crone, 2000). Parker (Parker, 1990) showed the northern part of the fault trace on the map and reported a dip of 85° SE with extensional slip during the formation of the Basin in Mesozoic (Stanford, 1995). Results of well, boring, and geophysical data showed movement of southeast side of the fault based on the thickened Pliocene gravel across the fault. The Pliocene gravel that thickens across the fault is overlain by late Pleistocene gravel, which is not offset by the fault, indicating the fault probably moved during Pliocene or early to middle Pleistocene (Stanford, 1995). Quaternary activity for the fault can not be demonstrated and the fault slip rate is unknown. Additionally, no paleoseismological study has been performed on the thickness of the Pliocene gravel to determine seismic creep, or different episodes of seismic faulting (Crone, 2000). According to Wheeler (Wheeler, 2005), Kingston Fault was assigned to a Class C feature. No seismicity has been associated with the fault. Based on the information above, the Kingston fault is not considered to be a capable tectonic source.

#### **2.5.1.1.4.4.2.5 New York Bight Fault (Class C)**

On the basis of seismic surveys, the New York Bight Fault is characterized as an approximately 31 mile (50 km) long, north-northeast-striking fault, located offshore of Long Island, New York (Schwab, 1997) (Hutchinson, 1985) (Feature 5 in Figure 2.5-15). It is 115 mi (185 km) east of the BBNPP Site, and parallel to the New Jersey coast (Hutchinson, 1985). Seismic reflection

data of the fault showed a length of at least 19 mi (30 km) and extended southward. Based on the results, the fault had offset Upper Cretaceous rocks and lower Tertiary and Quaternary deposits, therefore it may be as young as the Quaternary (Hutchinson, 1985).

The fault was mapped (Crone, 2000) along 24 mi (39 km) of its northern extension and dips almost vertically and was traced to within 6 mi (10 km) of the Long Island Coast (Lotto, 1997). Cretaceous to Eocene strata have been offset by the fault, but an unconformity which separates the Eocene and Miocene strata (and Miocene strata overlying it) are not offset sufficiently within the resolution of the available seismic profiles (Hutchinson, 1985). Ongoing seismic reflection work indicated that middle to late Quaternary sediments overlay Cretaceous and Tertiary strata at the fault (Lotto, 1997). These Quaternary sediments are not offset more than 3 ft (0.9 m) which is the resolution of the measurement (Crone, 2000).

The seismicity near the fault show small magnitudes (less than 3.00) have been located within 13 mi (21 km) from the fault. The location error for the offshore earthquakes exceeds 6 mi (10 km) (Yang, 1981) because of the seismic station distribution; therefore the location of these events is not reliable. Crone and Wheeler (Crone, 2000) and Wheeler (Wheeler, 2006) classify the New York Bight Fault Zone as a Class C feature, based on the lack of Quaternary activity evidence. Information on this feature was reviewed for this study, and this feature is considered to be a capable source, but not a characteristic tectonic source for the BBNPP site.

#### **2.5.1.1.4.4.2.6 Mosholu Fault (Class C)**

The Mosholu Fault is located in Bronx County, New York City, New York (Wheeler, 2006). The fault is approximately 135 mi (217 km) east of the BBNPP (Feature 4 in Figure 2.5-15). The Mosholu Fault is 5.6 mi (9 km) long northwest trending right-lateral oblique-slip fault with a steep dip that crosses the Bronx River channel. Merguerian (Merguerian, 1996) suggested that the fault showed postglacial age uplift forming the buried ridge. The ridge caused the creation of a lake where clay settled in the Bronx River and overlay the glacial deposits detected north of the bedrock barrier. The fault has also been mapped and renamed by Baskerville (in Crone, 2000). The fault has been located by geological mapping and subsurface data (Crone, 2000). Localized surface deformation of post glacial times may have occurred in the area where the Bronx River crosses the Mosholu Fault (Merguerian, 1997). The blockage of the Webster Avenue lowland, which has caused the diversion of the Bronx River, resulted from neotectonic uplift of a block of bedrock along the NE side of Mosholu fault. Crone (Crone, 2000) argues that while attributing the uplift to post glacial slip on the Mosholu fault, it is not demonstrated. Merguerian (Merguerian, 1997) could not prove that the uplift occurred seismically. Additionally Merguerian (Merguerian, 1997) mentioned that none of the previous New York City's magnitude ~ 5.0 earthquakes of 1737, 1783, and 1884 have been connected with surface blockage of crustal rocks. Merguerian (Merguerian, 1997) was not able to associate historic earthquakes with the faults in New York City area. However, based on the circumstantial evidence from the Bronx River, Merguerian suggests that NW-trending faults in New York City area, such as the Mosholu Fault and the Dobbs Ferry Fault, are seismically capable. Earthquakes have not been associated with the Mosholu Fault and the fault has not been studied in detail for paleoseismological evidence of possible Quaternary activity. The Mosholu Fault has been assigned to Class C by Crone (Crone, 2000). Information on this feature was reviewed for this study, and this feature is considered to be a capable source, but not a characteristic tectonic source for the BBNPP site.

#### **2.5.1.1.4.4.2.7 New Castle County Faults (Class C)**

The New Castle County faults are characterized as 146 mi (235 km) long buried north and northeast-striking faults that displace an unconformable contact between Precambrian to Paleozoic bedrock and overlying Cretaceous deposits. The faults are located in northern Delaware, near New Castle, about 150 mi (242 km) southeast of the BBNPP site (Feature 9 in Figure 2.5-15). Based on research (Spoljaric, 1973), a graben is present in New Castle County with a northeastern strike near Delaware City. The graben is bounded by faults that are part of the basement fault that underlies the Coastal Plain of Northern Delaware. The bounded faults have shown displacements ranging from 32 to 98 ft (10 to 30 m) across the basement-Cretaceous boundary (Spoljaric, 1972). Along this fault zone, earthquakes have occurred and have showed magnitudes as high as 3.8. According to studies completed by the Delaware Geological Survey (DGS) (McLaughlin, 2002), a subsurface investigation utilizing seismic reflection and seismic refraction, subsurface drilling, geophysical logging and trench excavation was performed to potentially locate displacement from faults near New Castle, Delaware. No shallow faults were detected during the subsurface drilling program and trench excavation. Seismic section identified extensive faulting in the investigation site where the New Castle fault is projected. DGS concluded that minimal, if any, modern fault activities occurred in the area of New Castle County. Wheeler (Wheeler, 2005) characterizes the New Castle County faults as a Class C features. Based on the above information, including the recent strong evidence as presented by McLaughlin (McLaughlin, 2002), the New Castle County Faults is considered to be a capable source, but not a characteristic tectonic source for the BBNPP site..

#### **2.5.1.1.4.4.2.8 Upper Marlboro Faults (Class C)**

The Upper Marlboro Faults are located in Prince Georges County, Maryland approximately 150 mi (241 km) southeast of the BBNPP site (Feature 25 in Figure 2.5-15). The faults are a series of features, which cut the Coastal Plain sediments (Crone, 2000). The faults have a low angle dip, which is more consistent with a surficial origin, and extend to hypocentral depths (Crone, 2000). Wheeler (Wheeler, 2006) related the faults to surficial land slides based on low angle dips. The faults are assigned to Class C feature by Crone and Wheeler (Crone, 2000) and Wheeler (Wheeler, 2006) because no evidence of Quaternary activity has been presented for the faults. Seismicity has not been associated with the fault. Accordingly, the Upper Marlboro Faults are not considered to be a capable tectonic source.

#### **2.5.1.1.4.4.2.9 Dobbs Ferry Fault (Class C)**

The Dobbs Ferry Fault is located in Westchester County, New York about 6 mi (10 km) north-northeast of New York City (Wheeler, 2006). The fault is approximately 155 mi (249 km) east of the BBNPP (Feature 3 in Figure 2.5-15). The Dobbs Ferry fault zone is a zone of abundant fractures and joints that extends southeastward from the east bank of the Hudson River and crosses the Bronx River. The fault had dextral slip during the Mesozoic as part of the Pangaea separation (Crone, 2000). Different orientations of superimposed slickenside show more than one episode of slip on the fault. Sinistral slip can be inferred from the majority of slip sense indicators, which is consistent with the present day, east-northeast, regional orientation of the maximum horizontal compressive stress. Some indicators are dextral and older, and perhaps date from Mesozoic extension (Seeber, 1998). The October 19, 1985, Ardsley earthquake, with a magnitude of 4.1 along with six aftershocks ranging from 4 to 4.5 in magnitude, occurred in southern Westchester County, New York, approximately 19 mi (30 km) north of central Manhattan, approximately 125 mi (201 km) east of the BBNPP Site. The location of the first six aftershocks (within a week of the main shock) defined a vertical northwest trending rupture zone with an approximate diameter of 2300 ft (700 m) and a depth ranging from 2.8 to 3.4 mi (4.5 to

5.5 km). The rupture zone corresponds directly to a 0.6 mi (1 km) segment of the Dobbs Ferry fault (Hough, 1991). First motion data of the main shock and this group of aftershocks yielded well constrained focal mechanism solutions indicating sinistral slip on a northwest striking plane (Hough, 1991) (Crone, 2000). Later aftershocks defined a northeast striking plane. These results led Seeber (Seeber, 1998) to conclude that the earthquakes probably occurred on the fault zone without surface rupture along its trace. Crone (Crone, 2000) assigns the Dobbs Ferry fault zone to a Class C feature because no paleoseismological evidence for Quaternary seismic activity has been reported for the fault. The Dobbs Ferry Fault is considered to be a capable source, but not a characteristic tectonic source for the BBNPP site.

#### **2.5.1.1.4.4.2.10 Lancaster Seismic Zone (Class C)**

According to Armbruster (Armbruster, 1987), the Lancaster Seismic Zone (LSZ) is located in southeast Pennsylvania. The seismic zone is of circular shape with a diameter of about 31 mi to 38 mi (50 km to 61 km) and is 55 mi (88 km) south of the BBNPP site (Feature 8 in Figure 2.5-15). The Lancaster Seismic Zone consists of short discontinuous north striking faults. The faults transect Triassic and Paleozoic rocks. The age of the faulting occurred during the early Mesozoic extension relating to the Atlantic margin. It also crosses the Newark-Gettysburg Triassic rift basin which consists of extensional faults associated with Mesozoic rifting. Most well-located epicenters in the Lancaster seismic zone lie directly outside the Gettysburg-Newark Basin (Scharnberger, 2006).

In the western part of the Lancaster Seismic Zone, the epicenters of 11 events from 1889 through 1994 (with magnitudes 3.04 to 4.61) define a north-south trend that intersects the juncture between the Gettysburg and Newark sub-basins (Armbruster, 1987). The highest magnitude earthquakes were at Marticville in 1984 at a magnitude of 3.7 and Cacoosing Valley in 1994 at a magnitude of 4.6. The earthquakes occurred 31 mi (50 km) apart and were related to tectonic fault lines.

The 1984 Marticville earthquake was the second-largest recorded event of the zone. Armbruster and Seeber (Armbruster, 1987) suggested a seismogenic shallow fault (centered at 2.5 mi (4.0 km) depth) based on the hypocentral distribution and first motion data from several recorded earthquakes of the 1984 sequence. While not in correlation with the trend of Paleozoic structures in the epicenter area, the 1984 rupture geometry conforms to the strike of the Jurassic dikes and their associated faults. Earthquakes in the zone may have been related with the Rockhill dike in particular (which bisects LSZ and is close to the 1984 rupture area), and its related faults (Armbruster, 1987).

The January 16, 1994 Cacoosing Valley earthquake which struck the northeastern edge of the zone, is known to be the largest earthquake of the LSZ (Seeber, 1998). The Cacoosing Valley earthquake was of magnitude 4.6 and Modified Mercalli Intensity (MMI) VI-VII. The associated earthquake sequence, from 1994 to 1997, is discussed in more detail in section 2.5.1.1.4.4.2.11. As is the case of 1984 Marticville earthquake, the 1994 Cacoosing Valley earthquake, and its aftershocks, provided seismological evidence for an active fault in the LSZ. The suggested faults, for both the 1984 and 1994 earthquakes, were based merely on seismological evidence. Geological evidence, such as surface rupture or liquefaction, has not been found in either case (Crone, 2000).

Some of the previous works summarized by Crone and Wheeler (Crone, 2000) indicate a seismically active fault, or fault zone, in the LSZ. Many studies provide evidence for high density north-striking, near surface structures and fracture zone in the LSZ. Some of these features of this seismic zone were explained in relation to the Newark Gettysburg Basin. The seismic zone,

as defined by Armbruster and Seeber (Armbruster, 1987), is located about 55 mi (88 km) south of the site and has been a source of seismicity for more than 2 centuries. The zone approximately coincides with Lancaster County, Pennsylvania. It marks the southwestern edge of the Newark Basin Seismic Zone (NBSZ). Thrust faults and folds (formed during the Paleozoic Appalachian orogeny) and extensional faults (associated with Mesozoic rifting) are among the main structures along the NBSZ and LSZ. Regional Jurassic dikes, striking north-northeast, traverse the LSZ. Coinciding with brittle faults, many of these dikes are perhaps among the youngest structures in the region which persist as large planar zones of weakness and cut through the crust (Seeber, 1998).

No other evidence of Quaternary faulting (e.g., paleoliquefaction) in the LSZ has been reported and it has been assigned to Class C (Crone, 2000). Crone and Wheeler (Crone, 2000) note that even short and shallow faults can host earthquakes as demonstrated by the unusually shallow seismicity of the LSZ. There might be other deeper geologic controls on seismicity which are reflected by the shallow faults of the LSZ (Wise, 1998). None of the EPRI study (EPRI, 1986) groups considered the LSZ as a separate seismic source zone. However, the maximum magnitudes (ranges from 5.2 to 7.2) that have been assigned by the groups to the regions in which the LSZ is situated are larger than any reported earthquakes in the LSZ. Therefore, the EPRI study (EPRI, 1986) adequately characterizes the LSZ in terms of the upper bound magnitude.

Accordingly, the Lancaster Seismic Zone is considered to be a capable source, but not a characteristic tectonic source for the BBNPP site.

#### **2.5.1.1.4.4.2.11 Cacoosing Valley Earthquake Sequence (Class C)**

The 1993 to 1997 Cacoosing Valley earthquake sequence occurred along the eastern margin of the Lancaster Seismic Zone with the main shock occurring on January 16, 1994, near Reading, Pennsylvania about 52 mi (84 km) south of the BBNPP site (Crone, 2000) (Feature 7 in Figure 2.5-15). The maximum magnitude earthquake associated with this sequence is an event of magnitude 4.6 (Crone, 2000). Forty one (41) aftershock hypocenters occurred around the rim within 1 mi (1.6 km) diameter and a depth ranging from 0 to 2 mi (0 to 3.2 km) and orientation of N43° W and 54° SW. The main shock occurred at a depth of 1 mi (1.6 km) and aftershocks occurred from that depth to the surface (Seeber, 1998).

The main shock occurred under an abandoned quarry. In December 1992, the quarry was allowed to flood with water rapidly. The unloading during the quarry process and increased pore pressure caused by subsequent flooding created the release of energy. The zone of the rupture, obtained from the aftershock locations, matched a nodal plane with reverse and left-lateral slip (strike 135°, dip 54° SW) of a focal mechanism obtained from aftershock first motions and main shock waveforms. However, the rupture did not correlate with any of the mapped faults in the area (Seeber, 1998), and the earthquake did not rupture the ground surface.

Wheeler (Wheeler, 2006) defines the seismic event as a feature where there is not sufficient evidence to demonstrate that other faulting occurred in the Quaternary and assigns the Cacoosing Valley earthquake sequence as a Class C feature. Additional information relative to the earthquake is in section 2.5.1.1.4.4.3.1.

This earthquake sequence is not considered to be related to a capable tectonic source.

#### **2.5.1.1.4.4.2.12 Moodus Seismic Zone (Class C)**

The Moodus Seismic Zone is located in Middlesex and New London County, Connecticut (Wheeler, 2005). The seismic zone is approximately 190 mi (306 km) northeast of the BBNPP (Feature 1 in Figure 2.5-15). The town of Moodus is located about 20 mi (32 km) southeast of Hartford, Connecticut. The area around the towns of Moodus and Haddam has been the most seismically active region in Connecticut, and the earliest records of earthquake activity in the area dates back to the second half of the 16th century (Ebel, 1982). The largest historical earthquake occurred in 1791 and had an estimated MMI in the range of VI-VIII (Ebel, 1982) (Crone, 2000). Four shallow microearthquakes occurred in the Moodus Seismic Zone in 1980's (Ebel, 1982) (Ebel, 1989). The seismic sources of the earthquakes were reviewed utilizing geological and geophysical methods (Koch, 1978) (Crone, 2000), but no causative fault has been identified (Crone, 2000). No evidence of liquefaction or paleoliquefaction has been found in the reconnaissance efforts in the Moodus area (Gelinas, 1993). Accordingly Crone and Wheeler (Crone, 2000) assign the seismic zone as a Class C feature. There is no new information about the seismic zone in Wheeler's (Wheeler, 2005) compilation of possible Quaternary features of the CEUS. Information on the Moodus seismic zone was reviewed for this study, and this feature is considered to be a capable source, but not a characteristic tectonic source for the BBNPP site.

#### **2.5.1.1.4.4.2.13 Clarendon-Linden Fault Zone (Class C)**

A major north-south trending system of thrust faults forms the Clarendon-Linden fault zone. The fault zone is approximately 150 mi (241 km) from the BBNPP site (Feature 20 in Figure 2.5-15) where it extends over 62 mi (100 km) from western New York State to northern Allegany County, New York (Fletcher, 1977). According to recent works it further extends into central Allegany County, near the Pennsylvania border (Crone, 2000). The presence of a north-south striking fault was inferred from the surface geology of the area (Herrmann, 1978). Since then, the fault zone has been mapped geologically at the surface and characterized by geophysical methods such as seismic reflection at subsurface (Crone, 2000).

The August 12, 1929 Attica earthquake, with an estimated magnitude of 5.2 (Street, 1977) and epicentral density of VIII, was located near the Clarendon-Linden Fault Zone. Herrmann (Herrmann, 1978) studied two other events (both with epicentral intensities VI) that occurred on January 1, 1966 and June 12, 1967 in the Attica region, and found shallow depth 1 mi to 2 mi (1.6 km to 3.2 km) for both events. The historical seismicity also shows a diffuse east-west trend that does not correlate with the north-south trend of the structure around Attica. Fault plane solutions for 1966 and 1967 events showed similar nodal planes for both events striking about 120° and 20°. The fault plane on the NNE nodal plane is parallel to the Clarendon-Linden structure (Herrmann, 1978).

There is no paleoseismological evidence of Quaternary slip on the fault zone so the zone has been assigned to Class C feature by Crone and Wheeler (Crone, 2000). Tuttle et al. (Tuttle, 2002) concluded that the Clarendon-Linden Fault Zone has not generated large events (moment magnitude > 6) during the past 12,000 years. Based on the lack of earthquake-induced liquefaction features along the fault zone, including the area of the 1929 Attica earthquake, the fault zone is probably the source of the 1929 shock. Tuttle (Tuttle, 2002) concluded that the Clarendon-Linden Fault Zone is capable of producing future M 5 events, and it is considered a seismic source zone by the Bechtel and Dames & Moore EPRI (1986) Groups. Information on this feature was reviewed for this study, and this feature is considered to be a capable source, but not a characteristic tectonic source for the BBNPP site.

#### **2.5.1.1.4.4.2.14 Offset Glaciated Surfaces (Class C)**

Small steeply dipping faults offset glacially smoothed rock surfaces at different locations in the northeastern U.S. including Maine, Massachusetts, New Hampshire, New York, and Vermont (Crone, 2000). The closest offset glaciated surfaces are located 75 mi (121 km) northeast of the BBNPP site (Feature 23 in Figure 2.5-15). The offsets are typically small in the range of millimeter to decimeter, with some exceptions of larger displacements up to one (1) meter. Crone and Wheeler (Crone, 2000) summarizes a few studies of localities in which such surfaces are located, but indicates that no systematic and comprehensive field or literature search has been done for such localities in the Northeast. Although some studies favor a tectonic origin over other frosting processes (Oliver, 1970), frost heaving was the likely origin for the offset glaciated surfaces (Crone, 2000). In Hudson River Valley of eastern New York, Quebec, and New Brunswick, the small faults show offsets that are uniform in size over distance and are parallel to the strike of cleavage in the heaved irregular size blocks. According to Crone (Crone, 2000), tectonic origin of the small faults in the present-day stress field is unlikely because these small faults have been found in all directions (Crone, 2000). Ratcliffe (Ratcliffe, 1982) studied Paleozoic slates, which had been summarized as also favors frost heaving over recent tectonic activity as the origin of the examined small faults.

Crone and Wheeler (Crone, 2000) conclude that evidence supports frost-wedging more than any other process as the likely origin of small faults. Even by assuming a tectonic origin, these small faults do not penetrate deep downward and therefore do not have a significant effect on seismic hazard. Crone and Wheeler (Crone, 2000) mention that offset glaciated surfaces are Quaternary in age but classify them as Class C features based on the conclusion that the small faults with limited length and depth extent do not affect seismic hazard significantly. These small faults have been observed over a large area in 5 states and there has been no seismicity associated with them. Based on the available information and evidence, these small faults are not considered to be capable tectonic sources.

#### **2.5.1.1.4.4.2.15 Fall Lines of Weems (Class C)**

Weems (Weems, 1998) identified numerous short stream segments or fall zones with steep gradients while examining longitudinal profiles of major rivers that flow southeastward or northwestward across the Piedmont and Blue Ridge provinces of North Carolina and Virginia. The northeastern tip of the Blue Ridge fall line is located about 140 mi (225 km) southwest of the BBNPP site (Feature 26 in Figure 2.5-15). He noticed the alignment of fall zones of different streams and used the term "fall lines" for the curvilinear trend of the alignments. He defined seven fall lines that trend northeastward, paralleling the regional tectonic fabric and gravitational gradient of the Appalachian Orogeny. The fall lines tend to merge northeastward. Weems (Weems, 1998) states that "limited available evidence favors a neotectonic origin" for the fall lines and rules out climate control. Wheeler (Wheeler, 2005) argues that the identification of fall zones is subjective; therefore Weems' (Weems, 1998) arguments and conclusions depend on the choice of the fall zones. Wheeler (Wheeler, 2005) concludes that tectonic faulting is not yet demonstrated for the fall lines and assigns the fall lines to Class "C." The fall lines of Weems have not been associated with seismicity and they are not considered to be capable tectonic sources.

#### **2.5.1.1.4.4.2.16 Everona Fault-Mountain Run Fault Zone (Class C)**

The Mountain Run Fault Zone is a regional geologic and tectonic feature of central Virginia, which extends from the eastern margin of the Mesozoic Culpeper Basin near the Rappahannock River southwestward, to near Charlottesville, VA (Pavrides, 1986) (Pavrides,

1994). The fault zone that is located about 180 mi (290 km) southwest of the BBNPP site, trends northeast and extends for 63 mi to 94 mi (100 km to 150 km) (Wheeler, 2006). See Feature 16 on Figure 2.5-15. The Fault zone forms part of the southeast boundary of Early Mesozoic Culpeper Basin. The Everona Fault occurs in close proximity to the Mountain Run Fault Zone (about 0.6 mi (1.0 km)) and has an estimated age of late Cenozoic (Crone, 2000) (Bobyarchick, 2007). The fault zone is a reverse fault that dips about 20° NW and truncates layers of rocks in the footwall of the Mountain Run fault zone (Bobyarchick, 2007). This small fault displaces the base of late Tertiary or Pleistocene gravel located about 0.6 mi (1.0 km) west of the Mountain Run Fault Zone (Wheeler, 2006).

Thrust faulting along the Mountain Run Fault Zone started at the end of Ordovician. Subsequent strike-slip movement in the fault zone occurred prior to middle Mesozoic, since undeformed basaltic dikes of Jurassic age cut the Mountain Run Fault Zone rocks (Pavrides, 1994). To the northeast, two scarps occur along the fault zone: (1) Mountain Run scarp, located on the southeast side of the Mountain Run extends for 8 mi (13 km), and (2) Kellys Ford scarp, located on the northeast part of Mountain Run Fault Zone, bounds the Culpeper Basin. Kelly Ford scarp is 1 mi (1.6 km) long and is related to the southeastern border fault of Mesozoic Culpeper Basin. Pavrides (Pavrides, 1994) argued that rugged topography of both scarps implies Cenozoic or possibly Pleistocene age. Based on the displaced saprolites in the area, the Everona Fault is a structure of tectonic origin involving the basement, and is not confined to the overlying surficial deposits. (Bobyarchick, 2007).

Crone (Crone, 2000) and Wheeler (Wheeler, 2006) conclude that while the faulting at Everona is likely to be of Quaternary age, no Quaternary activity has been demonstrated for the feature. They classify the feature as Class C, but mention that this feature did not have a detailed paleoseismological study to determine whether it has been a site of Quaternary earthquake. Additional investigation has been done for the North Anna Early Site Permit which shows that Mountain Run Fault Zone has not been active during the Quaternary (Bobyarchick, 2007). Information on this feature was reviewed for this study, and this feature is considered to be a capable source, but not a characteristic tectonic source for the BBNPP site.

#### **2.5.1.1.4.4.2.17 Stafford Fault System (Class C)**

The Stafford fault system approaches within 180 mi (290 km) south of the site (Feature 18 in Figure 2.5-15d). The 42 mi (68 km) long fault system strikes approximately N35°E for a distance of 45 mi (72 km) along the west bank of the Potomac River in northeastern Virginia (Mixon, 1977). The Stafford Fault System consists of five northwest dipping, high-angle reverse faults and follows the inner margin of the Coastal Plain province. Four faults, the Dumfries Fault Zone, Fall Hill Fault, Hazel Run Fault, and unnamed fault, strike northeast. The fifth fault is the Brooke Fault Zone, northeast of the unnamed fault. Toward its northernmost end, the Brooke Fault Zone is named the Tank Creek Fault (Wheeler, 2005). The Stafford Fault System was originally activated in the Early Cretaceous time. The fault was reactivated at the Fall Hill Fault showing displacement in the Pliocene-Pleistocene sandy gravel, and Cretaceous strata (Mixon, 2000).

Recurrent movement has been demonstrated on the Stafford Fault System by displacements that decrease upward in the Coastal Plain (Mixon, 1977). None of the reports and maps used by Wheeler (Wheeler, 2005) documented Quaternary activity on any faults of the Stafford Fault System. The youngest movement, demonstrating late Tertiary activity, has been documented on the Fall Hill Fault, of the Stafford Fault System, which offsets Pliocene-Pleistocene sandy gravel (Wheeler, 2005). The Stafford Fault System was assigned to Class C based on lack of evidence of Quaternary slip.



Marple (Marple, 2004) suggested a significantly longer Stafford Fault System which extends from Fredericksburg, Virginia to New York City (Marple, 2004). It was proposed as part of a northeastern extension of the East Coast Fault System, previously postulated by Marple and Talwani (Marple, 2000). Existing data do not support the extended Stafford Fault System beyond its previous extent and, despite the suggested correlation of some historical earthquakes with the northern extension of the fault system by Marple and Talwani (Marple, 2000), seismicity data show a poor association between earthquake epicenters and extended segment of the Stafford Fault System.

Based on the foregoing discussion, the Stafford Fault System may not be a capable tectonic source. However it was included as a source zone in the EPRI study (EPRI, 1986). The fault system has a probability of activity of 0.08 in both the Dames and Moore, and the LAW Engineering groups tectonic feature assessments. The numbers mainly reflect the low probabilities assigned by groups because of (1) poor association with seismicity and (2) lack of demonstrated Quaternary slip along the fault.

Information on this feature was reviewed for this study, and this feature is considered to be a capable source, but not a characteristic tectonic source for the BBNPP site.

#### **2.5.1.1.4.4.3 Relevant Tectonic Features with Associated Seismicity**

Within a 200 mi (322 km) radius of the BBNPP site, seven (7) Relevant Tectonic Features with Associated Seismicity (Shown on Figure 2.5-15) were identified. Descriptions follow for the seven (7) features that were evaluated as possible contributors to the seismic hazard at the BBNPP Site.

These features are labeled below with the reference numbers utilized on Figure 2.5-15 in parentheses:

1. New York-Alabama Lineament (28)
2. Hudson River Valley Trend (27)
3. Pittsburgh-Washington Lineament (30)
4. Tyrone-Mt. Union Lineament (29)
5. Bristol Block Geopotential Trends (31)
6. Reading Prong (51)
7. Peekskill-Stamford Seismic Boundary (52)

#### **2.5.1.1.4.4.3.1 New York-Alabama Lineament**

The New York-Alabama Lineament (NY-AL) is a northeast trending lineament characterized by aeromagnetic mapping and regional gravity data which extends more than 1,000 mi (1,609 km) from Alabama to New York (King, 1978). The closest approach of the NY-AL is approximately 30 mi (48 km) west of the BBNPP site (Feature 28 in Figure 2.5-15). The NY-AL in Pennsylvania has been disrupted or offset between two major features called the Tyrone-Mt. Union (TMU) lineament and the Pittsburgh-Washington (PW) lineament (Lavin, 1982). TMU and PW crustal features define the boundaries of a northwest trending feature called the Lake

Erie-Maryland crustal block. A right-lateral offset of 38 mi (61 km) along TMU is indicated by disruption of NY-AL in southwestern Pennsylvania (Muller et al. in Lavin, 1982). Earthquakes have occurred at different locations along the feature, however, association with the feature cannot be established (EPRI, 1986). Johnston et al. (Johnston, 1985a) concluded that between 80% and 90% of southern Appalachian earthquakes recorded from 1981 to 1983 lie between the NY-AL and a parallel structure to the southeast called Clingman-Ocoee lineament. Johnston et al. (Johnston, 1985a) further conclude that the NY-AL and the Clingman-Ocoee lineament do not appear to be seismogenic but rather bound the crustal block that generates the earthquakes. Appalachian seismicity occurs beneath the decollement which separates thrust and folded Paleozoic rocks from Precambrian basement rocks, indicating that Appalachian seismicity is not related to tectonic and geological features at the surface (Johnston, 1985b).

Kaufman and Long (Kaufmann, 1999) inverted travel time residuals from relocated earthquakes in southeastern Tennessee to obtain the velocity structure of the local upper crust. They stated that the results do not agree with the NY-AL as a linear feature extending through southeastern Tennessee parallel to contours in gravity anomalies. The southeastern Tennessee seismicity is not constrained by major crustal features but is rather associated with low velocity regions in midcrustal depths. Joint hypocenter-velocity inverted on the eastern Tennessee seismic zone suggest a strong low-velocity zone parallel to the seismicity with a northeast trend. The southern margin of this trend coincides with the NY-AL. According to Vlahoic (Vlahoic, 1998) research utilizing 3-D velocity earthquake, most earthquakes were located in regions of average velocity or small velocity anomalies and reject the association of eastern Tennessee seismicity with low velocity regions in the crust (Vlahoic, 1998).

Information on this feature was reviewed for this study, and this feature is considered to be a capable source, but not a characteristic tectonic source for the BBNPP site.

#### **2.5.1.1.4.4.3.2 Hudson River Valley Trend**

Also known as the Hudson River Line (HRL), this feature trends north-south for about 156 mi (251 km) along the Hudson River Valley Trend, coming as close as 120 mi (193 km) NE of the BBNPP Site (Feature 27 in Figure 2.5-15). The feature is weakly associated with the western part of the isostatic gravity low at the New Jersey border to the southeastern edge of Adirondack gravity high. Due to large uncertainty in subsurface geometry, the actual structure of the feature is not determined (EPRI, 1986). Based on early instrumentally recorded seismicity (Yang, 1981), the feature was seismically limited. Subsequent observations of seismicity indicate that few earthquakes are located along some parts of the trend. The moderate-sized earthquakes occur only near the edges of the feature. Therefore the overall seismicity does not indicate the localization of activity along the trend. Recent earthquakes west of the Hudson River range in estimated depth from 9 to 12 mi (14 to 19 km), indicating the possibility of a deep structure, but this is not supported by the isostatic gravity data (Yang, 1981). No evidence of the recent tectonic activity has been demonstrated and based on the forgoing discussion, the seismicity is only poorly associated with the feature. One of the six (6) EPRI (1986) Groups, Woodward-Clyde, considers this to be a capable tectonic source. Information on this feature was reviewed for this study, and this feature is considered to be a capable source, but not a characteristic tectonic source for the BBNPP site.

#### **2.5.1.1.4.4.3.3 Pittsburgh-Washington and Tyrone-Mt Union Lineament**

These two major lineaments have been identified from analysis of regional gravity and magnetic patterns, LANDSAT images and geological data (Lavin, 1982). Trending NW-SE, they cross the

Appalachian orogen to the vicinity of Lake Erie (EPRI, 1986). The Pittsburgh-Washington (PW) and Tyrone-Mt. Union (TMU) lineaments are expressions of deep crustal fracture zones which extend over a distance of 375 mi (604 km) across western Pennsylvania and parts of surrounding states. The PW-TMU lineaments are located approximately 115 mi (185 km) southwest of BBNPP site (Features 29 and 30 in Figure 2.5-15) (Rodgers, 1984). The TMU and PW lineaments are parallel and form the NE and SW boundaries of the Lake Erie-Maryland crustal block respectively (Lavin, 1982). Major crustal displacements have occurred along the TMU lineament during late Cambrian to early Ordovician time. There is no concentration of seismicity along the TMU lineament. Evidence for displacement along PW lineament is not as strong as TMU. Concentration of seismicity has been observed near the northern end of PW lineament in northeastern Ohio but earthquake activity is not localized along the feature in general (EPRI, 1986). The PW and TMU features have been evaluated and the probability of activity for both features is very low (EPRI, 1986). Therefore, it is concluded that the PW-TMU lineament system is not considered to be a capable tectonic source.

#### **2.5.1.1.4.4.3.4 Bristol Block Geopotential Trends**

The Bristol Block is an area of magnetic and gravity lows and extends from Tennessee to Pennsylvania. It is bordered by the New York-Alabama lineament on the west, and by the Clingman-Ocoee lineament on the east (Feature 31 on Figure 2.5-15). The northern portion of the Block is located about 80 mi (129 km) southwest of the site (EPRI, 1986). It includes a series of low gravity and magnetic anomalies associated with some earthquakes, since these anomalies extend over a large area. Small earthquakes occur within this block but not all the tectonic features within the block are associated with earthquakes. The Giles County, Virginia seismic zone, which is located within the Bristol block, has been considered separately as seismic source zone (EPRI, 1986). Even though there is associated seismicity with this feature it not considered a capable source, and it is not a characteristic tectonic source for the BBNPP site.

#### **2.5.1.1.4.4.3.5 Reading Prong**

The Reading Prong massif is a major Precambrian complex, one of several outcropping between the sedimentary rocks of the Appalachian basin and the Paleozoic metamorphic terrains of New England and the southeastern Piedmont (Wolf, 2003), with its closest approach to the site approximately 50 mi (80 km) east of the BBNPP (Feature 51 in Figure 2.5-15d). Metamorphism of the Reading Prong rocks occurred during the Grenville Orogeny, approximately 1.2 billion years ago, and the area was also extensively intruded by synorogenic granites during the same Grenville event. Subsequently, between the Late Proterozoic and the Mesozoic, the rocks throughout the region suffered periodic episodes of deformation that was especially intense during the Late Permian Alleghanian Orogeny when the region was pervasively fractured during the development of imbricated thrust sheets (Wolf, 2003). The rocks of the Reading Prong are allochthonous and represent an overlapping stack of thrust sheets that have been thrust over the Paleozoic Rocks of the Great Valley to the north (Senior, 2006). The thrusting juxtaposed the Precambrian rocks in a structurally high position relative to the thick section of lower Paleozoic rocks immediately to the west.

Sykes (Sykes, 2008) as a result of the observations on a catalog of 383 earthquakes in the Greater New York City – Philadelphia Area, refers that most hypocenters are concentrated in older terranes bordering the Mesozoic Newark basin (Reading, Manhattan, and Trenton prongs) and in similar rocks found at a shallow depth beneath the coastal plain from south of New York City across central New Jersey. The discussion about the seismicity of the reading prong is entirely base on his report.

As a result of Sykes (Sykes, 2008) observations a belt of activity of about 30 km (18.6 mi) is located to the northwest of the Newark basin and includes the 1881, 1951, 1957, and 2003 shocks plus numerous smaller earthquakes; most of those events, especially those located instrumentally, are concentrated in the eastern Reading prong between the Green Pond syncline and the Mesozoic Ramapo Fault. They also think that earthquakes occur in stronger basement rocks below those younger sediments, with most computed hypocenters in that area being 5-10 km (3-5 mi), and often poorly constrained.

Other activity in the reading prong extends as far northwest as the westernmost outcrops of Precambrian rock. A band of high activity strikes about N 100 W and extends from the Ramapo Fault across the entire Reading prong. The lower boundary of activity beneath the reading prong appears to dip southeasterly, maybe controlled by one or more Paleozoic imbricate thrust faults that sole into a master sub horizontal detachment fault whose depth is as great as the 12-15 km (7-9 mi) maximum depth of earthquakes beneath the Ramapo Fault.

Information on this feature was reviewed for this study, and this feature is considered to be a capable source, but not a characteristic tectonic source for the BBNPP site.

#### **2.5.1.1.4.4.3.6 Peekskill-Stamford Seismic Boundary**

Instrumental data recorded since 1971 shows that activity in the Manhattan prong cuts off abruptly along a nearly vertical, northwest-striking boundary that extends from Stamford, Connecticut, to Peekskill, New York (Feature 52 on Figure 2.5-15). The boundary is named the Peekskill-Stamford boundary (Sykes, 2008), and is considered a newly identified feature. It was possible to identify it after accurate locations and depths of earthquakes, especially those east of the Hudson River, became available with the installation of a seismic station near the New York–Connecticut border in 1971. The Peekskill-Stamford boundary coincides near its western end with the boundary between the Manhattan prong and the Hudson Highlands, and it may be considered as a ramification of the Ramapo Fault Zone. There has been no large magnitude seismicity associated with this feature, though micro seismic records indicate recent activity with defined epicenter locations. The Peekskill-Stamford boundary is incorporated into the BBNPP PSHA as part of area seismic sources that cover the Ramapo Fault Zone. The maximum magnitude considered for those areas exceeds that of any recorded events in this region.

Information on this feature was reviewed for this study, and this feature is considered to be a capable source, but not a characteristic tectonic source for the BBNPP site.

#### **2.5.1.1.4.4.4 Relevant Tectonic Features with No Associated Seismicity**

A total of 21 tectonic features have been identified with no associated seismicity.

Of these 21, five are located within the site vicinity (25 mi (40 km) radius). Tectonic structures and features closest to the BBNPP site are relatively more important because of their proximity and are discussed first. These five features and other structures listed in this subsection are shown Figure 2.5-27b (site vicinity), and Figure 2.5-15 (site region). Based on review of published literature and historical seismicity, there is no reported geomorphic expression, historical seismicity, or Quaternary deformation along any of the twenty-one features identified below. Thus they are not considered to be a capable tectonic source for calculating the seismic hazard for BBNPP.

- Berwick Anticlinorium (subsection 2.5.1.1.4.4.4.1)
  - Light Street Thrust Fault (subsection 2.5.1.1.4.4.4.1.1)
  - Berwick Fault (subsection 2.5.1.1.4.4.4.1.2)
- Lackawanna Synclinorium (subsection 2.5.1.1.4.4.4.2)
- Anthracite Region (subsection 2.5.1.1.4.4.4.3)
- Scranton Gravity High (subsection 2.5.1.1.4.4.4.4)
- Yellow Breeches Fault Zone (subsection 2.5.1.1.4.4.4.5)
- Rome Trough (subsection 2.5.1.1.4.4.4.6)
- Pleasant Valley-Huntingdon Valley Fault (subsection 2.5.1.1.4.4.4.7)
- Transylvania Fault Zone (subsection 2.5.1.1.4.4.4.8)
- Plimmers Island and Pleasant Grove Shear Zones (subsection 2.5.1.1.4.4.4.9)
- Newark-Gettysburg Basin (subsection 2.5.1.1.4.4.4.10)
- Hartford Basin (subsection 2.5.1.1.4.4.4.11)
- Connecticut Basin (subsection 2.5.1.1.4.4.4.12)
- Brandywine Fault System (subsection 2.5.1.1.4.4.4.13)
- Martic Fault (subsection 2.5.1.1.4.4.4.14)
- East Border Fault (subsection 2.5.1.1.4.4.4.15)
- Catawissa-McCauley Mountain Synclinorium (subsection 2.5.1.1.4.4.4.16)
- Broadtop Synclinorium (subsection 2.5.1.1.4.4.4.17)
- Sweet Arrow Fault (subsection 2.5.1.1.4.4.4.18)
- Chestnut Ridge Anticline (subsection 2.5.1.1.4.4.4.19)

#### **2.5.1.1.4.4.4.1 Berwick Anticlinorium**

The principal bedrock structure within the site area is the Berwick Anticlinorium (also referred to as the Montour Anticline (Pohn, 2001)), which has been described (Inners, 1978) as "a moderately complex, first order fold which trends in a northeast-southwest direction". The bedrock map and section of the Berwick Quadrangle (Inners, 1978) shows the formations at the BBNPP site area to consist of Silurian, Devonian, and Carboniferous rocks that have been gently folded, with limited faulting (Figure 2.5-23 and 2.5-27b). The BBNPP site is situated on the northern limb of the fold, with beds that are steeply dipping. Two faults have been mapped

in the vicinity: the Light Street fault located on the northern limb of the fold, and the Berwick fault, inferred to be on the southern limb of the fold. The northeast ends of both faults lie within the site area, but do not directly underlie the site. Both faults are considered folded faults, therefore there is limited chance for these to be reactivated in the contemporary stress regime.

#### **2.5.1.1.4.4.1.1 Light Street Thrust Fault**

Light Street Thrust Fault is the wedge fault located approximately 2 mi (3.2 km) west of the BBNPP site (Inners, 1978) and is the closest feature to the BBNPP site. The fault was active during the lower Devonian period. Based on studies, the fault dips to the south, at a small angle of 10 to 30 degrees to stratigraphic bedding. The strike of the fault has a northeast-southwest orientation. The fault is located in the north side of the Berwick anticlinorium and extends for about 20 mi (32 km) west of Berwick. The fault overlaps the Old Port and Keyser formations. Seismic reflection profiles indicate that the fault originated during the Triassic (Inners, 1978). Based on review of published literature and historical seismicity, there is no reported geomorphic expression, historical seismicity, or Quaternary deformation along the Light Street Thrust Fault. Thus, this feature is not considered to be a capable tectonic source.

#### **2.5.1.1.4.4.1.2 Berwick Fault**

In conjunction with the Light Street Fault as discussed above, the Berwick Fault is mapped as inferred and is based on limited surface data and a water well log drilled at the Berwick Lumber and Supply Company at 329 West Second Street in Berwick, PA (Inners, 1978). The BBNPP site investigators were unable to locate this well and it was possibly abandoned prior to the site investigation. The inferred Berwick Fault lies within the site area and comes to within approximately 3.5 mi (5.6 km) of the BBNPP site. The exact length of the Berwick Fault is not completely mapped and is believed to be a south-dipping reverse fault on the south flank of the BBNPP (Inners, 1978).

Inners (Inners, 1978) also states that the Berwick fault extends east-northeastward into an exposed third order anticline in the Marcellus-Mahantango interval in the Old Port and Keyser Formations. Inners (Inners, 1978) attributes the presence of the Berwick Fault to the folding and faulting actions that occurred at the site area during the Alleghanian Orogeny, approximately 250 million years ago. There is no pre-EPRI or post-EPRI study (EPRI, 1986) seismicity associated with this feature nor is there any geomorphic evidence of Quaternary deformation. The Berwick fault is not considered to be a capable tectonic source.

A review and interpretation of aerial photography and digital elevation models of the BBNPP site area identified few discontinuous north to northeast-striking lineaments. None of these lineaments were interpreted as fault-related, nor coincident with the Light Street Fault, Berwick Fault, or the other previously inferred Paleozoic structures mapped by Inners (Inners, 1978) and Berg (Berg, 1980). A review of regional geologic sections suggest that the features postulated by Inners (Inners, 1978), if present, are not moderate or prominent structures, and do not deform Quaternary strata. In summary, on the basis of regional and site geologic and geomorphic data, there are no known faults within the site area that pose a structural hazard to the site, including the poorly constrained Light Street and Berwick faults that lie within the southwestern section of the 5 mi (8 km) radius of the site.

#### **2.5.1.1.4.4.2 Lackawanna Synclinorium**

Lackawanna Synclinorium is a first order fold syncline and is mapped within the 5 mi (8 km) radius north-northeast of the BBNPP site area (Figure 2.5-15 and 2.5-27b) (Inners, 1978) as it

trends southwest-northeast. The Lackawanna Synclinorium is a 68 mi (110 km) long structural trough in the Appalachian foreland of northeastern Pennsylvania (Harrison, 2004).

There are numerous unnamed faults within the Lackawanna Synclinorium. These are folded faults and as result there is limited chance for these to be reactivated in the contemporary stress regime.

In map view (Figure 2.5-27b), its hinge displays a concave-to-the-foreland (concave towards the west) curve. The entire synclinorium was thought to be an Alleghanian thin-skinned contractional structure that formed similarly to the fold trains of the central Appalachian Valley and Ridge province (Harrison, 2004). Interpretation of seismic reflection data across the structure, however, suggests that the synclinorium formed primarily by the removal of salt. The trend of the central and northern synclinorium reflects the location of the Upper Silurian Salina salt. In the southern synclinorium, the structure translated northwestward over a thrust ramp joining detachments in the Cambrian Waynesboro Formation and the Salina Group (Harrison, 2004).

In addition to the flexural-slip mechanisms responsible for many of the folds in the mapped area of Figure 2.5-27b, flexural-flow folding was likely another strong component to their formation (Inners, 1978). Inners (Inners, 1978) identifies several features of flexural-slip folds evident in the Berwick Quadrangle including the common occurrence of slickenlines on bedding surfaces and maintenance of approximately the same bedding thickness across the folds. Flexural-flow folding characteristics, within any structure, include prominent cleavage in argillaceous rocks and thickening of beds within the hinges.

Inners (Inners, 1978) prepared a lithostratigraphic column along an almost 9 mi (14.5 km) long stretch of the Berwick Quadrangle that intersects much of the BBNPP site area. When these stratigraphic columns are compiled into a cross section (Figure 2.5-23a), they collectively provide an almost 9 mi (14 km) long, nearly continuous exposure of Silurian, Devonian, and Mississippian formations. Inners' (Inners, 1978) stratigraphic analysis indicates that these Paleozoic formations dip to the north on the north side of the Susquehanna River and dip to the south on the south side. Erosional processes of the Berwick Anticlinorium have produced two mountain ridges, Lee Mountain to the north of the site and Nescopeck Mountain to the south, and have produced similar topography on each mountain (Inners, 1978). The Light Street and Berwick faults are also mapped on this cross section, as seen in Figure 2.5-23a. The apparent structural relief of the Berwick Anticlinorium is approximately 12,000 ft (3,658 m) with a wavelength of about 8.2 mi (13.2 km) (Inners, 1978). Quaternary deposits overlying the site show no signs of faulting or folding.

There is no pre-EPRI or post-EPRI study (EPRI, 1986) seismicity associated with the features (Inners, 1978); the hypothetical features are not aligned or associated with gravity and magnetic anomalies, nor is there data to indicate that the features proposed by Inners (Inners, 1978) are capable tectonic sources.

#### **2.5.1.1.4.4.3 Anthracite Region**

The Anthracite region, located within a northeast plunging syncline, is the most faulted area of the Appalachian Ridge and Valley province. The asymmetric basin is evident in the contrast between the northwestern and southeastern sides, in terms of the intensity and manner of production of folds and tectonic structures (Hornberger, 2004) (Figure 2.5-15). The dominant faults consist of thrust faults, as part of the base décollement. The thrust faults are dominantly located in the cores of the anticlines. These faults tend to be low angle dipping and transect the

bedding planes along the northwest limbs of the folds (Faill 1999). Other faults associated with these structures are steeply dipping and are parallel to the stratigraphic beds located in the southeast limbs of the anticlines. The fault system was active during the post Carboniferous period and located 20 mi (32 km) north of the site (Berg, 1980). Based on review of published literature and historical seismicity, there is no reported geomorphic expression, historical seismicity, or Quaternary deformation along the Anthracite region. Although this feature is not considered to be a capable seismic source, there have been reported earthquake epicenters associated with mining activities (collapse, blasting).

#### **2.5.1.1.4.4.4 Scranton Gravity High**

The Scranton Gravity High (SGH) is located underneath the BBNPP site (Feature 34 in Figure 2.5-15d, and in Figure 2.5-17). The SGH extends about 250 mi (402 km) from Albany, New York, to Harrisburg, Pennsylvania. The large gravity anomaly covers portions of the Appalachian Plateaus and Ridge and Valley Provinces (Hawman, 1992). There is no apparent relationship to the seismicity in the area of the SGH, which is sparse. Earthquakes occur southeast of SGH in southeastern Pennsylvania and Western New Jersey, but there is no localization of events along the feature (EPRI, 1986). Therefore, the SGH is not considered to be a capable tectonic source.

#### **2.5.1.1.4.4.5 Yellow Breeches Fault Zone**

The northeast-striking Yellow Breeches Fault Zone is located within the northeastern portion of the Ridge and Valley Physiographic Province in southwestern Pennsylvania and extends to Virginia (Figure 2.5-15). The Martinsburg/Hamburg foreland segment is divided by the Yellow Breeches fault. The thrust rocks of the Cocalico terrane are oriented northward and are part of the overturned limb of the Lebanon Valley nappe which occurred during the Alleghanian Orogeny (Pohn, 2000). The east-dipping Yellow Breeches Fault, part of the Reading Prong nappe megasystem, is shown as several miles in length (Figure 2.5-15). This décollement represents an upper-level detachment above a deeper décollement about 5 mi (8 km) deep (Faill, 1999). The Yellow Breeches fault is exposed in outcrops of the Ordovician St. Paul Group located within the Ridge and Valley Province. The Yellow Breeches fault zone is not considered a capable tectonic source. Based on published literature, no seismicity is attributed to the Yellow Breeches fault zone and published literature does not indicate that it offsets late Cenozoic deposits or exhibits geomorphic expression indicative of Quaternary deformation. Therefore, this Paleozoic fault is not considered to be a capable tectonic source (Wheeler, 2006).

#### **2.5.1.1.4.4.6 Rome Trough**

The Rome Trough is a Cambrian extensional graben system that extends from northern Tennessee, northeastward through Kentucky, West Virginia, and into western Pennsylvania. This northeast-trending graben, which underlies the Appalachian Plateau Province, is mainly characterized by normal faults of early Paleozoic age. On the other hand complex folds and thrust faults of late Paleozoic age characterize the eastern Appalachian Plateau (Kulander, 2005). Kulander and Ryder (Kulander, 2005) studied data from seismic lines across the Rome trough in West Virginia, western Maryland, and southwestern Pennsylvania. Basement-involved thrust faults have been reported in some parts of the Rome trough and attributed it to regional compression dating 0.8-1.0 Billion years ago. Major normal movements along the Rome trough boundary faults occurred in the Early to Middle Ordovician, and no other movement seems to have occurred since then (Kulander, 2005). The details of basement structure in western Pennsylvania and interpretations of the faults accompanying the Rome trough are based on



limited data (Ryder, 2002). However, as Kulander and Ryder (Kulander, 2005) indicate, the latest movements seem to have occurred in the Middle Ordovician. The association of this feature with seismicity is also limited. Therefore, the Rome trough is not considered to be a capable tectonic source.

#### **2.5.1.1.4.4.7 Pleasant Valley-Huntingdon Valley Fault**

Pleasant Valley-Huntingdon Valley Fault is the eastward continuation of the Cream Valley fault. See Feature 12 in Figure 2.5-15. The fault separates the Baltimore Gneiss, overlain by the Chickies and Ledger Formation on the north, from Wissahickon Formation schist on the south. The fault was active during the post Ordovician period and located 120 mi (193 km) south of the site (Crawford, 1999). Major subvertical northeast striking faults, including the Brandywine Manor fault, the Cream Valley fault, the Pleasant Valley-Huntingdon Valley Fault, and the Rosemont fault, intersect through the blocks containing Grenville-age gneisses and juxtaposed them against younger rocks. The Pleasant Valley-Huntingdon Valley Fault borders the Piedmont Upland Section, Piedmont Lowland Section, and Gettysburg-Newark Lowland Section. Based on review of published literature and historical seismicity, there is no reported geomorphic expression, historical seismicity, or Quaternary deformation along the Huntingdon Valley Fault. Therefore, this feature is not considered to be a capable tectonic source.

#### **2.5.1.1.4.4.8 Transylvania Fault Zone**

The Transylvania Fault Zone, near the latitude 40° N, extends from the Early Mesozoic Gettysburg Basin (Feature 17 in Figure 2.5-15), in Pennsylvania, westward into Ohio and striking at roughly 270 degrees (Dodson, 2008). The fault system is located approximately 170 mi (274 km) west of the BBNPP site (Berg, 1980). The fault zone is mapped as large subvertical east-west trending faults extending through the Blue Ridge, Great Valley and Ridge and Valley provinces. Through the Appalachian Plateau, the fault zones are detected through subsurface records and geophysical studies. Root and Hoskins (Root, 1977) proposed a zone of east-west trending faults which extend from the eastern margin of the Blue Ridge to the Allegheny front near latitude 40° N, for about 75 mi (121 km). The fault zone transects strata nearly across the entire length of Pennsylvania. In the eastern part of the region two faults, 9 mi (14 km) apart, have been previously mapped for about 23 mi (37 km). These are Shippensburg and Carbaugh-Marsh Creek faults which extend east-west in parallel. Root and Hoskins (Root, 1977) describe the following faults in the zone: Sideling Hill, Breezewood, Everett gap, and Wills Mountain faults. Root and Hoskins (Root, 1977) do not consider the fault zone as a major transcurrent fault because the apparent strike-slip movement associated with the fault is no more than 2.5 mi (4 km). Root and Hoskins (Root, 1977) conclude that the Transylvania fault zone is a fundamental fracture, which possibly extends through the continental plate. The fault system originated in the Precambrian, and was reactivated during the Taconic Orogeny in the middle Ordovician and again in the Carboniferous Period during the Alleghanian Orogeny. The Transylvania fault also reactivated in the Early Jurassic (Root, 1977). This fault zone has been included by EPRI (EPRI, 1986) study teams as a tectonic feature but has not been associated with seismicity. Based on review of published literature and historical seismicity, there is no reported geomorphic expression, historical seismicity, or Quaternary deformation along the Transylvania fault zone. Therefore, this feature is not considered to be a capable tectonic source.

#### **2.5.1.1.4.4.9 Plummers Island and Pleasant Grove Shear Zones**

The Plummers Island Shear Zone and Pleasant Grove Shear Zone are thrust faults that include the Plummers Island fault, Pleasant Grove fault, Hyattstown fault and Martic fault (Kunk, 2004).

The faults strike in a northeast-southwest trend with a high angle westerly dipping orientation and are located in the south central Appalachians (Feature 13 in Figure 2.5-15). The fault zones are located 175 mi (282 km) south of the BBNPP site (Kunk, 2004). The faults were active during the Acadian Orogeny in the Early Devonian period and reactivated in the Carboniferous period during the Alleghanian Orogeny, and are part of the south central Appalachians. Based on review of published literature and historical seismicity, there is no reported geomorphic expression, historical seismicity, or Quaternary deformation along the Plummers Island shear zone or the Pleasant Grove shear zone. Therefore, these features are not considered to be capable tectonic sources (Wheeler, 2006).

#### **2.5.1.1.4.4.10 Newark-Gettysburg Basin**

The Newark-Gettysburg (NG) Basin extends from southeast New York through New Jersey to southern Pennsylvania, and is located south approximately 60 mi (97 km) of the BBNPP site (Faill, 1973) (Figure 2.5-15d). The NG Basin, one of the several Triassic basins in eastern North America, has been developed either by downthrown block and subsequent sediment deposition from the northwest and southeast direction or as fault-troughs or grabens with faulting and sedimentation occurring at the same time. The Newark and Gettysburg basins are two separate basins that formed the NG Basin. These two basins along with Culpeper Basin and Barbourville Basin (both in Virginia) are the remnants of a larger Triassic feature called the Birdsboro Basin (Faill, 2004). Faults and folds have tilted and deformed the Birdsboro Basin in Early Jurassic (Faill, 2004). The Newark Basin is bounded to the northwest by the Ramapo Fault system in northern New Jersey and connects (southward) to the Gettysburg Basin in Pennsylvania. Fault structures within the basin strike in the northeast direction. Border faults, normal faulting and wrench faulting are associated with the NG basin. The Furlong Fault, Hopewell Fault and Chalfont Fault were generated by wrenching faulting in the NG basin (Root, 1999).

The Woodward-Clyde Group (EPRI, 1986) concluded that the Newark Basin is seismically limited but the basin may be responsible for the localization of events in the region. Seismicity occurs near the edges of the basin with two notable concentrations of earthquakes. One concentration is near the transition from the Newark Basin into the Gettysburg Basin, which corresponds to the Lancaster Seismic Zone (LSZ) south of the narrowest section of the NG Basin (Armbruster, 1987). The other concentration occurs near the Maryland-West Virginia border, outside the southern edge of the basin. Most well-located epicenters in the LSZ are located just outside the NG Basin (Scharnberger (2006) (EPRI (1986). A detailed description for the LSZ is provided in section 2.5.1.1.4.4.2.10.

#### **2.5.1.1.4.4.11 Hartford Basin**

The Hartford Basin of Massachusetts and Connecticut, with two bounded faults is the largest Mesozoic-age graben in New England. The East Border Fault extends about 130 mi (209 km) from Knee, New Hampshire to New Haven, Connecticut and further to Long Island Sound. The fault is located approximately 180 mi (290 km) east of the BBNPP site (Figure 2.5-15d). The fault strikes generally north and dips west, and changes in strike to north-northeast from central Connecticut toward southern direction (Wheeler, 2005). Stratigraphy at Farm River marsh, nearly 1 mi (1.6 km) inland from the Long Island Sound, showed the downthrown block of the southeastern portion of the marsh with respect to the northwestern portion coincided spatially with the buried trace of the fault. The fault has been active since 2000 years ago and reactivated in the present day compressional field (Wheeler, 2005).

Wheeler (Wheeler, 2005) argues that no fault surface has been reported within the overlying marsh deposits. In addition, the downthrown block shows the displacement across a wide slope as opposed to a sharp offset on a fault plane. The area also lacks evidence for sudden movements that would imply tectonic faulting. Wheeler (Wheeler, 2005) concludes that evidence of faulting has not been reported in Quaternary sediments of the Farm River marsh and accordingly classifies the East Border fault as Class C. Seismicity has not been associated with the East Border fault. Therefore, it is concluded that the Hartford Basin is not a capable tectonic source.

#### **2.5.1.1.4.4.12 Connecticut Basin**

The Connecticut Basin is the largest onshore Mesozoic-age graben in New England (Wheeler, 2005). From Long Island Sound in the south, the basin crosses through central Connecticut and Massachusetts and extends to southern New Hampshire in the north. The basin is located approximately 175 mi (282 km) northeast of the site (Figure 2.5-15d). The sedimentary series is called Newark Super Group. Sedimentation continued until the early Jurassic period during which the basin also experienced intrusive volcanic activity (Bennington, 2006). The Connecticut Basin has been evaluated by Rondout Associates Inc. (RAI) tectonic team (EPRI, 1986). RAI considered the basin as seismically active, because they defined the Connecticut Basin feature as containing the Moodus Seismic Zone. Even though seismicity has been associated with the Connecticut Basin, Quaternary activity has not been demonstrated for the structures within the basin or for its boundary faults. Therefore it is concluded that the Connecticut Basin is not a capable tectonic source.

#### **2.5.1.1.4.4.13 Brandywine Fault System**

The Brandywine Fault System is located approximately 180 mi (290 km) south of the site and north of the Potomac River (Feature 19 in Figure 2.5-15). The 12 to 30 mi (19 to 48 km) long Brandywine fault system consists of a series of en echelon high angle reverse fault segments with associated flexing of the overlying Coastal Plain sedimentary strata. The fault system trends north-northeast with displacements ranging from a few feet to 250 ft (76 m). The Brandywine Fault System consists of the Cheltenham Fault and Danville Fault (Cumbest, 2000). The Brandywine Fault System was active in the Cretaceous and middle Eocene and middle Miocene (Mixon, 1977).

The Brandywine Fault System is located 6 mi to 12 mi (10 km to 19 km) east of the Stafford Fault Zone and strikes roughly parallel to it. Wheeler (Wheeler, 2005) considers Skinkers Neck and Brandywine as a single fault zone, southeast of Stafford Fault System. Compared to Stafford Fault System, the Skinkers Neck-Brandywine Fault Zone is less known and its boundary is shown by dashed line (inferred) in the map of Mixon (Mixon, 2000 in Wheeler, 2005). The last activity of the fault was during the Miocene. There is no seismicity associated and no evidence of Quaternary activity with the fault has been demonstrated. Therefore, the Brandywine Fault System is not considered as a capable tectonic source (Wheeler 2005).

#### **2.5.1.1.4.4.14 Martic Fault**

The Frederick Valley lies near the western edge of the Piedmont physiographic province of Maryland, with the Martic fault being the eastern boundary of this valley (Brezinski, 2007). This fault is located 95 mi (153 km) SSW of the BBNPP site (Feature 11 in Figure 2.5-15).

This fault juxtaposes low-grade metamorphic phyllites against Early Paleozoic carbonates. Little difference on topographic relief exists between the Early Paleozoic carbonates of the Frederick

Valley and the rocks of the Triassic basins because much of the Triassic bedrock is composed of detrital limestone and dolomite whose clasts are Early Paleozoic in age (Brezinski, 2007).

The Martic thrust fault cuts up section westward, placing Ijamsville Phyllite on rocks of the Urbana, Araby, and Frederick Formations; the fault is folded with the footwall rocks. The Barnesville-Monrovia thrust fault places rocks of the Sams Creek Formation on rocks of the Ijamsville Phyllite. This fault is early (now folded) and late (cut existing faults) thrust faults, reactivated thrust faults, and numerous small intraformational faults of limited displacement (Southworth, 2002). Rocks of the Ijamsville Phyllite of the Martic thrust sheet have a composite foliation that consists of a transposition foliation that is overprinted with phyllonitic foliation, and several generations of cleavage. Vein quartz that impregnates the rock has been sheared, transposed, and folded into steep isoclinal folds. These steep F1 folds were deformed by westward-verging, inclined F2 folds that plunge steep to gentle in all directions. These F2 folds have attendant northeast-dipping axial planar pressure solution crenulation cleavage.

There is no seismicity associated with this fault, and it is considered to be non-capable under the present stress regime.

#### **2.5.1.1.4.4.15 East Border Fault**

The East Border Fault is the eastern limit of the Hartford Basin 180 mi (290 km) east of the BBNPP site (Feature 22 in Figure 2.5-15).

Except for the Deep River basin and the Hartford basin the border faults are on the western sides of the exposed basins and dip easterly (Benson, 1992). As revealed by seismic reflection profiles the East Border Fault of the Hartford Basin strikes north and dips west at 180 and continues beneath the basin before merging with sub horizontal reflectors deep in the pre-Mesozoic metamorphic basement (Rodgers, 1985) (Benson, 1992).

Of the two border faults, the East Border fault has the larger net vertical slip (Rodgers, 1985). Where the fault crosses the shore of Long Island Sound, the fault curves more to strike east-northeast before returning to a northeast strike under the Sound (Rodgers, 1985).

Thomson (Thomson 2000 and 2001) reconstructed the past geographic distributions of high marsh, relative sea-level rise curves, and dated paleohistory of vertical offsets interpreted original horizontal peats. Same authors interpreted that the sea level rose in the southern part of the Farm River marsh more rapidly than in the northwestern part of the marsh. The stratigraphy of the southeastern part of the marsh is dominated by low marsh and mudflat deposits, while the stratigraphy of the northwestern marsh is predominantly high marsh deposits. The difference between the two stratigraphies implies to the authors that the southern part of the marsh was downthrown with respect to the northwestern marsh.

Several questions remain to be answered with respect to the Quaternary activity of this fault (Wheeler, 2005). Among these is mainly that there are no reported surface faulting within the till or the overlying marsh deposits, a possible erosive event a possible alternative explanation for parts of the transition, and (Obermeier, 1996) lack of evidence that individual increments of the downthrowing occurred suddenly, within the seconds to minutes that an earthquake rupture lasts, or at least more suddenly than could be produced by any alternative, non faulting mechanism.

These questions and possible alternative explanations mean that faulting has not been demonstrated in the Quaternary sediments of the Farm River marsh, and classified the East Border fault as a Class C feature.

#### **2.5.1.1.4.4.16 Catawissa-McCauley Mountain Synclinorium**

The only mention to this structure is in Inners (Inners, 1978) when he refers to the two first-order folds structures (Lackawanna Synclinorium and Catawissa-McCauley Mountain Synclinorium) shown in Figures 2.5-15 and 2.5-27. The mapped axes of both first-order folds, Inners (Inners, 1978) add, are conjectural when the exact location cannot be determined due to the lack of outcrop in the glaciated valleys.

Inners (Inners, 1978) identified in the extreme southeast corner of the Berwick Quadrangle, a synclinal axis that can be mapped in the Mauch Chunk Formation along Nescopeck Creek that may represent the main axis of the Catawissa-McCauley Mountain Synclinorium. Alternatively, Inners (Inners, 1978) also considers that this axis may define only one of several en echelon second-order synclines that form the Synclinorium in this area.

There is no historic or recorded earthquake epicenters associated with this structure. The Catawissa-McCauley Mountain Synclinorium is not considered to be a capable tectonic source.

#### **2.5.1.1.4.4.17 Broadtop Synclinorium**

The Broad Top synclinorium is a large regional synclinorium which extends from central Pennsylvania to western Virginia. It is the most western of the synclinoria in the Valley and Ridge province (Jacobein, 1974) and is located 100 mi (161 km) WSW of the BBNPP site on its closest position (Feature 48 in Figure 2.5-15). Although the synclinorium is broken into a series of folded and faulted structures, its basic structural style reflects both basement movement and thin-skinned tectonics. It is considered a first order folds which traverses the northwest portion of the quadrangle, the Jacks Mountain Anticline, which crosses the middle of the quadrangle, and part of the northwest flank of an unnamed 1st order synclinorium that is in the southeast portion of the quadrangle. Many smaller folds, from 3rd to 5th order, occur within the major folds (McElroy, 2007).

More than 125 mi of seismic surveys, 22 wells, and surface maps were analyzed by Jacobeen (Jacobein, 1974) to determine the precise relation of basement to thin-skinned tectonics. This relation, as interpreted, indicates that Taconic and older, tension-induced features have a pronounced effect on the localization and genesis of the ramping of décollements. This ramping, in turn, produces many of the large, prominent, first-order structural features in both the Valley and Ridge, and the Plateau provinces of the Appalachian basin. The Allegheny Front is an example of such a feature. The décollement ramping also induces formation of smaller, second-order features and other flexures and faults within the Broad Top synclinorium.

Folds in the Broad Top synclinorium strike in an average N30°E direction and have been mapped as long, continuous features, extending over 100 mi (161 km). The anticlines and synclines are shown roughly parallel and doubly plunging. Along their axes they rise and fall in a series of domes and saddles. Along the strike of the structures, outcrops indicate an en echelon structural trend of anticlinal and synclinal axes (Jacobein, 1974). The southwest plunge of the Broad Top Synclinorium results in the youngest rocks in the quadrangle, the Brallier Formation, is being exposed along its axis on the western border of the quadrangle. Dips are steeper on its southeast flank and there are smaller folds contained within it. Overturning of fold limbs occurs north and west of Garret Knob (McElroy, 2007).

The structural style of the Broad Top synclinorium reflects both basement movement and thin-skinned tectonics. The structure as seen today represents the interaction between Taconic block faulting and late Paleozoic thrusting. Such late Paleozoic thrusting, which dominates the structure in the Appalachians, is believed to be modified and localized by older structures in other areas.

There is no seismicity associated with this fault, and it is considered to be non-capable under the present stress regime.

#### **2.5.1.1.4.4.18 Sweet Arrow Fault**

The Sweet Arrow fault, located between Swatara Gap and New Ringgold, in east-central Pennsylvania, is located 50 miles (80 km) to the southeast of the site (Feature 49 on Figure 2.5-15). It is a thrust fault separating the overturned rocks to the north from the upright rocks to the south (Wood, 1960), and has a dip of 40° to 70° to the south and a throw that can reach several thousand feet. It was recognized by Wood (Wood, 1960) from the thinning and offsetting of stratigraphic units and the truncation of structural features in the area. The fault cuts through rocks of Silurian and Devonian ages, and is buried in many places beneath heavy talus slopes. Wood (Wood, 1960) never mentioned the cross-cutting of younger rocks, including rocks of Quaternary age, but the sense of the motion on the fault (thrust from the southeast) is consistent with the stress regime that existed during the Alleghanian Orogeny, when other similar faults were active. There is no seismicity associated with this fault, and it is considered to be non-capable under the present stress regime.

#### **2.5.1.1.4.4.19 Chestnut Ridge Anticline**

A widely accepted structural model for folds in the outer central Appalachian foreland is partially based on the geologic structure of the North Summit field. The model includes a simple surface anticline that is detached in Silurian Salina Group salts and cored by imbricated Devonian Tully-Helderberg rocks thrust inward toward a depressed axial low (Shumaker, 2002).

The Chestnut Ridge anticline is a major regional feature of the Appalachian fold belt (Feature 50 on Figure 15). Its western limb forms a 100 mi-long (185 km) segment of a physiographic break called the Intra-Plateau Front (Gwinn, 1964). The Front separates the Allegheny Mountains (Figure 1), with their high-relief open folds (1000 ft [305 m] of relief), from the lowlands and low folds of the Pittsburgh Plateau to the west (Shumaker, 2002).

Space problems within collapsing synclines above competent reservoir rocks, the Huntersville-Helderberg lithostructural unit, are resolved by distortion and evacuation of overlying, incompetent Hamilton rocks. Huntersville-Helderberg rocks deformed into a variety of structural shapes, not solely the imbricated model that traditionally has been applied to Plateau folds of the central Appalachian foreland.

#### **2.5.1.1.4.4.5 Seismic Sources as Interpreted by EPRI Groups**

In 1986, the Electric Power Research Institute (EPRI) developed a seismic source model for the Central and Eastern United States (CEUS), which included the BBNPP Site region (EPRI, 1986). The CEUS is a stable continental region (SCR) characterized by low rates of crustal deformation and no active plate boundary conditions. The EPRI source model included the independent interpretations of six Earth Science Teams and reflected the general state of knowledge of the geoscience community as of 1986. Each of these teams developed a tectonic framework, defined as the collection of tectonic features thought to have a non-negligible

probability of generating magnitude 5 (mb) or greater earthquakes in the present stress regime due to tectonic processes, using comprehensive geophysical and seismological databases compiled in initial stages of the project. In order to develop their individual framework each of the six groups did: (1) interpret the crustal stress regime, (2) identify the tectonic features that might produce moderate to large earthquakes, (3) list and evaluate criteria for assessing the likelihood of activity of those features, and (4) quantify the probability of activity of each feature. Using this tectonic framework they later extended the tectonic evaluations and seismic source zones assessment to the entire Central and Eastern United States, for which they estimated seismicity parameters based on advanced and traditional analyses of the historical earthquake data set by means of a consistent, systematic format with full documentation (EPRI, 1986).

Seismic source zones have been configured either to coincide with tectonic features, or to envelope tectonic features and adjacent patterns of observed seismicity. The original seismic sources identified by EPRI are thoroughly described in the EPRI study reports (EPRI, 1986) and details of the interpretations from each of the six Tectonic Evaluation Contractors (TECs) are summarized in Sections 2.5.1.1.4.4.5.1 through 2.5.1.1.4.4.5.6. Each of these sections presents particulars of the interpretations from each of the six Tectonic Evaluation Contractors (TECs): Weston Geophysical Corporation, Dames & Moore, Law Engineering Testing Company, Woodward-Clyde Consultants, Bechtel Group, Inc., and Rondout Associates Incorporated. The evaluations of each team are supported by figures that show their tectonic interpretation as well as the seismic zones that they specifically defined for the CEUS. All figures include plots of earthquakes with mb equal to or higher than 3.0 in the updated earthquake catalog, in order to illustrate the spatial relationships between seismicity and the tectonic features and their associated seismicity, if this connection exists. Every TEC was initially responsible for a particular region of the Central and Eastern United States. After these regional tectonics evaluations were made, each of the TECs extended the tectonic evaluations and source zone assessments to the entire Central and Eastern United States (EPRI, 1986).

Section 2.5.1.1.4.4 summarizes specific tectonic features and characteristics within these different source areas, their evolution in time and the current state of knowledge on the tectonic setting and tectonic structures in the site region and beyond. Section 2.5.2 gives details of the procedure by which each of the TEC's resulting seismic source zones distribution was evaluated and the calculation of their Gutenberg-Reuter parameters, as well as the integration of the resulting seismic hazard for the BBNPP Site.

#### **2.5.1.1.4.4.5.1 Tectonic Interpretations by Bechtel Group, Inc.**

Bechtel Group, Inc. tectonic interpretation and proposed seismic source zone configuration and dependencies are presented in corresponding Figures 2.5-16 and Figure 2.5-16f.

Their interpretation for tectonic features affecting the site region (Figure 2.5-16) includes: Charlevoix – La Malbaie (3), Clarendon-Linden Fault (11), Eastern Mesozoic basins (13), Ramapo Fault (14), Stafford Fault (17), Lebanon-Pennsylvania geopotential trend (23), Bristol Block geopotential trends (24), and the New York – Alabama geopotential lineament.

Seismic source zones contributing to 99 percent of BBNPP Site Hazard, illustrated in Figure 2.5-16f, are defined by six (6) source zones and three (3) background zones. These seismic source zones are first, and primarily, identified for each tectonic feature with an assessed activity of 0.05 or greater. Five (5) of the six sources zones were defined like this and were given a number designator; they include: Charlevoix – La Malbaie (3), Clarendon-Linden (11), Mesozoic Basins (13), Bristol Block (24), and the New York – Alabama Lineament (25). A second type of seismic source was defined for areas where either no adequate tectonic feature

could be identified, or where the features that were identified were not assessed with a high enough likelihood to adequately explain the location of seismicity in the area considered. The remaining seismic source was defined this way and was given a letter designator; this one is the Niagara (D). Three (3) background zones are left after all seismic sources have been considered: Southern Appalachians (BZ5), Southern Eastern Craton (BZ6), and the Northern Eastern Craton (BZ7) (EPRI, 1986).

#### **2.5.1.1.4.4.5.2 Tectonic Interpretations by Dames & Moore**

Dames & Moore tectonic interpretation and proposed seismic source zone configuration and dependencies are presented in corresponding Figure 2.5-16a and Figure 2.5-16g.

Their tectonic feature assessment for the site region (Figure 2.5-16a) includes: St. Lawrence River Valley Rift (1), La Malbaie/Charlevoix Zone (2), and Clarendon-Linden Fault Zone (27); Newark-Gettysburg Basin (14) as part of the Mesozoic/Cenozoic Rifts; Eocambrian Rifts including Scranton Gravity Anomaly (12) and Pennsylvania Aulacogen (13); and Ramapo Fault (39), as part of the Mesozoic/Cenozoic Faults.

Seismic source zones contributing to 99 percent of BBNPP Site Hazard, illustrated in Figure 2.5-16g, are defined by eight (8) source zones. These eight seismic source zones are: Adirondacks Zone (3), Paleozoic Fold Belts (4), Eastern Marginal Basin (8), Clarendon-Linden Zone (9), Southern Cratonic Margin (41, Default Source Zone), Newark-Gettysburg Basin (42), Southern Appalachian Mobile Belt (53, Default Source Zone), and La Malbaie/Charlevoix Zone (59) (EPRI, 1986).

#### **2.5.1.1.4.4.5.3 Tectonic Interpretations by Law Engineering Testing Company**

Law Engineering Testing Company tectonic interpretation and proposed seismic source zone configuration and dependencies are presented in corresponding Figures 2.5-16b and 2.5-16h.

Their interpretation for tectonic features affecting the site region (Figures 2.5-16b) comprises: Pennsylvania Aulacogen and Scranton Gravity High (Eocambrian Rifts); St. Lawrence Valley Rift and Charlevoix-La Malbaie Area (St. Lawrence Rift System as part of the Mesozoic Rifts); Adirondacks (Uplift as part of the Mesozoic Rifts); Stafford-Brandywine Fault Zones, Clarendon-Linden Structure, Ramapo Fault, and Latitude 400 N Fault Zone (Faults and Shear Zones).

Seismic source zones contributing to 99 percent of BBNPP Site Hazard, illustrated in Figure 2.5-16h, are defined by four (4) source zones. The four seismic source zones are: Buried East Coast Mesozoic Basins (8), Charlevoix-La Malbaie Area (12), Eastern Basement (17), and Reactivated Eastern Seaboard Normal Faults (22) (EPRI, 1986).

#### **2.5.1.1.4.4.5.4 Tectonic Interpretations by Rondout Associates, Inc.**

Rondout Associates, Incorporated tectonic interpretation and proposed seismic source zone configuration and dependencies are presented in corresponding Figures 2.5-16c and 2.5-16i.

Their interpretation for tectonic features affecting the site region (Figure 2.5-16c) includes: Clarendon-Linden (CL), Gravity Lineament (F), Inboard Mesozoic (IMEF), Niagara Magnetic Anomaly (NMA), New York – Alabama Lineament (NY-AL), Pittsburgh Washington Lineament (PW), Scranton Gravity High (SH), St. Lawrence Rift (SLR), Tyrone-Mt. Union lineament (TMU), and a Gravity Anomaly (X).



Seismic source zones contributing to 99 percent of BBNPP Site Hazard, illustrated in Figure 2.5-16i, are defined by six (6) source zones, and one (1) background zone. The six seismic source zones are: Shenandoah (30), Quakers (31), Niagara by the Lake (33), Nessmuk (34), La Malbaie (37), and Vermont (41). The Background Region Seismic Sources considered is Grenville Crust (50) (EPRI, 1986).

#### **2.5.1.1.4.4.5.5 Tectonic Interpretations by Weston Geophysical Corporation**

Weston Geophysical Corporation tectonic interpretation and proposed seismic source zone configuration and dependencies are presented in corresponding Figures 2.5-16d and 2.5-16j.

Their interpretation for tectonic features affecting the site region (Figure 2.5-16d) includes: Adirondack Mountains, The Charlevoix-La Malbaie, Clarendon-Linden Structure, N.Y.-Alabama /Clingman lineaments block, St. Lawrence Rift, Appalachian Plateau, Valley and Ridge Fold and Thrust Zone, Mesozoic Basins, and the Moodus Seismic Area.

Seismic source zones contributing to 99 percent of BBNPP Site Hazard, illustrated in Figure 2.5-16j, are defined by six (6) source zones and three (3) background zones. The six seismic source zones are: Charlevoix-La Malbaie Seismic Zone (1), Adirondack Mountains (6), Clarendon-Linden (8), New York Nexus (21), and the Zones of Mesozoic Basins (28B and 28E). The three Background Region Seismic Sources are: Appalachian Plateau (102), Southern Appalachian (103) and The Southern Coastal Plain (104) (EPRI, 1986).

#### **2.5.1.1.4.4.5.6 Tectonic Interpretations by Woodward-Clyde Consultants**

Woodward-Clyde Consultants tectonic interpretation and proposed seismic source zone configuration and dependencies are presented in corresponding Figures 2.5-16e and 2.5-16k.

Their interpretation for tectonic features affecting the site region (Figure 2.5-16e) includes: Charlevoix-La Malbaie Structure (12), Adirondack Dome/Uplift (18), Mohawk River Valley Trend (20a), New Jersey Isostatic Gravity Saddle (21), Hudson River Valley Trend (25), Western New York/Southern Ontario Crustal Block (33), Attica-NY intersection Feature (34), Tyrone-Mt. Union Lineament (61), and Pittsburgh-Washington Lineament (63).

Seismic source zones contributing to 99 percent of BBNPP Site Hazard, illustrated in Figure 2.5-16k, are defined by 9 seismic source zones equivalent to the tectonic features described in previous paragraph. There is one defined background zoning: Susquehanna Background (BZ16) (EPRI, 1986).

### 2.5.1.2 Site Geology

The U.S. EPR FSAR includes the following COL Item in Section 2.5.1.2:

Site-specific geology information will be addressed by the COL applicant.

This COL Item is addressed as follows:

{Sections 2.5.1.2.1 through 2.5.1.2.6 are added as a supplement to the U.S. EPR FSAR.

The following subsections provide a summary of geologic conditions in the BBNPP site area (5 mi (8 km) radius) and site location (0.6 mi (1km) radius). These subsections present information concerning the physiography, geomorphology, geologic history, stratigraphy, structural geology, geologic hazard evaluation, and engineering geology evaluation related to the BBNPP site. The information presented is based on a review of previous reports and documents, a review of geologic literature, communications with geologists and other researchers who are familiar with previous studies in the site area, and geotechnical and geologic field investigations conducted at and in the vicinity of the BBNPP site. A geologic map of site area (5 km (8 km) radius) is shown on Figure 2.5-27, and a geologic map of the site location (0.6 mi (1-km) radius) is shown on Figure 2.5-26.

#### 2.5.1.2.1 Site Area Physiography and Geomorphology

The BBNPP site area is located within the Susquehanna Lowland Section of the Ridge and Valley Physiographic Province and is bordered by the Appalachian Plateaus Province to the west and north, and the New England Province to the east (Figure 2.5-7 and Figure 2.5-8). Additionally, within the Ridge and Valley Physiographic Province, the Anthracite Valley Section of the Ridge and Valley Province borders the Susquehanna Lowland Section to the north of the site and the Anthracite Upland Section borders on the south.

The topography within 5 mi (8 km) of the site consists of low to moderately high, linear ridges and valleys that primarily follow structural trends with elevations ranging from about 260 ft (79 m) msl to 2,368 ft (722 m) msl (Figure 2.5-3). To the north of the site, elevation increases to 1,500 ft (457 m) msl at the peak of Lee Mountain. With the increase in elevation, the steepness in the slope, from the top of the mountain to the banks of the Susquehanna River, also increases to near vertical northeast of the site. Approximately 10,000 ft (3,048 m) south of the site, on the south side of the Susquehanna River is a steep embankment creating the base of Nescopeck Mountain, which reaches an elevation of approximately 2,368 ft (722 m) msl.

**The site is well-drained principally by Walker Run, which flows from north to south along the west border of the site then to the south and southwest until it reaches the Susquehanna River. A tributary to Walker Run flows to the northwest within the site area (0.6 mile (1km) radius). There is also a third unnamed, unmapped tributary that flows from east to west through the middle of the site, feeds the mapped wetlands area to the southwest of the, and ultimately drains into Walker Run. Drainage within 5 miles of the site is a primarily a moderately developed trellis pattern with ridges and valleys forming sub-parallel drainage systems. These connect at abrupt incisions nearly normal to the ridge, or at the termination of a ridge, and drain down gradient toward the Susquehanna River. Streams that incise into the glacial deposits establish a dendritic pattern that is not subject to the effect of geologic structure. The longest stream near the site is the previously mentioned Walker Run which is approximately 2 mi (3.2 km) long and ultimately drains into the Susquehanna River. Ephemeral stream channels in the vicinity of the**

BBNPP site flow directly into the Susquehanna River. These limited stream channels maintain dendritic patterns that are observed to be minimally affected by the geologic structure as they incise into the underlying Trimmers Rock and Mahantango formations and glacial sediments.

The site area bedrock geologic map (Figure 2.5-27), compiled by (Berg, 1980) and (King, 1974) indicates that most of the site area (5 mi (8km) radius) surrounding, and including, the BBNPP site are underlain by Late Silurian, Devonian, and Lower Mississippian bedrock. More specifically, the site bedrock is Upper Devonian Mahantango Formation. The most recent geologic influence on the site, not including sub aerial exposure and erosion, is the Wisconsinan glaciation that is discussed in Section 2.5.1.1.1. The relevance of this event to the physiography of the site is the deposition of glacial materials. Remnants of kame terrace and outwash labeled as stratified drift deposits of sand and gravel outwash depicted in Figure 2.5-9 are mapped as overlying site bedrock at approximate elevation 660 ft (210m) msl. This relationship is shown on the geologic cross section of the Berwick Quadrangle (Figure 2.5-23a)

Confers Lane and the western boundary for SSES form the eastern boundary of the BBNPP site. The SSES is located approximately 200 ft (61 m) above the Susquehanna River with a long, gradual slope leading from the SSES to the banks of the river. Private property borders the BBNPP site to the immediate south and west. Approximately 7,000 ft (2,134 m) south of the site lays the Susquehanna River. The BBNPP will be constructed at a final grade elevation of 674 ft (205 m) msl and will be set back approximately 7,500 ft (2,286 m) from the Susquehanna River bank.

An additional significant geomorphological/physiographic aspect within the site area is the formation of moderate karst geomorphology approximately five miles to the south and west within the Tonoloway and Keyser Formations. The location of these strata are shown to the west of the site in Figure 2.5-27. The Tonoloway Formation is a thinly bedded limestone up to 100 ft. (30 m) thick in the site area (5 mi (8 km) radius) and the Keyser Formation is a fossiliferous limestone up to 125 ft. (38 m) thick. Both lithologies are susceptible to the development of karst features due to dissolution of calcium carbonate within the rock. However, the formation of karst geomorphology is not applicable to the site location (0.6 mi (1 km) radius), as carbonate rocks are not at or near the surface of the site. Figure 2.5-21 establishes the depth to carbonate rocks at more than 1,500 ft.(457 m) beneath the surface at the site location (0.6 mi (1 km) radius). Core holes from field studies for the BBNPP site extend to depths as great as 150 ft. (46 m) and do not notes the **stratigraphic position of these formations beneath the site location (0.6 mi (1 km) radius)**. Quaternary processes affecting the site are limited to weathering and erosion of existing material, and subjection to the regional stress field that affects the eastern half of the continent. The site area (5.0 mile (8km) radius) exhibits no physiographic characteristic that would detrimentally affect the BBNPP site.

#### **2.5.1.2.2 Site Area Geologic History**

This subsection presents an overview of the geologic history of the site area. The overall geologic history and tectonic framework of the region are outlined in Subsections 2.5.1.1.2 and 2.5.1.1.4. A detailed discussion of the surface faulting within 25 mi (40 Km) of the BBNPP site area is provided in Subsection 2.5.3. The following geologic history of the area around the BBNPP site is summarized based on the recent detailed field studies, and literature review which includes compilations by the Pennsylvania Department of Conservation and Natural Resources and publications by (Braun, 2007). Each has been integral in characterizing the site area (5.0 mile (8km) radius).

The site area geologic history prior to the early Ordovician is inferred from scattered borehole data, geophysical surveys and a synthesis of published information. Limited geophysical and borehole data indicate that the basement rock beneath the site most likely consists of a crystalline metamorphosed greenschist or amphibolite (Gold, 2008). Although the basement has not been penetrated directly beneath the site with drill holes, regional geologic cross sections developed from geophysical, gravity and aeromagnetic, as well as limited deep borehole stratigraphic data beyond the site area suggest Precambrian (approximately +542 million years ago) rocks are most likely present at a depth of about 33,000 ft (10,058 m) beneath the site Figure 2.5-19 shows the inferred depth of the Precambrian basement rocks beneath the BBNPP site.

Tectonic models discussed in Section 2.5.1.1.2 and Section 2.5.1.1.4 establish that the crystalline basement was accreted to the Proterozoic North American Craton during the Grenville Orogeny. The site area during the Cambrian era represents passive continental margin marked by carbonate shelf deposition with clays and silts deposited beyond the carbonate shelf. These sediments, deposited over a period of nearly 50 million years (approximately 542-488 million years ago), comprise the shales and limestones overlying the Grenville basement at the BBNPP site. The depositional environment for the Cambrian bedrocks underlying the site is primarily a marine carbonate platform extending to a more distal setting upon which shales were deposited (Gao et al, 2000). The site area during the Ordovician (approximately 488-444 million years ago) represented an area of continued and increased deposition in an elongated basin along the continental margin. As discussed in Section 2.5.1.1.2.3, the Middle Ordovician to Late Ordovician (Hurowitz, 2004) with respect to the eastern margin of the North American craton, is affected by the Taconic Orogeny Deposition of material from the Taconic Allochthon, a suite of oceanic and arc terranes (Trembley, Bedard, and Lauziere 1997) as it approached from the east-southeast, yielded significant amounts of siliciclastic sediment to the basin. Shale and sandstone units that now outcrop in places throughout central Pennsylvania but underlie the site area by over thousands of feet (Thompson, 1999) were deposited at this time. Uplift of the area as the Taconic allochthon converged with the North American Craton shifted the depositional environment from that of carbonate platform to a much shallower basin receiving coarser siliciclastic source material. As the Taconic front converged with the craton, portions of the ocean floor and island arc of the Taconic front were obducted onto the craton at the eastern portion of what is now defined as the site region (200 mi (322 km) radius) while some ocean floor subducted under the continent. The site area (5 mi (8 km) radius), at the close of the Ordovician, is on the western portion of a depositional basin bounded on the west by the North American Craton and on the east by the Taconic Highlands (Thompson, 1999).

The continued erosion of the Taconic Mountains during the early part of the Silurian Period (444~416 million years ago) added to the sedimentation throughout the site area. This early period of sedimentation differed from that of the Ordovician, in that the sand and gravels deposited during the early Silurian were extremely quartz-rich in mineral composition (Barnes, 2002), thus creating a very erosion-resistant sandstone. After the Taconic Mountains were almost completely eroded away during the Late Ordovician, carbonate sedimentation continued through the site area creating shale and limestone intervals (Laughrey, 1999).

**The Devonian Period (approximately 416-359 million years ago) primarily marks the Acadian Orogeny (as discussed in Section 2.5.1.1.2.5), the result of the North American landmass colliding with the Avalon Terrane and Baltica (Eusden, et al, 2000). The site area remained a basin area while the Acadian mountain range, to the east of the site area, was subjected to erosional processes. These eroded sediments were deposited in the site area and are**

represented by the modern day black and gray shales underlying the site (Ver Straeten and Brett, 1999). The effect of the Acadian Orogeny is more pronounced in terms of metamorphism and magmatism in the northeastern United States (Maine, New Hampshire, Vermont), (Eusdan et al, 1999) but is significant to the site area in that the source material for the formation of the Devonian rocks that outcrop in the site area was primarily derived from the Acadian Highlands. (Ver Straeten and Brett, 1999) identify the progression of a terrigenous clastic wedge from the orogenic belt to the west during multiple phases of orogenic uplift during the Devonian that was the source of Devonian deposition.

The Carboniferous Period (approximately 359-299 million years ago) is best described in two distinct categories including the older Mississippian and younger Pennsylvanian sub-periods of the Carboniferous Era. The Mississippian (approximately 359-318 million years ago) was a time of continued sedimentation from the Acadian mountains but a change of meteorological climate is represented in the lack of oxidation of rocks from this time frame (Barnes, 2002). Near the end of the Mississippian the site and surrounding area likely became poorly drained and was an area of thick forests and swamps. During the early Pennsylvanian these forests deposited great amounts of organics which did not rot or oxidize due to the water rich environment of the swamps (Edmunds, 1999). These organics were then overlain by sediment deposits and compacted into the coal fields that occupy the Anthracite Valley Section to the north and the Anthracite Upland Section to the south of the site (Edmunds, 1999) (Figure 2.5-8). No Mississippian or Pennsylvanian rocks exist within the site location (0.6 mi (1km) radius) due to uplift and erosion throughout geologic time. Mississippian and Pennsylvanian rocks are found within the site area (5.0 mi (8km) to the north and to the south of the site location as depicted on Figure 2.5-27.

The most significant geologic event to affect the structure of the site area is the Alleghanian Orogeny (discussed in detail in Section 2.5.1.1.2.5). Beginning in the middle Pennsylvanian, continental convergence of North America and Africa began the multiple phases of the Alleghanian Orogeny which was the primary orogenic event to form the portion of the Appalachian Mountains that extend from northeast Alabama through Pennsylvania. In addition to the Appalachian Mountains, the orogeny formed numerous thrust faults, fractures, anticlines and synclines throughout the Ridge and Valley, Piedmont, and Blue Ridge Physiographic Provinces from Alabama through Pennsylvania including the site area. Convergence from the southeast to the northwest created a structural fabric that strikes primarily northeast and dips to the northwest and southeast. Deformation was both ductile and brittle depending on pressure/temperature conditions. Offset along faults within the Ridge and Valley is primarily reverse dip slip but often has a strike slip component (Pohn, 2001).

The site area was subjected to brittle deformation in the form of folding and thrusting that developed the structural makeup of the Ridge and Valley Province within which the site area lies. The Berwick and Light Street Faults, depicted on Figure 2.5-27 and discussed in Section 2.5.1.2.4.1 are examples of this Alleghanian deformation within the site area (5 mi (8 km) radius), and are not recognized as active. In addition, the Berwick Anticline, an east-northeast striking, gently northeast plunging anticline trends directly through the site area (5 mi (8 km) radius). The Berwick Anticline is an asymmetric structure in the site area (5 mi (8 km) radius) with the north-northwest limb dipping steeply to the north-northwest and the southern limb dipping more gently to the south-southeast. The structure imparted by the Alleghanian Orogeny is significant to the topography, drainage, and seismicity of the site area, defining the major landforms (elongated ridges and valleys), drainage patterns, and structural discontinuities within the Paleozoic strata. By the end of the Permian Period, the Appalachian Mountains had been subjected to significant erosion providing source material for an alluvial plane across the site

area (Edmunds, 1999). Publications (Pazzaglia, 2006,) describe the cycle of erosion and isostatic uplift as the dominant geologic process within the site area since the late Permian, with the exception of discrete glacial periods that significantly altered and even reversed the erosion and uplift process. Cenozoic faulting in the Piedmont, especially along the contact of the Piedmont and Coastal plain, where a hinge line of erosional unloading and depositional loading roughly from northwest to southeast (Weems, 1998) is known to exist. Cenozoic faulting is also recognized in the Piedmont (Wentworth and Mergner-Keefer 1981) approximately 50 miles from the site. In addition, Crone and Wheeler (Crone and Wheeler, 2000) present data for all Quaternary faults and possible tectonic features in the Eastern United States and report two sites in the Pennsylvania Ridge and Valley province, but no seismic events or tectonic structures are reported within the site area (5 mi (8km) or the site vicinity (25 mi (40 km) radius). Capable seismic sources, their location, and affect on the site area (5 mi (8 km) radius) are discussed in Section 2.5.1.2.4.

The Triassic (approximately 251-199 million years ago), Jurassic (approximately 199-145 million years ago), and the Cretaceous (approximately 145-65 million years ago) Periods, all within the Mesozoic Era, were all time zones of slow erosional processes for many areas of Pennsylvania, including the site area (Barnes, 2002). During these periods of erosion, new drainage patterns and streams were formed. By the end of the Cretaceous Period, chemical alteration became the dominant erosional source of rocks that likely comprised the site area, changing them into clays and saprolite (MacLachlan, 1999). Rifting of Pangea in the late Permian and throughout much of the Triassic affected the site region (200 mi, (322km) radius) by changing the stress field to an extensional regime. Mesozoic rift basins and intrusive diabase dikes along rift structures mark the extensional effects on the North American Atlantic margin. The aforementioned extension related features do not occur within the site vicinity (25 mi (40-km) radius) and site area (5 mi (8 km) radius) (Froelich and Gottfried, 1985).

Though there is little record as to what happened at or around the site area during the Cenozoic Period (approximately 65 million years to present), much can be inferred from the glacial deposits of the Quaternary Period (approximately 1.8 million years ago to present). It is theorized that during the Tertiary Period (approximately 65-1.8 million years ago) erosion at the site area continued with chemical alteration and erosion, primarily during the early Tertiary, but transitioned to intense physical erosion during the Late Tertiary (Barnes, 2002). This physical erosion was the result of cooler and drier conditions. It is also believed that during this time many of the modern day rivers and streams, such as the Susquehanna, established themselves (Sevon, 1999). During the Quaternary, continental glaciers covered Canada and advanced into a small portion of Pennsylvania, including the site area (5 mi (8 km) radius), and site vicinity (25 mi (40-km) radius) as shown in Figure 2.5-9. The site area was subjected to three different periods of glaciation with the earliest occurring approximately 770,000 years ago and the most recent occurring approximately 17,000 to 22,000 yrs ago. These periods of glacial advance and retreat had both erosional and depositional effects on the site area, the degree of which is difficult to quantify. The main effect of glaciation on the site area was the effect of glacial outwash enhancing drainage patterns that were already in progress (Sevon, 1999). During and after glacial retreat, the site has been an area of deposit for stratified drift which includes sand and gravel, kame terraces, and outwash. Surficial deposits are depicted on Figure 2.5-24 for the site location (0.6 mi (1 km) radius) and on Figure 2.5-25 for the Berwick Quadrangle.

The surface at the site is comprised of glacial deposits from the Wisconsinan glaciation that is discussed in Section 2.5.1.1.2. As briefly discussed in Section 2.5.1.2.1, with respect to physiography, the formation of karst is also part of the geologic history of the site area (5 mi (8 km) radius). Approximately five miles to the west within the Tonoloway and Keyser Formations,

as shown on Figure 2.5-27, karst feature development is exhibited. The Tonoloway Formation is a thinly bedded limestone up to 100 ft. (30m) thick in the site area (5 mi (8 km) radius) and the Keyser Formation is a fossiliferous limestone up to 125 ft. (38 m) thick. Both lithologies are susceptible to the development of karst features due to dissolution of calcium carbonate within the rock. However, the formation of karst geomorphology is not applicable to the site location (0.6 mi (1 km) radius), as carbonate rocks are not at or near the surface of the site (Inners, 1978). Figure 2.5-21 notes the stratigraphic position of these formations beneath the site location (0.6 mi (1 km) radius). The Tonoloway and Keyser Formations are at depths as great as 1,500 feet under the site location (0.6 mi (1 km) radius) due to the position of the site on the steeply dipping northern limb of the previously discussed northeast plunging Berwick Anticline. The dip of the northern limb and plunge of the axis of this anticline places the Keyser and Tonoloway at significant depths under the BBNPP site location (0.6 mi (1 km) radius). The geologic history of the site is complex, but the current geologic processes affecting the site are limited to weathering and erosion of existing material, and subjection to the regional stress field that affects the passive Atlantic margin (Figure 2.5-10b, (Heidbach, 2008)) shows the current stress fields in the eastern portion of North America, and minimal isostatic uplift. With respect to seismic stability and geologic hazards due to the site area (5 mi (8 km) radius) geologic history, the site area (5 mi (8 km) radius) and site location (0.6 mi (1 km) radius) are positioned in a stable geologic setting.

#### **2.5.1.2.3 Site Area Stratigraphy**

As discussed in Section 2.5.1.1, the BBNPP site is located near the deepest portion of the Appalachian structural basin (see Figure 2.5-14). Approximately 33,000 ft (10,058 m) of Paleozoic rocks overlie the Pre-Cambrian basement at the BBNPP site. The regional source of sediments, the types of sediments, and the deformation of these sediments that fill the Basin are described in Section 2.5.1.1. This section (2.5.1.2.3) provides more detailed information related to the local stratigraphy, including the nomenclature of the formations, the lithologies, and the thickness of formations. In addition, any characteristics of the geologic media underlying the site that could impact the design or the safety of the BBNPP are presented in this section. Based on the drilling, geologic observations, and laboratory testing results, the rock formations below the BBNPP site are very competent and physically suitable to serve as a stable foundation for the BBNPP.

Site-specific information on the stratigraphy of geologic materials underlying and directly adjacent to the BBNPP site is based on the geologic/geotechnical investigations performed for the BBNPP site and the historical investigations performed at the SSES site. The SSES site is located immediately adjacent to the BBNPP site and observations concerning its soil and rock formations are comparable with those from the BBNPP investigation borings. Given that the geologic conditions at the SSES and BBNPP sites are very similar, the available SSES data and information were used to supplement the BBNPP data and enhance the characterization of subsurface conditions at BBNPP.

A total of 45 boreholes and 41 monitoring wells were completed at BBNPP in late 2007 and early 2008 for sampling and standard penetration test (STP) purposes. In addition, 3 boreholes were performed for geophysical testing purposes. Of the 48 geotechnical and geophysical borings drilled during the BBNPP site investigation, 14 were 70 to 99 ft (21 to 30 m) deep, 24 were 100 to 199 ft (30 to 61 m) deep, six were advanced to a total depth between 200 and 299 ft (61 and 91 m) below ground surface (bgs), and four were drilled to a depth between 300 and 405 ft (91 and 123 m) bgs (Table 2.5-28). The deepest geotechnical boring advanced at the BBNPP site was boring G-301, which was drilled to a total depth of 405 ft (123 m) near the proposed centerpoint of the reactor building. All of the geotechnical and geophysical borings

penetrated the full thickness of the glacial overburden deposits and at least 21 ft (6.4 m) of the Middle Devonian Mahantango Formation. In addition, approximately 250 exploratory borings were drilled in soil and rock in late 1970, Spring 1971, and 1983 at or near the SSES site. Also, test pits were excavated at selected locations at or near the SSES site.

It is estimated that the full thickness of the Mahantango Formation at the BBNPP site is approximately 1,500 ft (457 m) thick (Inners, 1978). Thus, there are approximately 1,100 ft (335 m) or more of Mahantango Formation that lies beneath the deepest geotechnical boring at the BBNPP site. Geological information relative to the uppermost 400 ft (123 m) of geologic strata at the site was derived from cores and drilling information obtained during the on-site BBNPP subsurface investigations. Stratigraphic information regarding the deeper geologic strata was obtained primarily from publications by the Pennsylvania Topographic and Geologic Survey which focus primarily on this area (e.g., Inners, 1978; Crowl, 1980; and Williams, 1987).

Characteristics of the basement rocks in Pennsylvania are extrapolated from the exposed metamorphic rock in the Piedmont Physiographic Province to the southeast and several exploratory wells in western Pennsylvania (Saylor, 1999). Based on available data (Saylor, 1999), it is likely that the basement rock beneath the site is similar to the schists and gneisses found in the Piedmont Physiographic Province, which is located approximately 50 mi (80 km) to the southeast of the BBNPP site as shown in Figure 2.5-6.

Paleozoic rocks of the Ridge and Valley Province are representative of formations deposited in a foreland basin that have undergone numerous cycles of marine regression/transgression. The Province extends from eastern New Jersey to Alabama and has been subjected to multiple orogenies. Stratigraphic formations, and their depositional environments, are also discussed in Sections 2.5.1.1.2 and 2.5.1.1.3. A stratigraphic column was created specifically for the BBNPP site and is shown in Figure 2.5-21. Stratigraphic columns for the site area (5-mi (8-km) radius) are presented in Figure 2.5-21. The following Subsections describe each of the stratigraphic formations, in order of oldest to youngest.

#### **2.5.1.2.3.1 Cambrian Formations**

The oldest inferred Cambrian Formation underlying the site area is the Waynesboro Formation. The Waynesboro Formation consists of sandstone with interbedded red and green shales and has a thickness of approximately 1,000 ft (305 m) or more (Kauffman, 1999). Overlying the Waynesboro Formation is the Pleasant Hill Formation, which is primarily a limestone formation with interbedded sandy and silty layers throughout (Kauffman, 1999). Overlying the Pleasant Hill Formation is the Warrior Formation. It is a dark, fossiliferous, fine grained limestone interbedded with silty dolomite, with a thickness up to 1,340 ft (408 m) (Kauffman (1999). Overlying the Warrior Formation, and marking the Cambrian-Ordovician boundary, is the Gatesburg Formation. The Gatesburg Formation consists of a series of sequential sandstone and dolomite units that are fossiliferous (Ryder, 1992) and in excess of 1211 ft (369 m) thick (Gold, 2003). Both the Warrior and Gatesburg formations likely represent a shallow-water carbonate bank or shelf that was subjected to periodic episodes of near-drying conditions (Kauffman, 1999).

#### **2.5.1.2.3.2 Ordovician Formations**

Overlying the Upper Cambrian Gatesburg Formation are the Lower Ordovician formations that comprise the Beekmantown Group. These Lower Ordovician formations, from oldest to most recent, include the Stonehenge Formation, Nittany Dolomite, Axemann Limestone, and



Bellefonte Dolomite. They are composed primarily of dolomites and limestones (Harper, 2003) and reach a combined thickness of up 4,200 ft (1,280 m) (Thompson, 1999).

The Middle Devonian strata include the Loysburg Formation, the Black River Group, and the Trenton Group. The Loysburg Formation is typically a dolomitic and stromatalite rich limestone underlying a coarse grained, fossiliferous limestone (Thompson, 1999) with an average thickness of 263 to 475 ft (80 to 145 m). Overlying the Loysburg Formation is the Black River Group that consists of the Snyder and Linden Hall formations (Thompson, 1999) and attains a thickness of about 632 ft (193 m). These formations are composed primarily of siliciclastic clay and shale. Overlying the Black River Group are fine-grained, black, graded limestone and shales of the Solona and Coburn formations of the Trenton Group (Thompson, 1999).

Rocks of the Beekmantown Group, Loysburg Formation, Black River Group, Solona Formation, and Coburn Formation were deposited in marine to marginal-marine environments. Where a platform existed and the seas became progressively shallower, depositional environments became more intertidal (Thompson, 1999). The uppermost rocks within the Trenton Group belong to the Antes Formation, a fossiliferous, generally black, shale (Thompson, 1999) that was likely deposited in shallow water, above the wave base. The Antes, Coburn, and Solona formations collectively attain a thickness of approximately 842 ft (257 m).

Upper Ordovician strata include the Reedsville, Bald Eagle, and Juniata formations (in ascending order). The Reedsville Formation is approximately 600 to 1800 ft (183-549 m) thick and is composed primarily of interbedded shale and sandstone beds, with some limestone (Thompson, 1999; Gold, 2003). Like the Antes Formation underlying it, the Reedsville Formation was likely deposited in shallow water. The Bald Eagle Formation and the Juniata Formation are 700 to 1313 ft (213 to 400 m) and 600 to 1,125 ft (183 to 343 m) thick, respectively (Thompson, 1999; Gold, 2003). Both of these formations consist of nonfossiliferous sandstones, conglomerates, and mudstones, but differ in color with the Bald Eagle being gray and the Juniata red (Thompson, 1999). Unlike the Reedsville Formation, the Bald Eagle and Juniata formations are non-fossiliferous and non-marine. These formations were likely deposited near low-sinuosity streams on alluvial fans (Thompson, 1999).

#### **2.5.1.2.3.3 Silurian Formations**

The Tuscarora Formation typically marks the boundary between Upper Ordovician and Silurian Periods. The Lower Silurian Tuscarora Formation is quartzose, sublithic, and argillaceous sandstone with few shale beds throughout, and was deposited in a fluvial environment (Laughrey, 1999). The Tuscarora Formation ranges between 400 ft (122 m) and 700 ft (213 m) thick and is extremely resistant to erosion and weathering processes (Laughrey, 1999; Gold, 2003).

Overlying the Tuscarora Formation (in ascending order) are the Rose Hill, Keefer, Mifflintown, Bloomsburg, Wills Creek, Tonoloway, and Keyser formations, all of Silurian age (Figure 2.5-21). The Rose Hill Formation consists of olive shale with interbedded layers of hematitic sandstone, purplish shale, and fossiliferous limestone (Laughrey, 1999). Above the Rose Hill Formation lies the Keefer Formation, a quartzose and hematitic sandstone with some mudstone. The Rose Hill and Keefer formations combine for a thickness that ranges between 600 ft (183 m) and 1070 ft (326 m) (Gold, 2003). The Mifflintown Formation reaches a thickness of about 336 ft (102 m) (Gold, 2003) and is composed of sandy and silty shales, siltstone, and limestone of a shallow marine setting (Laughrey, 1999). The likely depositional environment for the Rose Hill, Keefer, and Mifflintown formations is that of a submarine ramp that deepened from the proximal

basin margin (Laughrey, 1999) during the Taconic Orogeny, as discussed in Sections 2.5.1.1.2.3 and 2.5.1.1.3.

Conformably overlying the Mifflintown Formation is the Bloomsburg Formation (Lower Silurian), which consists primarily of grayish-red clay-siltstone with some interbedded fine- to coarse-grained sandstone beds. The Bloomsburg Formation ranges in thickness from 85 ft (26 m) to 464 ft (141 m). It is slightly fossiliferous and probably represents sediments deposited in deltaic waters with a high enough salinity to allow some fauna to exist (Laughrey, 1999).

Upper Silurian strata include the Wills Creek, Tonoloway, and Keyser formations. The Wills Creek Formation conformably overlies the Bloomsburg Formation, and consists primarily of claystone to silty claystone with some interbedded grayish-red sandstone and fine-grained argillaceous dolostone and dolomitic limestone. The Wills Creek Formation is approximately 750 ft (229 m) thick (Inners, 1978) and represents deposition in lagoonal, intertidal, and supratidal environments. The Tonoloway Formation is primarily a thinly-bedded limestone deposit with a few thin beds of calcareous shale (Laughrey, 1999); it is about 100 ft (30 m) thick (Inners, 1978). Both the Wills Creek and Tonoloway formations represent numerous shallowing-upward cycles that have been interpreted as repeated progradational events on very large tidal flats (Laughrey, 1999).

The Keyser Formation conformably overlies the Tonoloway Formation and consists mainly of medium to dark gray, fine-grained, fossiliferous limestone with minor interbeds of dark-gray calcareous clay shale. Some dark gray cherty nodules are present toward the upper part of the formation. The Keyser Formation straddles the boundary between the Late Silurian and Early Devonian as the formation represents continuous carbonate sedimentation from both periods and has a thickness of about 125 ft (38 m) (Inners, 1978).

#### **2.5.1.2.3.4 Devonian Formations**

The Devonian system of rocks is described as a westward-thinning wedge of sediments with a thickness of almost 11,000 ft (3,353 m) throughout much of Pennsylvania (Harper, 1999). The Upper Keyser Formation makes up the basal unit for the Devonian period formations. Overlying the Keyser Formation is the Old Port Formation, which consists of (in ascending order) the Corriganville Limestone, the Mandata Shale, Shriver Chert, and Ridgeley Sandstone (Harper, 1999). The Corriganville Limestone consists of finely crystalline, thick to thinly bedded limestone and ranges from 10 ft (3 m) to 30 ft (9 m) thick (Harper, 1999). The Mandata Shale is dark gray to black, thinly bedded, siliceous shale, and ranges in thickness from 20 ft (6 m) to 100 ft (30 m) (Harper, 1999). Light colored cherty, mudstones and calcareous siltstones characterize the Shriver Chert (Harper, 1999), which ranges in thickness from 80 ft (24 m) to 170 ft (52 m). The Ridgeley Sandstone ranges in thickness from 8 ft (2 m) to 170 ft (52 m) and is generally white to light-gray, medium grained, quartzose sandstone (Harper, 1999). These units of the Old Port Formation represent the gradual deepening of the Appalachian basin and range in overall thickness from 100 ft (30 m) to 150 ft (46 m) in the site area (Inners, 1978).

The Silurian limestone formations (Wills Creek, Tonoloway, and Keyser) and the Devonian Old Port Formation contain minor dissolution features along bedding planes, fractures, and joints, which makes them relatively good aquifers at shallow depth (see Section 2.4.12). The total thickness of these formations is approximately 700 feet (213 m) and they are located at least 1,700 ft (518 m) below the BBNPP site. In general, due to their depth, dissolution features along bedding planes or fractures at this depth are very minor and will have no effect on the structural stability or strength of the overlying rock strata.

The Onondaga Formation (Middle Devonian) disconformably overlies the Old Port Formation and reaches a thickness of about 175 ft (53 m) in the site area (Inners, 1978). The Onondaga Formation consists of silty, shaley, and cherty limestones, and likely represents a shelf margin depositional environment (Harper, 1999). The upper part is a medium dark gray argillaceous, fine-grained, pyritic limestone with interbeds of medium dark gray calcareous clay shale (Selinsgrove Limestone Member) and the lower part is a medium to dark gray, calcareous, slightly silty, clay shale (Needmore Shale Member).

The Marcellus and Mahantango formations together form the Hamilton Group of Middle Devonian age (Figure 2.5-21). The Marcellus Formation, located below the Mahantango Formation, consists of approximately 350 ft (107 m) (Inners, 1978) of dark-gray to black shales that are carbonaceous, and contain pyrite and few fossils (Harper, 1999). The Marcellus Formation was likely deposited in a variety of shallow-water anoxic environments (Harper, 1999).

The Mahantango Formation overlies the Marcellus Shale and is the uppermost bedrock unit at the BBNPP site, as shown in Figure 2.5-21 and Figure 2.5-23. Harper (Harper, 1999) describes the Mahantango Formation as a complex series of interbedded shales, siltstones, and sandstones, ranging from 1,200 ft (366 m) to 2,200 ft (671 m) thick. Inners (Inners, 1978) reports the thickness of Mahantango Formation is approximately 1,500 ft (457 m) in the site area. According to Inners (Inners, 1978), the Mahantango Formation consists primarily of "medium dark to dark gray, silty to very silty claystone (95 percent), with some argillaceous, fine-grained limestone (5 percent) in the uppermost part. Claystone generally in poorly-defined beds 3 to 10 feet (1 to 3 m) thick; massive, often with profuse burrows; noncalcareous to very calcareous; medium-dark-gray, pyritic siderite and limestone nodules, 1 to 4 inches (2.5 to 10 cm) in diameter, are locally abundant and often form distinct bands; several calcareous, fossiliferous horizons 3 to 10 feet (1 to 3 m) thick and abounding in marine invertebrates, can be traced locally." The Tully Member (or Tully Limestone Member) is found at the top of the Mahantango Formation and is medium dark gray (light gray where weathered), fossiliferous, argillaceous, fine-grained limestone and calcareous clay shale. It is approximately 50 to 75 feet (15 to 23 m) thick, but has been eroded from the crest of the Berwick Anticline at the site (Inners, 1978). The Mahantango Formation and the Tully Member represent marine sediments deposited on a shallow subtidal shelf. Harper (1999) suggests that the Mahantango Formation was deposited as a prograding marine shoreline during the early stages of the Catskill Delta.

The shales and siltstones encountered during the BBNPP site investigation were typically dark gray, ranged in hardness from soft to moderately hard, increased progressively in the level of calcareous content with depth, and were slightly pyritic and fossiliferous throughout. The weathered shale that formed the uppermost bedrock surface was typically 10 ft (3 m) thick or less. Below the weathered shale, the rock becomes much harder and the rock quality designation (RQD) was relatively high; the RQD below the weathered zone generally ranges between 95 and 100 percent (see Figure 2.5-96, Figure 2.5-97, and figure 2.5-99). In some cases, the Mahantango Shale was so unfractured and competent that the shale cores were removed from the core barrels intact with no breaks at all (Figure 2.5-27a). As discussed in Section 2.5.4.2, the Mahantango shale is a hard, competent bedrock unit. According to Inners (Inners, 1978), the Mahantango Formation has "high foundation support strength, suitable for heavy structures if excavated to sound bedrock." Physical properties of the shale bedrock are discussed in detail in Section 2.5.4.2.

The Mahantango is the uppermost bedrock unit present throughout most of the BBNPP site, including all of the areas supporting safety-related structures. Another black shale (the Harrell Shale) is present only in the northernmost portion of the site and is about 120 ft (37 m) thick

(Inners, 1978). It conformably overlies the Mahantango Formation and marks the initial unit of the Upper Devonian strata. The Harrell Formation is typically represented by dark gray to grayish black, pyritic, organic-rich, clay shale and silty clay shales (Inners, 1978; Harper, 1999). The Harrell Shale was deposited in a poorly-oxygenated basinal marine environment. Where exposed today at the ground surface, the shale has moderately developed cleavage, is jointed, and forms splintery and platy fragments. The weathered surfaces can be covered by efflorescent secondary sulfate minerals. Because the shale has low resistance to weathering, it typically forms a swale at the foot of escarpments developed by the Trimmers Rock Formation (Inners, 1978). The swale that runs east to west along the northern edge of the SSES and the BBNPP sites (immediately south of Beach Grove Road) is shown as the outcrop area of Harrell Shale by Inners (Inners, 1978; Figure 2.5-23)

While the Mahantango and Harrell formations are the only two bedrock formations exposed at the BBNPP site, other Paleozoic Era formations that were deposited after and are younger than the Harrell Formation exist within the site area (5 mi (8 km) radius), as shown in Figure 2.5-23, Figure 2.5-26, and Figure 2.5-27. These formations comprise the ridges and bedrock outcrops to the north and south of the site, such as Lee and Nescopeck Mountains, respectively.

The Upper Devonian Trimmers Rock Formation lies immediately above the Harrell Shale. It is composed primarily of medium to dark gray siltstone and shale (60 percent), medium to medium dark gray, thinly-bedded, very fine to fine grained sandstone (25 percent), and medium to dark gray silty clay shale (Inners, 1978; Harper, 1999). The Trimmers Rock Formation has a calculated thickness of approximately 3,000 ft (914 m) (Inners, 1978) and likely represents a delta-fed submarine slope of the Appalachian Basin. The Trimmers Rock Formation is resistant to erosion and forms steep escarpments on the north and south sides of the Susquehanna River Valley. The east-west trending unnamed ridge that lies directly north of the BBNPP site is composed of Trimmers Rock Formation and the ground surface rises steeply going up the south flank of this ridge (Figure 2.5-4).

Above the Trimmers Rock Formation, within the site area, lie the members of the Upper Devonian Catskill Formation including (in ascending order) the Irish Valley, Sherman Creek, and Duncannon Members. Each member of the Catskill ranges in thickness from 150 ft (46 m) to 3,700 ft (1,128 m) and generally consists of gray to red mudstones, claystones, siltstones, and conglomerates that were deposited in mixed continental, fluvial-deltaic, and marginal-marine environments (Inners, 1978; Harper, 1999).

The uppermost unit of Devonian age rocks in the site area is the Spechty Kopf Formation, which also spans into, and identifies the beginning of the Carboniferous Period. The Spechty Kopf Formation has a thickness of about 575 ft (175 m) near the crest of Nescopeck Mountain, but is absent north of BBNPP on Lee Mountain (Inners, 1978). It is composed mainly of medium gray to olive sandstone with other components including siltstone, shale, and conglomerates (Berg, 1999). The likely depositional environment of the Spechty Kopf Formation was that of ephemeral lakes formed on the surface of the Catskill alluvial plain (Berg, 1999).

#### **2.5.1.2.3.5 Mississippian Formations**

Mississippian rocks of the site area represent a transition from the prograding deltas of the Late Devonian Period (Brezinkski, 1999) to Pennsylvanian rocks, which are primarily represent sedimentary environments within an elongate basin aligned in a northeast to southwest direction (Edmunds, 1999).

The Devonian-Mississippian boundary is marked by the presence of the upper Spechty Kopf Formation. Unconformably overlying the Spechty Kopf Formation is the Pocono Formation, which was likely deposited on a high-gradient alluvial plain or alluvial fan, is represented by the non-red beds of medium to coarse grained sandstone, siltstone, and conglomerates (Brezinski, 1999) with a thickness of about 600-650 ft (183-198 m) (Inners, 1978). Overlying the Pocono Formation within the 5 mile site area radius, is the Mauch Chunk Formation, easily recognizable by its red to reddish-brown mudstone and siltstone with reddish-brown and greenish-gray sandstones and conglomerates (Brezinski, 1999). The Mauch Chunk ranges in thickness throughout the site area but has been estimated to be between 3,000 ft (914 m) to 4,000 ft (1,219 m) thick (Brezinski, 1999). The depositional environment of the Mauch Chunk Formation was likely that of a broad alluvial plain in which deposits came from two distinct sources. The first source was red clastics, likely derived from the Taconic highlands, and the second was the non-red, quartz sand from the erosion of the previously deposited sandstones (Brezinski, 1999).

#### **2.5.1.2.3.6 Pennsylvanian Formations**

The Mississippian-Pennsylvanian boundary in the site area is generally considered to be the top of the Mauch Chunk Formation and bottom of the Pottsville Formation. The Pennsylvanian Pottsville Formation overlies the Mauch Chunk Formation conformably and ranges in thickness from 100 ft (30 m) to 1,600 ft (488 m) (Edmunds, 1999). The Pottsville Formation consists mainly of a cobble and pebble conglomerate with some sandstones and finer clastics and coal (Edmunds, 1999).

The youngest rock formation within a 5-mile (8-km) radius of the site area, and overlying the Pottsville Formation, is the Llewellyn Formation. The Llewellyn Formation reaches a thickness of approximately 3,500 ft (1,067 m) through other portions of Pennsylvania and generally consists of subgraywacke clastics, ranging from conglomerates to clay shale and containing numerous coal beds (Edmunds, 1999). The Llewellyn Formation is very resistant to erosion and forms the crest of Lee Mountain near the town of the Shickshinny, as seen in Figure 2.5-3.

#### **2.5.1.2.3.7 Quaternary Deposits**

The bedrock at the BBNPP site and the site area is covered with a variable thickness of Pleistocene sediments including kame, kame terrace, outwash, ground moraine and end moraine deposits (Figure 2.5-24). Quaternary deposits of the site area are primarily the result of glacial deposition during at least three known glacial events that are believed to have impacted the site area. Of these three events, Quaternary deposits from two of them comprise the soil overburdens present within the site area. The earliest deposit is of Late Illinoian age (unnamed) and can be stratigraphically correlated to that of the Titusville Till in northwestern Pennsylvania, as shown in Figure 2.5-9. These Unnamed and Titusville Tills are described as a thin, gray to brown and grayish-red clay and sand (Sevon, 2000). This was almost entirely eroded away during the next glacial advance through the site during the Wisconsinan Epoch (Crowl, 1999). The resulting glacial deposits from the Wisconsinan event is known as Olean Till, which is described as moderately thick, gray to grayish-red sandy till (Sevon, 2000).

In addition to glacial till covering the upland areas, the BBNPP site and areas in lower elevations (e.g., Susquehanna River floodplain) are covered with thick deposits of stratified drift, as shown in Figure 2.5-25. The stratified drift, as defined by Sevon (2000), is sand and gravel in eskers, kame terraces, and outwash. Stratified drift has been deposited the site area since the Late Illinoian (Sevon, 2000), during periods of glacial melting and retreat.

Up to 60 feet (18 m) of glacial deposits underlie the safety-related structures of BBNPP and contain mostly large boulders, gravel, and coarse sands. The gravel and boulder content diminishes with depth. The lowermost glacial deposits contain fine to medium sand that is better sorted. These glacial deposits are very difficult to drill through or to excavate. Except for the uppermost soils, these deposits are saturated, cohesionless, and very permeable. These glacial deposits will be excavated and removed from beneath the safety-related structures during the construction of BBNPP.

#### **2.5.1.2.4 Site Area Structural Geology**

The structural geology of the BBNPP site described in this section has been developed using the following sources:

- the BBNPP site reconnaissance and subsurface explorations performed for this study,
- results of earlier investigations performed at the SSES site, and
- published geologic literature.

There have been no geologic structures identified during this investigation that pose a threat to the BBNPP site operations. The following discussions will support that conclusion by presenting the information on the geologic structures present within 25 miles (40 km) of the BBNPP site, and with increasing detail within the 5 mile (8 km) and 0.6 mile (1 km) distances from the site.

The following discussions present the current understanding of the geologic structures within the site vicinity, starting with the deep crystalline basement, followed by the overlying folded and lithified sedimentary bedrock, and concluding with the on-site investigations conducted for this report. As reported in more detail in Section 2.5.1.1.4, there are faults and folds that occur within the 25 mi (40 km) BBNPP site vicinity (Figures 2.5-23 and 2.5-27b). These include the Light Street and Berwick faults within the Berwick Anticlinorium, the Lackawanna Synclinorium and the Catawissa-McCauley Mountain Synclinorium. Each will be discussed in turn in the following sections.

##### **2.5.1.2.4.1 Structures in the Crystalline Basement**

Available geophysical data in surrounding areas indicate that the basement likely consists of exotic crystalline magmatic arc material (Hansen, 1986; Glover, 1995). Regional geologic cross sections developed from geophysical, gravity, and aeromagnetic surveys, as well as limited deep borehole data from outside of the BBNPP site area, suggest that complexly deformed, metamorphosed crystalline igneous rocks of Precambrian and Paleozoic age are likely present at approximately -33,000 ft (-10,058 m) msl (Figures 2.5-14 and 2.5-19) (Crawford, 1999; King, 1974; and Gold, 2005). The basement map in Figure 2.5-14 (Gold, 2005) confirms the depth to the basement rocks as well as the relative featurelessness of that surface beneath the site.

To supplement the discussion of basement structures, regional and site vicinity maps of the gravity and magnetic fields are presented in Figures 2.5-17 and 2.5-17a (Kucks, 1999) and Figures 2.5-18 and 2.5-18a (Bankey, 2002). None of these data reveal new anomalies related to geologic structures. The following sections discuss the local gravity and magnetic anomalies, as presented in more detail in Section 2.5.1.1.4.3.

#### **Gravity Data and Features**

Anomalies in gravity data occur from density contrasts in size and depth of geologic structures. Gravity highs, or positive anomalies, are created by abundant thickness or shallow burial of dense features, while gravity lows are from mass deficiencies. Long wavelengths show deep structures or highly concentrated deep structures, and shorter wavelengths are created by shallower structures (Lavin, 1999).

As shown on Figure 2.5-17a, the only feature relevant to the site vicinity is the Scranton gravity high, that has a northeast trend that is sub-parallel to the folds in the current mountains. The long wavelength of this high and the shallow gradients along its margins indicates its deep burial. The Scranton gravity high is related to dense mafic material that was emplaced during late Precambrian rifting (Lavin, 1999).

### **Magnetic Data and Features**

Magnetic data compiled for the 2002 Magnetic Anomaly Map of North America reveal numerous northeast-southwest trending magnetic anomalies, generally parallel to the structural features of the Appalachian orogenic belt (Bankey, 2002) (Figure 2.5-18a). The magnetic map allows a visualization of the geological structure of the upper crust in the subsurface showing the spatial geometry of bodies of rock and the presence of faults and folds.

The BBNPP site is located in an area with few anomalies and with low gradients between anomalies. These conditions are consistent with the profound basement depths in the vicinity of the site. The BBNPP site is situated in the middle of a northeast-trending area of low magnetic relief bounded to the northwest by the New York-Alabama Lineament and to the southeast by the Clingman-Ocoee Lineament. The Clingman-Ocoee block is a block of Precambrian basement bounded by the NY-AL and Clingman-Ocoee lineaments (Johnston, 1985b).

The magnetic data published since the mid-1980's confirm the interpretations and provide additional documentation of previous observations (i.e., EPRI, 1986) across this region of eastern North America, and do not reveal any new anomalies.

#### **2.5.1.2.4.2 Berwick Anticlinorium**

The principal bedrock structure within the site area is the Berwick Anticlinorium (also referred to as the Montour Anticline (Pohn, 2001)), which has been described (Inners, 1978) as "a moderately complex, first order fold which trends in a northeast-southwest direction". The bedrock map and section of the Berwick Quadrangle (Inners, 1978) shows the formations at the BBNPP site area to consist of Silurian, Devonian, and Carboniferous rocks that have been gently folded, with limited faulting (Figure 2.5-23 and 2.5-23a). The BBNPP site is situated on the northern limb of the fold, with beds that are steeply dipping. Two faults have been mapped in the vicinity: the Light Street fault located on the northern limb of the fold, and the Berwick fault, inferred to be on the southern limb of the fold. The northeast end of both faults lies within the site area, but do not directly underlie the site.

Unconformably overlying this Paleozoic bedrock is a thin covering of a few flat-lying post-Paleozoic formations composed primarily of glacial till and colluvium overburden, laid down by multiple Quaternary glacial events (Inners, 1978) (Figures 2.5-24 and 2.5-25).

#### **Light Street Fault**

The 20 mile (32 km) long Light Street fault approaches to within 2 mi (3.2 km) of the BBNPP site (Figures 2.5-23 and 2.5-23a). The Light Street Fault brings into contact the Marcellus

Formation from the Onondaga and Old Port formations. Based on published data (Inners, 1978) and lack of historical seismicity (EPRI, 1986), the nature of this fault is described as being either (1) a reverse fault that dips in a southerly direction and eliminates a section of the stratigraphy between the Wills Creek and Marcellus formations; (2) the detachment of a major décollement that dips to the north; or (3) a combination of (1) and (2) (Inners, 1978). The simplest explanation for the occurrence in the area is a south-dipping reverse fault as mapped in Figures 2.5-23 and 2.5-23a.

The presence of the Light Street Fault is attributed to the folding and faulting actions that occurred at the site area during the Alleghanian Orogeny, approximately 250 million years ago (Inners, 1978). Figure 2.5-23a shows that the fault does not cut the Quaternary surface deposits overlying the area. This is confirmed by the surficial geologic map (Figure 2.5-25) that also shows no evidence of the fault penetrating those deposits.

There is no historical seismicity that is spatially related to this fault (Figure 2.5-15d) (EPRI, 1986). A paleoseismological study of geologic features thought to result from Quaternary tectonic faulting in the Eastern United States (Wheeler, 2006), did not identify this fault as being active.

Because of the age of last motion on this fault, and its lack of historical seismicity, the Light Street Fault is not considered to be a capable tectonic source and poses no threat to BBNPP.

#### **Berwick Fault**

The Berwick Fault identifies a possible shortening of the section between the Old Port and Keyser formations on the southern limb of the fold. The fault is inferred from limited surface data and from a water well log drilled at the Berwick Lumber and Supply Company at 329 West Second Street in Berwick, PA (Figures 2.5-23 and 2.5-23a) (Inners, 1978). The investigators were unable to locate this well and it appears to have been abandoned prior to the site investigation. The inferred Berwick Fault lies within the site area and its northeastern end comes to within approximately 3.5 mi (5.6 km) of the BBNPP site. The length of the Berwick Fault is not completely mapped and is believed to be a south-dipping reverse fault on the south flank of the Berwick Anticlinorium (Figure 2.5-23a) (Inners, 1978). The Berwick fault extends east-northeastward into an exposed third order anticline in the Marcellus-Mahantango interval (Inners, 1978).

This fault is like the Light Street fault, in respect to both its age and its lack of recent seismicity. Because the Berwick Fault is attributed to the folding and faulting actions that occurred approximately 250 million years ago (Inners, 1978), and because there is no seismicity that is spatially related to this feature (Figure 2-15d) (EPRI, 1986) (Wheeler, 2006), the Berwick fault is not considered a capable tectonic source, and poses no threat to BBNPP.

#### **2.5.1.2.4.3 Lackawanna and Catawissa-McCauley Mountain Synclinoria**

Sectional views of the site geology (Figures 2.5-19 and 2.5-23a) show these two folds to be synclines located beneath Lee and Huntington Mountains to the north (Lackawanna Synclinorium), and just beyond Nescopeck Mountain to the south (Catawissa-McCauley Mountain Synclinorium). The southern end of the Lackawanna Synclinorium (Inners, 1978) lies within the site area and a portion of it is shown in the northwest corner of Figure 2.5-23. The Lackawanna Synclinorium is a coal-bearing 68 mi (110 km) long structural trough in the Appalachian foreland of northeastern Pennsylvania (Harrison, 2004), that extends to the northeast of the site. The entire synclinorium was thought to be an Alleghanian thin-skinned



contractional structure that formed similarly to the fold trains of the central Appalachian Ridge and Valley province (Harrison, 2004). However, interpretation of seismic reflection data from across the structure suggests that the synclinorium formed primarily by the removal of salt (Harrison, 2004).

The Catawissa-McCauley Mountain synclinorium is represented by the synclinal axis in the extreme lower right corner of Figure 2.5-23, just south of Nescopeck Mountain. The axes of this synclinorium follow the regional trends to the southwest-northeast, parallel with the axis of the Berwick Anticlinorium. This syncline has been delineated from numerous outcrop searches (Inners, 1978), geophysical data and a limited number of deep boreholes that penetrate the crust (Pohn, 2001). This fold is the northern-most part of the Anthracite Region located to the south of the site. The Anthracite region, located within a northeast plunging syncline, is the most faulted area of the Appalachian Ridge and Valley Province (Hornberger, 2004) (Figure 2.5-27b). The dominant faults consist of thrust faults, as part of the base décollement. The fault system has not been active since the post Carboniferous period and is located 20 mi (32 km) to the south of the site (Berg, 1980).

These folded features were formed at the same time as the Berwick Anticline (approximately 250 Ma) and show limited faulting, and there is no historical seismicity that is spatially related to these features (EPRI, 1986) (USGS, 2008). Because of the age of the deformations and the lack of current seismicity related to them, these features are not considered to be capable tectonic sources and pose no threat to BBNPP.

#### **2.5.1.2.4.4 Site Investigations**

The investigations at the site include a subsurface investigation for the foundation study, a geophysical investigation, and a field reconnaissance to inspect surface features for evidence of deformation, including folding, faulting and liquefaction. A subsurface investigation was completed at the SSES site, adjacent to the BBNPP site to the east, as part of the FSAR process for that site. The conclusions of that investigation at the SSES location are relevant to the current study by reason of proximity and intended purpose.

##### **2.5.1.2.4.4.1 Subsurface Investigations at the BBNPP Site**

Geologic sections developed from geotechnical data collected from 45 boreholes as part of the BBNPP study (as discussed in Section 2.5.4) provide detailed information in the upper 400 ft (122 m) of strata for the presence of structures directly beneath the site. The investigation at the BBNPP site indicates that the site is underlain by unfaulted Middle Devonian shale dipping 15 to 85 degrees, and covered by a layer of undeformed glacial outwash and till (Figures 2.5-96, 2.5-97, 2.5-98, 2.5-99, 2.5-100, and 2.5-100a).

The top of the bedrock shale beneath the site is interpreted to be an erosional surface left by the glaciers, upon which the younger sediments were deposited. Figure 2.5-33 shows that the younger sediments form a blanket over the bedrock surface that mimics the shape of the underlying bedrock surface. Cross sections prepared oblique to the previously mapped northeast-trending Light Street and Berwick faults (Figures 2.5-28, 2.5-29, and 2.5-33) show a slightly rolling bedrock surface below the BBNPP site, without offsets that would be characteristic of a faulted area.

Although the bedrock formations underlying the BBNPP site are dipping and have experienced folding during the Alleghanian Orogeny (Williams, 1987; Faill, 1999), surficial sediments at the

site display no signs of faulting or folding during the Pleistocene to Holocene time period, and rest unconformably on the eroded surface of the tilted beds of the local shale bedrock of the Mahantango Formation.

#### **2.5.1.2.4.4.2 Geophysical Investigations**

Seismic refraction surveys were performed to support site characterization studies for the BBNPP (Section 2.5.4.2.2.2.3). Because earth materials exhibit characteristic wave propagation velocities, they can be classified simply in terms of their seismic velocity. Seismic refraction data were interpreted for this study to assist in characterizing the local subsurface geologic materials regarding depths to glacial till, to weathered or fractured bedrock, and to competent bedrock. Seismic refraction surveys were operated along 6 profile lines totaling 4,000 linear feet (1,219 m) of coverage. The data for the surveys were collected from January 7 through 10, 2008 using approved quality assurance procedures. The complete report of this survey (Weston, 2008) is included in COLA Part 11G.

Figure 2.5-103 is a map depicting the layout of the 6 lines used during the survey (along Lines 1 through 3, oriented north-south; and Lines A through C, oriented east-west). Seismic P-wave velocity profiles, as interpreted by the SeisOpt @2D™ software are presented on Figure 2.5-104. These profiles are plotted without vertical exaggeration, with the vertical scale measuring elevation in feet, msl. These interpreted velocity profiles indicate a generally flat-lying eroded bedrock surface overlain by a variably thin veneer of overburden material. Figure 2.5-105 is a representation of the surface of the bedrock, as indicated by an interpreted velocity of at least 14,000 fps (4,267 mps). The velocity model developed for the site depicts the bedrock surface to be nearly flat lying from west to east and dipping to the south.

The seismic profiles constructed from manual calculations are presented in Figures 2.5-105a through 2.5-105f. These differ from the software-derived profiles, in that they have fewer layers, depict the depths of the units more accurately, and do not show the lateral changes in velocities that may be an artifact of the processing by the SeisOpt @2D™ software. The manual profiles also show the tops of the local bedrock as determined from borings installed during the site subsurface investigation. As is evident on the Figures, these tops compare favorably with the tops of the bedrock surface determined by seismic methods. The map of the top of the Mahantango Shale from the boring log data (Figure 2.5-105g) shows the same east-west strike and southward dip as the surface shown in Figure 2.5-105 that was developed from the geophysical data.

A review of the data in the Figures shows no offsets or abrupt changes in the tops of the bedrock that could indicate recent deformation. The close correlation with the boring log data further confirms the interpretation that there is no evidence of faulting in the Quaternary deposits underlying the site, in that no evidence of faulting was found in the core samples.

The somewhat irregular surface of the bedrock noted in the Refraction Survey and the boring logs is not interpreted as significant offsets of the top of the Mahantango Formation that could be attributed to faulting. Instead, the irregular bedrock surface is mimicked by the surface of the overlying glacial till layer, suggesting a glacially eroded bedrock surface. Most likely, the apparent irregular surface is the result of glacial scour, an erosional feature, within the upper part of the Mahantango Shale.

#### **2.5.1.2.4.4.3 Interpretation of Aerial Photography and LIDAR Imagery**

Aerial reconnaissance within a 25 mi (40 km) radius of the site was conducted by various personnel using aerial photographs from numerous publications. Figure 2.5-105r is a sample of the aerial imagery used, and it contains selected way points from the field reconnaissance. LIDAR imagery of the BBNPP site vicinity was also acquired for review and interpretation. The central portion of the LIDAR image contains elevation data with a 2 ft (0.6 m) contour interval. For clarity, the remainder of the image is a shaded relief representation without contours. The aerial reconnaissance investigated geomorphology and targeted numerous previously mapped geologic features and potential seismic sources (e.g., Berwick fault, Light Street fault, and Berwick Anticlinorium).

As shown in the following section, field reconnaissance coupled with interpretation of remote imagery (review and inspection of features preserved on the images) shows that there are no geomorphic features indicative of potential Quaternary activity along trends of the postulated folds and faults interpreted by Inners and Williams (Inners, 1978; Williams, 1987). No features suggestive of tectonic deformation were observed in the Quaternary glaciofluvial deposits, and no potential liquefaction features were observed along the Susquehanna River.

#### **2.5.1.2.4.4.4 Field Reconnaissance**

Information developed from the literature and the imagery interpretation was supplemented by field reconnaissance within a 25 mi (40 km) radius of the site. These field-based studies were performed to verify, where possible, the existence of mapped bedrock faults in the BBNPP site area and to assess the presence or absence of geomorphic features suggestive of potential Quaternary fault activity along the mapped faults, or previously undetected faults. Features reviewed during the field reconnaissance and office-based analysis of aerial photography and LIDAR imagery were based on a compilation of existing regional geologic information in the vicinity of the BBNPP site. As shown on topographic section B-B' on figure 2.5-105s there is no topographic offset to indicate recent movement of either Light Street or Berwick Faults.

Field reconnaissance was conducted by geologists in teams of two or more. Field reconnaissance visits in 2007 and 2008 focused on exposed portions of the Mahantango Formation, other formation exposures along the faces of Lee and Nescopeck Mountains, and roads traversing the site vicinity. Key observations and discussion items were documented in field notebooks and photographs. Field locations were logged by hand on detailed topographic base maps and with hand-held Global Positioning System (GPS) receivers. There were no faults or other forms of deformation noted in the field. No surface expression of either the Berwick or Light Street faults was noted, consistent with the conclusions documented in the literature. Figures 2.5-105h and 2.5-105i (Waypoint 12 on Figure 2.5-105t) show that there is no offset in the Quaternary deposits along Syber Creek, where the trace of Light Street Fault crosses it. Photos of the shale bedrock on the site show the steeply dipping nature of the bedding (Figure 2.5-105n and Waypoint WF3 on Figure 2.5-105r). Outcrops in a nearby borrow area show an undeformed contact between the glacial overburden and the shale bedrock (Figures 2.5-105o, 2.5-105p, and 2.5-105q and Waypoint WF5 on Figure 2.5-105r).

A third reconnaissance was conducted during the fall of 2008, to investigate the occurrence of potential liquefaction features along the Susquehanna River. The field reconnaissance was carried out by a team of geologists and engineers from Paul C. Rizzo Associates, Inc, and John Sims & Associates from both the land and water approaches to the river banks. The investigation was conducted for the course of the river for a reach of 25 miles (40 km) upstream and downstream of the site (Figure 2.5-105t). Because of the prevalent bedrock exposures in

both the river banks and the river bottoms, they found few locations where liquefaction conditions were possible and no evidence that liquefaction had occurred. Figures 2.5-105j through 2.5-105m show the rocky nature of the riverbed and its banks and some of the typical exposures found during the investigation (for Wapoints WP1, WP10, WP20, and WP22 respectively).

#### **2.5.1.2.4.4.5 Prior Work Completed at the SSES Units 1 and 2**

A detailed subsurface exploration near the BBNPP site was performed earlier as part of the original SSES FSAR (SSES FSAR, 2003) for the SSES Units 1 and 2 foundation and supporting structures. That exploration covered an area that is located adjacent to the east side of the BBNPP site and addressed the same issues as the current BBNPP FSAR investigation. The high level of detail in that earlier study, and the close proximity to the BBNPP site make that study especially relevant to the current effort. The level of effort for that earlier FSAR study included drilling 250 geotechnical boreholes, collecting down-hole geophysical data, and acquiring seismic refraction data from across the site. Previous site investigations performed for the existing units are summarized in the SSES Final Safety Analysis Report (FSAR) (SSES FSAR, 2003). As cited in the SSES FSAR, these previous investigations provide the following results documenting the absence of Quaternary faults at and within the area of the BBNPP site:

- Interpretation of satellite photos and topographic maps. This interpretation revealed no evidence of surface rupture, surface warping, or offset of geomorphic features indicative of active faulting.
- Interviews with personnel from government agencies and private organizations. These interviews concluded that no known faults are present beneath the existing SSES Units 1 and 2 site area.
- Seismicity Analysis -This analysis showed that no microseismic activity has occurred in the site area; the site is located in a region that has experienced only infrequent minor earthquake activity approximately 35 mi (56 km) northeast of the BBNPP site, between Lackawanna and Wyoming Counties; the closest fault-related epicentral location (the Anthracite Zone) is greater than 25 mi (40 km) away. No earthquake within 50 mi (80 km) of the SSES site has been large enough to cause significant damage in the time the region has been populated, approximately 270 years.
- Approximately 250 exploratory boreholes were drilled at the SSES Units 1 and 2 site area. Borehole data have provided evidence for the lateral continuity of strata across the existing SSES site area (SSES FSAR, 2003). The inspection of soil samples has revealed no adverse effects indicative of geologically recent or active faulting.
- Field reconnaissance of many surface outcrops at the site and within the 25 mi (40 km) radius of the site, coupled with geophysical surveys, provided no evidence for faulting at the SSES site.
- At the time of the original studies for the SSES FSAR (SSES FSAR, 2003), published maps showing bedrock faults within a 5 mi (8 km) radius of the SSES site identified only the Light Street fault. The closest significant bedrock faults mapped prior to 1975 were faults located about 80 mi (129 km) southwest of the SSES site near Lewistown, PA (SSES FSAR, 2003).

#### **2.5.1.2.5 Site Area Geologic Hazard Evaluation**

This subsection identifies and discusses potential geologic hazards that could be present in the site area (within 5-mi (8-km) radius of the site). These potential geologic hazards include rock dissolution features (caves and karst features), active deformation zones, landslides, abandoned underground mine cavities, and volcanism.

There are no caves or recognized karst features in the site area, and none were discovered during the investigations for BBNPP and SSES. Small- to medium-scale dissolution features occur in the site area where carbonate bedrock formations occur. Formations containing carbonate beds in the area include the Wills Creek, Tonoloway, Keyser, Old Port, and Onondaga formations of Silurian and Lower Devonian strata. These formations are at least 1,600 ft (488 m) below ground surface (bgs) at the BBNPP site and have not been penetrated by any borings at the site. These formations crop out or are located closer to the ground surface approximately 3 to 8 miles (5 to 13 km) southwest of BBNPP. Water wells in the Berwick area (southwest of BBNPP) are screened in the limestone formations and obtain ground water from dissolution features located along joints, fractures, and bedding planes (Inners, 1978). For example, the public water supply wells located in Berwick are screened in the Keyser Formation at depths of 87 to 160 ft (27 to 49 m) bgs and yield up to 500 gallons per minute (gpm) (1893 liters per minute (lpm)). However, one well drilled to a depth of 515 feet (157 m) at the same location produced no water. This is an example supporting the general belief that fractures, dissolution features, and secondary permeability of the rock decreases with depth because the confining pressure within the rock increases with depth and causes the fractures to close. This is discussed in greater detail in Section 2.4.12.1.2.8. Because the carbonate formations are located at least 1,600 ft (488 m) bgs, the frequency and magnitude of fracturing and dissolution features should be minimal.

Inventories of caves and karst features in Pennsylvania show no caves or karst to be present in Luzerne or Columbia counties; all significant caves and karst features are located in southeast, central, and southwestern Pennsylvania. The nearest significant cave is Crystal Cave near Kutztown, Pennsylvania, which is located approximately 46 mi (74 km) from BBNPP. In summary, based on the absence of limestone dissolution features in the 0.6-mile (1 km) site radius and the depth of 1,600 ft (488 m) to limestone, karst is not a geologic hazard at the BBNPP site.

Major springs, defined as having flow rates of 100 gpm (379 lpm) or more, can be an indicator of carbonate formations and karst features. An inventory of major springs in the Ridge and Valley Province of Pennsylvania by the USGS (Saad, 1990) identified only one major spring in Luzerne or Columbia counties; this spring (Lu 8) discharges from the Mauch Chunk Formation, approximately 10 miles (16 km) southeast of BBNPP (Saad, 1990). The Mauch Chunk Formation consists of clastic sedimentary rocks ranging from claystone up to conglomeratic sandstones. There are no carbonate strata in this formation and this formation does not underlie the BBNPP site. Thus, there are no documented cases of springs in Luzerne or Columbia counties that infer the presence of karst or significant dissolution features in carbonate rocks.

The last major tectonic events that generated large-scale earthquakes, faults, and deformation along the eastern coast of the United States occurred in the Mesozoic Era (Triassic and Jurassic, approximately 250 Ma). Active deformation processes and seismic activity within the site vicinity have been minimal since the Mesozoic Era, except for the Charlevoix Seismic Zone in Canada and the Charleston Seismic Zone in South Carolina. Thus, active deformation activities are not a source of significant geologic hazard at BBNPP (see also Section 2.5.1.1.4).

Because there are steep slopes and high topographic relief present in this portion of the Ridge and Valley Physiographic Province, landslides and other mass movements (e.g., soil slumping) can occur. Approximately 7.5 miles (12 km) north-northeast of the BBNPP site, is the location of one of the largest landslides in Pennsylvania. Approximately 4,000 years ago (Ka), a rock block landslide on the south side of Schickshinny Mountain caused 20,260,000–27,450,000 yd<sup>3</sup> (15,490,000–20,987,000 m<sup>3</sup>) of rock to move 1,250 ft (381 m) onto the Susquehanna River floodplain and partially diverted the Susquehanna River (Inners, 1988). Varnes (Varnes, 1978) defines a rock block slide as “a translational slide in which the moving mass consists of a single unit that is not greatly deformed”. Another much smaller landslide located 2 miles (3.2 km) northeast of the first (9.5 miles (15.3 km) northeast of BBNPP), was witnessed in 1947 in which rainfall, that deposited 6 inches of rain within 2 hours, likely caused approximately 122,000 yd<sup>3</sup> (93,277 m<sup>3</sup>) to move downslope within a minute or two (Inners, 1988). Including the aforementioned landslides, thirteen rock block slides have been mapped between Nanticoke, PA and Schickshinny, PA (a distance of approximately 9 miles (14.5 km)) along the south side of Schickshinny Mountain, with a total volume of about 56,000,000 yd<sup>3</sup> (42,815,000 m<sup>3</sup>) (Inners, 1988). All of these landslides, with the exception of the 1947 landslide, are prehistoric, having a maximum age of approximately 11 Ka, and were the likely results of a combination of the dip slope of Schickshinny Mountain being ultimately underlain by a weak mudstone, a relatively low dipping angle of the rock beds on the slope (approximately 20°), and the undercutting of the sandstone-mudstone bedding planes by the Susquehanna River (Inners, 1988). Inners (Inners, 1988) suggests that even though porewater pressure as a result of high moisture conditions in the area was the most likely cause of many of these historic rock block landslides, the larger landslides probably required a longer ‘wet’ season and/or multiple year high-moisture conditions. All of these rockslides occurred at least 7.5 miles (12 km) upstream along the banks of the Susquehanna River. No landslides (historic or pre-historic) of this proportion have been recognized or mapped in the BBNPP site area.

Underground coal mining has occurred in the Pennsylvania anthracite fields since the early 1800s. Hundreds of miles of underground workings are located in four different anthracite basins. While underground mining is currently very limited, the abandoned workings still result in mine subsidence immediately over the mine workings. The nearest anthracite basin (the Northern Anthracite Basin) is located at least four miles (6 km) distance from BBNPP, and all coal-bearing formations within the site area have been eroded long ago, so coal mining and mine subsidence at this distance will have no potential effect on the construction or operations of the BBNPP.

Tectonic activities and volcanism have not occurred in the eastern United States since the Mesozoic Era. Volcanic activity during the past 2,000 years has only occurred in the Western United States. The area of greatest volcanic activity is associated with the Cascade Range in the states of Washington, Oregon and California. The last eruption was Mount St. Helens in 1980. This is over 2500 mi (4023 km) away from the BBNPP site. Therefore, volcanism and related hazards are not a geologic hazard at the BBNPP site or vicinity.

Based on the discussion above, there are no geologic hazards that represent a risk to the construction or operation of the BBNPP.

#### **2.5.1.2.6 Site Engineering Geology Evaluation**

This section addresses the engineering significance of geologic and geotechnical characteristics of features and materials within the site area (5 mi (8 km) radius) and site location (0.6 mi (1 km) radius).

#### **2.5.1.2.6.1 Engineering Soil Properties and Behavior of Foundation Materials**

The overburden soil, based on engineering properties defined in Section 2.5.4., is not an adequate foundation substrate for safety related structures or facilities. Specifically, kame terrace and outwash as depicted on Figures 2.5-24 and 2.5-25 (Inners 1978) is comprised of stratified sands and gravels which are liquefiable under seismic loading. As a result, the overburden soil will be removed in order to found safety related structures or facilities on bedrock. Sound bedrock of the Mahantango Formation is encountered within 12.5 and 62.0 ft (3.8 – 18.9 m) below ground surface (bgs) at the BBNPP site location (0.6 mi (1 km) radius). Selected Category 1 structures will be founded on sound rock and the rock surface will be improved with removal of weathered rock, cleaning, and localized concrete fill as necessary. Engineering soil properties, including index properties, static and dynamic strength, and compressibility are discussed in Section 2.5.4. Variability and distribution of properties for the foundation bearing soils will be evaluated and mapped as the excavation is completed. Settlement monitoring will be based on analyses performed for the final design.

#### **2.5.1.2.6.2 Unrelieved Residual Stresses in Bedrock**

Due to the fact that bedrock within the site area (5 mi (8 km) radius) and the site location (0.6 mi (1 km) radius) has been buried under a cover of glacial sediments since the Wisconsinan glacial event (17,000-22,000 my), it is expected that removal of the glacial overburden will cause unloading effects such as exfoliation jointing or opening of existing joints and subsequent slaking of the rock. The effects of isostatic rebound as a result of rapid erosion due to glacial melt and rapid reduction in normal stress due to glacial retreat is well established in the literature (Pelletier, 2004; Stern et al, 2005). In addition, geologic mapping of rock outcrops in the site vicinity (5 mi (8 km) radius) revealed very continuous, low angle exfoliation fractures in the Mahantango Formation within ten feet of the surface. The presence of these fractures is a typical reaction to stress relief, and is an indication that unloading the rock by excavating overburden will result in the propagation of some low angle fractures that may necessitate further excavation or rock anchoring in some areas.

#### **2.5.1.2.6.3 Zones of Alteration, Weathering, and Structural Weakness**

Below the top of competent bedrock as defined by weathering and rock mass rating (RMR) scores, no unusual weathering profiles have been encountered during the site investigation. However, calcareous zones, thin pyrite lenses and zones exhibiting increased fracture density at depth were observed in isolated areas. Despite these discreet zones the preponderance of the rock mass is very competent. The site area bedrock geologic map and site location geologic map (Figures 2.5-27 and 2.5.27b, respectively), compiled by Berg (1980) and King (1974) indicates that most of the site area (5 mi (8km) radius) surrounding, and including, the BBNPP site are underlain by Late Silurian, Devonian, and Lower Mississippian bedrock. More specifically, the site location (0.6 mi (1 km) radius) bedrock is Upper Devonian Mahantango Formation which serves as the foundation for SSES Units 1 and 2. (PPL, 2004). The Mahantango Formation is part of the Hamilton Group as shown on Figure 2.5-27. The most recent geologic influence on the site, not including sub aerial exposure and erosion, is the Wisconsinan glaciation that is discussed in Section 2.5.1.1.1. The relevance of this glacial event with respect alteration, weathering, and structural weakness is complex. Glacial scour removed weathered rock from the surface of the site area (5 mi (8 km) radius) and glacial deposits left behind after glacial retreat have altered the weathering rate at the bedrock interface. The result is a slightly to moderately weathered bedrock surface. Evidence of a very thin to absent weathering profile is provided in the BBNPP site boring logs, which exhibit an average rock quality designation (RQD) of over 90% in the first ten feet of coring from the top of rock. Based

on the performance of the Mahantango Formation as the SSE Units 1 and 2 foundation, it is evident that, with proper excavation and foundation preparation, this rock is a suitable foundation bearing material for Category 1 structures. As noted previously, any zones of alteration, weathering, or increased joint or fracture density will be mapped during excavation and evaluated for removal, cleaning, and treatment with dental concrete.

#### **2.5.1.2.6.4 Deformational Zones**

The Light Street fault (DCNR, 2007) and the Berwick Anticlinorium (Inners, 1978) have been mapped at or within the 5 miles (8 km) radius of the BBNPP site. The Berwick Anticlinorium is an east-northeast striking, gently northeast plunging anticlinal structure with an axial trace that trends directly through the site area (5 mi (8 km) radius) and site location (0.6 mi (1 km) radius). The Berwick Anticline is an asymmetric structure in the site area (5 mi (8 km) radius) with the north-northwest limb dipping steeply to the north-northwest and the southern limb dipping more gently to the south-southeast. The relevance of the Berwick Anticlinorium is that the dip of the limbs have the potential to provide sliding planes within an excavation. In addition, axial plane cleavage may have the potential to result in toppling failure from an excavation slope. Any excavations into bedrock or bedrock slopes will be mapped and monitored during excavation and backfill. Field mapping efforts did not successfully identify surface expression of the Berwick Fault or the Light Street Fault. Deformation including fracturing and folding was mapped in outcrop and is discussed in Section 2.5.3.2. In addition, a thorough literature search was conducted to identify previous studies that have identified any form of deformation in the derock or identified in the stratified glacial deposits.

In addition, two other zones of deformation have been noted within the site area (5 mi (8 km) radius). These zones of deformation include the Lackawanna Synclinorium (Inners, 1978), which is located approximately 4 miles (6.4 km) northwest of the site, and the inferred Berwick fault (Inners, 1978) (DCNR, 2007), which is inferred to lie approximately 3 miles (4.8 km) southwest of the site (see Section 2.5.1). Neither of these features are considered capable tectonic sources, as defined in Appendix A of Regulatory Guide 1.165 (NRC, 1997). Based on geologic mapping, interpretation of borehole data and interpretation of geophysical investigation, no deformation zones were encountered in the site investigation for BBNPP that have a detrimental effect the engineering properties of the rock mass. In addition, no deformational zones were encountered during the exploration or excavation for SSES Units 1 and 2. Detailed mapping of the excavation for the proposed excavation for the BBNPP Category 1 structure will document the presence, orientation, continuity, and classification of any deformational zone. Each deformational zone or structure will be evaluated as to its potential affect on the foundation.

#### **2.5.1.2.6.5 Prior Earthquake Effects**

Two earthquakes have occurred within 25 miles (40 km) of the site (to the south and southwest) as shown on Figure 2.5-87. There are no reported earthquakes in the BBNPP site area (5 mi (8 km) radius). Studies of the SSES Unit 1 and 2 excavation, available outcrops within the BBNPP site area, and extensive exposures along the road cuts of Nescopeck and Lee Mountain, have not indicated any evidence for earthquake activity that affected the Paleozoic bedrock or Quaternary surficial deposits within the site area. Also, there are no historic accounts of earthquakes in the site area (5 mi (8 km) radius).



#### **2.5.1.2.6.6 Effects of Human Activities**

Previous investigations performed for the SSES Unit 1 and 2, and for this study identified no activities that would adversely affect the site. However, stones for construction material in the form of crushed rock have been quarried in the site vicinity. Sand and gravel also have been quarried, mostly along the Susquehanna River bank primarily for highway construction. Even though no mineral resources of economic significance and no mine workings are noted at BBNPP site, there are deep anthracite coal mining operations located at 11 miles east of the BBNPP site near the town of Shickshinny. Additionally, oil and gas reserves have been explored near Harveys Lake, approximately 25 miles (40 km) north of the BBNPP site. There is limited withdrawal of groundwater or recharge through impoundments of water in the site area, and adverse effects related to pore pressure change from groundwater withdrawal or injection are not expected to occur at BBNPP site.

#### **2.5.1.2.6.7 Site Groundwater Conditions**

A detailed discussion of groundwater conditions is provided in Section 2.4.12. Groundwater in the site area is founded in both the Glacial Overburden bedrock formations. In the vicinity of the BBNPP site, the glacial overburden aquifer is the most capable aquifer for transmitting groundwater, and it is the source aquifer for many wells and springs in the county. Groundwater in the bedrock formations is present primarily in secondary openings, including fractures, joints, and bedding plane separations. As a result, the ability of the bedrock to store groundwater or yield to well is typically less than the overburden for formations. Only the Glacial Overburden aquifer at site would be affected by construction and operation of the BBNPP. However, active dewatering methods within the sites are expected to limit the adverse effects in site overburden aquifer (see section 2.4.12.5.1).

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## 2.5.2 Vibratory Ground Motion

The U.S. EPR FSAR includes the following COL Items for Section 2.5.2:

A COL applicant that references the U.S. EPR design certification will review and investigate site-specific details of the seismic, geophysical, geological, and geotechnical information to determine the safe shutdown earthquake (SSE) ground motion for the site and compare site-specific ground motion to the Certified Seismic Design Response Spectra (CSDRS) for the U.S. EPR.

This COL Item is addressed as follows:

This section provides a detailed description of the vibratory ground motion assessment that was carried out for the {BBNPP} site, resulting in the development of the {BBNPP} site Safe Shutdown Earthquake (SSE) ground motion response spectra. {Starting points for this site assessment are the United States Geological Service (USGS) documentation of the studies for the 2002 and 2008 National Seismic Hazard maps (USGS, 2002)(USGS, 2008), the EPRI-SOG probabilistic seismic hazard analysis (PSHA) methodology outlined in EPRI NP-4726-A 1988 (EPRI, 1988), and the Early Site Permit (ESP) Application for the Clinton Nuclear Power Plant site (EGC, 2006) submitted to the NRC on April 14, 2006 by Exelon Generation Company (EGC). The following is a review of the approaches outlined in NRC Regulatory Guides 1.165 and 1.208 for conducting the vibratory ground motion studies used for the BBNPP site.

Nuclear Regulatory Commission (NRC) Regulatory Guide 1.208, "A Performance-Based Approach to Define Site-Specific Earthquake Ground Motion," March, 2007, (NRC, 2007a) states in Section B, Discussion:

"The CEUS is considered to be that part of the United States east of the Rocky Mountain front or east of Longitude 105 West (Refs. 13, 14). A Probabilistic Seismic Hazard Analysis (PSHA) in the Central Eastern United States (CEUS) must account for credible alternative seismic sources through the use of a decision tree with appropriate weighting factors that are based on the most up-to-date information and relative confidence in alternative characterizations for each seismic source. Seismic sources identified and characterized by Lawrence Livermore National Laboratory (LLNL) (Refs. 13-15) and the Electric Power Research Institute (EPRI) (Ref. 16, 17) were used for CEUS studies in the past. In addition to the LLNL and EPRI resources, the United States Geological Survey maintains a large database of seismic sources for both the CEUS and the WUS. The characterization of specific seismic sources found in these databases may still represent the latest information available at the time that a PSHA is to be undertaken. However, if more up-to-date information is available, it should be incorporated."

Regulatory Guide 1.165 (NRC, 1997a) provides the framework for assessing the appropriate SSE ground motion levels for new power generating nuclear plants. Regulatory Guide 1.165 also notes that an acceptable starting point for the SSE assessment at sites in the Central and Eastern United States (CEUS) is the PSHA conducted by the Electric Power Research Institute (EPRI) for the Seismicity Owners' Group (SOG) in the 1980's. Regulatory Guide 1.165 further specifies that the adequacy of the EPRI-SOG hazard results must be evaluated in light of more recent data and evolving knowledge pertaining to seismic hazard evaluation in the CEUS.

Reference 16 of the NRC Regulatory Guide 1.208 is Electric Power Research Institute, "Probabilistic Seismic Hazard Evaluations at Nuclear Power Plant Sites in the Central and Eastern United States," NP-4726, All Volumes, 1989-1991. The title and number of the

referenced document are not in agreement. The title of EPRI-4726 is "Seismic Hazard Methodology for the Central and Eastern United States." No document could be found that had the title provided by the NRC. In lieu of the reference 16, Section 2.5.2 of this document has used concepts from and interpretations presented in EPRI NP-4726, "Seismic Hazard Methodology for the Central and Eastern United States," (EPRI 1986); EPRI-4726-A, "Seismic Hazard Methodology for the Central and Eastern United States," (EPRI 1988); and EPRI NP-6395-D-1989, (EPRI, 1989a).

As stated in Regulatory Guide 1.208, the PSHA should incorporate the detailed guideline from NUREG/CR-6372 "Recommendations for Probabilistic Seismic Hazard Analysis: Guidance on Uncertainty and Use of Experts" Vol. 1 and 2 (NRC, 1997b). However, RG-1.208 does not limit the procedure to conduct the PSHA to the approach described in EPRI-4926, "Seismic Hazard Methodology for the Central and Eastern United States" (EPRI, 1986). The USGS information is also included in Regulatory Guide 1.208 as a potential starting point. USGS information can be used not only to define seismic sources but also to implement the PSHA procedure. In addition, the PSHA results developed by the USGS (Frankel, 1995) are prescribed in several building codes such as Minimum Design Loads for Buildings and Other Structures (ASCE/SEI 7-05)(ASCE, 2005a), and the International Building Code. These building codes are widely accepted by the engineering community.

Frankel's smoothed seismicity approach was developed to be applied in the calculation of annual probabilities of exceedance as low as  $10E-05$ . In his original paper, Frankel (Frankel, 1995) shows that his smoothed seismicity methodology reproduces the hazards obtained at 30 nuclear power plants sites following the EPRI methodology. He also shows that at four sites, the PSHA results obtained by his and the EPRI methodologies are very similar down to hazards of  $10E-04$  to  $10E-05$ .

The USGS and the EPRI PSHA methodologies are essentially the same. Their most noticeable difference is in their approach to calculating the seismicity parameters. Even in this step, both methodologies rely mainly on the historical seismicity, including estimates of incompleteness, and using a Gaussian smoothing procedure. The USGS, as the consultant of the NRC to review the EPRI NP-4726 report expressed several concerns about the EPRI PSHA methodology in calculating the seismicity parameters. However, after discussions among EPRI, USGS, and NRC staffs, they concluded that both the USGS and the EPRI approaches for calculating the seismicity parameters of the source zones provided satisfactory hazard results. Neither approach is superior to the other in performing PSHA especially in the CEUS.

Accordingly, the evaluation of vibratory ground motion made for the BBNPP site addresses seismic hazard update requirements in Regulatory Guide 1.165 (NRC, 1997a) and meets the SSE requirements given in paragraph (d) of 10 CFR 100.23 (CFR, 2007). Following the recommendation of Regulatory Guide 1.165 (NRC, 1997a), the 1989 EPRI study, EPRI NP-6395-D (EPRI, 1989a) provides a basis to start seismic hazard calculations. A Probabilistic Seismic Hazard Analysis (PSHA) determines the annual frequency of exceedance as a function of minimum ground motion. This annual frequency results from the integration of hazard contributions of seismic sources characterized by spatial extent and location, magnitude, frequency recurrence, and propagating the ground motion from the sources to the site. These calculations incorporate parametric variability, including alternative models and parametric distributions, as well as consideration of statistical uncertainties.

The following subsections summarize the procedure followed and results from the vibratory ground motion studies that were carried out for the BBNPP Site.

1. As a starting step, the EPRI-SOG tectonic interpretations in EPRI NP-4726 (EPRI, 1986) were examined in light of more recent geological, seismological, and geophysical data under the guidance of NRC Regulatory Guide 1.208, (NRC, 2007a). Sections 2.5.2.1 through 2.5.2.3 document this review and update of the EPRI-SOG seismicity, seismic source, and ground motion models.
2. Section 2.5.2.4 develops PSHA parameters at the site assuming the very hard rock foundation conditions implied by currently accepted ground motion attenuation models.
3. Section 2.5.2.5 summarizes information about the seismic wave transmission characteristics of the BBNPP site with reference to more detailed discussion of all engineering aspects of the subsurface in Section 2.5.4.
4. Section 2.5.2.6 describes the development of the horizontal SSE ground motion for the BBNPP site.

The selected SSE ground motion is based on the risk-consistent/performance-based approach of Regulatory Guide 1.208 (NRC, 2007a), with reference to NUREG/CR-6728 (NRC, 2001), NUREG/CR-6769 (NRC, 2002), and ASCE/SEI 43-05 (ASCE, 2005b). Horizontal ground motion amplification factors are developed using site-specific data and estimates of near-surface soil and rock properties. These amplification factors are then used to scale the hard rock spectra to develop Uniform Hazard Spectra accounting for site-specific conditions using Approach 2B of NUREG/CR-6728 (NRC, 2001) and NUREG/CR-6769 (NRC, 2002). Horizontal SSE spectra are developed from these soil Uniform Hazard Response Spectra using the performance-based approach of ASCE/SEI 43-05 (ASCE, 2005b), as implemented in Regulatory Guide 1.208 (NRC, 2007a). The SSE motion is defined at the free ground surface of a hypothetical outcrop at the base of the nuclear island foundation. See Sections 2.5.4 and 2.5.2.5 for further discussion of the subsurface conditions. Section 2.5.2.6 also describes vertical SSE spectra developed by scaling the horizontal SSE by a frequency-dependent vertical-to-horizontal (V/H) factor.

The SSE spectra described herein are considered performance goal-based (risk-informed) site specific safe shutdown earthquake response spectra. As discussed below, the SSE spectra for the BBNPP Site have been developed following the graded performance-based, risk-consistent method described in ASCE/SEI Standard 43-05 (ASCE, 2005b). The method specifies the level of conservatism and rigor in the seismic design process such that the performance of structures, systems, and components (SSC) of the plant achieve a uniform seismic safety performance consistent with the NRC's safety goal policy statement.

The SSE spectra, and its specific location at a free ground surface, reflect the seismic hazard in terms of a PSHA and geologic characteristics of the site and represent the site-specific ground motion response spectra (GMRS) of Regulatory Guide 1.208 (NRC, 2007a). These spectra are expected to be modified as appropriate to develop ground motion for design considerations.

The SSE developed in this section meets the requirements of paragraph (d) of 10 CFR 100.23 (CFR, 2007).}



### **2.5.2.1 Seismicity**

The U.S. EPR FSAR includes the following COL Item in Section 2.5.2.1:

Seismicity is site specific and will be addressed by the COL applicant.

This COL Item is addressed as follows:

{Current probabilistic hazard methodologies consider that the activity in area seismic sources can be adequately represented by the Gutenberg-Richter (G-R) recurrence equation in terms of body wave magnitude,  $m_b$ . A quantitative derivation of the G-R parameters is based on historical seismicity, i.e., on catalogs of seismic events. The seismic hazard analysis conducted by EPRI as delineated in NP-6395-D 1989 (EPRI, 1989a) relied, in part, on an analysis of historical seismicity in the Central and Eastern United States (CEUS) to estimate seismicity parameters (rates of activity,  $a$ , and slope  $b$ -values of the Gutenberg-Richter equation) for individual seismic sources. The historical earthquake catalog used in the EPRI analysis was complete through 1984.

As recognized in NRC Regulatory Guide 1.208 (NRC, 2007a), the United States Geological Survey (USGS) maintains a large database of seismic sources for both the Central and Eastern United States, as part of their efforts to develop National Seismic-Hazard Maps (USGS, 1996)(USGS, 2002)(USGS, 2008). In Open-File Report 96-532, entitled "National Seismic-Hazard Maps: Documentation June 1996 (USGS, 1996) the USGS states that their CEUS catalog was primarily based on a catalog by Seeber (Seeber, 1991), who conducted a refinement of the EPRI 1986 catalog (EPRI, 1986).

Section 2.5.2.1.1 and Section 2.5.2.1.2 are added as a supplement to the U.S. EPR FSAR.}

#### **2.5.2.1.1 Regional Seismicity Catalog**

{Many seismic networks record earthquakes in the CEUS. A large effort is continuously made by the USGS to examine and combine available data on historical earthquakes and to develop a homogeneous earthquake catalog that contains all recorded earthquakes for the region. "Homogeneous" means that estimates of  $m_b$  for all earthquakes are consistent, duplicate earthquakes have been eliminated, non-earthquakes (e.g., mine blasts and sonic booms) have been eliminated, and significant events in the historical record have not been missed. Thus, the USGS catalog forms a strong basis on which to estimate seismicity parameters. The USGS catalog updated up to 2001 has been used because this is the latest year for which the  $m_b$  units were reported. The use of  $m_b$  is required in view that the Gutenberg-Richter equation that describes the seismicity in area sources is considered to be valid in  $m_b$  units. Table 2.5-1 lists the earthquakes in the USGS Earthquake Catalog for the CEUS with  $m_b$  values greater than or equal to 3.0.

#### **2.5.2.1.2 Updated Seismicity Data**

Regulatory Guide 1.165 (NRC, 1997a) specifies that earthquakes of a Modified Mercalli Intensity (MMI) greater than or equal to IV or of a magnitude greater than or equal to 3.0 should be listed for seismic sources, "any part of which is within a radius of 200 mi (322 km) of the site (the site region)." The USGS catalog and methodology for determining seismicity parameters consider precisely the minimum magnitude of  $m_b$  equal to 3.0.

The USGS updated catalogs are compiled by examining and combining events listed in several CEUS source catalogs (Mueller, 1997). In this effort, the USGS intent is to develop a catalog dominated by entries from the best-researched sources and they use this priority to choose the best location and magnitude from among multiple source catalogs for each earthquake. In addition, the secondary events have been filtered as explained in a recent USGS publication (USGS, 2008). Traditionally, most CEUS earthquake magnitudes are reported as a short-period surface-wave magnitude ( $m_{bLg}$ ) and the ground-motions used in the hazard analysis are predicted based on  $m_{bLg}$ . In most cases a preferred magnitude from a catalog was assumed by the USGS to be equivalent to  $m_{bLg}$ , calling it  $m_b$ .

The catalog of use is the USGS catalog updated to the year 2007. Recent applications, such as the Clinton ESP (EGC, 2008), conclude that the update in the seismicity from the 1984 EPRI-SOG study do not significantly affect the seismicity parameters, i.e. the slope of the G-R (Gutenberg-Richter) equation (b parameter) and seismic rate (recurrent rate) in their respective regions of study. The same conclusion is reached in relation to the geometry of the seismic sources. The only relevant updates that were identified are the maximum magnitude of the Wabash Valley area source and the introduction of the New Madrid characteristic cluster events. A cluster model is required to represent the events that occurred in the three-series cluster with large magnitude ( $>7.5M$ ). The seismic parameters related to the New Madrid events are not in agreement with the general G-R equation utilized in the area source hazard computation, and therefore need to be treated separately. The New Madrid events occurred in a cluster of three events. The event shown in the catalog is considered to be the main New Madrid event for the 1811-1812 cluster set. The other two events are considered as the foreshock/aftershock events and are filtered out from the catalog by USGS. This New Madrid event is treated as the New Madrid Characteristic Cluster events in the PSHA since it does not follow the G-R (Gutenberg-Richter) relationship for seismicity rate. Therefore, this 1811-1812 main event (shown in the catalog) is ignored when calculating the seismicity rate for the New Madrid area source.

The 2002 USGS catalog is used to update more current seismicity data required for the PSHA at the BBNPP Plant Site. This catalog is a significant update to the EPRI-SOG 1984 seismic catalog. Seismicity recorded up to the year 2007 has been accounted for by adding the only two events within the site region recorded in the period between 2002 and 2007. These two events have been added to the end of Table 2.5-1 and are used in the BBNPP PSHA.

The 2002-2007 update to the catalog does not have a significant effect on the b-parameter, seismic occurrence rate, and/or the entire PSHA study at the BBNPP site. The use of the updated earthquake catalog results in a marginal reduction of the seismic occurrence rates, when compared to the USGS 2002 catalog. The resulting ground motion levels are marginally lower. The USGS 2002 G-R parameters are selected for the BBNPP PSHA since the effect on the seismic hazard of the 2002-2007 update is a marginal reduction in the ground motion levels, which is deemed to be less conservative.

A sensitivity analysis was performed to understand the effects of the 2002-2007 seismicity in an area broader than the 200 mi (322 km) region. For this analysis, an earthquake catalog search for 2002-2007 was performed with the following criteria:

- Use of a 500 mi (805 km) circular area centered at the BBNPP site 16W (41.08(N), -76.16(W)) for the USGS earthquake database; selection of the USGS/NEIC (PDE) '1973 - Present' database for magnitudes from 3.0 to 10.0 between the year 2002 and 2007 (USGS web database: [http://neic.usgs.gov/neis/epic/epic\\_circ.html](http://neic.usgs.gov/neis/epic/epic_circ.html)) (USGS, 2008);

- Use of a rectangular boundary zone with limits from 33.88(N) to 48.27(N) (latitude) and -66.60(W) to -85.70(W) (longitude) for the Advanced National Seismic System (ANSS) earthquake database for magnitude 3.0 to 10.0 between years 2002 to 2007 (ANSS composite web database: <http://www.ncedc.org/anss/catalog-search.html>) (ANSS, 2008);
- The sensitivity study extended the 2001 catalog to the year 2007. The update methodology adopted the approach described in the USGS open file report 2008-1128 "The 2008 Update of the United States National Seismic Hazard Maps" that USGS uses for catalog updates to 2006 to reduce the additional epistemic uncertainty related to the magnitude conversion. Magnitude may be adopted from the reported values. In addition differences resulting from unit conversion are negligible for small magnitude earthquakes. If more than one magnitude is listed in the original source-catalog record for an earthquake, a preferred magnitude is selected. Only about 10 percent of the earthquakes do not have a listed  $m_{blg}$  or  $m_b$  (USGS 2008); in these cases, the preferred magnitude is assumed to be  $m_{blg}$ . For the same event, if the magnitude between databases differs, the larger magnitude is selected;
- Consider only those earthquakes within the circular 500 mi (805 km) area.

As with the 200 mi (322 km) case, this analysis shows that the 2002-2007 update does not have a significant effect on the PSHA study at the BBNPP site. The result is a marginal reduction of ground motion levels, between one and two percent, when compared to the USGS 2002 catalog. Therefore, the more conservative USGS 2002 G-R parameters are selected for use in the BBNPP PSHA.

Earthquakes outside the region that have significant contribution to the hazard at the BBNPP site have been properly accounted for. Those with significant contribution include the Charlevoix Zone, the New Madrid Fault Zone, and the Charleston Earthquake. These contribute to the low frequency ground motion levels.

Major sources of potential seismic activity such as the New Madrid Seismic Zone (NMSZ) and the Charleston Seismic Zone (CSZ) are located beyond 200 mi (322 km) from the site. However, based on new paleo-seismology data, updated characteristic earthquake models have been recently formulated for the NMSZ and the CSZ. A sensitivity analysis for the BBNPP Site using these updated models showed that characteristic earthquake events from both sources are significant contributors to low frequency ground motion at the site. The sensitivity analysis also showed that Charlevoix seismic zone (in Canada) is a contributor to the hazard at low frequencies. Therefore these three sources have been included in the PSHA study for the site.

A PSHA result showed that the New Madrid Seismic Zone (NMSZ) is a contributor to the hazard at the BBNPP site at low frequencies. As such, characterization of this seismic source was added to the PSHA input for BBNPP. After a literature review of the existing NMSZ models, the characteristic earthquake model as described in the Clinton ESP (EGC, 2006) was selected as the input.

The Clinton ESP (EGC, 2006) also documented paleoliquefaction evaluations where evidence of soil liquefaction that occurred in prehistoric times is inferred from features such as sand boils or blows, dikes, and sills. By estimating the date and geographical distribution of these features, it is possible to infer the magnitude of the earthquake that originated the features. Earlier

investigations of paleoliquefaction features in the southern Illinois basin and in parts of Indiana, Illinois, and Missouri have identified paleoliquefaction occurrences that could have been caused by Holocene and latest Pleistocene earthquakes with estimated moment magnitudes (M) of 6 to 7.8. Details about the paleoliquefaction reconnaissance carried out for the Clinton ESP Site seismic hazard evaluation are given in Section 2.1.4 and Attachment 1 of Appendix B of the Clinton ESP document (EGC, 2006). These details include a discussion of each of the identified features, pictures of the features, results of radiocarbon dating, and criteria for differentiating seismic versus non-seismic liquefaction features. The Clinton ESP was issued by the NRC in March 2007 (NRC, 2007c).

These paleoliquefaction studies have been utilized for developing improved representations of characteristic earthquakes in the New Madrid Fault System. It was also concluded, from these paleoliquefaction evaluations, that the range of maximum magnitude earthquakes assigned to a random background earthquake in the PSHA for the Clinton ESP Site must include events comparable to that estimated for the Springfield, IL earthquake which occurred approximately 22 mi (35 km) northeast of Springfield, IL and 30 mi (48 km) southwest of the Clinton ESP site, that is, M 6.2 to 6.8.5.

The update in the introduction of the New Madrid characteristic cluster events is performed. These updates are adapted to the BBNPP site from the Clinton ESP Application. The updated New Madrid Model is discussed in Section 2.5.2.2.3.1.

Another significant source of severe seismic events in the East Coast of the United States is the Charleston seismic source that is about 620 mi (998 km) from the BBNPP site. Despite this long distance, it was considered that this source could still have some significant contribution to the hazard at the BBNPP site, particularly at low frequency ground motion. Thus, the Charleston seismic source has been included in the present PSHA. Since publication of the EPRI seismic hazard analyses, paleoliquefaction investigations and other studies have impacted the characterization of the geometry,  $M_{\max}$ , and recurrence in the Charleston seismic source. Paleoliquefaction studies in the area of the 1886 Charleston earthquake date back to Cox and Talwani (Cox, 1983) who discovered evidence for earthquake induced liquefaction features preserved in the South Carolina Coastal Plain sediments. Following this discovery, USGS conducted intensive studies to identify the spatial extent of paleoliquefaction features. USGS studies led to the discovery of sand blows that predated 1886, hence providing a basis for estimating the recurrence interval of large earthquakes in the Charleston area (Obermeier, 1987). More recent studies and interpretations have led to the refinement of the Charleston source zone parameters (Johnston, 1996; Bakun, 2004; Marple, 2000; Talwani, 2001; USGS, 2002; USGS, 2008). For example, radiocarbon dating techniques in new studies account for the fluctuation of atmospheric C-14 over time while previous studies assumed that the amount of C-14 has remained constant (Talwani, 2001). Based on the new interpretations, alternative geometries have been used for this zone. Marple (Marple, 2000) proposed a postulated East Coast Fault System (ECFS) in the Coastal Plain of the Eastern US and argued that the southern segment of this fault system is probably the source of the 1886 Charleston earthquake. In their 2008 version of National Seismic Hazard Maps, USGS (USGS, 2008) extended the Charleston area source offshore to include the Helena Banks fault zone as a possible source. USGS (USGS, 2008) also define another (elongated) area source which encloses the Woodstock lineament. This area source envelops half of the southern segment of ECFS. These two area sources have equal weights. Bechtel has examined these new data and developed an Updated Charleston Seismic Source (UCSS) model. The UCSS model has been used in development of the FSAR for the Vogtle ESP (SNOC, 2008). The UCSS model as described in the Vogtle ESP

(SNOC, 2008) has been adopted here after review and comparison with other models of the Charleston seismic source.

The mean of  $M_{\max}$  distribution used in the UCSS model ( $M_w$  7.1) is very close to that of USGS ( $M_w$  7.2). The source geometry used by UCSS considers four source zones with different weights. Beside the area of strong shaking during 1886 Charleston earthquake, this source zone combination accounts for the liquefaction features that are distributed far from the epicentral area. It also includes the southern segment of ECFS (Marple, 2000) as a possible source of the 1886 Charleston earthquake with a low weight of 0.1.

The recurrence interval of the Charleston characteristic earthquake in the UCSS model is based on the work of Talwani (Talwani, 2001). While Talwani (Talwani, 2001) argues that only the 2000 year record of paleoliquefaction data is complete, the UCSS model uses a combination of 2000 year and 5000 year records, thereby considering the possibility that the paleoliquefaction features may have been preserved in the 5000 year data. The 5000 year record, however, has a lower weight (0.2) than the 2000 year data (0.8). Based on comparisons between the UCSS model and other Charleston characteristic earthquake models including the USGS (USGS, 2008) model, it is concluded that the UCSS model better addresses the epistemic uncertainty in source zone parameters (including recurrence times) and therefore is used here to characterize the Charleston seismic source. This model is discussed in section 2.5.2.2.3.2. The description of the UCSS model is based on the Vogtle ESP FSAR (SNOC, 2008).

As a result of the investigations performed, relevant updates in maximum magnitude and geometry have been performed for the New Madrid cluster events and the UCSS. These events are distant from the BBNPP site but they still contribute to the hazard at the low frequencies.

The Charlevoix seismic zone is the most seismically active region of Eastern North America. It is located about 60 mi (97 km) downstream from Quebec City. This seismic zone has been the location of numerous small to medium earthquakes as well as five  $M \geq 6.0$  earthquakes in the last 350 years (Mazzotti, 2005). Among the larger events, only the 1925 earthquake,  $M_s = 6.2$ , (Bent, 1992) has been recorded by seismographs. The Canadian Geological Survey (CGS) conducted two field surveys in 1970 and 1974 to define the extensions of the seismic zone. The results of these two surveys delineated an active seismic zone about 20 by 55 mi (32 by 86 km) along the Saint Lawrence River (Earthquakes Canada, 2008). Maximum magnitudes of  $M \geq 7.0$  have been assigned to this seismic zone (Mazzotti, 2005). All six EPRI teams have considered the Charlevoix seismic zone in their source zone models and assigned mean maximum magnitudes close to or larger than 7.0. This seismic zone is located beyond the 200 mile radius of the Bell Bend site, but a sensitivity analysis showed that it is a contributor to the seismic hazard at the Bell Bend site lower frequencies. Therefore, this seismic source was included in the final PSHA for the BBNPP site.}

### **2.5.2.2 Geologic and Tectonic Characteristics of Site and Region**

The U.S. EPR FSAR includes the following COL Item in Section 2.5.2.2:

Geologic and tectonic characteristics are site specific and will be addressed by the COL applicant.

This COL Item is addressed as follows:

{As described in Section 2.5.1, a comprehensive review of available geological, seismological, and geophysical data has been performed for the BBNPP Site region and adjoining areas. As discussed in Section 2.5.1.2.5.2, excavation mapping is required during construction and any noted deformational zones will be evaluated and NRC notified when excavations are open for inspection. The seismotectonic characteristics of the region constitute the basis for defining the seismic source zones that affect the BBNPP Site.

This section summarizes the geologic structure and activity that could potentially result in seismic-induced vibratory ground motion at the BBNPP Site. The summary addresses Regulatory Positions 1 and 2 within Regulatory Guide 1.165 (NRC, 1997a), which requires that investigation of seismic sources to be performed within a 200 mi (322 km) radius of the site. The following sections summarize the seismic source interpretations (EPRI, 1986) that lie at least partially within this radius, relevant post-EPRI seismic source characterization studies and updated interpretations based on new data. The evaluation identified no new information which involved a change to the catalog that could impact the outcome of the PSHA.

Three major updates on seismic sources and characteristic earthquake models include:

The East Coast Fault System (ECFS) represents a new postulated seismic source along the Atlantic Seaboard (Section 2.5.1.1.4.4.1.2). The southern segment of the ECFS has been proposed by Marple (Marple, 2000) as being the source of the 1886 Charleston earthquake;

The average recurrence interval for large magnitude earthquakes in the Charleston characteristic model has been updated to 550 years based on paleoliquefaction data. The Charleston seismic source geometry also has been updated to include the southern segment of the ECFS as a possible source of the 1886 earthquake;

Assessments of magnitude, location, and return periods of large characteristic earthquakes of the New Madrid Seismic Zone (NMSZ) have been updated.

Detailed discussions of the updated source models are presented in the sections 2.5.2.2.3.1 and 2.5.2.2.3.2.

Section 2.5.2.2.1 and Section 2.5.2.2.2 are added as a supplement to the U.S. EPR FSAR.

#### **2.5.2.2.1 Summary of EPRI Seismic Sources**

The evaluations of new information examined in previous ESPs (EGC, 2006)(NRC, 2005) concluded that the EPRI-SOG seismic sources remain appropriate for assessing seismic hazards in CEUS. Therefore, the seismic sources defined in the 1989 EPRI/SOG study (EPRI, 1989a) have been adopted for updating the BBNPP site PSHA. However, it is noted that updates and adjustments are required for the maximum magnitude distribution for the area

sources and that characteristic earthquake models must be used to properly account for more recent information on the seismic activity in the New Madrid and Charleston seismic zones.

In the 1986 EPRI study (EPRI, 1986), six independent Earth Science Teams (ESTs) evaluated geological, geophysical, and seismological data to develop seismic sources in the CEUS. These sources were used to model the occurrence of future earthquakes and evaluate earthquake hazards at nuclear power plant sites across the CEUS. The six ESTs involved in the EPRI project were: Bechtel Group, Dames & Moore, Law Engineering, Rondout Associates, Weston Geophysical Corporation, and Woodward-Clyde Consultants. Each team produced a report, included in EPRI NP-4726, (EPRI, 1986), that provides detailed descriptions of how they identified and defined seismic sources.

The EPRI/SOG ESTs also determined recurrence parameters and maximum magnitudes for each source in  $m_b$  or magnitude units, including their corresponding weights. These models were implemented into a probabilistic seismic hazard analysis (PSHA) reported in EPRI NP-6395-D (EPRI, 1989a). EPRI NP-6452-D (EPRI, 1989b) summarized the parameters used in the final PSHA calculations, and this reference is the primary source for defining the geometry of area seismic sources for the BBNPP PSHA presented herein. For the computation of hazard, some of the 1989 EPRI seismic source parameters were updated, as discussed below.

The following sections list the seismic source interpretations in the 1989 EPRI PSHA study (EPRI, 1989a), relevant post-EPRI seismic source characterization studies, and updated interpretations provided by the more recent data. The summary of seismic sources and parameters was developed from the 1989 EPRI project EPRI NP-6452-D (EPRI, 1989b). The listed area seismic sources are those that at least partially lie within the "site region," i.e. within the circle with a 200-mi (322-km) radius centered at the BBNPP Site. The list includes the code used by each team to designate each source, the name of the source, the assigned recurrence parameter  $b$  and the assigned maximum magnitude, and weights assigned to each value of the parameter  $b$  and of the maximum magnitude.

Figures 2.5-34 through Figure 2.5-39a present the geometry of the seismic sources selected to estimate the hazard at the BBNPP Site, including plots of earthquakes with  $m_b$  equal to or higher than 3.0 in the updated earthquake catalog, to illustrate the spatial relationships between seismicity and seismic sources. Earthquake epicenters in the updated earthquake catalog include events from the period between 1627 and 2007, as listed in Table 2.5-1. Following the 1989 EPRI study (EPRI, 1989a) and the 1996, 2002 and 2008 USGS studies (USGS, 1996; USGS, 2002; USGS, 2008), the recurrence parameters for area seismic sources were computed for each one-degree latitude and longitude cell that intersects any portion of a seismic source.

The PSHA conducted in the EPRI-SOG study employed three strong ground motion attenuation relationships developed by Boore and Atkinson (Boore, 1987) and McGuire and others (McGuire, 1988) combined with the response spectral shapes by Newmark and Hall (Newmark, 1982) which are based on Western North America earthquake records. More recently-developed ground motion attenuation models (EPRI, 2004) are supported by a better understanding of earthquake generation and indicate that significant differences in the crustal properties between western and eastern North America lead to significant differences in the frequency content of ground motions between the two regions. In addition, the more recent ground motion models include an improved assessment of variability about median estimates, and thus have been used for this evaluation.

#### **2.5.2.2.1.1 Sources Used for EPRI PSHA – Bechtel Group**

The seismic sources and recurrence parameters identified by the Bechtel EPRI/SOG EST (EPRI, 1989a) that are within 200 mi (322 km) of the BBNPP Site are listed in Table 2.5-3.

Figure 2.5-34 illustrates the locations and geometries of the Bechtel Group seismic sources contributing to 99% of the seismic hazard along with plots of earthquakes with  $m_b$  equal to or higher than 3.0 between 1627 and 2007 (within 500 miles (805 km) of the BBNPP site for 2002 – 2007). Figure 2.5-34a shows the same magnitude events within the Bechtel Group seismic source zones for a 50 mi (80 km) radius of the BBNPP site.

#### **2.5.2.2.1.2 Sources Used for EPRI PSHA – Dames & Moore**

The seismic sources and recurrence parameters identified by the Dames & Moore EPRI/SOG EST (EPRI, 1989a) that are within 200 mi (322 km) of the BBNPP Site are listed in Table 2.5-4.

Figure 2.5-35 illustrates the locations and geometries of the Dames and Moore seismic sources contributing to 99% of the seismic hazard along with plots of earthquakes with  $m_b$  equal to or higher than 3.0 between 1627 and 2007 (within 500 miles (805 km) of the BBNPP site for 2002 – 2007). Figure 2.5-35a shows the same magnitude events within the Dames & Moore seismic source zones for a 50 mi (80 km) radius of the BBNPP site.

#### **2.5.2.2.1.3 Sources Used for EPRI PSHA – Law Engineering**

The seismic sources and recurrence parameters identified by the Law Engineering EPRI/SOG EST (EPRI, 1989a) that are within 200 mi (322 km) of the BBNPP Site are listed in Table 2.5-5.

Figure 2.5-36 illustrates the locations and geometries of Law Engineering seismic sources contributing to 99% of the seismic hazard along with plots of earthquakes with  $m_b$  equal to or higher than 3.0 between 1627 and 2007 (within 500 miles (805 km) of the BBNPP site for 2002 – 2007). Figure 2.5-36a shows the same magnitude events within the Law Engineering seismic source zones for a 50 mi (80 km) radius of the BBNPP site.

#### **2.5.2.2.1.4 Sources Used for EPRI PSHA – Rondout Associates**

The seismic sources and recurrence parameters identified by the Rondout Associates EPRI/SOG EST (EPRI, 1989a) that are within 200 mi (322 km) of the BBNPP Site are listed in Table 2.5-6.

Figure 2.5-37 illustrates the locations and geometries of Rondout seismic sources contributing to 99% of the hazard along with plots of earthquakes with  $m_b$  equal to or higher than 3.0 between 1627 and 2007 (within 500 miles (805 km) of the BBNPP site for 2002 – 2007). Figure 2.5-37a shows the same magnitude events within the Rondout Associates seismic source zones for a 50 mi (80 km) radius of the BBNPP site.

#### **2.5.2.2.1.5 Sources Used for EPRI PSHA – Weston Geophysical Consultants**

The seismic sources and recurrence parameters identified by the Weston Geophysical EPRI/SOG EST (EPRI, 1989a) that are within 200 mi (322 km) of the BBNPP Site are listed in Table 2.5-7.



Figure 2.5-38 illustrates the locations and geometries of Weston seismic sources contributing to 99% of the hazard along with plots of earthquakes with  $m_b$  equal to or higher than 3.0 between 1627 and 2007 (within 500 miles (805 km) of the BBNPP site for 2002 – 2007). Figure 2.5-38a shows the same magnitude events within the Weston Geophysical seismic source zones for a 50 mi (80 km) radius of the BBNPP site.

#### **2.5.2.2.1.6 Sources Used for EPRI PSHA – Woodward-Clyde Consultants**

The seismic sources and recurrence parameters identified by the Woodward-Clyde Consultants EPRI/SOG EST (EPRI, 1989a) that are within 200 mi (322 km) of the BBNPP Site are listed in Table 2.5-8.

Figure 2.5-39 illustrates the locations and geometries of Woodward-Clyde seismic sources contributing to 99% of the hazard along with plots of earthquakes with  $m_b$  equal to or higher than 3.0 between 1627 and 2007 (within 500 miles (805 km) of the BBNPP site for 2002 – 2007). Figure 2.5-39a shows the same magnitude events within the Woodward-Clyde seismic source zones for a 50 mi (80 km) radius of the BBNPP site.

#### **2.5.2.2.2 Post-EPRI Seismic Source Characterization Studies**

Seismic hazard evaluations more recent than the EPRI/SOG study have identified new information that could affect the assessment of the seismic hazard at the BBNPP Site. Specifically, updated data and information can have an impact on:

- Characterization of the rate of earthquake occurrences;
- Estimates of the maximum magnitude for seismic sources;
- Updated earthquake ground motions for the CEUS.

Studies that have used new data and information are described with emphasis on the items relevant for the evaluation of the seismic hazard at the BBNPP Site. These descriptions are provided in Section 2.5.2.2.2.1 through Section 2.5.2.2.2.5.

##### **2.5.2.2.2.1 USGS Studies for the United States National Maps**

Between 1996 and 2008, the USGS produced updated seismic hazard maps for the United States based on updated seismological, geophysical, and geological information (USGS, 1996)(USGS, 2002)(USGS, 2008). Each map reflects changes to the source models used to construct the previous version of the national seismic hazard maps. Among the most significant modifications to the CEUS portion of the source models are changes in the recurrence, maximum magnitude ( $M_{max}$ ), and geometry of the Charleston and New Madrid sources. Unlike the EPRI models that incorporate many local sources, the USGS source model in the BBNPP site region (200-mi (322 km) radius) includes only three sources that are important to the site hazard: the Extended Margin background, Stable Craton background, and New Madrid. Except for the New Madrid zone, where earthquake recurrence is modeled as characteristic earthquakes, the hazard for the large background or "maximum magnitude" zones is largely based on historical seismicity and the variation of that seismicity.

Since 1996, the USGS considered the occurrence of large events in the New Madrid as a characteristic rupture model with a characteristic moment magnitude  $M$  of 8.0, similar to the estimated magnitudes of the largest events in 1811-12 (USGS, 1996). The geometry of the New

Madrid source was modeled as three S-shaped parallel faults encompassing the area of highest historic seismicity. The USGS study used an average recurrence time of 1000 years for the New Madrid characteristic earthquakes.

The 1996 USGS study (USGS, 1996) also recognized that several paleoearthquakes have been identified in the Wabash Valley area. This seismic activity was modeled as an area zone with a maximum magnitude of M 7.5. For background zones, values of the Gutenberg-Richter (G-R) parameter "a" were determined in the 1996 USGS study (USGS, 1996) by counting the number of  $m_b=3$  and larger events within the zone since 1927 and adjusting the rate to equal that since 1976. The area-normalized a-value was then disaggregated into a set of grid cells to calculate the hazard considering the smoothed historic seismicity. The G-R parameter "b" was assigned a value of 0.95, based on calculations for the entire CEUS (USGS, 1996).

Some changes in the 2002 USGS study (USGS, 2002) that most affected the hazard estimates for the BBNPP site were the use of an updated mean recurrence time, characteristic magnitude, and spatial concentration to characterize the New Madrid sources of large earthquakes. A shorter mean recurrence time of 500 years was adopted and logic trees were developed for the characteristic magnitude related to the same configuration of three fictitious fault sources as in the 1996 maps, giving to the central fictitious source twice the weight of each of the faults to the sides. These changes markedly increased the probabilistic ground motions for the 10% probability of exceedance around the New Madrid area, compared to the 1996 results.

The documentation reported by the USGS for its 2008 (USGS, 2008) update of the national seismic hazard maps points out the following changes related to the Central and Eastern U.S.:

- Revise catalog and account for magnitude uncertainty

- Develop a logic tree for New Madrid (lower recurrence on northern arm and reduced magnitude)

- Implement a cluster model for New Madrid earthquakes

- Modify hypothetical fault geometry for New Madrid

- Develop a logic tree for  $M_{max}$  area sources

The USGS basic methodology for hazard estimates in the CEUS for the 2008 hazard maps is similar to that implemented in the 1996 and 2002 maps. Such methodology includes background-seismicity and fault source models (USGS, 1996)(USGS, 2002)(USGS, 2008). Background sources account for random earthquakes that occur off known faults and moderate size earthquakes that occur on modeled faults. The USGS, 2008 background source model (USGS, 2008) is composed of three smoothed gridded seismicity models, a large regional zone model, and local special seismicity-based zones. The gridded seismicity models are based on recorded historical earthquakes and account for the observation that larger earthquakes occur in regions that have experienced previous smaller earthquakes. Large regional zones account for low potential of random seismicity in areas without historical seismicity and establish a floor to the seismic hazard calculations. The special local zones allow for local variability in the G-R seismicity parameters. Fault models account for earthquakes on mapped active faults that have paleoseismic or historical evidence of repeated large earthquakes. One of the four CEUS fault models considered in 2008 by the USGS is the New Madrid Fault System (NMFS).

The USGS gridded seismicity, large regional zone and the local seismicity models require a declustered earthquake catalog for calculation of earthquake rates. The USGS develops these gridded seismicity rates from their seismic catalog for the Central and Eastern United States. The truncated Gutenberg-Richter (Gutenberg, 1944) magnitude-frequency distribution is used to model rates for different sizes of earthquakes in each grid cell or source zone. The USGS estimates completeness levels from the earthquake catalog, and calculates Gutenberg-Richter (G-R) parameters of the magnitude-rate relationship (intercept  $a$  and slope  $b$ ) using a maximum-likelihood method (Weichert, 1980) that accounts for variable completeness. The rates in the gridded cells are spatially smoothed using a two-dimensional Gaussian smoothing operator.

In 2008, the USGS (USGS, 2008) has used five fictitious parallel fault traces, each one having three arms. This is meant to represent the aleatory uncertainty in the locations of future large magnitude earthquakes in New Madrid, in a way similar to the three traces used in their 2002 model. The center of the five traces most closely follows the seismicity pattern and is assigned a weight of 0.7; the traces just outside of the central traces are weighted 0.1 each, and the outer traces are weighted 0.05.

USGS studies have also continuously incorporated developments in ground motion modeling (attenuation equations). In 1996 (USGS, 1996), the USGS adopted attenuation relationships derived for "hard rock conditions" recognizing that most attenuation relations for the CEUS published at that time were based on those site conditions. The USGS noted that it was less problematic to convert these to a firm-rock condition instead of converting them to soil conditions, since there would be less concern over possible non-linearity for the firm-rock site compared to the soil site.

The USGS 2008 study (USGS, 2008) includes several new simulation-based attenuation relations that were not available in 2002. While in 1996 and 2002 the USGS used ground motion models based on a single corner model (USGS, 1996; USGS, 2002), a double corner and hybrid models were incorporated in the 2008 study (USGS, 2008). The following is a list of the eight attenuation relationships used by the USGS in 2008, along with their assigned weights:

Single corner - finite fault

Toro and others (Toro, 1997), weight 0.2

Silva and others (Silva, 2002), constant stress drop with saturation, weight 0.1

Single corner - point source with Moho bounce

USGS (USGS, 1996), weight 0.1

Dynamic corner frequency

Atkinson and Boore (Atkinson, 2006), 140 bar stress drop, weight 0.1

Atkinson and Boore (Atkinson, 2006), 200 bar stress drop, weight 0.1

Full waveform simulation

Somerville and others (Somerville, 2001), for large earthquakes, weight 0.2

Hybrid empirical

Campbell (Campbell, 2003), weight 0.1

Tavakoli and Pezeshk (Tavakoli, 2005), weight 0.1

The 2002 and 2008 USGS efforts (USGS, 2002; USGS, 2008) have produced ground motion maps for a return period of 2475 years for building code applications.

#### **2.5.2.2.2 The Lancaster Seismic Zone**

The Lancaster seismic zone (LSZ) is located in southeastern Pennsylvania and is known as a post-EPRI study seismic zone located about 55 mi (89 km) south of the BBNPP Site (Section 2.5.1.1.4.4.2.10). The largest known earthquake of the LSZ is the January 16, 1994 Cacoosing Valley earthquake of  $m_{bLg} = 4.6$  near Reading Pennsylvania (Seeber, 1998). This event was located about 52 mi (84 km) south of the BBNPP Site. The Cacoosing Valley event has been attributed to unloading during a quarry process (Seeber, 1998) but it has not been removed from the standard earthquake catalogs used in PSHA studies. The LSZ is not included in the original EPRI source zone model (EPRI, 1986) as a separate source zone. However, the range of  $M_{max}$  values assigned to other EPRI source zones, adequately characterizes the LSZ in terms of the upper bound magnitude (Section 2.5.1.1.4.4.2.10). Therefore no update is required for the EPRI (EPRI, 1986) seismic source zone model within the region of BBNPP Site.}

#### **2.5.2.2.3 Updated Interpretations Based on New Data**

##### **2.5.2.2.3.1 Updated New Madrid Model**

As previously noted, seismologic, geologic, and geophysical studies have associated faults within the New Madrid region with the large-magnitude historical earthquakes that occurred during 1811 and 1812. In particular, paleoliquefaction studies indicate that large-magnitude events have occurred on these faults more frequently than the seismicity rates specified in the EPRI/SOG source characterizations. Thus, the updated seismic source evaluations focus on the characteristic large-magnitude events along the New Madrid Fault System.

##### Fault Geometry

As reported by the USGS and illustrated on Figure 2.5-74, very significant seismic activity occurs in the area of the New Madrid Fault System (NMFS). The severe 1811 through 1812 earthquakes are thought to have ruptured the Reelfoot Fault and fault segments to the south and the north. The precise locations of these three large events are not entirely known. The only evidence of surface rupture appears along the Reelfoot Fault and earthquake locations are generally constrained only by intensity and paleoseismic data. However, the available information indicates that the seismic activity at the NMFS can be attributed to the following three sources:

New Madrid South (NS) Fault;

New Madrid North (NN) Fault; and

Reelfoot Fault (RF).

The USGS studies for updating the 1996 U.S. national seismic map (USGS, 1996) considered this seismic activity as a "characteristic" rupture model. The USGS 1996 study (USGS, 1996) included a moment magnitude  $M$  of 8.0 and a recurrence time of 1000 years for such an event. Later, in the USGS work for the 2002 update of the national seismic hazard map (USGS, 2002), significant changes were introduced in mean recurrence time, characteristic magnitude, and spatial concentration of New Madrid sources of large earthquakes. It was recognized that the locations of these three large events are generally constrained only by intensity (felt) and paleoseismic data and a logic-tree approach was introduced to represent optional interpretations of fault locations and magnitudes of the New Madrid characteristic events. This logic tree was meant to characterize the range of expert opinions on the magnitude of the largest events of the 1811-12 sequence. The 2002 USGS study (USGS, 2002) represented the NMFS as three hypothetical sources: one fault with trace matching the observed microearthquake activity and two adjacent sources situated near the borders of the Reelfoot Rift. Also, a shorter mean return time of 500 years for characteristic earthquakes was considered by the USGS in the development of the 2002 maps (USGS, 2002). The end result was that the probabilistic ground motions for the 10% probability of exceedance level increased markedly around the New Madrid area, compared to the 1996 maps.

It is important to note that in the 1996 and 2002 models, the USGS (USGS, 1996; USGS, 2002) employed a single large earthquake that affects all three of the hypothetical faults, since these source models assumed that all earthquakes were independent.

Very recently, for the 2008 update of the hazard maps (USGS, 2008), the USGS takes into account the uncertainty in the locations of previous earthquakes by using five fictitious parallel fault-traces, similar to those used in the 2002 model (USGS, 2002). The central trace is weighted 0.7, the traces just outside of the central traces are weighted 0.1 each, and the outer traces are weighted 0.05 each. The USGS summarized expert opinions on the magnitudes of the 1811-1812 events, which shows that the estimated magnitudes range from  $M$  7.0 up to  $M$  8.1. Of the three largest New Madrid earthquakes, the one in January 1812 is the most likely to have ruptured the northern arm of the seismic zone (Figure 2.5-74). The three leading sets of magnitude estimates for the New Madrid sequence suggest that the January earthquake was  $0.2 \pm 0.1$  magnitude units smaller than the December shock (Johnston, 1996)(Hough, 2000)(Bakun, 2004).

Based on the updated information, for developing the 2008 maps, the USGS has assigned magnitudes for the northern section of the NMFS that are 0.2 units lower than those assigned for the central and southern sections (USGS, 2008). For the northern arm model the USGS applies the following weighting:  $M$  7.1 (wt 0.15),  $M$  7.3 (wt 0.2),  $M$  7.5 (wt 0.5),  $M$  7.8 (wt 0.15). The central and southern segments remain characterized as in 2002, i.e.,  $M$  7.3 (wt 0.15),  $M$  7.5 (wt 0.2),  $M$  7.7 (wt 0.5),  $M$  8.0 (wt 0.15).

Regarding large earthquake recurrence for the NMFS, the USGS 2008 study (USGS, 2008) has used paleoliquefaction data indicating a 500 year recurrence. Three large earthquake sequences are recognized from cross-cutting relationships and radiometric dating of sandblows (liquefaction effects). The USGS refers to Tuttle and others (Tuttle, 2002) who have recognized that events about 900 A.D., 1450 A.D., and 1811-1812 A.D. have occurred. These dates agree with a 500-year mean recurrence. However, citing lack of certainty on whether or not the northern portion of the fault system ruptured in the 1450 A.D. sequence, the USGS consider the possibility of 750-year and 500-year recurrences, equally weighted, for the northern arm of New Madrid. The 500 year recurrence for the southern and central sections remained unchanged in view that Tuttle and others (Tuttle, 2002) published evidence that all three of the sequences affected those arms.

Another relevant modification made by the USGS in their 2008 New Madrid source modeling (USGS, 2008) is that in addition to an unclustered model, as used in the 1996 and 2002 studies, a clustered large earthquake model was included. A clustered model postulates that the 1811-1812 earthquakes involved a sequence of three large earthquakes. This hypothesis is supported by geologic data of Tuttle and others (Tuttle, 2002) showing evidence that pre-historic earthquakes on the NMFS typically occur in sequences of three large earthquakes similar to those observed in 1811-1812. The relevance of this consideration is that a particular site will have a larger probability of exceeding a ground motion level if it is affected by three dependent events rather than one independent event.

The USGS 2008 study assigns equal weight to a clustered model for the NMFS characteristic earthquake and to a 2002-type unclustered source model. In addition, a more extensive logic tree was used to represent the rates and location of seismic activity at the NMFS.

The recent ESP submitted by Exelon for the Clinton Site (EGC, 2006) also recognizes that seismologic, geologic, and geophysical studies have associated faults within the New Madrid region with the large-magnitude historical earthquakes that occurred during 1811 and 1812. The Clinton Site was included in the 1989 EPRI/SOG study; however, the Clinton ESP notes that paleoliquefaction studies indicate that large-magnitude events have occurred on the NMFS more frequently than the seismicity rates specified in the EPRI/SOG source characterizations. Thus, Exelon decided to update the seismic source evaluations for the Clinton Site focusing on the characteristic large-magnitude events along New Madrid. To this end, Exelon supported a vast paleoseismicity investigation to develop an improved model for the characteristic events at the NMFS. This investigation provides the most complete available representation of New Madrid characteristic events, particularly regarding the development of logic trees for representing various rupture scenarios and optional recurrence models. Details of the Clinton ESP characterization of the NMFS are presented in subsequent sections of this document. Due to its proximity to the Clinton Site, Exelon conducted comprehensive studies for characterizing the seismic activity in the NMSZ as presented in the 2006 Clinton ESP application (EGC, 2006). The Exelon efforts included a thorough review of the technical literature as well as paleoliquefaction studies to identify the fault source geometry and to estimate recurrence parameters in the NMSZ. It was recognized that paleoliquefaction studies indicate that clustered large-magnitude earthquakes have occurred in this zone which can be properly modeled as characteristic events. Recent work for characterizing seismic activity at the NMSZ has also been conducted by the USGS for the 2008 update of the National Seismic Hazard Maps (USGS, 2008). The USGS also considers a temporal clustering model for the large NMSZ earthquakes and has developed a logic tree for representing these events. A detailed review of the Exelon and the 2008 USGS characteristic earthquake models for the NMSZ was conducted. Estimates of the locations, potential magnitudes, and recurrence of the characteristic events are similar in both models, even though the Exelon model is appreciably more detailed. The main difference is that the USGS also considers an un-clustered model giving only 50% weight to their clustered model while the Exelon model incorporates only the clustered model with a 100% weight. Thus the Exelon approach is appreciably more conservative and it is considered that this level of conservatism is adequate for assessing the seismic hazard for critical facilities such as nuclear power plants. The Exelon model for the NMSZ has also been adopted by EPRI in the 2004 update of the seismic hazard for nuclear power plant sites in the CEUS.

Based on the Clinton ESP (EGC, 2006), the logic tree used to represent the uncertainty in the model for the NMFS characteristic events is shown on Figure 2.5-42. The first two levels of the logic tree take into account the uncertainty in the location and extent of the faults that can

rupture in an earthquake sequence, by considering alternative geometries for the NS, RF and NN Faults. The considered fault locations are displayed on Figure 2.5-74. Distances to the BBNPP Site for the various options are listed in Table 2.5-9.

For the New Madrid South fault arm, two alternatives are considered:

1. Blytheville arch/Bootheel lineament (BA/BL); weight 0.6, length 82 mi (132 km), and
2. Blytheville arch/Blytheville fault (BA/BFZ); weight 0.4, length 71 mi (115 km).

Two alternative total lengths are considered for the New Madrid North fault arm:

1. With a weight of 0.7, rupture of the NN 37-mi (60-km) segment, and
2. The 60 mi (97 km) length including the NN and NNE is given a 0.3 weight.

Two possible alternatives are considered for the Reelfoot arm:

1. A full length segment including the northwest part, with weight 0.7, and
2. A central segment, excluding the northwest part, with 0.3 weight.

#### New Madrid Characteristic Earthquake Magnitude

Table 2.5-10 contains expected moment magnitudes for characteristic earthquake ruptures for each fault within the New Madrid Fault System along with their corresponding weights. As considered in the Clinton ESP, the size of the next characteristic earthquake is assumed to vary randomly about the expected value following a uniform distribution over a range of  $\pm 0.25$  moment magnitude units, to represent the aleatory variability in the size of individual characteristic earthquakes.

For the Clinton ESP, constraints on recurrence of characteristic NMFS events were derived from paleoliquefaction and paleoseismic investigations of the Reelfoot fault scarp and associated fold. It was concluded that the NMFS has generated temporally clustered large earthquakes in AD  $900 \pm 100$ , AD  $1450 \pm 150$  years and in 1811 to 1812; the time between clustered events may be from 200 to 800 years, with an estimated average of 500 years. Thus, a quantitative assessment of the uncertainty in the dates for prehistoric New Madrid earthquakes was developed, using a Monte Carlo simulation of constraints on the possible dates for the prehistoric earthquakes. The time intervals between these simulated dates were then fit with poissonian and renewal recurrence models. Table 2.5-10 lists the discrete distribution for equivalent annual frequency for characteristic New Madrid earthquakes. In this table, for Model A, all ruptures are similar in size to the 1811 and 1812 earthquakes. In Model B, 1/3 of the sequences consider a smaller (lower magnitude) rupture of the New Madrid North fault and 1/3 of the sequences assume a smaller rupture of the New Madrid South fault. The difference in magnitude from the 1811 and 1812 ruptures was set to be no more than 1/2 magnitude unit, and no magnitude ruptures are considered to be less than M 7. Model A and Model B were assigned weights of 2/3 and 1/3, respectively.

#### New Madrid Characteristic Earthquake Recurrence

The recurrence estimates, based on the poissonian and renewal models, used to represent the occurrence of characteristic New Madrid earthquakes in the Clinton ESP have been used

herein, as well as their corresponding weights, as summarized in Table 2.5-10. Since the site is affected by three dependent events, the frequency of exceedance,  $v(z)$ , of a spectral value  $z$  from a characteristic earthquake sequence is where:

$$v(z)_{\text{characteristic}} = \lambda_{\text{rate of cluster}} (1 - (1 - P_1)(1 - P_2)(1 - P_3))$$

$v(z)_{\text{characteristic}}$  is the probability of exceeding ground motion  $z$ ,

$\lambda_{\text{rate of cluster}}$  is the equivalent mean annual rate of occurrence of the event cluster, and

$P_1$ ,  $P_2$ , and  $P_3$  are the probabilities of exceeding the ground motion level  $z$ , when an earthquake of specified magnitude and distance occurs.

The values and weights for the rate of cluster are included in Table 2.5-10.

#### New Madrid Characteristic Earthquake Ground Motion Assessment

Consistent with the hazard calculation for area sources, the contribution of the New Madrid characteristic events was conducted using the CEUS ground motions developed by EPRI (EPRI, 2004). Figure 2.5-75 shows the logic tree structure defined by EPRI to represent the uncertainty in the median ground motion equation and in the aleatory variability about the median. As noted in the previous sections for area sources, the EPRI 2004 Report defines four clusters of median ground motion models to represent the alternative modeling approaches. All four clusters have been used for assessing the hazard from the New Madrid characteristic earthquakes, as illustrated on Figure 2.5-75. The rift option was selected for the fourth cluster, instead of the non-rift option that is used for area sources.

The three branches of the second level of the logic tree on Figure 2.5-75 represent the epistemic uncertainty in the median attenuation relationship for each cluster. The branches incorporate a three-point discrete distribution with weights of 0.63, 0.185 and 0.185 for the median, the 5<sup>th</sup> and the 95<sup>th</sup> percentiles, respectively. The third branching level addresses the uncertainty in the model for the aleatory variability in ground motions about the median attenuation relationship. Models 1A and 1B, as well as their weights, are those proposed by Abrahamson (EPRI, 2006a)(EPRI, 2006b) to account for inter-event and intra-event variability for events with distances longer than 12.4 mi (20 km) (termed  $s_1$  by Abrahamson). The additional standard deviation,  $s_2$ , developed by Abrahamson to incorporate additional variability at short distances, is not applicable for the distances between the BBNPP Site and any of the arms of the New Madrid faults.

The EPRI 2004 ground motion attenuation relationships use either the closest distance to the rupture plane or closest distance to the surface projection of the rupture plane (Joyner-Boore distance). Thus, the EPRI 2004 document also presents adjustments for use when the hazard integration is conducted based on point-source distances. These adjustments were unnecessary in the hazard calculations due to the New Madrid characteristic events, since the specific closest or Joyner-Boore distance was calculated for each fault arm, for input to the EPRI ground motion models.

#### **2.5.2.2.3.2 Updated Charleston Seismic Source (UCSS) Model**

Results of several post-EPRI studies have demonstrated that the parameters of the Charleston seismic source need to be updated. These parameters include the geometry, the maximum magnitude and the recurrence of characteristic events. Recent models of the Charleston



characteristic earthquake are significantly different from the 1986 EPRI characterizations. The most recent and detailed study of the Charleston characteristic events has been conducted for the Vogtle ESP (SNOC, 2008) producing the so-called Updated Charleston Seismic Source (UCSS) Model. The present PSHA for the BBNPP has adopted the UCSS model that was developed for the Vogtle ESP (SNOC, 2008) and was also used in the seismic hazard studies that support the recent FSAR for the CCNPP Unit 3 (UniStar Nuclear, 2007). The following description of the UCSS model is based on Section 2.5.2 of the CCNPP Unit 3 FSAR. The information with the largest relevance for the BBNPP Site is the assessment of characteristic magnitude. The exact location is less important in view of the large distance, more than 500 mi (805 km), to the BBNPP Site. The selection of the UCSS model has been based on the review of current literature related to the geometry (Marple, 2000; USGS 2002; USGS 2008), maximum magnitude (Johnston, 1996; Bakun 2004; USGS 2008), and recurrence intervals (Obermeier, 1987; Talwani, 2001). Based on this literature review (including the UCSS model), it was concluded that the UCSS model better captures the epistemic uncertainty in recurrence intervals and source zone geometries. The mean of maximum magnitude distribution is very similar to other models. For this reason, the UCSS model was selected as the preferred model to characterize the characteristic earthquake for the Charleston seismic source.

#### UCSS Geometry

The UCSS model includes four mutually exclusive source zone geometries (A, B, B', and C; Figure 2.5-82). These geometries have been defined based on the current understanding of geologic and tectonic features and shaking intensity in the region affected by the 1886 Charleston earthquake; on the distribution of seismicity; and on the geographic distribution, age, and density of liquefaction features associated with both the 1886 and prehistoric earthquakes. These features indicate that most of the evidence related to the Charleston source is concentrated in the Charleston area and is not widely distributed throughout South Carolina.

#### Geometry A

Geometry A is a northeast-oriented area centered at the 1886 Charleston meizoseismal area (Figure 2.5-82). This geometry encompasses the 1886 earthquake MMI X isoseismal (Bollinger, 1977), most identified Charleston area tectonic features and inferred fault structures and the majority of reported 1886 liquefaction features. Geometry A excludes outlying liquefaction features, because liquefaction occurs as a result of strong ground shaking that may extend well beyond the aerial extent of the tectonic source.

Existing evidence indicates that the seismic source for the 1886 Charleston earthquake was located in a relatively restricted zone defined by Geometry A. This zone envelops the local tectonic features, the area of ongoing concentrated seismicity, the area of high density liquefaction features, and the meizoseismal area of the 1886 earthquake. These observations suggest that future earthquakes with magnitudes comparable to the 1886 Charleston earthquake will likely occur within the area of Geometry A. Thus, a weight of 0.7 has been assigned to Geometry A (Figure 2.5-83).

#### Geometries B, B', C

Geometries B, B', and C are defined to capture the possibility that future earthquakes may not be restricted to Geometry A. The distribution of liquefaction features along the entire coast of South Carolina suggests that the Charleston source could extend beyond Geometry A. Therefore, Geometries B and B' represent larger source zones, while Geometry C represents the southern segment of the hypothesized East Coast Fault System source zone. Geometry B'

is a subset of B and defines the onshore coastal area as a source thus restricting the earthquakes in such onshore regions.

#### Geometry B - Coastal and Offshore Zone

Geometry B is a coast-parallel source including Geometry A in its entirety and elongated to the northeast and southwest to capture more distant liquefaction features in coastal South Carolina. The source also extends to the southeast region to include the offshore Helena Banks fault zone. This geometry is assigned a weight of 0.1.

#### Geometry B' - Coastal Zone

Geometry B' is a coast-parallel source that also incorporates all of Geometry A, as well as the majority of reported paleoliquefaction features. However, it does not include the Helena Banks Fault Zone. A weight of 0.1 has been assigned to this geometry.

#### Geometry C - East Coast Fault System (ECFS South Segment)

Geometry C envelops the southern segment of the proposed East Coast Fault System (Marple, 2000) as a possible source for the 1886 Charleston earthquake. A weight of 0.1 has been assigned to geometry C.

#### UCSS Maximum Magnitude Return Period

Based on currently available data and interpretations regarding modern  $M_{\max}$  estimates (Table 2.5-18), the UCSS model modifies the USGS magnitude distribution (USGS, 2002) to include a total of five discrete magnitude values each separated by 0.2 M units (Figure 2.5-83). The UCSS  $M_{\max}$  distribution includes a discrete value of M 6.9 to represent the Bakun best estimate of the 1886 Charleston earthquake magnitude, as well as a lower value of M 6.7 to capture a low probability that the 1886 earthquake was smaller than the Bakun mean estimate of M 6.9 (Bakun, 2004).

The UCSS magnitudes and weights are as follows:

<u>M</u>	<u>Weight</u>
6.7	0.10
6.9	0.25
7.1	0.30
7.3	0.25
7.5	0.10

This, results in a weighted mean Maximum magnitude of M 7.1 for the UCSS. This is slightly lower than the mean magnitude of M 7.2 in the USGS model (USGS, 2002).

The UCSS model incorporates geologic data to characterize the return period of  $M_{\max}$  earthquakes. Identifying and dating paleoliquefaction data provides a basis for estimating the recurrence of large earthquakes. Recent estimates of  $M_{\max}$  recurrence intervals are significantly

shorter than estimates in the EPRI models. Details regarding the processing, aging, and completeness of Charleston paleoliquefaction data can be found in Talwani (Talwani, 2001) and the CCNPP Unit 3 FSAR (UniStar Nuclear, 2007).

Records along two different time intervals (2000 yr and 5000 yr) are used in the UCSS model. Return periods derived from recorded paleoliquefaction features assume that these features were produced by large  $M_{\max}$  events and that both the 2000-year and 5000-yr records are complete.

The UCSS model calculates two average recurrence intervals covering two different time intervals, which are used as two recurrence branches on the logic tree (Figure 2.5-83). The first average recurrence interval is based on four events that occurred in the past 2000 years. The second average recurrence interval is based on events that occurred within the last 5000 years. The 2000 and 5000 records have been assigned weights of 0.8 and 0.2, respectively.

#### **2.5.2.2.3.3 The Ramapo Fault**

Reactivation of the Ramapo Fault during the Quaternary period has not been demonstrated. Results of core analyses in six localities of Ramapo and other basin-border faults showed that the most recent slip was extensional at each locality. The extensional tectonic episode did not extend beyond the Mesozoic and there is no evidence of post-Jurassic displacement (Sykes, 2008).

There is an apparent discrepancy between the distribution of earthquakes in the Ramapo Seismic zone and the lack of displacement in the last 150 Ma in the localities where the cores have been taken (Sykes, 2008). As a possible explanation for this discrepancy, it is assumed that earthquakes may originate from other preexisting faults which may or may not strike similar to the Ramapo fault. The term "Ramapo Seismic Zone" (RSZ) is used for the seismically active 7.5 mi (12 km) wide eastern area of the Reading Prong (Sykes, 2008). Since post 1974 earthquakes have been located with higher accuracy, this 7.5 km (12 km) width cannot be attributed to location errors. Therefore, it is concluded that more than one fault must be involved in generating the earthquakes (Sykes, 2008). This instrumental data suggests that activity in the Manhattan prong cuts off abruptly along a nearly vertical, northwest-striking boundary that extends from Stamford, Connecticut, to Peekskill, New York. The Peekskill-Stamford boundary is considered a newly identified feature (Sykes, 2008). It was possible to identify it after accurate locations and depths of earthquakes, especially those east of the Hudson River, however, became available with the installation of a seismic station near the New York–Connecticut border in 1971.

It is not clear which faults are active. The seismicity data used and processed cannot be solely used to delineate a single fault or multiple faults. Geologic evidence of Holocene fault movements is also very hard to be found in the study area. It is not possible to conclude that the Ramapo fault is an active feature. Therefore, even though it cannot be ruled out as a possible source for some of the observed earthquakes, the cause of the earthquakes is still unknown. This is a conclusion that has been reached by most previous studies (Kafka, 1985).

Recent research (Sykes 2008) has extrapolated the Gutenberg-Richter (GR) recurrence law for the earthquakes to magnitudes 6 and 7 and obtained the repeat times of 670 and 3400 years respectively. This has been done based on the observed seismicity of the whole study area and not just the Ramapo fault. It is not clear which magnitude will probably represent the maximum magnitude for the area or for the Ramapo fault assuming that it is active. The largest observed

earthquake in the area is the 1884 (offshore) New York earthquake with  $m_{bLg} = 5.25$ . Based on the observed seismicity a maximum magnitude of 6 or slightly higher is adequate.

The EPRI seismic source model does not consider the Ramapo fault system as a separate seismic source. For the BBNPP PSHA the maximum magnitudes of the EPRI seismic source zones (that encompass the Ramapo seismic zone) range from 5.3 to 7.1. These values have been presented in different logic trees by six EPRI teams and adequately characterize the upper level seismicity.

Even though a larger source zone (currently used in the BBNPP PSHA) tends to diffuse the hazard, this would only have an effect if the RSZ was located closer to the site. Based on the available data and information the current PSHA results adequately reflect the hazard at the BBNPP from the Ramapo fault zone. The Ramapo fault is located 82-93 mi (132-150 km) away from the BBNPP. Some branches of the Ramapo fault system (not the Ramapo fault itself) may extend into southeastern Pennsylvania. Those branches have similar distances from the site. Even if we assume the Ramapo fault is an active source (with  $M_{max} \sim 6.0$ ), it cannot produce significant ground motion at the site because of the relatively large distance.

#### **2.5.2.2.3.4 The St. Lawrence Zone**

Recent research has been performed in the southeastern corner of Lake Ontario and the western Lake Ontario in a search for evidence of neotectonic faulting (Wallach, 2002). In the Rochester Basin, vertical separations of layers of unconsolidated sediments and the underlying Paleozoic bedrock had been recognized. It may be possible to interpret that the observed displacements of the units are due to the recent tectonic faulting (Wallach 2002). However, such interpretation remains uncertain. Finding Quaternary faulting in Eastern North America (ENA) will continue to be difficult. Evidently, more rigorous study and work is needed in the study area to favor or to refute these interpretations.

In addition, historical seismicity, including instrumentally recorded data, show significant seismic activity in 3 areas of the St. Lawrence Seismic Zone (LSZ): (1) the Lower St. Lawrence Seismic Zone, (2) the Charlevoix Seismic Zone, and (3) the area of the 1944 Cornwall-Messana earthquake. The study region in the Lake Ontario south of Cornwall has not experienced high seismicity. According to the Earthquakes Canada website the region of southern Great Lakes has a low to moderate level of seismicity compared to more active zones to the east in general and to Charlevoix in specific. A map of seismicity on the website shows even fewer events for the Lake Ontario (Earthquakes Canada, 2008). It may be argued that the seismicity history is short but other zones in the Saint Lawrence Seismic Zone show high levels of activities in the same short period of time. According to the same website, on average, 2 to 3 magnitude 2.5 and larger earthquakes have been recorded in the southern Great Lakes region in the past 30 years. It is postulated that (Wallach, 2002):

“Because earthquakes of  $M \sim 5.5-7.0$  have been spatially related to the Saint Lawrence fault zone northeast of study area, similar sized earthquakes might reasonably be expected beneath Lake Ontario”.

The current level of seismic activity beneath Lake Ontario does not suggest potential for large earthquakes such as  $M7.0$ . Earthquakes with such magnitudes have been attributed to the Charlevoix Seismic Zone which has a distinctive seismic activity in the area. There is not sufficient evidence that could support the westward extension of the seismic zone. The lower St. Lawrence Zone and the Charlevoix zone are properly accounted for by the BBNPP PSHA.

#### **2.5.2.2.3.5 The New England Seismic Source Zone and the Cape Ann Earthquake**

The area of seismic activity of the New England region in northeastern United States is called the New England seismic zone. The seismic zone has experienced small to moderate historical seismicity. The 1755, M 5.9, Cape Ann earthquake (Ebel, 2006) is an example of the moderate historical events in this seismic zone. There is no evidence that earthquakes with magnitudes M 7.0 or larger may have happened in this seismic zone. The estimated magnitude of the 1638 New Hampshire earthquake (Estimated M=6.5-7.0) (Ebel, 1996) is also uncertain. The Weston Observatory of Boston College has operated a seismic network to monitor the earthquake activity in the New England seismic zone and adjacent areas. The seismic network can provide more accurate earthquake locations that can be used to explore the possibility of correlation between the epicenters and known tectonic and geologic features. Moderate magnitude events (M=6.0) from this seismic zone (which is located beyond the 200 mile radius) will not result in significant ground motion levels at the BBNPP site.

The Cape Ann, Massachusetts earthquake of 1755 was felt over an extensive area along the East Coast around the New England Region. The most likely location of its epicenter is 25 mi (40 km) East North East of Cape Ann, Massachusetts (Ebel, 2006). The estimated moment magnitude is M 5.9 and it is believed that peak ground accelerations were as high as 0.08 to 0.12 g in some soil locations. Due to the large distance and moderate (estimated) magnitude for this event, it is concluded that such earthquakes would not significantly affect the seismic hazard at the BBNPP site.

Recent research (Ebel, 2006) aims at refining the earthquake location and providing estimates of ground motion levels at a few locations using different methods for the 1755 earthquake. Three different methods have been used to estimate ground motion from the 1755 earthquake (Ebel, 2006): (1) using attenuation equations for CEUS assuming a magnitude and distance for the earthquakes; (2) MMI to ground motion conversions; and (3) estimates from chimney and unreinforced masonry damage. Ground motion estimates from the three methods have been given for PGA and 0.3 sec spectral acceleration assuming 5% damping for soil and rock sites at several locations.

The most distant location (Ebel, 2006) analysis of ground motion using attenuation relationships is New Haven, CT with an estimated distance of 175 mi (282 km) from the epicenter. The most conservative estimates for PGA and SA (0.3 sec) among all possibilities are 0.01g and 0.014g respectively. The soil acceleration for the same hazard level is 0.21 g. The Bell Bend site is located at a larger distance from the proposed epicenter of the 1755 event. The PSHA of the BBNPP site results in a uniform hard rock PGA of 0.1 g for 1E-4 hazard level.

The Cape Ann earthquake and associated seismic area do not have a significant contribution to the seismic hazard at the BBNPP site. The Cape Ann earthquakes are included in the PSHA in the USGS 2002 catalog and as part of the Law Engineering EST EPRI source zones.

Another earthquake of interest is M 5.6, 1727 Newbury event. The estimated magnitude for the 1727 event is M 5.6 (Ebel, 2000), which is a moderate event. This event is distant from the BBNPP site and would not have a significant contribution to the hazard.

#### **2.5.2.2.3.6 Charlevoix**

As previously discussed, the Charlevoix seismic zone is the most seismically active region of Eastern North America. It is located about 60 mi (97 km) downstream from Quebec City. All six EPRI teams have considered the Charlevoix seismic zone in their source zone models and

assigned maximum magnitudes close to or larger than 7.0. This seismic zone is located beyond the 200 mile radius of the BBNPP site, but a sensitivity analysis showed that it is a contributor to the seismic hazard at Bell Bend site lower frequencies.

Section 2.5.2.4.2 provides a discussion on related to the Gutenberg-Richter b parameter update from the original EPRI-SOG work. The BBNPP evaluation considered that an update is required to the b value of this zone. Values of b for the Charlevoix seismic zone were computed using the USGS 2001 (USGS, 2002) catalog and source zone geometries of Dames and Moore and Woodward-Clyde and a value of 0.7 was obtained in both cases. This is consistent with the b value obtained by other EPRI EST teams. Therefore, the b value for the Charlevoix seismic zone for those two teams was changed to 0.7. A b value of 0.7 is slightly more conservative than 0.79. Due to the long distance of the Charlevoix seismic zone from the BBNPP site, the new b value equal to 0.7 does not have a significant impact on the PSHA results.

An event of interest near the Charlevoix Zone is the November 25, 1988 Saguenay, Quebec earthquake, which occurred 22 mi (35 km) south of Chicoutimi, Quebec and 47 mi (75 km) north (northwest) of the Charlevoix seismic zone (Earthquakes Canada, 2008). The earthquake had an estimated magnitude of 5.9 in both body wave and moment magnitude units ( $m_b$  5.9,  $M$  5.9). A depth of 18 mi (29 km) was estimated for the earthquake. This earthquake was located outside the Charlevoix seismic zone and according to Earthquakes Canada it occurred in a relatively aseismic region. The earthquake was felt over a wide area with a maximum epicentral intensity of VII-VIII. The maximum recorded accelerations on rock were 0.156 g on horizontal component (at 40 mi (64 km) distance) and 0.102 g on vertical component (at 27 mi (43 km) distance).

The epicenter of Saguenay earthquake (48.12 N, 71.18 W) was located about 540 mi (870 km) from the BBNPP site. According the intensity map of Earthquakes Canada website, this event was felt with intensity III (MMI scale) in Pennsylvania.

The Saguenay earthquake epicenter is outside all defined areas of Charlevoix seismic zone for all 6 ESTs. The maximum magnitude estimates for the zone that contain the 1988 earthquake may have been underestimated by the EPRI teams. Since there has been no significant seismicity in the epicentral area of 1988 prior to the Saguenay earthquakes, most ESTs have defined small source zones for the Charlevoix based on the well defined area of significant seismicity. Weston has defined an alternative zone to consider the probability that some events outside Charlevoix may indeed belong to this zone. Rondout Associates defined a larger zone for Charlevoix and the 1988 is outside but very close to the western edge of this zone. The followings are two modifications suggested for the EPRI source zone model:

1. Define a broader Charlevoix seismic zone to cover the epicentral area of the 1988 earthquake. It is not clear how far the source zone should be extended beyond the epicenter of 1988 event unless other data become available. This alternative might diffuse seismicity and result in lower hazard. However, a sensitivity analysis was performed.

2. Increase the maximum magnitude of the current source zones that contain the epicenter of the 1988 earthquake. A maximum magnitude with the following distribution is suggested:

<u>Mmax</u>	<u>Weight</u>
5.9	0.33
6.3	0.34
6.7	0.33

A sensitivity analysis was performed with updated maximum magnitudes.

For the BBNPP case, both sensitivity analyses were performed and no significant change in the PSHA results occurred (less than one percent reduction in the low frequency range of the spectra). This is attributed to the large distance (540 mi (870 km)) of the event from the site. Regardless of the sensitivity results, the b G-R value was updated as discussed in Section 2.5.2.4.2.}

### **2.5.2.3 Correlation of Earthquake Activity with Seismic Sources**

The U.S. EPR FSAR includes the following COL Item in Section 2.5.2.3:

Correlation of earthquake activity with seismic sources is site specific and will be addressed by the COL applicant, consistent with the guidance of RG 1.208 and RG 1.165, as appropriate.

This COL Item is addressed as follows:

{Following Regulatory Guide 1.165 (NRC, 1997a) and 10 CFR 100.23 (CFR, 2007), a PSHA was conducted to determine the SSE and to account for uncertainties in the seismological and geological evaluations for the BBNPP site. The probabilistic approach was based on the PSHA conducted by the EPRI for CEUS in the mid to late 1980s (EPRI, 1989a) with changes to incorporate updated data. Expert opinion was incorporated following a Senior Seismic Hazard Analysis Committee (SSHAC) approach (NRC, 1997b).

The location of earthquakes was accounted for by an updated USGS catalog (USGS, 2002), covering events between 1627 and 2007. The updated catalog has been adopted for assessing the BBNPP site seismic hazard. This update is a refinement of the EPRI SOG catalog that listed earthquakes between 1627 and 1984 (EPRI, 1988). Figure 2.5-34 through Figure 2.5-39a show the distribution of earthquake epicenters from both the EPRI (EPRI, 1986) and updated 2001 USGS (USGS, 2002) earthquake catalogs in comparison to the seismic sources identified by each of the EPRI ESTs. These figures include updates for seismic activity in the zones for 2002 – 2007 out to a radius of 500 mi (805 km) from the BBNPP site. The comparison of earthquake distributions from both earthquake catalogs supports the following conclusions:

- The updated catalog does not show any earthquakes within the site region that can be associated with a known geologic or tectonic structure.
- The updated catalog does not show a unique cluster of seismicity that would suggest a new seismic source outside of the EPRI seismic source model (EPRI, 1986).
- The updated catalog does not show a pattern of seismicity that would require significant revision to the EPRI seismic source geometry.

- Two events were added to the 2001 USGS catalog in the period of 2002-2007. This update does not impact the result of the PSHA.}

#### **2.5.2.4 Probabilistic Seismic Hazard Analysis and Controlling Earthquake**

The U.S. EPR FSAR includes the following COL Item in Section 2.5.2.4:

The probabilistic seismic hazard analysis is site specific and will be addressed by the COL applicant, consistent with the guidance of NUREG/CR-6372, RG 1.165 and RG 1.208, as appropriate.

This COL Item is addressed as follows:

{Sections 2.5.2.4.1 through 2.5.2.4.6 are added as a supplement to the U.S. EPR FSAR.

##### **2.5.2.4.1 1989 EPRI Probabilistic Seismic Hazard Analysis**

The seismic hazard for the BBNPP was calculated using the original EPRI EST teams area sources, plus the New Madrid and Charleston characteristic earthquakes, and with the updated ground motion model and aleatory uncertainty model. This calculation was first made for hard rock conditions, and these results were then modified to account for local site conditions.

The analysis of seismic hazard consists of calculating annual frequencies of exceeding different amplitudes of ground motion, for all combinations of seismic sources, seismicity parameters, maximum magnitudes, ground motion equations, and ground motion aleatory uncertainties. This calculation is made separately for the New Madrid zone, for the Charleston zone and for the seismic sources defined by each of the six EPRI EST teams and results in a family of seismic hazard curves. The alternative assumptions on seismic sources, seismicity parameters, maximum magnitudes, ground motion equations, and ground motion aleatory uncertainties are weighted, resulting in a combined weight associated with each hazard curve. From the family of hazard curves and their weights, the mean hazard (and the distribution of hazard) can be calculated.

The quantification of the Probabilistic Seismic Hazard at hard rock utilized Rizzo's in-house software, ProHazard. This code uses the definition of site area seismic sources, the seismic potential of these sources in terms of generating future earthquakes, and the ground motion models, to estimate the annual exceedance probabilities for various levels of spectral accelerations at different spectral frequencies.

The technical methodology utilized in ProHazard follows the approach implemented in the 1989 Electric Power Research Institute study for Nuclear Power Plant Sites in the Central and Eastern United States (EPRI, 1989). This methodology is generally based on the early work of Cornell (Cornell, 1968) (Cornell, 1971) and integrates the product of the conditional probability that a ground motion measure will be exceeded given the earthquake magnitude and distance, and the probability distribution of magnitude and distance over all sources that can significantly contribute to the site seismic ground motion. This is expressed as:

$$v(z) = \sum \alpha_n(m_o) \int f(m) \left[ \int [f(r|m)P(Z > z|m, r)dr]dm \right]$$



where:

$Z$  is the peak ground acceleration or the spectral pseudo-acceleration at prescribed natural frequencies,

$P(Z > z | m, r)$  is the conditional probability that  $Z$  will exceed a value  $z$ , given the earthquake magnitude,  $m$ , and distance,  $r$ ,

$f(m)$  and  $f(r)$  are the probability density functions for magnitude and distance, and

$\alpha_n(m_0)$  is the number of earthquakes per year above a prescribed minimum magnitude  $m_0$ , in the  $n$ -th seismic source.

The integration over magnitude is performed from  $m_0$  to an upper bound magnitude  $m_u$ , and the integration over distance is performed usually over a prescribed radius from the site, typically larger than 186 mi (300 km). The probability density function for distance assumes that earthquakes can occur randomly over the source areas or faults. The functions  $f(m)$  and  $\alpha(m_0)$  define the recurrence relationships for the respective source zones.

The conditional probability in the above equation represents the random uncertainties in the natural phenomenon (aleatory). Additionally, ProHazard addresses epistemic uncertainties in sources and recurrence parameters and the ground motion attenuation resulting from limitations in the available data and alternative interpretations of this data. Alternative assumptions on seismic sources, seismicity parameters, maximum magnitudes, ground motion equations, and ground motion aleatory uncertainties are weighted, resulting in a combined weight associated with each hazard curve. The mean hazard and the distribution of hazard (i.e., median and fractiles) are obtained from the resulting family of hazard curves and the associated weights.

The attenuation relationships developed in 2004 by EPRI (EPRI, 2004) for the CEUS have been implemented in ProHazard. This model was the outcome of several workshops that convened a panel of six ground motion experts who developed a consensus-based ground motion model consisting of weighting of several attenuation relationships. The ground motion model relates spectral accelerations at frequencies of 100 Hz (equivalent to peak ground acceleration (PGA)), 25 Hz, 10 Hz, 5 Hz, 2.5 Hz, 1 Hz, and 0.5 Hz for generic hard rock conditions to moment magnitude at and distance to a given source. Epistemic uncertainty is represented using multiple ground motion equations and multiple estimates of aleatory uncertainty (sigma), all with associated weights. Further, EPRI (EPRI, 2006) corrects the excessive aleatory uncertainties in the 2004 study, particularly for low frequencies.

The above uncertainties are implemented in the ProHazard analysis utilizing the logic tree formalism. The logic trees represent discrete alternatives of models and model parameters and assign relative weights to the alternatives. These weights were developed from statistical analysis of the data and represent the best judgment of experts. Thus, several analyses reflecting various scenarios quantitatively assess the modeling uncertainties.

ProHazard has been subjected to the verification and validation procedures stipulated in Rizzo's Quality Assurance Manual. Computer software control for ProHazard has been done according to Rizzo's Quality Assurance Manual. The Quality Assurance Manual addresses software activities including software acquisition and development, tracking installation of design and analysis software on individual computers, program verification and validation, in-use testing, software usage, change control, configuration control, error notice documentation and

distribution, software maintenance, virus protection, software retirement, records and monitoring.

The method used to verify and validate the capabilities of ProHazard are in accordance with methods accepted by the NRC and as described in EPRI NP-4726 (EPRI, 1986). Validation test problems are selected for testing specific combination of analysis capabilities of the ProHazard. The problems are modeled and analyzed using ProHazard. The analysis results obtained from ProHazard are then compared with the benchmark solutions from published technical literatures.

#### **2.5.2.4.2 Effects of New Regional Earthquake Catalog**

A sensitivity study was done in order to determine if the activity rates have changed. Seismicity rates in the EPRI study (EPRI, 1986) were based on an earthquake catalog that extended through 1984. The USGS 2001 catalog (USGS, 2002) has 17 more years of data and it was updated to include seismicity data up through the year 2007. Using the USGS 2001 catalog and completeness periods the  $b$  values for some of the EPRI source zones were computed and compared to the  $b$  values obtained by EPRI teams. The differences between the two sets of  $b$  values are small and can be attributed to using different catalogs and different completeness periods. For example, the EPRI  $b$  value for Rondout source zone 31 is 0.96 while the estimated  $b$  value using USGS 2001 catalog is 1.02. For many of the source zones the differences are less than 0.05. Therefore, the EPRI  $b$  values do not need any update with exception of the Charlevoix seismic zone. Except for the Dames and Moore and Woodward-Clyde teams, other EPRI EST teams have derived a  $b$  value between 0.70 and 0.79 for the Charlevoix seismic zone. The USGS used a  $b$  value of 0.76 (USGS, 1996) (USGS, 2002) (USGS, 2008) for this source zone based on the work of Adams and others in the Geological Survey of Canada. Values of  $b$  for the Charlevoix seismic zone were computed using the USGS 2001 (USGS, 2002) catalog and source zone geometries of Dames and Moore and Woodward-Clyde and a value of 0.7 was obtained in both cases. This is consistent with the  $b$  value obtained by other EPRI EST teams. Therefore, the  $b$  value for the Charlevoix seismic zone for those two teams was changed to 0.7. A  $b$  value of 0.7 is slightly more conservative than 0.79. However, considering the long distance of the Charlevoix seismic zone from the BBNPP site, the new  $b$  value equal to 0.7 does not have a significant impact on the PSHA results. No other changes in EPRI seismicity parameters are required.

#### **2.5.2.4.3 New Maximum Magnitude Information**

The upper magnitude,  $M_{\max}$ , utilized in the magnitude recurrence equation could significantly affect the low probability seismic hazard, in particular from the near field events. In the 1989 EPRI/SOG (EPRI, 1989a) study, each EST developed alternative values of  $M_{\max}$  for each seismic source in a body wave (mb) unit. More recent studies (USGS, 2008; Bakun, 2004), however, have revised  $M_{\max}$  for the Charleston, New Madrid, and local sources. In addition, it has been recognized that large historical events have occurred at the New Madrid and the Charleston fault systems that cannot be adequately modeled by the G-R equation. Instead, the concept of characteristic earthquakes (Schwartz, 1984; Youngs, 1985) has been introduced to more appropriately represent the seismic activity at New Madrid and Charleston. Thus, characteristic events have been adopted in the calculation of the hazard at the BBNPP Site. Moment magnitudes  $M$  between 7 and 8.1 were considered for the New Madrid source and between 6.5 and 6.7 for the Charleston source. Tables 2.5-3 through Table 2.5-8 list revised maximum magnitudes and their corresponding weights for the seismic sources selected for the BBNPP Site PSHA.

The EPRI/SOG ESTs defined the maximum magnitude for each of their seismic sources using either body wave magnitude,  $m_b$ , or seismic moment magnitude,  $M$ . Furthermore, the G-R parameters  $a$  and  $b$  are derived in terms of  $m_b$ , while the equations for ground motion models are functions of  $M$ . Therefore, conversions from body wave magnitude into moment magnitude are required. The three magnitude-conversion relationships shown in Table 2.5-2 were used in the BBNPP PSHA and the three of them were assigned equal weight.

#### **2.5.2.4.4 New Seismic Source Characterizations**

New characteristic earthquake New Madrid and Charleston source models have been adopted to reflect updated estimates of the possible geometries and maximum magnitude at both fault zones. The Gutenberg-Richter (G-R) equation (Gutenberg, 1944) has been used to describe recurrence in area seismic sources. This equation was truncated at the maximum magnitude,  $M_{max}$ . The  $a$  and  $b$  parameters characterizing the potential of area seismic sources have been updated as well as their maximum earthquake magnitude. As noted before, each EPRI EST (EPRI, 1989a) developed G-R parameters  $a$  and  $b$  for each of their seismic sources, identifying their selected smoothing options and their corresponding weights. Smoothing allows incorporation of the variation of the G-R parameters  $a$  and  $b$  within the seismic source. For the BBNPP PSHA, the smoothing approach developed by USGS (USGS, 2002) (USGS, 2008) has been used. This approach considers only the variation of the intercept parameter  $a$  for prescribed constant values of the slope parameter  $b$ . The constant values of  $b$  have been taken as the averages of the  $b$ -values adopted for each seismic source by each EPRI EST (EPRI, 1989a), along with the corresponding weights for each smoothing option. Tables 2.5-3 through Table 2.5-8 present the values of the average seismic parameter  $b$  used as input to the BBNPP PSHA.

Four smoothing options are considered for characterizing the recurrence parameter  $a$  in the USGS 2008 approach. Each of the first three smoothing options is based on an incompleteness period, a minimum incompleteness magnitude, and a smoothing correlation distance. The fourth option is considered only for the background seismic source since it has negligible effect on main sources such as New Madrid or Charleston that have a much smaller area than the background source. The information for each model is listed in Table 2.5-16.

#### **2.5.2.4.5 New Ground Motion Models**

Once the earthquake sources are defined, attenuation relations relate the source characteristics of the earthquake and propagation path of the seismic waves to the ground motion at a site. Predicted ground motions are typically quantified in terms of a median value (a function of magnitude, distance, site condition, and other factors) and a probability density function of peak horizontal ground acceleration or spectral accelerations.

The estimation of strong ground motion for specified magnitude, distance, and site conditions in the CEUS is difficult due to the paucity of physical data. Most of the available data correspond to  $M < 5.8$  and distances exceeding about 31 mi (50 km). Considerable effort has been directed to developing appropriate attenuation relations for the CEUS conditions. In general, the attenuation relationships utilize standard forms to regress on recorded data in the region, augmented by data from other similar tectonic regimes and stochastic time histories tied to source types and styles of faulting.

Since publication of the 1989 EPRI study (EPRI, 1989a), much work has been done to evaluate strong earthquake ground motion in the CEUS. In 2004, EPRI completed a study on strong ground motion prediction in the CEUS following the SSHAC (NRC, 1997b) guidelines for a Level

III Analysis. A panel of six ground motion Experts was reconvened during several workshops to provide advice to a Technical Integrator (TI) on the adequacy of available CEUS ground motion relationships. On this basis, the TI developed a representation of the current scientific understanding on the subject, consisting of "clusters" of ground motion relationships with associated weights to represent the uncertainty in predicting the median ground motion, in terms of moment magnitude. Each cluster corresponds to relationships based on a similar approach for ground motion modeling. The uncertainty in the median model for each ground motion cluster is defined by two additional models: one representing the 5<sup>th</sup> percentile of the median uncertainty distribution and the other corresponding to the 95<sup>th</sup> percentile.

Epistemic uncertainty is modeled using multiple ground motion equations and multiple estimates of aleatory uncertainty (sigma), all with associated weights. Different sets of equations are recommended for sources that represent rifted versus non-rifted parts of the earth's crust. Equations are available for spectral frequencies of 100 Hz (equivalent to PGA), 25 Hz, 10 Hz, 5 Hz, 2.5 Hz, 1 Hz, and 0.5 Hz, and these equations apply to hard rock conditions, i.e., rock with a shear wave velocity of 9200 ft/sec (2804 m/sec).

EPRI has published updated estimates of aleatory uncertainty (EPRI, 2006a). This update reflected the observation that sources of the aleatory uncertainties in the original EPRI attenuation study (EPRI, 2004) were probably too large, resulting in over-estimates of seismic hazard. The 2006 EPRI study (EPRI, 2006a) recommends a revised set of aleatory uncertainties (sigmas) with weights that can be used to replace the original aleatory uncertainties published in the 2004 EPRI study (EPRI, 2004).

In accordance with Regulatory Guide 1.208 (NRC, 2007a), the hazard curves from the PSHA have to be defined for generic hard rock conditions as defined in the development of the attenuation equations. The 2004 EPRI ground motion models correspond to a shear wave velocity ( $V_s$ ) of 9200 ft/sec (2804 m/sec). These EPRI 2004 equations have been adopted for median ground motion estimates, and the Abrahamson log-sigma model (EPRI, 2006a) is used to incorporate aleatory variability. Within this context, Figure 2.5-76 shows the logic tree for general area sources such as background or local source, and Figure 2.5-75 shows the logic tree for non-general sources such as New Madrid and Charleston. Adopting the EPRI 2004 ground motion model implies that the seismic hazard is calculated at the location where the rock reaches a  $V_s$  of 9200 ft/sec (2804 m/sec).

EPRI TR-1014381 (EPRI, 2006a) was used in lieu of the Regulatory Guide 1.208 cited document, i.e. EPRI Report 1013105 (EPRI, 2006b). EPRI Report 1013105 (EPRI, 2006b) was an Update Report while EPRI TR-1014381 (EPRI, 2006a) is the final report. For the purposes of revised estimates of aleatory uncertainty in the CEUS, there is no technical difference between the documents. The "Recommended CEUS Sigma" values and "Conclusions" of both reports are identical.

Earthquakes occurring within the area seismic sources were treated as point sources. Thus, the adjustments to the ground motion equations developed in EPRI (EPRI, 2004) to account for this point-source representation were incorporated in the hazard calculations.

#### **2.5.2.4.6 Updated EPRI Probabilistic Seismic Hazard Analysis Deaggregation, and 1 Hz, 2.5 Hz, and 10 Hz Spectral Accelerations Incorporating Significant Increases Based on the Above Sensitivity Studies**

Figures 2.5-67 through Figure 2.5-73 and Tables 2.5-19 through 2.5-25 present the resulting updated probabilistic seismic hazard hard rock curves for the seven spectral ordinates (100 Hz (equivalent to PGA), 25 Hz, 10 Hz, 5.0 Hz, 2.5 Hz, 1.0 Hz, and 0.5 Hz). The mean and fractile (5%, 16%, 50% (median), 84% and 95%) hazard curves are indicated.

Figure 2.5-43 shows mean and median uniform hazard spectra for  $10^{-4}$ ,  $10^{-5}$ , and  $10^{-6}$  annual frequencies of exceedance from these calculations at seven structural frequencies. Numerical values of these spectra are documented in Table 2.5-14.

The mean rock hazard has been de-aggregated for the  $10^{-4}$ ,  $10^{-5}$  and  $10^{-6}$  levels of probability of exceedance. The magnitude and distance bins for the de-aggregation table were taken from Regulatory Guide 1.208 (NRC, 2007a). The results have been plotted in Figures 2.5-44 through Figure 2.5-47, Figure 2.5-85 and Figure 2.5-86, for the required low frequency (1 and 2.5 Hz), the high frequency (5 and 10 Hz) ranges, and for the  $10^{-4}$ ,  $10^{-5}$ , and  $10^{-6}$  levels of probability of exceedance, respectively. These figures depict the percent contribution of each magnitude-distance bin to the total hazard.

Approach 2B of NUREG/CR-6728 (NRC, 2001) was used to derive the controlling events at the BBNPP Site. First, the controlling events were identified using the de-aggregation results. Table 2.5-11 lists the de-aggregated controlling events. Each de-aggregated earthquake (DE) is prescribed as a pair of distance and its associated contribution to a high frequency (HF) or low frequency (LF) response. DEL indicates the low end of the distance range, while DEM and DEH refer to the middle and high ends, respectively. Using the magnitude-distance pairs for each sub-controlling event, DEL, DEM, and DEH, the CEUS single corner spectral shapes from NUREG/CR-6728 (NRC, 2001) were adopted to develop the corresponding spectral shapes. Then, the response spectra of each sub-controlling event were scaled to match the rock UHRS at 1.75 or 7.5 Hz for low frequency and high frequency events (Reference Events, REF), respectively. The resulting scaled response spectra are presented on Figure 2.5-77 and Figure 2.5-78 for hazard levels of  $10^{-4}$  and  $10^{-5}$ , respectively.

The de-aggregation of the total hazard clearly reveals that the nearby area sources largely govern the hazard at the BBNPP Site. The influence of local earthquakes is more appreciable in the HF motion. Each of the controlling earthquakes (DEL, DEM, and DEH) of both LF and HF was taken as input for the seismic site amplification analyses as described in the following section. For each sub-controlling event, all selected time histories have been scaled and modified to match their calculated response spectra with the target scaled response spectrum. Figure 2.5-79 compares the target response spectrum with the response spectra of selected time histories, after performing the spectral matching for the  $10^{-4}$  hazard low frequency controlling event of the sub-controlling DEL and DEH.}

### **2.5.2.5 Seismic Wave Transmission Characteristics of the Site**

The U.S. EPR FSAR includes the following COL Item in Section 2.5.2.5.

Seismic wave transmission characteristics are site specific and will be addressed by the COL applicant.

This COL Item is addressed as follows:

{The uniform hazard spectra developed through Section 2.5.2.4 and displayed on Figure 2.5-43 are defined on hard rock (shear-wave velocity of 9200 ft/sec (2804 m/sec)). Rock layers with shear-wave velocities of such value are located at depths between 190 ft (57.9 m) below the foundation level at the BBNPP Site. To determine the SSE at the ground surface, it is necessary to adjust the uniform hazard spectra for amplification or de-amplification as the vibratory ground motion propagates through the soil media. As mentioned above, the adjustment was made by conducting Site Response Analyses following Approach 2B described in NUREG/CR-6728 (NRC, 2001). These analyses consist in defining the shear wave velocity and material damping characteristics in the soil and rock profile between the ground surface and the depth of hard rock. Then uni-dimensional site analyses are conducted using equivalent linear procedure (Schnabel, 1972). The results are used to derive site amplification factors for modifying the response spectra at rock on account of the seismic wave transmission characteristics of the soil layers. This section describes the various steps involved in the calculation and application of the site amplification factors. The seismic wave transmission characteristics and effects of this thick soil column on hard rock ground motions are described in this section.

Section 2.5.2.5.1 is added as a supplement to the U.S. EPR FSAR.

#### **2.5.2.5.1 Development of Site Amplification Functions**

##### **2.5.2.5.1.1 Methodology**

The calculation of site amplification factors is performed in the following 4 steps:

1. Develop a best estimate soil and rock column in which mean low-strain shear wave velocities and material damping values, and strain-dependencies of these properties, are estimated for relevant layers from the surface to the hard rock horizon. At the BBNPP site, hard rock ( $V_s = 9200$  ft/sec (2804 m/sec)) is at sloping depths between 190 ft (57.9 m) and 237 ft (72.2 m);
2. Develop a probabilistic model that describes the uncertainties in the above properties, locations of layer boundaries, and correlation between the velocities in adjacent layers, and generate a set of 60 artificial "randomized" profiles;
3. For each of the sub-controlling earthquakes (DEL, DEM, and DEH) of  $10^{-4}$  and  $10^{-5}$  annual frequencies of exceedance for both LF and HF earthquakes, use the corresponding controlling time histories for input into dynamic response analysis as the outcrop motion at the hard rock elevation;
4. Use an equivalent-linear time-history site-response formulation to calculate the dynamic response of the site for each of the 60 artificial profiles, and calculate the mean of site

response. This step is repeated for each de-aggregated earthquake of the four input motions ( $10^{-4}$  and  $10^{-5}$  annual frequencies, HF and LF events).

These steps are described in the following subsections. The calculation of site effects was performed with an in-house version of the computer program SHAKE (Schnabel, 1972). This program computes the response in a system of viscous-elastic, horizontally layered, soil units, overlying a uniform half-space, subjected to transient, vertical travelling shear waves.

The analytical method implemented in SHAKE is based on the solution of the wave equation and the Fast Fourier Transform algorithm. The nonlinearity of the shear modulus and damping is accounted for by the use of equivalent linear soil properties within an iterative procedure to obtain values for modulus and damping compatible with the effective strains in each layer. Therefore, for any set of layer properties, SHAKE performs a linear analysis.

The motion used as basis for the analysis (i.e., the motion that is considered to be known) can be applied to any layer in the system. An iterative procedure is used to account for the nonlinear behavior of the soils. The object motion can be specified at the top of any sub-layer within the soil profile or at the corresponding outcrop.

It is noted that the solution of a particular problem requires use of realistic ground motions (loading), modeling site dynamics (response), and the interpretation and prediction of soil behavior subject to dynamic loading (analysis). To facilitate conducting and verifying these tasks, modifications incorporated in Rizzo's in-house version include the following:

- The number of sub-layers was increased to up to 500 to allow a more accurate representation of deeper and/or softer soil deposits;
- Modulus reduction and damping relationships can be specified by the user, up to 13 different curves;
- User specified periods are allowed for calculating spectral ordinates;
- The code can accept input data to generate random soil/rock columns by utilizing best estimates of the mean and the standard deviation along with prescribed probabilistic distributions for material properties (stiffness, mass and damping) and for layer thickness.

Computer software control for SHAKE has been done according to Rizzo's Quality Assurance Manual. The Quality Assurance Manual addresses software activities including software acquisition and development, tracking installation of design and analysis software on individual computers, program verification and validation, in-use testing, software usage, change control, configuration control, error notice documentation and distribution, software maintenance, virus protection, software retirement, records and monitoring.

To verify and validate the reliability and functionality of the Rizzo's in-house version of SHAKE, six validation problems are chosen. Each function in the program is verified at least once by the sample problems. One of the sample problem intents to verify the capability of the number of soil layers of 500 in the in-house version. The results calculated by the program are compared to analytical solutions from public sources. The validation and verification presents a good agreement between SHAKE computational solution and analytical solution for each sample problem.

#### **2.5.2.5.1.2 Base Case Soil/Rock BBNPP and Uncertainties**

Development of a best estimate soil/rock column is described in detail in Section 2.5.4. Summaries of the low strain shear wave velocity, material damping, and strain-dependent properties of the base case materials are provided below in this section. These parameters are used in the site response analyses.

The total depth of approximately 386 ft (117.7 m) of the BBNPP Site was investigated using test borings and geophysical methods. The geotechnical investigation is described in detail in Section 2.5.4.

The layers in the 386 ft (117.7 m) of the site consist of the following stratigraphic units:

- Overburden Soils:
  - Glacial Overburdens
- Rock Formations:
  - Mahantango Shale

A layer of concrete with an average thickness of 10 ft (3 m) below the center line of the planned nuclear reactor facility will be built on top of the Mahantango Shale. This concrete layer is placed between the power block basemat and the bedrock. Section 2.5.4 provides detailed contour information related to the position of the bedrock below the power block's footprint.

The compressional and shear-wave velocities are taken from geophysical field tests using two different techniques:

1. Four sets of downhole tests,
2. Four sets of suspension logging tests

Of the eight geophysical measurements, two borings, G301 and B301, provide the deeper site-specific geophysical information collected during the geotechnical investigation. P-S Suspension and downhole tests were performed down to a depth of approximately 400 ft (122 m) (GeoVision, 2008; NGA, 2008). These two locations are at the center line of the projected containment footprint.

The downhole profiles consist of average compressional and shear-wave velocities for thicknesses varying from 13 ft (4 m) to 120 ft (36.6 m). The suspension logging profiles provide detailed discrete compressional and shear-wave velocities for thickness of approximate 1.5 ft (0.5 m).

Resonant Column and Torsional Shear Laboratory Tests were performed on soil and backfill samples. The complete set of results from these tests is reported in Section 2.5.4.2.3. Generic cohesionless soil curves (EPRI, 1993) were adopted to describe the strain dependencies of shear modulus and damping for the backfill based on available results from the site investigation. As required by Regulatory Guide 1.208 (NRC, 2007a) the damping curves for soils were truncated at 15 percent for the site response analysis.



In these areas, there are numerous records of deep gamma ray surveys and geologic columns with lithologic descriptions. The analysis of shear wave velocities at depths beyond the reach of the boring exploration program became irrelevant since the 9200 ft/sec (2804 m/sec) horizon was clearly encountered by the geophysical exploration program.

The Mahantango Formation reached such shear wave velocity above a depth of 350 ft (107m). Past reports place the total thickness of the Mahantango Formation at approximately 1,500 ft (457 m) (Inners, 1978).

The geologic column at the site is an extension of the Mahantango Shale, which is a dark gray to black formation, with few to no fractures. Some distinctive features are the presence of calcareous zones, the presence of thin pyrite lenses that increase in abundance with depth, and the presence of calcite veins perpendicular to the bedding plane that are micro-faulted. The upper surface of the Mahantango Formation shows the effects of solution and weathering in a few areas, but it is predominantly very competent and indurated.

For the Site Response Analyses, the concrete and Mahantango Shale is assumed to behave linearly during earthquake shaking. "Free-Free" Direct arrival tests were performed on undisturbed rock samples by the University of Texas. The "Free-Free" Direct arrival test results are provided in Table 2.5-42. The tests provided material velocity and damping values associated with shear-waves as well as those associated with compressional waves.

The average of the laboratory test results for damping for the Mahantango Shale is 0.86 percent. Lower values, 0.8 and 0.7 percent, are conservatively used for the analysis. The Mahantango Shale has a very high Rock Quality Designation (RQD) and as a rock mass is capable of transmitting shear waves very efficiently with small amounts of damping. Therefore, the lower reported laboratory values are selected for the analysis. The RQD of the Mahantango Shale is reported in the field boring logs.

#### **2.5.2.5.1.3 Site Properties Representing Uncertainties and Correlations**

To account for variations in shear-wave velocity across the site, 60 artificial profiles were generated using the stochastic model developed by Toro (Toro, 1996), with the approximation of the standard deviation of  $\ln V_s$  as the coefficient of variation of  $V_s$  (Ang and Tang, 1975). These artificial profiles represent the soil column from the top of the ground surface to the top of bedrock with a shear-wave velocity of 9,200 ft/s (2804 m/sec). The model uses as inputs the following quantities:

- The best estimate of the shear-wave velocity profile and other soil properties described above;
- The coefficient of variation of the shear wave velocity as a function of depth, developed using available site data (refer to Section 2.5.4);
- Correlation coefficients between  $V_s$  in adjacent layers, determined using correlation results for the USGS site characterization category (Toro, 1996);
- The probabilistic characterization of layer thickness as a function of depth, computed assuming a normal distribution;
- The depth to bedrock, which is randomized assuming a normal distribution to account for epistemic uncertainty in the bedrock-depth data described in Section 2.5.4.

Figure 2.5-50 shows the best estimate  $V_s$  value and corresponding coefficient of variation as a function of depth of the downhole tests and suspension logging tests at different boreholes.

The coefficient of variation of shear wave velocities calculated from the best estimate soil/rock column is used as the standard deviation of  $\ln(V_s)$  as a function of depth. Figure 2.5-51 shows the coefficient of variation of shear-wave velocity, which were used to generate multiple profiles. The correlation coefficients between shear wave velocities in adjacent layers were determined using USGS empirical relationships.

The randomly generated thicknesses of layer were computed assuming a normal distribution using the coefficient of variation of 0.10 to 0.15 for thickness of each layer. For consistency with the site-specific data, the generated  $\ln$ -velocities and the generated thicknesses were truncated at  $\pm 2s$  according to the recommendations of Toro (Toro, 1996).

Figure 2.5-52 illustrates the  $V_s$  profiles generated for profiles 1 through 60, using the median, logarithmic standard deviation, and correlation models described. These profiles include uncertainty in depth to bedrock. In total, 60 profiles were generated. Figure 2.5-53 compares the mean of these 60  $V_s$  profiles to the best estimate  $V_s$  profile described in the previous section, indicating very good agreement. This figure also shows the  $\pm 1$  standard deviation values of the 60 profiles, reflecting the coefficient of variations indicated on Figure 2.5-51.

Mean values of shear stiffness ( $G/G_{MAX}$ ) and damping for each geologic unit are described in Section 2.5.4. Uncertainties in the properties for each soil unit are characterized using the values obtained by Costantino (Costantino, 1996). Figure 2.5-54 and Figure 2.5-55 illustrate the shear stiffness and damping curves generated for backfill, although that is not present in the Best Estimate soil column model. Stiffness and damping of soils depend on the strain level during ground shaking. However, for significantly stiff materials such as concrete and the Mahantango Shale, these properties are independent of the strain level during earthquake ground motion. Both properties retain their "low-strain" values. These values are also subject to the random variation procedure.

This set of 60 profiles, consisting of  $V_s$  versus depth, depth to bedrock, stiffness, and damping, are used to calculate and quantify site response and its uncertainty, as described in the following sections.

#### **2.5.2.5.1.4 Development of Smooth Uniform Hazard, Controlling, and Reference Response Spectra**

In order to derive smooth spectra corresponding to the  $10^{-4}$  and  $10^{-5}$  amplitudes, the magnitude and distance pairs of both controlling and reference earthquakes summarized in Table 2.5-11 were used as described below.

The magnitudes and distances were applied to spectral shape equations from NUREG/CR-6728 (NRC, 2001) to determine realistic spectral shapes for the four representative earthquakes (at spectral frequency 0.5, 1.75, 7.5, and 25 Hz) of  $10^{-4}$  and  $10^{-5}$  events.

For smooth Uniform Hazard Response Spectra (UHRS), the 25 Hz smooth shapes were utilized and scaled to the Uniform Hazard Spectra mean values for  $10^{-4}$  or  $10^{-5}$  between 25 Hz and 100 Hz. The 7.5 Hz smooth shapes were utilized and scaled to the Uniform Hazard Spectra mean values for  $10^{-4}$  or  $10^{-5}$  between 5 Hz and 10 Hz. The 1.75 Hz smooth shapes were utilized and scaled to the Uniform Hazard Spectra mean values for  $10^{-4}$  or  $10^{-5}$  between 0.5 Hz and 2.5 Hz.

Below 0.5 Hz, the 0.5 Hz smooth shapes were scaled and utilized without any modification. The smooth UHRS are presented in Table 2.5-61, and in Figure 2.5-48 for  $10^{-4}$ ,  $10^{-5}$ , and  $10^{-6}$ .

For the reference response spectra, the HF reference spectra shapes for  $10^{-4}$  or  $10^{-5}$  at the spectral frequency above 5 Hz were the same as the smooth UHRS. The spectral shape at 7.5 Hz was extrapolated from 5 Hz without regard to Uniform Hazard Spectra amplitudes at lower frequencies. The LF reference spectra shapes were scaled to the smooth UHRS values for  $10^{-4}$  or  $10^{-5}$  for frequency less than 2.5 Hz. Above 2.5 Hz, the spectral shape was extrapolated from 2.5 Hz, without regard to Uniform Hazard Spectra amplitudes at higher frequencies by using the smooth spectral shape at frequency of 1.75 Hz.

Creation of smooth  $10^{-4}$  and  $10^{-5}$  reference spectra (REF) in this way ensures that the HF spectra match the  $10^{-4}$  and  $10^{-5}$  Uniform Hazard Spectra values at high frequencies (5 Hz and above), and ensures that the LF spectra match the  $10^{-4}$  and  $10^{-5}$  Uniform Hazard Spectra values at low frequencies (2.5 Hz and below). In between calculated values, the spectra have smooth and realistic shapes that reflect the magnitudes and distances dominating the seismic hazard, as reflected in Table 2.5-12. The smooth reference spectra are presented in Figure 2.5-80 and Figure 2.5-81. For controlling response spectra, the smooth spectra shapes for  $10^{-4}$  and  $10^{-5}$  events, LH and HF, and sub-event, DEL, DEM, and DEH were developed directly from the NUREG/CR-6728 using the magnitudes and distances in Table 2.5-11 without any modification. These smooth spectra then scaled to match the smooth UHRS at 1.75 Hz for LF events and at 7.5 Hz for HF events. The smooth controlling response spectra are presented on Figure 2.5-77 and Figure 2.5-78.

#### **2.5.2.5.1.5 Controlling Time Histories**

Four initial time histories were selected from the rock time histories database from NUREG/CR-6728 (NRC, 2001) for sub-controlling earthquakes (DEL, DEM, and DEH) for the  $10^{-4}$  and  $10^{-5}$  levels and for both LF and HF events according to their deaggregated magnitudes and distances. These time histories were then modified according to the spectral matching criteria set for time histories in Appendix F of Regulatory Guide 1.208 (NRC, 2007a) to match their target smooth controlling response spectra. The selected time histories are listed in Table 2.5-17.

#### **2.5.2.5.1.6 Site Response Analysis**

The site response analysis performed for the BBNPP Site used a time history-based procedure in conjunction with the following assumptions:

- Vertically-propagating shear waves are the dominant contributor to site response.
- An equivalent-linear formulation of soil nonlinearity is appropriate for the characterization of site response.

Sixty response analyses were performed using the program SHAKE (Schnabel, 1972) to calculate the site amplification function for each de-aggregation earthquake. The 60 randomized velocity profiles were paired with the 60 sets of randomized modulus reduction and damping curves (one profile with one set of modulus). Sixty response analyses were performed using the program SHAKE (Schnabel, 1972) as modified by Rizzo to calculate the site amplification function for each de-aggregation earthquake. The 60 randomized velocity profiles were paired with the 60 sets of randomized modulus reduction and damping curves (one profile with one set of modulus reduction and damping curves) to define 60 soil columns, each

characterized by a set of shear wave velocities, modulus reduction curves, and material damping curves. Each of the four scaled time histories corresponding to a de-aggregated earthquake was used to compute the response of fifteen profile-soil property curve sets.

For each analysis, the response spectrum for the computed motion at the top of the concrete was divided, frequency by frequency, by the response spectrum for the input motion at the hard rock to obtain a site amplification function. The arithmetic mean of these 60 individual response spectral ratios was taken as the mean site amplification function for each de-aggregated earthquake.

The following figures and table describe the site amplification factors for the high and low frequencies and  $10^{-4}$  and  $10^{-5}$  input motions:

- Figure 2.5-56: mean site amplification factor and coefficient of variation at the top of concrete for  $10^{-4}$  HF DEM input motion;
- Figure 2.5-57: maximum strains vs. depth for  $10^{-4}$  HF DEM input motion;
- Figure 2.5-58: mean site amplification factor and coefficient of variation at the top of concrete for  $10^{-4}$  LF DEM input motion;
- Figure 2.5-59: maximum strains vs. depth for  $10^{-4}$  LF DEM input motion;
- Figure 2.5-60: mean site amplification factor and coefficient of variation at the top of concrete for  $10^{-5}$  HF DEM input motion;
- Figure 2.5-61: maximum strains vs. depth for  $10^{-5}$  HF DEM input motion
- Figure 2.5-62: mean site amplification factor and coefficient of variation at the top of concrete for  $10^{-5}$  LF DEM input motion; and
- Figure 2.5-63: maximum strains vs. depth for  $10^{-5}$  LF DEM input motion.
- Table 2.5-13: amplification factors for  $10^{-4}$  and  $10^{-5}$  input motions and HF and LF rock spectra}

#### **2.5.2.6 Ground Motion Response Spectra**

The U.S. EPR FSAR includes the following COL Item in Section 2.5.2.6:

A COL applicant that references the U.S. EPR design certification will verify that the site-specific seismic parameters are enveloped by the CSDRS (anchored at 0.3 g PGA) and the 10 generic soil profiles discussed in Section 2.5.2 and Section 3.7.1 and summarized in Table 3.7.1-6.

This COL Item is addressed as follows:

This section and Section 3.7.1 describe the reconciliation of the site-specific parameters for the {BBNPP} and demonstrates that these parameters are enveloped by the Certified Seismic Design Response Spectra (CSDRS), anchored at 0.3 g PGA, and the 10 generic soil profiles used in the design of the U.S. EPR.

Table 5.0-1 of the U.S. EPR FSAR identifies shear wave velocity as a required parameter to be enveloped, defined as "Minimum shear wave velocity of 1000 feet per second (Low strain best estimate average value at bottom of basemat)."

{Figure 2.5-84 compares the 10 generic soil profile cases used for the U.S. EPR and the best estimate shear wave velocity profile that was adopted for the BBNPP site.

Reconciliation of the BBNPP site-specific seismic parameters with the U.S. EPR certified seismic design response spectra (CSDRS) and the 10 generic soil profiles used for the U.S. EPR is addressed in Section 3.7.1. The evaluation guidelines in U.S. EPR FSAR Section 2.5.2.6 are used to perform the reconciliation.

The steps and conclusions of the seismic parameter reconciliation are summarized below. Summaries of select U.S. EPR structures, systems, and components evaluations which confirm they are adequate for the BBNPP site are also provided as required by seismic reconciliation Step 9.

The seismic reconciliation steps and conclusions:

1. Step 1 of the U.S. EPR FSAR Section 2.5.2.6 seismic reconciliation guidelines is confirmation that the peak ground acceleration for the ground motion response spectrum (GMRS) is less than 0.3g. The BBNPP site-specific GMRS are described in Section 3.7.1. The peak ground acceleration for the BBNPP site-specific GMRS is confirmed to be less than 0.3g.
2. Step 2 of the U.S. EPR FSAR Section 2.5.2.6 seismic reconciliation guidelines is confirmation that the low strain, best estimate value of the shear wave velocity at the bottom of the foundation basemat of the NI Common Basemat Structures is 1000 fps, or greater. The low strain, best estimate value of the BBNPP site-specific shear wave velocity at the bottom of the foundation basemat of the NI Common Basemat Structures is confirmed to be greater than 1000 fps.
3. Step 3 of the U.S. EPR FSAR Section 2.5.2.6 seismic reconciliation guidelines is confirmation that the foundation input response spectra (FIRS) are enveloped by the certified seismic design response spectra (CSDRS). Comparison of the BBNPP site-specific GMRS/FIRS with the U.S. EPR CSDRS is described in Section 3.7.1. The site-specific horizontal and vertical GMRS/FIRS exceed the envelope of the U.S. EPR CSDRS ground motions, primarily in the high frequency region. The BBNPP design ground motion response spectra are as described in Section 3.7.1, instead of the CSDRS, because the GMRS/FIRS exceed the CSDRS. This represents a departure from the U.S. EPR FSAR, as described in Section 3.7.1.
4. Step 4 of the U.S. EPR FSAR Section 2.5.2.6 seismic reconciliation guidelines is confirmation that the site-specific soil profile is laterally uniform. Horizontal soil layering is confirmed for the BBNPP site-specific soil profile.
5. Step 5 of the U.S. EPR FSAR Section 2.5.2.6 seismic reconciliation guidelines is confirmation that the idealized site soil profile is similar to or bounded by the 10 generic soil profiles used for the U.S. EPR. The BBNPP idealized site soil profile is described in Section 3.7.1. The BBNPP idealized site soil profile is not considered bounded by the U.S. EPR 10 generic soil profiles. This represents a departure from the U.S. EPR FSAR, as described in Section 3.7.1.

6. Step 6 of the U.S. EPR FSAR Section 2.5.2.6 seismic reconciliation guidelines is confirmation that the conditions of Steps 1 through 5 are met. The conditions of Steps 3 and 5 are not met for the BBNPP site because the BBNPP site-specific GMRS/FIRS exceed the envelope of the U.S. EPR CSDRS and the BBNPP site-specific idealized site soil profile is not bounded by the 10 generic soil profiles used for the U.S. EPR. Because the conditions of Steps 3 and 5 are not met for the BBNPP site, seismic reconciliation guideline Step 7 is performed.
7. Step 7 of the U.S. EPR FSAR Section 2.5.2.6 seismic reconciliation guidelines is performance of intermediate-level studies, such as evaluation of the site-specific motion at the top the of the basemat, to demonstrate that the site is bounded by the design of the U.S. EPR. BBNPP site-specific response spectra are developed for the NI Common Basemat Structures basemat and the footprints of the EPGB and ESWB and are compared to the corresponding U.S. EPR design certification spectra. The BBNPP site-specific spectra exceed the envelope of the U.S. EPR certified design spectra; therefore, seismic reconciliation guideline Step 8 is performed.
8. Step 8 of the U.S. EPR FSAR Section 2.5.2.6 seismic reconciliation guidelines is performance of site-specific soil-structure interaction (SSI) analyses, development of in-structure response spectra (ISRS), and confirmation that the BBNPP site-specific ISRS do not exceed the ISRS for the U.S. EPR design certification by more than 10% at the key building locations. BBNPP site-specific SSI analyses are performed and site-specific ISRS are developed for comparison to the U.S. EPR design certification ISRS. The U.S. EPR design certification SSI analysis methodology is used to perform the site-specific SSI analyses, except as noted in Sections 3.7.1 and 3.7.2. Performance of the SSI analyses and comparison of the BBNPP site-specific ISRS with the U.S. EPR design certification ISRS is described in Section 3.7.1. The BBNPP site-specific ISRS exceed the envelope of the U.S. EPR certified design ISRS by more than 10% at some of the specified key building locations. This represents a departure from the U.S. EPR FSAR, as described in Section 3.7.1. Therefore, seismic reconciliation guideline Step 9 is performed.
9. Step 9 of the U.S. EPR FSAR Section 2.5.2.6 seismic reconciliation guidelines is performance of additional evaluations to confirm that safety-related structures, systems, and components of the U.S. EPR at the building locations where BBNPP site-specific ISRS exceed the ISRS for the U.S. EPR design certification by more than 10% are not affected. These evaluations, summarized below, confirm that the safety-related structures, systems and components of the U.S. EPR are not affected.

The BBNPP average shear wave velocity profile shown in the above figure is for soils below elevation +638 ft (195 m) (bottom of the basemat). Soils such as Glacial Overburdens will not be used for support of foundations of the BBNPP Nuclear Island. Therefore, shear wave velocity measurements in the BBNPP site soils above elevation +638 ft (195 m) regardless of value, are excluded from this evaluation as they lie above the basemat. Results from the above figure indicate that:

1. The BBNPP average shear wave velocity profile is bounded by the 10 generic profiles used for the U.S. EPR.
2. The BBNPP average shear wave velocity profile offers a shear wave velocity at the bottom of the basemat (approximate elevation +638 ft (195 m)(or depth = 0 in the above figure)) of 7,240 ft/sec (2,207 m/sec).

3. The minimum shear wave velocity from the BBNPP average shear wave velocity profile is 6,800 ft/sec (2,073 m/sec).

On the above basis, it is concluded that the BBNPP site shear wave velocity profile is bounded by the 10 generic soil profiles used for the U.S. EPR and meets the minimum 1,000 ft/sec (305 m/sec) criterion identified in the U.S. EPR FSAR.}

As described in Section 2.5.2.4, the end results of the site response analysis are weighted average site amplification factors. In this section, these factors are used to develop the ground motion response spectra (GMRS) by modifying the spectra at rock. The GMRS was developed in accordance with the performance-based approach described in Regulatory Position 5 of Regulatory Guide 1.208 (NRC, 2007a).

The Safe Shutdown Earthquake (SSE) ground motion was developed starting from the  $10^{-4}$  and  $10^{-5}$  rock Uniform Hazard Spectra. At high frequencies, the appropriate ( $10^{-4}$  or  $10^{-5}$ ) HF mean amplification factor was applied to the  $10^{-4}$  or  $10^{-5}$  HF smooth rock spectrum, to calculate site spectral amplitudes for  $10^{-4}$  and  $10^{-5}$  annual frequencies of exceedance. At low frequencies, a similar technique was used with the LF mean amplification factors. At intermediate frequencies, the larger of the HF and LF site spectral amplitudes was used.

{Figure 2.5-64 illustrates the resulting site spectra. At high frequencies, the HF spectral amplitudes are always greater, and at low frequencies, the LF spectral amplitudes are always greater.

This procedure implements Approach 2B in NUREG/CR-6728 (NRC, 2001) and NUREG/CR-6769 (NRC, 2002), where in the rock Uniform Hazard Spectra (for example, at  $10^{-4}$ ) is multiplied by a mean amplification factor at each frequency to estimate the  $10^{-4}$  site Uniform Hazard Spectra. Note that the amplification factors plotted on Figure 2.5-56, Figure 2.5-58, Figure 2.5-60, and Figure 2.5-62 are logarithmic mean amplification factors, which correspond approximately to the median. The amplification factors used to prepare Figure 2.5-64 are arithmetic mean amplification factors, which are slightly higher than the median.

The low-frequency character of the spectra on Figure 2.5-64 reflects the low-frequency amplification of the site, as shown in the amplification factors of Figure 2.5-56, Figure 2.5-58, Figure 2.5-60, and Figure 2.5-62. That is, there is a fundamental site resonance at about 0.22 Hz, with a dip in site response at about 0.4 Hz, and this dip occurs for all 60 of the site profiles that were used to characterize the site profile. As a result, there is a dip in the site spectra for  $10^{-4}$  and  $10^{-5}$  at 0.4 Hz that reflects the site characteristics.

The ASCE (ASCE, 2005b) performance-based approach was used to derive an SSE from the  $10^{-4}$  and  $10^{-5}$  site spectra. The SSE spectrum is derived at each structural frequency as follows:

$$A_R = SA(10^{-5})/SA(10^{-4})$$

$$DF = 0.6 A_R^{0.8}$$

$$SSE = \max(SA(10^{-4}) \times \max(1.0, DF), 0.45 \times SA(10^{-5}))$$

The last term in the above equation was not published in this form in ASCE (ASCE, 2005) but is a supplemental modified form, as presented in NRC Regulatory Guide 1.208 (NRC, 2007a). The resulting horizontal SSE spectrum is plotted in Figure 2.5-65.

A vertical SSE spectrum was constructed from the horizontal SSE spectrum following the approach described in NUREG/CR-6728 (NRC, 2001) by deriving vertical-to-horizontal (V/H) ratios and applying them to the horizontal SSE. As background and for comparison purposes, V/H ratios were obtained by the following methods:

The vertical SSE spectrum was constructed from the horizontal Design Response Spectrum (DRS) using vertical to horizontal (V/H) response spectral ratios appropriate for the BBNPP Site. The V/H ratios are developed following the approach described in NUREG/CR-6728 (NRC, 2001). Figure 2.5-66 shows the V/H ratios recommended for CEUS rock sites as a function of spectral frequency and the level of peak ground acceleration (PGA) for the horizontal component. Figure 2.5-66 shows the weighted average of these V/H ratios based on the PGA for the de-aggregated earthquakes (DE's) that make up the high-frequency (HF) and low-frequency (LF) mean  $10^{-4}$  reference earthquakes (RE's). The weights assigned to the DE are listed in Table 2.5-11. The weighted V/H ratios are essentially the same for the HF and LF mean  $10^{-4}$  DE.

The EPRI 2004 ground motion model for CEUS is defined at the hard rock or at the elevation that the shear wave velocity in the material is approximately 9200 ft/sec (2804 m/sec). Only the horizontal component of the ground motion is defined in this ground motion model, not the vertical component. Consequently, the PSHA is done at the hard rock level for the horizontal ground motion component. The site response analysis is performed to bring the ground motion from the hard rock elevation to the ground surface or top of competent material to define the GMRS according to Regulatory Guide 1.208 (NRC, 2007a). The end result of the site response analysis is the horizontal ground motion at free field or top of competent material. In order to define the vertical ground motion component, Regulatory Guide 1.208 (NRC, 2007a) Section C5.2 permits using the procedure described in NUREG/CR-6728 (NRC, 2001) for the CEUS soil site. The procedure begins by calculating the V/H ratio of the rock site in the CEUS via the set of equations provided in NUREG/CR-6728 (NRC, 2001). The transfer function calculated from the ratio of V/H ratio of the soil site with respect to the V/H ratio of the rock site in the WUS soil site is applied the V/H ratio of the rock site in the CEUS to obtain V/H ratio for the soil site in the CEUS. The Clinton ESP (EGC, 2006) also performed the GMRS calculation according to this procedure. The Clinton ESP application has been accepted by the NRC (NRC, 2007c).

The vertical DRS is obtained by scaling the horizontal DRS by the soil V/H ratios shown on Figure 2.5-66. A smooth spectrum enveloping the vertical DRS was then constructed. The resulting vertical SSE is shown on Figure 2.5-65 and is tabulated in Table 2.5-12 along with the horizontal SSE spectrum.

Refer to Sections 3.7.1 and 3.7.2 for a description of the soil-structure interaction analyses performed for the U.S. EPR design certification.}

#### CAV Filtering In Surface Ground Motions

The use of a lower bound magnitude in the calculation of the probabilistic seismic hazard could result into some excessive conservatism as a consequence of including the effects of non-damaging earthquakes. The reason is that, according to probabilistic methodologies and current attenuation equations, small magnitude near site events could occur very frequently having a significant contribution to the integrated hazard. However, it has been found that facilities designed and built with sound engineering practices do not suffer damage from this type of events (EPRI, 1988a). Examining this issue, the Cumulative Absolute Velocity (CAV) was proposed as a parameter for quantifying the damage potential associated to an earthquake



record (EPRI, 1988). For a given accelerogram,  $a(t)$ , the CAV is calculated with the following equation:

$$CAV = \sum_i H(pga_i - 0.025g) \int_i |a(t)| dt$$

$pga$  is peak ground acceleration

$g$  is gravity

$H(x)$  is the Heaviside function (unity for  $x > 0$  and 0 otherwise)

It should be noted that the surface ground motion  $a(t)$  is used to calculate the CAV. It has been observed that no damage occurs on well designed and built structures when the CAV is equal to or lower than 0.16g-sec (EPRI, 2006).

Recently, EPRI (EPRI, 2006) has published methodologies for incorporating the CAV filter into seismic hazard calculations. The most direct method consists in including the probability of exceeding the 0.16g-sec threshold into the integral to calculate the hazard. This, however, would require that site effects be included in the hazard integration, for instance, in the attenuation equations. In addition, the computation time would be significantly increased. Thus, EPRI (EPRI, 2006) has also developed a more efficient method for applying the minimum as a post-processing procedure to the hazard calculation. EPRI TR-1014099 (EPRI, 2006) was used in lieu of the Regulatory Guide 1.208 (NRC, 2007a) cited document (EPRI Report 1012965). EPRI Report 1012965 was an update report for CAV research while EPRI TR-1014099 (EPRI, 2006) is the final report. For the purposes of revised calculation of the CAV in the CEUS, there is no technical difference between the documents. The methodologies of calculation of the CAV of both reports are identical. This approach uses the hazard curve and the de-aggregation obtained in the PHSA at rock to calculate the rate of occurrence,  $v(z_k, i, j)$ , of the spectral acceleration around a small acceleration range close to  $z_k$ , due to a magnitude-distance pair ( $M_i, R_j$ ). Equations developed by EPRI to estimate the CAV in terms of  $M$  and peak ground acceleration (PGA) can then be used to calculate the probability that  $P(CAV > 0.16)$  for the corresponding  $M_i, R_j$  pair, and the filtered hazard  $v'(S)$  is calculated as follows:

$$v'(S > z) = \sum_i \sum_j \sum_k v(z_k, i, j) P(CAV > 0.16)$$

The CAV filtering is implemented by first breaking the hazard curve at rock into rates of occurrence of scenario earthquakes ( $M, R, PGA$ ). We can then compute the probability that this scenario will lead to a CAV value greater than 0.16g-sec. This probability is then multiplied by the rate of the scenario, and the sum of the filtered rates furnishes the CAV filtered hazard. The spectral value can be related to a corresponding PGA using the uniform hazard spectrum shape at the corresponding exceedance rate.

Following details presented in EPRI (EPRI, 2006), the CAV filtering was incorporated as a post-processing application into the hazard calculation at the BBNPP Site. Very modest reductions in spectral values were obtained, particularly for the  $10^{-5}$  hazard. The explanation is that after applying the site amplification factors, the PGA values corresponding to this hazard level are relatively high (about 0.4g) and, consequently, almost certainly damaging. In fact, CAV reductions on the GMRS were negligible.

### 2.5.2.7 Conclusions

{This section is added as a supplement to the U.S. EPR FSAR.

An updated evaluation of the vibratory ground motion has been conducted for the BBNPP Site. A Probabilistic Seismic Hazard Analysis (PSHA) was selected as the appropriate basis for evaluating the vibratory ground motion accounting for all credible alternative seismic sources. The alternative seismic sources identified by the Electric Power Research Institute (EPRI) for the Central and Eastern United States (CEUS), Seismic Hazard Methodology for the Central and Eastern United States (EPRI, 1986) issued in 1986 are still considered to constitute an adequate definition of seismic area sources. However, updated information available from databases maintained by the United States Geological Survey has been used to determine recurrence parameters. Since the New Madrid Fault System (NMFS) and the Charleston Seismic Source (CSS) have some contribution to the seismic hazard at the BBNPP Site, updated logic-tree representations of the clustered characteristic earthquakes at the NMFS and the un-clustered CSS have been incorporated into the PSHA. The NMFS characterization is provided by Exelon in the Clinton ESP application (EGC, 2006) and the CSS characterization is the one presented in the CCNPP Unit 3 FSAR (UniStar Nuclear, 2007). Both characterizations have been verified with USGS modeling of the New Madrid and Charleston Faults. The PSHA for the BBNPP Site makes use of a decision tree approach with appropriate weighting factors that are based on the most up-to-date information and relative confidence in alternative characterizations for each area and characteristic seismic source.

The guidance of Regulatory Guide 1.208, "A Performance-Based Approach to Define the Site-Specific Earthquake Ground Motion," (NRC, 2007a) was used to develop the Ground Motion Response Spectra (GMRS) at the BBNPP Site. This GMRS adequately represents the regional and local seismic hazards and accurately includes the effects of the local soils at the BBNPP Site.

It is concluded that the performance-based approach outlined in Regulatory Guide 1.208 (NRC, 2007a) constitutes an advancement over the solely hazard-based reference probability approach recommended in Regulatory Guide 1.165 (NRC, 1997a) and used it where appropriate in the determination of the GMRS. The performance-based approach uses not only the seismic hazard characterization of the site from the PSHA but also basic seismic fragility SSC modeling in order to define a ground motion that directly targets a structural performance frequency value. It is concluded that the application for the BBNPP Site is acceptable from a geologic and seismologic standpoint and meets the requirements of 10 CFR 100.23(d) (CFR, 2007). Deviations from the NRC guidance in Regulatory Guide 1.165 (NRC, 1997a), Regulatory Guide 1.208 (NRC, 2007a), or review criteria in Standard Review Plan 2.5.2 (NRC, 2007b) have been identified and acceptable alternatives, including technical justification, have been provided.}

#### 2.5.2.8 References

{This section is added as a supplement to the U.S. EPR FSAR.

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### **2.5.3 Surface Faulting**

The U.S. EPR FSAR includes the following COL Item in Section 2.5.3:

A COL applicant that references the U.S. EPR design certification will investigate site-specific surface and subsurface geologic, seismic, geophysical, and geotechnical aspects within 25 miles around the site and evaluate any impact to the design. The COL applicant will demonstrate that no capable faults exist at the site in accordance with the requirements of 10 CFR 100.23 and 10 CFR 50, Appendix S. If non-capable surface faulting is present under foundations for safety-related structures, the COL applicant will demonstrate that the faults have no significant impact on the structural integrity of safety-related structures, systems or components.

This COL Item is addressed as follows:

{There is no potential for tectonic fault rupture and there are no capable tectonic sources within a 25 mi (40 km) radius of the BBNPP site.} A capable tectonic source is a tectonic structure that can generate both vibratory ground motion and tectonic surface deformation, such as faulting or folding at or near the earth's surface in the present seismotectonic regime (NRC, 1997). The following Sections provide the data, observations, and references to support this conclusion. Information contained in these Sections was developed in accordance with RG 1.165 (NRC, 1997), and is intended to satisfy 10 CFR 100.23, "Geologic and Seismic Siting Criteria" (CFR, 2007a) and 10 CFR 50, Appendix S, "Earthquake Engineering Criteria for Nuclear Power Plants" (CFR 2007b).

Sections 2.5.3.1 through Section 2.5.3.9 are added as a supplement to the U.S. EPR FSAR.

#### **2.5.3.1 Geological, Seismological, and Geophysical Investigations**

The following investigations were performed to assess the potential for surface fault rupture out to a 25 mile (40 km) radius of the BBNPP site:

- Site subsurface investigations including geotechnical borings and seismic refraction surveys. (Section 2.5.3.1.1)
- Interpretation of aerial photography and satellite imagery. (Section 2.5.3.1.2)
- Field reconnaissance. (Section 2.5.3.1.3)
- Compilation and review of existing geologic and seismologic data and literature, including a review of seismic data. (Section 2.5.3.1.4)

#### 2.5.3.1.1 Subsurface Investigations at the BBNPP Site

Geologic sections developed from geotechnical data collected from 45 boreholes as part of the BBNPP study (as discussed in Section 2.5.4) provide detailed information in the upper 400 ft (122 m) of strata for the presence of structures directly beneath the site. The interpretations developed from the previous investigation at the SSES site confirm the interpretation of the new borehole data at the BBNPP site:

unfaulted Middle Devonian shale dipping steeply to the south-southeast, and covered by a layer of undeformed glacial outwash and till (Figures 2.5-96, 2.5-97, 2.5-98, 2.5-99, 2.5-100, and 2.5-100a).

Although the bedrock formations underlying the BBNPP site are steeply dipping and have experienced folding during the Alleghanian Orogeny (Williams, 1987; Faill, 1999), surficial sediments of the site display no signs of faulting or folding during the Pleistocene to Holocene time period (Figure 2.5-101), and rest unconformably on the eroded surface of the tilted beds of the local shale bedrock.

Geotechnical data collected to the southern portion of the BBNPP site was compiled along section D-D' shown in Figure 2.5-99. Although these geotechnical boreholes are limited in depth (from 99.5 ft to 200 ft (30.3 m to 60.9 m)), they provide additional evidence of the lateral continuity between the Pleistocene glacial outwash deposits and the Devonian Mahantango Formation. Figures 2.5-96, 2.5-98, and 2.5-99 display a general thickening of surficial sediments across the site (from north to south), support detailed published reports of the site area local geology, and confirm that surficial faulting and folding are absent in the interpreted sections within the BBNPP site.

Seismic refraction surveys were performed to support site characterization studies for the BBNPP (Section 2.5.4.2.2.2.3). Because earth materials exhibit characteristic wave propagation velocities, they can be classified simply in terms of their seismic velocity. Seismic refraction data were interpreted for this study to assist in characterizing the local subsurface geologic materials regarding depths to glacial till, to weathered or fractured bedrock, and to competent bedrock. Seismic refraction surveys were performed along 6 profile lines totaling 4,000 linear feet (1,219 m) of coverage. The data for the surveys were collected from January 7 through 10, 2008 using approved quality assurance procedures. The complete report of this survey (Weston, 2008) is included in COLA Part 11G.

Figure 2.5-103 is a map depicting the layout of the 6 lines used during the survey (along Lines 1 through 3, oriented north-south; and Lines A through C, oriented east-west). Seismic P-wave velocity profiles, as interpreted by the SeisOpt @2D™ software are presented on Figure 2.5-104. These profiles are plotted without vertical exaggeration, with the vertical scale measuring elevation in feet, msl. These interpreted velocity profiles indicate a generally flat-lying eroded bedrock surface overlain by a variably thin veneer of overburden material. Figure 2.5-105 is a representation of the surface of the bedrock, as indicated by an interpreted velocity of at least 14,000 fps (4,267 mps). The velocity model developed for the site depicts the bedrock surface to be nearly flat lying from west to east and dipping to the south.

The subsurface profiles constructed from manual calculations are presented in Figures 2.5-105a through 2.5-105f. These differ from the software-derived profiles, in that they have fewer layers, depict the depths of the units more accurately, and do not show the lateral changes in velocities that may be an artifact of the processing by the SeisOpt @2D™ software. The manual profiles also show the tops of the local bedrock as determined from borings installed during the site

subsurface investigation. As is evident on the Figures, these tops compare favorably with the tops of the bedrock surface determined by seismic methods. The map of the top of the Mahantango Shale from the boring log data (Figure 2.5-105g) shows the same east-west strike and southward dip as the surface shown in Figure 2.5-105 that was developed from the geophysical data.

A review of the data in the Figures shows no offsets or abrupt changes in the tops of the bedrock that could indicate recent deformation. The close correlation with the boring log data further confirms the interpretation that there is no evidence of faulting in the quaternary deposits underlying the site, in that no evidence of faulting was found in the core samples.

The somewhat irregular surface of the bedrock noted in the Refraction Survey and the boring logs is not interpreted as significant offsets of the top of the Mahantango Formation that could be attributed to faulting. Instead, the irregular bedrock surface is mimicked by the surface of the overlying glacial till layer, suggesting a glacially eroded bedrock surface. Based on the nature of the observed interface between the glacial deposits and the shale, the lack of faulting evidence in the boreholes, and the un-deformed condition of the glacial deposits over the inferred fault traces, the apparent irregular surface is the result of glacial scour, an erosional feature, across the upper part of the Mahantango Shale and supports the conclusion that faulting is not present.

#### **2.5.3.1.2 Interpretation of Aerial Photography and LIDAR Imagery**

Aerial reconnaissance within a 25 mi (40 km) radius of the site was conducted by various personnel using aerial photographs from numerous publications. Figure 2.5-105r is a sample of the aerial imagery used, and it contains selected way points from the field reconnaissance. LIDAR imagery of the BBNPP site vicinity was also acquired for review and interpretation. The central portion of the LIDAR image contains elevation data with a 2 ft (0.6 m) contour interval. For clarity, the remainder of the image is a shaded relief representation without contours. The aerial reconnaissance investigated geomorphology and targeted numerous previously mapped geologic features and potential seismic sources (e.g., Berwick fault, Light Street fault, and Berwick Anticlinorium).

As shown in the following section, field reconnaissance coupled with interpretation of remote imagery (review and inspection of features preserved on the images) shows that there are no geomorphic features indicative of potential Quaternary activity along trends of the postulated folds and faults interpreted by Inners and Williams (Inners, 1978; Williams, 1987). No features suggestive of tectonic deformation were observed in the Quaternary glaciofluvial deposits, and no potential liquefaction features were observed along the Susquehanna River.

#### **2.5.3.1.3 Field Reconnaissance**

Information developed from the literature and the imagery interpretation was supplemented by field reconnaissance within a 25 mi (40 km) radius of the site. These field-based studies were performed to verify, where possible, the existence of mapped bedrock faults in the BBNPP site area and to assess the presence or absence of geomorphic features suggestive of potential Quaternary fault activity along the mapped faults, or previously undetected faults. Features reviewed during the field reconnaissance and office-based analysis of aerial photography and LIDAR imagery were based on a compilation of existing regional geologic information in the vicinity of the BBNPP site. As shown on topographic section B-B' on figure 2.5-105s there is no topographic offset to indicate recent movement of either Light Street or Berwick Faults.

Field reconnaissance was conducted by geologists in teams of two or more. Field reconnaissance visits in 2007 and 2008 focused on exposed portions of the Mahantango Formation, other formation exposures along the faces of Lee and Nescopeck Mountains, and roads traversing the site vicinity. Key observations and discussion items were documented in field notebooks and photographs. Field locations were logged by hand on detailed topographic base maps and with hand-held Global Positioning System (GPS) receivers. There were no faults or other forms of deformation noted in the field. No surface expression of either the Berwick or Light Street faults was noted, consistent with the conclusions documented in the literature. Figures 2.5-105h and 2.5-105i (Waypoint 12 on Figure 2.5-105t) show that there is no offset in the Quaternary deposits along Syber Creek, where the trace of Light Street Fault crosses it. Photos of the shale bedrock on the site show the steeply dipping nature of the bedding (Figure 2.5-105n and Waypoint WF3 on Figure 2.5-105r). Outcrops in a nearby borrow area show an undeformed contact between the glacial overburden and the shale bedrock (Figures 2.5-105o, 2.5-105p, and 2.5-105q and Waypoint WF5 on Figure 2.5-105r).

A third reconnaissance was conducted during the fall of 2008, to investigate the occurrence of potential liquefaction features along the Susquehanna River. The field reconnaissance was carried out by a team of geologists and engineers from Paul C. Rizzo Associates, Inc, and John Sims & Associates from both the land and water approaches to the river banks. The investigation was conducted for the course of the river for a reach of 25 miles (40 km) upstream and downstream of the site (Figure 2.5-105t). Because of the prevalent bedrock exposures in both the river banks and the river bottoms, they found few locations where liquefaction conditions were possible and no evidence that liquefaction had occurred. Figures 2.5-105j through 2.5-105m show the rocky nature of the riverbed and its banks and some of the typical exposures found during the investigation (for Waypoints WP1, WP10, WP20, and WP22 respectively).

#### **2.5.3.1.4 Compilation and Review of Existing Data and Literature**

The existing body of geologic and geotechnical data regarding faulting in the vicinity of the proposed BBNPP site is contained in the following principal sources, as discussed in the three sections below:

1. Work performed for the existing Susquehanna Steam Electric Station (SSES) Units 1 and 2 and complementary structures (SSES FSAR, 2003).
2. Published geologic mapping performed primarily by the USGS and Pennsylvania Department of Conservation and Natural Resources (PA DCNR).
3. Seismicity data compiled and analyzed in published journal articles.

##### **2.5.3.1.4.1 Work at the SSES Units 1 and 2**

The most detailed previous subsurface exploration near the BBNPP site was performed as part of the original SSES FSAR (SSES FSAR, 2003) for the SSES Units 1 and 2 foundation and supporting structures. That exploration covered an area that is located adjacent to the east side of the BBNPP site and was completed to address the same issues as the current BBNPP FSAR investigation. The high level of detail in that earlier study, and the close proximity to the BBNPP site make that study especially relevant to the current effort. The level of effort for that earlier FSAR study included drilling 250 geotechnical boreholes, collecting down-hole geophysical data, and acquiring seismic refraction data from across the site. Previous site investigations performed for the existing units are summarized in the SSES Final Safety Analysis Report

(FSAR) (SSES FSAR, 2003). As cited in the SSES FSAR, these previous investigations provide the following results documenting the absence of Quaternary faults:

- Interpretation of satellite photos and topographic maps. This interpretation revealed no evidence of surface rupture, surface warping, or offset of geomorphic features indicative of active faulting.
- Interviews with personnel from government agencies and private organizations. These interviews concluded that no known faults are present beneath the existing SSES Units 1 and 2 site area.
- Seismicity Analysis -This analysis showed that no microseismic activity has occurred in the site area; the site is located in a region that has experienced only infrequent minor earthquake activity approximately 35 mi (56 km) northeast of the BBNPP site, between Lackawanna and Wyoming Counties; the closest fault-related epicentral location (the Anthracite Zone) is greater than 25 mi (40 km) away. No earthquake within 50 mi (80 km) of the SSES site has been large enough to cause significant damage in the time the region has been populated, approximately 270 years.
- Approximately 250 exploratory boreholes were drilled at the SSES Units 1 and 2 site area. Borehole data have provided evidence for the lateral continuity of strata across the existing SSES site area (SSES FSAR, 2003). The inspection of soil samples has revealed no adverse effects indicative of geologically recent or active faulting.
- Field reconnaissance of many surface outcrops at the site and within the 25 mi (40 km) radius of the site, coupled with geophysical surveys, provided no evidence for faulting at the SSES site.
- At the time of the original studies for the SSES FSAR (SSES FSAR, 2003), published maps showing bedrock faults within a 5 mi (8 km) radius of the SSES site identified only the Light Street fault. The closest significant bedrock faults mapped prior to 1975 were faults located about 80 mi (129 km) southwest of the SSES site near Lewistown, PA (SSES FSAR, 2003).

#### **2.5.3.1.4.2 Published Geologic Mapping**

Since the late 1960s, extensive mapping of the BBNPP site region within the Ridge and Valley Province has been performed by the Pennsylvania Geological Survey (PGS) and other governmental agencies to improve knowledge of the Ridge and Valley stratigraphy and other geologic structures within the region (Inners, 1978; USGS, 2001; Wheeler, 2006). Local mapping includes geologic mapping across the BBNPP site area (Inners, 1978) (Figure 2.5-88), a developed geologic section through the central Appalachian Basin (Ryder, 1992) (Figures 2.5-89, 2.5-90, and 2.5-91), and a Precambrian Basement Map (Gold, 2005) based on borehole and seismic reflection data (Figure 2.5-92). This compilation of previous mapping and exploration studies provides the principal basis for the few bedrock faults recognized within the site area.

A local geologic section, oriented north-south within the site area (5 mi (8 km) radius), depicts slightly faulted anticlinal Silurian-Mississippian bedrock that is unconformably overlain by unconsolidated Pliocene-Holocene deposits, as shown in Figure 2.5-88a (DCNR, 2007; Inners, 1978). A review of the SSES FSAR reported the presence of the Light Street Fault but failed to uncover evidence, through either published reports or field investigations, to support the

existence of the inferred Berwick fault (SSES FSAR, 2003). Folds, as reported by Inners (Inners, 1978), are prevalent structures throughout the bedrock of the BBNPP site, mainly in second- and third-order. The major structure of the area is the Berwick Anticlinorium, a moderately complex, first order fold that passes through the center of Figure 2.5-88 (Inners, 1978).

However, the most detailed section of the site area was created by Inners (Inners, 1978) as part of a study conducted on behalf of the Commonwealth of Pennsylvania Department of Environmental Resources-Bureau of Topographic and Geologic Survey. This geologic cross-section, Figure 2.5-88a, was developed extending from just north of Lee Mountain, northwest of the BBNPP site, to near Black Creek, just south of Nescopeck Mountain, south of the BBNPP site. This section depicts moderately dipping, un-deformed geologic contacts between the Middle Devonian Mahantango Formation, the overlying Middle Devonian Harrell Formation, and underlying Marcellus Formation.

As shown on Figure 2.5-88 and Figure 2.5-94, the Light Street Fault (DCNR, 2007) and the Berwick Anticlinorium (Inners, 1978) have been mapped within the 5 mi (8 km) radius of the BBNPP site. In addition, two other structures have been proposed within the 5 mi (8 km) radius of the site, the Lackawanna Synclinorium (Inners, 1978), approximately 4 mi (6.4 km) northwest of the BBNPP site, and the inferred Berwick fault (Inners, 1978) (DCNR, 2007), approximately 3 mi (4.8 km) southwest of the BBNPP site. All of these structural features are consistent with published evidence of the intense folding and faulting that occurred to the bedrock formations during the Alleghanian Orogeny, as discussed in Section 2.5.1 (Faill, 1999; Harper, 1999; Way, 1999). The Light Street fault, inferred Berwick fault (Inners, 1978) (DCNR, 2007) and inferred folds (Inners, 1978) are described previously in Section 2.5.1. None of these features are considered capable tectonic sources, as defined in Appendix A of Regulatory Guide 1.165 (NRC, 1997).

Considering the evidence provided above, as well as the previous site investigations (SSES FSAR, 2003) discussed in Section 2.5.3.1.1, and the field reconnaissance interpretation undertaken for the BBNPP study, no evidence of Quaternary deformation has been reported in the literature or observed on the site for the Light Street or Berwick faults, or the Berwick Anticlinorium.

In summary, numerous investigations of the BBNPP site vicinity have found no syncline-related faulting in the basement directly beneath the BBNPP site area, and no signs of tectonic deformation within the exposed Quaternary deposits near the BBNPP site. Collectively, the published geologic information for the BBNPP site area, combined with regional geologic sections (Inners, 1978; Williams, 1987) and site and aerial reconnaissance, indicate the absence of Pleistocene and younger faulting and folding. A review of local and regional geologic sections (Figure 2.5-88a and Figure 2.5-96) suggest that the features, if present, are not prominent structures and are not developed within the Quaternary age units. In conclusion, there are no known tectonically active faults within the site area, and the Light Street fault and Berwick fault (if present) have been last active in the Late Permian (Inners, 1978).

#### **2.5.3.1.4.3 Recent Seismicity Compilations**

The USGS completed a compilation of all Quaternary faults, liquefaction features, and possible tectonic features in the eastern U.S. (Crone, 2000) (Wheeler, 2005) (Wheeler, 2006). These compilations do not show any Quaternary faults or features within a 25 mi (40 km) radius of the site, as shown in Figure 2.5-94. The nearest reported capable Quaternary feature (Crone, 2000) is the Cacoosing Valley earthquake (Number 7 on Figure 2.5-93), part of the Lancaster

seismic zone (Number 8), approximately 52 mi (84 km) south of the BBNPP site. The closest documented paleo-liquefaction site is known as the Newbury liquefaction features (Number 21), and is located over 260 mi (418 km) from the BBNPP site, in northeastern Massachusetts (Crone, 2000).

As was reported in Section 2.5.2, the locations of earthquakes were accounted for by an updated USGS catalog (USGS, 2008), for assessing the BBNPP site seismic hazard. This update is a refinement of the EPRI SOG catalog (EPRI, 1986) that brings the catalog of recorded seismic events current to the year 2007. Figure 2.5-87 (USGS, 2001) shows two earthquakes lying within the 25 mi (40 km) radius of the BBNPP site. Further analysis of these mapped earthquakes revealed that their sources were quarry blasts or mine collapses (Faill, 2004). Neither of these events was fault related or was associated with bedrock/ basement rock deformation. No publications reported vibratory ground motions being felt as a result of either of these seismic events. Based on this information, there are no significant hazard potential faults within a 25 mi (40 km) radius of the BBNPP site.

#### **2.5.3.2 Geological Evidence, or Absence of Evidence, for Surface Deformation**

Based on the discussions in the preceding sections, no evidence for surface deformation has been found for this site vicinity. The evidence against that deformation is as follows:

- No subsurface offsets within the site boundary, on the adjacent property for the SSES, or in the site vicinity where literature sources infer faults to exist.
- No capable faults mapped within 50 miles (80 km) of the site (at the local or regional scale).
- No seismic activity recorded within 50 miles (80 km) of the site.
- No evidence of prehistoric ground motion (liquefaction features) within the site area.

#### **2.5.3.3 Correlation of Earthquakes with Capable Tectonic Sources**

As presented in Figure 2.5-87, two earthquake epicenters occur within the 25 mi (40 km) radius of the BBNPP site, further analysis revealed that the sources for these earthquakes were quarry blasts or mine collapses (Faill, 2004). No reported historical earthquake epicenters have been associated with either of the bedrock faults (Light Street or Berwick faults) located within the 25 mi (40 km) radius of the BBNPP site vicinity. The distribution of earthquakes from the catalog supports the following conclusions:

- There are no earthquakes within the site vicinity that are associated with a known geologic or tectonic structure.
- The updated catalog of earthquake events (USGS, 2008) does not indicate a unique cluster of seismicity in the area that would suggest a local seismic source outside of the EPRI seismic source model (EPRI, 1986).
- The catalog does not show a pattern of seismicity that would require significant revision to the EPRI seismic source geometry.

#### **2.5.3.4 Ages of Most Recent Deformations**

As presented in Section 2.5.1.2.4, the local tectonic features located within 5 mi (8 km) of the BBNPP site (the Light Street fault, the inferred Berwick fault, the Berwick Anticlinorium, and the Lackawanna Synclinorium) do not exhibit evidence of Quaternary activity. They are interpreted to have formed during the Paleozoic Era as part of the regional Alleghanian Orogeny and have been inactive since that time (Inners, 1978). Based on a review of available published geologic literature, field reconnaissance, and interpretation of aerial photography, tectonic deformation associated with the local structures are constrained to the Late Permian Period, and do not affect Quaternary aged deposits (Inners, 1978) (Williams, 1987) (DCNR, 2007).

#### **2.5.3.5 Relationship of Tectonic Structures in the Site Area to Regional Tectonic Structures**

All four of the features evaluated within the 5 mi (8 km) radius of the BBNPP site (Light Street fault, Berwick fault, Berwick Anticlinorium, and Lackawanna Synclinorium) have been linked with regional tectonic events, mainly the Alleghanian Orogeny, and have been inactive since Late Permian time. Tectonic models hypothesize that the crystalline basement underlying the BBNPP site was accreted to a pre-Taconic North American margin during the Precambrian. Episodes of continental collisions have produced a series of accreted terrains that are separated, in part, by low-angle detachment faults, as discussed in detail in Section 2.5.1.1.2. In association with these continental collisions, the Paleozoic bedrocks of eastern North America, including the Ridge and Valley Province, contain a number of generally northeast striking thrust faults (Schlische, 2003) such as the Light Street fault and the inferred Berwick fault. The Berwick Anticlinorium and Lackawanna Synclinorium are both results of regional extension and compression due to the orogenic events discussed in detail in Section 2.5.1.1.2.

#### **2.5.3.6 Characterization of Capable Tectonic Sources**

None of the mapped bedrock structures located within the Site vicinity is considered to be a capable tectonic source as defined by Regulatory Guide 1.208 (NRC, 2007). This conclusion is based on the following observations:

- The Light Street and the inferred Berwick faults are inactive, and were most recently active in the Late Permian, approximately 250 Ma.
- There was no deformation of Quaternary age deposits identified in either the literature review or during the site investigations for this study.
- There is no association of historical earthquakes with any local fault or any cluster of seismicity within the site area that would indicate the presence of a capable source.

#### **2.5.3.7 Designation of Zones of Quaternary Deformation Requiring Detailed Fault Investigation**

There are no zones of Quaternary deformation requiring detailed investigation within the BBNPP site area. A review and interpretation of digital elevation models coupled with aerial reconnaissance identified few discontinuous north to northeast-striking lineaments. None of these lineaments are interpreted as fault-related, or are coincident with the Light Street fault or the inferred Berwick fault.



#### **2.5.3.8 Potential for Tectonic or Non-Tectonic Deformation at the Site**

The potential for tectonic deformation at the BBNPP site is negligible. This conclusion is based on the following:

1. No evidence for faulting in the subsurface investigation of the BBNPP site.
2. No evidence of faulting in either the detailed subsurface investigation at the adjacent SSES site, or in outcrops within the site area.
3. No evidence of paleoliquefaction found at the site or surrounding site area.
4. No historical seismic activity or mapped capable faults within the site vicinity.

Collectively, these data support the interpretation for the absence of any Quaternary surface faults or capable tectonic sources within the BBNPP site area. In addition, there is no evidence of non-tectonic deformation at the site, such as glacially induced faulting, collapse structures, growth faults, salt migration, or volcanic intrusion.

### 2.5.3.9 References

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**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
6.3	-70.1	47.7	0	1534	1	1	0	0	0	DNAG
6.5	-71.8	44.4	0	1638	6	11	19	0	0	Ebel
7	-70.1	47.6	0	1663	2	5	22	30	0	Ebel
3.5	-71.5	42.5	0	1668	12	19	0	0	0	Ebel
3.5	-70.8	42.8	0	1685	2	18	21	0	0	Ebel
3	-70.8	42	0	1697	2	20	11	15	0	Ebel
3.3	-73.5	41.4	0	1702	1	1	0	0	0	NCEER
3.3	-71.1	42.4	0	1705	6	27	0	0	0	NCEER
5.1	-70.6	42.8	0	1727	11	10	3	40	0	NCEER
3.3	-70.6	42.8	0	1728	5	16	0	0	0	NCEER
3.3	-70.6	42.8	0	1728	7	30	15	0	0	NCEER
3.5	-70.6	42.8	0	1729	2	10	14	0	0	NCEER
3.3	-70.6	42.8	0	1729	3	30	19	0	0	NCEER
3.3	-73.5	41.4	0	1729	8	6	0	0	0	NCEER
3.3	-70.6	42.8	0	1729	9	19	20	30	0	NCEER
3.3	-70.6	42.8	0	1729	10	10	21	30	0	NCEER
3.3	-70.6	42.8	0	1729	11	25	13	0	0	NCEER
3.3	-70.6	42.8	0	1730	2	20	1	0	0	NCEER
3.3	-70.6	42.8	0	1730	4	24	1	0	0	NCEER
3.3	-70.6	42.8	0	1730	12	7	1	20	0	NCEER
3.3	-70.6	42.8	0	1731	1	13	0	0	0	NCEER
3.3	-70.6	42.8	0	1731	7	16	10	0	0	NCEER
3.3	-70.6	42.8	0	1731	10	13	4	0	0	NCEER
3.3	-70.6	42.8	0	1732	2	19	0	0	0	NCEER
5.8	-73.6	45.5	0	1732	9	16	16	0	0	NCEER
3.3	-70.6	42.8	0	1734	11	23	5	0	0	NCEER
3.3	-70.6	42.8	0	1736	2	13	22	45	0	NCEER
3.3	-70.6	42.8	0	1736	10	12	6	30	0	NCEER
3.3	-70.6	42.8	0	1736	11	23	7	0	0	NCEER
3.3	-71	42.4	0	1737	2	17	21	30	0	NCEER
3.3	-70.6	42.8	0	1737	9	20	15	20	0	NCEER
5.2	-74	40.8	0	1737	12	19	3	45	0	NCEER
3.3	-70.6	42.8	0	1739	8	13	7	30	0	NCEER
3.3	-70.6	42.8	0	1741	2	5	20	50	0	NCEER
3.3	-71.2	46.8	0	1744	5	27	0	0	0	NCEER
4.6	-70.9	42.5	0	1744	6	14	15	15	0	NCEER
3.6	-76.3	40	0	1752	12	17	23	30	0	SRA
5.8	-70.3	42.7	0	1755	11	18	9	12	0	NCEER
3.3	-71.1	42.3	0	1757	7	8	19	15	0	NCEER
3.5	-76.5	38.9	0	1758	4	25	2	30	0	NCEER
3.3	-71	42.35	0	1759	2	2	7	0	0	NCEER
4.3	-71	42.5	0	1761	3	12	7	15	0	NCEER
3.3	-71.5	43.1	0	1761	11	2	1	0	0	NCEER
3.5	-66	45.3	0	1764	9	30	0	0	0	NCEER
3.3	-70.3	43.7	0	1766	1	23	10	0	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

(Page 2 of 65)

$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.3	-70.77	43.07	0	1766	3	2	8	0	0	NCEER
3.3	-70.8	43.1	0	1766	12	17	11	48	0	NCEER
3.3	-70.3	43.7	0	1769	10	19	0	0	0	NCEER
4.6	-77.4	37.2	0	1774	2	21	19	0	0	NCEER
3.3	-78.8	37.7	0	1775	3	16	19	15	0	NCEER
4	-82	39.9	0	1776	1	1	0	0	0	NCEER
3.3	-83	35.2	0	1776	11	5	0	0	0	NCEER
3.3	-84	36	0	1777	11	16	7	0	0	NCEER
3.8	-87.2	30.4	0	1780	2	6	0	0	0	NCEER
3.3	-70.9	42.5	0	1780	11	29	0	0	0	NCEER
4.9	-74.5	41	0	1783	11	30	3	50	0	NCEER
3.3	-71.2	46.8	0	1784	1	2	10	0	0	NCEER
3.3	-78.8	37.7	0	1791	1	13	9	0	0	NCEER
3.3	-77.5	37.5	0	1791	1	15	10	0	0	NCEER
4.5	-72.4	41.5	0	1791	5	16	13	22	0	NCEER
6	-70.5	47.4	0	1791	12	6	20	0	0	NCEER
3.3	-72.5	41.5	0	1792	8	29	3	0	0	NCEER
3.3	-72.5	41.5	0	1794	3	6	19	0	0	NCEER
3.4	-89.9	39	0	1795	1	8	9	0	0	NCEER
4	-79	42.9	0	1796	12	26	11	0	0	NCEER
4.4	-80	32.9	0	1799	4	11	8	20	0	NCEER
3.9	-76.39	40.12	0	1800	11	20	5	0	0	NCEER
3.3	-72.3	43.7	0	1800	12	20	0	0	0	NCEER
3.3	-71.1	41.9	0	1800	12	25	0	0	0	NCEER
3.3	-70.8	43.1	0	1801	3	1	20	30	0	NCEER
3.5	-79.1	37.4	0	1802	8	23	10	0	0	NCEER
3.3	-70.9	42.5	0	1803	1	18	14	50	0	NCEER
4.4	-87.8	42	0	1804	8	20	20	10	0	USHIS
4.2	-89	42	0	1804	8	24	20	10	0	NCEER
3.3	-70.9	42.5	0	1805	4	25	0	0	0	NCEER
3.3	-69	44.5	0	1805	6	12	12	30	0	NCEER
3.3	-72.5	41.5	0	1805	12	30	11	0	0	NCEER
3.3	-71.1	43	0	1807	1	14	4	0	0	NCEER
3.5	-79.1	37.4	0	1807	5	1	9	0	0	NCEER
3.3	-70.5	43.5	0	1807	5	6	18	0	0	NCEER
3.5	-69	44.4	0	1808	6	26	2	50	0	NCEER
3.9	-70.9	43	0	1810	11	10	2	15	0	NCEER
3.3	-80.2	36.1	0	1811	11	27	8	0	0	NCEER
3.3	-77.4	37.6	0	1812	2	2	9	30	0	NCEER
7.4	-89.6	36.5	0	1812	2	7	9	45	0	NCEER <sup>1</sup>
3.3	-77.5	37.5	0	1812	4	22	4	0	0	NCEER
4	-70.3	43.7	0	1814	11	29	0	14	0	NCEER
3	-89.5	36.6	0	1816	7	25	15	0	0	NCEER
5.2	-73.6	45.5	0	1816	9	9	0	0	0	NCEER
5	-80	32.9	0	1817	1	8	9	0	0	USHIS

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

(Page 3 of 65)

$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
4.7	-67.2	45	0	1817	5	22	20	0	0	NCEER
4.2	-71.2	42.5	0	1817	10	5	16	45	0	NCEER
3.1	-84.5	38.5	0	1817	12	11	0	0	0	NCEER
3	-90.2	38.6	0	1818	4	11	20	0	0	NCEER
3.3	-71.2	46.9	0	1818	10	11	0	0	0	NCEER
3.3	-76.5	44	0	1818	12	7	0	0	0	NCEER
3.4	-89.7	37.7	0	1819	9	2	8	0	0	NCEER
3.1	-89.8	38.1	0	1819	9	17	4	0	0	NCEER
3	-89.5	36.6	0	1820	1	1	0	0	0	NCEER
3.3	-79.3	33.4	0	1820	9	3	8	30	0	NCEER
3.4	-89.5	37.3	0	1820	11	9	22	0	0	NCEER
3.8	-68.8	44.8	0	1821	5	5	12	30	0	NCEER
3.3	-70	43.9	0	1823	3	7	15	0	0	NCEER
3.5	-68.8	44.8	0	1823	6	10	17	0	0	NCEER
3.8	-70.6	42.9	0	1823	7	23	11	55	0	NCEER
3.5	-66.5	46.5	0	1824	7	9	0	0	0	NCEER
4.1	-80.5	39.7	0	1824	7	15	16	20	0	NCEER
3.3	-81.56	30.08	0	1826	2	26	14	0	0	NCEER
3.3	-81.2	36.1	0	1827	5	11	0	0	0	NCEER
4.8	-87.5	38	0	1827	7	5	11	30	0	NCEER
4.8	-88	38	0	1827	8	7	4	30	0	NCEER
4	-85.8	38.3	0	1827	8	7	7	0	0	NCEER
3.3	-72.1	41.4	0	1827	8	23	0	0	0	NCEER
4.8	-80	37	0	1828	3	9	0	0	0	NCEER
3.3	-70	43.9	0	1828	7	25	11	0	0	NCEER
3.3	-69.8	44.2	0	1829	8	27	21	45	0	NCEER
5.2	-70.5	47.3	0	1831	5	8	0	0	0	NCEER
5	-70.1	47.6	0	1831	7	14	0	0	0	NCEER
4	-85.6	42.3	0	1833	2	4	0	0	0	NCEER
3.3	-70.17	47.65	0	1833	3	1	0	0	0	NCEER
3.3	-70.2	47.7	0	1833	4	1	0	0	0	NCEER
4.6	-78	37.7	0	1833	8	27	11	0	0	NCEER
3.8	-76.14	39.85	0	1834	2	5	22	30	0	NCEER
3.4	-86	38	0	1834	11	20	19	40	0	NCEER
3.1	-81.7	41.5	0	1836	7	9	2	15	0	NCEER
3.3	-70.9	42.5	0	1837	1	15	7	0	0	NCEER
3.5	-72.7	41.7	0	1837	4	12	0	0	0	NCEER
5	-88	38.5	0	1838	6	9	14	45	0	NCEER
3.1	-83.8	38.6	0	1839	9	5	0	0	0	NCEER
3.7	-75	43	0	1840	1	16	20	0	0	NCEER
3.7	-72.9	41.5	0	1840	8	9	20	30	0	NCEER
4	-79.85	43.2	0	1840	9	10	0	0	0	NCEER
3.3	-74.25	40.79	0	1841	1	25	5	30	0	NCEER
4.2	-89.2	36.6	0	1841	12	28	5	50	0	NCEER
3.1	-89.2	36.6	0	1842	5	28	5	0	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.4	-89.2	36.6	0	1842	11	4	6	30	0	NCEER
3.8	-73.2	46	0	1842	11	9	0	0	0	NCEER
5.4	-89.6	35.5	0	1843	1	5	2	45	0	NCEER
3.1	-90.3	38.8	0	1843	2	16	0	0	0	NCEER
4.4	-90.5	35.5	0	1843	2	17	5	0	0	NCEER
3.3	-72.5	44.4	0	1843	3	14	0	0	0	NCEER
4.1	-87.1	35.6	0	1843	8	9	0	0	0	NCEER
3.3	-71.2	41.1	0	1843	10	24	0	0	0	NCEER
3.5	-78.33	43.05	0	1844	10	22	7	0	0	NCEER
3.3	-73.6	45.5	0	1844	11	1	0	0	0	NCEER
3.6	-83.27	35.79	0	1844	11	28	8	0	0	NCEER
3.8	-73.67	41.22	0	1845	10	26	23	15	0	NCEER
3.3	-70.3	42.7	0	1846	5	30	18	30	0	NCEER
4.1	-70.8	42.5	0	1846	8	25	9	45	0	NCEER
3.3	-69.1	44.2	0	1847	2	2	0	0	0	NCEER
4.2	-70.1	41.7	0	1847	8	8	15	0	0	NCEER
3.7	-82.53	39.65	0	1848	4	6	0	0	0	NCEER
4.4	-73.85	41.11	0	1848	9	9	0	0	0	NCEER
3.4	-89.2	36.6	0	1849	1	24	0	0	0	NCEER
4.3	-88	37	0	1850	4	5	2	5	0	NCEER
3.1	-81.7	41.5	0	1850	10	1	10	25	0	NCEER
3.3	-78.4	37.3	0	1850	10	17	0	0	0	NCEER
3.3	-69.6	44.6	0	1851	1	4	4	30	0	NCEER
3.3	-71.4	41.2	0	1852	1	10	11	40	0	NCEER
4.9	-81.6	36.6	0	1852	4	29	18	0	0	NCEER
3.3	-82	33.48	0	1852	8	25	2	40	0	NCEER
4.4	-78.6	37.6	0	1852	11	2	23	35	0	NCEER
3.7	-70.9	43	0	1852	11	28	4	45	0	NCEER
4.5	-75.5	43.7	0	1853	3	12	7	0	0	NCEER
3.3	-79.4	43.1	0	1853	3	13	10	0	0	NCEER
4.4	-79.5	38.5	0	1853	5	2	14	20	0	NCEER
4	-81.96	33.49	0	1853	5	20	5	10	0	NCEER
3.3	-70.2	43.5	0	1853	7	17	10	30	0	NCEER
3.3	-71.9	43	0	1853	11	28	0	0	0	NCEER
4.1	-89.2	36.6	0	1853	12	12	0	0	0	NCEER
3	-83.8	37.2	0	1854	2	12	0	0	0	NCEER
3.6	-84	37.6	0	1854	2	28	0	0	0	NCEER
3.1	-85.2	38.2	0	1854	3	8	0	0	0	NCEER
3.3	-83.62	32.82	0	1854	3	20	6	15	0	NCEER
3.3	-72.3	42.9	0	1854	10	24	0	0	0	NCEER
3.4	-70.8	43	0	1854	12	11	5	30	0	NCEER
4	-71	44	0	1855	1	16	23	0	0	NCEER
3.9	-78.6	37	0	1855	2	2	8	0	0	NCEER
3.3	-69.6	44.6	0	1855	2	19	0	0	0	NCEER
3.1	-89.2	37	0	1855	5	3	3	33	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.3	-71.6	44.7	0	1855	5	29	10	0	0	NCEER
3.3	-65.5	44.7	0	1855	6	1	0	0	0	NCEER
3.3	-73.7	43.3	0	1855	12	17	19	0	0	NCEER
3.3	-78.2	39.2	0	1856	1	16	8	0	0	NCEER
3.3	-72.6	41.4	0	1856	3	13	3	0	0	NCEER
4.1	-89.5	36.6	0	1856	11	9	10	0	0	NCEER
3.1	-74.75	40.08	0	1857	2	10	23	30	0	NCEER
3.9	-81.05	42.22	0	1857	2	27	20	30	0	NCEER
3.3	-80.6	41.8	0	1857	3	1	1	40	0	NCEER
3.4	-83.18	33	0	1857	3	1	22	45	0	NCEER
5.1	-89.2	38.7	0	1857	10	8	10	0	0	NCEER
4	-78.97	42.74	0	1857	10	23	20	15	0	NCEER
3.3	-68	46.7	0	1857	12	8	20	0	0	NCEER
4.1	-80.73	32.78	0	1857	12	19	8	50	0	NCEER
3.9	-70.2	44.1	0	1857	12	23	18	30	0	NCEER
3.3	-78.5	42.9	0	1858	1	1	7	0	0	NCEER
3.1	-81.3	41.7	0	1858	4	10	11	30	0	NCEER
3.3	-72.1	45.5	0	1858	5	17	20	0	0	NCEER
3.3	-73	41.3	0	1858	7	1	3	45	0	NCEER
4	-89.2	36.5	0	1858	9	21	0	0	0	NCEER
3.1	-81.5	37.1	0	1859	3	22	0	0	0	NCEER
4.5	-94.8	46	0	1860	1	1	0	0	0	NCEER
4.3	-80.57	33.68	0	1860	1	19	18	0	0	NCEER
3.3	-70.5	42.2	0	1860	3	17	2	30	0	NCEER
4	-83.64	35.08	0	1860	4	24	20	0	0	NCEER
4.3	-87.5	37.5	0	1860	8	7	15	30	0	NCEER
6	-70.1	47.5	0	1860	10	17	11	15	0	NCEER
3.2	-82.64	34.13	0	1860	10	22	5	0	0	NCEER
4.3	-83.36	35.09	0	1861	1	3	16	30	0	NCEER
5	-75.4	45.4	0	1861	7	12	0	0	0	NCEER
5.2	-82.3	36.18	0	1861	8	31	5	0	0	NCEER
3.5	-73.7	45.6	0	1861	10	1	0	0	0	NCEER
3.3	-72.5	41.5	0	1862	2	3	1	0	0	NCEER
3.3	-73	44.5	0	1863	6	9	21	30	0	NCEER
3.8	-71.2	46.9	0	1864	4	20	18	15	0	NCEER
3.3	-73.6	45.5	0	1864	10	21	9	10	0	NCEER
4.6	-89.5	36.5	0	1865	8	17	15	0	0	NCEER
3.3	-71.2	46.8	0	1866	11	9	16	10	0	NCEER
5.2	-96.3	39.2	0	1867	4	24	20	22	0	NCEER
3.1	-95.8	40.7	0	1867	4	28	0	0	0	NCEER
4.7	-75.15	44.65	0	1867	12	18	8	0	0	NCEER
3.3	-80	32.9	0	1869	1	1	0	0	0	NCEER
3.4	-84.5	38.1	0	1869	2	20	0	0	0	NCEER
5.1	-67.2	45	0	1869	10	22	11	0	0	NCEER
4	-70.5	47.5	0	1869	12	1	0	0	0	NCEER



**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

(Page 6 of 65)

$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.3	-69.8	44.1	0	1870	2	8	0	0	0	NCEER
3.3	-66.5	45.5	0	1870	3	17	11	0	0	NCEER
6.5	-70.5	47.4	0	1870	10	20	16	30	0	NCEER
3	-89.2	36.6	0	1870	12	14	0	0	0	NCEER
3.5	-74.6	45.6	0	1871	1	3	0	0	0	NCEER
3.3	-71.2	46.8	0	1871	5	20	7	0	0	NCEER
3.3	-71.5	43.2	0	1871	7	20	0	0	0	NCEER
3.4	-90	38.5	0	1871	7	25	6	40	0	NCEER
3.8	-75.5	39.7	0	1871	10	9	14	40	0	NCEER
3.4	-83.8	43.5	0	1872	2	6	14	0	0	NCEER
3	-89.2	37	0	1872	2	8	11	0	0	NCEER
3.7	-78	37.7	0	1872	6	5	3	0	0	NCEER
3	-83.22	33.06	0	1872	6	17	14	30	0	NCEER
3.1	-93.5	39.8	0	1872	7	9	2	30	0	NCEER
3	-73.8	40.9	0	1872	7	11	10	25	0	NCEER
3.6	-97	42.7	0	1872	10	9	16	0	0	NCEER
3.4	-71.6	43.2	0	1872	11	18	19	0	0	NCEER
3	-84.2	39.7	0	1873	4	23	4	14	0	NCEER
3.5	-74.2	44.8	0	1873	4	25	19	0	0	NCEER
3.1	-79.9	43.3	0	1873	4	30	0	0	0	NCEER
3.3	-74.7	45	0	1873	4	30	0	0	0	NCEER
3	-97.7	30.2	0	1873	5	1	4	30	0	NCEER
3.7	-89.6	36	0	1873	5	3	21	0	0	NCEER
4	-78.94	42.69	0	1873	7	6	0	0	0	NCEER
3.3	-73.2	45.5	0	1873	9	30	11	50	0	NCEER
3.3	-76	46.5	0	1873	9	30	11	50	0	NCEER
3.6	-78.2	37.2	0	1873	10	3	12	45	0	NCEER
3.2	-83.9	33	0	1873	10	4	0	0	0	NCEER
3.3	-71.2	43.6	0	1874	1	6	0	0	0	NCEER
3.3	-71.4	42.6	0	1874	1	25	17	0	0	NCEER
3.3	-71.5	43	0	1874	1	26	7	0	0	NCEER
3.3	-82.1	35.7	0	1874	2	22	0	0	0	NCEER
3.7	-67.28	45.18	0	1874	2	28	3	40	0	NCEER
3.3	-82.1	35.7	0	1874	3	17	0	0	0	NCEER
3.3	-82.1	35.7	0	1874	4	14	0	0	0	NCEER
3	-89.2	37	0	1874	7	9	22	0	0	NCEER
3.3	-69.1	48.6	0	1874	7	31	9	0	0	NCEER
3	-70.9	42.7	0	1874	11	24	0	0	0	NCEER
3.4	-73.8	40.9	0	1874	12	11	3	25	0	NCEER
3.4	-82.51	35.29	0	1875	4	10	0	0	0	NCEER
4.7	-84	40.2	0	1875	6	18	13	43	0	NCEER
3.4	-73	41.9	0	1875	7	28	9	10	0	NCEER
4.1	-89.6	36.1	0	1875	10	7	0	0	0	NCEER
3.1	-90	35.1	0	1875	10	28	3	0	0	NCEER
4.8	-82.9	33.49	0	1875	11	1	22	30	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.8	-95.7	39	0	1875	11	8	10	40	0	NCEER
3.6	-84	36	0	1875	11	12	7	0	0	NCEER
3.3	-72.3	42.9	0	1875	12	1	9	0	0	NCEER
4.8	-78.5	37.6	0	1875	12	23	4	45	0	NCEER
3.4	-84.2	40.4	0	1876	6	1	0	0	0	NCEER
3.1	-99.6	44.1	0	1876	8	17	5	25	0	NCEER
3.6	-71.3	41.5	0	1876	9	22	4	30	0	NCEER
4.7	-87	38.5	0	1876	9	25	6	15	0	NCEER
3.3	-80	32.9	0	1876	12	12	0	0	0	NCEER
3.3	-77.5	37.4	0	1876	12	23	4	45	0	NCEER
3.4	-83.5	38.8	0	1877	1	23	21	0	0	NCEER
3.6	-84	36	0	1877	5	25	0	0	0	NCEER
3	-87.9	38.2	0	1877	5	26	21	0	0	NCEER
4.2	-89.7	36.8	0	1877	7	15	0	40	0	NCEER
3.1	-83.3	42.3	0	1877	8	17	16	50	0	NCEER
3.3	-74.9	40.3	0	1877	9	10	14	59	0	NCEER
4.7	-73.9	45.2	0	1877	11	4	0	0	0	NCEER
5	-97	41	0	1877	11	15	17	45	0	NCEER
3.7	-84	35.5	0	1877	11	16	7	38	0	NCEER
3	-89.2	37	0	1877	11	19	11	10	0	NCEER
3.5	-76.85	45.7	0	1877	12	18	10	0	0	NCEER
3	-89.2	37	0	1878	1	9	4	30	0	NCEER
3.9	-89.1	36.8	0	1878	3	12	10	0	0	NCEER
3.4	-74	41.5	0	1878	10	4	7	30	0	NCEER
5.2	-90.7	35.5	0	1878	11	19	5	52	0	NCEER
3.3	-99.1	39.6	0	1879	3	1	0	0	0	NCEER
3.3	-75.5	39.2	0	1879	3	26	0	30	0	NCEER
3.3	-73.6	45.6	0	1879	6	11	0	0	0	NCEER
3.3	-79.2	43.2	0	1879	8	21	8	0	0	NCEER
3.7	-90.3	35.3	0	1879	9	26	3	10	0	NCEER
3.3	-71.5	43	0	1879	10	26	2	30	0	NCEER
3.1	-81.08	34.37	0	1879	10	26	20	0	0	NCEER
3.6	-97.3	42.9	0	1879	12	29	6	30	0	NCEER
3.4	-71	42.7	0	1880	5	12	12	45	0	NCEER
3.3	-75.3	45.2	0	1880	5	31	0	0	0	NCEER
3.7	-90.3	35.3	0	1880	7	14	2	30	0	NCEER
3.3	-71.5	43	0	1880	7	20	0	0	0	NCEER
3.3	-73.8	45.2	0	1880	9	6	5	30	0	NCEER
3.3	-70.5	47.45	0	1880	11	28	13	30	0	NCEER
3	-97.2	49	0	1880	12	28	7	15	0	NCEER
3.3	-70	44	0	1881	1	21	2	40	0	NCEER
3.1	-85.8	41.6	0	1881	4	20	0	0	0	NCEER
4	-89.1	41.3	0	1881	5	27	0	0	0	NCEER
3.3	-70.2	47.6	0	1881	10	1	6	40	0	NCEER
3.3	-71.6	43.2	0	1881	10	6	5	3	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

(Page 8 of 65)

$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.1	-90	35.1	0	1881	10	7	16	52	0	NCEER
3.1	-84.2	40.4	0	1882	2	9	20	0	0	NCEER
3.3	-71.7	43.2	0	1882	4	17	0	0	0	NCEER
3.4	-89.2	36.9	0	1882	7	20	10	0	0	NCEER
3.9	-90.6	37.6	0	1882	7	28	0	0	0	NCEER
3.3	-67.4	49.3	0	1882	8	15	15	30	0	NCEER
4.4	-89.5	38.7	0	1882	9	27	10	20	0	NCEER
4.8	-95.6	33.6	0	1882	10	22	22	15	0	NCEER
6.2	-105.5	40.5	0	1882	11	8	1	30	0	USHIS
3.1	-79.25	43	0	1882	11	27	23	30	0	NCEER
3.6	-71.4	43.2	0	1882	12	19	22	24	0	NCEER
3.5	-67	45	0	1883	1	1	2	55	0	NCEER
3.3	-67.7	44.6	0	1883	1	1	7	58	0	NCEER
4.7	-88.5	37	0	1883	1	11	7	12	0	NCEER
4.7	-85.6	42.3	0	1883	2	4	11	0	0	NCEER
3.3	-71.2	43.6	0	1883	2	4	20	5	0	NCEER
3.6	-71.3	41.5	0	1883	2	28	3	30	0	NCEER
3.1	-76.4	39.5	0	1883	3	11	23	57	0	NCEER
3.3	-74.5	45.1	0	1883	3	12	0	0	0	NCEER
4.5	-89.2	37	0	1883	4	12	8	30	0	NCEER
3.1	-82.6	38.4	0	1883	5	23	4	30	0	NCEER
4	-90	35.1	0	1883	6	11	18	16	0	NCEER
3.7	-89.1	37	0	1883	7	14	7	30	0	NCEER
3.1	-90.2	38.7	0	1883	11	15	3	14	0	NCEER
4	-91.2	35.7	0	1883	12	5	15	20	0	NCEER
3.3	-71.7	43.2	0	1884	1	18	7	0	0	NCEER
3.7	-77.59	34.59	0	1884	1	18	8	0	0	NCEER
3.1	-100.7	41.1	0	1884	3	17	20	0	0	NCEER
3.6	-83.05	32.8	0	1884	3	21	4	30	0	NCEER
5.2	-74	40.6	0	1884	8	10	19	7	0	NCEER
3.3	-83.83	36.07	0	1884	8	24	19	45	0	NCEER
4.8	-84.1	40.7	0	1884	9	19	20	14	0	NCEER
3.7	-71.7	43.2	0	1884	11	23	5	30	0	NCEER
3.5	-89.7	35.5	0	1884	11	30	5	0	0	NCEER
3.3	-71.5	43.7	0	1884	12	17	7	0	0	NCEER
3.9	-77.5	39.2	0	1885	1	3	2	12	0	NCEER
3.4	-73.9	41.3	0	1885	1	4	11	6	0	NCEER
3.5	-81.6	41.16	0	1885	1	18	10	30	0	NCEER
3.3	-81.1	36.9	0	1885	2	2	12	10	0	NCEER
3.3	-66.1	45.1	0	1885	6	1	0	0	0	NCEER
3.4	-81.83	36.12	0	1885	8	6	9	0	0	NCEER
4.4	-78.8	37.7	0	1885	10	10	4	35	0	NCEER
3.6	-82.71	33.17	0	1885	10	17	18	20	0	NCEER
3.3	-71.5	42.9	0	1886	1	6	0	10	0	NCEER
3.3	-73.8	41.6	0	1886	1	25	0	4	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

(Page 9 of 65)

$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.5	-88	32.8	0	1886	2	5	1	0	0	NCEER
3.3	-81.52	35.93	0	1886	2	5	2	0	0	NCEER
3	-85.5	39	0	1886	3	1	16	0	0	NCEER
4.2	-89.2	37	0	1886	3	18	5	59	0	NCEER
3.8	-82.24	39.36	0	1886	5	3	3	0	0	NCEER
3.3	-67.4	49.3	0	1886	5	18	19	30	0	NCEER
3.3	-74	46	0	1886	8	12	0	0	0	NCEER
3	-86.1	39.7	0	1886	8	14	0	0	0	NCEER
3.8	-80.14	33.38	0	1886	8	27	8	30	0	NCEER
3.3	-81.7	30.4	0	1886	9	1	0	0	0	NCEER
6.8	-80	32.9	0	1886	9	1	2	51	0	NCEER
4.4	-81.94	33.93	0	1886	9	1	8	55	0	NCEER
4.2	-82.86	34.3	0	1886	9	1	9	45	0	NCEER
4.3	-81.56	33.41	0	1886	9	1	23	50	0	NCEER
3.6	-81.23	34.72	0	1886	9	2	23	0	0	NCEER
3.3	-72.5	41.5	0	1886	9	5	0	0	0	NCEER
4.1	-80.97	33.05	0	1886	9	7	11	42	0	NCEER
4.3	-80.68	32.7	0	1886	9	7	17	0	0	NCEER
3.7	-80.96	32.23	0	1886	9	19	0	0	0	NCEER
3.2	-80.05	36.7	0	1886	9	25	2	0	0	NCEER
3.4	-81.62	34.7	0	1886	9	27	22	0	0	NCEER
4.1	-81.66	34.71	0	1886	10	22	0	0	0	NCEER
5.2	-81.01	33.87	0	1886	10	22	14	45	0	NCEER
3.9	-80.39	33.9	0	1886	10	31	14	20	0	NCEER
5.3	-80.42	33.4	0	1886	11	5	12	25	0	NCEER
3.8	-81.06	33.04	0	1886	12	2	2	20	0	NCEER
3.6	-82.06	34.18	0	1886	12	11	16	0	0	NCEER
3.6	-97.06	30.15	0	1887	1	5	17	57	0	SRA
3.1	-82.42	34.35	0	1887	1	12	6	0	0	NCEER
4.5	-88.5	39	0	1887	2	6	22	15	0	NCEER
3.1	-80	45.35	0	1887	2	19	0	0	0	NCEER
3.3	-67.4	49.3	0	1887	2	22	22	59	0	NCEER
3.3	-70.5	47.5	0	1887	3	11	0	0	0	NCEER
3.7	-80.37	33.9	0	1887	5	22	20	45	0	NCEER
4	-70.5	47.45	0	1887	5	27	6	15	0	NCEER
3.6	-81	34	0	1887	6	3	8	45	0	NCEER
3.3	-71.5	43.2	0	1887	6	30	22	0	0	NCEER
3.7	-80.77	33.74	0	1887	7	10	13	5	0	NCEER
3.8	-78.83	34.41	0	1887	8	2	1	0	0	NCEER
4.6	-88.5	37.2	0	1887	8	2	18	36	0	NCEER
3.9	-79.95	33.83	0	1887	8	10	7	1	0	NCEER
4.2	-80.62	32.49	0	1887	8	26	23	45	0	NCEER
4	-81.22	33.52	0	1887	8	27	4	56	0	NCEER
3.9	-80.86	33.66	0	1887	8	28	22	57	0	NCEER
3.3	-77.1	45.8	0	1888	1	11	9	0	0	NCEER

Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
4.4	-80.17	34.18	0	1888	1	12	9	55	0	NCEER
3.3	-70.1	44.65	0	1888	2	1	0	0	0	NCEER
3	-82.5	36.4	0	1888	3	17	0	0	0	NCEER
3.3	-70.5	47.45	0	1888	4	19	5	30	0	NCEER
3.3	-70	44.3	0	1888	8	15	1	15	0	NCEER
3.1	-81.08	34.37	0	1888	8	15	18	30	0	NCEER
3.5	-83.35	33.42	0	1888	9	17	21	30	0	NCEER
3.1	-90.4	35.4	0	1888	11	3	0	0	0	NCEER
3.3	-68.7	48.5	0	1888	12	7	14	25	0	NCEER
3.9	-79.2	33.16	0	1889	2	5	19	40	0	NCEER
3.3	-71.6	43.5	0	1889	3	8	0	0	0	NCEER
4.1	-76	40	0	1889	3	8	23	40	0	NCEER
3.7	-88.1	35.9	0	1889	6	6	16	25	0	NCEER
4.4	-80.33	32.4	0	1889	7	11	21	47	0	NCEER
4	-90	35.2	0	1889	7	20	1	32	0	NCEER
3.3	-73.7	43.4	0	1889	8	10	0	0	0	NCEER
3.5	-84.87	35.03	0	1889	9	29	0	0	0	NCEER
3.1	-82.6	34.72	0	1889	10	24	10	0	0	NCEER
3.8	-95.2	31.7	0	1891	1	8	6	0	0	NCEER
3	-90	35.1	0	1891	1	14	0	0	0	NCEER
3.7	-71.6	43.2	0	1891	5	2	0	10	0	NCEER
4.2	-87.5	37.9	0	1891	7	27	2	28	0	NCEER
5.5	-88.5	38.3	0	1891	9	27	4	55	0	NCEER
3.3	-80	32.9	0	1891	10	13	5	55	0	NCEER
3.3	-71.7	44.3	0	1892	12	11	16	30	0	NCEER
3.3	-74	40.6	0	1893	3	9	5	30	0	NCEER
3.3	-72.7	42.3	0	1893	3	14	0	0	0	NCEER
3.3	-81.7	30.4	0	1893	6	21	7	7	0	NCEER
3.3	-80	32.9	0	1893	7	5	8	10	0	NCEER
3.3	-80	32.9	0	1893	9	19	7	5	0	NCEER
3.3	-80	32.9	0	1893	11	8	4	40	0	NCEER
5.2	-73.3	45.5	0	1893	11	27	16	50	0	NCEER
3.3	-80	32.9	0	1893	12	27	6	51	0	NCEER
3.3	-66.8	49.7	0	1894	1	11	9	0	0	NCEER
3.3	-80	32.9	0	1894	1	30	4	5	0	NCEER
3.3	-72.5	41.6	0	1894	4	10	0	0	0	NCEER
3.3	-80	32.9	0	1894	6	16	2	16	0	NCEER
4.3	-106.3	42.9	0	1894	6	25	0	0	0	NCEER
3.3	-80	32.9	0	1894	12	11	5	27	0	NCEER
3.3	-73.8	42.5	0	1894	12	17	0	0	0	NCEER
3.3	-80	32.9	0	1895	1	8	5	40	0	NCEER
3.3	-80	32.9	0	1895	4	27	7	40	0	NCEER
3.3	-80	32.9	0	1895	7	25	4	1	0	NCEER
3	-88.2	35.2	0	1895	7	27	0	0	0	NCEER
4.3	-74.3	40.46	0	1895	9	1	11	9	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

(Page 11 of 65)

$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.3	-80	32.9	0	1895	10	6	6	25	0	NCEER
3.5	-77.5	35.9	0	1895	10	7	4	30	0	NCEER
3.5	-103.3	43.9	0	1895	10	11	23	55	0	NCEER
5.4	-89.4	37	0	1895	10	31	11	8	0	NCEER
3.3	-80	32.9	0	1895	11	12	23	33	0	NCEER
3.3	-78.6	36.3	0	1896	2	11	1	45	0	NCEER
3.1	-84.2	40.3	0	1896	3	15	7	0	0	NCEER
3.3	-80	32.9	0	1896	3	19	8	22	0	NCEER
3.8	-67.2	45.2	0	1896	3	23	0	56	0	NCEER
3.3	-66.6	45.9	0	1896	5	16	4	0	0	NCEER
3.3	-80	32.9	0	1896	8	11	5	58	0	NCEER
3.3	-71.8	44.3	0	1896	10	22	10	30	0	NCEER
3.3	-80	32.9	0	1896	11	14	8	15	0	NCEER
3.3	-66.8	44.5	0	1897	1	28	0	0	0	NCEER
3.1	-79.2	43.1	0	1897	3	7	0	0	0	NCEER
5	-73.6	45.5	0	1897	3	23	23	7	0	NCEER
3.7	-89.6	35.8	0	1897	4	26	4	0	0	NCEER
3.3	-89	37	0	1897	5	1	4	0	0	NCEER
4.5	-73.5	44.5	0	1897	5	28	3	16	0	NCEER
5	-80.7	37.3	0	1897	5	31	18	58	0	NCEER
3.3	-71.6	43.7	0	1897	7	1	9	20	0	NCEER
3.3	-72.5	41.5	0	1897	9	5	0	0	0	NCEER
4.1	-68.7	44.7	0	1897	9	25	18	5	0	NCEER
4.1	-81.1	36.9	0	1897	10	22	3	20	0	NCEER
4.5	-106.3	42.9	0	1897	11	14	0	0	0	NCEER
3.3	-77.5	37.7	0	1897	11	27	20	56	0	NCEER
4	-97.7	37.7	0	1897	12	2	7	10	0	NCEER
4.6	-77.5	37.7	0	1897	12	18	23	45	0	NCEER
3.3	-74.3	45.1	0	1898	1	7	6	0	0	NCEER
3.3	-66.8	44.7	0	1898	1	11	9	0	0	NCEER
3.1	-90.6	34.6	0	1898	1	27	1	35	0	NCEER
4.3	-81	37	0	1898	2	5	20	0	0	NCEER
3.3	-72.6	42.8	0	1898	6	11	6	45	0	NCEER
4	-88.7	36.5	0	1898	6	14	15	6	0	NCEER
3.1	-97.3	42.6	0	1898	9	16	9	59	0	NCEER
4.3	-81	37	0	1898	11	25	20	0	0	NCEER
4.4	-81	37	0	1899	2	13	9	30	0	NCEER
3.3	-76.3	36.9	0	1899	3	3	0	0	0	NCEER
3.3	-80	32.9	0	1899	3	10	5	45	0	NCEER
4.6	-87.4	38.5	0	1899	4	30	2	5	0	NCEER
3.6	-72.6	41.6	0	1899	5	17	1	15	0	NCEER
3.3	-69.5	44	0	1899	10	5	11	30	0	NCEER
3.2	-86.5	42.1	0	1899	10	11	4	0	0	NCEER
3.1	-83	39.3	0	1899	11	12	14	0	0	NCEER
3.1	-94.4	36.8	0	1899	12	1	18	50	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

(Page 12 of 65)

$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.3	-80	32.9	0	1899	12	4	12	48	0	NCEER
3.5	-99	44.5	0	1899	12	6	12	0	0	NCEER
3	-89.5	45.5	0	1900	3	14	3	0	0	NCEER
4	-81.8	41.4	0	1900	4	9	13	0	0	NCEER
3.5	-81.7	30.4	0	1900	10	31	16	15	0	NCEER
3.1	-96.8	36	0	1900	12	1	0	0	0	NCEER
3.5	-94	37.8	0	1901	1	4	3	12	0	NCEER
3.8	-90	36	0	1901	2	15	0	15	0	NCEER
4.3	-82.66	38.95	0	1901	5	17	7	0	0	NCEER
3.3	-80	32.9	0	1901	12	2	0	26	0	NCEER
4.3	-89	42.3	0	1902	1	24	10	18	0	NCEER
3	-85.2	39.9	0	1902	3	10	6	0	0	NCEER
3.4	-80.6	37.3	0	1902	5	18	4	0	0	NCEER
4.2	-85.3	35.1	0	1902	5	29	7	30	0	NCEER
3.3	-81.4	40.3	0	1902	6	14	7	0	0	NCEER
4.2	-97.5	42.5	0	1902	7	28	18	0	0	NCEER
3.6	-97.6	30.1	0	1902	10	9	19	0	0	SRA
3.2	-85.3	35	0	1902	10	18	22	0	0	NCEER
4.1	-80	32.9	0	1903	1	24	1	0	0	NCEER
4.8	-89.3	37.8	0	1903	2	9	0	21	0	NCEER
3	-89.5	39.1	0	1903	3	17	11	50	0	NCEER
3	-71	42.7	0	1903	4	24	12	30	0	NCEER
3.1	-86.3	39.4	0	1903	9	20	0	0	0	NCEER
3.1	-88.1	38.7	0	1903	9	21	0	0	0	NCEER
3.7	-90.2	38.3	0	1903	10	5	2	56	0	NCEER
3	-89.3	37.8	0	1903	11	3	18	0	0	NCEER
4.9	-89.8	36.5	0	1903	11	4	19	14	0	NCEER
3.9	-89.5	37	0	1903	11	27	7	0	0	NCEER
3.5	-75.5	44.7	0	1903	12	25	12	30	0	NCEER
3.4	-83.5	35.7	0	1904	3	5	0	30	0	NCEER
5	-67.2	45	0	1904	3	21	6	4	0	NCEER
3.5	-100.2	37.5	0	1904	10	28	4	30	0	NCEER
3.4	-91.1	30.5	0	1905	2	3	0	0	0	NCEER
3.4	-87.7	45.1	0	1905	3	13	16	30	0	NCEER
3.6	-91.6	40.4	0	1905	4	13	16	30	0	NCEER
4.4	-70	44.2	0	1905	7	15	10	10	0	NCEER
4.5	-88.4	47.3	0	1905	7	27	0	20	0	NCEER
5.2	-89.3	37.2	0	1905	8	22	5	8	0	NCEER
3.4	-70.7	43.1	0	1905	8	30	22	40	0	NCEER
3.3	-72.2	44.9	0	1905	10	22	0	0	0	NCEER
3.3	-71.3	41.5	0	1905	11	26	0	30	0	NCEER
4.9	-96.5	39.2	0	1906	1	8	0	15	0	NCEER
3.1	-91.4	39.7	0	1906	3	6	0	0	0	NCEER
3.4	-83.6	40.7	0	1906	4	23	7	12	0	NCEER
3.6	-85.8	39.5	0	1906	5	8	6	58	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.3	-72.5	41.5	0	1906	5	8	13	30	0	NCEER
3	-75.7	38.7	0	1906	5	8	17	41	0	NCEER
3.1	-85.9	39.2	0	1906	5	9	6	38	0	NCEER
4.2	-101.3	43	0	1906	5	10	0	27	0	NCEER
3.3	-87.2	38.5	0	1906	5	11	6	15	0	NCEER
3.4	-88.4	38.7	0	1906	5	21	19	0	0	NCEER
3.3	-88.4	47.3	0	1906	5	26	14	42	0	NCEER
3.2	-81.6	40.4	0	1906	6	27	21	10	0	NCEER
3.7	-88.4	47.3	0	1906	8	8	0	0	0	NCEER
3.1	-86.8	39.7	0	1906	8	13	13	19	0	NCEER
3.8	-87.7	38.2	0	1906	9	7	16	33	0	NCEER
4	-70.5	43.5	0	1906	10	20	16	0	0	NCEER
3.3	-75.41	45.61	0	1906	11	17	14	0	0	NCEER
3.3	-77.1	41.2	0	1907	1	10	9	45	0	NCEER
3.1	-97	37.1	0	1907	1	11	7	45	0	NCEER
3.4	-86.6	39.5	0	1907	1	30	5	30	0	NCEER
3.6	-89.5	38.9	0	1907	1	31	5	30	0	SRA
4	-78.3	37.7	0	1907	2	11	13	22	0	NCEER
3.9	-80	32.9	0	1907	4	19	8	30	0	NCEER
3.3	-70.5	43.5	0	1907	6	29	0	0	0	NCEER
3.2	-90.4	37.8	0	1907	7	4	9	0	0	NCEER
3.3	-70.16	47.65	0	1907	8	5	12	43	0	NCEER
3.5	-71	42.8	0	1907	10	16	0	10	0	NCEER
3.3	-76.68	45.47	0	1907	11	14	5	0	0	NCEER
3.1	-89.8	42.3	0	1907	11	28	16	30	0	NCEER
3.1	-90.2	38.6	0	1907	12	11	4	32	0	NCEER
3.3	-70.5	47.45	0	1908	3	10	0	0	0	NCEER
3.1	-75.5	40.6	0	1908	5	31	17	42	0	NCEER
3.5	-74.8	45.1	0	1908	6	16	20	41	0	NCEER
3.3	-76.35	45.43	0	1908	7	17	7	10	0	NCEER
3.5	-67.6	46.3	0	1908	8	8	12	0	0	NCEER
3.5	-77.9	37.5	0	1908	8	23	9	30	0	NCEER
3.5	-89.6	36.6	0	1908	9	28	19	34	0	NCEER
3.6	-89.2	37	0	1908	10	28	0	27	0	NCEER
3.8	-93.2	38.7	0	1908	11	12	12	0	0	SRA
3.3	-71.7	43.5	0	1908	11	23	13	0	0	NCEER
3.7	-88	37.5	0	1908	12	27	21	15	0	NCEER
3.4	-88.6	47.2	0	1909	1	23	3	15	0	NCEER
3.3	-97.8	42.3	0	1909	1	26	20	15	0	NCEER
3.3	-73.57	45.51	0	1909	2	1	8	20	0	NCEER
3.5	-78	39.4	0	1909	4	2	7	25	0	NCEER
3.3	-74.3	46.1	0	1909	5	10	1	20	0	NCEER
5.5	-104	49	0	1909	5	16	4	15	0	NCEER
5	-88.1	41.6	0	1909	5	26	14	42	0	USHIS
3.3	-74.28	46.05	0	1909	6	8	8	25	0	NCEER



**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
4.3	-90.7	40.3	0	1909	7	19	4	34	0	NCEER
3.8	-90.1	38.3	0	1909	8	16	22	45	0	NCEER
3.7	-86.5	38.7	0	1909	9	22	0	0	0	NCEER
4.8	-87.4	39.5	0	1909	9	27	9	45	0	NCEER
3.3	-85	34.9	0	1909	10	8	10	0	0	NCEER
3.1	-90.6	37.6	0	1909	10	22	22	0	0	NCEER
3.3	-89.7	41.8	0	1909	10	22	22	30	0	NCEER
4.3	-89.5	37	0	1909	10	23	7	10	0	NCEER
3.9	-87.8	39	0	1909	10	23	9	47	0	NCEER
3.3	-75.6	45.4	0	1909	12	10	6	24	10	NCEER
3.8	-70	48	0	1910	2	1	0	0	0	NCEER
3.2	-78.7	38.8	0	1910	2	8	14	0	0	NCEER
3.1	-79.8	43.2	0	1910	2	25	0	0	0	NCEER
3.8	-97.4	41.4	0	1910	2	26	8	0	0	NCEER
3.5	-78.4	37.7	0	1910	5	8	21	10	0	NCEER
3.2	-96	30.1	0	1910	5	12	0	0	0	NCEER
4.3	-109.3	41.5	0	1910	7	26	1	30	0	DNAG
3.3	-71.1	42.7	0	1910	8	21	18	45	0	NCEER
3.3	-72.1	43.4	0	1910	8	30	14	30	0	NCEER
3.3	-68.8	44.3	0	1910	10	20	21	50	0	NCEER
4	-69.8	47.6	0	1910	10	25	9	30	0	NCEER
3.3	-79.4	36.6	0	1911	2	10	10	22	0	NCEER
3.1	-90.3	38.7	0	1911	2	28	9	0	0	NCEER
3.3	-71.5	43.2	0	1911	3	2	21	30	0	NCEER
4.6	-91.8	34	0	1911	3	31	16	57	0	NCEER
3.5	-92.2	33.8	0	1911	3	31	18	10	0	NCEER
3.3	-75.5	38.3	0	1911	4	8	1	0	0	NCEER
3.5	-82.7	35.1	0	1911	4	20	22	0	0	NCEER
4.2	-98.2	44.2	0	1911	6	2	22	34	0	NCEER
3.3	-87.6	41.8	0	1911	7	29	0	0	0	NCEER
4.7	-89	42.3	0	1912	1	2	16	21	0	NCEER
3.4	-79.7	43.2	0	1912	5	27	12	52	0	NCEER
4.9	-80	32.9	0	1912	6	12	10	30	0	NCEER
3.5	-81	32	0	1912	6	20	0	0	0	NCEER
3.3	-78.4	37.7	0	1912	8	8	1	0	0	NCEER
3	-89.1	42.3	0	1912	9	25	0	0	0	NCEER
3.3	-68	49.5	0	1912	10	23	0	0	0	NCEER
3.3	-83.5	32.7	0	1912	10	23	1	15	0	NCEER
3.3	-80	32.9	0	1912	11	17	12	30	0	NCEER
3.3	-81.7	34.7	0	1912	12	7	19	10	0	NCEER
5	-81.7	34.7	0	1913	1	1	18	28	0	NCEER
3.3	-85	34.5	0	1913	3	13	5	0	0	NCEER
4	-83.7	36.2	0	1913	3	28	21	50	0	NCEER
3.7	-84.2	35.3	0	1913	4	17	16	30	0	NCEER
4.4	-75.33	44.87	0	1913	4	29	0	28	57	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.3	-74.4	45.68	0	1913	6	8	6	30	0	NCEER
3.7	-88.9	35.8	0	1913	6	9	15	30	0	NCEER
3.8	-84	36	0	1913	8	3	16	45	0	NCEER
3.4	-74	44	0	1913	8	10	5	15	0	NCEER
3.7	-89.7	41.8	0	1913	10	17	2	15	0	NCEER
3.3	-71.5	41.5	0	1913	11	3	14	30	0	NCEER
3.1	-85.8	38.2	0	1913	11	11	14	0	0	NCEER
3.3	-67.2	45.1	0	1914	1	13	8	0	0	NCEER
3.5	-84.5	35.6	0	1914	1	24	3	24	0	NCEER
5.5	-75	46	0	1914	2	10	18	31	0	NCEER
3.5	-73.6	46.4	0	1914	2	14	9	34	0	NCEER
3.8	-70.5	45	0	1914	2	22	19	15	0	NCEER
4.6	-83.5	33.5	0	1914	3	5	20	5	0	NCEER
3.3	-79.8	34.2	0	1914	3	7	1	20	0	NCEER
3.3	-67.61	49.31	0	1914	4	12	0	0	0	NCEER
3.3	-80	32.9	0	1914	7	14	1	53	0	NCEER
4.3	-80	32.9	0	1914	9	22	7	4	0	NCEER
3.1	-89.4	43.1	0	1914	10	7	21	0	0	NCEER
3.5	-95.9	30.5	0	1914	12	30	1	0	0	NCEER
3.6	-82.2	36.6	0	1915	1	14	9	20	0	NCEER
3.1	-88.6	37.7	0	1915	2	5	6	55	0	NCEER
3.1	-89.2	37.1	0	1915	2	19	4	35	0	NCEER
3.5	-71.4	42.7	0	1915	2	21	1	20	0	NCEER
3.3	-73.4	44.7	0	1915	2	21	23	41	0	NCEER
3	-88.4	47.3	0	1915	3	3	7	45	0	NCEER
3.7	-88.1	38.7	0	1915	4	15	13	20	0	NCEER
3.1	-89.5	36.5	0	1915	4	28	23	40	0	NCEER
3.1	-103.6	48.1	0	1915	8	8	15	15	0	NCEER
3	-99.3	42.8	0	1915	9	16	19	0	0	NCEER
3.3	-88.4	47.3	0	1915	10	4	14	2	0	NCEER
3.7	-95.3	35.7	0	1915	10	8	16	50	0	NCEER
3.8	-101.5	43.8	0	1915	10	23	6	5	0	NCEER
3.4	-88.6	36.7	0	1915	10	26	7	40	0	NCEER
4.4	-90	36	0	1915	12	7	18	40	0	NCEER
4	-73.7	43.7	0	1916	1	5	13	56	0	NCEER
3.7	-87	39.1	0	1916	1	7	19	45	0	NCEER
3.8	-74	43	0	1916	2	3	4	26	0	NCEER
5.2	-83.55	35.62	0	1916	2	21	22	39	0	Chapman
3.3	-70.9	46.8	0	1916	2	29	5	15	0	NCEER
3.3	-82.7	34.5	0	1916	3	2	5	2	0	NCEER
4	-77	47	0	1916	4	24	16	7	45	NCEER
3.6	-89.5	36.6	0	1916	5	21	18	24	0	NCEER
3.3	-73.8	41	0	1916	6	8	21	15	0	NCEER
3.5	-89.2	37	0	1916	8	24	9	0	0	NCEER
3.6	-81	36	0	1916	8	26	19	36	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
5.2	-86.2	33.5	0	1916	10	18	22	3	40	NCEER
3.5	-73.7	43.3	0	1916	11	2	2	32	0	NCEER
3.7	-89.2	36.6	0	1916	12	19	5	42	0	NCEER
3.4	-83.5	36.1	0	1917	1	25	21	15	0	NCEER
3.5	-74.5	46.8	0	1917	1	26	19	35	0	NCEER
3.1	-95	47.9	0	1917	2	6	17	26	0	NCEER
3.3	-72.5	41.5	0	1917	2	16	9	0	0	NCEER
3.4	-84	36	0	1917	3	5	2	7	0	NCEER
3.6	-83.5	36.1	0	1917	3	25	19	15	0	NCEER
3.8	-101.3	35.3	0	1917	3	27	19	56	0	NCEER
4.9	-90	37	0	1917	4	9	20	52	0	NCEER
3.1	-90.2	38.1	0	1917	4	9	23	38	0	NCEER
3.9	-90.4	36.8	0	1917	5	9	9	0	0	SRA
3.9	-75.6	45.1	0	1917	5	22	9	0	26	NCEER
3.8	-89.4	36.8	0	1917	6	9	13	14	0	NCEER
4	-68	49	0	1917	6	12	2	0	0	NCEER
3.8	-83	36	0	1917	6	21	0	0	0	NCEER
3.4	-87.5	32.7	0	1917	6	30	1	23	0	NCEER
4.2	-94.8	46.3	0	1917	9	3	21	30	0	NCEER
3	-97.7	35.5	0	1918	1	1	0	0	0	NCEER
3.3	-83.9	35.9	0	1918	1	16	15	45	0	NCEER
3.7	-89.2	37	0	1918	2	17	8	10	0	NCEER
3.1	-84.2	42.8	0	1918	2	22	0	0	0	NCEER
4.7	-78.4	38.7	0	1918	4	10	2	9	0	NCEER
3.5	-84.1	36.1	0	1918	6	22	1	0	0	NCEER
3.1	-91.4	39.7	0	1918	7	1	19	2	0	NCEER
3.3	-71.35	46.85	0	1918	7	23	12	0	0	NCEER
4.2	-70.5	44.2	0	1918	8	21	4	11	54	USHIS
3.6	-98	35.5	0	1918	9	10	16	30	0	NCEER
4	-91.1	35	0	1918	10	4	9	21	0	NCEER
3.5	-91	36.1	0	1918	10	13	9	30	0	NCEER
4.2	-90	36	0	1918	10	16	2	15	0	NCEER
3.5	-87.5	37.8	0	1919	2	11	3	37	0	NCEER
3	-91.3	36.2	0	1919	4	8	12	30	0	NCEER
3.7	-89.2	36.6	0	1919	5	23	12	30	0	NCEER
3.8	-87.5	38.3	0	1919	5	25	9	45	0	NCEER
3.7	-97.3	37.7	0	1919	5	27	4	0	0	NCEER
3.7	-89.5	36.4	0	1919	5	28	13	45	0	NCEER
3.3	-70	43.9	0	1919	7	11	1	40	0	NCEER
3.3	-70.3	43.7	0	1919	7	23	11	50	0	NCEER
3.6	-97.3	37.7	0	1919	7	26	13	55	0	NCEER
3.8	-78.2	38.8	0	1919	9	6	2	46	0	NCEER
3.3	-70	47.6	0	1919	10	26	10	28	0	NCEER
3.6	-91	36.3	0	1919	11	3	20	40	0	NCEER
3.3	-69.71	48.15	0	1920	2	6	0	0	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.9	-93.3	37.2	0	1920	2	29	3	5	0	NCEER
3.7	-88.2	36.3	0	1920	4	7	20	45	0	NCEER
3.5	-89.1	38.6	0	1920	4	30	15	12	0	NCEER
3.9	-89.6	38	0	1920	5	1	15	15	0	NCEER
3.3	-71.5	43.1	0	1920	5	23	8	0	0	NCEER
3.3	-70.5	43.5	0	1920	6	7	8	0	0	NCEER
3.5	-103.2	43.2	0	1920	7	14	23	0	0	NCEER
3.3	-78.4	38.7	0	1920	7	24	0	0	0	NCEER
3.7	-94.3	38.6	0	1920	10	3	14	15	0	NCEER
3.4	-73.43	46.01	0	1920	11	8	0	0	0	NCEER
3.3	-67.1	45	0	1920	11	9	0	40	0	NCEER
3.4	-85	36	0	1920	12	24	7	30	0	NCEER
3.3	-89.5	36.4	0	1921	1	9	21	54	0	NCEER
3.2	-74.91	40.01	0	1921	1	26	23	40	0	NCEER
3.7	-89.2	37	0	1921	2	27	22	16	0	NCEER
4	-88	40	0	1921	3	14	12	15	0	NCEER
3	-96.7	43.5	0	1921	3	16	23	45	0	NCEER
3.1	-87.8	37.9	0	1921	3	31	20	3	0	NCEER
3.3	-70.4	42.5	0	1921	7	29	21	14	0	NCEER
3.8	-78.4	37.8	0	1921	8	7	6	30	0	NCEER
3.4	-76	47	0	1921	8	27	8	12	16	NCEER
3.5	-90.1	38.3	0	1921	9	9	3	0	0	NCEER
3.1	-98.7	43.7	0	1921	9	24	0	30	0	NCEER
3.5	-88.6	37.7	0	1921	10	1	9	0	0	NCEER
3.7	-90.1	38.3	0	1921	10	9	7	50	0	NCEER
3.3	-67	44.8	0	1921	10	10	13	0	0	NCEER
3.4	-84.6	35.8	0	1921	12	15	13	20	0	NCEER
4	-99.3	43.8	0	1922	1	2	14	50	0	NCEER
3.7	-87.8	37.9	0	1922	1	11	3	42	0	NCEER
4.6	-88.4	37.9	0	1922	3	22	22	30	0	NCEER
4.6	-89.4	37.4	0	1922	3	23	2	22	0	USHIS
3.8	-88.9	37	0	1922	3	23	21	45	0	NCEER
3.8	-90.4	36.7	0	1922	3	28	16	42	0	NCEER
3.1	-86.7	35.5	0	1922	3	30	1	20	0	NCEER
3.3	-82.3	36.8	0	1922	3	30	3	21	0	NCEER
3.9	-89.6	36.1	0	1922	3	30	16	53	0	NCEER
3.3	-82.2	36.5	0	1922	3	30	22	20	0	NCEER
3.3	-71.4	43.4	0	1922	5	7	22	40	0	NCEER
3.5	-66.6	46.5	0	1922	7	2	22	25	35	NCEER
3.4	-88.5	43.8	0	1922	7	7	0	0	0	NCEER
4.6	-88.2	37.4	0	1922	11	27	3	31	0	NCEER
3.5	-75.1	44.4	0	1922	12	8	21	24	0	NCEER
3.9	-89.4	38.9	0	1923	3	9	2	45	0	SRA
3.7	-89.7	34.6	0	1923	3	27	8	0	0	NCEER
3.7	-89.2	37	0	1923	5	6	7	50	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3	-96.2	41.7	0	1923	9	10	6	30	0	NCEER
4.1	-90	35.3	0	1923	10	28	17	10	0	NCEER
3.3	-89.9	40	0	1923	11	10	4	0	0	NCEER
3.6	-90.4	35.5	0	1923	11	26	23	25	0	NCEER
3.1	-89.2	37	0	1923	11	29	23	20	0	NCEER
3.3	-82.5	34.8	0	1923	12	31	20	6	0	NCEER
4.3	-90	36	0	1924	1	1	3	5	0	NCEER
3.3	-78.1	39.1	0	1924	1	5	0	0	0	NCEER
3.5	-70.2	47.8	0	1924	3	4	19	15	0	NCEER
4	-88.8	37	0	1924	4	2	11	15	0	NCEER
3.7	-89.8	36.5	0	1924	6	7	5	42	0	NCEER
3.7	-76.5	45.7	0	1924	7	15	0	10	0	NCEER
3.4	-104.5	36	0	1924	8	13	4	23	0	NCEER
3.1	-100.1	40.9	0	1924	9	24	11	0	0	NCEER
4.4	-82.6	35	0	1924	10	20	8	30	0	NCEER
4	-82.2	36.6	0	1924	11	13	10	30	0	NCEER
3.3	-76.3	45.5	0	1924	11	14	1	32	0	NCEER
3.5	-79.9	37.3	0	1924	12	26	4	30	0	NCEER
3.6	-103.5	43.5	0	1924	12	30	22	10	0	NCEER
3.9	-70.6	42.6	0	1925	1	7	13	7	0	NCEER
3.6	-91.7	36.2	0	1925	1	27	22	42	0	NCEER
6.6	-69.84	47.76	9	1925	3	1	2	19	14.7	NCEER
3.3	-71.5	42.9	0	1925	3	9	0	0	0	NCEER
3.4	-83.9	39.5	0	1925	3	27	4	6	0	NCEER
3.6	-70.8	41.7	0	1925	4	24	7	56	0	NCEER
4.9	-88.2	38	0	1925	4	27	4	5	0	NCEER
3.3	-70.9	42.5	0	1925	5	4	17	51	0	NCEER
3.6	-88.6	36.7	0	1925	5	13	12	0	0	NCEER
3.5	-77.5	37.3	0	1925	5	16	1	30	0	NCEER
3.5	-93.2	36.2	0	1925	7	8	16	0	0	NCEER
3.8	-90	38.8	0	1925	7	13	0	0	0	NCEER
3.3	-77.5	37.6	0	1925	7	14	21	20	0	NCEER
3.1	-101.2	34.5	0	1925	7	29	11	30	0	NCEER
3.4	-100.3	34.5	0	1925	7	30	8	0	0	NCEER
4.8	-101.3	35.4	0	1925	7	30	12	17	0	NCEER
3.1	-97.4	42.8	0	1925	8	25	6	27	0	NCEER
4.5	-87.2	37.9	0	1925	9	2	11	55	0	NCEER
3.3	-71.22	46.82	0	1925	10	9	5	0	0	NCEER
4	-71.1	43.7	0	1925	10	9	13	55	0	NCEER
3.3	-70.2	44.1	0	1925	10	18	21	30	0	NCEER
3.5	-73	47	0	1925	10	19	12	5	17	NCEER
3.4	-72.4	41.7	0	1925	11	14	13	4	0	NCEER
3.3	-72.7	41.8	0	1925	11	16	6	20	0	NCEER
4.3	-107	44.6	0	1925	11	18	1	50	0	NCEER
3.3	-71.8	41.6	0	1926	1	4	0	0	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
4	-94.9	35.6	0	1926	1	20	0	0	0	NCEER
3.5	-75	40	0	1926	1	26	23	40	0	NCEER
3.3	-74.1	44.3	0	1926	1	27	0	0	0	NCEER
3.3	-71	47.7	0	1926	2	19	20	20	0	NCEER
3.5	-71.8	42.8	0	1926	3	18	21	9	0	NCEER
3.5	-88.6	37.8	0	1926	3	22	14	30	0	NCEER
3.9	-89	36.2	0	1926	4	28	2	16	0	SRA
3	-73.9	40.9	0	1926	5	12	3	30	0	NCEER
4	-94.9	35.6	0	1926	6	20	14	20	0	NCEER
3.5	-82.1	35.9	0	1926	7	8	9	50	0	NCEER
3.3	-71.5	47	0	1926	7	18	6	0	0	NCEER
3.3	-77.1	45.8	0	1926	8	23	16	40	0	NCEER
3.4	-70	44.7	0	1926	8	28	21	0	0	NCEER
3.5	-90.4	36.7	0	1926	10	27	16	22	0	NCEER
3.4	-83.6	41.7	0	1926	10	28	8	42	0	NCEER
3.8	-82.1	39.1	0	1926	11	5	16	53	0	USHIS
3.3	-67.5	45	0	1926	11	24	19	30	0	NCEER
3.7	-89.4	36.7	0	1926	12	13	23	3	0	NCEER
3.5	-89.5	36.4	0	1926	12	17	0	0	0	NCEER
3.6	-97.7	38.3	0	1927	1	7	9	30	0	NCEER
3.9	-89.7	37.4	0	1927	2	2	1	30	0	SRA
3.8	-90.4	36.7	0	1927	2	3	8	0	0	SRA
3.1	-82.5	40.7	0	1927	2	17	5	30	0	NCEER
3.4	-71.4	43.3	0	1927	3	9	4	8	0	NCEER
3.3	-75.2	44.6	0	1927	3	12	22	12	0	NCEER
3.9	-95.3	39.9	0	1927	3	18	17	25	0	NCEER
3.3	-72.8	41.7	0	1927	3	30	0	0	0	NCEER
3.5	-89.5	36.3	0	1927	4	18	10	30	0	NCEER
4.7	-90.2	36	0	1927	5	7	8	28	0	NCEER
3.9	-74	40.3	0	1927	6	1	12	23	0	NCEER
3.6	-79	38	0	1927	6	10	7	16	0	NCEER
3.6	-86	34.7	0	1927	6	16	12	0	0	NCEER
4	-71	47.3	0	1927	7	25	0	56	0	NCEER
4.1	-89.5	36.4	0	1927	8	13	16	10	0	NCEER
3.3	-85.3	35.1	0	1927	10	8	4	30	0	NCEER
3.1	-98.9	41.6	0	1927	10	14	16	10	0	NCEER
3.3	-73.8	44.7	0	1927	10	24	11	0	0	NCEER
3.3	-76.2	36.3	0	1927	10	27	0	0	0	NCEER
3.4	-81.2	40.9	0	1927	10	29	0	0	0	NCEER
3.1	-79.06	43.1	0	1927	11	13	0	50	0	NCEER
3.4	-90.2	32.3	0	1927	11	13	16	21	0	NCEER
3.3	-78	33.9	0	1927	11	23	0	50	0	NCEER
3.5	-89.4	28.9	0	1927	12	15	4	30	0	NCEER
3.3	-71.6	41.2	0	1928	1	13	19	50	0	NCEER
3.1	-90	42	0	1928	1	23	9	19	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.3	-70.2	48	0	1928	1	27	0	0	0	NCEER
3.5	-69	45.3	0	1928	2	8	0	0	0	NCEER
3.5	-87	35.6	0	1928	3	7	2	45	0	NCEER
3.3	-90.2	38.6	0	1928	3	17	21	15	0	NCEER
4	-74.3	44.5	0	1928	3	18	15	25	0	NCEER
3.3	-69	45.3	0	1928	3	22	13	30	0	NCEER
3.1	-89.5	36.6	0	1928	4	15	11	0	0	NCEER
3.1	-89.5	37.3	0	1928	4	15	15	5	0	NCEER
3.1	-89.2	36.5	0	1928	4	23	11	0	0	NCEER
3.8	-71.2	44.5	0	1928	4	25	23	38	0	NCEER
3.1	-89.5	36.6	0	1928	5	31	22	40	0	NCEER
3	-84.1	40.4	0	1928	10	27	0	0	0	NCEER
3.4	-77.5	37.5	0	1928	10	30	11	45	0	NCEER
4.6	-82.83	36.11	0	1928	11	3	4	2	49.8	NCEER
3.1	-89.1	39.5	0	1928	11	8	14	15	0	NCEER
3.1	-91.1	36.1	0	1928	11	10	6	20	0	NCEER
3.5	-103.7	44.1	0	1928	11	16	13	45	0	NCEER
3.2	-67.2	45	0	1928	11	20	2	30	0	NCEER
3.7	-82.3	35.8	0	1928	11	20	3	45	0	NCEER
4	-81.5	50	0	1928	12	1	0	0	0	NCEER
3.3	-80.3	35.3	0	1928	12	23	2	30	0	NCEER
3.1	-93.9	47.6	0	1928	12	23	6	10	0	NCEER
3.1	-91.1	36.1	0	1928	12	26	3	25	0	NCEER
3.3	-80.3	33.9	0	1929	1	3	12	5	0	NCEER
3.3	-70.3	44	0	1929	2	5	19	9	0	NCEER
3.2	-87.6	38.3	0	1929	2	14	20	12	0	NCEER
3.1	-90.6	37.6	0	1929	2	26	8	15	0	NCEER
3.7	-84.2	40.4	0	1929	3	8	9	6	0	NCEER
3.3	-71.9	45.4	0	1929	5	11	9	30	0	NCEER
3.5	-89.5	36.4	0	1929	5	13	3	50	0	NCEER
3.4	-89.4	28.9	0	1929	7	28	17	0	0	NCEER
5.2	-78.4	42.91	9	1929	8	12	11	24	48.7	NCEER
3.9	-96.6	39	0	1929	9	23	11	0	0	NCEER
3.4	-97.4	42.8	0	1929	10	6	12	30	0	NCEER
3.3	-82.4	34.3	0	1929	10	28	2	15	0	NCEER
3.1	-99.8	37.2	0	1929	11	26	16	20	0	NCEER
3.4	-96.6	39.2	0	1929	12	7	8	2	0	NCEER
3.6	-78.5	38.1	0	1929	12	26	2	56	0	NCEER
4	-97.9	35.5	0	1929	12	28	0	30	0	NCEER
4.6	-65.83	46.73	0	1930	1	4	14	30	38	NCEER
3.1	-91.1	36.1	0	1930	1	26	21	0	0	NCEER
3	-90.2	37	0	1930	2	25	12	45	0	NCEER
3.3	-71.6	43.3	0	1930	3	19	0	15	0	NCEER
3.5	-90	35.1	0	1930	3	26	8	56	0	SRA
3.1	-89.7	36.1	0	1930	4	2	9	39	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.6	-71.22	45.73	0	1930	6	19	12	6	56	NCEER
3.2	-84	40.5	0	1930	6	26	21	45	0	NCEER
3.1	-83.2	40.6	0	1930	7	11	0	15	0	NCEER
3.1	-69.83	47.5	0	1930	7	13	4	52	39.3	NCEER
3.1	-70.8	41.5	0	1930	8	1	2	0	0	NCEER
3.1	-91.4	39.7	0	1930	8	8	18	31	0	NCEER
3.5	-89.1	37	0	1930	8	29	6	26	54	NCEER
3.5	-84.4	35.9	0	1930	8	30	9	28	0	NCEER
3.7	-89.4	36.6	0	1930	9	1	20	27	24	NCEER
3.5	-84.3	40.3	0	1930	9	30	20	40	0	NCEER
3.9	-68.7	48.93	0	1930	10	8	1	8	41	NCEER
3.6	-83.9	36	0	1930	10	16	21	50	0	NCEER
4.2	-91	30.1	0	1930	10	19	12	12	0	NCEER
3.3	-76.5	39.1	0	1930	11	1	1	34	0	NCEER
3.3	-92.8	34.3	0	1930	11	16	12	30	0	NCEER
3.5	-70.17	47.65	0	1930	12	13	23	18	23.7	NCEER
3.2	-90.7	38.5	0	1930	12	23	14	44	0	NCEER
3.1	-80.3	34.5	0	1930	12	26	3	0	0	NCEER
3.5	-87	39	0	1931	1	6	2	51	0	SRA
5.4	-70.4	47.3	0	1931	1	8	0	13	0.3	NCEER
3.1	-98.7	43.7	0	1931	1	17	18	45	0	NCEER
3	-84.2	40.4	0	1931	3	21	15	48	0	NCEER
3.5	-88.3	36.9	0	1931	4	1	23	20	9	NCEER
3.1	-89	36.8	0	1931	4	6	15	37	3	NCEER
4.8	-73.78	43.47	5	1931	4	20	19	54	30.6	NCEER
3.1	-78.9	42.9	0	1931	4	22	0	0	0	NCEER
4.2	-86.6	33.7	0	1931	5	5	12	18	0	NCEER
3	-73.4	41.6	0	1931	7	1	2	45	0	NCEER
3.3	-89.5	36.6	0	1931	7	18	14	52	0	NCEER
3.3	-65.77	44.62	0	1931	8	7	0	0	0	NCEER
3.8	-94.7	39.1	0	1931	8	9	6	18	37	NCEER
4.7	-84.27	40.43	5	1931	9	20	23	4	54	NCEER
4.5	-76.07	47	0	1931	9	23	22	47	37	NCEER
3.4	-70.17	47.33	0	1931	11	14	14	2	29.5	NCEER
3.3	-89.9	35.9	0	1931	12	10	8	11	36	NCEER
4.7	-89.8	34.1	0	1931	12	17	3	36	0	NCEER
3.1	-78.4	37.6	0	1932	1	5	4	5	0	NCEER
3.1	-81.6	41.1	0	1932	1	21	0	0	0	NCEER
3.8	-99.6	39	0	1932	1	29	0	15	0	NCEER
3.8	-74.67	46.47	0	1932	3	9	5	23	38.8	NCEER
3.6	-96.4	31.7	0	1932	4	9	10	15	0	NCEER
3.4	-90.2	36	0	1932	11	22	7	56	42	NCEER
3.2	-74.1	44.4	0	1932	12	7	3	15	0	NCEER
3.2	-70.5	47.45	0	1933	1	11	23	32	0	NCEER
3.8	-74.65	45.3	0	1933	1	21	16	4	39.5	NCEER



**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.3	-74.7	40.2	0	1933	1	25	2	0	0	NCEER
3.8	-99.9	39.8	0	1933	2	20	17	0	0	NCEER
3.3	-84.2	40.3	0	1933	2	23	3	20	0	NCEER
3.4	-69.93	47.43	0	1933	2	25	9	43	2.7	NCEER
3.1	-90.4	36.7	0	1933	3	11	12	48	0	NCEER
3.4	-83.7	38.6	0	1933	5	28	15	10	0	NCEER
3.3	-83.5	33.3	0	1933	6	9	11	30	0	NCEER
3	-73.8	41	0	1933	6	26	14	10	0	NCEER
3.3	-89.9	37.9	0	1933	7	13	14	42	39	NCEER
3.9	-75.7	45.42	0	1933	7	14	4	48	40	NCEER
3.1	-89.9	37.9	0	1933	8	4	4	34	15	NCEER
3.3	-103.7	41.9	0	1933	8	8	0	0	0	NCEER
3.4	-98	35.5	0	1933	8	19	19	30	0	NCEER
3.1	-73.7	43	0	1933	10	29	0	0	0	NCEER
3.3	-90.6	38.6	0	1933	11	16	9	29	1	NCEER
3.1	-89.2	42.9	0	1933	12	7	5	55	0	NCEER
4	-90.2	35.8	0	1933	12	9	8	50	0	NCEER
3.5	-80	32.9	0	1933	12	23	9	40	0	NCEER
3.1	-97.7	45.9	0	1934	1	29	12	30	0	NCEER
3.3	-72.6	41.8	0	1934	1	30	10	30	0	NCEER
3.6	-72.7	44	0	1934	4	11	3	0	0	NCEER
3.9	-95.5	33.9	0	1934	4	12	1	40	0	SRA
3.9	-73.8	44.7	0	1934	4	15	2	58	0	NCEER
3.3	-74.3	44.8	0	1934	4	15	18	5	0	NCEER
3.2	-98.7	41.5	0	1934	5	11	10	40	0	NCEER
3	-89.9	37.9	0	1934	5	15	14	28	0	NCEER
3.1	-90	35.2	0	1934	7	3	3	10	41	NCEER
4.3	-103	42.7	0	1934	7	30	7	20	0	NCEER
3.2	-70.3	43.7	0	1934	8	2	14	59	0	NCEER
4.3	-89.2	36.9	0	1934	8	20	0	47	0	NCEER
3	-67	44.9	0	1934	8	26	11	36	0	NCEER
3.1	-99.1	43.4	0	1934	8	30	3	50	0	NCEER
3.2	-80.2	42	0	1934	10	29	20	7	0	NCEER
3.3	-88.5	37.5	0	1934	10	30	2	25	47	NCEER
3.4	-100.2	42.6	0	1934	11	8	4	45	0	NCEER
3.9	-90.5	41.5	0	1934	11	12	14	45	0	NCEER
3.3	-80	32.9	0	1934	12	9	9	0	0	NCEER
3.4	-83.6	35.1	0	1935	1	1	8	15	0	NCEER
3	-90.6	41.5	0	1935	1	5	18	40	0	NCEER
3.3	-77.4	37.2	0	1935	2	10	23	45	0	NCEER
4.8	-96.2	40.3	0	1935	3	1	10	59	44	NCEER
3.3	-70.2	42.2	0	1935	4	24	1	24	0	NCEER
3.1	-89.5	36.4	0	1935	7	24	1	38	0	NCEER
6.2	-79.07	46.78	0	1935	11	1	6	3	40	NCEER
3.3	-78.9	38.9	0	1935	11	1	8	30	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
4.9	-78.17	47.23	0	1935	11	2	14	31	58	NCEER
3.3	-81.7	29.6	0	1935	11	14	3	30	0	NCEER
3.1	-83.2	41.2	0	1936	1	31	19	30	0	NCEER
3.1	-89.7	36.2	0	1936	2	17	5	5	8	NCEER
3.4	-95.2	34	0	1936	3	14	17	20	0	NCEER
4	-70.25	47.33	0	1936	3	29	0	49	23.4	NCEER
3.4	-71.5	43.5	0	1936	6	14	5	40	0	NCEER
4.4	-100.77	35.31	5	1936	6	20	3	24	3.5	NCEER
3.3	-74.2	44.7	0	1936	6	21	4	20	0	NCEER
3	-102.9	36.9	0	1936	7	12	0	23	0	NCEER
3.9	-89	36.7	0	1936	8	2	22	16	25	NCEER
3.3	-84.4	39.3	0	1936	10	8	16	30	0	NCEER
3.1	-103.5	43.5	0	1936	10	30	10	30	0	NCEER
3.3	-71.4	43.6	0	1936	11	10	2	46	0	NCEER
3.2	-71.7	44.7	0	1936	11	10	4	2	0	NCEER
3.3	-89.7	36.2	2	1937	1	30	8	57	9	NCEER
3.6	-78.7	37.7	0	1937	2	3	1	26	0	NCEER
4.9	-84.27	40.49	2	1937	3	2	14	47	33.3	NCEER
3.3	-75.2	44.6	0	1937	3	10	5	29	0	NCEER
4	-90.6	36.1	0	1937	5	17	0	49	46	NCEER
3.2	-96.9	35.3	0	1937	6	8	14	26	0	NCEER
3.5	-73.71	40.72	0	1937	7	19	3	51	0	NCEER
4.5	-65.43	47.8	5	1937	9	30	7	58	3.4	NCEER
4	-75.82	46.78	24	1937	11	6	14	31	20.6	NCEER
3.7	-74.47	46.1	1	1937	11	12	16	57	31.3	NCEER
4	-89.1	38.6	0	1937	11	17	17	4	0	NCEER
3.5	-98.2	44.5	0	1938	1	2	17	5	0	NCEER
3.2	-75.18	44.9	0	1938	1	6	13	28	42.2	NCEER
3	-76.27	45.57	0	1938	1	24	5	29	2	NCEER
3.8	-87	41.6	0	1938	2	12	6	27	0	NCEER
3.2	-75.4	46.38	0	1938	2	23	17	56	35.7	NCEER
3.1	-83.2	42.4	0	1938	3	13	16	10	0	NCEER
3.3	-103.4	42.7	0	1938	3	24	13	11	0	NCEER
3.6	-83.5	35.6	0	1938	3	31	10	10	0	NCEER
3.2	-79.08	46.72	0	1938	4	12	18	55	47	NCEER
3.1	-93.5	34.2	0	1938	4	26	5	42	0	NCEER
3	-74.5	45.37	0	1938	5	5	0	33	0.3	NCEER
3.9	-68	49	0	1938	5	17	18	32	0	NCEER
3	-66.8	46.5	0	1938	6	15	5	7	43	NCEER
3.2	-78.43	40.68	1	1938	7	15	22	46	12	NCEER
3.1	-73.7	41.08	0	1938	8	2	9	2	30	NCEER
3.8	-68.79	44.89	5	1938	8	22	12	48	9.4	NCEER
3.8	-74.36	40.05	21	1938	8	23	5	4	53.4	NCEER
3	-74.9	45.87	0	1938	9	7	23	18	18.9	NCEER
4.8	-90.254	35.413	1	1938	9	17	3	34	28.3	USHIS

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.2	-72.2	41.5	0	1938	9	20	0	0	0	NCEER
4.1	-69.58	48.78	0	1938	9	28	4	33	16	NCEER
3.8	-99.3	43.8	0	1938	10	1	22	15	0	NCEER
3.8	-96.7	43.5	0	1938	10	11	9	37	0	NCEER
3.3	-98.9	43.2	0	1938	11	4	22	10	0	NCEER
3.5	-75.25	44.75	0	1938	11	18	22	19	6	NCEER
4.2	-76.2	47.03	0	1938	11	26	7	47	57.5	NCEER
3.9	-75.4	47.6	0	1938	12	25	7	46	0	NCEER
3.3	-79.85	43.25	0	1939	1	14	8	10	16	NCEER
3.6	-95.8	46.8	0	1939	1	28	17	55	0	NCEER
3	-78.3	42.9	0	1939	2	24	0	20	0	NCEER
3.6	-77.5	46.4	0	1939	3	16	20	21	0	NCEER
3.3	-84	40.4	0	1939	3	18	14	3	0	NCEER
3.2	-89.4	36.8	0	1939	4	15	17	25	0	NCEER
3.3	-85.8	33.7	0	1939	5	5	2	45	0	NCEER
3.9	-96.4	35	0	1939	6	1	7	30	0	NCEER
3.1	-98.9	43	0	1939	6	10	18	30	0	NCEER
3.1	-84	40.3	0	1939	6	18	3	20	0	NCEER
4.1	-92.6	34.1	0	1939	6	19	21	43	12	NCEER
3.4	-86.6	34.7	0	1939	6	24	10	27	0	NCEER
4.5	-69.98	47.59	14	1939	6	24	17	20	18.3	NCEER
5.6	-69.8	47.8	15	1939	10	19	11	53	58	NCEER
4.1	-70.5	47.8	0	1939	11	7	2	40	32	NCEER
3.8	-75.05	39.58	3	1939	11	15	2	53	48.7	NCEER
4.9	-90.14	38.18	0	1939	11	23	15	14	52	NCEER
3.5	-76.6	39.5	0	1939	11	26	5	20	0	NCEER
3.6	-71.4	47.97	0	1939	12	8	1	17	47	NCEER
3	-79.08	46.72	0	1940	1	5	0	34	14	NCEER
3.4	-70.8	41.6	0	1940	1	28	23	12	0	NCEER
4	-76.83	46.5	0	1940	2	10	20	57	17.3	NCEER
3.5	-78.5	38.8	0	1940	3	26	0	1	0	NCEER
3.8	-70.73	47.73	0	1940	4	13	8	13	34	NCEER
3.6	-73.2	45.8	0	1940	5	16	14	0	17.1	NCEER
3.4	-88.6	37.1	0	1940	5	31	19	3	0	NCEER
3.1	-82.3	40.9	0	1940	6	16	4	30	0	NCEER
3.1	-74.78	46.25	0	1940	8	4	16	20	52	NCEER
3	-74.83	45.77	0	1940	8	7	23	57	35.3	NCEER
3.5	-71.13	47	0	1940	9	11	1	6	55.4	NCEER
4.7	-69.8	47.8	0	1940	10	13	19	50	51	NCEER
3.3	-85.1	34.7	0	1940	10	19	5	54	0	NCEER
5	-90.1	38.2	0	1940	11	23	21	15	0	NCEER
3.1	-94	33	0	1940	12	2	16	16	0	NCEER
5.5	-71.37	43.87	10	1940	12	20	7	27	26.2	NCEER
3.2	-82.9	35.9	0	1940	12	25	6	5	0	NCEER
3.6	-87.3	37.9	0	1940	12	29	2	30	0	SRA

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.4	-83.9	36	0	1941	3	4	6	15	0	NCEER
3	-75.5	46.27	0	1941	3	5	7	29	23.2	NCEER
3.3	-73.92	44.73	0	1941	4	4	8	10	43.7	NCEER
3.3	-82.6	35.6	0	1941	5	10	11	12	0	NCEER
3.7	-103.5	43.5	0	1941	5	25	6	25	0	NCEER
3	-70.34	47.39	0	1941	6	22	9	59	31	NCEER
4.1	-76.83	47.4	0	1941	6	26	4	5	44.9	NCEER
3	-90.8	32.3	0	1941	6	28	18	30	0	NCEER
3.7	-67.9	46.1	0	1941	8	30	10	21	0	NCEER
3	-85.3	35	0	1941	9	8	9	45	0	NCEER
4	-70.73	47.63	0	1941	10	6	16	34	27.6	NCEER
3.3	-89.7	36.2	0	1941	10	8	7	51	0	NCEER
3	-72.3	42.3	0	1941	10	11	8	15	0	NCEER
3.2	-99	35.4	0	1941	10	18	7	48	0	NCEER
3.3	-74.8	44.77	0	1941	10	21	6	10	41	NCEER
3.3	-89.1	37	0	1941	10	21	16	53	0	NCEER
3.6	-74.3	45.7	0	1941	10	24	14	13	59.3	NCEER
3.1	-90	35.1	0	1941	11	15	3	7	0	NCEER
4.2	-89.7	35.5	0	1941	11	17	3	8	0	NCEER
3.3	-90.3	38.4	0	1942	1	14	18	5	0	NCEER
3.3	-81	26.5	0	1942	1	19	0	0	0	NCEER
3.1	-74.77	46.83	0	1942	2	18	7	55	12	NCEER
3.6	-89.7	41.2	0	1942	3	1	14	43	10	NCEER
3.2	-70.4	44.2	0	1942	3	8	23	37	0	NCEER
3	-103.5	44.4	0	1942	3	11	16	55	0	NCEER
3	-88.6	37.7	0	1942	3	29	12	43	0	NCEER
4.4	-74.67	45.77	0	1942	5	20	12	19	22.8	NCEER
3.9	-73.8	44.7	0	1942	5	24	11	33	0	NCEER
3.5	-97.9	36.4	0	1942	6	12	4	50	0	NCEER
3.7	-77.5	46.8	0	1942	8	26	17	54	0	NCEER
3.1	-89.2	37	0	1942	8	31	9	28	0	NCEER
3.1	-71.5	46.97	0	1942	9	5	14	30	24.1	NCEER
3.4	-99.3	38.8	0	1942	9	10	9	0	0	NCEER
3.7	-67.4	49.22	0	1942	9	11	11	5	13	NCEER
3.3	-76	46.78	0	1942	9	15	22	32	46	NCEER
3	-73.8	42.57	0	1942	10	2	22	29	50.5	NCEER
3.3	-78.4	37.6	0	1942	10	7	2	15	0	NCEER
3.4	-75.25	40.97	0	1942	10	24	17	27	3.6	NCEER
3.6	-75.05	46.42	0	1942	11	16	0	13	29.4	NCEER
3	-90.2	38.6	0	1942	11	17	18	18	0	NCEER
4.2	-76.07	46.97	0	1942	12	5	21	10	51.2	NCEER
4.4	-69.33	45.16	0	1943	1	14	21	32	38	NCEER
3.7	-75.77	46.5	0	1943	2	28	16	40	1.2	NCEER
4.4	-81.31	41.63	7	1943	3	9	3	25	24.9	NCEER
3.9	-71.6	43.7	0	1943	3	14	14	2	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.1	-85.8	38.3	0	1943	4	13	17	0	0	NCEER
3.2	-73.83	44.77	0	1943	5	9	11	3	12.5	NCEER
3.1	-103.5	43.5	0	1943	5	16	19	40	0	NCEER
3	-90.4	38.6	0	1943	6	8	19	50	0	NCEER
4	-105	48.5	0	1943	6	25	4	25	0	NCEER
3.8	-73.03	44.84	22	1943	7	6	22	10	16	NCEER
3.3	-70.65	47.55	0	1943	9	25	5	52	36.1	NCEER
3.8	-70.4	47.27	0	1943	9	28	16	30	25.2	NCEER
3.9	-70.08	47.38	0	1943	11	6	0	6	40.5	NCEER
3.2	-74.87	47.68	0	1943	12	6	7	19	40	NCEER
3	-69.6	44.6	0	1943	12	19	9	0	44	NCEER
3.3	-80.2	33	0	1943	12	28	10	25	0	NCEER
3.2	-89.7	37.5	0	1944	1	7	5	18	0	NCEER
4.2	-75.5	39.8	0	1944	1	8	0	0	0	DNAG
4.3	-76.78	45.83	0	1944	1	22	21	55	9.1	NCEER
3.3	-80	32.9	0	1944	1	28	17	30	0	NCEER
4	-70.5	47.4	0	1944	2	5	12	37	52.5	NCEER
3.7	-76.2	40.8	0	1944	2	5	16	22	0	NCEER
4.1	-78.9	46.7	0	1944	3	8	12	49	0	NCEER
4.9	-67.4	49.9	0	1944	4	9	12	44	0	NCEER
3.7	-70.28	47.3	0	1944	6	9	15	19	8.7	NCEER
5.1	-67.75	49.42	0	1944	6	23	6	37	53	NCEER
3.7	-74.25	46	0	1944	6	24	23	48	38.5	NCEER
5.8	-74.72	44.96	12	1944	9	5	4	38	45.7	NCEER
5	-107.5	39	0	1944	9	9	4	12	20	DNAG
3.9	-90	37.9	0	1944	9	25	11	37	23	NCEER
4.2	-67	48.5	0	1944	10	14	13	26	17	NCEER
4.4	-80.8	48.7	0	1944	11	5	19	7	0	NCEER
4.1	-84.4	40.4	0	1944	11	13	11	52	0	NCEER
3.1	-87.1	45.7	0	1944	12	10	11	0	0	NCEER
3.6	-72.8	41.6	0	1944	12	14	3	15	0	NCEER
3.1	-89.7	36.2	0	1944	12	23	7	23	0	NCEER
3.2	-90.2	37.8	0	1945	1	16	2	0	0	NCEER
3.3	-80	32.9	0	1945	1	30	20	20	0	NCEER
3.7	-90.2	38.6	0	1945	3	28	1	45	58	NCEER
3.1	-76.4	43	0	1945	4	15	13	15	0	NCEER
3.3	-89.7	36.5	0	1945	5	2	11	22	0	NCEER
3.1	-90.2	38.6	0	1945	5	21	7	51	0	NCEER
4.3	-75.4	47.08	0	1945	6	12	7	58	15.1	NCEER
3.6	-84.5	35	0	1945	6	14	3	25	0	NCEER
4.7	-71.09	47.34	5	1945	6	18	15	20	4.7	NCEER
3.9	-76.8	48.47	0	1945	7	2	13	29	52.1	NCEER
3.3	-67	44.9	0	1945	7	15	10	44	0	NCEER
4	-81.38	33.75	5	1945	7	26	10	32	16.4	NCEER
3.1	-89.8	37	0	1945	9	23	6	22	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
4.7	-69.81	47.99	5	1945	10	9	13	18	42	NCEER
3.3	-78.5	37.5	0	1945	10	12	19	0	0	NCEER
3.1	-97.9	43	0	1945	11	10	8	0	0	NCEER
3.7	-89.2	37	0	1945	11	13	8	21	0	NCEER
3	-74.9	45	0	1945	12	2	15	22	32	SRA
4.3	-68.7	49.4	0	1946	1	17	8	4	52	NCEER
3.3	-89.1	38.6	0	1946	2	25	0	52	0	NCEER
3.8	-84.9	35.2	0	1946	4	7	5	0	0	NCEER
3.6	-73.43	45.73	0	1946	4	21	5	5	55.5	NCEER
4	-90.8	36.6	0	1946	5	15	6	10	0	NCEER
3	-74.53	44.65	0	1946	6	27	21	6	22	NCEER
4	-98.6	44.1	0	1946	7	23	6	45	0	NCEER
3.3	-71.47	47.33	0	1946	9	1	4	39	41	NCEER
3	-74.88	44.9	0	1946	9	4	19	30	0	NCEER
3.2	-75	47.72	0	1946	9	19	0	53	28.8	NCEER
3.4	-72.15	46.43	0	1946	9	26	21	19	8.2	NCEER
4	-90.6	37.5	0	1946	10	8	1	12	2	NCEER
3.1	-103.6	48.1	0	1946	10	26	20	37	0	NCEER
3.6	-76.6	41.5	0	1946	10	28	20	36	0	NCEER
3.1	-77.45	42.87	0	1946	11	10	11	41	23.1	NCEER
3	-74.9	45	0	1946	11	11	10	20	47	SRA
3.1	-74.68	45.17	0	1946	11	24	10	20	47.2	NCEER
3	-74.9	44.9	0	1946	12	25	4	48	3	NCEER
3.3	-73.6	41	0	1947	1	4	18	51	0	NCEER
3.9	-76.7	46.8	0	1947	1	19	0	45	1.7	NCEER
4.2	-70.53	47.67	0	1947	2	2	16	50	32.3	NCEER
3.1	-88.3	42.1	0	1947	3	16	15	30	0	NCEER
4	-88.4	37	0	1947	3	26	0	0	0	NCEER
4	-70.23	47.37	0	1947	3	29	12	28	52.4	NCEER
3.5	-87.9	43	0	1947	5	6	21	27	0	NCEER
3.1	-100.9	46	0	1947	5	14	5	2	0	NCEER
3	-100.3	44.4	0	1947	5	16	5	45	0	NCEER
3.4	-84	36	0	1947	6	6	12	55	0	NCEER
4.2	-90.2	38.4	0	1947	6	30	4	23	53	NCEER
4.4	-81.1	46.5	0	1947	8	8	5	39	0	NCEER
4.5	-85	41.93	2	1947	8	10	2	46	41.3	NCEER
3.1	-98.9	43.1	0	1947	8	25	14	0	0	NCEER
4.3	-81.3	47	0	1947	9	14	19	29	0	NCEER
3.3	-92.6	31.9	0	1947	9	20	21	30	0	NCEER
3.8	-70.72	47.55	0	1947	10	22	9	36	38.3	NCEER
3.3	-80	32.9	0	1947	11	2	4	30	0	NCEER
4.5	-81.2	45.7	0	1947	11	3	19	51	0	NCEER
3.7	-90.6	36.7	0	1947	12	1	7	47	33	NCEER
4	-90.1	35.6	0	1947	12	16	3	27	0	SRA
3.2	-85.3	35	0	1947	12	27	19	0	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

(Page 28 of 65)

$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.8	-69.2	45.2	0	1947	12	28	19	58	0	NCEER
4.5	-70.4	47.3	0	1948	1	1	18	33	45.3	NCEER
3.3	-78.3	37.7	0	1948	1	5	2	45	0	NCEER
3.5	-78.5	37.5	0	1948	1	5	3	20	0	NCEER
3.1	-89.1	38.6	0	1948	1	6	1	34	0	NCEER
3.1	-89.7	43.1	0	1948	1	15	17	40	0	NCEER
3.7	-69	50	0	1948	1	16	6	2	56	NCEER
3.4	-84.1	36.4	0	1948	2	10	0	4	0	NCEER
4.6	-102.48	36.22	5	1948	3	12	4	29	6.3	NCEER
3.1	-97.3	37.7	0	1948	4	3	3	0	0	NCEER
3.1	-91.8	41.7	0	1948	4	20	14	17	0	NCEER
3.2	-71.8	41.4	0	1948	5	4	2	23	0	NCEER
4	-73.69	45.86	3	1948	5	7	12	2	27.3	NCEER
3.7	-73.87	45.23	0	1948	6	9	3	4	12.2	NCEER
3.3	-82.2	26.5	0	1948	11	8	17	44	0	NCEER
3.5	-70.3	46.7	0	1948	11	13	16	49	56.6	NCEER
3	-69.2	45.2	0	1948	11	29	4	56	0	NCEER
3.5	-89.7	36.4	0	1949	1	14	3	49	0	NCEER
3.3	-80	32.9	0	1949	2	2	10	52	0	NCEER
3.6	-77.6	37.6	0	1949	5	8	11	1	0	NCEER
3.3	-99	42.5	0	1949	5	13	4	15	0	NCEER
3.1	-100	45	0	1949	6	3	0	0	0	NCEER
3.3	-80	32.9	0	1949	6	27	6	53	0	NCEER
3.2	-83	36.7	0	1949	9	17	9	30	0	NCEER
4.7	-70.58	44.84	20	1949	10	5	2	33	47.8	NCEER
4.2	-74.9	45.49	14	1949	10	16	23	33	45.4	NCEER
3.4	-72.12	46.47	0	1949	10	30	20	51	13.7	NCEER
5	-109.5	35.7	0	1950	1	17	0	51	0	DNAG
5.3	-110.5	40.5	0	1950	1	18	1	55	51	USHIS
3.7	-92.7	37.7	0	1950	2	8	10	37	0	NCEER
3.3	-95.2	46.1	0	1950	2	15	10	5	0	NCEER
4	-74.5	46	0	1950	3	6	16	14	11.8	NCEER
3.1	-97.1	33.5	0	1950	3	20	13	24	0	NCEER
3.3	-75.8	41.5	0	1950	3	20	22	55	11.5	NCEER
4.9	-75.5	47.83	0	1950	4	14	18	20	48.5	NCEER
3.1	-84.2	39.8	0	1950	4	20	0	0	0	NCEER
3.5	-84	35.8	0	1950	6	19	4	19	0	NCEER
4.3	-68.1	49.9	0	1950	6	29	9	13	33	NCEER
3.2	-70.25	47.33	0	1950	8	4	6	45	21	NCEER
4	-74.72	45.2	0	1950	8	4	14	29	28.7	NCEER
3	-89.9	35.7	0	1950	9	17	5	48	0	NCEER
3	-77.12	45.82	0	1950	10	29	5	59	26	NCEER
3.3	-78.3	37.7	0	1950	11	26	7	45	0	NCEER
3.2	-72.5	41.5	0	1951	1	26	3	27	0	NCEER
3.3	-80	32.9	0	1951	3	4	2	55	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.5	-77.6	37.6	0	1951	3	9	7	0	0	NCEER
3.7	-71.53	41.52	5	1951	6	10	17	20	37.7	NCEER
4.3	-103.04	35.22	1	1951	6	20	18	37	11.1	NCEER
4.2	-67.5	50	0	1951	6	28	1	3	57	NCEER
3.3	-71.37	47.2	0	1951	7	25	0	22	51.5	NCEER
3.3	-74.67	45.93	0	1951	8	8	9	36	24.1	NCEER
3.6	-73.86	41.35	18	1951	9	3	21	26	24.8	NCEER
4.3	-66.25	49.3	0	1951	9	19	8	19	38	NCEER
3.3	-89.9	38.7	0	1951	9	20	2	38	0	NCEER
3.7	-75.37	46.22	0	1951	9	25	15	45	0	NCEER
3.8	-74.73	45.27	0	1951	10	25	7	7	52.8	NCEER
3.9	-73.55	44.92	31	1951	11	6	17	54	45.9	NCEER
3.3	-75.5	40.6	0	1951	11	23	6	45	0	NCEER
3.3	-80	32.9	0	1951	12	30	7	55	0	NCEER
3.1	-73.2	44.5	0	1952	1	30	4	0	0	NCEER
3.3	-69.38	46.33	0	1952	2	18	20	56	7	NCEER
3.9	-89.5	36.4	0	1952	2	20	22	34	39	NCEER
3.7	-70.2	46.8	0	1952	2	26	0	56	0	NCEER
3.8	-76.17	47.1	0	1952	3	17	4	14	41	NCEER
4.1	-69.88	47.83	0	1952	3	30	13	11	7	NCEER
5.1	-97.85	35.52	10	1952	4	9	16	29	28.4	NCEER
3.8	-70.58	47.47	0	1952	4	19	2	50	52.8	NCEER
3.7	-78.5	47	0	1952	4	26	4	59	0	NCEER
3.3	-89.7	36.6	0	1952	5	28	9	54	14	NCEER
3.3	-82.3	36.3	0	1952	6	11	20	20	0	NCEER
3.9	-82.02	39.64	9	1952	6	20	9	38	8.6	NCEER
4	-89.6	36.2	0	1952	7	16	23	48	10	NCEER
4.3	-75.84	46.87	1	1952	7	19	1	16	17.2	NCEER
3.2	-74.5	43	0	1952	8	25	0	7	0	NCEER
3.1	-96.5	35.1	0	1952	10	8	4	15	0	NCEER
4.3	-74	41.7	0	1952	10	8	21	40	0	DNAG
4.9	-69.8	47.8	0	1952	10	14	22	3	44.8	NCEER
3.1	-89.6	36.2	0	1952	10	17	4	16	18	NCEER
3.1	-93.7	30.1	0	1952	10	17	15	48	0	NCEER
3.1	-103.5	44.1	0	1952	11	15	0	0	0	NCEER
3.3	-84.6	30.6	0	1952	11	18	20	12	0	NCEER
3.1	-80	32.9	0	1952	11	19	0	0	0	NCEER
3.1	-81	43.8	0	1952	12	25	0	0	0	NCEER
3.6	-89.8	35.9	0	1952	12	25	4	23	24	NCEER
4.6	-66	49.07	0	1953	1	24	9	58	37	NCEER
3.1	-89.5	36	0	1953	1	26	23	18	0	NCEER
3.3	-78.1	37.7	0	1953	2	7	7	5	0	NCEER
3.3	-89.5	36.5	0	1953	2	11	10	50	54	NCEER
3.1	-89.8	36.1	0	1953	2	17	11	45	0	NCEER
3.5	-74.43	48.07	0	1953	2	28	6	24	2.5	NCEER



Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
4.2	-98	35.6	0	1953	3	17	14	25	0	NCEER
3.3	-81.4	28.6	0	1953	3	26	0	0	0	NCEER
3	-73.5	41.1	0	1953	3	27	8	50	0	NCEER
3.6	-73	43.7	0	1953	3	31	2	50	0	NCEER
3	-73.5	44.7	0	1953	4	26	1	17	0	NCEER
3.1	-90.3	35.6	0	1953	5	12	18	50	0	NCEER
3.1	-96.7	34.7	0	1953	6	6	17	40	0	NCEER
3.5	-83.6	41.7	0	1953	6	12	0	0	0	NCEER
4.3	-110.163	38.997	0	1953	7	30	5	45	0	DNAG
3.1	-74	41	0	1953	8	17	4	22	50	NCEER
3.9	-90.1	38.8	0	1953	9	11	18	26	28	NCEER
4.4	-65.2	49.1	0	1953	9	14	22	52	57	NCEER
3.3	-83.9	36	0	1953	10	11	4	0	0	NCEER
3	-102.9	45.2	0	1953	12	21	22	43	0	NCEER
3.3	-89.1	38.6	0	1953	12	30	22	0	0	NCEER
3.1	-99.3	43.1	0	1953	12	31	20	30	0	NCEER
3.8	-83.2	37.3	0	1954	1	1	2	30	0	NCEER
4.4	-83.7	36.6	0	1954	1	2	3	25	0	NCEER
3.2	-76	40.3	0	1954	1	7	7	25	0	NCEER
3.1	-68.23	49.17	0	1954	1	10	21	4	30	NCEER
3.1	-89.4	36	0	1954	1	17	7	15	0	NCEER
3.4	-105.5	41.5	0	1954	1	20	20	50	1	NCEER
3.1	-84.4	35.3	0	1954	1	23	1	0	0	NCEER
3.1	-77.3	42.9	0	1954	1	31	12	30	0	NCEER
3.3	-76.65	43.03	0	1954	2	1	0	37	50	NCEER
4.3	-90.3	36.7	0	1954	2	2	16	53	0	NCEER
3.8	-70.25	47.6	0	1954	2	7	20	24	16	NCEER
3.5	-70.62	47.67	0	1954	2	21	9	0	37	NCEER
3.1	-96.4	35	0	1954	4	11	0	0	0	NCEER
4.3	-76.12	47	6	1954	4	12	21	22	0.1	NCEER
3.1	-73.5	44.7	0	1954	4	21	15	45	0	NCEER
4.1	-79.2	43.1	0	1954	4	27	2	14	8	NCEER
3.9	-90	35.1	0	1954	4	27	4	9	0	NCEER
3	-74.2	45	0	1954	5	20	22	0	0	NCEER
3.7	-70.12	47	0	1954	6	30	7	41	7	NCEER
3.5	-70.7	42.81	1	1954	7	29	19	56	56	NCEER
3.1	-87.3	38.5	0	1954	8	9	0	0	0	NCEER
3.3	-76	40.3	0	1954	8	11	3	40	0	NCEER
3.6	-68.37	49.03	0	1954	9	8	1	29	53	NCEER
4.6	-75.66	47.18	28	1954	9	11	18	55	55.6	NCEER
3	-74.6	44.6	0	1954	12	13	3	53	0	NCEER
3.1	-88.4	47.3	0	1955	1	5	20	0	0	NCEER
3.1	-82.2	36.6	0	1955	1	6	20	30	0	NCEER
3.4	-88.6	47.1	0	1955	1	7	5	0	0	NCEER
3.3	-78.4	37.3	0	1955	1	17	12	37	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
4	-73.8	43	0	1955	1	21	8	40	0	DNAG
4.3	-89.83	36.07	8	1955	1	25	7	24	39.1	NCEER
3.1	-83.9	36	0	1955	1	25	19	34	0	NCEER
4	-70.5	47.67	0	1955	2	1	12	40	27	NCEER
4.1	-89.1	30.4	0	1955	2	1	14	45	0	NCEER
4.3	-107	40.5	0	1955	2	10	17	30	0	NCEER
3.3	-98.6	41.3	0	1955	2	25	1	45	0	NCEER
4.3	-89.78	38.23	11	1955	4	9	13	1	23.3	NCEER
3.2	-81.4	41.3	0	1955	5	26	18	9	0	NCEER
3.1	-81.4	41.3	0	1955	6	29	1	16	33	NCEER
3	-79.63	43.77	0	1955	6	29	1	17	40	NCEER
3.5	-78.3	42.9	0	1955	8	16	7	35	0	NCEER
3.4	-89.5	36	0	1955	9	6	1	45	0	NCEER
3.1	-89.5	36.4	0	1955	9	24	18	45	0	NCEER
3.5	-81.3	36.6	0	1955	9	28	7	1	41.5	NCEER
3.5	-73.9	45.22	0	1955	10	7	18	9	52	NCEER
3.4	-70.2	48.93	0	1955	10	20	21	31	6	NCEER
3.5	-75.87	46.5	0	1955	11	1	7	45	52	NCEER
3.4	-89.5	36	0	1955	12	13	7	43	0	NCEER
3.3	-82.4	34.3	0	1956	1	5	8	0	0	NCEER
4	-98.35	37.58	29	1956	1	6	11	58	7.4	NCEER
3.8	-94.8	29.3	0	1956	1	7	23	30	0	SRA
3.3	-75.47	45.67	0	1956	1	10	12	8	18	NCEER
3.7	-84	40.5	0	1956	1	27	11	3	27	NCEER
3.9	-89.8	35.76	16	1956	1	29	4	44	15.5	NCEER
3.7	-71.17	47.05	0	1956	1	30	9	43	13	NCEER
3.1	-74.82	45.45	0	1956	2	2	19	24	16	NCEER
3.9	-97.3	35.4	0	1956	2	16	23	30	0	NCEER
3.1	-75.38	44.85	0	1956	3	6	23	38	10	NCEER
3.7	-90.4	40.5	0	1956	3	13	15	5	0	SRA
3.5	-95.6	34.2	0	1956	4	2	16	3	18	NCEER
3.3	-82.4	34.3	0	1956	5	19	19	0	0	NCEER
3.9	-76.43	47.1	0	1956	6	15	0	53	37	NCEER
3.1	-87.7	43.6	0	1956	7	18	21	30	0	NCEER
3.4	-73.78	44.7	0	1956	7	27	1	34	44	NCEER
3.6	-66.17	49.42	0	1956	8	3	12	52	9	NCEER
4	-83.79	36.44	5	1956	9	7	13	35	50.8	NCEER
4.2	-84	35.5	0	1956	9	7	13	49	29	NCEER
3.2	-86.7	35.8	0	1956	9	9	22	45	0	SRA
3.8	-88.4	31.9	0	1956	9	27	14	15	0	NCEER
3.1	-87.9	42.9	0	1956	10	13	0	0	0	NCEER
3.4	-69	48.25	0	1956	10	27	14	40	6	NCEER
3.4	-89.7	36.1	0	1956	10	29	9	23	44	NCEER
4	-95.8	36.2	0	1956	10	30	10	36	21	USHIS
4	-75.42	45.96	1	1956	11	4	11	53	29.2	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
4.3	-90.39	36.91	1	1956	11	26	4	12	43.3	NCEER
3	-88.8	43.5	0	1957	1	8	16	0	0	NCEER
4	-83.7	36.6	0	1957	1	25	18	15	0	NCEER
3.5	-69.93	48.4	0	1957	2	19	5	18	33	NCEER
3.3	-74.9	44.9	0	1957	2	20	15	45	0	NCEER
4	-94.7	32.6	0	1957	3	19	16	37	38	NCEER
3.5	-74.8	40.6	0	1957	3	23	19	2	0	NCEER
3.1	-88.4	37	0	1957	3	26	8	27	6	NCEER
4.3	-86.72	33.77	5	1957	4	23	9	23	39	NCEER
3.3	-72	44.42	0	1957	4	24	0	41	59	NCEER
4.4	-70.25	43.53	5	1957	4	26	11	40	8.6	NCEER
4.1	-82.14	35.8	5	1957	5	13	14	24	51.1	NCEER
3.3	-84.1	35.95	5	1957	6	23	6	34	16	NCEER
3.1	-81.3	42.9	0	1957	6	29	11	25	9	NCEER
3.9	-82.7	35.6	0	1957	7	2	9	33	1	NCEER
3.7	-67.08	46.58	0	1957	8	4	12	40	58	NCEER
4	-70.42	47.48	0	1957	8	6	23	50	38	NCEER
3.3	-70.12	46.73	0	1957	8	17	1	30	7	NCEER
3.1	-89.5	36.2	0	1957	8	17	23	0	0	NCEER
3	-76.17	44.8	0	1957	8	21	2	40	33	NCEER
3.1	-69.9	48.42	0	1957	10	9	14	16	58	NCEER
3.2	-78.75	46.38	0	1957	10	27	8	48	27	NCEER
3.3	-84	36	0	1957	11	7	17	15	0	NCEER
3.5	-69.55	48.67	0	1957	11	13	20	49	19	NCEER
3.9	-83.5	35	0	1957	11	24	20	6	17	NCEER
3.1	-98.2	43.8	0	1957	12	3	7	30	0	NCEER
4.5	-87.9	38.4	0	1958	1	8	2	41	43	DNAG
3.3	-74.9	44.9	0	1958	1	11	16	36	0	NCEER
3.5	-81.3	45	0	1958	1	24	17	10	0	NCEER
3.8	-89.7	36.1	0	1958	1	26	16	55	37	NCEER
3.9	-89.2	37.1	0	1958	1	28	5	56	40	NCEER
3.9	-76.03	46.9	0	1958	3	1	17	41	49	NCEER
3.4	-77.8	34.2	0	1958	3	5	11	53	43	NCEER
3.1	-77.13	46	0	1958	3	19	6	39	25	NCEER
3.4	-67.12	45.55	0	1958	3	23	22	4	17	NCEER
3.4	-89.2	36.3	0	1958	4	8	22	25	33	NCEER
3.2	-89.5	36.4	0	1958	4	26	7	30	0	NCEER
3.1	-81.8	41.5	0	1958	5	1	22	46	31	NCEER
3.7	-70.32	48.57	0	1958	5	6	16	2	49	NCEER
5	-76.82	47.09	1	1958	5	14	17	41	16.7	NCEER
3.3	-82.6	35.6	0	1958	5	16	22	30	0	NCEER
3.1	-90.4	35.5	0	1958	5	20	1	25	0	NCEER
3.2	-71.4	46.7	0	1958	7	18	23	56	27	NCEER
4.4	-79.5	43	0	1958	7	22	1	46	40	NCEER
3.8	-75.8	46.57	0	1958	7	25	3	45	11	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

(Page 33 of 65)

$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3	-70.3	47.32	0	1958	7	27	8	58	0	NCEER
3.8	-80	43.13	0	1958	8	4	20	25	58	NCEER
3.1	-106	41.1	0	1958	8	7	0	46	43	NCEER
3.6	-70.38	47.93	0	1958	8	8	22	15	3	NCEER
3.9	-69.38	48.6	0	1958	8	12	3	22	12	NCEER
3.6	-79	43	0	1958	8	22	14	25	5	NCEER
3	-70.2	43.6	0	1958	9	19	17	45	0	NCEER
3.3	-69.27	48.38	0	1958	9	29	10	45	29	NCEER
3.7	-73.73	45.18	0	1958	9	30	0	13	58	NCEER
3.5	-82.7	34.5	0	1958	10	20	6	16	0	NCEER
4.1	-68	49.6	0	1958	10	21	9	32	51	NCEER
3	-81.9	37.2	5	1958	10	23	2	29	44.3	NCEER
3.1	-90.1	29.9	0	1958	11	6	23	8	0	NCEER
4.4	-88.01	38.44	5	1958	11	8	2	41	12.6	NCEER
3.3	-91.2	30.5	0	1958	11	19	18	15	0	NCEER
3.3	-71.7	44	0	1958	11	21	23	30	0	NCEER
3.7	-69.82	46.98	0	1958	12	23	23	14	16	NCEER
3.1	-98.1	44.9	0	1959	1	12	13	0	0	NCEER
3.1	-89.5	36.3	0	1959	1	21	15	35	0	NCEER
3.1	-81	43	0	1959	2	9	0	0	0	NCEER
4.2	-100.9	35.5	0	1959	2	10	20	5	0	NCEER
3.2	-89.5	36.1	0	1959	2	13	8	37	0	NCEER
3.4	-73.27	41.92	0	1959	4	13	21	20	19	NCEER
3.5	-70.33	47.12	0	1959	4	16	16	36	25	NCEER
3.7	-80.68	37.39	1	1959	4	23	20	58	39.5	NCEER
3.9	-76.45	46.55	0	1959	5	21	9	38	51	NCEER
3.5	-79.2	48.8	0	1959	5	24	10	52	0	NCEER
3.6	-84.3	35.4	0	1959	6	13	1	0	0	SRA
3.7	-96.7	34.7	0	1959	6	15	12	45	0	NCEER
4.1	-98.06	34.64	5	1959	6	17	10	27	10.6	NCEER
3.1	-80.7	37.3	0	1959	7	7	23	17	0	NCEER
4.1	-68.32	48.42	0	1959	8	1	13	52	49	NCEER
3.9	-86.56	34.79	5	1959	8	12	18	6	1.4	NCEER
3.2	-80.7	37.3	0	1959	8	21	17	20	0	NCEER
3.2	-70.78	46.95	0	1959	8	22	3	52	30	NCEER
3.4	-93.1	29.8	0	1959	10	15	15	45	0	NCEER
4	-80.2	34.5	0	1959	10	27	2	7	28	NCEER
3.4	-89.34	36.03	5	1959	12	21	16	23	39.6	NCEER
4.3	-106.2	41.1	0	1959	12	25	9	50	0	NCEER
3.7	-75.67	46.97	0	1960	1	20	20	7	40	NCEER
3.4	-75.5	41.5	0	1960	1	22	20	53	22	SRA
3.2	-89.5	36	0	1960	1	28	21	38	0	NCEER
3.3	-70.38	47.8	0	1960	2	6	0	44	2	NCEER
4.2	-80.12	33.07	9	1960	3	12	12	47	44	NCEER
3.5	-84	35.8	0	1960	4	15	10	10	10	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

(Page 34 of 65)

$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.4	-89.5	36	0	1960	4	21	10	45	0	NCEER
4	-70.34	47.88	5	1960	4	23	11	47	47.1	NCEER
3.1	-92	34.2	0	1960	5	4	16	31	32	NCEER
3.7	-80	32.9	0	1960	7	24	3	37	30	NCEER
3.3	-79.3	37.4	0	1960	9	4	18	40	0	NCEER
5.5	-107.6	38.3	49	1960	10	11	8	5	30.5	USHIS
3.6	-95.5	34.9	0	1961	1	11	1	40	0	NCEER
3.8	-66.93	46.38	0	1961	1	29	0	49	39	NCEER
3.7	-83.3	41.2	0	1961	2	22	9	45	3	NCEER
3.2	-75.28	45.17	0	1961	3	13	10	55	45	NCEER
3.6	-99.77	39.98	1	1961	4	13	21	14	55.2	NCEER
3.2	-74.8	45	0	1961	4	20	13	13	0	NCEER
3.6	-95	34.6	0	1961	4	26	7	5	0	NCEER
3.8	-95.3	34.9	0	1961	4	27	7	30	0	NCEER
3.4	-70.5	47.33	0	1961	8	22	18	55	51	NCEER
3.8	-90.19	35.96	5	1961	9	9	22	42	55	NCEER
4.3	-75.5	40.8	0	1961	9	15	2	16	56	DNAG
3.1	-74.9	44.9	0	1961	9	29	6	30	0	NCEER
3.8	-76.58	48.67	0	1961	10	7	22	36	51	NCEER
3.9	-94.24	39.32	9	1961	12	25	12	58	16.8	NCEER
3.3	-74.8	40.5	0	1961	12	27	17	6	0	NCEER
4	-100.72	44.25	23	1961	12	31	16	36	5.8	NCEER
3.8	-74.85	45.92	0	1962	1	27	12	11	17	NCEER
3.5	-67.13	47.5	0	1962	1	31	14	32	38	NCEER
4.3	-89.51	36.37	4	1962	2	2	6	43	30	NCEER
3	-88.7	37	0	1962	2	16	0	0	0	NCEER
3.3	-69.47	47.18	0	1962	3	23	2	2	21	NCEER
4	-66.02	47.57	0	1962	3	25	5	15	5	NCEER
3	-79.3	43	0	1962	3	27	6	35	0	NCEER
4.3	-72.97	44.11	5	1962	4	10	14	30	45.2	NCEER
3.3	-98.6	35.3	0	1962	4	28	6	9	11	NCEER
3	-89.5	36.5	0	1962	5	24	0	0	0	NCEER
3.2	-90.39	35.38	1	1962	6	1	11	23	38.6	NCEER
3.2	-72.64	45.44	1	1962	6	21	2	6	47	NCEER
5.4	-88.64	37.9	0	1962	6	27	1	28	59.3	NCEER
3.2	-89.82	36.56	1	1962	7	14	2	23	44	NCEER
3.6	-89.4	36.04	8	1962	7	23	6	5	15.7	NCEER
3.9	-70.67	47.25	0	1962	7	27	17	56	57	NCEER
3.2	-97.4	34.8	0	1962	8	10	20	47	19	SRA
3.6	-70.05	47.53	0	1962	8	11	3	5	16	NCEER
3.3	-77.7	39.5	0	1962	9	4	23	40	0	NCEER
3.3	-78.2	39.7	0	1962	9	7	14	0	0	NCEER
3.1	-110.89	39.2	7	1962	9	7	16	50	23.8	SRA
3.2	-98.4	34.7	0	1962	9	7	22	53	44	SRA
3.3	-74.3	44.8	0	1962	10	2	23	45	0	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3	-69.13	45.57	0	1962	12	1	21	29	23	NCEER
3.4	-110.42	39.36	7	1962	12	11	10	28	13.5	SRA
3	-71.7	42.8	0	1962	12	29	6	19	0	NCEER
3.3	-80.1	37.3	0	1963	1	17	14	26	50.8	NCEER
3	-75.9	44	0	1963	1	30	14	50	0	NCEER
3.4	-92.1	34.4	0	1963	2	7	21	18	36	NCEER
4.3	-109.2	42.6	33	1963	2	25	18	45	16.5	SRA
3	-79.57	43.2	0	1963	2	27	6	0	0	NCEER
3.4	-75.73	41.51	0	1963	3	2	20	24	32	NCEER
4.8	-90.05	36.64	15	1963	3	3	17	30	10.6	NCEER
3.1	-95.9	34.6	0	1963	3	13	9	33	34	NCEER
3.9	-109.8	45.1	33	1963	4	3	9	55	12.6	SRA
3.1	-89.58	36.46	6	1963	4	6	8	12	22.7	NCEER
3.1	-89.54	36.67	10	1963	5	2	1	9	21.4	NCEER
3.3	-80.19	32.97	5	1963	5	4	21	1	50.3	NCEER
3	-96.4	34.3	0	1963	5	7	20	3	29	SRA
3.5	-75.2	43.5	0	1963	5	19	19	14	0	NCEER
4.4	-104	39.3	0	1963	6	5	0	13	50.6	NCEER
4	-104.4	36.6	0	1963	6	6	8	5	33	SNMX
3.3	-73.75	42.37	0	1963	7	1	19	59	12	NCEER
3.1	-90.47	36.97	0	1963	7	8	23	51	42.1	NCEER
3	-66.5	46.8	0	1963	8	1	6	34	16	NCEER
4.4	-88.77	36.98	7	1963	8	3	0	37	49.1	NCEER
3.5	-73.95	45.18	0	1963	8	26	16	29	35	NCEER
4.3	-111.22	38.1	7	1963	9	30	9	17	39.3	SRA
3.2	-82.5	33.9	0	1963	10	8	6	1	43.4	SRA
3.6	-78.197	39.655	0	1963	10	10	14	59	52.3	SRA
4.5	-108.3	42.2	30	1963	10	14	8	31	23	SRA
4.2	-77.47	46.37	8	1963	10	15	13	59	50.8	NCEER
3.8	-70.42	42.4	14	1963	10	16	15	30	59.7	NCEER
3.7	-81	36.7	0	1963	10	28	22	38	0.3	NCEER
3.4	-70.8	42.7	0	1963	10	30	22	36	57.9	NCEER
3.2	-71.6	43.6	9	1963	12	4	21	32	34.8	NCEER
3.2	-86.97	37.15	1	1963	12	5	6	51	0.5	NCEER
3.4	-104.133	35.133	0	1963	12	19	16	47	28	SNMX
3.9	-77.53	46.23	0	1964	1	8	10	3	26	NCEER
3.3	-89.46	36.84	0	1964	1	16	5	9	58	NCEER
3.1	-70.7	46.92	27	1964	1	20	18	57	45.3	NCEER
3	-89.5	36.5	0	1964	1	25	19	54	10	NCEER
3.4	-99.7	35.1	0	1964	2	2	8	23	0	NCEER
3.3	-77.96	40.38	1	1964	2	13	19	46	40.8	NCEER
3.3	-85.4	34.7	0	1964	2	17	22	47	0	SRA
3.3	-82.39	33.72	5	1964	3	7	18	2	58.6	NCEER
3.9	-83.31	33.19	1	1964	3	13	1	20	17.5	NCEER
3.5	-89.6	36.2	0	1964	3	17	2	16	6	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.5	-103.5	43.5	0	1964	3	24	6	12	0	NCEER
3.4	-104.1	42.7	0	1964	3	28	3	0	0	NCEER
4.5	-101.8	43	30	1964	3	28	10	8	46.5	NCEER
4.3	-74.9	44.9	0	1964	3	29	9	16	0	NCEER
3.1	-71.5	43.6	0	1964	4	1	11	21	34	NCEER
3.8	-81.1	46.4	0	1964	4	5	13	21	0	NCEER
3.7	-81.1	33.84	3	1964	4	20	19	4	44.1	NCEER
3.6	-93.81	31.42	5	1964	4	24	7	33	51.9	NCEER
3.2	-76.41	40.3	1	1964	5	12	6	45	10.7	NCEER
3.9	-90.02	36.58	3	1964	5	23	11	25	34.5	NCEER
4.2	-94	31.3	0	1964	6	2	23	0	0	SRA
3	-94	31	0	1964	6	3	9	37	0	NCEER
3.3	-74.3	40.9	0	1964	6	16	0	0	0	NCEER
3.8	-71.68	43.4	1	1964	6	26	11	4	49	NCEER
3.7	-79.2	47.8	0	1964	6	27	19	17	0	NCEER
3.8	-67.42	49.43	0	1964	7	1	21	41	30	NCEER
3.4	-71.41	46.72	0	1964	7	12	0	0	41	NCEER
3.3	-76.25	46.65	0	1964	7	24	10	34	11	NCEER
3	-83.9	36	0	1964	7	28	0	0	0	NCEER
4	-106	39.7	0	1964	8	4	11	13	25.2	NCEER
3	-110.92	38.95	7	1964	8	5	15	17	56.2	SRA
3	-93.8	31.4	0	1964	8	16	11	35	31	NCEER
4.5	-104.7	42.9	0	1964	8	22	3	28	11	NCEER
3	-102.25	43.77	20	1964	8	26	16	58	55.1	NCEER
3.1	-73.87	48.4	0	1964	9	9	6	16	26	NCEER
4.1	-107.8	41.9	33	1964	9	10	6	19	50.7	SRA
4	-109.7	50	33	1964	9	19	20	51	5	DNAG
3	-91.1	37.1	0	1964	9	24	8	9	34	NCEER
3.4	-96.4	44	0	1964	9	28	15	41	0	NCEER
3	-89.8	47.4	0	1964	10	10	8	30	0	NCEER
3	-90.3	47.3	0	1964	10	10	11	30	0	NCEER
3.2	-83.9	36	0	1964	10	13	16	30	0	NCEER
3.9	-67.25	47.67	0	1964	10	17	14	13	7	NCEER
3.1	-73.7	41.2	0	1964	11	17	17	8	0	NCEER
3.6	-81.698	37.394	6	1964	11	25	2	50	6.4	SRA
3.1	-110.916	38.923	7	1964	12	16	21	39	25.2	DNAG
3	-72	43.5	0	1965	1	3	17	5	1	NCEER
3.5	-78.5	48	0	1965	1	8	12	29	45	NCEER
3.3	-110.35	39.44	7	1965	1	14	12	30	10.8	SRA
3.4	-103.8	35.1	0	1965	2	3	11	32	34	SNMX
3	-103	31.9	0	1965	2	3	19	59	32	SRA
3.3	-89.59	36.52	3	1965	2	11	3	40	24.8	NCEER
3	-93.3	36.9	0	1965	2	14	20	3	20	NCEER
3.1	-71.25	47.5	0	1965	3	1	2	22	8	NCEER
3.2	-78.83	47.72	0	1965	3	5	12	11	1	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
4	-91.03	37.4	7	1965	3	6	21	8	50.3	NCEER
3.1	-83	45	0	1965	3	6	21	13	0	NCEER
3.1	-67.53	49.77	0	1965	3	18	12	9	5	NCEER
3.7	-89.52	36.46	3	1965	3	25	12	59	27.7	NCEER
3.4	-80.5	46	0	1965	4	1	6	30	20	NCEER
3.5	-81.6	37.32	5	1965	4	26	15	26	19.7	NCEER
3.3	-89.9	36.1	0	1965	5	25	7	15	43	NCEER
4.3	-106.3	39.4	33	1965	5	30	17	31	4.1	SRA
3	-89.5	36.5	0	1965	6	1	7	24	57	NCEER
4.7	-106.5	43.6	33	1965	6	3	19	30	25.8	SRA
3.3	-89.5	36.5	0	1965	7	8	7	3	50	NCEER
3	-78.08	43.04	18	1965	7	16	11	6	57	NCEER
3.1	-109.9	39.5	33	1965	7	18	3	55	51.4	SRA
3.9	-109.8	44.7	33	1965	8	6	15	39	49.2	SRA
3.8	-89.31	37.23	1	1965	8	14	13	13	56.9	NCEER
3.3	-110.6	42.3	33	1965	8	22	17	54	33.3	SRA
3.1	-78.1	43	0	1965	8	28	1	55	0	SRA
3.5	-102.3	32.1	0	1965	8	30	5	17	38	NCEER
3.2	-65.28	46	0	1965	8	31	8	38	44	NCEER
3.2	-81.2	34.7	0	1965	9	9	14	42	20	NCEER
3.8	-79.05	46.72	0	1965	9	15	17	56	28	NCEER
3.3	-74.4	41.4	0	1965	9	29	20	57	0	NCEER
3.9	-67.66	49.78	0	1965	10	5	14	36	55	NCEER
3.3	-79.75	40.08	0	1965	10	8	2	17	27	NCEER
3.1	-97.7	36.1	0	1965	10	10	23	51	33	NCEER
4.9	-90.94	37.48	5	1965	10	21	2	4	39.1	NCEER
3	-70.1	41.3	0	1965	10	24	17	45	0	NCEER
3.4	-90.92	37.03	4	1965	11	4	7	43	37.9	NCEER
4.2	-76.36	47.25	10	1965	11	7	20	57	41.8	NCEER
3.3	-83.2	33.2	0	1965	11	8	12	58	1	SRA
3.7	-76.28	46.93	0	1965	11	24	21	28	1	NCEER
3.2	-71.4	41.7	0	1965	12	8	3	2	0	NCEER
4.1	-70.6	47.83	0	1965	12	16	13	53	19	NCEER
3.8	-89.76	36.03	1	1965	12	19	22	19	12	NCEER
3.9	-78.25	42.84	2	1966	1	1	13	23	39	NCEER
3.9	-67.47	48.9	0	1966	1	14	15	29	25	NCEER
3.6	-89.87	35.95	1	1966	2	12	4	32	12.8	NCEER
3.5	-87	33.6	0	1966	2	13	6	29	43	SRA
3.6	-90.9	37.04	6	1966	2	13	23	19	37.8	NCEER
3	-90	36.2	0	1966	3	13	14	24	42	NCEER
3.2	-76.16	46.5	0	1966	3	20	23	45	33	NCEER
3.4	-103.333	35.283	0	1966	4	21	14	14	19	SNMX
3.3	-71.9	44.1	0	1966	4	28	12	2	0	NCEER
4.3	-78.13	37.66	2	1966	5	31	6	18	59.5	NCEER
3.1	-88.2	38.6	0	1966	6	22	11	27	53	NCEER



**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.4	-73.83	45.16	0	1966	6	25	0	5	51	NCEER
3.1	-103.43	44.3	2	1966	6	26	11	59	43.1	NCEER
3.3	-66	49.5	0	1966	7	12	1	6	38	NCEER
3.7	-101.33	35.64	3	1966	7	20	9	4	58.8	NCEER
3.2	-70	47.75	0	1966	7	20	20	8	29	NCEER
3	-67.6	44.5	0	1966	7	24	1	59	0	NCEER
3.7	-68.55	49.63	0	1966	7	24	22	19	46	NCEER
3.1	-110.36	39.44	7	1966	7	30	3	25	31	SRA
4.3	-102.339	32.115	3	1966	8	14	15	25	53.7	USHIS
3	-84	35.8	0	1966	8	24	6	0	0	NCEER
4.2	-107.6	38.3	33	1966	9	4	9	52	34.5	SRA
3.1	-98.81	41.3	27	1966	9	9	9	50	34.2	NCEER
3.1	-80.3	39.3	0	1966	9	28	0	0	0	NCEER
3.2	-65.25	46.92	0	1966	9	28	20	11	35	NCEER
3	-70.33	47.66	0	1966	10	1	17	23	55	NCEER
4.5	-104.1	37.4	0	1966	10	3	2	26	2.3	NCEER
3	-104.6	39.3	0	1966	10	13	0	33	0	SRA
3.2	-71.8	43	0	1966	10	23	23	5	0	NCEER
3.9	-106.9	40.2	33	1966	11	1	7	40	28	SRA
3.6	-76.3	47	0	1966	11	13	15	43	0	NCEER
3	-92.8	38.9	0	1966	12	6	8	0	47	NCEER
3.4	-68.17	49	0	1966	12	12	21	4	12	NCEER
3.7	-107.017	36.983	0	1966	12	16	2	0	40	SNMX
3.3	-106.5	39	5	1966	12	19	20	52	33.3	SRA
4.4	-107.51	38.98	33	1967	1	12	3	52	6.2	SRA
4.1	-107.86	37.67	33	1967	1	16	9	22	45.9	SRA
3.8	-107.05	40.05	33	1967	1	18	6	12	0.6	SRA
3.6	-109.77	44.74	33	1967	1	21	0	18	16	SRA
3.1	-84.6	42.7	0	1967	2	2	6	30	0	NCEER
3.3	-71.4	41.4	0	1967	2	2	13	40	9	NCEER
3	-110.1	39.55	33	1967	2	5	10	7	16.6	SRA
3.1	-90	36	0	1967	2	12	0	0	0	SRA
3.4	-110.37	39.27	5	1967	2	15	15	2	16.5	SRA
3.1	-110.28	42.05	7	1967	3	10	2	20	33.2	SRA
3.6	-109.9	45.16	33	1967	3	28	20	31	35.4	SRA
3	-107.75	38.32	33	1967	4	4	22	53	39.5	SRA
3.7	-82.53	39.65	1	1967	4	8	5	40	30.5	NCEER
3	-89.7	36.1	0	1967	4	11	23	44	45	NCEER
4.7	-108.77	43.41	5	1967	4	26	10	17	59.4	SRA
4.8	-105.9	43.66	0	1967	5	11	21	15	6.6	NCEER
3.2	-69.9	42.3	0	1967	5	15	22	47	12	NCEER
4.3	-90.84	33.55	6	1967	6	4	16	14	12.6	NCEER
3.7	-75.03	46.58	0	1967	6	11	1	49	39	NCEER
3.9	-78.23	42.84	1	1967	6	13	19	8	55.5	NCEER
3.8	-69.9	44.4	0	1967	7	1	16	11	18.9	SRA

Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.4	-90.4	35.8	0	1967	7	6	16	43	51	NCEER
4.6	-90.44	37.44	15	1967	7	21	9	14	48.8	NCEER
3.1	-91.1	37.1	0	1967	8	25	19	15	18	NCEER
3.4	-70.7	46.93	0	1967	9	23	16	27	55	NCEER
4.7	-65.63	49.44	15	1967	9	30	22	39	48	NCEER
3	-89.5	36.5	0	1967	10	18	5	8	36	NCEER
3.4	-80.22	32.8	19	1967	10	23	9	4	2.5	NCEER
3.2	-73.8	41.2	0	1967	11	22	22	10	0	NCEER
3.5	-99.6	43.56	1	1967	11	23	6	23	42.1	NCEER
3.5	-81.6	37.36	2	1967	12	16	12	23	33.4	NCEER
3.1	-95.55	34.85	0	1968	1	4	22	30	0	NCEER
3.8	-106.8	42.7	33	1968	1	9	2	16	39.3	SRA
3.3	-89.8	36.2	0	1968	1	23	16	16	0	NCEER
3.8	-89.86	36.52	7	1968	2	10	1	34	30.6	NCEER
3.2	-110.61	41.72	7	1968	2	20	6	34	26.4	SRA
3.4	-80.77	37.28	8	1968	3	8	5	38	15.7	NCEER
3.1	-70.49	47.94	18	1968	3	30	15	28	59	NCEER
4.5	-89.85	38.02	1	1968	3	31	17	58	9.6	NCEER
3.5	-70.4	47.6	0	1968	4	11	9	18	0	NCEER
3.8	-102.1	37.8	0	1968	4	21	7	8	7	NCEER
3.3	-66.66	46.9	18	1968	5	27	19	21	56	NCEER
3.5	-89.5	36.5	0	1968	5	30	1	59	33	NCEER
3.3	-110.45	39.21	7	1968	6	2	18	59	23.2	SRA
3	-110.47	41.93	7	1968	6	14	21	11	15.3	SRA
3.8	-107.41	39.31	33	1968	6	23	20	16	13	SRA
3.7	-100.74	46.59	27	1968	7	8	16	50	14.7	NCEER
3.1	-89.5	36.5	0	1968	7	14	4	21	25	SRA
3	-90.8	35.7	0	1968	7	15	4	21	25	NCEER
3.1	-71.3	47.01	18	1968	7	24	23	16	37	NCEER
3	-84.2	40.4	0	1968	7	26	15	2	53.7	SRA
3.1	-81.48	34.11	1	1968	9	22	21	41	18.2	NCEER
3.3	-69.45	45.17	18	1968	9	23	15	38	50	NCEER
3.4	-81.66	45.8	18	1968	10	10	20	10	41	NCEER
3.5	-96.8	34	0	1968	10	14	14	42	54	NCEER
3.2	-74.1	45.3	0	1968	10	19	10	37	0	NCEER
3.6	-70.6	47.5	0	1968	10	20	2	36	0	NCEER
3.7	-83	43	0	1968	10	31	0	0	0	NCEER
3.3	-72.5	41.4	0	1968	11	3	8	33	0	NCEER
3.1	-76.3	46.17	18	1968	11	3	20	50	49	NCEER
5.5	-88.37	37.91	21	1968	11	9	17	1	40.5	NCEER
3.3	-77.9	34.1	0	1968	11	25	20	0	0	NCEER
3.2	-74.6	39.7	0	1968	12	10	9	12	0	NCEER
3.4	-87.6	37.8	0	1968	12	11	15	0	0	NCEER
3	-85.8	38.3	0	1968	12	11	16	0	0	NCEER
4.4	-92.69	34.99	7	1969	1	1	23	35	38.7	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.4	-90.4	37.8	0	1969	1	20	19	25	0	NCEER
3.5	-96.3	34.2	0	1969	4	13	6	27	51	NCEER
3.3	-96.31	35.29	8	1969	5	2	11	33	21.7	NCEER
3.6	-70.65	47.47	18	1969	5	10	18	43	29	NCEER
3.5	-82.58	33.95	0	1969	5	18	0	0	0	NCEER
3.1	-78.245	39.61	0	1969	5	22	14	59	51.6	SRA
3.5	-104.4	40.4	0	1969	5	26	1	30	8.6	NCEER
3	-97.8	34.8	0	1969	5	30	14	8	5	NCEER
3.1	-81.45	49.67	18	1969	6	4	9	36	2	NCEER
3	-97	37.4	0	1969	7	1	3	36	58	NCEER
4.1	-83.69	36.12	1	1969	7	13	21	51	9.8	NCEER
3.8	-70.09	47.83	18	1969	7	14	3	6	59	NCEER
3.1	-89.5	36.5	0	1969	7	27	0	0	0	SRA
3.5	-71.4	43.8	0	1969	8	6	16	2	0.5	NCEER
3.2	-70.07	47.49	18	1969	8	31	7	20	27	NCEER
3.3	-74.6	41.1	0	1969	10	6	0	0	0	NCEER
3.1	-106.58	48.29	18	1969	10	6	20	24	53	NCEER
4.5	-75.06	46.31	2	1969	10	10	0	7	4.9	NCEER
4.6	-80.93	37.45	5	1969	11	20	1	0	9.3	NCEER
3.6	-77.67	37.84	1	1969	12	11	23	44	37.4	NCEER
3.7	-82.85	35.04	6	1969	12	13	10	19	29.7	NCEER
3.1	-89.9	35.2	0	1970	1	7	17	45	0	NCEER
3.8	-103.417	35.9	0	1970	1	12	11	21	15	SNMX
3.1	-97	31	0	1970	2	3	0	0	0	NCEER
4	-108.31	37.92	33	1970	2	3	5	59	35.6	SRA
3.4	-90.6	37.9	0	1970	2	6	4	53	2	NCEER
3.1	-77.78	48.24	18	1970	2	27	8	8	36	NCEER
3.3	-89.54	36.6	5	1970	3	27	3	44	29.2	NCEER
3.1	-81.22	49.7	18	1970	4	25	0	46	27	NCEER
3.2	-78.275	39.619	0	1970	5	27	17	59	41.4	SRA
3.7	-82.206	36.99	12	1970	7	30	15	15	16.9	SRA
3.5	-83.4	37.7	0	1970	7	31	0	31	0	SRA
3.3	-66.12	45.8	18	1970	8	8	0	10	30	NCEER
3.2	-82.05	38.23	10	1970	8	11	6	14	25.5	NCEER
3.2	-70.3	47.92	18	1970	9	7	21	39	27	NCEER
3.1	-81.42	36.02	1	1970	9	10	1	41	5.2	NCEER
3	-71.03	48.72	18	1970	10	9	16	35	1	NCEER
3.3	-76.25	47.07	18	1970	10	15	18	56	11	NCEER
3	-90	36	0	1970	11	5	10	25	35	NCEER
4.4	-89.95	35.86	16	1970	11	17	2	13	54.1	NCEER
3	-89.5	36.3	0	1970	11	30	4	46	53	NCEER
3	-89	36	0	1970	12	8	23	16	0	NCEER
4.9	-107.55	43.96	15	1970	12	12	15	57	19.1	SRA
3.4	-89.55	36.71	15	1970	12	24	10	17	56.8	NCEER
3	-75.96	47.17	18	1971	1	6	6	22	8	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

(Page 41 of 65)

$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.8	-107.31	39.49	33	1971	1	7	20	39	52.1	SRA
3.1	-75.18	46.92	18	1971	1	19	13	44	25	NCEER
3.1	-87.85	38.5	15	1971	2	12	12	44	27.5	NCEER
3	-83.2	37.1	0	1971	2	19	23	11	42	NCEER
3.6	-87.84	33.18	12	1971	3	14	17	27	54.6	NCEER
3.7	-88.3	32.8	0	1971	3	16	2	37	28	NCEER
4.4	-106.97	40.7	10	1971	3	18	9	8	59.9	SRA
3	-81.6	37.4	0	1971	4	1	5	5	11	NCEER
3	-90.1	35.8	0	1971	4	13	14	0	51	NCEER
3.2	-73.37	45.1	18	1971	5	14	6	20	9	NCEER
3.7	-80.66	33.36	1	1971	5	19	12	54	3.6	NCEER
4.1	-74.48	43.9	2	1971	5	23	6	24	27.9	NCEER
3	-76.28	46.55	18	1971	7	6	17	47	49	NCEER
3.1	-81.2	46.74	18	1971	7	9	5	5	26	NCEER
3.7	-109.6	40.24	7	1971	7	10	17	22	36.8	SRA
3.4	-84	36	0	1971	7	13	2	3	0	SRA
3.4	-84.3	36	0	1971	7	13	3	3	0	NCEER
3.6	-83	34.8	0	1971	7	13	11	42	26	NCEER
3.3	-75.6	39.7	0	1971	7	14	0	0	0	NCEER
3.6	-103.17	31.64	5	1971	7	30	1	45	51.4	NCEER
3.8	-80.63	33.34	4	1971	7	31	20	16	55	NCEER
3.6	-77.59	38.15	5	1971	9	12	0	6	27.6	NCEER
3.2	-70.24	47.56	18	1971	9	12	8	31	43	NCEER
3	-103.2	31.6	0	1971	9	24	1	1	54	SRA
3.2	-75.17	45.71	18	1971	9	27	8	47	23	NCEER
4.1	-90.49	35.77	9	1971	10	1	18	49	38.5	NCEER
3.7	-83.37	35.8	8	1971	10	9	16	43	32.7	NCEER
3	-89.6	36.7	0	1971	10	18	6	39	31	NCEER
3.7	-101.26	43.69	17	1971	10	19	21	7	37.4	NCEER
3.4	-71.2	42.7	0	1971	10	21	0	54	0	NCEER
3.3	-83	36	0	1971	10	22	21	55	0	SRA
3	-67.13	49.23	18	1971	10	27	7	13	24	NCEER
4	-108.68	38.91	5	1971	11	12	9	30	44.6	SRA
3	-73.87	45.06	18	1971	11	15	10	38	55	NCEER
3	-76.28	47.24	18	1971	11	22	5	29	7	NCEER
3	-76.62	45.83	18	1971	11	23	16	32	30	NCEER
4.1	-110.34	42.49	7	1971	12	3	7	44	59.2	SRA
4	-74.67	46.01	13	1971	12	18	15	36	24.5	NCEER
3.3	-75.6	39.7	0	1971	12	29	0	0	0	NCEER
3.7	-81.6	37.4	0	1972	1	9	23	24	29	NCEER
3.7	-90.85	36.37	3	1972	2	1	5	42	9.5	NCEER
4.5	-80.58	33.31	2	1972	2	3	23	11	9.7	NCEER
3.3	-75.6	39.7	0	1972	2	11	0	16	0.3	NCEER
4.3	-105.12	44.29	0	1972	2	18	11	43	38.1	NCEER
3.7	-89.74	36.12	7	1972	3	29	20	38	31.7	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.3	-75.99	46.67	18	1972	4	25	3	24	25	NCEER
3.4	-89.97	35.93	1	1972	5	7	2	12	8.7	NCEER
3.9	-82.2	37	0	1972	5	20	19	39	6	NCEER
3.1	-90.37	37.62	12	1972	6	9	19	15	18.9	NCEER
4.5	-89.08	37	13	1972	6	19	16	15	18.8	NCEER
3	-77.86	47.9	18	1972	7	17	1	58	46	NCEER
4.3	-104.93	49.35	5	1972	7	26	3	58	19	NCEER
3.1	-104.033	32.65	0	1972	7	26	4	35	40	SNMX
3.3	-75.6	39.7	0	1972	8	14	1	9	0	NCEER
3	-81.4	33.2	0	1972	8	14	15	5	19	NCEER
3.9	-66.47	49.54	18	1972	8	22	19	17	48	NCEER
3.4	-77.7	37.6	0	1972	9	5	16	0	0	NCEER
3.2	-77.56	46.18	18	1972	9	12	9	15	38	NCEER
4.4	-89.37	41.64	10	1972	9	15	5	22	15.9	NCEER
3.7	-99.6	42.3	0	1972	10	16	5	47	33	NCEER
3	-74.56	44.76	0	1972	11	2	5	15	8.8	SRA
3.3	-112.74	49.12	18	1972	11	21	6	8	46	DNAG
3.5	-76.24	40.14	2	1972	12	8	3	0	33.3	NCEER
4.1	-108.39	43.65	20	1972	12	8	18	47	39.4	SRA
3.9	-75.1	45.64	10	1972	12	16	19	1	37.2	NCEER
3.2	-87.22	37.4	14	1973	1	7	22	56	6.2	NCEER
3.5	-90.6	33.8	0	1973	1	8	9	11	37	NCEER
3.2	-90.48	37.89	17	1973	1	12	11	56	56.2	NCEER
3.1	-70	47.98	10	1973	1	28	13	7	50	NCEER
3.2	-110.425	36.43	5	1973	2	9	17	38	37	SRA
3	-70	44.5	0	1973	2	26	13	42	0	SRA
3.8	-75.43	39.69	12	1973	2	28	8	21	33.2	NCEER
3.4	-77.7	37.3	0	1973	4	9	23	11	0	NCEER
4.8	-107.85	42.64	33	1973	4	22	6	7	12.4	SRA
3.4	-90.8	33.9	0	1973	5	25	14	40	14	NCEER
3.1	-66.5	49.48	18	1973	6	14	15	9	55	NCEER
4.8	-71.12	45.31	12	1973	6	15	1	9	5.1	NCEER
3.3	-75.7	39.7	0	1973	7	10	4	38	0.2	NCEER
3.5	-74.47	43.87	2	1973	7	15	8	20	30.7	NCEER
3.1	-66.96	49.56	18	1973	7	20	17	6	39	NCEER
4.2	-104.57	37.15	5	1973	9	23	3	58	54.9	NCEER
3.4	-90.05	35.87	6	1973	10	3	3	50	19.8	NCEER
3.8	-89.62	36.49	3	1973	10	9	20	15	26.5	NCEER
3.5	-80.65	28.48	5	1973	10	27	6	21	2	NCEER
3.5	-84.12	35.76	1	1973	10	30	22	58	39	NCEER
3.1	-69.01	49.6	18	1973	11	15	17	31	35	NCEER
3.1	-70.29	47.55	10	1973	11	16	1	36	34	NCEER
3.1	-94.7	35	0	1973	11	18	10	3	53	NCEER
4.6	-83.99	35.89	12	1973	11	30	7	48	40.5	NCEER
3	-80.27	32.97	6	1973	12	19	10	16	8.7	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

(Page 43 of 65)

$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.1	-89.69	36.14	10	1973	12	20	10	45	0.9	NCEER
3.1	-98.3	29	0	1973	12	25	2	46	0	NCEER
3.9	-89.47	36.18	7	1974	1	8	1	12	38.1	NCEER
4.5	-100.69	36.4	0	1974	2	15	13	33	49.2	NCEER
3.8	-93.04	34.03	10	1974	2	15	22	49	4.4	NCEER
3	-67.09	49.54	18	1974	2	17	12	57	28	NCEER
3	-90.41	35.69	5	1974	3	4	14	24	28.1	NCEER
3.2	-89.8	35.64	5	1974	3	12	12	30	29.2	SRA
3.5	-107.05	40.7	5	1974	3	31	11	58	47.1	SRA
4.7	-88.07	38.55	15	1974	4	3	23	5	2.8	NCEER
3	-98	29	0	1974	4	20	23	46	10	SRA
3.2	-75.907	40.974	0	1974	4	27	14	45	39.9	SRA
3.3	-75.6	39.8	0	1974	4	28	14	19	0	NCEER
3.8	-89.36	36.74	4	1974	5	13	6	52	18.7	NCEER
3.6	-80.54	37.46	5	1974	5	30	21	28	35.3	NCEER
3.2	-84.75	38.48	10	1974	6	5	0	16	40.2	NCEER
3.2	-89.91	38.65	12	1974	6	5	8	6	10.7	NCEER
3.4	-98	29	0	1974	6	24	18	3	10	SRA
3.4	-67.22	49.58	10	1974	7	2	4	46	51	NCEER
3	-98	29	0	1974	8	1	13	33	10	SRA
4.1	-82.53	33.91	4	1974	8	2	8	52	11.1	NCEER
3.2	-76.08	45.93	18	1974	8	8	11	55	33	NCEER
3.2	-91.16	36.93	6	1974	8	11	14	29	45.4	NCEER
4.4	-107.38	44.11	10	1974	9	19	15	36	11.4	SRA
3	-83.49	41.21	1	1974	9	29	2	26	19.1	NCEER
3.1	-82.4	33.9	0	1974	10	8	23	22	28	NCEER
3.8	-81.61	39.06	4	1974	10	20	15	13	55.6	NCEER
3.2	-75.48	46.08	10	1974	10	23	22	52	57	NCEER
3	-81.92	33.79	0	1974	10	28	11	33	0	NCEER
3.2	-75.03	46.07	10	1974	11	2	13	47	56	NCEER
3.7	-82.22	33.73	0	1974	11	5	3	0	0	NCEER
4.3	-80.16	32.92	6	1974	11	22	5	25	56.7	NCEER
3.3	-79.11	43.33	0	1974	11	27	10	28	51.7	SRA
4	-104.017	32.633	0	1974	11	28	3	35	20	SNMX
3.5	-75.5	46.25	10	1974	12	2	10	58	5	NCEER
3.6	-82.5	33.95	0	1974	12	3	8	25	0	NCEER
3.2	-87.46	31.23	18	1974	12	10	6	1	35	NCEER
3.1	-91.86	34.49	3	1974	12	13	5	3	55.5	NCEER
3	-69.8	42.37	0	1974	12	22	20	46	48.7	SRA
3.5	-67.44	49.14	18	1974	12	27	0	50	12	NCEER
3.7	-103.11	30.92	5	1974	12	30	8	5	27.1	NCEER
3	-90.9	34.9	0	1975	1	2	9	19	0	NCEER
3.1	-74.6	44.9	0	1975	1	15	19	16	0	NCEER
4.1	-108.65	39.27	5	1975	1	30	14	48	40.3	SRA
3	-103.1	35.067	0	1975	2	2	20	39	23	SNMX

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

(Page 44 of 65)

$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.3	-83.2	41.3	0	1975	2	3	10	31	0	NCEER
3.4	-89.59	36.55	3	1975	2	13	19	43	58	NCEER
3	-82.35	38.88	4	1975	2	16	23	21	34.4	NCEER
3.5	-87.98	33.55	18	1975	3	1	11	50	0.2	NCEER
3	-80.48	37.32	5	1975	3	7	12	45	13.5	NCEER
4.8	-108.1	42.67	10	1975	3	25	14	59	58	SRA
3.1	-74.24	45.73	5	1975	4	3	19	3	17	NCEER
3	-80.22	33	10	1975	4	28	5	46	52.6	NCEER
3.3	-98.5	42.07	1	1975	5	13	7	53	40	NCEER
3.1	-103.7	43.2	0	1975	5	16	5	57	1	NCEER
3.2	-75.19	47.23	18	1975	5	29	21	19	16	NCEER
3.7	-108.8	41.91	5	1975	6	7	4	36	21.7	SRA
3.5	-73.65	44.87	11	1975	6	9	18	39	22.7	NCEER
3.9	-89.68	36.54	9	1975	6	13	22	40	27.5	NCEER
3.5	-89.44	36.21	7	1975	6	20	7	29	6.6	DNAG
3.7	-87.84	33.7	4	1975	6	24	11	11	36.6	NCEER
3	-79.77	43.4	10	1975	6	30	20	15	23	NCEER
4.6	-96.1	45.5	8	1975	7	9	14	54	21.3	NCEER
4.2	-76.31	46.54	17	1975	7	12	12	37	13.8	NCEER
3.1	-66.81	49.16	18	1975	7	18	4	21	6	NCEER
3.1	-70.18	47.44	5	1975	8	21	4	29	37	NCEER
3.6	-78.118	48.123	1	1975	8	21	21	42	51.1	PDE
3	-89.84	36.05	11	1975	8	25	7	11	8	NCEER
3	-65.34	46.8	18	1975	8	27	22	28	22	NCEER
4.4	-86.59	33.66	4	1975	8	29	4	22	52.1	NCEER
3.3	-69.74	48.29	2	1975	9	2	6	21	17	NCEER
3.8	-104.38	48.37	0	1975	9	5	20	47	40.7	NCEER
3	-89.3	30.7	0	1975	9	9	11	52	44	NCEER
3.2	-97.22	34.13	5	1975	9	13	1	25	5.6	NCEER
3.2	-97.7	35.5	5	1975	10	12	2	58	11.5	NCEER
3.1	-65.89	45.11	18	1975	10	15	3	26	17	NCEER
3.3	-83	34.9	0	1975	10	18	4	31	0	NCEER
3.1	-68.13	49.13	18	1975	10	21	20	50	2	NCEER
4.1	-68.62	49.83	2	1975	10	23	21	17	48.7	NCEER
4	-74.65	43.91	5	1975	11	3	20	54	55.3	NCEER
3.5	-87.33	33.31	4	1975	11	7	23	39	31.7	NCEER
3.2	-80.89	37.22	1	1975	11	11	8	10	37.6	NCEER
3.2	-82.9	34.93	10	1975	11	25	15	17	34.8	NCEER
3.5	-97.42	34.68	14	1975	11	29	14	29	44.9	NCEER
3.3	-94.62	38.24	0	1975	12	4	18	59	59.9	SRA
3.6	-78.89	47.01	0	1975	12	19	15	25	0.1	NCEER
3.5	-107.65	42.85	0	1975	12	19	23	26	19.5	SRA
4.6	-108.212	35.817	0	1976	1	5	6	23	33.9	SNMX
3.4	-92.16	35.9	7	1976	1	16	19	42	56.9	NCEER
3.3	-103.1	31.9	0	1976	1	19	4	3	30	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.8	-83.86	36.87	1	1976	1	19	6	20	39.6	NCEER
3.3	-100.087	31.9	4	1976	1	25	4	48	28.5	SRA
3.4	-83.73	41.88	5	1976	2	2	21	14	2.3	NCEER
3.6	-84.7	34.97	14	1976	2	4	19	53	53	NCEER
3.5	-71.21	41.56	0	1976	3	11	8	29	32.2	NCEER
3.1	-69.97	41.66	0	1976	3	14	23	12	24.6	NCEER
3.1	-95.6	35.43	5	1976	3	16	7	39	45.3	NCEER
3.5	-104.27	49.39	5	1976	3	25	0	12	16	NCEER
4.9	-90.48	35.58	17	1976	3	25	0	41	20.8	NCEER
3.3	-67.86	49.34	18	1976	3	29	21	23	27	NCEER
3	-86.7	39.3	0	1976	4	8	7	38	53	NCEER
3.2	-74.03	40.8	0	1976	4	13	15	39	12.9	NCEER
3.3	-87.31	37.38	4	1976	4	15	7	3	34.4	NCEER
3.5	-99.79	36.04	8	1976	4	19	4	42	46.9	NCEER
3.5	-109.1	35.39	5	1976	4	19	23	35	45.5	SRA
3	-103.14	32.27	0	1976	5	1	11	13	40.8	SRA
3.1	-73.9	49.56	18	1976	5	5	3	1	4	NCEER
3.1	-79.9	39.6	0	1976	5	6	18	46	8.1	NCEER
3.3	-68.62	49.84	3	1976	5	15	21	6	52	NCEER
3.2	-89.83	36.03	9	1976	5	22	7	40	46.1	NCEER
3	-104.02	37.41	5	1976	5	30	1	43	37.3	SRA
3.3	-81.6	37.34	1	1976	6	19	5	54	13.4	NCEER
3	-103.283	35.617	0	1976	6	24	15	27	32	SNMX
3.1	-74.1	45.18	9	1976	7	13	3	51	14	NCEER
3.1	-110.3	40.75	7	1976	7	30	22	19	0.2	SRA
3	-103.02	31.57	0	1976	8	5	18	53	9	SRA
3.1	-74.98	49.77	18	1976	8	7	7	50	11	NCEER
3.4	-106.57	45.03	5	1976	8	10	13	54	57.5	SRA
4.5	-106.15	44.04	10	1976	9	3	4	18	16.2	SRA
4.3	-80.77	36.62	9	1976	9	13	18	54	38	NCEER
3	-103.1	32.21	0	1976	9	17	2	47	45.4	SRA
3.1	-102.5	31.4	0	1976	9	17	3	56	29	DNAG
3.4	-67.1	49.36	18	1976	9	18	0	40	32	NCEER
3.5	-90.47	35.58	8	1976	9	25	14	6	55.8	NCEER
3.5	-106.57	45.03	0	1976	10	8	13	54	0	SRA
3	-97.06	35.38	5	1976	10	22	17	15	50.5	NCEER
3	-78.05	48.16	18	1976	10	22	18	50	56	NCEER
3.1	-88.98	32	10	1976	10	23	0	40	59.2	NCEER
4.2	-69.78	47.82	18	1976	10	23	20	58	18	NCEER
3	-75.96	47.11	18	1976	11	6	6	9	29	NCEER
4.2	-91.04	38.1	0	1976	12	11	7	5	1.1	NCEER
3.5	-90.26	37.81	9	1976	12	13	8	35	55.1	NCEER
3.7	-82.5	32.06	14	1976	12	27	6	57	15.2	NCEER
3.6	-89.71	37.58	5	1977	1	3	22	56	48.5	NCEER
3	-80.17	33.06	1	1977	1	18	18	29	14.1	NCEER



**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
4.5	-66.73	46.88	18	1977	2	6	9	1	19	DNAG
3.1	-70.42	47.54	8	1977	2	14	0	35	4.1	NCEER
3.4	-78.63	37.9	0	1977	2	27	20	5	34.6	NCEER
3.1	-109.903	44.608	1	1977	3	2	18	7	24.2	DNAG
3.8	-107.15	41.24	5	1977	3	3	17	50	28	SRA
4.2	-108.222	35.748	0	1977	3	5	3	0	55.8	SNMX
3.7	-107	44.6	0	1977	3	24	8	55	0	DNAG
3.3	-103.1	31.9	0	1977	4	26	9	3	7	NCEER
3.3	-88.44	31.96	0	1977	5	4	2	0	24.3	NCEER
3.6	-94.17	34.56	10	1977	6	2	23	29	10.6	NCEER
3.2	-84.71	40.71	1	1977	6	17	15	39	46.9	NCEER
3.1	-70.16	47.77	8	1977	6	20	5	5	54.7	NCEER
3.4	-74.38	46.04	4	1977	7	14	7	39	29.8	NCEER
3	-102.7	31.8	0	1977	7	22	4	1	10	DNAG
3.5	-84.41	35.42	5	1977	7	27	22	3	20.8	NCEER
3.1	-80.7	33.37	9	1977	8	4	4	20	7.7	NCEER
3.9	-67.05	49.77	18	1977	8	8	23	8	40	NCEER
3.1	-80.69	33.39	0	1977	8	25	4	20	7	NCEER
3.5	-107.31	39.31	5	1977	9	24	11	16	48.4	SRA
4.8	-110.47	40.47	6	1977	9	30	10	19	20.4	USHIS
3	-73.82	46.52	1	1977	10	16	21	29	19.1	NCEER
3	-67.05	47	18	1977	10	24	18	9	12	NCEER
3.4	-89.173	33.928	16	1977	11	4	11	21	10.2	SRA
3	-75.15	46.27	5	1977	11	7	20	48	52.7	NCEER
3	-75.86	46.69	1	1977	11	25	18	47	24.4	NCEER
3.1	-92.91	34.39	10	1977	11	26	4	18	18.1	NCEER
3	-80.18	32.88	9	1977	12	15	19	16	43.1	NCEER
3.1	-70.68	41.79	0	1977	12	20	17	44	23.8	NCEER
3.5	-76.91	46.84	11	1977	12	22	14	57	1.3	NCEER
3.2	-71.64	43.2	0	1977	12	25	15	35	53.5	NCEER
3.2	-70.55	44.07	9	1978	1	4	19	28	10.8	NCEER
3.1	-88.21	32.7	1	1978	1	8	11	34	23.4	NCEER
3.8	-81.6	28.1	0	1978	1	12	21	10	0	DNAG
3	-105.31	42.43	5	1978	1	16	3	50	3.1	SRA
3	-88.01	38.244	5	1978	1	28	16	40	58.8	DNAG
3.3	-109.7	42.5	30	1978	2	7	5	3	10.4	SRA
4.1	-74.11	46.35	7	1978	2	18	14	48	25	NCEER
3.5	-102.5	31.55	1	1978	3	2	10	4	53	NCEER
3.1	-90	36.63	9	1978	4	3	12	24	21.5	NCEER
3	-69.89	47.751	22	1978	4	8	8	21	45	PDE
3.1	-78.24	39.7	15	1978	4	26	19	30	23.3	SRA
3.6	-101.95	42.26	38	1978	5	7	16	6	23	NCEER
3.2	-69.99	47.72	3	1978	5	26	2	31	0.4	NCEER
3	-107.32	39.28	5	1978	5	29	16	45	18	SRA
3.2	-88.46	38.41	20	1978	6	2	2	7	28.9	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.8	-107.83	43.63	5	1978	6	6	21	23	34.7	SRA
3.3	-88.595	32.042	2	1978	6	9	23	15	19.6	SRA
3.4	-101.94	31.05	0	1978	6	29	20	58	45.1	DNAG
3.2	-68.39	48.75	18	1978	6	30	0	17	0	NCEER
3.1	-76.22	39.9	0	1978	7	16	6	39	29.7	NCEER
3.1	-105.04	34.68	0	1978	7	21	5	2	36.2	SNMX
3.4	-65.61	49.3	18	1978	7	29	13	56	43	NCEER
3.6	-74.44	45.68	7	1978	7	30	10	54	44	NCEER
3	-70.22	47.669	18	1978	8	14	22	55	58	PDE
3.1	-74.51	44.52	0	1978	8	21	8	47	0	NCEER
3	-76.43	47.17	5	1978	8	26	3	54	34.1	NCEER
3.5	-111.48	48.49	5	1978	8	30	16	33	21.2	SRA
3.5	-89.44	36.09	1	1978	8	31	0	31	0.6	NCEER
3	-90.28	38.58	1	1978	9	20	12	24	8.9	NCEER
3.1	-91.92	33.96	33	1978	9	23	7	34	3.7	NCEER
3	-76.15	40.08	0	1978	10	6	19	25	47.4	NCEER
3.8	-82.65	30.2	0	1978	11	6	23	0	0	DNAG
3.3	-67.62	48.96	18	1978	12	2	8	36	44	NCEER
3.5	-88.37	38.56	23	1978	12	5	1	48	2	NCEER
3.5	-88.47	31.91	3	1978	12	11	2	6	50.1	NCEER
3.3	-107.86	40.82	5	1979	1	20	6	59	8.4	SRA
3	-74.26	40.32	0	1979	1	30	16	30	52	NCEER
3.2	-90.1	35.84	10	1979	2	5	5	31	9.4	NCEER
3.3	-113.299	49.233	5	1979	2	24	15	49	15.5	PDE
3.4	-91.2	35.96	10	1979	2	27	22	54	54.8	NCEER
3.2	-74.5	40.72	0	1979	3	10	4	49	39.6	NCEER
3	-76.46	45.84	18	1979	3	18	16	31	12	NCEER
3.1	-108.9	40.18	2	1979	3	19	14	59	29.7	SRA
3.2	-70.1	47.69	10	1979	3	23	22	53	5	NCEER
3.2	-112.41	48.59	5	1979	4	14	9	39	6.4	SRA
3.1	-95.54	46.7	20	1979	4	16	6	40	16.7	SRA
4	-69.79	43.97	17	1979	4	18	2	34	15.3	NCEER
3.2	-66.03	45.24	18	1979	4	20	10	32	49.2	NCEER
3.1	-71.24	43.04	0	1979	4	23	0	5	45.7	NCEER
3.8	-111.02	37.88	7	1979	4	30	2	7	10.3	SRA
3.3	-75.68	46.18	1	1979	5	26	21	58	32.8	NCEER
3	-74.99	45	2	1979	5	29	20	48	49.1	NCEER
3.3	-67.54	49.37	18	1979	6	5	8	58	20	NCEER
3	-99.76	35.22	2	1979	6	7	7	39	36.3	NCEER
3.1	-73.86	44.43	0	1979	6	7	13	45	53.3	NCEER
3.8	-89.64	36.15	15	1979	6	11	4	12	17.1	NCEER
3	-110.904	37.861	7	1979	6	16	1	8	44.7	DNAG
3	-74.38	41.35	0	1979	6	20	19	20	17.8	SRA
3	-90.45	35.56	7	1979	6	25	17	11	13.8	NCEER
3	-97.29	39.92	7	1979	6	30	20	46	42.3	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.8	-76.6	46.87	18	1979	7	8	1	29	18	NCEER
3.1	-89.31	36.91	2	1979	7	8	12	35	15.5	NCEER
3.7	-74.65	46.54	18	1979	7	9	8	16	26	NCEER
3.2	-100.32	40.18	4	1979	7	16	0	3	48.4	NCEER
3.1	-110.591	37.592	7	1979	7	25	23	56	11.7	DNAG
3.5	-70.44	43.29	11	1979	7	28	23	29	12	NCEER
3	-81.358	34.333	3	1979	8	7	19	32	17.2	SRA
3.8	-111.47	48.49	5	1979	8	9	17	12	55.4	SRA
3.7	-84.36	35.21	10	1979	8	13	5	18	56.8	NCEER
4.6	-69.9	47.67	10	1979	8	19	22	49	30.4	NCEER
3.7	-82.956	34.916	1	1979	8	26	1	31	45	USHIS
3.8	-91.5	36.3	0	1979	8	26	11	28	0	DNAG
3.2	-83.24	35.3	10	1979	9	6	20	38	16.3	NCEER
3.2	-83.91	35.58	12	1979	9	12	6	24	4	NCEER
3.4	-99.47	35.19	1	1979	9	13	0	49	21.5	NCEER
3.6	-82.08	36.44	5	1979	10	8	8	53	52.8	SRA
3.5	-110.93	37.89	7	1979	10	23	4	17	19.9	SRA
3.2	-91.04	36.46	6	1979	11	5	16	35	25.9	NCEER
3.6	-82.81	38.49	1	1979	11	9	21	29	59.8	NCEER
3.3	-98.41	35.63	5	1979	11	27	9	10	36.7	NCEER
3.8	-66.72	49.43	18	1979	12	19	18	58	10	NCEER
3.5	-91.22	49.62	18	1980	2	27	6	13	41	NCEER
3.1	-74.2	42.58	12	1980	2	29	5	53	56.1	SRA
3.7	-71.87	46.79	18	1980	3	11	4	15	55	NCEER
3.3	-75.09	40.15	0	1980	3	11	6	0	26.9	NCEER
3	-88.44	37.89	20	1980	3	13	2	23	13	NCEER
3.3	-86.76	37.6	9	1980	3	23	21	38	16.2	NCEER
4	-67.95	48.77	18	1980	4	3	16	57	24	NCEER
3.2	-68.36	44.71	0	1980	4	10	15	36	43.8	NCEER
4.1	-81.64	49.64	18	1980	4	13	22	40	23	NCEER
3.6	-112.34	48.79	5	1980	4	14	3	27	33.8	SRA
3	-81.324	34.329	3	1980	4	24	6	16	57.2	SRA
3	-75	40.3	0	1980	5	2	19	2	0	NCEER
3	-75.25	45.26	19	1980	5	19	23	40	50	DNAG
3.4	-74.55	44.89	0	1980	5	23	8	39	44	NCEER
3.5	-75.23	43.56	1	1980	6	6	13	15	51.9	NCEER
3.4	-101.01	35.48	1	1980	6	9	22	37	12.3	NCEER
3	-82.81	35.46	1	1980	6	10	23	47	32.2	NCEER
3.3	-84.03	35.73	1	1980	6	25	18	2	1.6	NCEER
3.4	-70.75	47.56	10	1980	7	1	3	6	38	NCEER
3.1	-70.33	47.3	10	1980	7	2	7	50	33	NCEER
3.5	-89.6	36.56	4	1980	7	5	8	54	40.1	NCEER
3.2	-99.7	35.18	5	1980	7	18	14	29	46.8	NCEER
3.1	-74.17	45.14	5	1980	7	25	6	22	34	NCEER
3.1	-87.44	33.94	0	1980	7	25	15	30	12.5	SRA

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
5.2	-83.89	38.19	16	1980	7	27	18	52	21.4	NCEER
3.2	-81.364	34.351	1	1980	7	29	1	10	22.7	SRA
3	-82.04	49.9	18	1980	7	31	10	10	4	NCEER
3.1	-74.15	40.43	8	1980	8	2	17	20	59.7	SRA
3.3	-75.16	43.54	0	1980	8	11	14	54	46.1	NCEER
3.2	-82.99	41.99	1	1980	8	20	9	34	53.4	NCEER
3.1	-84.87	37.98	1	1980	8	23	3	49	3.7	NCEER
3	-74.9	39.8	0	1980	8	30	9	19	0	NCEER
3.2	-73.78	41.11	13	1980	9	4	4	30	55.8	SRA
3.2	-69	44.67	8	1980	9	8	5	59	55.2	NCEER
3.2	-105.12	41.18	0	1980	9	12	22	33	55.4	SRA
3.2	-74.02	43.63	0	1980	9	21	20	52	45.1	DNAG
3	-69.9	47.67	6	1980	9	30	18	26	1	NCEER
3.4	-80.57	43.15	5	1980	10	14	0	58	56.4	NCEER
3.1	-72.9	41.3	0	1980	10	24	17	27	38.2	NCEER
3	-97.76	35.46	1	1980	11	2	10	0	48.9	NCEER
3	-79.9	38.18	4	1980	11	5	21	48	14.7	NCEER
3.8	-89.43	36.17	5	1980	12	2	8	59	29.7	NCEER
3.4	-78.44	37.72	6	1981	2	11	13	44	16.4	NCEER
3.8	-91.8	30	0	1981	2	13	2	15	0	DNAG
3.3	-74.93	45.96	18	1981	2	19	7	7	10	NCEER
4.3	-104.96	39.91	8	1981	4	2	16	10	6.4	NCEER
3.5	-89.38	38.87	1	1981	4	8	1	53	13	SRA
3.3	-82.05	35.51	0	1981	4	9	7	10	31.2	NCEER
3.7	-65.7	45.93	18	1981	4	13	17	31	38	NCEER
3.5	-82.42	35.33	10	1981	5	5	21	21	56.7	NCEER
3	-91.63	36.76	1	1981	5	25	22	50	18.2	SRA
3	-110.37	36.83	1	1981	5	29	3	9	2.2	SRA
3	-81.67	36.18	1	1981	6	3	20	54	22.4	SRA
3.2	-94.32	31.99	5	1981	6	9	1	46	32.7	NCEER
3.4	-89.03	37.82	19	1981	6	9	14	15	47.8	NCEER
3.8	-89.9	43.9	0	1981	6	12	15	30	0	DNAG
3.7	-70	47.47	8	1981	6	16	17	55	4	NCEER
3.5	-90.07	35.85	9	1981	6	26	8	33	27	NCEER
3.1	-71.55	43.57	0	1981	6	28	22	42	35	NCEER
3.7	-74.62	45.14	13	1981	7	4	23	16	32	NCEER
3.5	-97.73	34.85	5	1981	7	11	21	9	21.8	NCEER
3.7	-66.8	49.82	18	1981	7	13	4	48	4	NCEER
3	-110.31	36.82	0	1981	7	14	19	29	51	SRA
4	-89.18	36.03	11	1981	8	7	11	53	44	NCEER
3	-73.54	44.07	1	1981	8	10	23	6	59.3	DNAG
3.3	-72.24	48.66	5	1981	8	23	23	17	20	NCEER
3.3	-80.59	43.15	1	1981	8	28	10	51	33	NCEER
3	-85.17	34.63	3	1981	9	4	17	21	44.5	NCEER
3.1	-81.41	42.8	9	1981	9	5	5	49	21	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

(Page 50 of 65)

$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.1	-100.52	42.89	5	1981	9	7	0	38	9.1	SRA
3.1	-110.56	37.5	2	1981	9	10	7	55	9	SRA
3.4	-101.85	43.04	5	1981	9	13	22	16	29.7	NCEER
3.1	-66.11	49.53	18	1981	9	18	2	24	11	NCEER
3.5	-75.02	46.11	0	1981	9	18	7	16	7	NCEER
3.5	-75.56	46.37	0	1981	9	30	23	41	39	NCEER
3.3	-98.54	41.17	5	1981	10	9	21	54	27.8	NCEER
3	-112.92	49.29	18	1981	10	20	3	47	36	DNAG
3.7	-72.57	41.14	5	1981	10	21	16	49	6.9	NCEER
3.9	-65.25	49.83	18	1981	10	28	19	56	14	NCEER
3.2	-95.26	32.02	5	1981	11	6	12	36	40.5	NCEER
3	-89.39	36.09	12	1981	11	8	17	11	19	NCEER
3.4	-77.04	46.98	18	1981	11	12	18	40	14	NCEER
3.7	-66.61	47.03	5	1981	11	28	5	12	3	NCEER
3.3	-72.64	45.38	3	1981	12	6	16	11	27	NCEER
3.1	-86.43	35.18	13	1982	1	2	2	0	26.2	NCEER
3.9	-102.49	31.18	5	1982	1	4	16	56	8.1	NCEER
5.7	-66.6	47	5	1982	1	9	12	53	52	NCEER
4.7	-71.62	43.51	7	1982	1	19	0	14	42.6	NCEER
4.3	-92.22	35.22	0	1982	1	24	3	22	44.4	NCEER
3.3	-70.38	47.45	6	1982	1	27	1	35	56	NCEER
3	-70.94	41.87	0	1982	1	27	18	50	4.6	NCEER
3.4	-81.39	32.98	7	1982	1	28	4	52	51.9	SRA
3	-67.48	49.18	18	1982	1	30	15	44	35	DNAG
3.1	-90.06	35.92	10	1982	2	2	9	26	46.3	NCEER
3.3	-104.03	48.51	18	1982	3	9	13	10	50.1	NCEER
3	-82.48	46.65	1	1982	3	13	4	34	32	NCEER
3.1	-103.27	35.36	5	1982	3	16	11	3	2.7	NCEER
3	-79.88	46.3	18	1982	3	19	16	48	13	NCEER
3.4	-113.25	49.07	18	1982	3	21	21	43	21	DNAG
3	-98.46	29.85	5	1982	3	28	23	24	32.9	NCEER
3	-82.04	36.51	3	1982	4	13	13	4	13.3	NCEER
3	-111.3	38.22	9	1982	4	17	6	0	12.5	SRA
3.5	-92.24	35.18	0	1982	4	21	21	17	55	NCEER
3.1	-96.47	33.99	5	1982	5	3	7	54	48.6	NCEER
3.4	-109.7	44.64	11	1982	5	9	21	7	36.5	SRA
3	-77.96	40.41	0	1982	5	12	18	29	33	SRA
3.5	-92.23	35.2	2	1982	5	31	18	21	19.7	NCEER
3.5	-76.95	47.38	18	1982	6	23	0	22	0	NCEER
3.5	-92.21	35.22	6	1982	7	5	4	13	52	NCEER
3.6	-96.72	44.01	5	1982	7	11	19	42	28.4	NCEER
3.8	-74.55	46.09	17	1982	7	13	2	18	49	NCEER
3	-69.02	46.08	6	1982	7	15	7	27	55.4	SRA
3.1	-81.55	34.32	2	1982	7	16	14	16	2.9	SRA
3.7	-75.46	45.89	19	1982	8	6	6	29	10	NCEER

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

(Page 51 of 65)

$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.2	-92.24	35.19	4	1982	8	9	11	12	31.6	SRA
3.1	-88.73	37.25	5	1982	8	11	10	32	38.8	NCEER
3.7	-78.61	46.67	18	1982	8	13	1	6	42	NCEER
4.3	-105.38	49.06	18	1982	8	17	4	49	25	NCEER
3.4	-70.38	47.37	20	1982	8	29	2	7	11	NCEER
3	-74.19	43.2	5	1982	8	31	10	16	28.4	NCEER
3.2	-108.85	42.72	5	1982	8	31	22	2	18.5	SRA
3	-82.9	34.96	3	1982	9	2	21	52	45.5	SRA
3.7	-76.61	45.67	12	1982	9	3	23	14	3	NCEER
3.2	-84.51	35.19	13	1982	9	5	10	11	9.4	NCEER
3.4	-84.25	35.68	8	1982	9	24	22	19	16.9	NCEER
3.5	-92.23	35.21	5	1982	9	25	23	17	5.5	SRA
3	-73.057	43.125	8	1982	9	28	22	24	12.5	DNAG
3.9	-102.57	36.1	5	1982	10	14	12	52	46.3	NCEER
3.3	-65.3	49.82	18	1982	10	29	21	50	57	NCEER
3.1	-84.89	32.64	0	1982	10	31	3	12	12.2	NCEER
3.3	-108.695	35.305	0	1982	11	3	17	54	1.9	SNMX
3.1	-100.2	35.2	0	1982	11	7	0	4	19	SRA
4.3	-97.85	43.01	5	1982	11	15	2	58	22.9	NCEER
3.5	-92.08	35.25	0	1982	11	21	16	35	31	NCEER
3	-73.43	45.34	5	1982	11	24	7	34	39	PDE
3	-71.52	43.62	6	1982	12	1	22	52	22.9	SRA
3.9	-70.22	47.54	16	1982	12	4	16	8	32	NCEER
3	-83.53	32.85	0	1982	12	11	0	25	0	NCEER
3.3	-78.83	46.82	18	1983	1	10	21	31	27	NCEER
4.1	-67.06	49.11	18	1983	1	17	19	35	52	NCEER
3.9	-92.16	35.28	0	1983	1	19	2	30	42	NCEER
3.1	-83.45	48.72	18	1983	1	20	9	16	45	NCEER
3.1	-67.86	47.46	15	1983	1	20	14	17	21	NCEER
3.3	-81.02	41.75	10	1983	1	22	7	46	58	NCEER
3.5	-83.56	32.85	0	1983	1	26	14	7	44.7	NCEER
3.1	-83.63	36.06	13	1983	1	27	22	9	35.1	NCEER
3.3	-110.674	37.778	7	1983	1	27	23	37	11.8	SRA
3.2	-88.31	34.73	0	1983	2	5	13	8	19	NCEER
3.5	-68.33	48.98	18	1983	2	11	15	46	56	NCEER
4	-105.729	42.232	5	1983	2	13	13	44	44	SRA
3.5	-112.373	48.539	14	1983	2	16	6	22	9.3	SRA
3.6	-89.6	36.19	1	1983	2	23	8	51	27	NCEER
3	-73.66	41.55	7	1983	2	26	19	59	35.4	NCEER
4.4	-99.41	44.21	5	1983	3	4	6	32	18.6	NCEER
3.4	-71.72	42.96	1	1983	3	24	14	27	20.4	NCEER
3.3	-82.46	35.33	12	1983	3	25	2	47	11.1	NCEER
3.2	-92.15	35.2	0	1983	3	30	4	15	26	NCEER
3.4	-102.38	35.32	5	1983	4	3	4	55	24.2	NCEER
3	-66.98	49.34	18	1983	4	13	16	6	53	DNAG

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3	-110.633	38.305	2	1983	5	3	12	43	37.7	SRA
3.3	-102.198	42.955	5	1983	5	6	6	14	46.9	SRA
3.9	-66.6	47	5	1983	5	13	23	40	57	NCEER
4.3	-89.57	38.77	0	1983	5	15	5	16	22	NCEER
3.8	-69.89	47.7	11	1983	5	16	2	1	57	NCEER
3	-92.36	38.48	0	1983	5	16	14	3	4	NCEER
3.7	-69.46	45.54	10	1983	5	27	23	4	35.2	NCEER
4.4	-70.4	44.49	3	1983	5	29	5	45	49.9	NCEER
3.4	-70.22	47.45	10	1983	6	2	6	30	23	NCEER
3	-69.65	47.46	10	1983	6	4	5	0	23	NCEER
3.5	-66.68	47.04	5	1983	6	28	8	5	49	NCEER
3	-90.94	37.1	0	1983	7	8	9	41	40	NCEER
3.4	-84.15	35.55	10	1983	7	8	19	29	5.9	NCEER
3	-74.91	46.06	18	1983	7	17	22	47	45	NCEER
3	-70.95	46.53	10	1983	7	23	3	25	38	NCEER
3.4	-98.131	28.743	5	1983	7	23	15	24	38.2	SRA
3.7	-67.68	44.97	12	1983	8	12	14	8	47.6	NCEER
3.5	-82.77	38.47	10	1983	8	17	14	3	15	NCEER
3.4	-104.314	37.469	5	1983	8	17	15	3	27.6	SRA
3.1	-83.82	36.68	18	1983	8	28	22	45	7.4	NCEER
3.1	-104.43	34.922	0	1983	9	15	23	25	37.5	SNMX
4.1	-108.837	40.789	5	1983	9	24	16	57	45.7	SRA
3.1	-79.79	43.44	2	1983	10	4	17	18	40	NCEER
5.2	-74.31	44.03	7	1983	10	7	10	18	47	NCEER
4.2	-75.77	45.21	15	1983	10	11	4	10	55	NCEER
3.1	-75.05	45.62	11	1983	10	16	3	0	47	NCEER
3.8	-93.39	30.24	5	1983	10	16	19	40	50.8	NCEER
3.2	-66.31	47.21	5	1983	10	17	22	58	56	NCEER
3.1	-77.97	48.14	1	1983	10	24	1	0	6	NCEER
3.5	-73.9	45.68	18	1983	11	1	10	16	52	NCEER
3.3	-80.16	32.94	10	1983	11	6	9	2	19.8	NCEER
3	-105.955	43.016	5	1983	11	15	12	33	12.1	SRA
3.8	-66.6	47	5	1983	11	17	15	32	18	NCEER
3.3	-69.16	45.19	2	1983	12	4	10	48	33.6	NCEER
3	-67.17	45.11	7	1983	12	8	12	23	5	NCEER
3	-92.704	33.183	5	1983	12	9	20	52	10.5	SRA
3.2	-76.29	46.69	18	1983	12	14	1	52	3	NCEER
3.1	-73.97	45.24	18	1983	12	21	15	4	44	NCEER
3.4	-76.33	47.01	18	1983	12	28	12	24	21	NCEER
3	-89.75	37.59	2	1984	1	12	2	48	15.7	SRA
3.5	-67.16	44.88	18	1984	1	14	9	9	32	NCEER
3	-83.43	41.65	0	1984	1	14	20	14	31	NCEER
3.3	-110.845	47.149	5	1984	1	16	19	50	25.2	SRA
3.1	-75.12	45.56	19	1984	1	17	19	4	46	NCEER
3.2	-89.92	36.61	1	1984	1	28	21	29	22.1	SRA

Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.2	-97.36	34.67	5	1984	2	3	4	38	28	NCEER
3.6	-83.74	36.13	10	1984	2	14	20	54	30.9	NCEER
3.6	-89	37.21	0	1984	2	14	22	56	10	NCEER
3.7	-66.6	47	5	1984	2	24	3	17	14	PDE
3.2	-108.638	41.539	2	1984	3	1	18	13	0.9	SRA
3.8	-98.461	28.852	5	1984	3	3	1	3	26.5	SRA
3	-84.05	35.83	7	1984	3	17	23	26	11.4	NCEER
3	-66.49	46.91	5	1984	3	27	22	56	24	DNAG
3.2	-66.46	49.61	18	1984	3	29	22	52	50	DNAG
3.4	-102.4	35.32	0	1984	4	3	4	55	24	SRA
3.8	-67.52	49.3	18	1984	4	11	19	7	42	DNAG
3.1	-66.6	47	5	1984	4	13	15	35	51	DNAG
3.4	-88.44	38.38	0	1984	4	17	4	44	44	NCEER
3.1	-107.19	39.281	5	1984	4	22	17	30	56.7	SRA
4.1	-76.37	39.95	4	1984	4	23	1	36	0	NCEER
3.2	-107.228	39.322	5	1984	5	14	10	14	17.2	SRA
3.4	-102.4	35.4	0	1984	5	21	13	30	14	SRA
3.1	-102.228	35.067	5	1984	5	21	13	31	13.5	SRA
3.6	-102.155	39.22	5	1984	5	27	23	30	19.3	SRA
3.2	-66.33	49.6	18	1984	5	28	21	4	52	DNAG
3.5	-80.78	46.63	1	1984	6	20	16	10	22	PDE
3.2	-89.39	36.1	12	1984	6	26	15	15	19.9	NCEER
3.1	-75.68	46.23	0	1984	6	28	3	8	49	DNAG
3.8	-88.47	37.7	2	1984	6	29	7	58	29.3	NCEER
3	-66.6	47	5	1984	7	2	5	24	54	DNAG
4.1	-81.17	46.53	1	1984	7	6	17	24	52	PDE
3	-89.53	36.5	7	1984	7	16	3	50	53.5	NCEER
4	-87.07	39.22	10	1984	7	28	23	39	27.4	NCEER
3	-90.92	37.82	7	1984	7	30	7	33	46.5	NCEER
3	-67.05	45.32	18	1984	8	3	13	41	11	DNAG
3	-98.362	29.133	5	1984	8	8	1	31	27.3	SRA
3.2	-86.3	34.62	8	1984	8	9	2	42	35.8	NCEER
4.2	-78.324	37.868	8	1984	8	17	18	5	46.9	SRA
3.2	-73.48	44.875	11	1984	8	20	10	58	17	SRA
3.1	-87.45	39.11	10	1984	8	29	6	50	59.5	NCEER
3.1	-84.34	35.57	13	1984	8	30	16	26	28.4	NCEER
5	-106.11	44.138	15	1984	9	8	0	59	31.1	USHIS
3.2	-100.697	31.991	5	1984	9	11	14	47	33.5	SRA
3.2	-108.582	41.61	2	1984	9	14	19	4	26.3	SRA
3.4	-92.21	35.25	5	1984	9	27	13	3	6	NCEER
3	-91.7	35.72	5	1984	9	27	13	16	22.9	NCEER
4.2	-85.2	34.75	12	1984	10	9	11	54	26.9	NCEER
3	-66.59	47.08	5	1984	10	13	1	45	15	DNAG
3.3	-65.66	44.72	18	1984	10	13	12	53	45	DNAG
5.4	-105.735	42.317	22	1984	10	18	15	30	22	USHIS



**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.2	-81.68	36.36	11	1984	10	22	18	58	41.7	NCEER
3.5	-73.93	43.59	0	1984	10	23	6	26	21.9	NCEER
4.7	-108.919	42.534	5	1984	11	3	9	30	8.4	USHIS
3.1	-97.41	34.71	5	1984	11	20	10	57	31.9	NCEER
3.2	-75.05	45.19	14	1984	11	26	9	3	49	PDE
3.7	-66.58	46.98	14	1984	11	30	5	54	22	PDE
3	-89.7	36.16	11	1984	12	3	11	55	44.6	NCEER
3.6	-66.04	47.52	1	1984	12	9	18	12	21	DNAG
3.5	-82.6	46.5	0	1984	12	17	9	38	36	PDE
3	-70.25	47.4	19	1984	12	22	12	46	30	DNAG
3	-89.91	35.93	9	1985	1	30	9	35	12.4	SRA
3.1	-89.51	36.29	7	1985	2	7	23	44	35.3	NCEER
3	-87.5	38.42	3	1985	2	13	10	22	24	SRA
3.3	-89.34	37.23	6	1985	2	15	15	56	9.9	NCEER
3.2	-70.48	47.39	14	1985	3	3	12	15	17	PDE
3.2	-105.85	38.558	5	1985	3	16	21	55	2.4	SRA
3.1	-69.96	47.52	12	1985	4	10	5	52	57	DNAG
3.1	-70.704	45.364	2	1985	4	12	5	27	30.5	SRA
3.2	-80.4	41.59	18	1985	4	14	11	39	54	DNAG
3.4	-108.92	35.26	0	1985	4	14	21	48	2.9	SNMX
3.1	-90.77	36.27	9	1985	5	4	7	7	12.5	SRA
3.2	-75.9	46.83	18	1985	5	16	13	39	7	DNAG
3.2	-80.485	37.248	11	1985	6	10	12	22	38.3	SRA
3.6	-82.038	37.222	1	1985	6	19	22	28	8.9	SRA
3	-85.156	35.198	3	1985	7	12	18	20	28.4	SRA
3.3	-92.202	35.219	5	1985	8	3	4	23	11	DNAG
3.5	-108.649	41.817	5	1985	8	13	20	57	0.8	SRA
4.3	-108.06	42.813	10	1985	8	16	6	5	22.6	SRA
3.1	-67.67	49.3	0	1985	8	16	22	48	37	DNAG
3	-110.232	46.109	5	1985	8	22	2	12	5	SRA
3.1	-76.64	45.67	18	1985	8	24	6	4	2	PDE
3.6	-93.118	35.809	10	1985	9	6	22	17	2.8	SRA
3	-88.014	41.848	5	1985	9	9	22	6	31	SRA
3.3	-97.051	33.548	5	1985	9	18	15	54	4.6	SRA
4	-66.6	47	5	1985	10	5	5	34	14	PDE
3	-109.498	40.407	21	1985	10	7	20	33	40.1	SRA
3.1	-71.471	42.528	13	1985	10	15	20	0	38.4	SRA
3.9	-73.829	40.983	6	1985	10	19	10	7	40.3	USHIS
3.3	-73.45	45.29	5	1985	11	1	23	33	39	PDE
3.3	-92.188	35.223	4	1985	11	8	19	56	48.5	SRA
3.2	-113.36	49.18	18	1985	12	4	8	38	14	DNAG
3.8	-89.99	35.88	5	1985	12	5	22	59	41.2	SRA
3	-104.665	35.437	0	1985	12	15	7	14	52.6	SNMX
3.1	-66.6	47	5	1985	12	21	6	3	11	DNAG
3.3	-83.72	35.701	13	1985	12	22	0	56	5	SRA

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.5	-88.965	38.552	5	1985	12	29	8	56	56.3	SRA
3.2	-84.762	35.609	22	1986	1	7	1	26	43.3	SRA
3.3	-77.32	45.8	18	1986	1	10	9	59	48	PDE
4	-70.18	47.7	18	1986	1	11	13	30	28	PDE
3.3	-100.693	32.066	5	1986	1	30	22	26	37	SRA
4.9	-81.162	41.65	2	1986	1	31	16	46	42.3	USHIS
3.5	-82.907	34.793	5	1986	2	13	11	35	45.3	SRA
3.1	-102.514	35.308	5	1986	3	3	11	45	17.4	SRA
3.5	-66.6	47	5	1986	3	6	8	34	51	PDE
3	-85.51	35.187	27	1986	4	19	7	40	53	SRA
4.4	-87.347	33.335	1	1986	5	7	2	27	0.4	SRA
3.3	-66.14	46.54	18	1986	5	9	9	4	33	PDE
3.2	-110.319	37.294	8	1986	5	14	15	2	55.7	SRA
3	-92.217	35.178	5	1986	5	24	8	16	1.5	SRA
3.4	-89.88	36.58	10	1986	5	24	12	48	13.5	SRA
3.4	-98.289	43.937	5	1986	5	25	7	13	22.1	SRA
3.4	-66.6	47	5	1986	6	1	14	53	14	PDE
3	-99.781	39.344	5	1986	6	2	4	4	5.2	SRA
3.3	-75.09	46.34	8	1986	6	5	12	13	22	PDE
3	-105.694	42.397	20	1986	6	12	15	14	34	SRA
3.8	-84.987	34.937	13	1986	7	11	14	26	14.8	USHIS
4.5	-84.371	40.537	10	1986	7	12	8	19	37.9	USHIS
3.4	-68.198	46.17	9	1986	7	12	20	32	48.4	SRA
3.5	-75.22	46.37	18	1986	8	6	11	19	36	PDE
3.3	-74.246	45.131	24	1986	8	13	4	55	18.4	PDE
4	-110.574	37.42	5	1986	8	22	13	26	33.3	SRA
3.7	-89.79	38.32	5	1986	8	26	16	41	24.8	SRA
3.1	-105.17	35.12	0	1986	8	27	18	6	58	SNMX
3.5	-107.09	38.912	5	1986	9	3	6	20	50.9	SRA
4.2	-70.32	47.3	22	1986	9	19	15	53	1	PDE
4.1	-66.6	47	5	1986	10	17	14	47	59	PDE
3	-101.372	37.918	5	1986	10	20	4	32	49	SRA
3.9	-71.59	43.399	5	1986	10	25	17	16	38.4	SRA
3.3	-108.896	41.922	5	1986	11	3	0	23	45	SRA
3	-110.297	37.43	1	1986	11	7	1	31	53.7	SRA
3	-82.88	34.898	9	1986	12	11	14	7	11.5	SRA
3.5	-89.58	36.42	14	1986	12	30	7	15	19.1	SRA
3.5	-103.482	42.788	5	1987	1	1	8	2	24	PDE
3	-89.978	35.893	5	1987	1	16	3	25	35.7	PDE
3.1	-98.097	35.828	5	1987	1	24	16	8	17	PDE
3.8	-110.616	40.442	1	1987	3	5	3	2	50.4	PDE
4.2	-84.229	35.567	19	1987	3	27	7	29	30.4	USHIS
5.2	-87.954	38.713	10	1987	6	10	23	48	54.8	USHIS
4.1	-89.686	36.605	5	1987	6	13	21	17	12.8	PDE
3.6	-89.173	36.839	5	1987	7	7	19	19	5.7	PDE

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3	-98.292	44.332	10	1987	7	9	22	6	45.4	PDE
3.4	-83.817	36.103	25	1987	7	11	0	4	29.4	PDE
3.8	-80.767	41.896	5	1987	7	13	5	49	17.4	PDE
3.4	-79.472	43.491	6	1987	7	23	9	32	28.5	PDE
3.1	-89.688	38.308	5	1987	8	31	17	12	35.5	PDE
3.3	-84.311	35.623	19	1987	9	22	17	23	50.1	PDE
3.7	-74.517	44.375	10	1987	9	26	17	44	6.9	PDE
4.5	-89.21	36.84	5	1987	9	29	0	4	57.2	USHIS
3.6	-107.381	45.771	5	1987	10	5	18	54	49.3	PDE
3.8	-88.793	37.049	5	1987	10	14	15	49	39.5	PDE
3	-98.599	44.472	5	1987	10	15	10	54	33.8	PDE
3.5	-83.099	36.848	14	1987	11	27	18	58	29.5	PDE
3.7	-98.024	36.055	5	1987	12	8	1	42	40.3	PDE
3	-82.628	34.244	5	1987	12	12	3	53	28.7	PDE
3.3	-84.201	35.275	12	1988	1	9	1	7	40.7	PDE
3.6	-89.621	46.559	5	1988	1	14	17	23	36.5	PDE
3.3	-80.157	32.935	7	1988	1	23	1	57	16.3	PDE
4.9	-65.58	48	18	1988	1	28	8	38	28	PDE
3.5	-90.465	35.681	10	1988	1	31	0	12	43.4	PDE
3.3	-108.532	40.626	5	1988	2	14	18	32	40.5	PDE
3.2	-82.304	36.561	5	1988	2	16	15	26	54.5	PDE
3.5	-83.853	35.366	5	1988	2	18	0	37	45.9	PDE
3.3	-66.6	47	5	1988	3	6	18	13	18.1	PDE
3.9	-75.716	46.341	18	1988	3	10	14	42	55.2	PDE
3.4	-99.155	39.093	5	1988	4	14	9	39	31.4	PDE
4.1	-81.987	37.238	0	1988	4	14	23	37	31.1	PDE
3.4	-66.6	47	5	1988	5	9	1	23	3.6	PDE
3.5	-75.58	45.17	7	1988	5	15	6	10	5.6	PDE
3.3	-92.77	37.288	5	1988	5	20	23	6	22.6	PDE
3.3	-110.448	36.374	5	1988	7	15	0	38	9.5	PDE
3.5	-74.955	44.995	10	1988	8	9	13	57	26.9	PDE
5.4	-110.869	39.128	10	1988	8	14	20	3	3.9	USHIS
3.8	-66.59	46.99	5	1988	8	26	5	59	10.2	PDE
4.6	-83.878	38.143	10	1988	9	7	2	28	9.5	USHIS
3.5	-87.931	38.69	5	1988	10	5	0	38	52.2	PDE
3.9	-71.158	44.539	5	1988	10	20	13	9	50.1	USHIS
3.8	-70.386	44.424	5	1988	11	14	6	15	43.1	PDE
5.8	-71.183	48.117	28	1988	11	25	23	46	4.5	PDE
3.3	-92.702	34.189	13	1988	12	25	15	57	57.7	PDE
3.5	-69.342	44.514	10	1988	12	28	6	28	44.4	PDE
3.5	-89.428	36.185	5	1988	12	31	14	24	20.5	PDE
3.8	-67.357	49.264	18	1989	1	1	17	55	53.6	PDE
3.1	-112.862	49.056	5	1989	1	4	18	50	9.7	PDE
3.4	-104.103	35.183	0	1989	1	29	5	7	15.6	SNMX
3.8	-101.898	42.685	5	1989	2	9	5	15	45.8	PDE

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.5	-87.092	33.643	0	1989	2	28	17	31	50.8	PDE
4.4	-69.9	47.7	18	1989	3	11	8	31	52.1	PDE
3.5	-71.144	44.511	5	1989	4	6	2	35	51.3	PDE
3.7	-105.602	47.716	5	1989	4	7	8	26	48.9	PDE
3	-89.711	36.557	5	1989	4	15	16	39	51.1	PDE
4.4	-89.768	36.006	10	1989	4	27	16	47	49.8	USHIS
3.7	-89.71	36.74	2	1989	5	14	0	16	9.5	PDE
3.9	-99.477	39.165	5	1989	6	8	18	18	43.3	PDE
3.1	-83.569	38.607	10	1989	7	15	0	8	2.6	PDE
3.1	-98.876	36.434	5	1989	7	20	6	7	50.4	PDE
3.3	-79.53	43.21	18	1989	8	5	21	7	59.1	PDE
3.5	-65.82	46.65	18	1989	8	10	21	17	43.5	PDE
3.4	-87.086	33.632	0	1989	8	13	20	16	2.9	PDE
3.9	-87.645	34.736	10	1989	8	20	0	3	17.8	USHIS
3	-70.899	41.614	5	1989	8	24	15	56	59.3	PDE
3	-108.948	47.547	5	1989	8	31	4	2	38.3	PDE
3.4	-89.62	36.545	11	1989	9	14	17	31	28	PDE
3	-107.027	41.207	5	1989	11	2	6	23	56.2	PDE
4	-76.59	46.57	18	1989	11	16	9	24	52	PDE
3	-107.767	38.055	5	1989	11	19	3	21	13.6	PDE
3.3	-99.908	45.317	5	1989	11	26	1	6	14.6	PDE
3.2	-90.744	35.245	5	1989	12	25	8	29	26.9	PDE
3.9	-86.434	38.133	5	1990	1	24	18	20	24.4	PDE
4	-102.504	43.313	5	1990	1	28	4	59	59.1	PDE
3.6	-89.219	38.868	10	1990	3	2	7	1	47.7	PDE
3	-91.49	36.72	5	1990	3	18	16	22	33	PDE
3.5	-68.23	47.28	18	1990	3	30	1	54	9	PDE
3.1	-112.37	48.717	6	1990	4	4	21	42	33.6	PDE
3.5	-109.519	40.082	3	1990	4	7	15	37	54.8	PDE
3	-84.852	40.46	5	1990	4	17	10	27	34.7	PDE
3	-88.23	39.556	10	1990	4	24	9	41	24.3	PDE
3	-110.828	38.952	11	1990	6	25	17	15	33.5	PDE
3	-98.954	41.507	5	1990	7	18	2	47	3.9	PDE
3	-89.24	36.85	6	1990	8	7	5	5	56.4	PDE
3.8	-83.34	36.794	10	1990	8	17	21	1	17.9	PDE
3.5	-89.66	35.83	13	1990	8	29	19	34	59.9	PDE
3.3	-83.731	38.061	5	1990	9	8	0	3	57.4	PDE
3	-106.206	39.701	5	1990	9	12	21	38	57.6	PDE
4.9	-89.577	37.165	12	1990	9	26	13	18	51.3	PDE
3	-101.505	41.815	5	1990	9	30	0	6	24	PDE
3.9	-75.19	46.32	17	1990	10	7	8	47	30.5	PDE
4.9	-75.59	46.47	13	1990	10	19	7	1	57.4	PDE
3	-75.506	39.512	10	1990	10	23	1	34	48.2	PDE
3.3	-88.99	38.31	5	1990	10	24	8	20	4.3	PDE
3.9	-98.472	43.794	5	1990	10	25	6	25	25.5	PDE

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.5	-89.62	36.54	8	1990	11	9	3	39	15.9	PDE
3.2	-80.136	32.947	3	1990	11	13	15	22	13	PDE
3.8	-97.59	34.76	5	1990	11	15	11	44	41.4	PDE
4.1	-76.219	47.128	18	1990	11	15	13	47	15.7	PDE
3.5	-66.6	47	5	1990	12	12	5	15	7.1	PDE
3.2	-87.044	40.068	10	1990	12	17	5	24	59.1	PDE
3.6	-86.671	39.57	10	1990	12	20	14	4	17.1	PDE
4.3	-72.556	47.579	18	1990	12	31	3	53	58.3	PDE
3	-88.86	37.946	5	1991	1	23	9	25	23.5	PDE
3	-97.3	36.378	5	1991	1	24	5	0	26.9	PDE
3.3	-81.453	41.536	5	1991	1	26	3	21	22.6	PDE
3.4	-111.429	37.681	9	1991	1	26	21	49	38	PDE
3	-89.95	35.98	14	1991	2	11	0	0	6.1	PDE
3.3	-109.483	40.091	1	1991	3	2	8	41	37.4	PDE
3.9	-76.874	46.282	18	1991	3	6	5	26	53.6	PDE
3.8	-77.916	37.746	17	1991	3	15	6	54	8.2	PDE
3.9	-66.594	49.698	18	1991	3	21	4	10	59.3	PDE
3	-106.857	42.031	5	1991	4	13	19	8	5.1	PDE
3.5	-80.207	37.941	14	1991	4	22	1	1	20.2	PDE
3.6	-66.6	47	5	1991	4	23	3	19	19	PDE
4.7	-89.823	36.564	5	1991	5	4	1	18	54.9	PDE
3.6	-74.4	45.5	18	1991	5	17	18	8	47	PDE
3.5	-99.4	39.2	5	1991	5	30	22	7	44	PDE
3.6	-112.007	48.374	5	1991	6	5	9	24	7.6	PDE
4.3	-76.7	47	18	1991	6	16	16	46	53	PDE
4	-74.678	42.63	5	1991	6	17	8	53	16.7	PDE
3	-110.358	37.209	1	1991	6	25	21	2	13.6	PDE
3.2	-81.668	38.276	5	1991	6	28	18	34	51.9	PDE
3.3	-91.71	37.49	5	1991	7	2	3	49	1.7	PDE
3.8	-73.896	45.232	18	1991	7	5	1	47	36.7	PDE
3.9	-91.643	36.658	5	1991	7	7	21	24	2.6	PDE
3	-89.44	36.14	12	1991	7	8	23	49	7.4	PDE
3.5	-98.042	28.908	10	1991	7	20	23	38	19.2	PDE
3.5	-108.861	43.502	5	1991	8	7	12	49	16.6	PDE
3	-77.657	40.786	1	1991	8	15	7	16	7.1	PDE
3.4	-100.533	42.162	5	1991	8	26	11	49	15.4	PDE
3.1	-84.095	35.711	5	1991	9	24	7	21	6.4	PDE
3.1	-89.432	36.841	5	1991	10	3	11	46	4.8	PDE
3	-73.578	41.07	10	1991	10	28	20	58	26.1	PDE
3.8	-87.894	38.713	10	1991	11	11	9	20	47.4	PDE
3	-90.27	35.72	9	1991	11	13	9	43	15.9	PDE
3	-108.895	47.952	10	1991	12	5	10	10	0.7	PDE
4	-69.8	47.7	18	1991	12	8	3	0	30	PDE
3.1	-106.917	41.936	5	1991	12	18	21	36	47.9	PDE
3.4	-106.715	45.82	5	1991	12	23	20	32	27.2	PDE

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
5	-103.187	32.302	0	1992	1	2	11	45	35.3	SNMX
3.2	-82.465	33.946	5	1992	1	3	4	21	22.2	PDE
3	-74.341	40.363	7	1992	1	9	8	50	45.2	PDE
3.5	-81.245	41.911	5	1992	3	15	6	13	55.2	PDE
3.3	-89.479	35.828	12	1992	4	3	3	6	3.9	PDE
3.2	-104.773	37.335	5	1992	4	15	22	46	5	PDE
3.1	-90.41	36.92	5	1992	4	30	0	1	30.9	PDE
3.2	-70.407	47.446	2	1992	5	1	0	37	51.4	PDE
3.1	-104.778	37.378	5	1992	5	2	10	19	29.8	PDE
3.7	-74.964	46.444	18	1992	5	19	5	59	41	PDE
3.3	-99.549	38.76	5	1992	7	15	2	56	40.7	PDE
4.2	-80.116	33.05	10	1992	8	21	16	31	55.1	PDE
3	-102.708	32.173	5	1992	8	26	3	24	52.6	PDE
3.3	-89.68	37.63	5	1992	8	26	5	41	38.4	PDE
3.6	-107.041	43.825	5	1992	8	31	1	40	14.2	PDE
3.4	-71.578	43.324	5	1992	10	6	15	38	4	PDE
4	-108.242	42.819	5	1992	10	10	15	40	56.2	PDE
3	-104.389	42.74	5	1992	11	2	6	54	10.3	PDE
3.3	-112.611	49.001	5	1992	11	17	3	37	22.9	PDE
4.2	-74.862	45.764	18	1992	11	17	3	58	0.9	PDE
3.5	-97.581	34.744	5	1992	12	17	7	18	4.2	PDE
3.2	-89.63	37.5	5	1992	12	27	10	12	58.9	PDE
3	-82.09	35.877	3	1993	1	1	5	8	5.3	PDE
3.3	-112.19	48.897	5	1993	1	1	15	57	41.9	PDE
3.5	-90.03	35.83	21	1993	1	8	13	1	18.8	PDE
3.1	-98.275	36.595	5	1993	1	14	17	6	10.4	PDE
3.1	-84.974	35.075	1	1993	1	15	2	2	51.8	PDE
3	-89.617	36.222	13	1993	1	21	19	46	19.3	PDE
3.4	-112.403	49.212	5	1993	1	22	6	2	32.7	PDE
3.2	-89.04	39.038	5	1993	1	29	13	56	23.2	PDE
3.5	-89.73	36.66	7	1993	2	6	2	9	45.5	PDE
3.5	-101.461	42.83	5	1993	2	20	13	8	10.1	PDE
3.7	-106.062	44.932	5	1993	2	25	3	44	15.5	PDE
3.1	-89.49	36.67	8	1993	3	2	0	29	11.8	PDE
3.2	-106.617	43.399	5	1993	3	10	3	54	31.1	PDE
3.2	-90.55	35.67	10	1993	3	16	7	38	10.2	PDE
3	-104.438	35.15	0	1993	3	24	2	32	5.9	SNMX
3.3	-89.42	36.79	5	1993	3	31	20	23	21.2	PDE
4.2	-98.124	28.811	5	1993	4	9	12	29	19.1	PDE
3.6	-89.44	36.19	7	1993	4	28	22	40	1.9	PDE
3.5	-75.5	46.3	18	1993	5	6	1	23	25.9	PDE
3.8	-107.575	42.304	5	1993	6	1	21	33	22.9	PDE
4.1	-96.293	45.674	10	1993	6	5	1	24	53	PDE
3	-105.373	42.985	5	1993	6	30	6	50	57.8	PDE
3.1	-106.715	39.227	5	1993	7	8	4	3	52.2	PDE

Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.7	-88.341	31.747	5	1993	7	16	10	54	32.8	PDE
3.7	-105.703	42.478	5	1993	7	23	6	30	23.8	PDE
3.9	-74.12	45.26	8	1993	7	30	22	30	54	PDE
3	-89.88	36	11	1993	8	5	7	21	37.4	PDE
3.2	-81.595	33.633	5	1993	8	8	9	24	31.1	PDE
3	-109.921	43.576	5	1993	8	23	5	29	47.6	PDE
3	-106.837	42.033	5	1993	8	23	13	12	13.8	PDE
3.3	-90.36	38.09	16	1993	8	27	0	8	34	PDE
3.5	-75.05	46.457	15	1993	8	30	5	15	28.5	PDE
3.7	-74.605	46.065	18	1993	9	23	6	45	28.4	PDE
3	-103.56	35.568	0	1993	9	29	2	1	28.5	SNMX
3.7	-105.868	42.421	5	1993	10	10	4	17	46.7	PDE
3.5	-81.012	41.698	5	1993	10	16	6	30	5.3	PDE
3.5	-107.384	43.884	5	1993	11	16	7	26	4	PDE
4.2	-73.495	45.182	17	1993	11	16	9	31	44.2	PDE
3.3	-103.157	35.808	0	1993	11	30	3	7	36.3	SNMX
3.5	-70.06	47.53	8	1993	12	1	12	47	15	PDE
3.5	-105.499	42.333	5	1993	12	13	14	51	3	PDE
4.3	-75.606	46.506	18	1993	12	25	16	44	22.3	PDE
4.7	-110.132	43.483	8	1993	12	28	21	2	28.7	PDE
3.8	-70.367	47.453	7	1993	12	30	23	1	47.5	PDE
4.6	-76.037	40.33	5	1994	1	16	1	49	16.2	PDE
3.3	-100.141	42.627	5	1994	1	25	2	44	39.8	PDE
4.2	-89.18	37.37	16	1994	2	5	14	55	37.7	PDE
3.1	-95	45	5	1994	2	9	8	45	35.5	PDE
3.2	-82	36.8	5	1994	2	12	2	40	24.5	PDE
3.5	-77.876	42.782	1	1994	3	12	10	43	15.7	PDE
3.6	-65.74	48.99	18	1994	3	28	16	28	23	PDE
3.2	-85.493	34.961	5	1994	4	5	22	21	59	PDE
3.1	-89.27	38.123	10	1994	4	6	17	38	55.8	PDE
3	-87.174	34.198	5	1994	5	4	9	12	2.7	PDE
3.2	-92.671	33.013	5	1994	6	10	23	34	2.9	PDE
4	-66.6	47	5	1994	7	14	12	41	52	PDE
3.7	-76.751	35.067	5	1994	8	6	19	54	9.9	PDE
4.1	-111.333	48.489	5	1994	8	16	11	3	41.7	PDE
3.5	-91.058	36.136	5	1994	8	20	10	45	44.6	PDE
3.5	-84.604	42.798	5	1994	9	2	21	23	6.5	PDE
3	-69.232	43.861	5	1994	9	5	14	13	52.2	PDE
4.5	-107.976	38.151	10	1994	9	13	6	1	23	PDE
3.6	-68.223	45.306	5	1994	9	16	4	22	42.5	PDE
4.2	-69.96	47.77	17	1994	9	25	0	53	28	PDE
3.6	-88.935	36.929	5	1994	9	26	14	23	22	PDE
3.6	-72.277	42.347	10	1994	10	2	11	27	22.5	PDE
3.4	-108.269	40.04	5	1994	11	3	11	40	10.1	PDE
4	-104.811	39.29	10	1994	12	25	19	6	7.5	PDE

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
4.1	-97.596	34.774	5	1995	1	18	15	51	39.4	PDE
5.2	-109.64	41.529	1	1995	2	3	15	26	10.6	PDE
3.1	-94.952	40.505	5	1995	2	11	5	54	10.1	PDE
3.5	-75.04	45.9	18	1995	2	15	15	53	57	PDE
3.6	-83.47	39.12	10	1995	2	19	12	57	6	PDE
3	-74.426	44.233	4	1995	3	2	5	33	51.4	PDE
3.3	-112.35	48.65	10	1995	3	5	12	17	11.5	PDE
3.3	-84.922	35.425	17	1995	3	18	22	6	21	PDE
3.3	-104.212	35	5	1995	3	19	18	36	43.9	PDE
4.1	-108.925	40.179	5	1995	3	20	12	46	16.3	PDE
3.9	-80.068	32.947	10	1995	4	17	13	45	57.8	PDE
3.3	-67.73	49.15	15	1995	4	20	4	37	5	PDE
3.9	-66.6	47	5	1995	5	6	7	51	35	PDE
3.5	-89.43	36.17	6	1995	5	27	19	51	10.4	PDE
3.4	-87.827	33.191	1	1995	5	28	15	28	36.9	PDE
3	-96.732	34.287	5	1995	6	1	4	49	29.3	PDE
3.6	-76.29	47.02	18	1995	6	3	22	44	32	PDE
3.8	-71.915	44.286	5	1995	6	16	12	13	11.4	PDE
3.5	-81.452	36.747	5	1995	6	26	0	36	17	PDE
3.8	-104.814	36.246	5	1995	7	4	3	59	4.5	PDE
3.7	-84.212	35.366	10	1995	7	5	14	16	44.4	PDE
3	-81.873	36.515	11	1995	7	7	21	1	2.8	PDE
3.3	-87.665	33.478	1	1995	7	15	1	3	28.3	PDE
3	-89.632	36.528	5	1995	7	20	2	10	34.4	PDE
3	-74.953	46.168	20	1995	7	28	5	47	37.1	PDE
3.1	-89.409	36.102	5	1995	8	17	23	18	50.8	PDE
3	-73.28	45.41	18	1995	8	20	16	15	26	PDE
3.7	-74.43	45.61	18	1995	9	12	3	59	5	PDE
3.9	-98.69	36.87	5	1995	9	15	0	31	33.2	PDE
3.1	-74.21	45.08	18	1995	9	21	23	3	27	PDE
3.3	-78.77	46.42	18	1995	10	10	7	19	20	PDE
3.7	-96.864	45.788	5	1995	10	20	15	57	18.7	PDE
3.6	-104.917	38.732	5	1995	12	23	6	51	48.8	PDE
4.2	-110.878	39.12	0	1996	1	6	12	55	58.6	PDE
3.6	-97.542	42.513	5	1996	2	6	15	10	28.2	PDE
3.7	-103.729	43.981	5	1996	2	6	16	8	36.7	PDE
4.3	-74.43	45.99	18	1996	3	14	10	42	26	PDE
3.1	-71.242	41.69	11	1996	3	22	20	22	12.5	PDE
3.3	-102.601	35.61	5	1996	3	25	6	43	46.8	PDE
3.3	-88.671	32.131	5	1996	3	25	14	15	50.5	PDE
3.7	-104.102	43.069	5	1996	4	9	2	48	8.1	PDE
3.3	-91.162	34.969	5	1996	4	11	21	54	57.6	PDE
4	-81.95	37.187	1	1996	6	29	19	30	42.6	PDE
3.8	-104.247	37.398	5	1996	8	1	5	44	22.7	PDE
3.4	-90.874	33.577	10	1996	8	11	18	17	49.8	PDE



**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

(Page 62 of 65)

$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.5	-112.405	49.076	5	1996	8	15	20	7	29.9	PDE
3.6	-82.92	49.21	18	1996	8	16	4	56	46	PDE
3.7	-71.352	44.184	10	1996	8	21	7	54	14	PDE
4.2	-106.056	43.09	5	1996	10	19	13	27	57.9	PDE
3.7	-109.27	42.549	5	1996	10	21	13	51	39.5	PDE
3.2	-104.232	37.349	5	1996	11	1	3	9	28.3	PDE
3	-100.504	35.04	5	1996	11	23	10	54	18.5	PDE
4.2	-89.927	35.919	20	1996	11	29	5	41	33.6	PDE
3.6	-89.37	36.29	5	1996	11	29	10	47	9	PDE
3.4	-107.693	42.369	5	1996	12	11	3	55	44	PDE
3.1	-87.4	39.5	5	1996	12	16	1	58	31.3	PDE
3	-100.89	34.947	5	1997	2	12	23	53	10.7	PDE
3.2	-100.569	34.973	5	1997	2	15	9	8	55.4	PDE
3.4	-93.435	34.209	5	1997	3	16	19	7	27.9	PDE
3.8	-98.054	27.717	5	1997	3	24	22	31	34.5	PDE
3.5	-72.33	45.98	5	1997	4	3	4	44	12	PDE
3.1	-108.732	42.683	5	1997	4	25	10	39	6.7	PDE
3.7	-112.65	49.13	5	1997	5	1	21	38	36	PDE
3.1	-87.4	31	5	1997	5	4	3	39	12.9	PDE
4.1	-74.421	45.978	10	1997	5	24	18	52	6.3	PDE
3.4	-95.966	33.182	5	1997	5	31	3	26	41.3	PDE
3.5	-84.808	35.056	10	1997	7	19	17	6	34.3	PDE
3.8	-83.509	36.436	5	1997	7	30	12	29	23.3	PDE
3.2	-75.37	43.624	5	1997	7	31	7	15	29.7	PDE
3.4	-97.185	41.795	5	1997	8	9	17	46	3.9	PDE
3.7	-70.29	47.53	18	1997	8	20	9	12	4	PDE
4.3	-96.435	34.66	5	1997	9	6	23	38	0.9	PDE
3.8	-90.457	35.619	5	1997	9	17	18	16	31.6	PDE
3.1	-90.924	37.179	5	1997	9	20	5	55	50.4	PDE
3.2	-89.817	36.545	5	1997	9	24	4	20	24.8	PDE
3	-89.484	36.201	5	1997	9	27	12	14	9.3	PDE
3	-74.968	44.36	4	1997	10	13	23	6	40.2	PDE
4.9	-87.339	31.118	10	1997	10	24	8	35	17.8	PDE
4.7	-69.91	47.67	11	1997	10	28	11	44	18	PDE
4.9	-71.41	46.8	22	1997	11	6	2	34	33	PDE
3	-76.252	40.146	5	1997	11	14	3	44	11	PDE
4	-87.306	33.466	1	1997	12	12	8	42	20.2	PDE
3.4	-103.408	37.828	5	1998	1	2	15	47	16.4	PDE
3	-89.712	36.123	10	1998	2	12	9	37	49.5	PDE
3	-107.513	44.333	5	1998	2	13	2	28	4	PDE
3.6	-76.36	46.07	18	1998	2	26	14	20	31	PDE
3.9	-81.07	46.49	1	1998	3	9	5	5	58	PDE
3.1	-111.35	38.25	3	1998	3	29	12	12	42	PDE
3.2	-89.02	36.94	13	1998	4	8	18	16	49	PDE
3.8	-80.466	34.61	5	1998	4	13	9	56	11.3	PDE

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.9	-74.99	45.57	18	1998	4	18	16	22	52	PDE
3.2	-102.383	35.453	5	1998	4	27	15	22	46.2	PDE
4.2	-98.416	34.782	5	1998	4	28	14	13	1.6	PDE
3.7	-81.174	46.457	1	1998	5	25	15	47	2	PDE
3.2	-80.821	35.479	5	1998	6	5	2	31	1.9	PDE
3.4	-73.72	44.75	4	1998	6	9	8	53	51	PDE
3.6	-84.405	35.926	10	1998	6	17	8	0	23.4	PDE
3.4	-103.003	42.622	5	1998	6	18	16	26	38.3	PDE
3.4	-87.954	32.501	5	1998	6	24	15	20	1.3	PDE
3.2	-97.589	34.719	5	1998	7	7	18	44	44.4	PDE
3.1	-101.111	43.554	5	1998	7	12	16	28	49.6	PDE
3.1	-89.52	36.69	13	1998	7	15	4	24	51	PDE
4	-66.61	47.02	5	1998	7	15	7	8	4	PDE
3.6	-104.706	48.37	5	1998	7	29	3	31	58.9	PDE
4.1	-74.73	46.16	10	1998	7	30	8	57	22	PDE
3.6	-107.19	41.953	10	1998	8	6	18	22	7.1	PDE
4.9	-80.388	41.495	5	1998	9	25	19	52	52	PDE
3.4	-111.091	36.033	5	1998	10	18	7	13	10.6	PDE
3.5	-78.367	37.381	13	1998	10	21	5	56	47.2	PDE
4.1	-66.88	49.34	18	1998	10	22	9	43	35	PDE
3.5	-97.6	36.8	5	1998	10	30	17	41	22.2	PDE
3.3	-104.032	48.548	5	1998	11	11	11	59	37.6	PDE
3.3	-77.93	43.83	18	1998	12	25	13	30	26	PDE
3	-99.378	38.674	5	1999	1	7	5	16	26.9	PDE
3	-70.98	42.84	2	1999	1	10	10	52	16.1	PDE
3	-83.691	36.854	5	1999	1	17	18	38	4.7	PDE
4.3	-87.255	33.405	1	1999	1	18	7	0	53.4	PDE
3.4	-80.939	49.267	18	1999	2	1	22	22	5.6	PDE
3.8	-69.52	44.48	3	1999	2	26	3	38	43	PDE
4	-104.63	32.591	1	1999	3	14	22	43	17.9	PDE
4.8	-66.32	49.61	18	1999	3	16	12	50	48	PDE
4.3	-107.741	41.451	10	1999	4	6	0	41	9.5	PDE
3	-94.7	39.1	5	1999	5	13	14	18	22.7	PDE
3.9	-104.664	32.575	10	1999	5	30	19	4	25.6	PDE
3.1	-108.459	42.559	5	1999	7	21	2	36	6.9	PDE
3.1	-89.503	36.264	9	1999	8	23	12	12	41.1	PDE
3.5	-89.433	41.721	5	1999	9	2	16	17	29.7	PDE
3.8	-91.02	36.49	19	1999	10	21	8	18	0	PDE
3	-99.659	36.846	26	1999	10	25	23	19	58.3	PDE
3.9	-74.32	45.85	18	1999	10	31	20	14	10	PDE
3.5	-105.467	45.512	10	1999	11	3	13	28	52	PDE
3.1	-107.477	43.479	5	1999	11	9	8	17	41.1	PDE
3.6	-78.997	43.71	12	1999	11	26	22	33	1.4	PDE
3.8	-87.253	33.416	1	1999	11	28	11	0	9.3	PDE
3	-69.37	44.94	5	1999	12	25	0	21	41	PDE

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
5	-78.93	46.888	18	2000	1	1	11	22	57	PDE
3.5	-70.17	44.31	9	2000	1	3	21	5	50	PDE
3	-110.34	42.1	12	2000	1	8	22	43	37	PDE
3.5	-70.44	44.57	16	2000	1	17	8	16	20	PDE
3.5	-83.214	32.993	5	2000	1	18	22	19	31.9	PDE
3	-71.18	43	1	2000	1	27	14	49	40	PDE
4.4	-109.679	41.464	1	2000	1	30	14	46	51.3	PDE
3	-106.732	42.24	5	2000	2	1	22	15	45.3	PDE
3	-106.666	40.601	5	2000	2	7	17	24	54.3	PDE
3.3	-105.813	42.409	5	2000	4	13	18	17	31.7	PDE
3.6	-86.75	39.76	5	2000	4	14	3	54	20	PDE
3.9	-74.257	43.949	5	2000	4	20	8	46	55.4	PDE
3.1	-79.099	43.806	18	2000	5	24	10	22	46.2	PDE
4	-107.57	42.196	5	2000	5	26	21	58	46.6	PDE
3	-87.82	33.809	5	2000	5	28	11	32	7	PDE
3.7	-69.81	47.67	10	2000	6	15	9	25	54	PDE
3.3	-72.82	42.1	9	2000	6	16	4	2	53	PDE
3	-109.31	40.69	1	2000	6	20	17	55	46	PDE
3.8	-92.75	35.8	0	2000	6	27	1	28	45	PDE
3	-88.87	37.13	4	2000	6	27	6	2	57	PDE
4.2	-71.1	47.52	10	2000	7	12	15	1	49	PDE
4.1	-74.97	46.19	18	2000	8	6	8	52	24	PDE
3.9	-101.814	35.39	5	2000	8	17	1	8	5.4	PDE
3.2	-108.26	42.554	5	2000	8	19	2	55	43.7	PDE
3.9	-91.106	36.492	8	2000	8	22	20	12	14	PDE
3.2	-69.382	44.355	5	2000	9	7	10	7	40.7	PDE
3.8	-74.02	45.13	18	2000	10	6	13	59	4	PDE
3	-107.693	43.437	5	2000	11	8	2	16	49.9	PDE
3.7	-109.23	40.28	5	2000	11	11	21	17	53	PDE
3.9	-87.66	37.973	5	2000	12	7	14	8	49.4	PDE
3.9	-101.8	35.4	5	2000	12	16	22	8	54	PDE
4.3	-80.802	41.942	5	2001	1	26	3	3	20	PDE
3.2	-77.394	42.345	0	2001	2	3	20	15	15	PDE
3	-92.66	33.19	5	2001	3	3	10	46	13	PDE
3.2	-84.81	35.51	6	2001	3	7	17	12	25	PDE
3.9	-76.28	47.05	18	2001	3	19	10	40	17	PDE
3.2	-85.439	34.857	3	2001	3	21	23	35	35	PDE
3.1	-93.327	37.933	5	2001	3	30	17	13	55.6	PDE
3	-83.34	36.53	0	2001	4	13	16	36	20.7	PDE
4.5	-92.194	35.205	10	2001	5	4	6	42	12.6	PDE
3.3	-103.141	32.334	5	2001	6	2	1	55	53.7	PDE
3.2	-80.767	41.905	5	2001	6	3	22	36	46.4	PDE
3	-89.396	36.279	14	2001	7	7	20	45	42.7	PDE
3.1	-105.129	39.022	5	2001	7	22	19	22	45.5	PDE
3	-97	37.7	5	2001	7	24	14	2	35	PDE

**Table 2.5-1 {USGS Earthquake Catalog for the CEUS with  $m_b \geq 3.0$ }**

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$m_b$	Longitude (degree)	Latitude (degree)	Depth (km)	Year	Month	Day	Hour	Minute	Second	Catalog Reference
3.2	-83.575	35.932	5	2001	7	26	5	26	44.7	PDE
3.1	-93.213	34.292	5	2001	8	4	1	13	25.3	PDE
4	-107.378	39.66	5	2001	8	9	22	38	54.5	PDE
4.5	-104.618	37.143	5	2001	9	5	10	52	7.8	PDE-W
4.3	-110.051	43.459	5	2001	9	27	22	5	21.7	PDE-W
3.2	-68.67	45.2	9	2001	10	25	0	24	29.8	PDE-W
3.4	-107.384	38.851	1	2001	11	5	8	34	23	PDE-W
3.3	-100.208	39.996	5	2001	11	13	1	56	13.1	PDE-W
3.1	-102.631	31.786	5	2001	11	22	0	7	8	PDE-W
3.1	-107.374	38.813	1	2001	12	4	18	20	9.1	PDE-W
3.9	-86.245	34.735	5	2001	12	8	1	8	21.5	PDE-W
3.3	-104.797	36.859	5	2001	12	15	7	58	31.3	PDE-W
3.8	-76.49	46.87	18	2001	12	24	16	58	21	PDE-W
3.8	-75.11	40.61	3	2003	8	26	18	24	18	PDE
3.1	-74.12	42.6	15	2007	7	24	1	56	49	PDE

1. The New Madrid events occurred in a cluster of three events. The event shown in the catalog is determined by USGS as the New Madrid event for the 1811-1812 cluster set. Two other events are considered as the foreshock/aftershock events and are filtered out from the catalog by USGS. The event shown in the catalog is not considered in the general are source hazard integration since its magnitude is above the maximum magnitude limit considered. This New Madrid event is accounted for in the PSHA in the New Madrid Characteristic Cluster events. The following events are the New Madrid set considered for the characteristic earthquake analysis:

Longitude (degree)	Latitude (degree)	$m_b$	Year	Month	Day	Hour	minute	Event Group
-90	36	7.2	1811	12	16	8	15	1
-90	36	7.0	1811	12	16	14	15	
-89.6	36.3	7.1	1812	1	23	15	0	2
-89.6	36.5	7.4	1812	2	7	9	45	3

**Table 2.5-2 {Conversion Between Body-Wave ( $m_b$ ) and Moment (M) Magnitude}**

Conversion Model	$m_b$ to M	M to $m_b$	Weight
Atkinson and Boore 1987	$M = 2.715 - 0.277m_b + 0.127 m_b^2$	$m_b = -4.249 + 3.3935M^2 + 0.0168M^3$	0.333
EPRI 1993	$M = 0.3281 + 1.9437m_b - 0.43m_b^2 + 0.0419m_b^3$	$m_b = -10.23 + 6.105M - 0.7632M^2 + 0.03436M^3$	0.334
Johnston 1996	$M = 1.14 + 0.24m_b + 0.0933m_b^2$	$m_b = -0.919 + 1.7864M - 0.1351M^2 + 0.0052M^3$	0.333

Notes:

$m_b$  indicates body wave magnitude

M indicates moment magnitude

**Table 2.5-3 {Summary of Bechtel Group Seismic Sources}**

Code	Name	b value	b weight	M <sub>max</sub> (m <sub>b</sub> )	M <sub>max</sub> weight
3	Charlevoix	0.707	0.33	6.4	0.10
		0.498	0.34	6.7	0.40
		0.808	0.33	7.0	0.40
				6.6	0.10
11	Clarendon - Linden	0.497	0.33	5.4	0.10
		0.498	0.34	5.7	0.40
		0.808	0.33	6.0	0.40
				6.6	0.10
13	Mesozoic Basins	0.991	0.33	5.4	0.10
		0.985	0.34	5.7	0.40
		1.037	0.33	6.0	0.10
				6.6	0.10
24	Bristol Block	0.844	0.33	5.7	0.10
		0.836	0.34	6.0	0.40
		0.979	0.33	6.3	0.40
				6.6	0.10
25	NY-AL Lineament	1.057	0.33	5.4	0.10
		1.055	0.34	5.7	0.40
		1.043	0.33	6.0	0.40
				6.6	0.10
D	Niagara	0.952	0.33	5.4	0.10
		0.953	0.34	5.7	0.40
		0.998	0.33	6.0	0.40
				6.6	0.10
BZ5 Background	Southern Appalachian	0.912	0.33	5.7	0.10
		0.920	0.34	6.0	0.40
		0.937	0.33	6.3	0.40
				6.6	0.10
BZ6 Background	Southern Eastern Craton	1.074	0.33	5.4	0.10
		1.073	0.34	5.7	0.40
		1.084	0.33	6.0	0.40
				6.6	0.10
BZ7 Background	Northern Eastern Craton	1.066	0.33	6.0	0.10
		1.075	0.34	6.3	0.40
		1.087	0.33	6.6	0.50

**Table 2.5-4 {Summary of Dames & Moore Seismic Sources}**

Code	Name	b value	b weight	M <sub>max</sub> (m <sub>b</sub> )	M <sub>max</sub> weight
3	Adirondacks	1.043	0.75	6.3	0.80
		1.047	0.25	7.2	0.20
4	Paleozoic Fold Belts	1.042	0.75	6.0	0.80
		1.047	0.25	7.2	0.20
8	Eastern Marginal Basin	1.042	0.75	5.6	0.80
		1.052	0.25	7.2	0.20
9	Clarendon - Linden	1.035	0.37	6.5	0.75
		1.011	0.12		
		1.031	0.38	7.2	0.25
		1.004	0.13		
41	Southern Cratonic	1.035	0.75	6.1	0.80
		1.028	0.25	7.2	0.20
42	Newark-Gettysburg Basin	1.015	0.75	6.3	0.75
		0.947	0.25	7.2	0.25
53	Southern Appalachian Mobile Belt	1.043	0.75	5.6	0.80
		1.053	0.25	7.2	0.20
59	Charlevoix	0.70*	1.00	7.2	1.00
* Updated using USGS 2001 catalog and Dames and Moore Source geometry for the source zone.					

**Table 2.5-5 {Summary of Law Engineering Seismic Sources}**

Code	Name	b value	b weight	M <sub>max</sub> (m <sub>b</sub> )	M <sub>max</sub> weight
8-17	East Coast Mesozoic Basements	1.057	1.00	6.8	1.00
12	Charlevoix	0.756	1.00	6.4	0.20
				7.4	0.80
17	Eastern Basements	0.992	1.00	5.7	0.20
				6.8	0.80
22	Eastern Seaboard Normal Faults	1.054	1.00	6.8	1.00

**Table 2.5-6 {Summary of Rondout Associates Seismic Sources}**

Code	Name	b value	b weight	M <sub>max</sub> (m <sub>b</sub> )	M <sub>max</sub> weight
30	Shenandoah	1.010	1.00	5.2	0.30
				6.3	0.55
				6.5	0.15
31	Quakers	0.960	1.00	5.8	0.15
				6.5	0.60
				6.8	0.25
33	Niagara	1.000	1.00	5.2	0.30
				6.3	0.55
				6.5	0.15
34	Nessmuk	0.920	1.00	5.2	0.30
				6.3	0.35
				6.5	0.15
37	Charlevoix	0.700	1.00	7.1	0.10
				7.3	0.80
				7.4	0.10
41	Vermont	1.100	1.00	5.2	0.30
				6.3	0.55
				6.5	0.15
50-2	Greenville - Background	1.010	1.00	4.8	0.20
				5.5	0.60
				5.8	0.20



**Table 2.5-7 {Summary of Weston Geophysical Seismic Sources}**

Code	Name	b value	b weight	M <sub>max</sub> (m <sub>b</sub> )	M <sub>max</sub> weight
1	Charlevoix	0.79	1.00	7.2	1.00
6	Adirondack Mountains	0.933	0.50	5.4	0.38
		0.936	0.50	6.0	0.46
				6.6	0.16
8	Clarendon - Linden	0.846	1.00	5.4	0.26
				6.0	0.50
				6.6	0.24
21	New York Nexus	0.934	1.00	5.4	0.62
				6.0	0.29
				6.6	0.09
28 B	Zone of Mesozoic Basin	0.854	1.00	5.4	0.65
				6.0	0.25
				6.6	0.10
28 E	Zone of Mesozoic Basin	0.918	1.00	5.4	0.65
				6.0	0.25
				6.6	0.10
102 Background	Appalachian Plateau	1.007	0.20	5.4	0.62
		1.007	0.80	6.0	0.29
				6.6	0.09
103 Background	Southern Appalachian	0.993	0.20	5.4	0.26
		0.996	0.80	6.0	0.58
				6.6	0.16
104 Background	Southern Coastal Plain	0.997	0.20	5.4	0.24
		0.997	0.80	6.0	0.61
				6.6	0.15

**Table 2.5-8 {Summary of Woodward-Clyde Consultants Seismic Sources}**

Code	Name	b value	b weight	M <sub>max</sub> (m <sub>b</sub> )	M <sub>max</sub> weight
12	Charlevoix	0.70*	1.00	6.5	0.33
				7.0	0.34
				7.5	0.33
18	Adirondack Uplift	1.006	0.25	5.4	0.33
		1.001	0.25	6.3	0.34
		0.930	0.25	6.9	0.33
		0.861	0.25		
20 A	Mohawk River Trend	1.010	0.33	5.5	0.33
		0.922	0.34	6.0	0.34
		0.836	0.33	7.0	0.33
21	NJ Gravity Saddle	0.781	0.25	5.3	0.33
		0.950	0.25	6.5	0.34
		0.873	0.25	6.9	0.33
		0.798	0.25		
25	Hudson River Trend	1.201	0.25	5.5	0.33
		1.056	0.25	6.3	0.34
		0.989	0.25	6.8	0.33
		0.926	0.25		
33	W. NY- S Ontario Seismic Zone	0.951	0.25	5.5	0.33
		0.977	0.25	6.5	0.34
		0.924	0.25	7.0	0.33
		0.873	0.25		
34	Attica, NY Intersection	0.780	0.33	5.6	0.33
		0.725	0.34	6.3	0.34
		0.672	0.33	7.4	0.33
61	Tyrone- Mt. Union Lineament	1.030	0.33	5.4	0.33
		0.939	0.34	6.5	0.34
		0.850	0.33	7.1	0.33
63	Pittsburgh-Washington Lineament	0.997	0.33	5.4	0.33
		0.904	0.34	6.3	0.34
		0.811	0.33	7.1	0.33
B16	Susquehanna Background	0.924	0.25	4.9	0.17
		0.993	0.25	5.4	0.28
		0.903	0.25	5.8	0.27
		0.814	0.25	6.5	0.28

**Table 2.5-9 {Alternative New Madrid Fault Locations}**  
(Metric Units)

Fault arm	New Madrid South		Reelfoot		New Madrid North		Combined Weight
Faults set	Distance to Site (km)	Weight	Distance to Site (km)	Weight	Distance to Site (km)	Weight	
1	1268	0.6	1259	0.7	1213	0.7	0.294
2					1185	0.3	0.126
3			1259	0.3	1213	0.7	0.126
4					1185	0.3	0.054
5	1280	0.4	1259	0.7	1213	0.7	0.196
6					1185	0.3	0.084
7			1259	0.3	1213	0.7	0.084
8					1185	0.3	0.036

**Table 2.5-9—{Alternative New Madrid Fault Locations}**  
(SI Units)

Fault arm	New Madrid South		Reelfoot		New Madrid North		Combined Weight
Faults set	Distance to Site (mile)	Weight	Distance to Site (mile)	Weight	Distance to Site (mile)	Weight	
1	788	0.6	782	0.7	754	0.7	0.294
2					736	0.3	0.126
3			782	0.3	754	0.7	0.126
4					736	0.3	0.054
5	796	0.4	782	0.7	754	0.7	0.196
6					736	0.3	0.084
7			782	0.3	754	0.7	0.084
8					736	0.3	0.036

**Table 2.5-10 {Earthquake Frequencies for Repeating New Madrid Earthquake Sequences}**

Rupture Set	Weight	Magnitude, M			Weight	Rupture Model	Weight	Combined Weight
		NMS	RF	NMN				
1	0.1667	7.8	7.7	7.5	1	A	0.667	0.1112
		7.8	7.7	7.5	0.333	B	0.333	0.0185
		7.3	7.7	7.5	0.333			0.0185
		7.8	7.7	7.0	0.333			0.0185
2	0.1667	7.9	7.8	7.6	1	A	0.667	0.1112
		7.9	7.8	7.6	0.333	B	0.333	0.0185
		7.4	7.8	7.6	0.333			0.0185
		7.9	7.8	7.1	0.333			0.0185
3	0.2500	7.6	7.8	7.5	1	A	0.667	0.1668
		7.6	7.8	7.5	0.333	B	0.333	0.0277
		7.1	7.8	7.5	0.333			0.0277
		7.6	7.8	7.0	0.333			0.0277
4	0.0833	7.2	7.4	7.2	1	A	0.667	0.0556
		7.2	7.4	7.2	0.333	B	0.333	0.0092
		7.0	7.4	7.2	0.333			0.0092
		7.2	7.4	7.2	0.333			0.0092
5	0.1667	7.2	7.4	7.0	1	A	0.667	0.1112
		7.2	7.4	7.0	0.333	B	0.333	0.0185
		7.0	7.4	7.0	0.333			0.0185
		7.2	7.4	7.0	0.333			0.0185
6	0.1667	7.3	7.5	7.0	1	A	0.667	0.1112
		7.3	7.5	7.0	0.333	B	0.333	0.0185
		7.0	7.5	7.0	0.333			0.0185
		7.3	7.5	7.0	0.333			0.0185

Notes:

- 1 ) Adapted from the Seismic Hazards Report for the EGC ESP Site  
See Table 4.1-2 of Appendix B (EGC, 2006)
- 2) NMS indicates New Madrid South Arm
- 3) RF indicates New Madrid Reelfoot Arm
- 4) NMN indicates New Madrid North Arm

**Table 2.5-11 {Controlling Earthquakes for BBNPP}**  
(Metric Units)

Hazard	Reference Earthquake (RE)		De-aggregation Earthquakes (DE)			
	Mag. (M)	Dist. (km)	Event	Mag. (M)	Dist. (km)	Weight
Mean 10 <sup>-4</sup> 1 and 2.5Hz	7.19	286.25	DEL	5.55	16.84	0.203
			DEM	6.45	112.09	0.249
			DEH	7.36	339.47	0.547
Mean 10 <sup>-4</sup> 5 and 10 Hz	5.83	43.65	DEL	5.40	17.62	0.551
			DEM	6.18	107.41	0.357
			DEH	7.09	307.72	0.091
Mean 10 <sup>-5</sup> 1 and 2.5Hz	7.30	287.56	DEL	5.79	13.67	0.387
			DEM	6.86	105.56	0.211
			DEH	7.38	340.39	0.402
Mean 10 <sup>-5</sup> 5 and 10 Hz	5.71	15.90	DEL	5.54	11.62	0.862
			DEM	6.77	100.70	0.122
			DEH	7.11	291.43	0.015

Note: Distance range of each event

DEL: 0 to 50 km

DEM: 50 to 200 km

DEH: > 200 km

**Table 2.5-11—{Controlling Earthquakes for BBNPP}**  
(SI Units)

Hazard	Reference Earthquake (RE)		De-aggregation Earthquakes (DE)			
	Mag. (M)	Dist. (mile)	Event	Mag. (M)	Dist. (mile)	Weight
Mean 10 <sup>-4</sup> 1 and 2.5Hz	7.19	171.40	DEL	5.55	10.46	0.203
			DEM	6.45	69.66	0.249
			DEH	7.36	210.98	0.547
Mean 10 <sup>-4</sup> 5 and 10 Hz	5.83	26.14	DEL	5.40	10.95	0.551
			DEM	6.18	66.76	0.357
			DEH	7.09	191.25	0.091
Mean 10 <sup>-5</sup> 1 and 2.5Hz	7.30	172.19	DEL	5.79	8.49	0.387
			DEM	6.86	65.61	0.211
			DEH	7.38	211.55	0.402
Mean 10 <sup>-5</sup> 5 and 10 Hz	5.71	9.52	DEL	5.54	7.22	0.862
			DEM	6.77	62.59	0.122
			DEH	7.11	181.13	0.015

Note: Distance range of each event

DEL: 0 to 31 mile

DEM: 31 to 125 mile

DEH: > 125 mile

**Table 2.5-12 {Recommended Horizontal and Vertical SSE Amplitudes and common V/H Ratio}**

Freq (Hz)	Horizontal SSE (g)	Vertical SSE (g)	V/H
0.1000	0.0056	0.0032	0.5818
0.1269	0.0084	0.0049	0.5818
0.1610	0.0123	0.0071	0.5818
0.2043	0.0174	0.0101	0.5818
0.2593	0.0242	0.0141	0.5818
0.3290	0.0327	0.0190	0.5818
0.4175	0.0433	0.0252	0.5818
0.5000	0.0529	0.0308	0.5818
0.6723	0.0598	0.0348	0.5818
0.8532	0.0616	0.0358	0.5818
1.0000	0.0588	0.0342	0.5818
1.3738	0.0757	0.0449	0.5932
1.7500	0.0896	0.0539	0.6020
2.2122	0.1035	0.0633	0.6114
2.5000	0.1101	0.0679	0.6169
3.5622	0.1488	0.0942	0.6330
4.5204	0.1830	0.1200	0.6558
5.0000	0.2091	0.1402	0.6705
7.5000	0.2870	0.2078	0.7240
10.0000	0.3630	0.2718	0.7487
11.7210	0.4152	0.3221	0.7757
14.8735	0.4743	0.3823	0.8060
18.8739	0.5089	0.4192	0.8238
25.0000	0.5471	0.4904	0.8963
30.3920	0.5265	0.4911	0.9328
38.5662	0.4987	0.4999	1.0025
48.9390	0.4420	0.4620	1.0452
62.1017	0.3515	0.3738	1.0635
78.8046	0.2509	0.2527	1.0071
100.0000	0.2090	0.1851	0.8857

**Table 2.5-13 {Amplification Factors for 10<sup>-4</sup> and 10<sup>-5</sup> Input Motions and HF and LF  
Rock Spectra}**

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Freq(Hz)	10 <sup>-4</sup> LF				10 <sup>-4</sup> HF			
	DEL	DEM	DEH	Ave.	DEL	DEM	DEH	Ave.
	0.0133	0.0387	0.9481		0.0535	0.0880	0.8585	
0.1000	1.0468	1.0040	1.0014	1.0113	1.1959	1.0478	1.0000	1.1251
0.1269	1.0817	1.0003	1.0001	1.0168	1.0980	1.0117	1.0004	1.0583
0.1610	1.0211	0.9989	1.0002	1.0042	1.0414	1.0056	1.0001	1.0248
0.2043	1.0036	1.0002	1.0001	1.0008	1.0307	1.0050	1.0003	1.0188
0.2593	0.9989	1.0000	1.0002	0.9999	1.0167	1.0006	1.0006	1.0095
0.3290	0.9984	1.0013	1.0001	1.0000	1.0027	0.9999	1.0001	1.0015
0.4175	0.9988	1.0005	1.0006	1.0002	1.0053	1.0012	1.0002	1.0033
0.5000	1.0021	1.0015	1.0006	1.0011	1.0048	1.0010	1.0006	1.0031
0.6723	1.0016	1.0011	1.0010	1.0012	1.0020	1.0014	1.0012	1.0017
0.8532	1.0025	1.0024	1.0019	1.0021	1.0038	1.0030	1.0021	1.0034
1.0000	1.0052	1.0022	1.0034	1.0035	1.0057	1.0037	1.0025	1.0047
1.3738	1.0085	1.0057	1.0044	1.0056	1.0079	1.0074	1.0049	1.0075
1.7500	1.0100	1.0123	1.0077	1.0093	1.0099	1.0109	1.0073	1.0100
2.2122	1.0155	1.0133	1.0142	1.0142	1.0142	1.0147	1.0122	1.0142
2.5000	1.0191	1.0172	1.0145	1.0161	1.0234	1.0173	1.0168	1.0206
3.5622	1.0373	1.0357	1.0329	1.0345	1.0365	1.0373	1.0323	1.0364
4.5204	1.0551	1.0608	1.0540	1.0559	1.0594	1.0609	1.0543	1.0595
5.0000	1.0575	1.0642	1.0646	1.0631	1.0642	1.0692	1.0621	1.0658
7.5000	1.1091	1.1297	1.1284	1.1248	1.1330	1.1113	1.1244	1.1245
10.0000	1.1981	1.1916	1.1823	1.1878	1.2029	1.1867	1.1969	1.1966
11.7210	1.2085	1.2067	1.2211	1.2150	1.2254	1.2141	1.2105	1.2200
14.8735	1.2176	1.2033	1.2228	1.2169	1.2184	1.2216	1.2189	1.2196
18.8739	1.1757	1.1840	1.1933	1.1874	1.1883	1.1775	1.1927	1.1848
25.0000	1.1946	1.1653	1.1615	1.1692	1.1788	1.1599	1.1667	1.1710
30.3920	1.1641	1.1681	1.1643	1.1652	1.1627	1.1509	1.1583	1.1581
38.5662	1.1361	1.1507	1.1696	1.1581	1.1541	1.1598	1.1560	1.1563
48.9390	1.1440	1.1631	1.1148	1.1328	1.1599	1.1471	1.1384	1.1533
62.1017	1.1506	1.1433	1.1171	1.1304	1.1563	1.1489	1.1193	1.1503
78.8046	1.1368	1.1661	1.1585	1.1560	1.1760	1.1572	1.1760	1.1693
100.0000	1.2615	1.2905	1.3655	1.3256	1.2993	1.3190	1.3514	1.3111
Note: Distance range of each event DEL: 0 to 50 km (0 to 31 mile) DEM: 50 to 200 km (31 to 125 mile) DEH: > 200 km (> 125 mile)								

**Table 2.5-13—{Amplification Factors for 10-4 and 10-5 Input Motions and HF  
and LF Rock Spectra}**

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Freq(Hz)	10 <sup>-5</sup> LF				10 <sup>-5</sup> HF			
	DEL	DEM	DEH	Ave.	DEL	DEM	DEH	Ave.
	0.0332	0.0370	0.9297		0.1532	0.0811	0.7657	
0.1000	1.0389	0.9990	1.0005	<b>1.0150</b>	1.0421	0.9978	1.0014	<b>1.0360</b>
0.1269	1.0107	1.0009	1.0000	<b>1.0043</b>	1.0662	0.9997	1.0000	<b>1.0570</b>
0.1610	0.9975	1.0008	1.0001	<b>0.9992</b>	1.0123	1.0014	1.0002	<b>1.0107</b>
0.2043	0.9989	1.0006	1.0000	<b>0.9997</b>	1.0097	0.9998	1.0004	<b>1.0083</b>
0.2593	1.0024	1.0004	1.0003	<b>1.0011</b>	1.0065	1.0011	1.0008	<b>1.0058</b>
0.3290	1.0009	1.0002	1.0002	<b>1.0005</b>	1.0023	1.0007	1.0006	<b>1.0021</b>
0.4175	0.9997	1.0004	1.0003	<b>1.0001</b>	1.0012	1.0003	1.0002	<b>1.0010</b>
0.5000	1.0010	1.0005	1.0008	<b>1.0008</b>	1.0013	1.0009	1.0003	<b>1.0012</b>
0.6723	1.0017	1.0011	1.0009	<b>1.0012</b>	1.0012	1.0015	1.0009	<b>1.0012</b>
0.8532	1.0020	1.0018	1.0018	<b>1.0019</b>	1.0042	1.0013	1.0025	<b>1.0038</b>
1.0000	1.0031	1.0031	1.0034	<b>1.0032</b>	1.0038	1.0030	1.0030	<b>1.0036</b>
1.3738	1.0048	1.0048	1.0045	<b>1.0047</b>	1.0062	1.0055	1.0046	<b>1.0061</b>
1.7500	1.0095	1.0078	1.0066	<b>1.0080</b>	1.0056	1.0106	1.0078	<b>1.0062</b>
2.2122	1.0131	1.0111	1.0137	<b>1.0129</b>	1.0121	1.0122	1.0121	<b>1.0121</b>
2.5000	1.0178	1.0158	1.0149	<b>1.0162</b>	1.0197	1.0185	1.0167	<b>1.0195</b>
3.5622	1.0419	1.0329	1.0334	<b>1.0366</b>	1.0435	1.0366	1.0321	<b>1.0425</b>
4.5204	1.0571	1.0530	1.0565	<b>1.0560</b>	1.0580	1.0587	1.0550	<b>1.0580</b>
5.0000	1.0683	1.0616	1.0674	<b>1.0666</b>	1.0798	1.0635	1.0623	<b>1.0775</b>
7.5000	1.1315	1.1294	1.1270	<b>1.1292</b>	1.1416	1.1336	1.1237	<b>1.1403</b>
10.0000	1.2165	1.1996	1.1868	<b>1.2010</b>	1.1953	1.1968	1.1923	<b>1.1955</b>
11.7210	1.2160	1.2149	1.2288	<b>1.2209</b>	1.2212	1.2044	1.2167	<b>1.2191</b>
14.8735	1.2122	1.2271	1.2197	<b>1.2184</b>	1.2143	1.2027	1.2168	<b>1.2130</b>
18.8739	1.1868	1.1851	1.1783	<b>1.1830</b>	1.1702	1.1808	1.1907	<b>1.1718</b>
25.0000	1.1708	1.1654	1.1568	<b>1.1640</b>	1.1740	1.1614	1.1635	<b>1.1723</b>
30.3920	1.1603	1.1644	1.1609	<b>1.1614</b>	1.1552	1.1577	1.1505	<b>1.1554</b>
38.5662	1.1500	1.1633	1.1645	<b>1.1586</b>	1.1547	1.1631	1.1645	<b>1.1558</b>
48.9390	1.1457	1.1284	1.1177	<b>1.1308</b>	1.1438	1.1573	1.1357	<b>1.1453</b>
62.1017	1.1395	1.1318	1.1297	<b>1.1339</b>	1.1424	1.1273	1.1439	<b>1.1406</b>
78.8046	1.1927	1.1437	1.1551	<b>1.1672</b>	1.1340	1.1439	1.1682	<b>1.1357</b>
100.0000	1.3012	1.3143	1.2837	<b>1.2969</b>	1.2311	1.2979	1.3019	<b>1.2404</b>
Note: Distance range of each event DEL: 0 to 50 km (0 to 31 mile) DEM: 50 to 200 km (31 to 125 mile) DEH: > 200 km (> 125 mile)								



**Table 2.5-14 {Uniform Hazard Response Spectra (Hard Rock Conditions)}**

Frequency (Hz)	10 <sup>-4</sup> SA, g		10 <sup>-5</sup> SA, g		10 <sup>-6</sup> SA, g	
	Mean	Median	Mean	Median	Mean	Median
0.5	0.0357	0.0219	0.1103	0.0560	0.2437	0.1321
1.0	0.0454	0.0320	0.1183	0.0757	0.2606	0.1741
2.5	0.0807	0.0653	0.2202	0.1914	0.5565	0.5330
5.0	0.1307	0.1238	0.4068	0.3990	1.1081	1.0739
10.0	0.1888	0.1850	0.6474	0.6366	1.8252	1.7707
25.0	0.2802	0.2648	1.0043	0.9680	3.1000	2.8950
100.0	0.0973	0.0945	0.3611	0.3516	1.0914	1.0350

**Table 2.5-15 {Earthquake Frequencies for Repeating New Madrid Earthquake Sequences}**

Model	Weight of model	Mean repeat time (years)	Equivalent annual frequency, $\lambda_{\text{rate of cluster}}$	Weight of $\lambda_{\text{rate of cluster}}$	Combined weight, $w_c$	$\lambda_{\text{rate of cluster times } w_c}$
Poisson	0.5	161	6.20E-03	0.10108	0.05054	0.00031
		262	3.82E-03	0.24429	0.12215	0.00047
		410	2.44E-03	0.30926	0.15463	0.00038
		694	1.44E-03	0.24429	0.12215	0.00018
		1563	6.40E-04	0.10108	0.05054	0.00003
Renewal $\alpha = 0.3$	0.5x0.2 = 0.1	333	3.39E-03	0.10108	0.01011	0.00003
		410	1.07E-03	0.24429	0.02443	0.00003
		485	3.02E-04	0.30926	0.03093	0.00001
		574	5.95E-05	0.24429	0.02443	0.00000
		709	4.30E-06	0.10108	0.01011	0.00000
Renewal $\alpha = 0.5$	0.5x0.5 = 0.25	316	4.85E-03	0.10108	0.02527	0.00012
		440	2.18E-03	0.24429	0.06107	0.00013
		573	8.89E-04	0.30926	0.07732	0.00007
		746	2.58E-04	0.24429	0.06107	0.00002
		1032	2.97E-05	0.10108	0.02527	0.00000
Renewal $\alpha = 0.7$	0.5x0.3 = 0.15	325	4.45E-03	0.10108	0.01516	0.00007
		506	2.25E-03	0.24429	0.03664	0.00008
		719	1.02E-03	0.30926	0.04639	0.00005
		1011	3.37E-04	0.24429	0.03664	0.00001
		1521	4.49E-05	0.10108	0.01516	0.00000
sums	1.00				1.00	0.00199
Average return period =					503	years

Note:

Adapted from the Seismic Hazards Report for the Clinton ESP Site  
See Figure 4.1-1 and Table 4.1-3 (EGC 2006)

**Table 2.5-16 {USGS 2008 Seismicity Smoothing Models}**

Model	Incompleteness Year	Minimum Magnitude, $m_b$	Correlation Distance		Weight	
			km	mile	Main Source	Background
1	1924	3.0	50	31	0.50	0.40
2	1860	4.0	75	47	0.25	0.20
3	1700	5.0	75	47	0.25	0.20
4	Uniformly distributed		-	-	-	0.20

**Table 2.5-17 {Selected Controlling Rock Motion Time Histories}**

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***10<sup>-4</sup> High Frequency Event***

Hazard	Controlling Event			Earthquake Acceleration Time History				
	Mag. (M)	Dist. (km)	Dist. (mi)	Event	Station	Mag. (M)	Dist. (km)	Dist. (mi)
10 <sup>-4</sup> Near field event, DEL	5.55	16.84	10.47	Livermore	San Roamon - eastman Kodak	5.4	17.6	10.9
						5.4	17.6	10.9
				Coalinga	Palmer Ave.	5.8	12.2	7.6
						5.8	12.2	7.6
10 <sup>-4</sup> Middle field event, DEM	6.45	112.09	69.66	San Fernando	Maricopa Array #3	6.6	113.0	70.2
						6.6	113.0	70.2
				Northridge	Riverside Airport	6.7	101.3	63.0
						6.7	101.3	63.0
10 <sup>-4</sup> Far field event, DEH	7.36	339.47	210.98	Landers	San Gabriel - E Gran Ave	7.3	141.6	88.0
						7.3	141.6	88.0
				Landers	Sun Valley - Sunland	7.3	162.6	101.1
						7.3	162.6	101.1

**Table 2.5-17—{Selected Controlling Rock Motion Time Histories}**

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***10<sup>-4</sup> High Frequency Event***

Hazard	Controlling Event			Earthquake Acceleration Time History				
	Mag. (M)	Dist. (km)	Dist. (mi)	Event	Station	Mag. (M)	Dist. (km)	Dist. (mi)
10 <sup>-4</sup> Near field event, DEL	5.40	17.62	10.95	Livermore	San Roamon - eastman Kodak	5.4	17.6	10.9
						5.4	17.6	10.9
				Coalinga	Palmer Ave.	5.8	12.2	7.6
						5.8	12.2	7.6
10 <sup>-4</sup> Middle field event, DEM	6.18	107.41	6.76	Borrego Mtn	San Onofre - So Cal Edison	6.8	124.7	77.5
						6.8	124.7	77.5
				San Fernando	San Onofre - So Cal Edison	6.6	122.0	75.8
						6.6	122.0	75.8
10 <sup>-4</sup> Far field event, DEH	7.09	307.72	191.25	Landers	San Gabriel - E Gran Ave	7.3	141.6	88.0
						7.3	141.6	88.0
				Tabas, Iran	Kashmar	7.4	199.1	123.7
						7.4	199.1	123.7

**Table 2.5-17—{Selected Controlling Rock Motion Time Histories}**

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**10<sup>-5</sup> High Frequency Event**

Hazard	Controlling Event			Earthquake Acceleration Time History				
	Mag. (M)	Dist. (km)	Dist. (mi)	Event	Station	Mag. (M)	Dist. (km)	Dist. (mi)
10 <sup>-5</sup> Near field event, DEL	5.79	13.67	8.50	Morgan Hill	Gilroy - Gavilan Coll.	6.2	16.2	10.1
						6.2	16.2	10.1
				Morgan Hill	Gilroy Array #6	6.2	11.8	7.3
						6.2	11.8	7.3
10 <sup>-5</sup> Middle field event, DEM	6.86	105.56	65.61	Landers	Duarte - Mel Canyon Rd	7.3	126.4	78.6
						7.3	126.4	78.6
				Landers	Villa Park - Serrano Ave	7.3	131.4	81.7
						7.3	131.4	81.7
10 <sup>-5</sup> Far field event, DEH	7.38	340.39	211.55	Landers	Villa Park - Serrano Ave	7.3	194.1	120.6
						7.3	194.1	120.6
				Landers	Sun Valley - Sunland	7.3	162.6	101.1
						7.3	162.6	101.1

**Table 2.5-17—{Selected Controlling Rock Motion Time Histories}**

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***10<sup>-5</sup> High Frequency Event***

Hazard	Controlling Event			Earthquake Acceleration Time History				
	Mag. (M)	Dist. (km)	Dist. (mi)	Event	Station	Mag. (M)	Dist. (km)	Dist. (mi)
10 <sup>-5</sup> Near field event, DEL	5.54	11.62	7.22	Whittier Narrows	Garvey Res. - Control Bldg.	6	12.1	7.5
						6	12.1	7.5
				Friuli, Italy	San Rocco	5.5	17.9	11.1
						5.5	17.9	11.1
10 <sup>-5</sup> Middle field event, DEM	6.77	100.70	62.59	San Fernando	Maricopa Array #3	6.6	113.0	70.2
						6.6	113.0	70.2
				Northridge	Riverside Airport	6.7	101.3	101.3
						6.7	101.3	101.3
10 <sup>-5</sup> Far field event, DEH	7.11	291.43	181.12	Landers	San Gabriel - E Gran Ave	7.3	141.6	88.0
						7.3	141.6	88.0
				Tabas, Iran	Kashmar	7.4	199.1	123.7
						7.4	199.1	123.7

**Table 2.5-18 Comparison of Post-EPRI NP-6395-D 1989 Magnitude Estimates for the 1886 Charleston Earthquake**

Study	Magnitude Estimation Method	Reported Magnitude Estimate	Assigned Weights	Mean Magnitude (M)
EPRI (1994)	worldwide survey of passive-margin, extended-crust earthquakes	<b>M</b> 7.56 ± 0.35	--	7.56
Martin (1994)	geotechnical assessment of 1886 liquefaction data	<b>M</b> 7 - 7.5	--	7.25
Johnston (1996)	isoseismal area regression, accounting for eastern North America anelastic attenuation	<b>M</b> 7.3 ± 0.26	--	7.3
Chapman (2002) (South Carolina Department of Transportation)	consideration of available magnitude estimates	<b>M</b> 7.1 <b>M</b> 7.3 <b>M</b> 7.5	0.2 0.6 0.2	7.3
Frankel et al. (2002) (USGS National seismic hazard mapping project)	consideration of available magnitude estimates	<b>M</b> 6.8 <b>M</b> 7.1 <b>M</b> 7.3 <b>M</b> 7.5	0.20 0.20 0.45 0.15	7.2
Bakun (2004)	isoseismal area regression, including empirical site corrections	<b>M</b> <sub>I</sub> 6.4 - 7.2	--	6.9
95% confidence interval estimate; <b>M</b> <sub>I</sub> (intensity magnitude) is considered equivalent to <b>M</b> (Bakun and Hopper 2004) Adopted from the Seismic Hazard Report for CCNPP Unit 3 FSAR 2007				

**Table 2.5-19 {Mean and Fractile Rock Hazard Curves for PGA}**

Ground motion level (g)	Annual Exceedance Frequency					
	Mean	5% Fractile	16% Fractile	50% Fractile	84% Fractile	95% Fractile
0.010	3.785E-03	2.404E-03	2.809E-03	3.664E-03	4.779E-03	5.584E-03
0.025	1.035E-03	5.204E-04	6.539E-04	9.649E-04	1.424E-03	1.789E-03
0.050	3.238E-04	1.620E-04	2.038E-04	3.016E-04	4.463E-04	5.616E-04
0.075	1.583E-04	8.525E-05	1.049E-04	1.494E-04	2.128E-04	2.618E-04
0.100	9.521E-05	5.413E-05	6.550E-05	9.067E-05	1.255E-04	1.519E-04
0.200	2.846E-05	1.744E-05	2.061E-05	2.741E-05	3.644E-05	4.308E-05
0.300	1.399E-05	8.279E-06	9.894E-06	1.340E-05	1.816E-05	2.170E-05
0.400	8.306E-06	4.607E-06	5.617E-06	7.878E-06	1.105E-05	1.347E-05
0.500	5.443E-06	2.817E-06	3.509E-06	5.102E-06	7.418E-06	9.240E-06
0.600	3.792E-06	1.836E-06	2.333E-06	3.511E-06	5.283E-06	6.713E-06
0.700	2.755E-06	1.252E-06	1.621E-06	2.519E-06	3.913E-06	5.067E-06
0.800	2.065E-06	8.845E-07	1.165E-06	1.864E-06	2.983E-06	3.930E-06
0.900	1.585E-06	6.422E-07	8.598E-07	1.414E-06	2.325E-06	3.112E-06
1.000	1.240E-06	4.771E-07	6.482E-07	1.093E-06	1.843E-06	2.505E-06
1.250	7.162E-07	2.451E-07	3.441E-07	6.136E-07	1.094E-06	1.536E-06
1.500	4.425E-07	1.370E-07	1.976E-07	3.693E-07	6.901E-07	9.959E-07
2.000	1.935E-07	5.082E-08	7.661E-08	1.543E-07	3.106E-07	4.682E-07
2.750	6.950E-08	1.519E-08	2.401E-08	5.243E-08	1.145E-07	1.810E-07
3.500	2.944E-08	5.606E-09	9.174E-09	2.125E-08	4.921E-08	8.054E-08
5.000	7.153E-09	1.113E-09	1.914E-09	4.821E-09	1.214E-08	2.088E-08



**Table 2.5-20 {Mean and Fractile Rock Hazard Curves for 25 Hz}**

Ground motion level (g)	Annual Exceedance Frequency					
	Mean	5% Fractile	16% Fractile	50% Fractile	84% Fractile	95% Fractile
0.010	8.506E-03	6.105E-03	6.856E-03	8.354E-03	1.018E-02	1.143E-02
0.025	3.601E-03	2.214E-03	2.614E-03	3.470E-03	4.606E-03	5.439E-03
0.050	1.566E-03	8.127E-04	1.011E-03	1.468E-03	2.132E-03	2.653E-03
0.075	8.854E-04	4.200E-04	5.370E-04	8.164E-04	1.241E-03	1.587E-03
0.100	5.707E-04	2.572E-04	3.338E-04	5.208E-04	8.125E-04	1.055E-03
0.200	1.806E-04	7.682E-05	1.014E-04	1.628E-04	2.615E-04	3.452E-04
0.300	8.872E-05	3.888E-05	5.088E-05	8.050E-05	1.273E-04	1.666E-04
0.400	5.315E-05	2.444E-05	3.153E-05	4.870E-05	7.522E-05	9.706E-05
0.500	3.562E-05	1.711E-05	2.179E-05	3.293E-05	4.975E-05	6.337E-05
0.600	2.565E-05	1.272E-05	1.605E-05	2.385E-05	3.546E-05	4.473E-05
0.700	1.940E-05	9.814E-06	1.231E-05	1.811E-05	2.665E-05	3.342E-05
0.800	1.521E-05	7.762E-06	9.709E-06	1.422E-05	2.083E-05	2.606E-05
0.900	1.225E-05	6.251E-06	7.819E-06	1.145E-05	1.678E-05	2.099E-05
1.000	1.008E-05	5.106E-06	6.401E-06	9.411E-06	1.384E-05	1.735E-05
1.250	6.622E-06	3.230E-06	4.095E-06	6.139E-06	9.203E-06	1.167E-05
1.500	4.655E-06	2.154E-06	2.775E-06	4.271E-06	6.576E-06	8.469E-06
2.000	2.609E-06	1.075E-06	1.432E-06	2.336E-06	3.810E-06	5.077E-06
2.750	1.318E-06	4.601E-07	6.423E-07	1.134E-06	2.004E-06	2.797E-06
3.500	7.563E-07	2.290E-07	3.324E-07	6.275E-07	1.185E-06	1.720E-06
5.000	3.095E-07	7.477E-08	1.152E-07	2.407E-07	5.031E-07	7.752E-07

**Table 2.5-21 {Mean and Fractile Rock Hazard Curves for 10 Hz}**

Ground motion level (g)	Annual Exceedance Frequency					
	Mean	5% Fractile	16% Fractile	50% Fractile	84% Fractile	95% Fractile
0.010	8.106E-03	5.951E-03	6.633E-03	7.979E-03	9.598E-03	1.070E-02
0.025	2.799E-03	1.770E-03	2.072E-03	2.708E-03	3.539E-03	4.141E-03
0.050	1.014E-03	5.615E-04	6.851E-04	9.619E-04	1.350E-03	1.648E-03
0.075	5.191E-04	2.812E-04	3.454E-04	4.904E-04	6.963E-04	8.553E-04
0.100	3.146E-04	1.750E-04	2.132E-04	2.985E-04	4.181E-04	5.094E-04
0.200	9.009E-05	5.639E-05	6.620E-05	8.703E-05	1.144E-04	1.343E-04
0.300	4.289E-05	2.801E-05	3.244E-05	4.166E-05	5.350E-05	6.196E-05
0.400	2.517E-05	1.653E-05	1.911E-05	2.447E-05	3.133E-05	3.622E-05
0.500	1.653E-05	1.073E-05	1.245E-05	1.604E-05	2.067E-05	2.399E-05
0.600	1.162E-05	7.400E-06	8.641E-06	1.126E-05	1.466E-05	1.712E-05
0.700	8.570E-06	5.329E-06	6.269E-06	8.271E-06	1.091E-05	1.284E-05
0.800	6.536E-06	3.966E-06	4.702E-06	6.286E-06	8.404E-06	9.964E-06
0.900	5.115E-06	3.028E-06	3.618E-06	4.902E-06	6.641E-06	7.936E-06
1.000	4.086E-06	2.361E-06	2.843E-06	3.901E-06	5.355E-06	6.447E-06
1.250	2.492E-06	1.359E-06	1.666E-06	2.357E-06	3.336E-06	4.089E-06
1.500	1.629E-06	8.418E-07	1.049E-06	1.526E-06	2.221E-06	2.767E-06
2.000	7.966E-07	3.744E-07	4.800E-07	7.332E-07	1.120E-06	1.436E-06
2.750	3.365E-07	1.400E-07	1.860E-07	3.019E-07	4.900E-07	6.510E-07
3.500	1.660E-07	6.221E-08	8.516E-08	1.454E-07	2.484E-07	3.400E-07
5.000	5.326E-08	1.672E-08	2.403E-08	4.462E-08	8.285E-08	1.191E-07

Table 2.5-22 {Mean and Fractile Rock Hazard Curves for 5 Hz}

Ground motion level (g)	Annual Exceedance Frequency					
	Mean	5% Fractile	16% Fractile	50% Fractile	84% Fractile	95% Fractile
0.010	6.489E-03	4.510E-03	5.119E-03	6.351E-03	7.880E-03	8.943E-03
0.025	1.854E-03	1.034E-03	1.259E-03	1.760E-03	2.461E-03	2.996E-03
0.050	5.974E-04	2.475E-04	3.292E-04	5.355E-04	8.710E-04	1.159E-03
0.075	2.900E-04	1.086E-04	1.487E-04	2.540E-04	4.339E-04	5.941E-04
0.100	1.693E-04	6.507E-05	8.844E-05	1.492E-04	2.517E-04	3.421E-04
0.200	4.336E-05	2.238E-05	2.790E-05	4.062E-05	5.915E-05	7.373E-05
0.300	1.895E-05	1.110E-05	1.331E-05	1.813E-05	2.470E-05	2.961E-05
0.400	1.037E-05	6.187E-06	7.375E-06	9.949E-06	1.342E-05	1.600E-05
0.500	6.414E-06	3.753E-06	4.501E-06	6.136E-06	8.364E-06	1.003E-05
0.600	4.285E-06	2.430E-06	2.943E-06	4.079E-06	5.654E-06	6.848E-06
0.700	3.020E-06	1.654E-06	2.025E-06	2.859E-06	4.036E-06	4.941E-06
0.800	2.214E-06	1.170E-06	1.448E-06	2.083E-06	2.996E-06	3.708E-06
0.900	1.673E-06	8.529E-07	1.067E-06	1.564E-06	2.291E-06	2.866E-06
1.000	1.295E-06	6.378E-07	8.063E-07	1.202E-06	1.793E-06	2.267E-06
1.250	7.387E-07	3.356E-07	4.346E-07	6.753E-07	1.049E-06	1.359E-06
1.500	4.578E-07	1.930E-07	2.554E-07	4.119E-07	6.642E-07	8.791E-07
2.000	2.065E-07	7.623E-08	1.048E-07	1.802E-07	3.100E-07	4.261E-07
2.750	8.009E-08	2.498E-08	3.597E-08	6.699E-08	1.247E-07	1.796E-07
3.500	3.715E-08	1.004E-08	1.503E-08	2.987E-08	5.937E-08	8.884E-08
5.000	1.090E-08	2.328E-09	3.702E-09	8.164E-09	1.800E-08	2.863E-08

**Table 2.5-23 {Mean and Fractile Rock Hazard Curves for 2.5 Hz}**

Ground motion level (g)	Annual Exceedance Frequency					
	Mean	5% Fractile	16% Fractile	50% Fractile	84% Fractile	95% Fractile
0.010	3.765E-03	2.187E-03	2.629E-03	3.598E-03	4.924E-03	5.918E-03
0.025	9.133E-04	3.131E-04	4.393E-04	7.828E-04	1.395E-03	1.957E-03
0.050	2.621E-04	4.515E-05	7.574E-05	1.829E-04	4.419E-04	7.413E-04
0.075	1.169E-04	1.470E-05	2.658E-05	7.297E-05	2.003E-04	3.621E-04
0.100	6.322E-05	7.472E-06	1.371E-05	3.855E-05	1.085E-04	1.989E-04
0.200	1.266E-05	2.401E-06	3.933E-06	9.125E-06	2.117E-05	3.468E-05
0.300	4.697E-06	1.457E-06	2.101E-06	3.923E-06	7.322E-06	1.056E-05
0.400	2.299E-06	9.031E-07	1.220E-06	2.036E-06	3.400E-06	4.592E-06
0.500	1.313E-06	5.464E-07	7.259E-07	1.178E-06	1.912E-06	2.540E-06
0.600	8.257E-07	3.369E-07	4.502E-07	7.377E-07	1.209E-06	1.615E-06
0.700	5.545E-07	2.161E-07	2.926E-07	4.902E-07	8.215E-07	1.112E-06
0.800	3.905E-07	1.444E-07	1.984E-07	3.410E-07	5.860E-07	8.051E-07
0.900	2.851E-07	1.000E-07	1.395E-07	2.458E-07	4.331E-07	6.038E-07
1.000	2.141E-07	7.142E-08	1.010E-07	1.822E-07	3.289E-07	4.650E-07
1.250	1.147E-07	3.407E-08	4.971E-08	9.465E-08	1.802E-07	2.629E-07
1.500	6.740E-08	1.809E-08	2.711E-08	5.407E-08	1.078E-07	1.617E-07
2.000	2.786E-08	6.277E-09	9.846E-09	2.121E-08	4.571E-08	7.170E-08
2.750	9.717E-09	1.768E-09	2.926E-09	6.908E-09	1.631E-08	2.700E-08
3.500	4.126E-09	6.286E-10	1.086E-09	2.760E-09	7.014E-09	1.212E-08
5.000	1.048E-09	1.199E-10	2.217E-10	6.316E-10	1.800E-09	3.326E-09

**Table 2.5-24 {Mean and Fractile Rock Hazard Curves for 1 Hz}**

Ground motion level (g)	Annual Exceedance Frequency					
	Mean	5% Fractile	16% Fractile	50% Fractile	84% Fractile	95% Fractile
0.010	1.435E-03	4.617E-04	6.593E-04	1.210E-03	2.221E-03	3.172E-03
0.025	3.220E-04	3.477E-05	6.510E-05	1.897E-04	5.529E-04	1.035E-03
0.050	8.283E-05	3.159E-06	7.386E-06	3.142E-05	1.337E-04	3.126E-04
0.075	3.311E-05	8.230E-07	2.093E-06	1.027E-05	5.043E-05	1.282E-04
0.100	1.611E-05	3.393E-07	8.904E-07	4.612E-06	2.389E-05	6.269E-05
0.200	2.256E-06	5.317E-08	1.366E-07	6.822E-07	3.407E-06	8.752E-06
0.300	6.488E-07	2.453E-08	5.744E-08	2.451E-07	1.046E-06	2.450E-06
0.400	2.665E-07	1.711E-08	3.592E-08	1.272E-07	4.507E-07	9.464E-07
0.500	1.346E-07	1.337E-08	2.552E-08	7.680E-08	2.312E-07	4.412E-07
0.600	7.744E-08	1.003E-08	1.801E-08	4.885E-08	1.326E-07	2.381E-07
0.700	4.858E-08	6.978E-09	1.223E-08	3.182E-08	8.282E-08	1.451E-07
0.800	3.240E-08	4.657E-09	8.160E-09	2.123E-08	5.524E-08	9.678E-08
0.900	2.261E-08	3.092E-09	5.482E-09	1.455E-08	3.863E-08	6.848E-08
1.000	1.634E-08	2.084E-09	3.755E-09	1.025E-08	2.798E-08	5.042E-08
1.250	8.088E-09	8.560E-10	1.610E-09	4.729E-09	1.389E-08	2.613E-08
1.500	4.466E-09	3.994E-10	7.804E-10	2.444E-09	7.657E-09	1.496E-08
2.000	1.673E-09	1.138E-10	2.360E-10	8.184E-10	2.838E-09	5.886E-09
2.750	5.237E-10	2.625E-11	5.805E-11	2.246E-10	8.692E-10	1.922E-09
3.500	2.050E-10	8.162E-12	1.892E-11	7.932E-11	3.325E-10	7.708E-10
5.000	4.632E-11	1.317E-12	3.263E-12	1.532E-11	7.198E-11	1.783E-10

**Table 2.5-25 {Mean and Fractile Rock Hazard Curves for 0.5 Hz}**

Ground motion level (g)	Annual Exceedance Frequency					
	Mean	5% Fractile	16% Fractile	50% Fractile	84% Fractile	95% Fractile
0.010	7.646E-04	1.285E-04	2.168E-04	5.292E-04	1.291E-03	2.180E-03
0.025	1.851E-04	8.350E-06	1.887E-05	7.577E-05	3.042E-04	6.875E-04
0.050	5.573E-05	8.105E-07	2.278E-06	1.327E-05	7.727E-05	2.172E-04
0.075	2.506E-05	2.383E-07	7.229E-07	4.793E-06	3.178E-05	9.638E-05
0.100	1.311E-05	1.027E-07	3.221E-07	2.263E-06	1.590E-05	4.987E-05
0.200	1.936E-06	1.215E-08	3.961E-08	2.968E-07	2.225E-06	7.250E-06
0.300	4.986E-07	3.141E-09	1.023E-08	7.659E-08	5.734E-07	1.867E-06
0.400	1.731E-07	1.181E-09	3.794E-09	2.776E-08	2.031E-07	6.524E-07
0.500	7.306E-08	5.662E-10	1.780E-09	1.254E-08	8.837E-08	2.778E-07
0.600	3.543E-08	3.227E-10	9.862E-10	6.624E-09	4.449E-08	1.360E-07
0.700	1.910E-08	2.096E-10	6.198E-10	3.936E-09	2.499E-08	7.390E-08
0.800	1.118E-08	1.499E-10	4.276E-10	2.553E-09	1.525E-08	4.349E-08
0.900	6.983E-09	1.147E-10	3.154E-10	1.768E-09	9.911E-09	2.724E-08
1.000	4.597E-09	9.164E-11	2.430E-10	1.281E-09	6.754E-09	1.791E-08
1.250	1.920E-09	5.513E-11	1.363E-10	6.382E-10	2.988E-09	7.389E-09
1.500	9.521E-10	3.139E-11	7.555E-11	3.377E-10	1.510E-09	3.633E-09
2.000	3.165E-10	9.298E-12	2.289E-11	1.063E-10	4.941E-10	1.216E-09
2.750	9.195E-11	2.030E-12	5.280E-12	2.694E-11	1.374E-10	3.575E-10
3.500	3.418E-11	6.008E-13	1.631E-12	8.951E-12	4.912E-11	1.333E-10
5.000	7.492E-12	1.035E-13	2.937E-13	1.738E-12	1.028E-11	2.918E-11

**Table 2.5-26 {Summary Of Thicknesses And Termination Elevations For Various Strata}**

English Units

	Thickness (feet)			Top (feet msl)		
	Minimum	Maximum	Average	Minimum	Maximum	Average
Glacial Overburden	12.5	62.0	38.7	656.6	801.3	678.7
Mahantango Formation	*	*	*	594.3	774.2	639.3

Note: \* This layer was not fully penetrated

**Table 2.5-26—{Summary Of Thicknesses And Termination Elevations For Various Strata}**

SI Units

	Thickness (m)			Top (m msl)		
	Minimum	Maximum	Average	Minimum	Maximum	Average
Glacial Overburden	3.8	18.9	11.8	200.2	244.3	206.9
Mahantango Formation	*	*	*	181.2	236.0	194.9

Note: \* This layer was not fully penetrated

**Table 2.5-27 {Summary Of Geotechnical Field Tests}**

Test	Specification	Quantity
Soil and Rock Borings	ASTM D1586/1587	48
Seismic Refraction Survey	NA	6
P-S Suspension Logging Surveys (boreholes)	NA	4
Downhole Velocity Measurements	NA	4
Pressuremeter Test (PMT)	ASTM D4719-00	8
SPT Hammer Energy Measurements	ASTM D4633	2

**Table 2.5-28 {Boring Locations And Surface Elevations}**

Structural (Foundation)			Boring	Horizontal Coordinates		El. (ft)	Boring Depth (ft)			Type of Test
				Plant N	Plant E		Soil	Rock	Total	
POWER BLOCK AREA	Nuclear Island and Reactor Auxiliary Building	UJA UJH UJK UFA UKA	B-301	339,151.79	2,405,430.68	666.38	41.0	359.0	400.0	SPT/Coring
			B-302	339,243.09	2,405,420.59	668.31	28.0	322.5	350.5	SPT/Coring
			B-303	339,142.99	2,405,338.53	664.18	26.5	125.0	151.5	SPT/Coring
			B-304	339,060.24	2,405,438.62	660.78	43.0	307.0	350.0	SPT/Coring
			B-305	339,160.19	2,405,520.63	668.32	45.0	54.7	99.7	SPT/Coring
			B-306	339,313.21	2,405,413.69	669.07	20.0	183.0	203.0	SPT/Coring
			B-307	339,193.33	2,405,276.10	665.31	26.8	73.2	100.0	SPT/Coring
			B-308	339,069.30	2,405,288.63	662.41	39.0	62.3	101.3	SPT/Coring
			B-309	338,998.80	2,405,333.72	659.64	45.0	156.2	201.20	SPT/Coring
			B-310	338,987.79	2,405,444.97	658.46	51.0	49.0	100.0	SPT/Coring
			B-311	339,099.65	2,405,592.49	665.43	50.0	50.8	100.8	SPT/Coring
			B-312	339,230.05	2,405,581.98	669.08	34.8	65.9	100.7	SPT/Coring
	Radioactive Waste Building	UKS	B-313	338,917.24	2,405,379.26	657.42	54.7	45.3	100.0	SPT/Coring
			B-314	338,916.47	2,405,288.20	658.29	62.0	138.0	200.0	SPT/Coring
			B-315	338,820.62	2,405,297.70	656.90	60.0	40.0	100.0	SPT/Coring
	Emergency Power Generation Building	UBP	B-316	338,882.24	2,405,513.39	656.82	54.0	46.0	100.0	SPT/Coring
			B-317	338,888.05	2,405,571.70	656.59	55.0	45.0	100.0	SPT/Coring
			B-318	339,436.13	2,405,520.54	672.30	27.0	73.6	100.6	SPT/Coring
			B-319	339,429.85	2,405,462.38	672.20	28.2	73.5	101.7	SPT/Coring
	ESWS Cooling Towers	URB	B-320	340,491.80	2,405,516.01	657.52	46.2	53.8	100.0	SPT/Coring
			B-321	338,752.66	2,405,830.21	658.02	41.3	58.7	100.0	SPT/Coring
			B-322	338,922.54	2,405,754.79	657.24	61.0	138.5	199.5	SPT/Coring
			B-323	338,927.92	2,405,815.58	657.68	58.5	41.5	100.0	SPT/Coring
			B-324	339,323.66	2,405,190.96	669.45	21.0	79.6	100.6	SPT/Coring
			B-325	339,328.29	2,405,252.94	668.90	20.9	79.6	100.5	SPT/Coring
			B-326	339,454.71	2,405,176.41	675.31	21.5	78.5	100.0	SPT/Coring
			B-327	339,452.09	2,405,233.06	675.38	20.5	180.5	201.0	SPT/Coring
	Turbine Building	UMA	B-328	339,176.76	2,405,699.74	667.16	57.6	42.3	99.9	SPT/Coring
			B-329	339,189.56	2,405,802.66	667.49	59.0	11.0	70.0	SPT/Coring
			B-330	339,200.55	2,405,916.03	668.91	61.2	8.8	70.0	SPT/Coring



**Table 2.5-28 {Boring Locations And Surface Elevations}**

Structural (Foundation)			Boring	Horizontal Coordinates		El. (ft)	Boring Depth (ft)			Type of Test
				Plant N	Plant E		Soil	Rock	Total	
OTHER STRUCTURES	ESWEMS	UHSp ond	B-331	339,872.65	2,406,407.01	683.70	41.0	29.0	70.0	SPT/Coring
			B-332	339,907.33	2,406,874.31	691.55	49.0	21.0	70.0	SPT/Coring
			B-333	339,667.26	2,406,421.42	678.29	34.0	36.3	70.3	SPT/Coring
			B-334	339,700.60	2,406,888.49	671.21	45.5	25.0	70.5	SPT/Coring
	Cooling Towers		B-335	340,767.34	2,405,475.56	793.71	19.0	51.0	70.0	SPT/Coring
			B-336	340,492.18	2,405,516.32	801.32	30.0	40.0	70.0	SPT/Coring
			B-337	340,239.46	2,405,528.14	790.57	19.0	80.5	99.5	SPT/Coring
			B-338	340,738.30	2,406,234.46	720.08	22.8	48.7	71.5	SPT/Coring
			B-339	340,479.98	2,406,149.45	710.66	18.0	52.0	70.0	SPT/Coring
			B-340	340,298.18	2,406,433.93	702.37	12.5	67.5	80.0	SPT/Coring
	UHS Pumphouse and Supply Line <sup>1</sup>		B-341	339,825.65	2,406,458.83	680.71	49.4	50.6	100.0	SPT/Coring
			B-342	339,721.50	2,406,467.46	678.15	38.0	62.0	100.0	SPT/Coring
			B-343	339,772.49	2,406,467.46	681.17	48.4	51.6	100.0	SPT/Coring
			B-344	339,761.98	2,406,301.53	683.12	42.0	58.0	100.0	SPT/Coring
			B-345	339,746.44	2,406,203.70	688.02	25.0	75.0	100.0	SPT/Coring
AUX	AuxiliaryGeophysics		G-301	339,151.79	2,405,430.68	666.81	40.0	365.0	405.0	SPT/Coring
			G-302	339,297.55	2,405,219.00	668.25	21.0	183.0	204.0	SPT/Coring
			G-303	338,698.96	2,405,865.49	659.18	42.5	162.5	205.0	SPT/Coring

**Table 2.5-29 {Summary of Hammer-Rod Energy Measurements}**

Drill Rig	Measurement in Boring No.	ETR Range (%)	Average ETR (%)	Energy Adjustment (ETR%/60%)
CME - 55 340665	B-336	75-87	84	1.40
CME 55 300 Carrier	B-327	60-80	78	1.30

Note: ETR = Percentage of theoretical hammer energy measured in the field

**Table 2.5-30 {Summary Of Field-Measured Standard Penetration Test (Spt) N-Values}**

Stratum	SPT N-values ( blows / feet )		
	Minimum	Maximum	Average
Glacial Overburden	0	131	35

**Table 2.5-31 {Summary Of Adjusted Spt N-Values Based On Energy Measurements}**

Stratum	Adjusted Minimum N-Value (blows/feet)	Adjusted Maximum N-Value (blows/feet)	Adjusted Average N-Value (blows/feet)	Recommended N-Value for Engineering Purposes (blows/ft)
Glacial Overburden	0	97	27	20

**Table 2.5-32 {Summary Of Borehole Pressuremeter Test Results}**

English Units

Boring	Depth (ft)	E <sub>p</sub> (ksf)	p <sub>creep</sub> (ksf)	p <sub>limit</sub> (ksf)
B-301	45	12932	*	*
B-301	55	29098	*	*
B-322	64	104006	*	*
B-322	74	9124	*	*
B-325	22	22908	*	*
B-325	31.5	25932	*	*
B-327	25	9986	*	*
B-327	35	11122	*	*

**Table 2.5-32—{Summary Of Borehole Pressuremeter Test Results}**

SI Units

Boring	Depth (m)	E <sub>p</sub> (kPa)	p <sub>l</sub> (kPa)	E <sub>s</sub> (kPa)
B-301	13.9	619200	*	*
B-301	17.0	1393200	*	*
B-322	19.8	4979800	*	*
B-322	22.9	436900	*	*
B-325	6.8	1096800	*	*
B-325	9.8	1241600	*	*
B-327	7.7	478100	*	*
B-327	10.8	532500	*	*

Note:

\* Tests terminated before creep and limit pressures could be established

p<sub>l</sub> - Pressuremeter limit pressureE<sub>p</sub> - Pressuremeter Modulus

**Table 2.5-33 {Summary Of Laboratory Tests}**

Test	Specification	Quantity
Engineering Classification	ASTM D2487-06 / ASTM D5878-05	114
Moisture Content	ASTM D2216-05	35
Unit Weight	From weight-volume relationship	19
Specific Gravity	ASTM D854-06 / ASTM D6473-99(2005)	13
Grain Size Analysis	ASTM D422-63 (2002)	114
Hydraulic Conductivity	ASTM D5084-03	3
Compaction Tests	ASTM D1557-07 Method C	2
Unconfined Compression	ASTM D7012-04	19
Resonant Column Torsional Shear	Technical Procedures for RCTS Tests <sup>(1)</sup>	5
Free-Free Test	Technical Procedures for URC Tests <sup>(2)</sup>	8
Sonic Pulse Test	ASTM D 2845-05	3
Organic Content	ASTM D2974-07	2
pH	ASTM D4792	38
Resistivity	ASTM G187-05	38
Chloride ion content	AASHTO T291	38
Sulphate ion content	AASHTO T290	38

<sup>(1)</sup> PBRCTS-1 Rev. 4, October 2004, University of Texas at Austin, Performed at FUGRO

<sup>(2)</sup> URC-1 Rev. 1, August 2004, University of Texas at Austin, Performed at FUGRO

**Table 2.5-34 {Summary Of Moisture Content}**

(Page 1 of 2)

Formation	Boring	Depth		Moisture Content (%)
		(ft)	(m)	
Glacial Overburden	B-301	5.8	1.8	3.8
	B-301	8.8	2.7	2.6
	B-302	0.8	0.2	16.7
	B-302	2.0	0.6	13.2
	B-302	8.8	2.7	3.3
	B-303	6.3	1.9	7.0
	B-303	9.3	2.8	13.6
	B-303	12.3	3.7	5.2
	B-303	2.5	0.8	9.9
	B-305	2.3	0.7	4.6
	B-305	3.8	1.1	8.6
	B-305	6.8	2.1	14.1
	B-308	5.8	1.8	10.2
	B-308	13.3	4.0	11.2
	B-309	5.3	1.6	20.9
	B-309	8.3	2.5	20.6
	B-309	11.3	3.4	18.9
	B-311	0.8	0.2	9.0
	B-311	5.3	1.6	4.7
	B-311	6.8	2.1	4.7
	B-317	0.8	0.2	21.2
	B-317	3.8	1.1	19.9
	B-317	14.3	4.3	25.6
	B-318	0.8	0.2	13.5
	B-318	3.8	1.1	13.2
	B-319	4.0	1.2	12.8
	B-324	2.3	0.7	6.3
	B-324	5.3	1.6	3.4
	B-324	8.3	2.5	3.2
	B-326	2.3	0.7	4.4
	B-326	6.8	2.1	10.8
	B-327	3.0	0.9	10.4
	B-333	5.8	1.8	12.2
	B-333	13.3	4.0	5.7
	B-303 U2	2.5	0.8	9.9
	B-319 U2	4.0	1.2	12.8
	B-327 ST2	3.0	0.9	10.4

10.8

**Table 2.5-34 {Summary Of Moisture Content}**

(Page 2 of 2)

Formation	Boring	Depth		Moisture Content (%)
		(ft)	(m)	
Mahantango Formation	B-301	197.8	60.3	NR
	B-301	152.4	46.4	NR
	B-302	318.1	97.0	0.3
	B-302	52.3	15.9	0.4
	B-302	214.1	65.3	0.4
	B-302	110.6	33.7	0.5
	B-303	142.7	43.5	0.0
	B-304	109.8	33.5	0.3
	B-304	170.8	52.1	0.3
	B-304	273.5	83.4	0.3
	B-309	54.7	16.7	0.6
	B-310	57.2	17.4	0.6
	B-318	93.5	28.5	0.0
	B-319	41.1	12.5	NR
	B-324	27.6	8.4	0.6
	B-326	41.1	12.5	1.2
	B-327	34.0	10.4	1.1
	B-331	67.5	20.6	0.9
	B-334	64.6	19.7	NR
				0.5

**Table 2.5-35 {Summary Of Unit Weight Tests Special Care Rock Samples And Undisturbed Samples}**

Formation	Boring	Depth (ft)	Dry Unit Weight (pcf)		Moist Unit Weight (pcf)	
Glacial Overburden	B-303 U2	2.5	99	102	109	121
	B-319 U2	4.0	118		133	
	B-327 ST2	3.0	86		95	
	B-331 ST2	17.0	105		126	
	B-310 U3	6.0	100		124	
Mahantango Formation	B-301, R20	197.8	NR	171	170	171
	B-301, R15	152.4	NR		171	
	B-302, R32	318.1	172		173	
	B-302, R3	52.3	171		172	
	B-302, R21	214.1	172		173	
	B-302, R11	110.6	172		173	
	B-303, R24	142.7	172		172	
	B-304, R8	109.8	172		173	
	B-304, R-14	170.8	172		173	
	B-304, R26	273.5	171		172	
	B-309, R7	54.7	NR		ND	
	B-310, R2	57.2	170		171	
	B-318, R14	93.5	172		172	
	B-319, R4	41.1	NR		172	
	B-324, R2	27.6	171		172	
	B-326, R1	41.1	166		168	
	B-327, R3	34.0	164		166	
	B-331, R7	67.5	170		172	
	B-334, R9	64.6	NR		172	

NR - Not Recorded

ND - Not Detected

**Table 2.5-35—{Summary Of Unit Weight Tests Special Care Rock Samples  
And Undisturbed Samples}**

Formation	Boring	Depth (m)	Dry Unit Weight (N/m <sup>3</sup> )	Moist Unit Weight (N/m <sup>3</sup> )
Glacial Overburden	B-303 U2	0.8	15550	17090
	B-303 U2	1.2	18470	20840
	B-319 U2	0.9	13510	14910
	B-331 ST2	5.2	16500	19810
	B-310 U3	1.8	15730	19540
Mahantango Formation	B-301, R20	60.3	NR	26770
	B-301, R15	46.4	NR	26890
	B-302, R32	97.0	27030	27110
	B-302, R3	15.9	26930	27030
	B-302, R21	65.3	27010	27110
	B-302, R11	33.7	27090	27220
	B-303, R24	43.5	27080	27080
	B-304, R8	33.5	27030	27110
	B-304, R-14	52.1	27020	27100
	B-304, R26	83.4	26920	27000
	B-309, R7	16.7	NR	ND
	B-310, R2	17.4	26700	26860
	B-318, R14	28.5	27030	27030
	B-319, R4	12.5	NR	26990
	B-324, R2	8.4	26900	27070
	B-326, R1	12.5	26080	26390
	B-327, R3	10.4	25730	26010
	B-331, R7	20.6	26730	26970
	B-334, R9	19.7	NR	26990

NR - Not Recorded

ND - Not Detected



**Table 2.5-36 {Summary Of Specific Gravity Tests Special Care Rock Samples And Undisturbed Samples}**

English Units

Formation	Boring	Depth (ft)	Depth (m)	SG
Glacial Overburden	B-303 U2	2.5	0.8	2.69
	B-319 U2	4.0	1.2	2.72
	B-327 ST2	3.0	0.9	2.69
Mahantango Formation	B-301 R20	197.8	60.3	2.77
	B-302 R3	152.4	46.4	2.76
	B-302 R11	318.1	97.0	2.76
	B-304 R8	52.3	15.9	2.76
	B-310 R2	214.1	65.3	2.77
	B-318 R14	110.6	33.7	2.76
	B-319 R4	142.7	43.5	2.83
	B-324 R2	109.8	33.5	2.76
	B-326 R1	170.8	52.1	2.67
	B-334 R9	273.5	83.4	2.77

**Table 2.5-37 {Chemical Test Results Of Soil And Rock Samples}**

Location		Depth		Formation	pH	Resistivity ( $\Omega$ -cm)		Sulfates* (ppm)	Chlorides* (ppm)
Boring	ID	(ft)	(m)			As Received	100% Saturated		
B-301	S-12	18.5	5.6	Glacial Overburden	7.7	18130	12432	94	20
B-301	R-1	41.0	12.5	Mahantango Formation	8.6	>3.11e6	566000	62	<10
B-302	S-10	17.0	5.2	Glacial Overburden	7.1	59570	41440	70	<10
B-302	R-1	28.5	8.7	Mahantango Formation	8.8	>2.13e6	776000	29	<10
B-303	S-3	7.0	2.1	Glacial Overburden	7.0	62160	41440	73	<10
B-303	S-5	10.0	3.0	Glacial Overburden	7.1	18907	17353	43	<10
B-303	R-1	30.0	9.1	Mahantango Formation	8.0	>5.90e6	4.34E+06	20	<10
B-308	S-2	3.5	1.1	Glacial Overburden	6.4	59570	38850	24	<10
B-310	S-5	12.0	3.7	Glacial Overburden	7.0	20202	18389	81	<10
B-310	S-27	45.0	13.7	Glacial Overburden	4.3	11396	7511	1443	13
B-311	R-2	50.8	15.5	Mahantango Formation	8.7	>3.2e6	1.89E+06	21	<10
B-312	S-5	6.0	1.8	Glacial Overburden	8.5	699300	108780	20	13
B-312	S-6	7.5	2.3	Glacial Overburden	7.3	284900	62160	26	<10
B-312	S-7	9.0	2.7	Glacial Overburden	7.1	602100	73590	23	13
B-312	S-8	10.5	3.2	Glacial Overburden	7.8	21238	20720	33	13
B-314	S-11	15.0	4.6	Glacial Overburden	8.0	22523	16725	21	13
B-314	S-12	16.5	5.0	Glacial Overburden	7.8	16056	9812	45	<10
B-314	S-13	18.0	5.5	Glacial Overburden	6.0	14272	12934	77	<10
B-314	S-14	19.5	5.9	Glacial Overburden	8.3	7136	5798	127	13
B-314	S-15	21.0	6.4	Glacial Overburden	7.7	21631	12934	111	13
B-315	S-10	22.5	6.9	Glacial Overburden	6.9	33670	28490	64	<10
B-315	S-11	25.0	7.6	Glacial Overburden	6.8	12950	10878	150	<10
B-315	S-12	27.5	8.4	Glacial Overburden	7.7	8697	8028	91	<10
B-317	S-2	1.5	0.5	Glacial Overburden	6.3	31080	25123	42	13
B-317	S-9	12.0	3.7	Glacial Overburden	7.5	13468	10360	74	13
B-317	S-36	52.5	16.0	Glacial Overburden	8.5	7770	5957	366	13
B-318	R-1	28.0	8.5	Mahantango Formation	8.1	>6.29e6	1.54E+06	133	<10
B-320	S-1	0.0	0.0	Glacial Overburden	6.7	2676	2453	19	20
B-320	S-2	1.5	0.5	Glacial Overburden	6.2	38850	20979	46	<10
B-322	S-16	22.5	6.9	Glacial Overburden	7.7	31080	25900	90	<10
B-322	S-17	24.0	7.3	Glacial Overburden	6.5	13209	12691	148	<10
B-322	S-18	25.5	7.8	Glacial Overburden	7.2	33670	28490	89	<10

**Table 2.5-37 {Chemical Test Results Of Soil And Rock Samples}**

Location		Depth		Formation	pH	Resistivity ( $\Omega$ -cm)		Sulfates* (ppm)	Chlorides* (ppm)
Boring	ID	(ft)	(m)			As Received	100% Saturated		
B-324	S-6	9.0	2.7	Glacial Overburden	6.9	468300	26760	71	40
B-324	R-2	22.0	6.7	Glacial Overburden	8.8	>3.73e6	780000	36	<10
B-327	S-5	15.0	4.6	Glacial Overburden	7.8	37910	33450	32	20
B-327	R-1	20.5	6.3	Mahantango Formation	6.5	>6.04e6	2.25E+06	530	<10
Test Pit Face <sup>(1)</sup>		NA	NA	Borrow area material	6.2	96690	82040	16	<10
Test Pit #5 <sup>(1)</sup>		NA	NA	Borrow area material	5.8	73250	52740	13	10

<sup>(1)</sup> Water Soluble Ion Content

<sup>1</sup> Borrow area fill Material

**Table 2.5-38 {Summary Of Unconfined Compressive Strength Tests Special Care Rock Samples}**

English Units

Page 1 of 2

Formation	Boring	Depth (ft)	UCS (psi)	UCS (psf)
Mahantango Formation	B-301, R20	197.8	3687	530900
	B-301, R15	152.4	9929	1429800
	B-302, R32	318.1	13833	1992000
	B-302, R3	52.3	8495	1223300
	B-302, R21	214.1	9042	1302000
	B-302, R11	110.6	8666	1247900
	B-303, R24	142.7	9207	1325800
	B-304, R8	109.8	12070	1738100
	B-304, R-14	170.8	8381	1206900
	B-304, R26	273.5	9924	1429100
	B-309, R7	54.7	(*)	(*)
	B-310, R2	57.2	12580	1811500
	B-318, R14	93.5	7556	1088100
	B-319, R4	41.1	6770	974900
	B-324, R2	27.6	(*)	(*)
	B-326, R1	41.1	(*)	(*)
	B-327, R3	34.0	9006	1296900
	B-331, R7	67.5	9535	1373000
	B-334, R9	64.6	(*)	(*)

(\*) Specimen broke during preparation

**Table 2.5-38—{Summary Of Unconfined Compressive Strength Tests  
Special Care Rock Samples}**

SI Units  
Page 2 of 2

Formation	Boring	Depth (m)	UCS (MPa)	
Mahantango Formation	B-301	60.3	25	64
	B-301	46.4	68	
	B-302	97.0	95	
	B-302	15.9	59	
	B-302	65.3	62	
	B-302	33.7	60	
	B-303	43.5	63	
	B-304	33.5	83	
	B-304	52.1	58	
	B-304	83.4	68	
	B-309	16.7	(*)	
	B-310	17.4	87	
	B-318	28.5	52	
	B-319	12.5	47	
	B-324	8.4	(*)	
	B-326	12.5	(*)	
	B-327	10.4	62	
	B-331	20.6	66	
	B-334	19.7	(*)	

(\*) Specimen broke during preparation

**Table 2.5-39 {Hydraulic Conductivity Test Results}**

English Units  
Page 1 Of 2

Soil Type	Boring	Depth (ft)	Eff. Confining Pressure (psf)	Hydraulic Conductivity (fps)
Glacial Overburden	B-303	2.5	500	6.89E-06
Glacial Overburden	B-319	4.0	500	9.84E-07
Glacial Overburden	B-327	3.0	500	2.89E-06

**Table 2.5-39—{Hydraulic Conductivity Test Results}**

SI Units  
Page 2 Of 2

Soil Type	Boring	Depth (m)	Eff. Confining Pressure (kPA)	Hydraulic Conductivity (cm/s)
Glacial Overburden	B-303	0.8	24	2.10E-04
Glacial Overburden	B-319	1.2	24	3.00E-05
Glacial Overburden	B-327	0.9	24	8.80E-05

**Table 2.5-40 {Dynamic Testing Program Samples}**

Location		Depth		Formation	Test
Boring	ID	(ft)	(m)		
B-331	ST2	17.0	5.2	Glacial Overburden	RCTS
B-310	U3	6.0	1.8	Glacial Overburden	RCTS
B-304	R2	41.0	12.5	Mahantango Formation	RCTS
Test Pit Face				Fill and Backfill Borrow Material	RCTS
Test Pit #5				Fill and Backfill Borrow Material	RCTS
B-301	R-1	48.3	14.7	Mahantango Formation	URC
B-301	R-10	108.4	33.0	Mahantango Formation	URC
B-309	R-34	189.6	57.8	Mahantango Formation	URC
B-301	R-25	242.2	73.8	Mahantango Formation	URC
B-302	R-27	272.1	82.9	Mahantango Formation	URC
B-304	R-30	305.8	93.2	Mahantango Formation	URC
B-302	R-35	344.1	104.9	Mahantango Formation	URC
B-301	R-42	395.0	120.4	Mahantango Formation	URC
B-303	R-4	44.3	13.5	Mahantango Formation	SP
B-309	R-25	147.0	44.8	Mahantango Formation	SP
B-313	R-11	95.9	29.2	Mahantango Formation	SP

RCTS: Resonant Column Torsional Shear

URC: Unconfined Resonant Column, Free-Free Test

SP: Sonic Pulse Test

**Table 2.5-41 {Resonant Column Low Strain Properties}**

English Units

Page 1 of 2

Sample	Location		Depth (ft)	Formation	Confining Pressures (psf)		$\gamma$ (psf)	$G_{max}$ (ksf)	$V_s$ (fps)	D (%)
	Boring	ID								
B-331-ST2	B-331	B-331	17	Glacial Overburden	$\sigma_o(min)$	288	126	472	347	1.32
					$\sigma_o(site)$	1008	126	980	500	1.12
					$\sigma_o(max)$	3888	126	2191	748	0.96
B-310-U3	U3	U3	6	Glacial Overburden	$\sigma_o(min)$	288	124	575	386	1.06
					$\sigma_o(site)$	1152	124	1091	531	0.84
					$\sigma_o(max)$	4752	124	2268	766	0.55
Test Pit Face				Borrow Area Fill	$\sigma_o(min)$	576	133	1093	515	2.07
					$\sigma_o(site)$	2448	133	2564	788	1.75
					$\sigma_o(max)$	9936	133	5735	1179	1.68
Test Pit #5				Borrow Area Fill	$\sigma_o(min)$	576	136	1307	556	1.64
					$\sigma_o(site)$	2448	136	2839	820	1.47
					$\sigma_o(max)$	9936	136	5877	1180	1.34

**Table 2.5-41—{Resonant Column Low Strain Properties}**

SI Units  
Page 2 of 2

Sample	Location		Depth (m)	Formation	Confining Pressures (kPa)		$\gamma$ (kN/m <sup>3</sup> )	$G_{max}$ (MPa)	$V_s$ (m/s)	D (%)
	Boring	ID								
B-331-ST2	B-331	B-331	5.18	Glacial Overburden	$\sigma_o(min)$	14	20	23	106	1.32
					$\sigma_o(site)$	48	20	47	153	1.12
					$\sigma_o(max)$	186	20	105	228	0.96
B-310-U3	U3	U3	1.83	Glacial Overburden	$\sigma_o(min)$	14	20	28	118	1.06
					$\sigma_o(site)$	55	20	52	162	0.84
					$\sigma_o(max)$	228	20	109	234	0.55
Test Pit Face				Borrow Area Fill	$\sigma_o(min)$	28	21	52	157	2.07
					$\sigma_o(site)$	117	21	123	240	1.75
					$\sigma_o(max)$	476	21	275	359	1.68
Test Pit #5				Borrow Area Fill	$\sigma_o(min)$	28	21	63	170	1.64
					$\sigma_o(site)$	117	21	136	250	1.47
					$\sigma_o(max)$	476	21	281	360	1.34



**Table 2.5-42 {"Free-Free" Test Results}**

English Units

Location		Depth (ft)	$\gamma$ (pcf)	$V_s$ (fps)	$D_s$ (%)	$V_c$ (fps)	$D_c$ (%)	$V_p$ (fps)
Boring	ID							
B-301	R-1	48.3	169.9	7680	2.51	12020	1.58	12940
B-301	R-10	108.4	171.0	9030	0.98	13590	1.26	13810
B-309	R-34	189.6	171.5	9360	0.75	14260	1.26	14690
B-301	R-25	242.2	170.9	9670	1.14	15030	1.37	15980
B-302	R-27	272.1	171.6	9290	0.86	14100	1.68	14490
B-304	R-30	305.8	170.8	9460	0.79	14390	1.46	14850
B-302	R-35	344.1	169.5	9600	0.7	14690	1.38	15120
B-301	R-42	395.0	171.1	9670	0.79	14870	1.16	15320
$\gamma$ - Unit Weight $V_s$ - Shear Wave Velocity $D_s$ - Shear Wave Damping $V_c$ - Compression Wave Velocity (Unconstrained Test) $D_c$ - Compression Wave Damping $V_p$ - Compression Wave Velocity (Constrained Test)								

**Table 2.5-42—{"Free-Free" Test Results}**

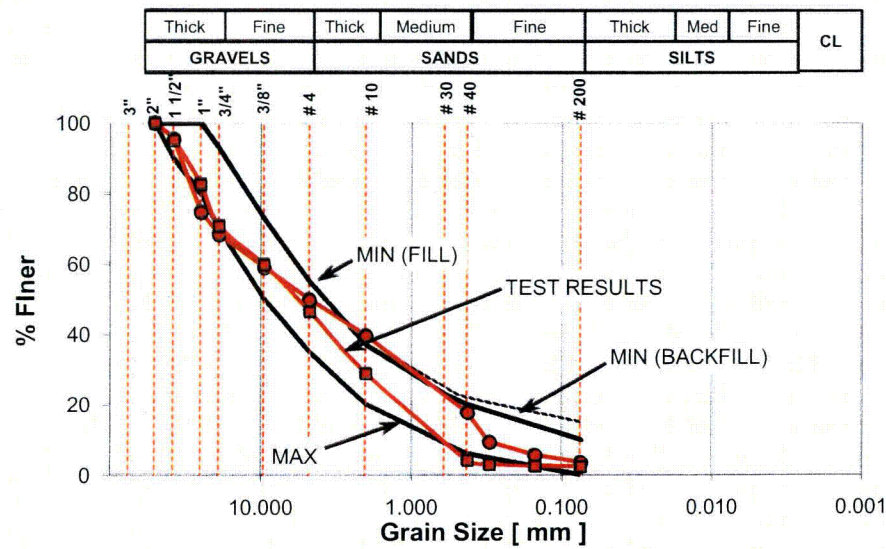
English Units

Location		Depth (m)	$\gamma$ (kN/m <sup>3</sup> )	$V_s$ (m/s)	$D_s$ (%)	$V_c$ (m/s)	$D_c$ (%)	$V_p$ (m/s)
Boring	ID							
B-301	R-1	14.7	26.7	2341	2.51	3665	1.58	3945
B-301	R-10	33.0	26.9	2753	0.98	4143	1.26	4210
B-309	R-34	57.8	26.9	2854	0.75	4348	1.26	4479
B-301	R-25	73.8	26.8	2948	1.14	4582	1.37	4872
B-302	R-27	82.9	27.0	2832	0.86	4299	1.68	4418
B-304	R-30	93.2	26.8	2884	0.79	4387	1.46	4527
B-302	R-35	104.9	26.6	2927	0.7	4479	1.38	4610
B-301	R-42	120.4	26.9	2948	0.79	4534	1.16	4671
$\gamma$ - Unit Weight $V_s$ - Shear Wave Velocity $D_s$ - Shear Wave Damping $V_c$ - Compression Wave Velocity (Unconstrained Test) $D_c$ - Compression Wave Damping $V_p$ - Compression Wave Velocity (Constrained Test)								

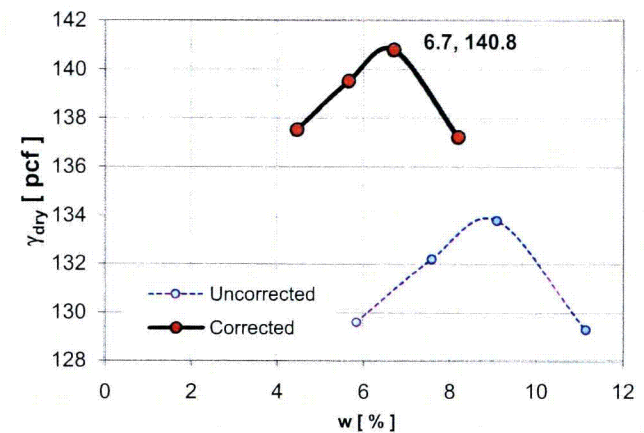
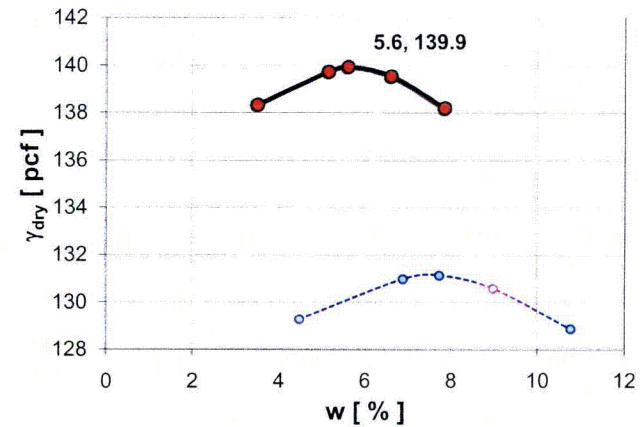
Table 2.5-43 {Category 1 Structural Fill and Backfill Properties}

English Units

Parameter				Category 1 Structural Fill <sup>1</sup>	Category 1 Backfill <sup>2</sup>
Grain Size Analysis	% Gravel		[ % ]	57.0	57.0
	% Sands		[ % ]	37.7	37.7
	% Fines		[ % ]	7.6	7.6
Mass ( in 1 ft <sup>3</sup> )	Solids	m <sub>s</sub>	[ lbm ]	133	126
	Water	m <sub>w</sub>	[ lbm ]	8	8
	Total	m <sub>t</sub>	[ lbm ]	141	134
Comb	Moisture Content	w	[ % ]	6.1	6.1
	Saturation	S	[ % ]	88.3	73.2
	Specific Gravity	SG	[ ]	2.6	2.6
Unit Weight	Dry	γ <sub>dry</sub>	[ pcf ]	133	126
	Moist	γ <sub>moist</sub>	[ pcf ]	141	134
	Saturated	γ <sub>sat</sub>	[ pcf ]	144	140



Modified Proctor Tests



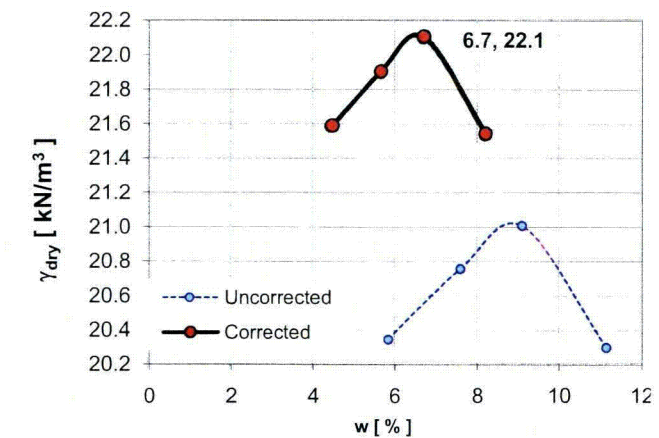
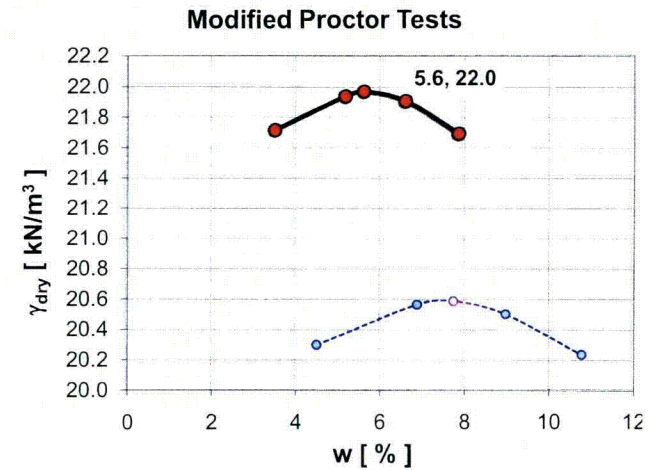
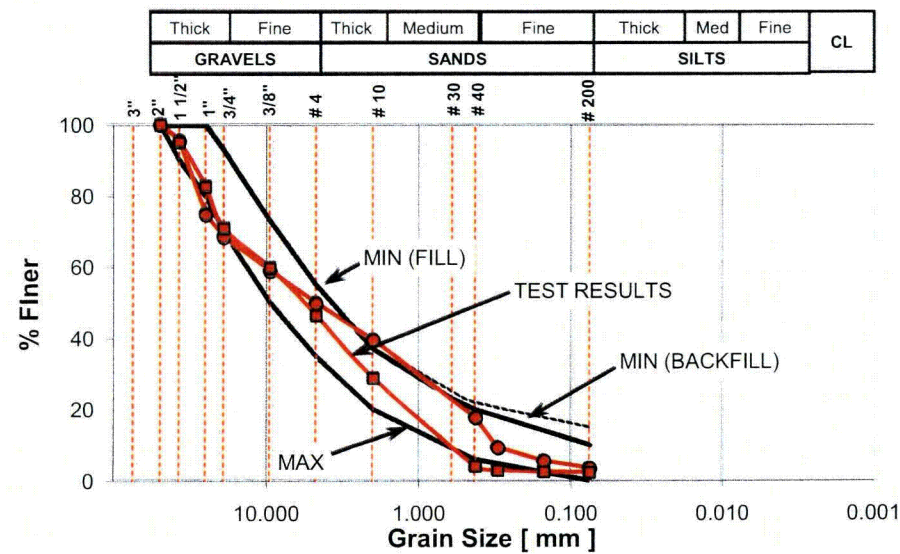
<sup>1</sup> 95% Modified Proctor

<sup>2</sup> 90% Modified Proctor

Table 2.5-44 {Category 1 Structural Fill and Backfill Properties}

SI Units

Parameter				Category 1 Structural Fill <sup>1</sup>	Category 1 Backfill <sup>2</sup>
Grain Size Analysis	% Gravel		[ % ]	57.0	57.0
	% Sands		[ % ]	37.7	37.7
	% Fines		[ % ]	7.6	7.6
Mass ( in 1 m <sup>3</sup> )	Solids	m <sub>s</sub>	[ kg ]	198	188
	Water	m <sub>w</sub>	[ kg ]	12	11
	Total	m <sub>t</sub>	[ kg ]	210	199
Comb	Moisture Content	w	[ % ]	6.1	6.1
	Saturation	S	[ % ]	88.3	73.2
	Specific Gravity	SG	[ ]	2.6	2.6
Unit Weight	Dry	γ <sub>dry</sub>	[ N/m <sup>3</sup> ]	20.9	19.8
	Moist	γ <sub>moist</sub>	[ N/m <sup>3</sup> ]	22.2	21.0
	Saturated	γ <sub>sat</sub>	[ N/m <sup>3</sup> ]	22.7	22.0



<sup>1</sup> 95% Modified Proctor

<sup>2</sup> 90% Modified Proctor

**Table 2.5-45 {Recommended Values Of Index Properties}**

English Units

Unit	USCS or URCS	Water Content (%)	Unit Weight (pcf)			Observations
			Dry	Moist	Sat	
Glacial Overburden	SW	11.0	109	121	144	-URCS Classification: (Weathering, Strength, Discontinuity, Weight) NA: Not Applicable
Mahantango Formation	ABAA	0.5	169	170	170	
Category 1 Granular Fill	SW	6.1	133	141	144	
Category 1 Granular Backfill	SW	6.1	126	134	140	
Concrete Fill	NA	NA	NA	150	NA	NA: Not Applicable

**Table 2.5-45—{Recommended Values Of Index Properties}**

SI Units

Unit	USCS or URCS	Water Content (%)	Unit Weight (kN/m <sup>3</sup> )			Observations
			Dry	Moist	Sat	
Glacial Overburden	SW	11.0	17.1	19.0	22.7	-URCS Classification: (Weathering, Strength, Discontinuity, Weight) NA: Not Applicable
Mahantango Formation	ABAA	0.5	26.6	26.7	26.7	
Category 1 Granular Fill	SW	6.1	20.9	22.2	22.7	
Category 1 Granular Backfill	SW	6.1	19.8	21.1	22.2	
Concrete Fill	NA	NA	NA	23.6	NA	NA: Not Applicable

**Table 2.5-46 {Rock Mass Rating For Mahantango Formation}**

English Units

Item	Value	Rating
Unconfined Compressive Strength (ksf)	1040	12
Rock Quality Designation (%)	83	17
Spacing of Discontinuities (m)	>0.61	20
Condition of Discontinuities	Slightly Rough Weathered Walls	25
Groundwater Conditions	Damp, less than 0.35 cf/min	10
Adjustment for Orientation	Favorable	-2
<b>Total Rating</b>		<b>82</b>
<b>Type of Rock</b>	<b>Very Good Rock</b>	<b>I</b>
<b>Equivalent Cohesion (ksf)</b>	<b>7.3</b>	
<b>Equivalent Friction (°)</b>	<b>40.0</b>	

**Table 2.5-46—{Rock Mass Rating For Mahantango Formation}**

SI Units

Item	Value	Rating
Unconfined Compressive Strength (ksf)	50	12
Rock Quality Designation (%)	83	17
Spacing of Discontinuities (m)	>2	20
Condition of Discontinuities	Slightly Rough Weathered Walls	25
Groundwater Conditions	Damp, less than 10 l/min	10
Adjustment for Orientation	Favorable	-2
<b>Total Rating</b>		<b>82</b>
<b>Type of Rock</b>	<b>Very Good Rock</b>	<b>I</b>
<b>Equivalent Cohesion (kPa)</b>	<b>7.3</b>	
<b>Equivalent Friction (°)</b>	<b>40.0</b>	

**Table 2.5-47 {Recommended Values For Strength Properties}**

English Units

Formation	SPT	c [ ksf ]	φ [ ° ]	q <sub>u</sub> [ ksf ]	Observations
Glacial Overburden	20	0	32.0	NA	<ul style="list-style-type: none"><li>• Friction obtained from SPT Correlation for Dense Sands and Gravels (Peck, 1974)</li><li>• q<sub>u</sub> determined from Unconfined Compressive Test</li><li>• For the Mahantango Formation equivalent cohesion and friction based on Rock Mass Rating (Bieniawski, 1989)- Concrete strength consistent with V<sub>s</sub> = 6800 fps</li><li>• Friction for granular fill based on common practice and conservatism</li><li>• NM: Not Measured</li><li>• NA: Not Applicable</li></ul>
Mahantango Formation	NA	7.3	40.0	1050	
Category 1 Granular Fill	NM	0.0	35.0	NA	
Category 1 Granular Backfill	NM	0.0	35.0	NA	
Concrete Fill	f'c = 5000				

**Table 2.5-47—{Recommended Values For Strength Properties}**

SI Units

Formation	SPT	c [ kPa ]	$\phi$ [ ° ]	q <sub>u</sub> [ kPa ]	Observations
Glacial Overburden	20	0	32.0	NA	<ul style="list-style-type: none"><li>• Friction obtained from SPT Correlation for Dense Sands and Gravels (Peck, 1974)</li><li>• q<sub>u</sub> determined from Unconfined Compressive Test</li><li>• For the Mahantango Formation equivalent cohesion and friction based on Rock Mass Rating (Bieniawski, 1989)- Concrete strength consistent with V<sub>s</sub> = 6800 fps</li><li>• Friction for granular fill based on common practice and conservatism</li><li>• NM: Not Measured</li><li>• NA: Not Applicable</li></ul>
Mahantango Formation	NA	350	40.0	50270	
Category 1 Granular Fill	NM	0.0	35.0	NA	
Category 1 Granular Backfill	NM	0.0	35.0	NA	
Concrete Fill	f'c = 34450				

**Table 2.5-48 {Recommended Values For Hydraulic Conductivity}**  
English Units

Formation	K (Laboratory) (fps)	K (Field) (fps)	K (Recommended) (fps)	Observations
Glacial Overburden	3.6E-06	1.2E-03	1.2E-03	NM: Not Measured Fill and Compacted Fill recommendation based on typical values for clean sands and gravels (Terzaghi, Peck 1967, 1996)
Mahantango Formation	NM	2.0E-05	2.0E-05	
Category 1 Granular Fill	NM	NM	3.3E-03	
Category 1 Granular Backfill	NM	NM	3.3E-03	
Concrete Fill	NM	NM	3.3E-12	

**Table 2.5-48—{Recommended Values For Hydraulic Conductivity}**  
SI Units

Formation	K (Laboratory) (cm/s)	K (Field) (cm/s)	K (Recommended) (cm/s)	Observations
Glacial Overburden	3.6E-06	1.2E-03	1.2E-03	NM: Not Measured Fill and Compacted Fill recommendation based on typical values for clean sands and gravels (Terzaghi, Peck 1967, 1996)
Mahantango Formation	NM	2.0E-05	2.0E-05	
Category 1 Granular Fill	NM	NM	3.3E-03	
Category 1 Granular Backfill	NM	NM	3.3E-03	
Concrete Fill	NM	NM	3.3E-12	

**Table 2.5-49 {Recommended Values For Elastic Modulus}**

English Units

Formation	E (ksf)						Observations
	Method 1	Method 2	Method 3	Method 4	Method 5	Recomm	
Glacial Overburden	2000	3600	2000	-	-	2000	Method 1: ASCE Typical Method 2: ASCE N Correlation Method 3: AASHTO Typical Method 4: Rock Mass Rating Method 5: Concrete strength ACI318
Mahantango Formation	-	-	523000	376000	-	376000	
Category 1 Granular Fill	-	-	2000	-	-	2000	
Category 1 Granular Backfill	-	-	2000	-	-	2000	
Concrete Fill	-	-	-	-	500000	500000	

**Table 2.5-49—{Recommended Values For Elastic Modulus}**

SI Units

Formation	E (MPa)						Observations
	Method 1	Method 2	Method 3	Method 4	Method 5	Recomm	
Glacial Overburden	96	172	96	-	-	96	Method 1: ASCE Typical Method 2: ASCE N Correlation Method 3: AASHTO Typical Method 4: Rock Mass Rating Method 5: Concrete strength ACI318
Mahantango Formation	-	-	25000	18000	-	18000	
Category 1 Granular Fill	-	-	96	-	-	96	
Category 1 Granular Backfill	-	-	96	-	-	96	
Concrete Fill	-	-	-	-	23900	23900	

**Table 2.5-50 {Recommended Values For Static Elastic Properties}**

English Units

Formation	E (ksf)	$\nu$	G (ksf)	Observations
Glacial Overburden	2000	0.40	710	Elastic Properties obtained from field measurements and published values References:- Peck, 1974 - AASHTO, 1998
Mahantango Formation	3.76E+05	0.30	144620	
Category 1 Granular Fill	2000	0.35	740	
Category 1 Granular Backfill	2000	0.35	740	
Concrete Fill	500000	0.20	208300	

**Table 2.5-50—{Recommended Values For Static Elastic Properties}**

SI Units

Formation	E (MPa)	$\nu$	G (MPa)	Observations
Glacial Overburden	100	0.40	36	Elastic Properties obtained from field measurements and published values References:- Peck, 1974- AASHTO, 1998
Mahantango Formation	18000	0.30	6923	
Category 1 Granular Fill	100	0.35	40	
Category 1 Granular Backfill	100	0.35	37	
Concrete Fill	23940	0.20	9975	



**Table 2.5-51 {Recommended Values For Low Strain Dynamic Elastic Properties At Center Of Nuclear Island Footprint}**

English Units

Unit		Depth (ft) (')		V <sub>s</sub> (fps)	V <sub>p</sub> (fps)	ρ (p <sub>m</sub> cf)	ν	G <sub>max</sub> (ksf)	E <sub>max</sub> (ksf)	DS <sub>o</sub> (%)	Observations
ROCK FORMATIONS	Glacial Overburden 1	0.0	20.0	1150	2550	126	0.37	5180	14220	1.00	$G_{max} = \rho V_s^2$ $E_{max} = 2(1 + \nu)G_s$ - Velocities determined fro, best estimate soil profile - Poisson's Ration determined from velocity ratio squared -Initial Shear Damping (DS <sub>o</sub> ): determined from "Free-Free" Testing for Rocks, and RCTS Testing for Overburden
	Glacial Overburden 2	20.0	40.0	2300	7000	126	0.44	20700	59590	1.00	
	Mahantango Formation Layer 1	40.0	60.0	6800	12900	170	0.31	244120	638430	0.80	
	Mahantango Formation Layer 2	60.0	100.0	7150	15300	170	0.36	269900	734290	0.80	
	Mahantango Formation Layer 3	100.0	155.0	7600	16100	170	0.36	304940	827390	0.70	
	Mahantango Formation Layer 4	155.0	220.0	8500	16100	170	0.31	381440	996910	0.70	
	Mahantango Formation Layer 5	220.0	270.0	8950	16750	170	0.30	422900	1099710	0.70	
FILLS	Mahantango Formation Layer 6	270.0	-	9600	16850	170	0.26	486560	1225840	0.70	NA: Not Applicable
	Category 1 Granular Fill/Backfill	0.0	10.0	600	1250	140	0.35	1570	4240	1.00	
		10.0	20.0	800	1670	140	0.35	2780	7510	1.00	
		> 20.0		1000	2100	140	0.35	4350	11770	1.00	
	Concrete	NA	NA	7240	11820	150	0.20	244180	585920	0.80	

**Table 2.5-51—{Recommended Values For Low Strain Dynamic Elastic Properties At Center Of Nuclear Island Footprint}**  
SI Units

Unit		Depth (m) (*)		V <sub>s</sub> (m/s)	V <sub>p</sub> (m/s)	ρ (kg/m <sup>3</sup> )	ν	G <sub>max</sub> (MPa)	E <sub>max</sub> (MPa)	DS <sub>o</sub> (%)	Observations
ROCK FORMATIONS SOILS	Glacial Overburden 1	0.0	6.1	351	777	2020	0.37	250	690	1.00	$G_{max} = \rho V_s^2$ $E_{max} = 2(1 + \nu)G_s$ - Velocities determined fro, best estimate soil profile - Poisson's Ration determined from velocity ratio squared -Initial Shear Damping (DS <sub>o</sub> ): determined from "Free-Free" Testing for Rocks, and RCTS Testing for Overburden
	Glacial Overburden 2	6.1	12.2	701	2134	2020	0.44	990	2850	1.00	
	Mahantango Formation Layer 1	12.2	18.3	2073	3933	2720	0.31	11690	30570	0.80	
	Mahantango Formation Layer 2	18.3	30.5	2180	4665	2720	0.36	12930	35180	0.80	
	Mahantango Formation Layer 3	30.5	47.3	2317	4909	2720	0.36	14600	39610	0.70	
	Mahantango Formation Layer 4	67.1	67.1	2729	5107	2720	0.30	20250	52660	0.70	
	Mahantango Formation Layer 5	82.3	82.3	2927	5137	2720	0.26	23300	58700	0.70	
FILLS	Category 1 Granular Fill/Backfill	0.0	3.0	183	381	2240	0.35	70	190	1.00	NA: Not Applicable
		3.0	6.1	244	509	2240	0.35	130	350	1.00	
		> 6.1		305	640	2240	0.35	210	570	1.00	
	Concrete	NA	NA	2207	3604	2400	0.20	11690	28050	0.80	

**Table 2.5-52 {Peak Ground Acceleration from FIRS Study}**

English Units

Structure		Foundation		Base to Rock (ft) <sup>(2)</sup>	Base to OGS (ft) <sup>(3)</sup>	Eng Fill <sup>4</sup> (ft)	Contact Soil <sup>5</sup>	Hor PGA (g)	Ver PGA (g)
		Depth (ft)	El. <sup>1</sup> (ft msl)						
Nuclear Island <sup>(6)</sup>	NI	36.0	638.0	12.6	-28.0	12.6	C-M	0.21	0.18
ESWEMS	URB1	22.0	652.0	3.0	-23.0	3.0	C	0.21	0.18
Cooling Towers ( URB )	URB2	22.0	652.0	4.5	-17.0	4.5	C	0.21	0.18
	URB3	22.0	652.0	58.5	-5.5	58.5	EF	0.21	0.21
	URB4	22.0	652.0	41.0	-5.5	41.0	EF	0.24	0.25
Emergency Power Building ( UBP )	UBP12 <sup>(7)</sup>	5.0	669.0	27.0	-3.0	27.0	C	0.21	0.19
	UBP12	5.0	669.0	27.0	-3.0	27.0	EF	0.30	0.33
	UBP34	5.0	669.0	67.0	12.0	67.0	EF	0.21	0.22
ESWEMS Pumphouse		33.0	641.0	7.0	-40.0	7.0	C	0.21	0.18

<sup>1</sup> Plant Grade El. (ft msl) 674

<sup>2</sup> Distance between rock and base of foundation

<sup>3</sup> Distance between foundation base original ground surface

<sup>4</sup> Engineered fill (Concrete for NI and soil for other facilities)

<sup>5</sup> Concrete (C), Mahantango Formation (M), Glacial Overburden (GO), Engineered Fill (EF)

<sup>6</sup> Depth given for containment building, other NI facilities have 41.3 ft

<sup>7</sup> Concrete fill is the preferred configuration for this facility

**Table 2.5-52—{Peak Ground Acceleration from FIRS Study}**

SI Units

Structure		Foundation		Base to Rock (m) <sup>(2)</sup>	Base to OGS (m) <sup>(3)</sup>	Eng Fill <sup>4</sup> (m)	Contact Soil <sup>5</sup>	Hor PGA (g)	Ver PGA (g)
		Depth (m)	El. <sup>1</sup> (m msl)						
Nuclear Island <sup>(6)</sup>	NI	11.0	194.5	3.8	-8.5	3.8	C-M	0.21	0.18
ESWEMS	URB1	6.7	198.8	0.9	-7.0	0.9	C	0.21	0.18
Cooling Towers ( URB )	URB2	6.7	198.8	1.4	-5.2	1.4	C	0.21	0.18
	URB3	6.7	198.8	17.8	-1.7	17.8	EF	0.21	0.21
	URB4	6.7	198.8	12.5	-1.7	12.5	EF	0.24	0.25
Emergency Power Building ( UBP )	UBP12 <sup>(7)</sup>	1.5	204.0	8.2	-0.9	8.2	C	0.21	0.19
	UBP12	1.5	204.0	8.2	-0.9	8.2	S	0.30	0.33
	UBP34	1.5	204.0	20.4	3.7	20.4	EF	0.21	0.22
ESWEMS Pumphouse		10.1	195.4	2.1	-12.2	2.1	C	0.21	0.18

<sup>1</sup> Plant Grade El. (m msl) 205

<sup>2</sup> Distance between rock and base of foundation

<sup>3</sup> Distance between foundation base original ground surface

<sup>4</sup> Engineered fill (Concrete for NI and soil for other facilities)

<sup>5</sup> Concrete (C), Mahantango Formation (M), Glacial Overburden (GO), Engineered Fill (EF)

<sup>6</sup> Depth given for containment building, other NI facilities have 12.6 m

<sup>7</sup> Concrete fill is the preferred configuration for this facility

**Table 2.5-53 {Soil Conditions For The U.S. EPR Standard Plant}**

English Units

Soil Case No.	Seismic Control Motion Applied	Soil Profile (Half Space or Layered)	Shear-wave Velocity of Soil <sup>(1)</sup> (fps)
1u	EUR Soft	Half-space	700
2u (A and B)	EUR Soft and Medium	Half-space	1,640
3u	EUR Medium	Half-space	2,625
4u (A and B)	EUR Medium and Hard	Half-space	3,937
5u	EUR Hard	Half-space	5,249
5a	EUR Hard	Half-space	13,123
1n2u	EUR Soft	Linear gradient within a 100 ft layer over a half-space	820 to 1,640
2sn4u	EUR Medium	49 ft uniform layer over a half- space	1,640/3,937
2n3u	EUR Medium	Linear gradient within a 200 ft layer over a half-space	1,640 to 2,625
3r3u	EUR Medium	20 ft uniform layer over 46 ft stiffer layer followed by soil half-space	2,625/5,249/2,625

<sup>(1)</sup> Shear wave velocities of generic soil profiles are taken as strain-compatible properties**Table 2.5-53—{Soil Conditions For The U.S. EPR Standard Plant}**

SI Units

Soil Case No.	Seismic Control Motion Applied	Soil Profile (Half Space or Layered)	Shear-wave Velocity of Soil <sup>(1)</sup> (m/s)
1u	EUR Soft	Half-space	215
2u (A and B)	EUR Soft and Medium	Half-space	500
3u	EUR Medium	Half-space	800
4u (A and B)	EUR Medium and Hard	Half-space	1,200
5u	EUR Hard	Half-space	1,600
5a	EUR Hard	Half-space	4,000
1n2u	EUR Soft	Linear gradient within a 30 m layer over a half-space	250 to 500
2sn4u	EUR Medium	15 m uniform layer over a half- space	500/1,200
2n3u	EUR Medium	Linear gradient within a 60 m layer over a half-space	500 to 800
3r3u	EUR Medium	6 m uniform layer over 14 m stiffer layer followed by soil half-space	800/1,600/800

<sup>(1)</sup> Shear wave velocities of generic soil profiles are taken as strain-compatible properties

**Table 2.5-54 {Soil Conditions For The U.S. EPR Standard Plant}**  
English Units

Site Class		Shear Wave Velocity (mean) (fps)	Location of GMRS	Geotechnical Analysis Requirements
<b>Rock Site (Rock is at the ground surface at the site)</b>	<b>Hard and Firm Rock Site</b>	$V_s > 3,500$	At top of Rock	Static and dynamic bearing capacity to be verified; no time-dependent settlement analysis required.
	<b>Soft Rock Site</b>	$2,400 < V_s < 3,500$	At top of Rock	Static and dynamic bearing capacity to be verified; no time-dependent settlement analysis required.
<b>Thin Soil Site (Rock is generally within 40 to 60 feet of the ground surface and the EPR Nuclear Island is founded on rock)</b>	<b>Thin Soil Site over Hard or Firm Rock</b>	At depth below Nuclear Island Basemat, $V_s > 3,500$	At Top of Outcropping Rock	Static and dynamic bearing capacity to be verified; no time-dependent settlement analysis required.
	<b>Thin Soil Site over Soft Rock</b>	At depth below Nuclear Island Basemat, $V_s < 3,500$	At Grade Elevation	Static and dynamic bearing capacity to be verified; no time-dependent settlement analysis required.
<b>Soil Sites (Foundation underlain by &lt; 200' of soil for Shallow and &gt; 200' feet for Deep)</b>	<b>Shallow Soil and Deep Soil Sites</b>	$1,000 < V_s < 3,500$ with soil below Nuclear Island Basemat of unlimited thickness	At free-field soil surface	Static and dynamic bearing capacity analysis required; requires verification that time-dependent settlement falls within EPR envelope.










**Table 2.5-54—{Soil Conditions For The U.S. EPR Standard Plant}**

SI Units

Site Class		Shear Wave Velocity (mean) (m/s)	Location of GMRS	Geotechnical Analysis Requirements
Rock Site (Rock is at the ground surface at the site)	Hard and Firm Rock Site	$V_s > 1,070$	At top of Rock	Static and dynamic bearing capacity to be verified; no time-dependent settlement analysis required.
	Soft Rock Site	$730 < V_s < 1,070$	At top of Rock	Static and dynamic bearing capacity to be verified; no time-dependent settlement analysis required.
Thin Soil Site (Rock is generally within 40 to 60 feet of the ground surface and the EPR Nuclear Island is founded on rock)	Thin Soil Site over Hard or Firm Rock	At depth below Nuclear Island Basemat, $V_s > 3,500$	At Top of Outcropping Rock	Static and dynamic bearing capacity to be verified; no time-dependent settlement analysis required.
	Thin Soil Site over Soft Rock	At depth below Nuclear Island Basemat, $V_s < 1,070$	At Grade Elevation	Static and dynamic bearing capacity to be verified; no time-dependent settlement analysis required.
Soil Sites (Foundation underlain by < 200' of soil for Shallow and > 200' feet for Deep)	Shallow Soil and Deep Soil Sites	$305 < V_s < 1,070$ with soil below Nuclear Island Basemat of unlimited thickness	At free-field soil surface	Static and dynamic bearing capacity analysis required; requires verification that time-dependent settlement falls within EPR envelope.


**Table 2.5-55 {Foundation Elevations}**

English Units  
(Page 1 of 2)

Structure		Foundation		Shape	Base to Rock (ft) <sup>2</sup>	Base to OGS (ft) <sup>3</sup>	Fill Thickness		Contact Soil <sup>5</sup>	Area (ft <sup>2</sup> )	Contact Pressure (ksf) <sup>6</sup>		
		Depth (ft)	El. <sup>1</sup> (ft msl)				In-Situ	Eng <sup>4</sup>			Average Static	Maximum Static	Dynamic
Nuclear Island <sup>(7)</sup>	NI	36.0	638.0		12.6	-28.0	0.0	12.6	C-M	80,170	14.7	22.0	25.0
Nuclear Auxiliary Building	UKA	41.5	632.5		19.5	-27.5	0.0	19.5	EF	12,510	9.8	22.0	25.0
Rad Waste Building	UKS	36.0	638.0		41.0	-19.5	0.0	41.0	EF	16,880	4.3	22.0	25.0
EWSW Cooling Towers	URB1	22.0	652.0		3.0	-23.0	0.0	3.0	C	22,120	5.4	22.0	25.0
	URB2	22.0	652.0		4.5	-17.0	0.0	4.5	C	22,120	5.4	22.0	25.0
	URB3	22.0	652.0		58.5	-5.5	0.0	58.5	EF	22,120	5.4	22.0	25.0
	URB4	22.0	652.0		41.0	-5.5	0.0	41.0	EF	22,120	5.4	22.0	25.0
Emergency Power Building (UBP)	UBP12	5.0	669.0		27.0	-3.0	0.0	27.0	C	12,650	3.2	22.0	25.0
	UBP34	5.0	669.0		67.0	12.0	0.0	67.0	EF	12,650	3.2	22.0	25.0

**Table 2.5-55 {Foundation Elevations}**

English Units  
(Page 2 of 2)

Structure		Foundation		Shape	Base to Rock (ft) <sup>2</sup>	Base to OGS (ft) <sup>3</sup>	Fill Thickness		Contact Soil <sup>5</sup>	Area (ft <sup>2</sup> )	Contact Pressure (ksf) <sup>6</sup>		
		Depth (ft)	El. <sup>1</sup> (ft msl)				In-Situ	Eng <sup>4</sup>			Average Static	Maximum Static	Dynamic
Turbine Building	UMA	33.0	641.0		33.0	-27.0	0.0	33.0	EF	100,000	4.2	10.0	12.0

<sup>1</sup> Plant Grade El. [ ft msl ] 674

<sup>2</sup> Distance between rock and base of foundation

<sup>3</sup> Distance between foundation base original ground surface

<sup>4</sup> Engineered fill (Concrete for NI and soil for other facilities)

<sup>5</sup> Concrete (C), Mahantango Formation (M), Glacial Overburden (GO), Engineered Fill (EF)










<sup>6</sup> As prescribed by the U.S. EPR FSAR and Foundation Interface Document

<sup>7</sup> Depth given for containment building, other NI facilities have 41.3 ft




**Table 2.5-55—{Foundation Elevations}**

SI Units  
(Page 1 of 2)

Structure		Foundation		Shape	Base to Rock (m) <sup>2</sup>	Base to OGS (m) <sup>3</sup>	Fill Thickness		Contact Soil <sup>5</sup>	Area (m <sup>2</sup> )	Contact Pressure (kPa) <sup>6</sup>		
		Depth (m)	EI. <sup>1</sup> (m msl)				In-Situ	Eng <sup>4</sup>			Average Static	Maximum Static	Dynamic
Nuclear Island	NI	11.0	194.5		3.8	-8.5	0.0	3.8	C-M	7,452	700	1,050	1,200
Nuclear Auxiliary Building	UKA	12.7	192.8		5.9	-8.4	0.0	5.9	EF	1,163	470	1,050	1,200
Rad Waste Building	UKS	11.0	194.5		12.5	-5.9	0.0	12.5	EF	1,569	210	1,050	1,200
ESWS Cooling Towers	URB1	6.7	198.8		0.9	-7.0	0.0	0.9	C	2,056	260	1,050	1,200
	URB2	6.7	198.8		1.4	-5.2	0.0	1.4	C	2,056	260	1,050	1,200
	URB3	6.7	198.8		17.8	-1.7	0.0	17.8	EF	2,056	260	1,050	1,200
	URB4	6.7	198.8		12.5	-1.7	0.0	12.5	EF	2,056	260	1,050	1,200
Emergency Power Building( UBP )	UBP12	1.5	204.0		8.2	-0.9	0.0	8.2	C	1,176	150	1,050	1,200
	UBP34	1.5	204.0		20.4	3.7	0.0	20.4	EF	1,176	150	1,050	1,200

**Table 2.5-55—{Foundation Elevations}**

SI Units  
(Page 2 of 2)

Structure		Foundation		Shape	Base to Rock (m) <sup>2</sup>	Base to OGS (m) <sup>3</sup>	Fill Thickness		Contact Soil <sup>5</sup>	Area (m <sup>2</sup> )	Contact Pressure (kPa) <sup>6</sup>		
		Depth (m)	El. <sup>1</sup> (m msl)				In-Situ	Eng <sup>4</sup>			Average Static	Maximum Static	Dynamic
Turbine Building	UMA	10.1	195.4		10.1	-8.2	0.0	10.1	EF	9,295	200	480	570

<sup>1</sup> Plant Grade El. [ m msl ] 205.5

<sup>2</sup> Distance between rock and base of foundation

<sup>3</sup> Distance between foundation base original ground surface

<sup>4</sup> Engineered fill (Concrete for NI and soil for other facilities)












<sup>5</sup> Concrete (C), Mahantango Formation (M), Glacial Overburden (GO), Engineered Fill (EF)

<sup>6</sup> As prescribed by the U.S. EPR FSAR and Foundation Interface Document

**Table 2.5-56 {Earth Pressure Coefficients}**

Formation	$\phi$ (°)	$k_a$	$k_p$	$k_o$	$k_{AE}$	$k_{PE}$	Observations
Glacial Overburden	35	0.27	3.69	0.43	0.58	8.12	NA - Not Applicable $k_a$ - Active Earth Pressure Coefficient $k_p$ - Passive Earth Pressure Coefficient $k_o$ - At Rest Earth Pressure Coefficient
Mahantango Formation	NA	NA	NA	NA	NA	NA	
Granular Fill/Backfill	35	0.27	3.69	0.43	0.58	8.12	
Concrete Fill	NA	NA	NA	NA	NA	NA	












**Table 2.5-57 {Bearing Capacity (Failure Controlled)}**  
English Units

Structure		Foundation		Footprint Shape	Simplified		Contact Surface <sup>2</sup>	Bearing Capacity (ksf)	
		Depth (ft)	El. <sup>1</sup> (ft msl)		B (ft)	L (ft)		Ultimate	Allowable FS=3.0
Nuclear Island	NI	36.0	638.0		284	284	C-M	720.0	240.0
Nuclear Auxiliary Building	UKA	41.5	632.5		105	120	C	239.6	79.9
Radioactive Waste Building	UKS	36.0	638.0		130	130	C	245.3	81.8
ESWS Cooling Towers	URB1	22.0	652.0		124	180	C	720.0	240.0
	URB2	22.0	652.0		124	180	C	720.0	240.0
	URB3	22.0	652.0		124	180	EF	222.8	74.3
	URB4	22.0	652.0		124	180	EF	222.8	74.3
Emergency Power Generation Building	UBP1	5.0	669.0		90	140	C	720.0	240.0
	UBP2	5.0	669.0		90	140	EF	139.1	46.4
Turbine Building	UMA	33.0	641.0		300	330	EF	436.6	145.5
10' x 10' Footing on Fill (3' Deep)		3.0	671.0		10	10	EF	22.9	7.6

<sup>1</sup> Plant Grade El. (ft msl) 674

<sup>2</sup> Concrete (C), Mahantango Formation (M), Glacial Overburden (GO), Engineered Fill (EF)

**Table 2.5-57—{Bearing Capacity (Failure Controlled)}**  
SI Units

Structure		Foundation		Footprint Shape	Simplified		Contact Surface <sup>2</sup>	Bearing Capacity (kPa)	
		Depth (m)	El. <sup>1</sup> (m msl)		B (m)	L (m)		Ultimate	Allowable FS=3.0
Nuclear Island	NI	11.0	194.5		87	87	C-M	34,470	11,490
Nuclear Auxiliary Building	UKA	12.7	192.8		32	37	C	11,470	3,820
Radioactive Waste Building	UKS	11.0	194.5		40	40	C	11,750	3,920
ESWS Cooling Towers	URB1	6.7	198.8		38	55	C	34,470	11,490
	URB2	6.7	198.8		38	55	C	34,470	11,490
	URB3	6.7	198.8		38	55	EF	10,670	3,560
	URB4	6.7	198.8		38	55	EF	10,670	3,560
Emergency Power Generation Building	UBP1	1.5	204.0		27	43	C	34,470	11,490
	UBP2	1.5	204.0		27	43	EF	6,660	2,220
Turbine Building	UMA	10.1	195.4		91	101	EF	20,900	6,970
10' x 10' Footing on Fill (3' Deep)		0.9	204.6		3	3	EF	1,100	370

<sup>1</sup> Plant Grade El. [ m msl ] 205

<sup>2</sup> Concrete (C), Mahantango Formation (M), Glacial Overburden (GO), Engineered Fill (EF)

**Table 2.5-58 {Elastic Settlement Analysis By Simplified Approximations}**  
English Units

Structure		Foundation		Contact Surface <sup>2</sup>	Service Load (ksf)	Settlement Approximation (in)					
		Depth (ft)	El. <sup>1</sup> (ft msl)			Janbu		Perloff		Kay & Cav.	
						Center	Edge	Center	Edge	Center	Edge
Nuclear Island	NI	36.0	638.0	C-M	14.7	0.0	-	0.0	0.0	0.0	0.0
Nuclear Auxiliary Building	UKA	41.5	632.5	C	9.8	0.3	-	1.1	0.4	0.4	0.2
Radioactive Waste Building	UKS	36.0	638.0	C	4.3	0.3	-	0.9	0.4	0.4	0.2
ESWS Cooling Towers	URB1	22.0	652.0	C	5.4	0.0	-	0.0	0.0	0.0	0.0
	URB2	22.0	652.0	C	5.4	0.0	-	0.0	0.0	0.0	0.0
	URB3	22.0	652.0	EF	5.4	0.6	-	1.8	0.7	0.9	0.4
	URB4	22.0	652.0	EF	5.4	0.4	-	1.1	0.4	0.6	0.3
Emergency Power Generation	UBP1	5.0	669.0	C	3.2	0.0	-	0.0	0.0	0.0	0.0
	UBP2	5.0	669.0	EF	3.2	0.4	-	0.9	0.4	0.7	0.3
Turbine Building	UMA	33.0	641.0	EF	4.2	0.4	-	0.7	0.3	0.3	0.1
10' x 10' Footing on Fill (3' Deep)		3.0	671.0	EF	2.0	0.0	-	0.1	0.0	0.1	0.0

<sup>1</sup> Plant Grade El. (ft msl) 674.0

<sup>2</sup> Concrete (C), Mahantango Formation (M), Glacial Overburden (GO), Engineered Fill (EF)

**Table 2.5-58—{Elastic Settlement Analysis By Simplified Approximations}**

SI Units

Structure		Foundation		Contact Surface <sup>2</sup>	Service Load (kPa)	Settlement Approximation (in)					
		Depth (m)	El. <sup>1</sup> (m msl)			Janbu		Perloff		Kay & Cav.	
						Center	Edge	Center	Edge	Center	Edge
Nuclear Island	NI	11.0	194.5	C-M	700	0.0	-	0.1	0.0	0.1	0.0
Nuclear Auxiliary Building	UKA	12.7	192.8	C	470	0.7	-	2.9	1.1	1.1	0.5
Radioactive Waste Building	UKS	11.0	194.5	C	210	0.8	-	2.3	0.9	1.1	0.5
ESWS Cooling Towers	URB1	6.7	198.8	C	260	0.0	-	0.0	0.0	0.0	0.0
	URB2	6.7	198.8	C	260	0.0	-	0.0	0.0	0.0	0.0
	URB3	6.7	198.8	EF	260	1.5	-	4.6	1.9	2.4	1.0
	URB4	6.7	198.8	EF	260	1.0	-	2.8	1.1	1.4	0.6
Emergency Power Generation	UBP1	1.5	204.0	C	150	0.0	-	0.0	0.0	0.0	0.0
	UBP2	1.5	204.0	EF	150	1.1	-	2.4	1.1	1.8	0.7
Turbine Building	UMA	10.1	195.4	EF	200	1.0	-	1.7	0.7	0.7	0.4
10' x 10' Footing on Fill (3' Deep)		0.9	204.6	EF	100	0.1	-	0.2	0.1	0.2	0.1

<sup>1</sup> Plant Grade El. (m msl) 205.5

<sup>2</sup> Concrete (C), Mahantango Formation (M), Glacial Overburden (GO), Engineered Fill (EF)

**Table 2.5-59 {Detailed Elastic Settlement Analysis}**

English Units

Structure		Depth (ft)	El. <sup>1</sup> (ft msl)	Contact Surface	Service Load (ksf)	Settlement	
						Total (in)	Differential (in/50ft)
Nuclear Island <sup>3</sup>	NI	36.0	638.0	C-M	14.7	< .5	<.1

1 Plant Grade El. (ft msl) 674.0

2 Concrete (C), Mahantango Formation (M), Glacial Overburden (GO), Engineered Fill (EF)

3 Depth shown for UJA (Containment Building). Depth is 41.5 feet around containment

**Table 2.5-59—{Detailed Elastic Settlement Analysis}**

SI Units

Structure		Depth (m)	El. <sup>1</sup> (m msl)	Contact Surface	Service Load (kPa)	Settlement	
						Total (cm)	Differential (cm/100 m)
Nuclear Island <sup>3</sup>	NI	11.0	194.5	C-M	700.0	<0.2	<2

1 Plant Grade El. (m msl) 205.5

2 Concrete (C), Mahantango Formation (M), Glacial Overburden (GO), Engineered Fill (EF)

3 Depth shown for UJA (Containment Building). Depth is 12.7 m around containment



**Table 2.5-60 {Factor Of Safety Against Sliding}**

		Circular Failure		Wedge Failure		Observations
		Static	Dynamain	Static	Dynamic	
<b>Permanent Slopes (ESWEMS Retention Pond)</b>	Section 1	5.6	2.9	7.9	3.3	
	Section 2	6.9	2.4	9.2	4.1	
	Section 3	4.4	2.0	5.9	2.2	
	Section 4	5.7	2.1	7.0	2.9	
<b>Temporary Slopes</b>	Section 1	1.3			1.4	Dynamic case not applicable for temporary slopes.
	Section 2	2.5		2.5		
<b>North Slope</b>	Section 1			5.7	1.6	North slope analyzed to examine risk of landslide on-site. Only wedge failure is applicable.

**Table 2.5-61 {Smooth Uniform Hazard Response Spectra}**

<b>Frequency (Hz)</b>	<b>Smooth UHRS 1E-4</b>	<b>Smooth UHRS 1E-5</b>	<b>Smooth UHRS 1E-6</b>
0.100	0.004	0.011	0.025
0.127	0.006	0.017	0.039
0.161	0.008	0.026	0.057
0.204	0.012	0.036	0.080
0.259	0.016	0.050	0.111
0.329	0.022	0.068	0.151
0.418	0.029	0.090	0.200
0.500	0.036	0.110	0.244
0.672	0.042	0.124	0.272
0.853	0.045	0.126	0.276
1.000	0.045	0.118	0.261
1.374	0.058	0.152	0.347
1.750	0.068	0.180	0.425
2.212	0.077	0.207	0.507
2.500	0.081	0.220	0.556
3.562	0.103	0.294	0.767
4.520	0.121	0.358	0.957
5.000	0.131	0.407	1.108
7.500	0.162	0.534	1.484
10.000	0.189	0.647	1.825
11.721	0.210	0.728	2.104
14.874	0.238	0.837	2.477
18.874	0.262	0.931	2.815
25.000	0.280	1.004	3.100
30.392	0.272	0.981	3.019
38.566	0.257	0.930	2.854
48.939	0.229	0.832	2.544
62.102	0.182	0.665	2.026
78.805	0.129	0.475	1.443
100.000	0.097	0.361	1.091

**Table 2.5-62 Relation of the Tectonic History of the Site Region to the Associated Tectonic Features**

(Page 1 of 3)

<b>Tectonic Event</b>	<b>Principal Duration of Event</b>	<b>Brief Description of Conditions during Event near Site Region</b>	<b>Named Features and Text Sections for Detailed Descriptions</b>	
Grenville Orogeny	Meso-Proterozoic (Precambrian)	Collision of N America and S America created the supercontinent of Rodinia, and created a band of highly metamorphosed rocks along the southern and eastern edge of the N American craton.		
Late Precambrian Rifting	Late Precambrian to Middle Ordovician	During the breakup of Rodinia, and the opening of the Iapetus Ocean, rift basins developed along the continental margin and the normal faults along the margins of the rifts.	<b>New Madrid Seismic Zone</b> (and Reelfoot Rift) <b>Charlevoix-La Malbaie Seismic Zone</b> (related to the St. Lawrence Rift Valley?) <b>St. Lawrence Rift Valley</b> <b>Bristol Block Geopotential Trends</b> Transylvania Fault Zone	<b>2.5.1.4.4.1.1</b> <b>2.5.1.4.4.1.3</b>  <b>2.5.1.4.4.1.4</b> <b>2.5.1.4.4.3.4</b> 2.5.1.4.4.4.8
Taconic Orogeny	Middle Ordovician to Middle Silurian	Collision of the N American continent with island arc terranes, and accretion of volcanics and sea floor crust to the continent. Folding, thrust faulting, and metamorphism were common.	Mountain Run Fault (thrust fault) Pittsburgh-Washington Lineament Tyrone-Mt. Union Lineament Rome Trough Pleasant Valley-Huntington Valley Fault Plummers Island and Pleasant Grove Shear Zones Broadtop Synclinorium	2.5.1.4.4.2.16 2.5.1.4.4.3.3 2.5.1.4.4.3.3 2.5.1.4.4.4.6 2.5.1.4.4.4.7 2.5.1.4.4.4.9 2.5.1.4.4.4.17
Acadian Orogeny	Devonian to Middle Mississippian	Collision of N American continent with Europe and Africa. Mainly affected New England area with intrusives and metamorphism, but the site area with deposition of clastics.	<b>Charlevoix-La Malbaie Seismic Zone</b> (meteorite impact) Berwick Anticlinorium, and the Lackawanna and Catawissa-McCauley Mtn Synclinoria Light Street Fault Berwick Fault Plummers Island and Pleasant Grove Shear Zones	<b>2.5.1.4.4.1.3</b> 2.5.1.4.4.4.1 (4.2, & 4.16) 2.5.1.4.4.4.1.1 2.5.1.4.4.4.1.2 2.5.1.4.4.4.9
Alleghanian Orogeny	Middle Pennsylvanian to Permian	Convergence of N American and N Africa continents (and others) to form the supercontinent of Pangaea. The principal effects were folding, thrusting and metamorphism and the reactivation of some older faults.	Anthracite Region Yellow Breeches Fault Zone Transylvania Fault Zone (reactivated?) Plummers Island and Pleasant Grove Shear Zones Sweet Arrow Fault Chestnut Ridge Anticline	2.5.1.4.4.4.3 2.5.1.4.4.4.5 2.5.1.4.4.4.8 2.5.1.4.4.4.9 2.5.1.4.4.4.18 2.5.1.4.4.4.19

**Table 2.5-62 Relation of the Tectonic History of the Site Region to the Associated Tectonic Features**

(Page 2 of 3)

Tectonic Event	Principal Duration of Event	Brief Description of Conditions during Event near Site Region	Named Features and Text Sections for Detailed Descriptions	
Mesozoic Rifting	Triassic to Jurassic	The breakup of Pangaea led to the formation of new rift basins along the crustal margin, when the Atlantic Ocean was opened. Intrusion of diabase dikes, and accumulation of volcanics	<b>Charlevoix-La Malbaie Seismic Zone</b> (reactivated b renewed rifting?)	<b>2.5.1.4.4.1.3</b>
			<b>Ramapo Fault</b> (Border fault of the Newark Basin)	<b>2.5.1.4.4.2.2</b>
			Furlong-Flemington Fault Syst (border fault for rift basin)	2.5.1.4.4.2.3
			Kingston Fault (border fault for the Newark Basin)	2.5.1.4.4.2.4
			<b>Dobbs Ferry Fault</b> (Multi episodes of movement?)	<b>2.5.1.4.4.2.9</b>
			<b>Lancaster Seismic Zone</b> (Rift basin boundary?)	<b>2.5.1.4.4.2.10</b>
			East Border Fault	2.5.1.4.4.4.15
			<b>Everona Fault-Mountain Run Fault Zone</b> (Culpepper Basin)	<b>2.5.1.4.4.2.16</b>
			Transylvania Fault Zone (reactivated?)	2.5.1.4.4.4.8
			Gettysburg, Culpepper, & Newark Basins	2.5.1.4.4.4.10
			Hartford Basin and East Border Fault	2.5.1.4.4.4.11
			Connecticut Basin	2.5.1.4.4.4.12

**Table 2.5-62 Relation of the Tectonic History of the Site Region to the Associated Tectonic Features**

(Page 3 of 3)

Tectonic Event	Principal Duration of Event	Brief Description of Conditions during Event near Site Region	Named Features and Text Sections for Detailed Descriptions	
Cenozoic Passive Margin	Late Cretaceous to Present	Uplift and erosion, followed by the advance of at least three glaciers to the site vicinity. Site area is in a Passive Continental Margin.	<b>New Madrid Seismic Zone</b> (and Reelfoot Rift) (as a reactivated fault along a Precambrian rift?)	<b>2.5.1.4.4.1.1</b>
			<b>Charleston Seismic Zone</b>	<b>2.5.1.4.4.1.2</b>
			<b>Charlevoix-La Malbaie Seismic Zone</b>	<b>2.5.1.4.4.1.3</b>
			<b>St Lawrence Rift Valley</b>	<b>2.5.1.4.4.1.4</b>
			<b>Newbury Liquefaction Features</b> (Class A)	<b>2.5.1.4.4.2.1</b>
			<b>Ramapo Fault</b> (seismic activity indicates reactivation?)	<b>2.5.1.4.4.2.2</b>
			Furlong-Flemington Fault Syst (border fault for rift basin)	2.5.1.4.4.2.3
			Kingston Fault (border fault for the Newark Basin)	2.5.1.4.4.2.4
			<b>New York Bight Fault</b> (offshore)	<b>2.5.1.4.4.2.5</b>
			<b>Moshulu Fault</b>	<b>2.5.1.4.4.2.6</b>
			<b>New Castle County Faults</b> (Delaware)	<b>2.5.1.4.4.2.7</b>
			Upper Marlboro Faults (Landslides?)	2.5.1.4.4.2.8
			<b>Dobbs Ferry Fault</b> (Multi episodes of movement?)	<b>2.5.1.4.4.2.9</b>
			<b>Lancaster Seismic Zone</b> (Pennsylvania)	<b>2.5.1.4.4.2.10</b>
			Cacoosing Valley EQ Sequence (Quarry unloading)	2.5.1.4.4.2.11
			<b>Moodus Seismic Zone</b>	<b>2.5.1.4.4.2.12</b>
			<b>Clarendon-Linden Fault Zone</b>	<b>2.5.1.4.4.2.13</b>
			Offset Glaciated Surfaces	2.5.1.4.4.2.14
			Fall Lines of Weems	2.5.1.4.4.2.15
			<b>Everona Fault-Mountain Run Fault Zone</b>	<b>2.5.1.4.4.2.16</b>
			<b>Stafford Fault System</b>	<b>2.5.1.4.4.2.17</b>
			<b>New York-Alabama Lineament</b>	<b>2.5.1.4.4.3.1</b>
			<b>Hudson River Valley Trend</b>	<b>2.5.1.4.4.3.2</b>
			Pittsburgh-Washington Lineament	2.5.1.4.4.3.3
			Tyrone-Mt. Union Lineament	<b>2.5.1.4.4.3.3</b>
			<b>Bristol Block Geopotential Trends</b>	<b>2.5.1.4.4.3.4</b>
			<b>Reading Prong</b> (west of the Ramapo)	<b>2.5.1.4.4.3.5</b>
			<b>Peekskill-Stamford Seismic Boundary</b>	<b>2.5.1.4.4.3.6</b>
			Scranton Gravity High	2.5.1.4.4.4.4
			Brandywine Fault System	2.5.1.4.4.4.13
			Martic Fault	2.5.1.4.4.4.14
			East Border Fault	2.5.1.4.4.4.15
			Broadtop Synclinorium	<b>2.5.1.4.4.4.17</b>

NOTE: Features that are considered to be capable are shown in **BOLD FONT**



Figure 2.5-1 Site Region Topographic Map 200 Mile (322 km) Radius

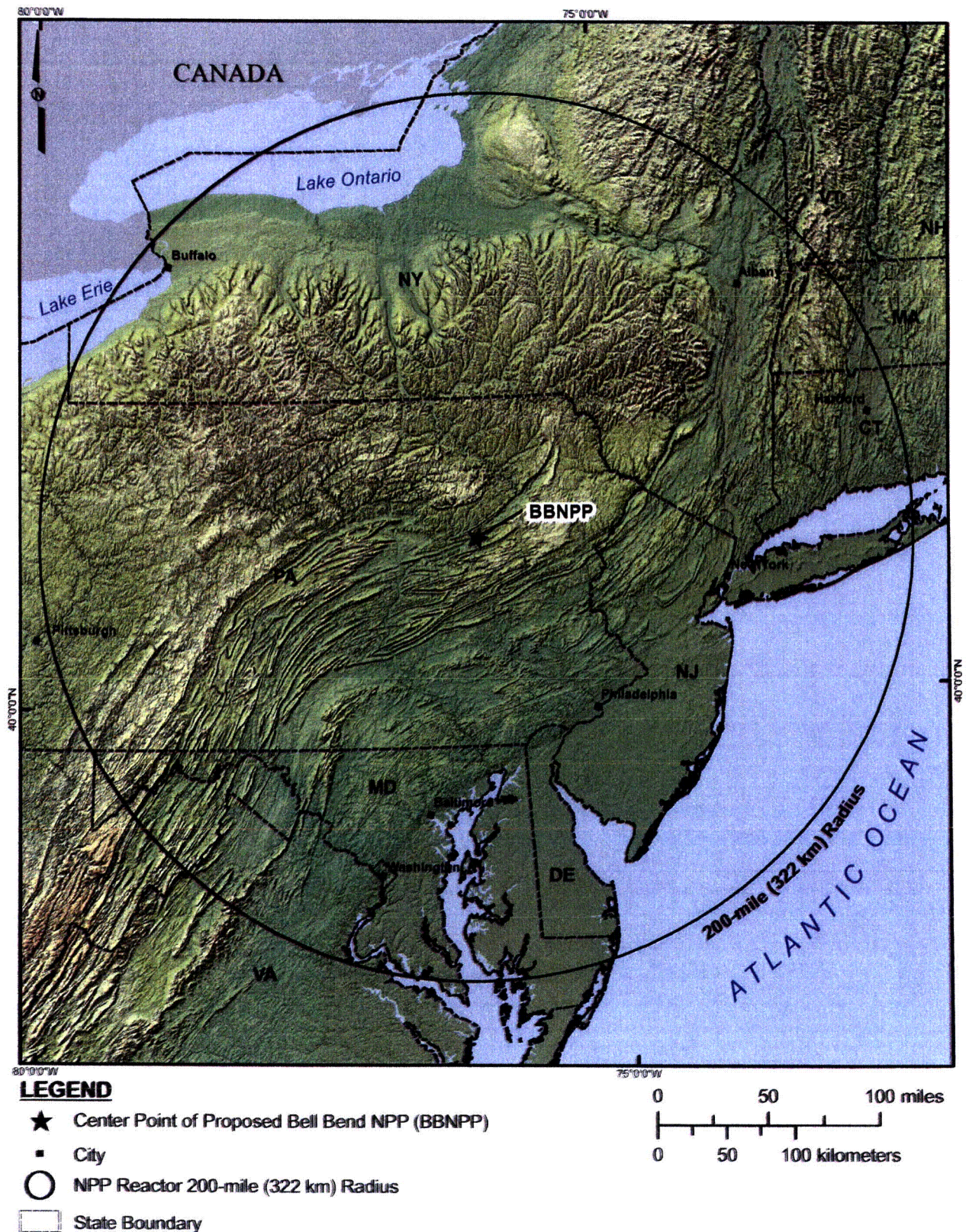
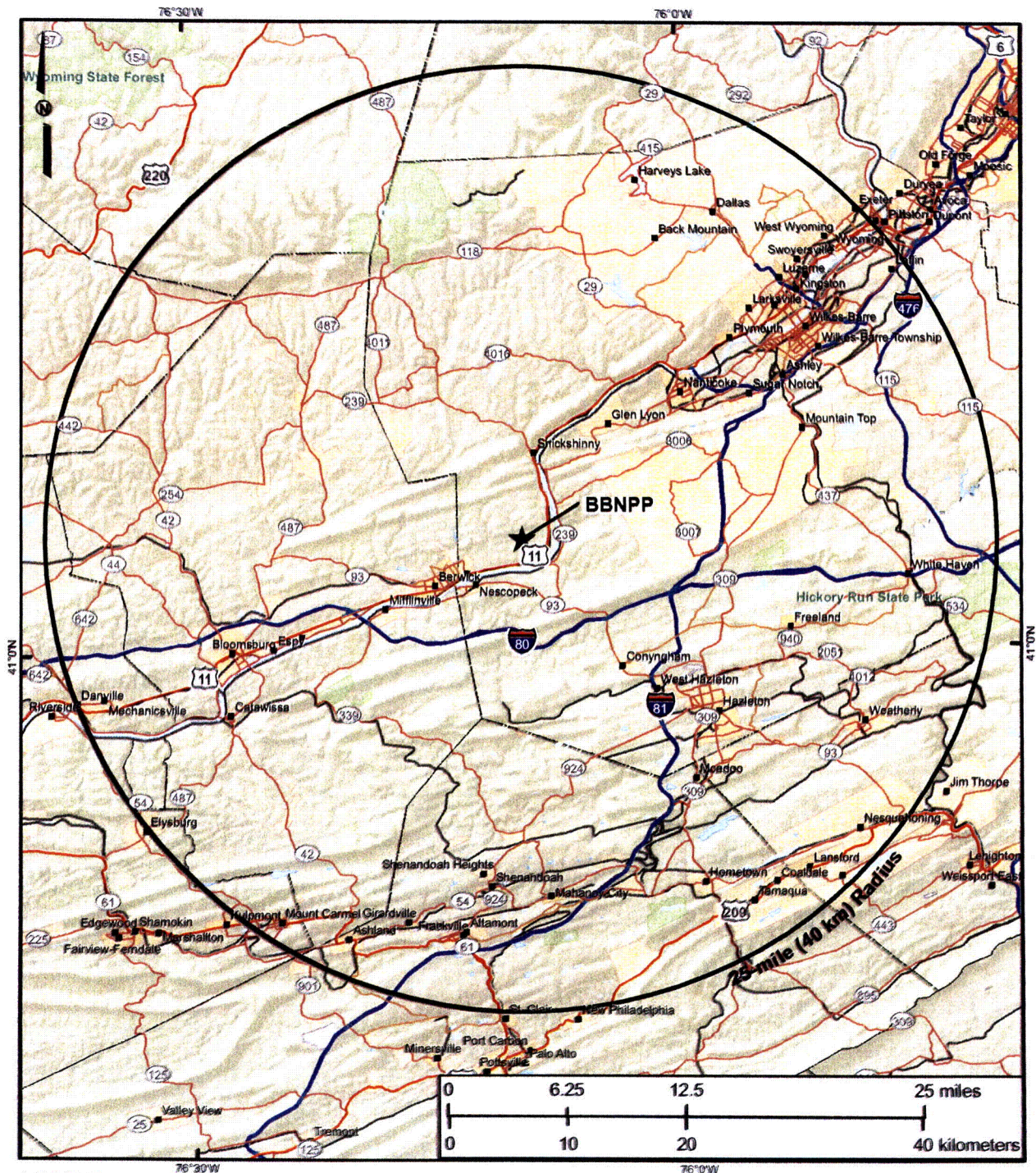




Figure 2.5-2 Site Vicinity Topographic Map 25 Mile (40 km) Radius



**LEGEND**

- ★ Center Point of Proposed Bell Bend NPP (BBNPP)
- NPP Reactor 25-mile (40 km) Radius

REFERENCE:  
• ESRI, 2007.



Figure 2.5-3 Site Topographic Map 5 Mile (8 km) Radius



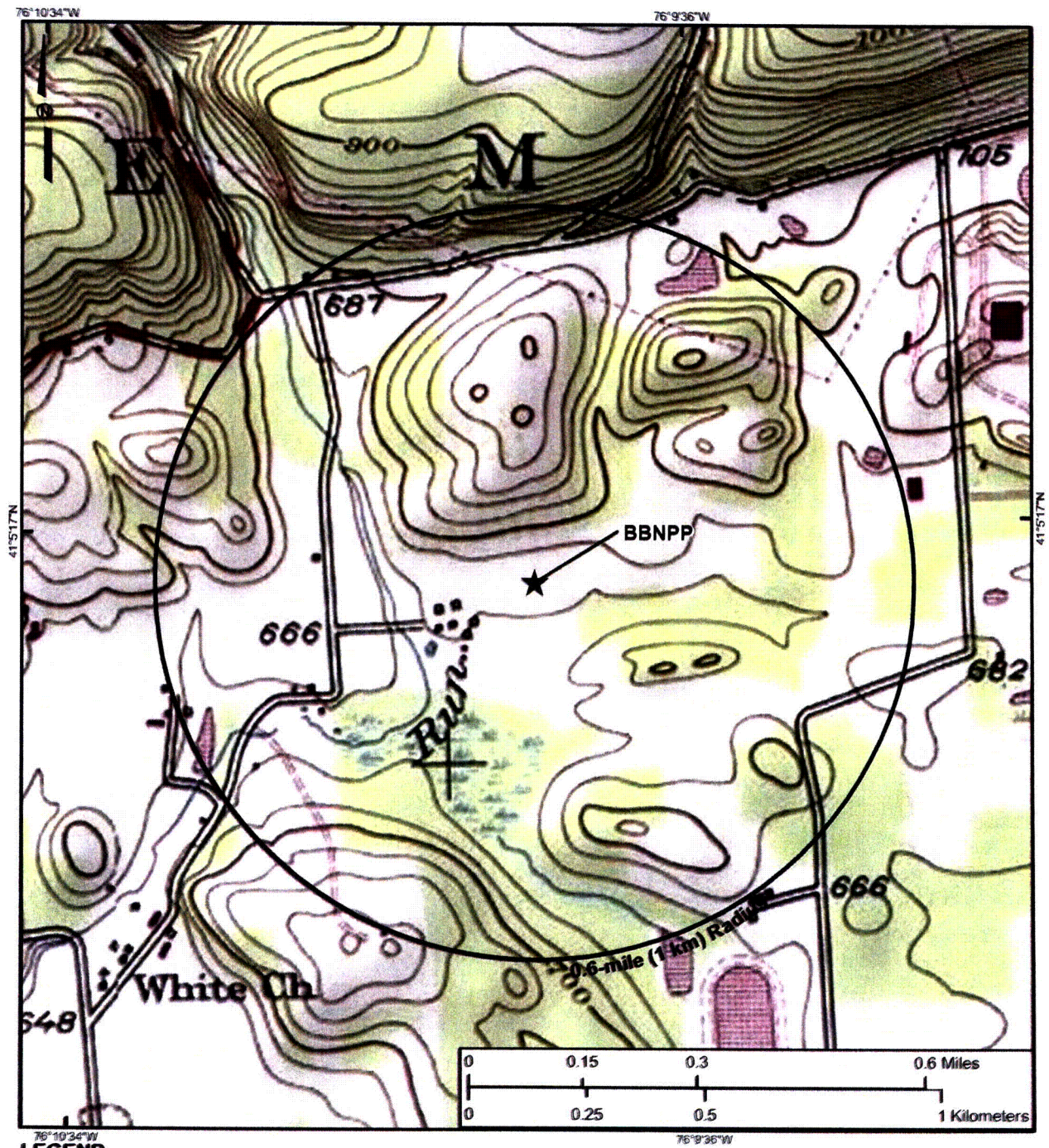
**LEGEND**

- ★ Center Point of Proposed Bell Bend NPP (BBNPP)
- NPP Reactor 5-mile (8 km) Radius

REFERENCE:  
• NGS, 2008.



Figure 2.5-4 Site Topographic Map 0.6 Mile (1 km) Radius



**LEGEND**

- ★ Center Point of Proposed Bell Bend NPP (BBNPP)
- NPP Reactor 0.6-mile (1 km) Radius

REFERENCE:  
- NGS, 2008.



Figure 2.5-5 Regional Geology Map 200 Mile (322 km) Radius

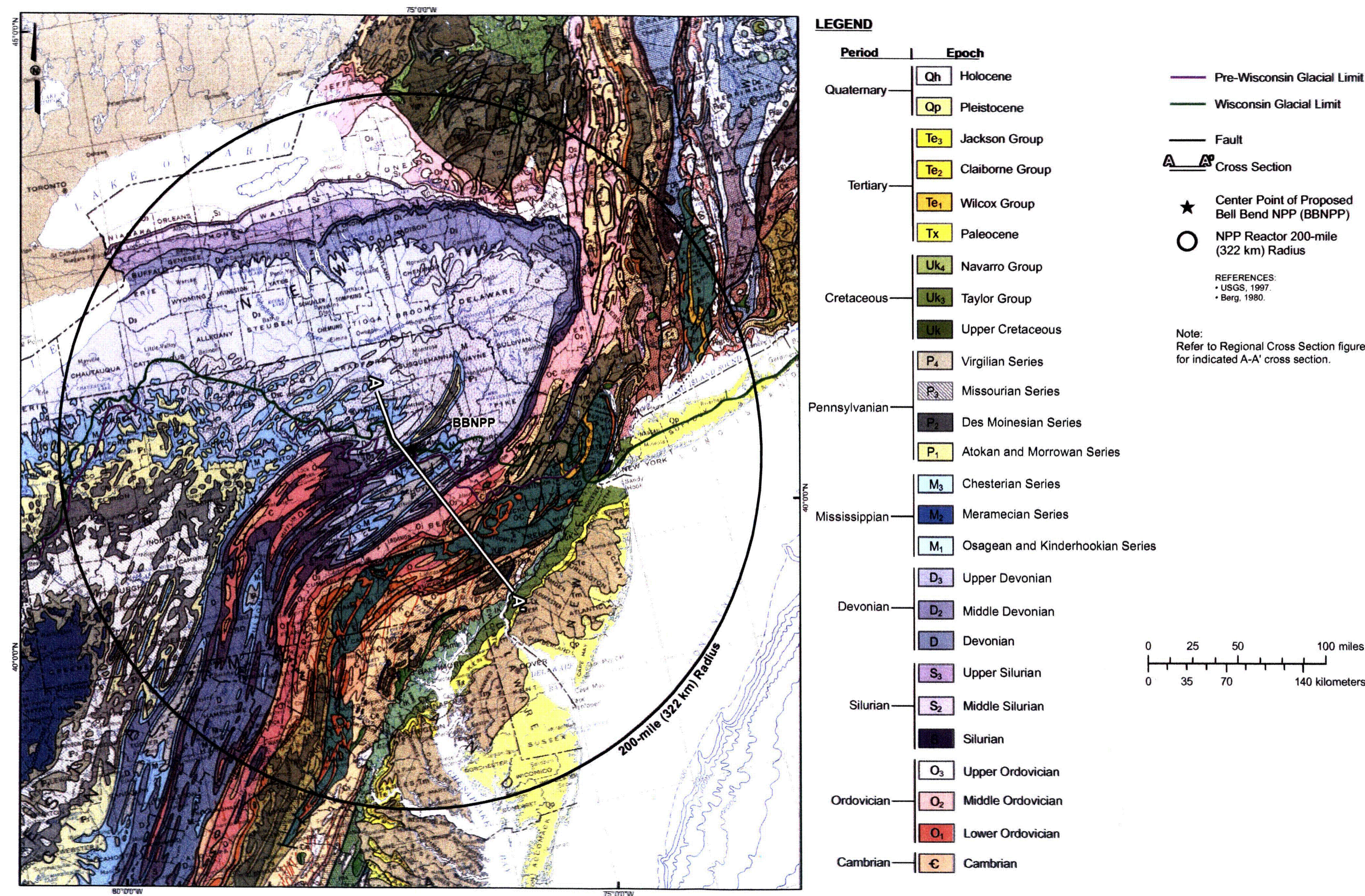
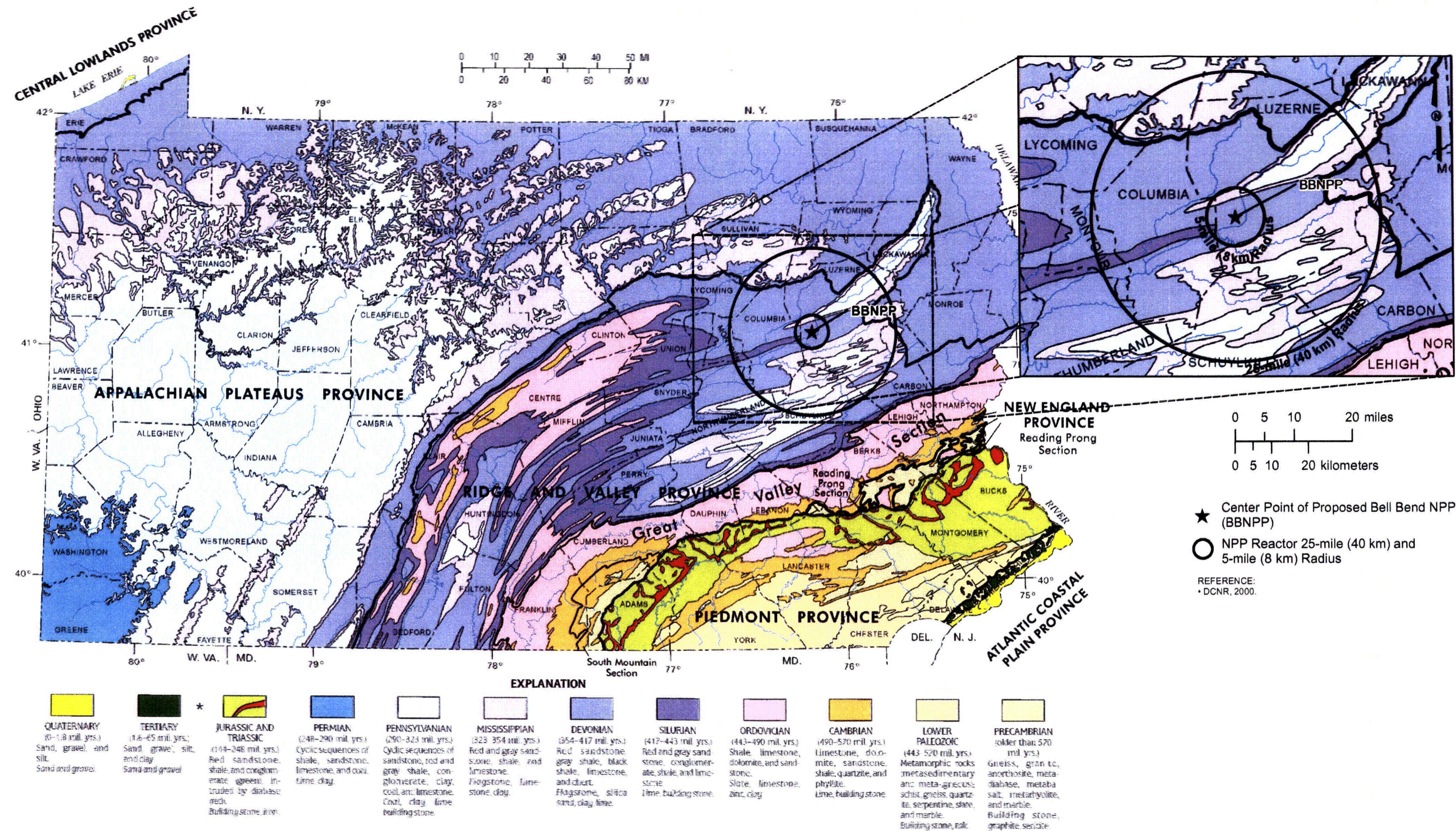




Figure 2.5-6 Physiographic Provinces of Pennsylvania (Geologic Map) with Sections 25 Mile (40 km) and 5 Mile (8 km) Radii



\*Cretaceous rocks, which are present in small areas of southern Montgomery County, cannot be shown at the scale of this map.



Figure 2.5-7 Physiographic Provinces (National) 200 Mile (322 km) Radius

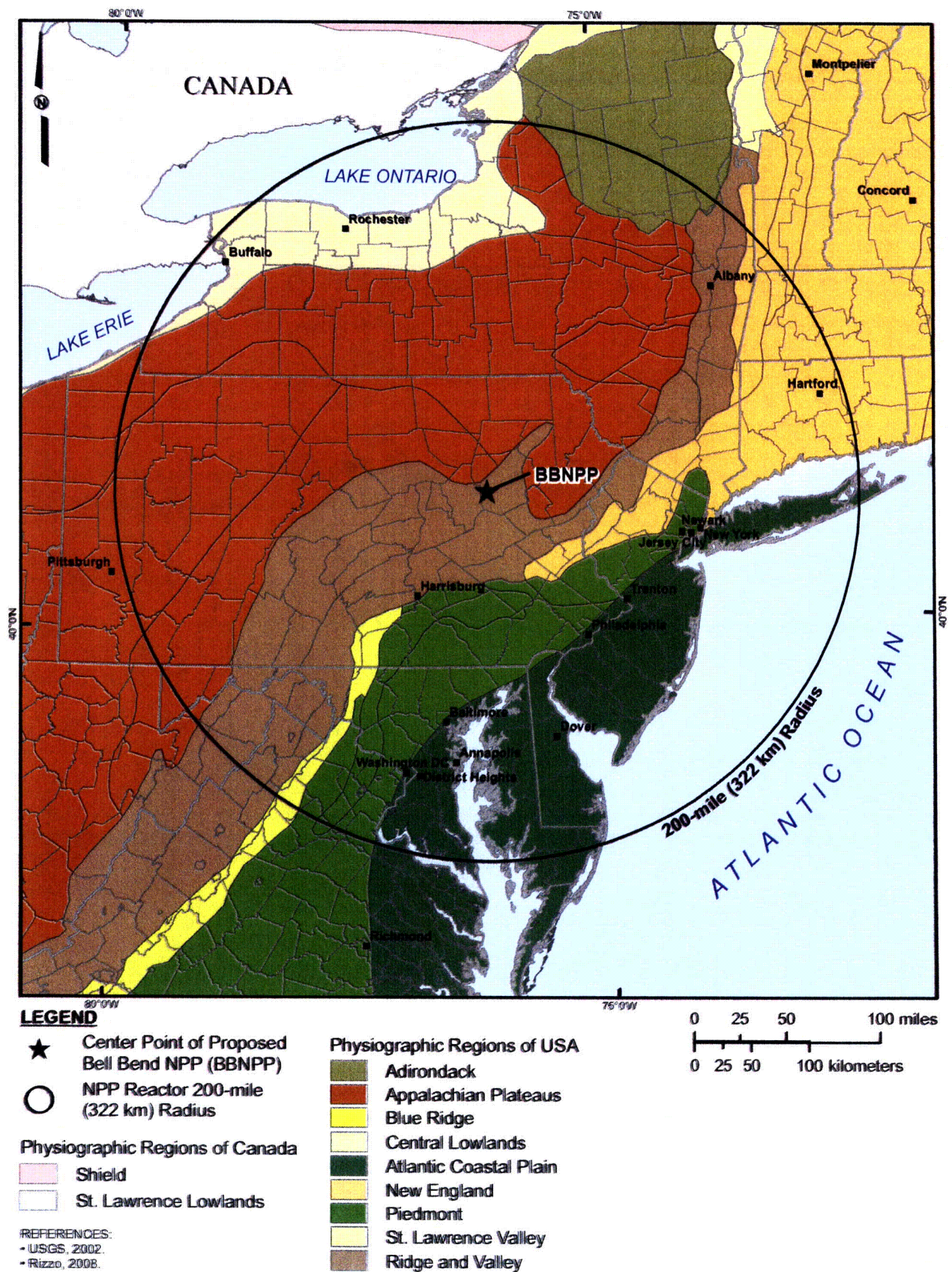
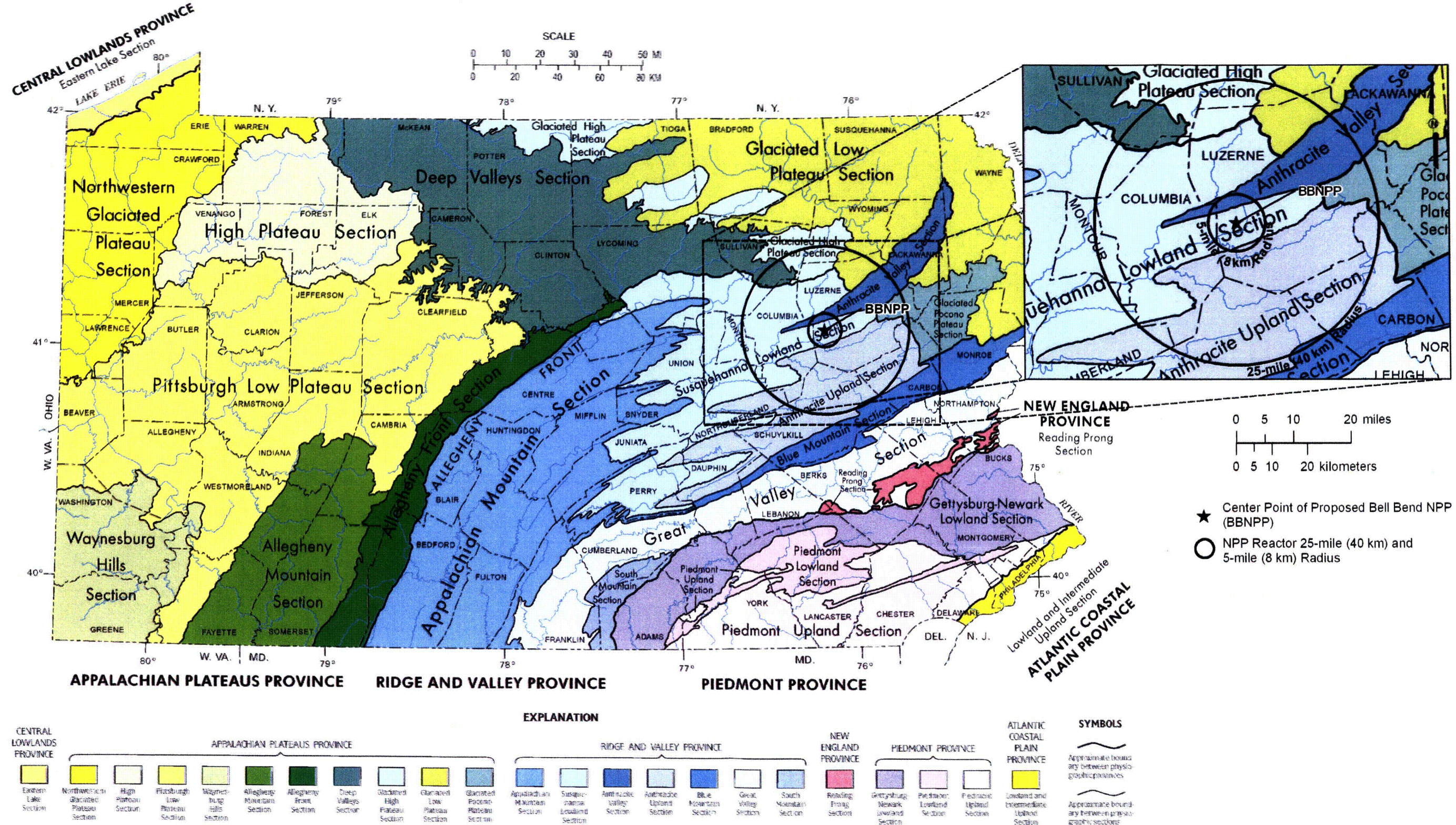




Figure 2.5-8 Physiographic Provinces of Pennsylvania with Sections 25 Mile (40 km) and 5 Mile (8 km) Radii



REFERENCE:  
Sevon, 2002.



Figure 2.5-9 Glacial Deposits of Pennsylvania 25 Mile (40 km) and 5 Mile (8 km) Radii

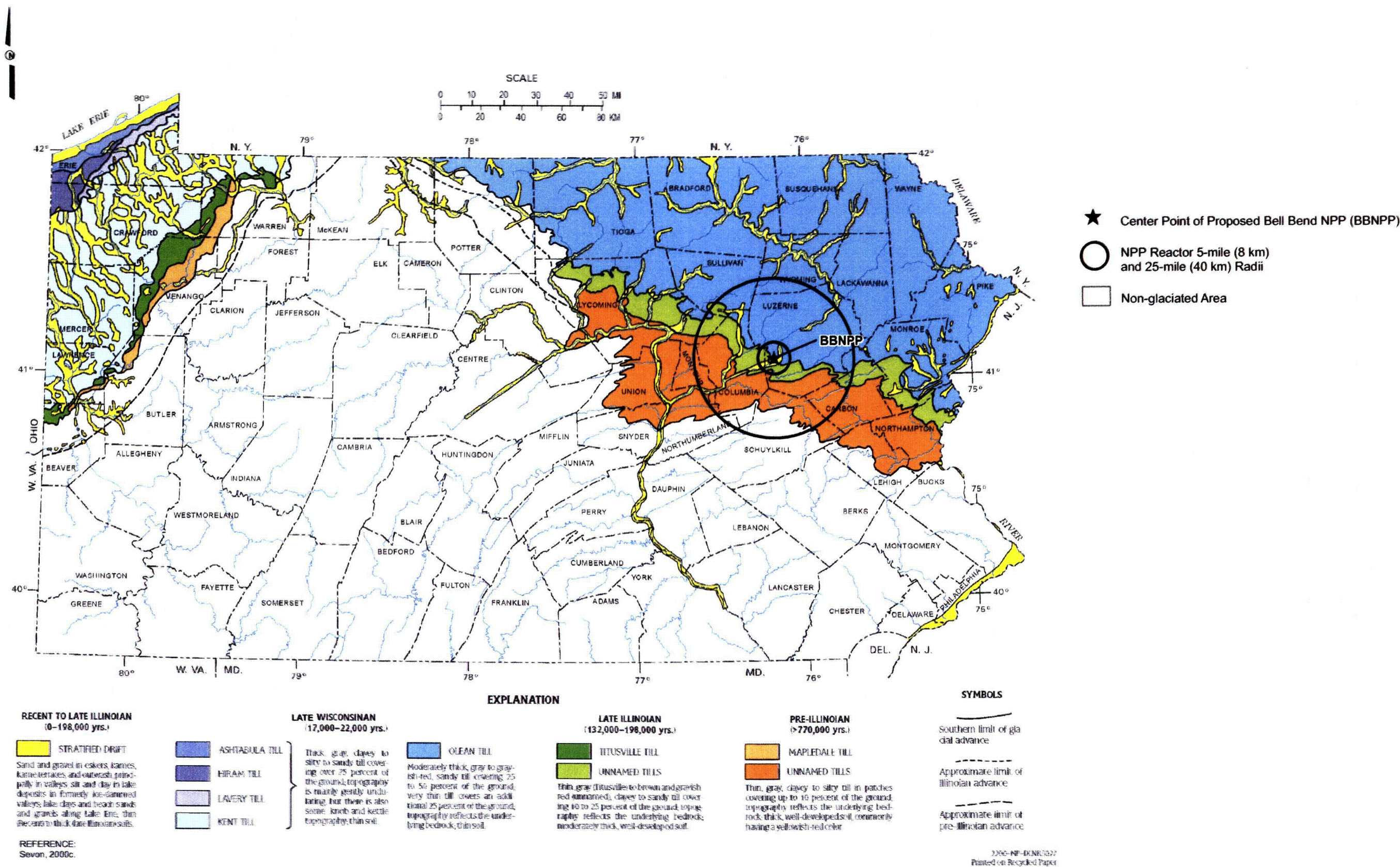
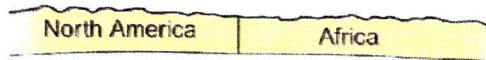


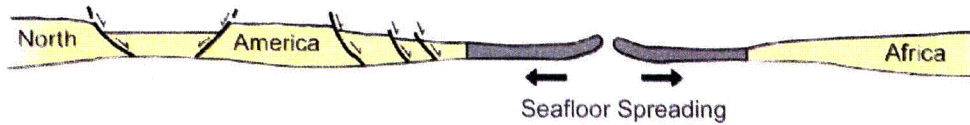


Figure 2.5-10 Evolution of the Appalachian Orogen

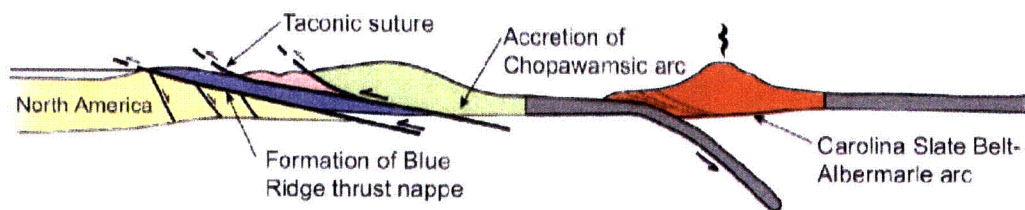
**Close of Grenville Orogeny (~1 Ga)**



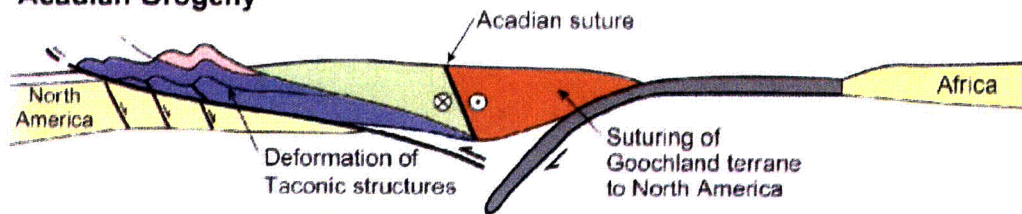
**Late Precambrian rifting; opening of Iapetus Ocean**



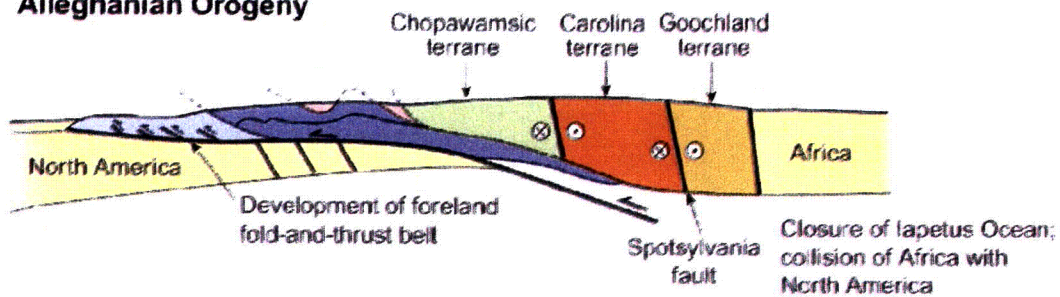
**Taconic Orogeny**



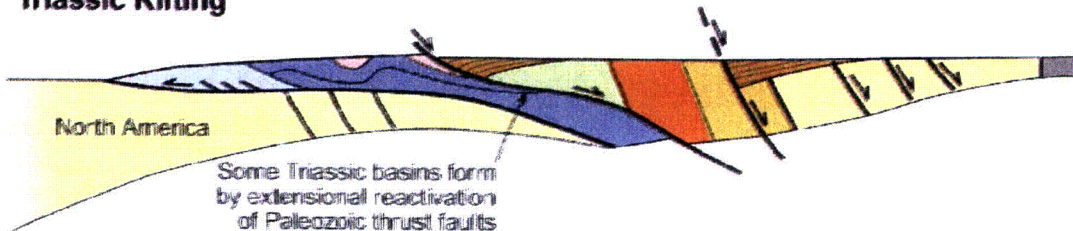
**Acadian Orogeny**



**Alleghanian Orogeny**



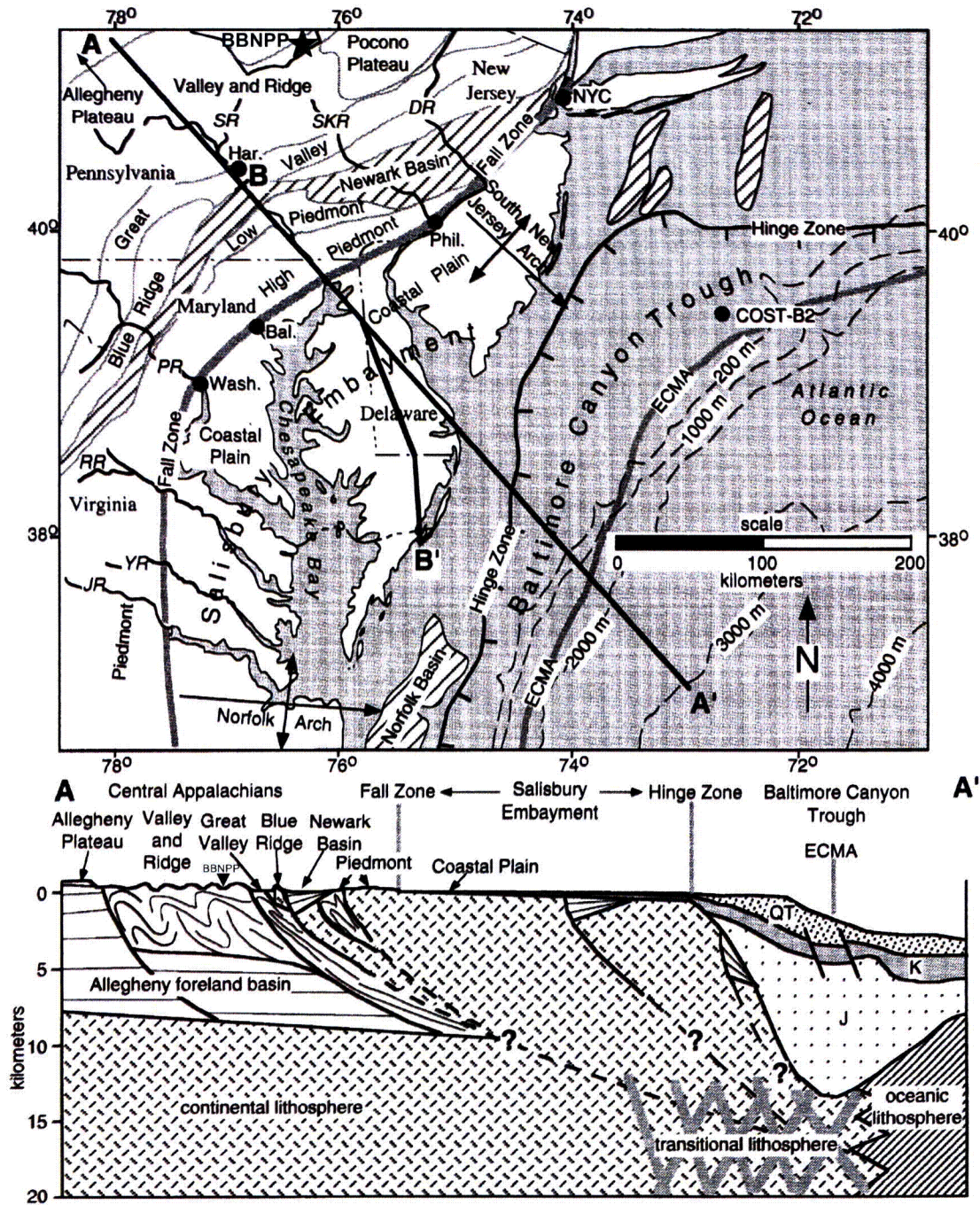
**Triassic Rifting**



Reference:  
• Hatcher, 1987.



Figure 2.5-10a Location Map and Cross Section of the Middle U.S. Atlantic Passive Margin



#### LEGEND

Syn-rift basins shown with the diagonal-line ruled pattern

DR, Delaware River; JR, James River; PT, Potomac River;  
RR, Rappahannock River; SKR, Schuylkill River;  
SR, Susquehanna River; YR, York River; Bal, Baltimore;  
Har, Harrisburg; NYC, New York City; Phil, Philadelphia;  
Wash, Washington

★ Bell Bend NPP

BBNPP Approximate Location of  
Bell Bend NPP

REFERENCE: Pazzaglia, 1994



**Figure 2.5-10b Maximum Horizontal Stress Directions in Eastern North America**

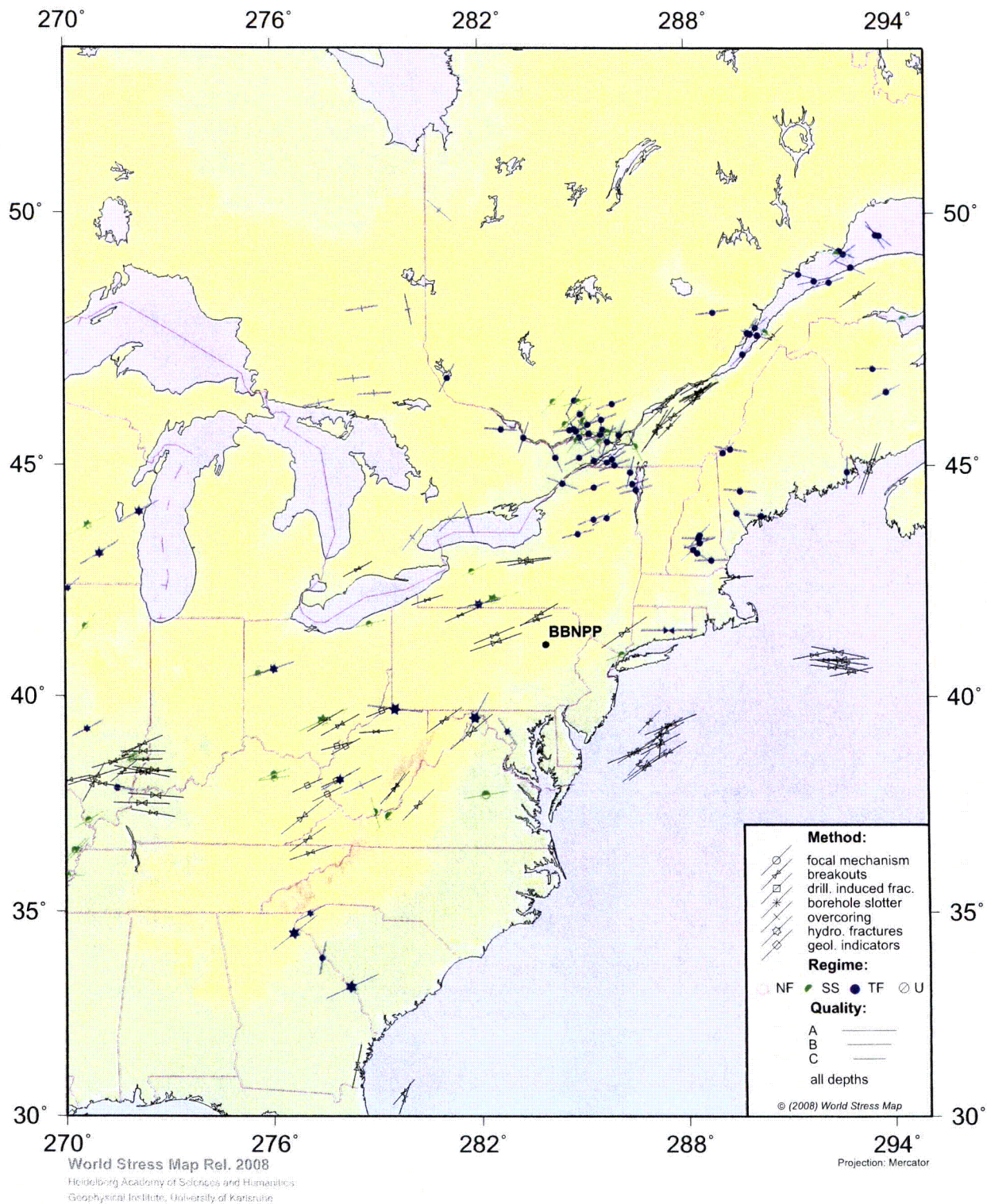




Figure 2.5-11 Cross Section Showing the Cambrian and Ordovician Rocks of the Appalachian Basin

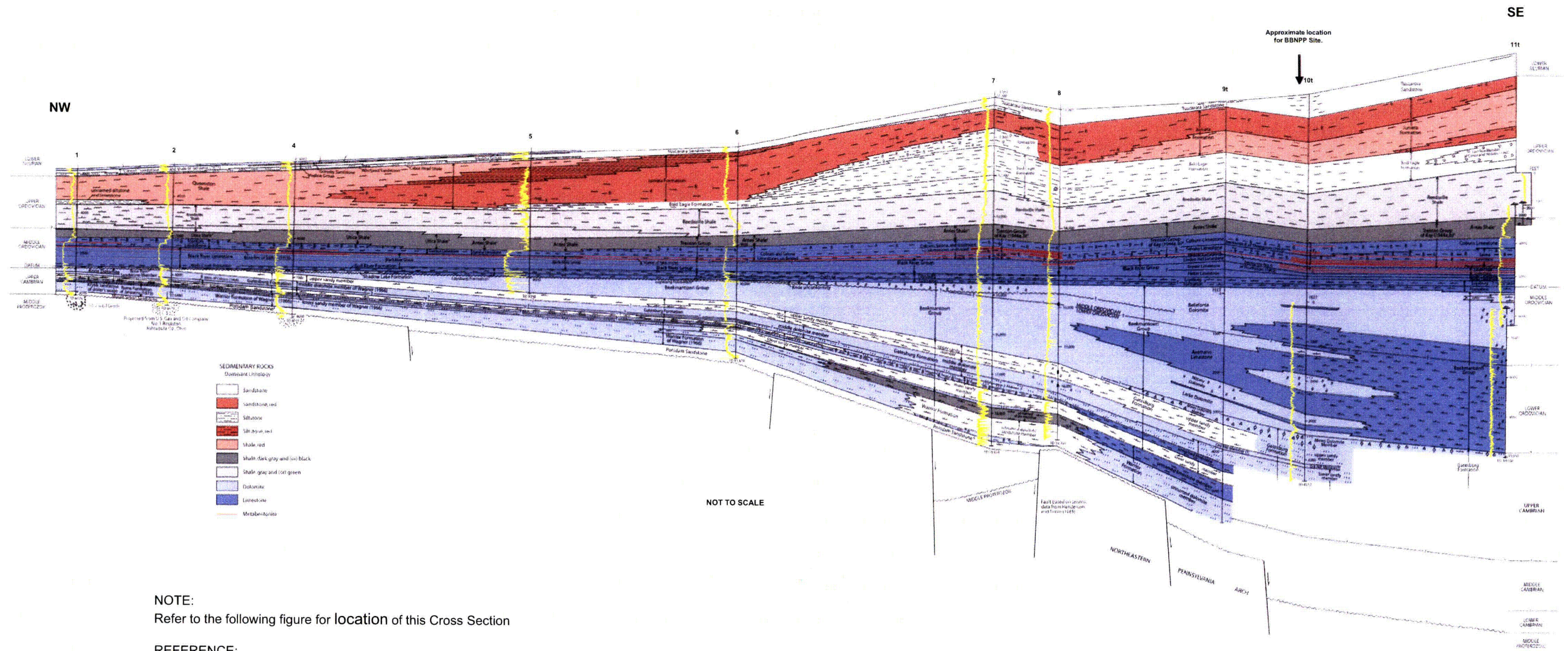
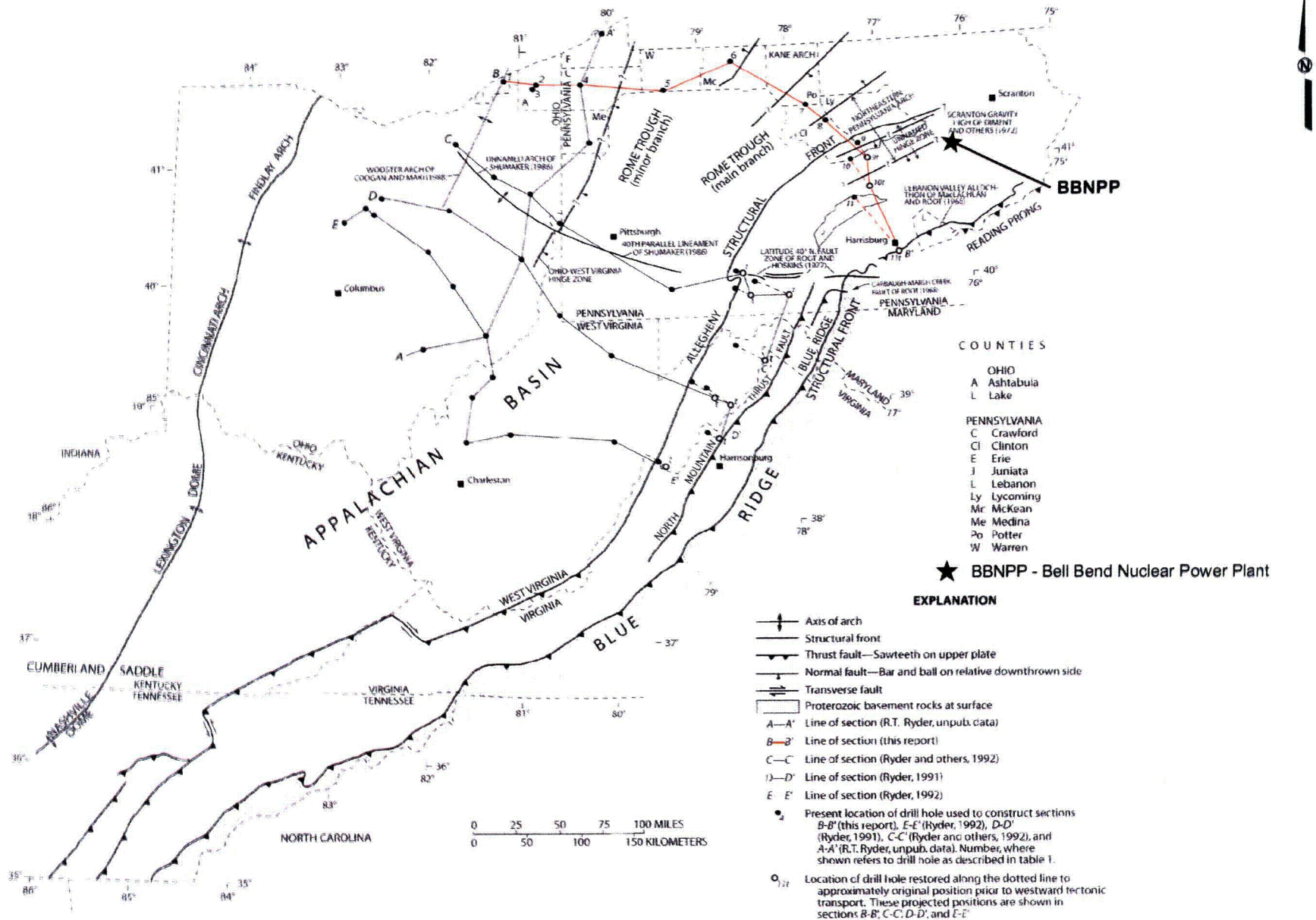




Figure 2.5-12 Map Showing the Location of the Appalachian Basin Cross Section



NOTE:  
Refer to previous figure.

REFERENCE:  
• Grangle, 2002.