2.4S.12 Groundwater

The following site-specific supplement addresses COL License Information Item 2.32.

This section describes the hydrogeologic conditions present at, and in the vicinity of, the STP 3 & 4 site. Regional and local groundwater resources that could be affected by the construction and operation of STP 3 & 4 are discussed. The regional and site-specific data on the physical and hydrologic characteristics of these groundwater resources are summarized in order to provide the basic data for an evaluation of impacts on the aquifers of the area.

The STP site covers an area of approximately 12,220 acres and is located on the coastal plain of southeastern Texas in Matagorda County. The power station lies approximately 10 mi north of Matagorda Bay. Nearby communities include Palacios, approximately 10 mi to the southwest and Bay City, approximately 12 mi to the northeast (Figure 2.4S.12-1). The closest major metropolitan center is Houston, approximately 90 mi to the northeast.

The 7000-acre Main Cooling Reservoir (MCR) is the predominant feature at the STP site, as shown in Figure 2.4S.12-2. The reservoir is fully enclosed with a compacted earth embankment, and it encompasses the majority of the southern and central portion of the site. The existing STP 1 & 2 facilities are located just outside of the MCR northern embankment. STP 3 & 4 is located further north of the embankment and to the northwest of STP 1 & 2.

The STP site, in general, has less than 15 ft of natural relief in the 4.5 mi distance from the northern to southern boundary. The northern section is at an elevation of approximately 30 ft above mean sea level (MSL). The southeastern section is at an elevation of approximately 15 ft above MSL. The Colorado River flows along the southeastern site boundary. There are also several unnamed drainages within the site boundaries, one of which feeds Kelly Lake.

Regional and local surface water features are described in Subsection 2.4S.1 and a geologic overview is presented in Subsection 2.5S.1.

2.4S.12.1 Description and Onsite Use

This section describes the regional and local groundwater aquifers and associated geologic formations, groundwater sources and sinks, and onsite use of groundwater.

2.4S.12.1.1 Regional Hydrogeologic Setting

The STP site islocated in Matagorda County and liesin the Gulf Coastal Plains physiographic province within the Coastal Prairies sub-province, which extends as a broad band parallel to the Texas Gulf Coast (Figure 2.4S.12-3). The Coastal Prairies sub-province is characterized by relatively flat topography with land elevation ranging from sea level along the coast to 300 ft above sea level along the western boundary. The geologic materials underlying the Coastal Prairies sub-province consist of deltaic sands and muds (Reference 2.4S.12-1).

The STP site is underlain by a thick wedge of southeasterly dipping sedimentary deposits of Holocene age through Oligocene age. The site overlies what has been referred to as the "Coastal Lowland Aquifer System" (Figure 2.4S.12-4). This aquifer system contains numerous local aquifers in a thick sequence of mostly unconsolidated Coastal Plain sediments of

alternating and interfingering beds of clay, silt, sand, and gravel. The sediments reach thicknesses of thousands of feet and contain groundwater that ranges from fresh to saline. Large amounts of groundwater are withdrawn from the aquifer system for municipal, industrial, and irrigation needs (Reference 2.4S.12-2).

The lithology of the aquifer system is generally sand, silt, and clay and reflects three depositional environments: continental (alluvial plain), transitional (delta, lagoon, and beach), and marine (continental shelf). The depositional basin thickens towards the Gulf of Mexico, resulting in a wedge-shaped configuration of hydrogeologic units. Numerous oscillations of ancient shorelines resulted in a complex, overlapping mixture of sand, silt, and clay (Reference 2.4S.12-2).

As part of the United States Geological Survey's (USGS) Regional Aquifer-System Analysis (RASA) program, the aquifer system was subdivided into five permeable zones and two confining units. The term "Gulf Coast Aquifer" is generally used in Texas to describe the composite of the sands, silts, and clays of the Coastal Lowland Aquifer System. Comparison of the USGS aquifer system nomenclature to that used in Texas is shown in Figure 2.4S.12-5. A cross-sectional representation is shown in Figure 2.4S.12-6 (Reference 2.4S.12-2).

The Texas nomenclature is used to describe the Gulf Coast Aquifer beneath the site. The hydrogeologic units commonly used to describe the aquifer system (from shallow to deep) are as follows (Figure 2.4S.12-5).

- -Chicot Aquifer
- -Evangeline Aquifer
- -Burkeville Confining Unit
- -Jasper Aquifer
- -Catahoula Confining Unit (restricted to where present in the Jasper Aquifer)
- -Vicksburg-Jackson Confining Unit

The base of the Gulf Coast Aquifer is identified as either its contact with the top of the Vicksburg-Jackson Confining Unit or the approximate depth where groundwater has a total dissolved solids concentration of more than 10,000 milligrams per liter (mg/L). The aquifer system is recharged by the infiltration of precipitation that falls on aquifer outcrop areas in the northern and western portion of the province. Discharge occurs by evapotranspiration, loss of water to streams and rivers as base flow, upward leakage to shallow aquifers in low lying coastal areas or in the Gulf of Mexico, and pumping.

In the shallow zones, the specific yield for sandy deposits generally ranges from between 10 percent and 30 percent. For the confined aquifer, the storage coefficient is estimated to range between 1 x 10^{-4} and 1 x 10^{-3} . The productivity of the aquifer system is directly related to the thickness of the sands in the aquifer system that contain freshwater. The aggregated sand thickness ranges from 0 ft at the up dip limit of the aquifer system to as much as 2000 ft in the east. Estimated values of transmissivity are reported to range from 5000 ft^2/day to nearly $35,000 \text{ ft}^2/\text{day}$ (Reference 2.4S.12-2).

2.4S.12.1.2 Regional Groundwater Aquifers

The STP site is located over the Gulf Coast Aquifer System as shown on Figure 2.4S.12-7 (Reference 2.4S.12-3). The Gulf Coast Aquifer has not been declared a Sole Source Aquifer (SSA) by the United States Environmental Protection Agency (EPA) (Reference 2.4S.12-4). A SSA is the sole or principal source of drinking water for an area that supplies 50 percent or more of drinking water with no reasonably available alternative source should the aquifer become contaminated. Figure 2.4S.12-8 shows the location of SSAs in EPA Region VI, which includes Texas. The nearest Texas SSA is the Edwards I and II Aquifer System, which is located approximately 150 mi northwest of STP. Based on a southeasterly groundwater flow beneath Matagorda County, toward the Gulf of Mexico, and the distances to the identified SSAs, STP 3 & 4 will not adversely impact the SSAs in EPA Region VI. The identified SSAs are beyond the boundaries of the local and regional hydrogeologic systems associated with the STP site.

The principal aquifer used in Matagorda County is the Chicot Aquifer, which extends to a depth of greater than 1000 ft in the vicinity of the STP site, as shown on Figure 2.4S.12-9. The Chicot Aquifer is comprised of Holocene alluvium in river valleys and the Pleistocene age Beaumont, Montgomery, and Bentley Formations, and the Willis Sand (Reference 2.4S.12-5). Groundwater flow is, in general, southeasterly from the recharge areas north and west of the county to the Gulf of Mexico. Numerous river systems and creeks flow south and southeasterly through Matagorda County. River channel incisions can act as localized areas of recharge and discharge to the underlying aquifer system, resulting in localized hydraulic sources and sinks.

The Chicot Aquifer geologic units used for groundwater supply in the STP site area are the Beaumont Formation and the Holocene alluvium in the Colorado River floodplain. The following sections describe the pertinent details of these units.

2.4S.12.1.2.1 Beaumont Formation

The Beaumont Formation consists of fine-grained mixtures of sand, silt, and clay deposited in alluvial and deltaic environments. In the upper portion of the Beaumont Formation, sands occur as sinuous bodies, representing laterally discontinuous channel deposits, while the clays and silts tend to be more laterally continuous, representing their deposition as natural levees and flood deposits. The deeper portion of the unit, the Deep Aquifer, is greater than about 250 ft below ground surface in the vicinity of the site and has thicker and more continuous sands. This portion of the Beaumont Formation is the primary groundwater production zone for most of Matagorda County. Well yields in this interval are typically between 500 gallons per minute (gpm) and 1500 gpm with yields of up to 3500 gpm reported (Reference 2.4S.12-6). Groundwater occurs in this zone under confined conditions.

2.4S.12.1.2.2 Holocene Alluvium

Holocene alluvium of the Colorado River floodplain occurs in a relatively narrow band surrounding the river. The alluvial deposits are typically coarser-grained than the materials found in the Beaumont Formation. The alluvium consists of silt, clay, fine- to coarse-grained sand, and gravel, along with wood debris and logs (Reference 2.4S.12-6). In the immediate site area, the alluvium is too thin to be a significant source of groundwater. Since the alluvial materials are deposited in a channel incised into the Beaumont Formation, it is likely that the alluvium is in contact with the shallow aquifer units in the Beaumont Formation.

2.4S.12.1.3 Local Hydrogeology

The local hydrogeologic system is identified in the STP site area and it includes areas of groundwater - surface water interactions within a few miles of the site. The Beaumont Formation within the Chicot Aquifer (and to a lesser, extent, the Holocene alluvium associated with the Colorado River floodplain) is the principal water-bearing unit used for groundwater supply in the vicinity of STP. Within this area, the Chicot Aquifer is divided into two aquifer units, the Shallow Aquifer and the Deep Aquifer. The base of the Shallow Aquifer is 90 ft to 150 ft deep in the site area. The Shallow Aquifer has limited production capability and is used for livestock watering and occasional domestic use. Potentiometric heads are generally within 15 ft of ground surface (Reference 2.4S.12-6). The Deep Aquifer is the primary groundwater production zone and lies below depths of 250 ft to 300 ft. An overlying zone of predominately clay materials, usually greater than 150 ft thick, separates the Shallow and Deep Aquifers.

Recharge to the Shallow Aquifer is considered to be a few miles north of the site. Discharge is to the Colorado River alluvial material east of the site. Recharge to the Deep Aquifer is further north in Wharton County where the aquifer outcrops. Discharge from the Deep Aquifer is to Matagorda Bay, groundwater production wells, and the Colorado River estuary, approximately 5 mi southeast of the site. Shallow Aquifer groundwater quality is generally inferior to that of the Deep Aquifer (Reference 2.4S.12-6).

The Shallow Aquifer has been subdivided into upper and lower zones over the site area. Both zones respond to pumping as confined or semi-confined aquifers with somewhat different potentiometric heads. The Upper Shallow Aquifer is comprised of interbedded sand layers to depths of approximately 50 ft below ground surface. The Lower Shallow Aquifer consists of the sand layers between depths of approximately 50 ft to 150 ft below ground surface.

Aquifer pumping tests performed at the site in support of STP 1 $\&$ 2 indicate well yields from 10 gpm to 300 gpm in the Shallow Aquifer. These tests also indicate a variable degree of hydraulic connection between the Upper and Lower Shallow Aquifer zones (Reference 2.4S.12-7). Analysis of the aquifer pumping tests indicates that groundwater occurs under confined conditions at the four test sites. A pumping test conducted at STP Production Well 5 confirmed confined conditions in the Deep Aquifer (Reference 2.4S.12-8).

2.4S.12.1.4 Site Specific Hydrogeology

A geotechnical and hydrogeological investigation was performed to provide information on the STP 3 & 4 site to depths of 600 ft below ground surface. Subsurface information was collected from over 150 geotechnical borings and cone penetrometer tests (CPTs). A detailed description of the geotechnical subsurface investigation, including the locations of these borings and CPTs, boring logs, and soil testing data is provided in Subsection 2.5S.4.

Twenty-eight (28) groundwater observation wells were installed in the vicinity of the STP 3 $\&$ 4 site. The wells were completed in the Upper and Lower Shallow Aquifer. The wells were located to a) supplement the existing STP piezometer network in order to provide an adequate distribution for determining groundwater flow directions and b) provide additional information on the hydraulic gradients beneath the site. Well pairs were installed at selected locations to determine vertical gradients. Figure 2.4S.12-10 shows the locations of observations wells and piezometers at the STP site. Table 2.4S.12-1 presents the installation information for the newly installed observation wells. Field hydraulic conductivity tests (slug tests) were conducted in each observation well. Monthly water level measurements from these groundwater observation wells began in December 2006.

The subsurface data collected in late 2006 and early 2007 as part of the STP 3 $\&$ 4 site subsurface investigation confirmed the aquifer conditions described for STP $1 \& 2$. The top of the uppermost sand layer within the Upper Shallow Aquifer is encountered at a depth of about 15 ft to 30 ft below ground surface at STP 3 & 4. The groundwater level is about 5 ft to 10 ft below ground surface. The unit is comprised of sand and silty sand, approximately 15 ft to 20 ft thick. Multiple sandy units that are separated by silts and clays define the Lower Shallow Aquifer. The groundwater level in these sand intervals is about 10 ft to 15 ft below ground surface beneath the STP $3 \& 4$ facility area.

2.4S.12.1.5 Groundwater Sources and Sinks

The natural regional flow pattern in the Beaumont Formation is from recharge areas, where the sand layers outcrop at the surface, to discharge areas, which are either at the Gulf of Mexico or the Colorado River Valley alluvium. The outcrop areas for the Beaumont Formation sands are in northern Matagorda County (Shallow Aquifer) and Wharton County (Deep Aquifer), to the north of Matagorda County. In the outcrop areas, precipitation falling on the ground surface can infiltrate directly into the sands and recharge the aquifer. Superimposed on this simplistic flow pattern is the influence of heavy pumping within the aquifer. Concentrated pumping areas can alter or reverse the regional flow pattern. Further discussion of regional groundwater use and flow patterns is presented in Subsection 2.4S.12.2.

The Holocene alluvium receives recharge from infiltration of precipitation and groundwater flow from the Shallow Aquifer in the Beaumont Formation. In the site area, flow paths in the alluvium are short due to the limited surface area. Discharge from the Holocene alluvium contributes to the base flow of the Colorado River. The Colorado River is dammed to the south of Bay City to supply irrigation water canals. During certain times of the year the only sources of water to the Colorado River below the dam are irrigation tail water releases and base flow created by seepage from the Holocene alluvium. Because there are no flow-gaging stations downstream of the dam, the amount of base flow contributed by seepage is not known (Reference 2.4S.12-6).

The MCR is unlined and may act as a local recharge source to the Shallow Aquifer at the STP site. The historical, normal maximum operating level of the 7000-acre MCR is at an elevation of 49 ft above mean sea level, imposing a head of up to 20 ft above ground surface. The capacity of the reservoir at this elevation is 202,700 acre-ft. The reservoir embankment dike is designed to lower the hydraulic gradient across the embankment to the extent that the potentiometric levels of the soil layers in the plant area stay below the ground surface. This is accomplished through the use of low permeability clay (compacted fill), relief wells, and sand drainage blankets. Discharge to the environment from the MCR occurs from seepage through the reservoir floor to the groundwater. Groundwater flow from the reservoir is intercepted, in

part, by the relief well system around the perimeter of the MCR, which is collected in toe and drainage ditches around the periphery of the reservoir embankment and then discharges to surface water features at various locations. Seepage discharge from the reservoir is composed of two parts: (a) seepage that is collected and discharged through about 700 relief wells that have been installed in the embankment around the reservoir to relieve excess hydrostatic pressure, and (b) seepage through the Upper Shallow Aquifer that bypasses the relief wells and continues down gradient. During the design stage, total seepage of the MCR was estimated to be 3530 gpm or approximately 5700 acre-ft/yr. Of this value, approximately 68 percent or 3850 acre-ft/yr would be discharged through the relief wells (Reference 2.4S.12-9).

2.4S.12.1.6 Plant Groundwater Use

Groundwater is currently used on the site to support STP $1 \& 2$ plant operations. The water is pumped from the Deep Aquifer using five production wells (Production Wells 5 through 8 and the Nuclear Training Facility [NTF] well). The production well depths are between 600 and 700 ft below ground surface with well capacities between 200 and 500 gpm as shown on Table 2.4S.12-2.

Figure 2.4S.12-10 shows the location of the existing site production wells. No sustained pumping is permitted within 4000 ft of the STP 1 $\&$ 2 plant area in order to minimize the potential for subsidence resulting from lowering of the Deep Aquifer zone potentiometric head. The exception is the NTF well, which was installed to provide water to the Nuclear Training Facility. (The NTF well only provides fire protection water to the NTF. Potable water for the NTF is supplied by Production Well 8.)

Groundwater use from these wells includes a makeup water source for the Essential Cooling Pond (ECP), makeup of demineralized water, the potable and sanitary water system, and the plant fire protection system (Reference 2.4S.12-9). Table 2.4S.12-3 presents the combined monthly groundwater withdrawals from the five production wells between 1995 and 2006. The table indicates that an annual groundwater usage of between 1200 and 1300 acre-ft is typical.

Groundwater is projected to be the main source of water for $STP 3 \& 4$ plant operations. Operation of STP 3 $\&$ 4 is predicted to require a typical groundwater consumption of 1077 gpm or 1738 acre-ft per year. The peak groundwater consumption (i.e., plant outage) for STP 3 $\&$ 4 is expected to be as great as 3935 gpm. The projected combined STP plant typical groundwater consumption for STP 1 $\&$ 2 and STP 3 $\&$ 4 is expected to be between 2938 acreft and 3038 acre-ft per year. The impacts to the local groundwater aquifer system are discussed in Subsection 2.4S.12.3.3.

The groundwater supply wells to be used for STP $3 \& 4$ are not a safety-related water source because the Ultimate Heat Sink (UHS) has a 30-day supply of water, which is sufficient to allow plant shutdown without additional water supply.

2.4S.12.2 Groundwater Sources

This section describes historical and projected groundwater use, groundwater flow directions, groundwater hydraulic gradients, temporal groundwater trends, aquifer properties, and hydrogeochemical characteristics. STP site groundwater information is based on groundwater observation wells installed at the site, as shown on Figure 2.4S.12-10.

2.4S.12.2.1 Historical and Projected Groundwater Use

Groundwater pumpage in the Gulf Coast Aquifer system was relatively small and constant from 1900 until the late 1930s. Pumping rates increased sharply between 1940 and 1960 and then increased relatively slowly through the mid 1980s. By the mid 1980s withdrawals were primarily from the east and central area of the aquifer system. This included the Houston area but some of the greatest pumpage was associated with rice irrigation centered in Jackson, Wharton, and portions of adjacent counties including Matagorda. The highest water demand was from the upper portion of the Deep Aquifer (Reference 2.4S.12-2).

Problems associated with groundwater pumpage, such as land subsidence, saltwater encroachment, stream base-flow depletion, and larger pumping lifts have caused pumpage to be curtailed in some areas. By the mid 1980s, the Texas Water Development Board (TWDB) had made projections of groundwater use to 2030. For the 10 counties that withdrew the largest amount of water from the Gulf Coast aquifer system during 1985, state officials projected a large decline in pumping from six counties, which included Matagorda County. The county was expected to experience a net decrease of 48 percent or 15 million gallons per day (mgd), with pumping rates decreasing from 31 mgd to approximately 16 mgd (Reference 2.4S.12-2). The water use projections undergo revisions and updating as technical and socioeconomic factors change. These factors are discussed later in this section.

The EPA monitors drinking water supply systems throughout the country and displays the results on their Safe Drinking Water Information System (SDWIS) website (Reference 2.4S.12- 10). Table 2.4S.12-4 presents a listing of SDWIS water supply systems in Matagorda County as of March 2007. Figure 2.4S.12-11 shows the locations of these water supply systems. A total of 40 systems are identified in Matagorda County by SDWIS with seven systems serving greater than 1000 people, 18 systems serving greater than 100 to less than 1000 people, and 15 systems serving less than or equal to 100 people. The closest SDWIS water supply systems are the onsite water supply (Water system ID TX1610051) and the Nuclear Training Facility water supply (Water system ID TX1610103). The nearest nonsite related SDWIS water supply system is the Selkirk Water System, which is located across the Colorado River from the STP, approximately 4 mi to the southeast (Water system ID TX1610027).

Groundwater use in the site area is controlled by the TWDB and locally (Matagorda County) by the Coastal Plains Groundwater Conservation District. The TWDB maintains a statewide database of wells called the Water Information Integration and Dissemination (WIID) system. This database includes water wells and petroleum production wells (Reference 2.4S.12-11). The Coastal Plains Groundwater Conservation District, in conjunction with the Coastal Bend Groundwater Conservation District (Wharton County), also maintains a database of water wells (Reference 2.4S.12-12).

Information from the TWDB database was used to prepare Figure 2.4S.12-12, which shows well locations near the STP site as of March 2007. Plate I in Appendix 2.4S.12-A shows known well locations in Matagorda County. This database includes water wells, driller's logs and petroleum wells, as designated on the figure and plate legends. Information for water wells contained in the database for Matagorda County is presented in tabular form in Appendices 2.4S.12-A1 and 2.4S.12-A2. The search area for wells was limited to Matagorda County because pumping effects in the Deep Aquifer and flow information in the Shallow Aquifer

suggest that groundwater impacts from groundwater use or accidents at STP would be limited to this area. The tables show a total of 838 water wells in Matagorda County. It should be noted that Appendix 2.4S.12-A2 (Driller's Report database) includes 18 wells identified as being in other counties, but the well coordinates plot within Matagorda County. It is not known if these entries have erroneous county names or location coordinates.

Figure 2.4S.12-13 presents the water well information from the Coastal Plains Groundwater Conservation District in the STP area as of March 2007. Plate II in Appendix 2.4S.12-A and Appendix 2.4S.12-A3 present the data for Matagorda County. The database includes 1989 water wells in Matagorda County and water use values for a portion of the wells. The larger number of wells in this database is primarily a result of including single-family domestic wells.

The TWDB conducts water use surveys throughout the state. The surveys are based on water user submitted information and may include estimated values. These surveys do not include single-family, domestic well groundwater use. The results of these surveys are divided up into use categories and water supply media (groundwater or surface water). Table 2.4S.12-5 presents regional historical groundwater and surface water use data for Matagorda County (Reference 2.4S.12-13). The table indicates that irrigation is the greatest groundwater user, followed by manufacturing, steam electric power generation, and municipal supply.

The TWDB also prepares estimates of future water use as part of water supply planning. These estimates have uncertainties associated with population growth projections, assumptions about climatic conditions (drought or wet years), and schedules for implementation of water conservation measures. The estimates of future water use for steam electric power generation include increased demand based on higher generation capacity and increased reservoir blowdown to maintain water quality. Table 2.4S.12-6 presents projected water use through the year 2060 (Reference 2.4S.12-14). This information was combined with historical water use to prepare the graphical representation of water use, as shown on Figure 2.4S.12-14. The relative percentages of water use categories are projected to remain the same as the historical data.

2.4S.12.2.2 Groundwater Flow Directions

A regional potentiometric surface map for the Deep Aquifer in Matagorda County in 1967 is presented on Figure 2.4S.12-15 (Reference 2.4S.12-6). Figure 2.4S.12-16 presents a potentiometric surface map for the Gulf Coast Aquifer from data collected between 2001 and 2005 (Reference 2.4S.12-15). Comparison of the figures suggests the regional flow direction of northwest to southeast is represented on both figures with localized flow disturbances caused by pumping. Comparison of the figures also suggests that groundwater elevations have increased in some parts of Matagorda County. In 1967, groundwater elevations above mean sea level were primarily located in the northern portion of the county. In the 2001-2005 potentiometric surface map, groundwater elevations in the northern and central portions of the county were above mean sea level. The hydraulic gradient in the STP site area for the 1967 potentiometric surface map is approximately 0.0006 ft/ft and for the 2001 to 2005 map is approximately 0.0002 ft/ft. Regional potentiometric surface maps are not available for the Shallow Aquifer due primarily to its limited regional use.

Using available information from the existing STP site piezometers, site-specific groundwater level measurements from November 1, 2005 and May 1, 2006 were used to develop

potentiometric surface maps for the Upper and Lower Shallow Aquifer (Figure 2.4S.12-17) and the Deep Aquifer (Figure 2.4S.12-18). The Upper Shallow Aquifer groundwater flow direction in the vicinity of STP $3 \& 4$ is generally toward the southeast. There is also an apparent southerly flow direction along the west side of the MCR. This southerly flow direction may be influenced by controlled leakage from the MCR or by the operation of the relief wells adjacent to the MCR dike. The groundwater flow direction in the vicinity of STP $3 \& 4$ in the Lower Shallow Aquifer is generally easterly. The Lower Shallow Aquifer flow direction turns southeasterly near the eastern edge of the site. Both the Upper and Lower Shallow Aquifer flow directions are consistent with flow toward the Holocene alluvium in the Colorado River floodplain. The potentiometric maps for the Deep Aquifer show the influence of onsite groundwater production, with a majority of the onsite groundwater flow toward the production wells. The onsite Deep Aquifer potentiometric surface suggests a reversal of the regional flow direction in the southern portion of the map, where flow is north towards the pumping wells, rather than toward the southeast.

The potentiometric surface maps were used to estimate hydraulic gradients at the site. For each map, a flow line originating in the area of STP $3 \& 4$ was drawn. The hydraulic gradient along these flow lines is estimated by dividing the head change along the flow line by the length of the flow line. The Upper Shallow Aquifer potentiometric surfaces indicate a hydraulic gradient of approximately 0.001 ft/ft. The Lower Shallow Aquifer maps indicate a hydraulic gradient of approximately 0.0004 ft/ft. The Deep Aquifer has a hydraulic gradient between approximately 0.0008 ft/ft and 0.002 ft/ft. The hydraulic gradient in the Deep Aquifer adjacent to STP 3 & 4 appears to be influenced primarily by changes in pumping at Production Well 6.

Monthly groundwater level measurements have been collected from the newly installed Shallow Aquifer observation wells for the STP $3 \& 4$ subsurface investigation. The measurements are presented on Table 2.4S.12-7. Well construction information is provided in Table 2.4S.12-1. The measurements were used to prepare the potentiometric surface maps shown on Figure 2.4S.12-19 for February and April of 2007. These maps indicate flow directions toward the southeast and southwest. The Upper Shallow Aquifer potentiometric surface map also shows seepage influence from the MCR and the duck pond/marsh located to the north of observation well pair OW-929U/L. The potentiometric surface maps indicate hydraulic gradients of approximately 0.001 ft/ft to 0.002 ft/ft for the southeast flow component in the Upper Shallow Aquifer and between approximately 0.0007 ft/ft and 0.0008 ft/ft for the southwest flow component. The Lower Shallow Aquifer hydraulic gradient is approximately 0.0004 ft/ft.

As part of the subsurface investigation program, well pairs screened in the Upper and Lower zones of the Shallow Aquifer were installed. These well pairs were used to estimate the vertical hydraulic gradient in the Shallow Aquifer. The vertical flow path length is assumed to be from the midpoint elevation of the Upper zone observation well screen to the midpoint elevation of the Lower zone observation well screen. Figure 2.4S.12-20 shows a generalized hydrogeologic section through the STP 3 $\&$ 4 area. This section shows the relationship between the Upper and Lower Shallow Aquifer zones and the interconnection of sand layers in the Lower Shallow Aquifer. The head difference over the vertical flow path is the difference in water level elevations between the two paired wells. The hydraulic gradient is estimated by dividing the head difference by the length of the flow path. Table 2.4S.12-8 presents the estimated vertical hydraulic gradients. All well pairs indicate a downward flow potential between the Upper and Lower zones in the Shallow Aquifer. The estimated vertical hydraulic gradients range from approximately 0.06 ft/ft to 0.29 ft/ft in a downward direction. Additional geologic and geotechnical cross-sections are provided in Section 2.5S.

A specific concern with respect to the groundwater flow direction in the Shallow Aquifer is the impact of the MCR on the groundwater system. Figure 2.4S.12-21 presents a conceptual hydrogeologic section extending from the MCR to the STP $3 \& 4$ area. This section suggests that the influence of the MCR is restricted to the area immediately downgradient (outside) of the reservoir dike. The combined effects of the relief wells and the toe drain act to reduce the head applied by the reservoir. Further evidence of the effectiveness of this drainage system is the absence of significant water ponding on the downgradient side of the MCR dike.

2.4S.12.2.3 Temporal Groundwater Trends

The TWDB has collected groundwater level data in Matagorda County since the 1930s (Reference 2.4S.12-16). Two observation wells near the STP were selected to prepare the regional hydrographs shown on Figure 2.4S.12-22. These wells monitor two different intervals in the Deep Aquifer. Well 8015402 monitors the heavy pumping interval at about 300 ft below ground surface. This well indicates that between 1957 and the early 1990s, a significant drop in groundwater level occurred. Since the early 1990s, the groundwater level has been recovering and has nearly returned to the 1957 level. The second well, 8015301, monitors the deeper zone of the Deep Aquifer, corresponding to the production zone in the STP onsite wells (well depths from 600 ft to 700 ft below ground surface). This well shows generally stable water levels over the period of record for the well. Due to the limited groundwater development potential in the Shallow Aquifer, regional temporal measurements of water levels have not been collected.

Groundwater levels are monitored in site observation wells as part of STP 1 & 2 operations. Selected observation wells in proximity to STP $3 \& 4$ were used to prepare hydrographs of the Shallow and Deep Aquifers, as shown on Figure 2.4S.12-23. The monitoring data set selected extends from March 1995 through May 2006. Upper Shallow Aquifer Wells 603B and 601 are located to the west and east, respectively, of STP $3 \& 4$ and well 602A, which is located immediately north of the STP 3 area. Well 603B shows some seasonal variability on the order of 1 ft to 2 ft, while Well 601 shows little seasonal variability. Well 602A shows some seasonal variability, with a peak groundwater elevation over the period of record of 25.8 ft MSL and with a long term variability of approximately 4 ft. Lower Shallow Aquifer wells 603A and 601A are located to the west and east, respectively, of STP $3 \& 4$. These wells show some seasonal variability with an overall decreasing trend in groundwater elevation. The elevation difference between the two wells suggests that they may be screened in different sand units within the Lower zone. Deep Aquifer observation wells 613 and 605 are located to the southwest and north, respectively, of STP $3 \& 4$. These wells show a notable increase in water level elevation between 1996 and 1998. Water levels in Well 613 show a slight declining trend between 2004 and 2006. Well 613 is located within the influence of STP Production Well 6, which may be the cause of the slight decrease in groundwater levels.

Shallow Aquifer observation wells installed as part of the STP $3 \& 4$ subsurface investigation program have been used for monthly water level measurements since December of 2006. Monthly groundwater levels will be collected through December 2007 from the STP $3 \& 4$

observation wells. Confirmatory information, based on the additional water level measurements, will be provided in a future COLA update in accordance with 10CFR50.71(e) (COM 2.4S-2). Three well series designations represent the following location areas.

- -OW-300 series wells are located in the proposed STP 3 facility area.
- -OW-400 series wells are located in the proposed STP 4 facility area.
- -OW-900 series wells include all of the wells located outside of the power block areas.

An "L" suffix on the well number indicates a Lower Shallow Aquifer well and a "U" suffix indicates an Upper Shallow Aquifer well.

Figure 2.4S.12-24 presents the hydrographs for these wells. These hydrographs suggest shortterm temporal variations in the Upper Shallow Aquifer on the order of 1 ft to 2 ft. The Upper Shallow Aquifer wells show consistently higher groundwater elevations than the adjacent Lower Shallow Aquifer wells. Within the STP 3 $\&$ 4 power block area, depth to groundwater is approximately 5 ft below ground surface.

Based on the water level elevations collected to date, the groundwater depth in both power block areas is below the maximum groundwater level of 61 cm (2 ft) below ground surface as specified in DCD/Tier 2 Table 2.0-1 for the ABWR. The plant ground floor grade elevation for safety-related structures is anticipated to be 35 ft MSL. Based on this observation, a permanent dewatering system will not be needed at STP 3 & 4.

2.4S.12.2.4 Aquifer Properties

Between 1951 and 1980 the average annual precipitation in the general area of STP was about 42 inches, and the corresponding average annual runoff is estimated as about 12 inches (Reference 2.4S.12-2). The difference of approximately 30 inches is either evaporated, consumed by plants, or percolates into the vadose zone to recharge the shallow aquifers. Much of the water is returned to the atmosphere by evapotranspiration (Reference 2.4S.12-2).

The vadose zone is considered to be relatively thin and limited at the site. The first saturated sand zone is encountered at a general depth of approximately 20 ft below ground surface, and it is classified as part of the Upper Shallow Aquifer. The aquifer zone exhibits semi-confined to confined conditions. The potentiometric head is under pressure, rising to within 5 ft to 10 ft of ground surface as measured in the onsite observation wells. The soils overlying the sand are generally described as clay (CL to CH, USCS Groups). From the geotechnical data listed in Subsection 2.5S.4, measured natural moisture contents from samples collected to a depth of 20 ft ranged from approximately 5 percent to 29 percent. The majority of the values ranged between 15 percent and 25 percent. Dry unit weights for the materials sampled ranged from approximately 92 pounds per cubic foot (pcf) to 115 pcf. Wet densities, when measured, ranged from approximately 97 pcf to 133 pcf.

The properties of the aquifer materials at the STP site are divided into hydrogeological and geotechnical derived parameters. The hydrogeological parameters include transmissivity and storage coefficient measurements from aquifer pumping tests and hydraulic conductivity values determined from historical aquifer pumping tests and the slug tests performed in December

2006 as part of the STP 3 $\&$ 4 site subsurface investigation. The geotechnical parameters derived from laboratory testing include bulk density (or dry unit weight), porosity, effective porosity, and permeability from grain size.

The following are definitions of hydrogeological parameters adapted from Reference 2.4S.12- 17:

- - Transmissivity - The rate at which a fluid of a specified density and viscosity is transmitted through a unit width of an aquifer or confining bed under a unit hydraulic gradient and is a function of the properties of the fluid, the porous medium, and the thickness of the porous medium.
- - Storativity (Storage Coefficient) - The volume of water released from or taken into storage per unit surface area of the aquifer per unit change in head.
- - Hydraulic Conductivity (permeability) - A coefficient of proportionality describing flow per unit time under a unit hydraulic gradient through a unit area of a porous medium and is a function of the properties of the fluid and the porous medium.

2.4S.12.2.4.1 Hydrogeological Parameters

Regional aquifer properties have been collected by the TWDB (Reference 2.4S.12-6). Data for the area in proximity to the STP site is presented in Table 2.4S.12-9. Deep Aquifer transmissivity ranges from 10,500 gpd/ft to 195,300 gpd/ft and storage coefficient ranges from 4.6×10^{-5} to 1.4 \times 10⁻³. Although several of the wells in the table have screened intervals that encompass the depth interval associated with the Shallow Aquifer at the STP site, the screened intervals also extend into the Deep Aquifer, thus the test results cannot be applied to the Shallow Aquifer. Aquifer pumping tests have been performed on the STP site (Reference 2.4S.12-7 and Reference 2.4S.12-8) at three of the Deep Aquifer production wells and four test wells in the Shallow Aquifer in support of STP $1 \& 2$. The results of these tests are summarized in Table 2.4S.12-10. Transmissivity ranges from 1100 gpd/ft to 50,000 gpd/ft and the storage coefficient ranges from 2.2 x 10^{-4} to 1.7 x 10^{-3} .

Figure 2.4S.12-25 presents a graphical comparison of regional and site-specific measurements using box and whisker plots. The box and whisker plot, also known as a boxplot, is a graphical representation of the data based on dividing the data set into quartiles. The data range of the solid portion of the box encompasses 50 percent of the data and the data range of each whisker contains 25 percent of the data. The ends of the whiskers represent the minimum and maximum values in the data set. Examination of the transmissivity plot indicates that the regional and STP deep values fall within the same data range, while the STP Shallow Aquifer data range falls below the regional range. This is caused by two Upper Shallow Aquifer tests that have transmissivity values of 1100 gpd/ft and 12,500 gpd/ft. The plot for storage coefficient indicates that the regional, STP Deep Aquifer, and STP Shallow Aquifer all fall within the same data range. The Shallow Aquifer values fall within the upper portion of the regional range of data. This may be a result of aquitard leakage influencing the Shallow Aquifer tests.

Hydraulic conductivity can be determined from aquifer pumping tests by dividing the transmissivity by the saturated thickness. There is uncertainty associated with this method, because assumptions are made regarding the amount of permeable material present within the screened interval of the test well. The pumping wells have screened intervals ranging from 16 ft to 819 ft in length, and the saturated thickness is apportioned across this screened interval (possibly underestimating the hydraulic conductivity for the more permeable sand units crossed by the well screen intervals). Hydraulic conductivity values from the aquifer pumping tests are included in Table 2.4S.12-9 and Table 2.4S.12-10.

Hydraulic conductivity can also be determined by the slug test method. This method measures the water level response in the test well to an instantaneous change in water level in the well. A disadvantage of this method is that it measures hydraulic conductivity only in the immediate vicinity of the test well. However, because the slug test requires minimal equipment and can be performed rapidly, slug tests can be performed in many wells, allowing a determination of spatial variability in hydraulic conductivity. Table 2.4S.12-11 presents a summary of slug tests performed in observation wells installed as part of the STP 3 & 4 subsurface investigation program. The test results indicate a range of hydraulic conductivity from 9 gpd/ft² to 561 gpd/ft^2 . The slug test results for the Upper and Lower zones of the Shallow Aquifer were contoured, as shown on Figure 2.4S.12-26 to delineate spatial trends. The Upper Shallow Aquifer contour map indicates areas of higher hydraulic conductivity in the vicinity of STP 3 and to the northwest of STP 4. The surrounding measurements suggest these areas are localized. The Lower Shallow Aquifer map indicates an area of higher hydraulic conductivity between STP 3 $&$ 4 and extending to the south of the units. This area corresponds to the area of higher groundwater elevation identified on the February 22, 2007 potentiometric surface map for the Lower Shallow Aquifer shown on Figure 2.4S.12-19. The correspondence between a higher hydraulic conductivity area and higher potentiometric elevation suggests the presence of a flow pathway, such as a paleochannel, from the MCR toward STP $3 \& 4$.

Box and whisker plots comparing hydraulic conductivity from regional aquifer pumping tests, STP site aquifer pumping tests, STP site slug tests, and grain size data are shown on Figure 2.4S.12-27. The grain size derived hydraulic conductivity is discussed in Subsection 2.4S.12.2.4.2. The plots indicate that the slug tests have the greatest range of hydraulic conductivity. However, the geometric means for the aquifer pumping test derived hydraulic conductivity values and the slug test results are not significantly different $(337 \text{ gpd}/\text{ft}^2)$ versus 205 gpd/ $\rm ft^2$).

2.4S.12.2.4.2 Geotechnical Parameters

The geotechnical investigation component of the STP $3 \& 4$ subsurface investigation program included the collection of soil samples for laboratory determination of soil properties. These tests are discussed in Section 2.5S.4. A summary of the test results is presented in Table 2.4S.12-12. The results have been arranged to reflect the properties of the various hydrogeologic units present at the site. Basic soil properties are used to estimate the hydrogeologic properties of the materials such as porosity, effective porosity (specific yield), and permeability. Bulk density values were measured by the laboratory thus no further processing of the data was necessary.

Porosity is determined from a conversion of the void ratio to porosity. The effective porosity (or specific yield) is some fraction of porosity. In general terms, the effective porosity of sands or gravels approximates porosity, while the effective porosity of silts and clays is much less than their porosity. Figure 2.4S.12-28 (from Reference 2.4S.12-18) is a graph that shows the

relationship between porosity, specific yield, and specific retention for various median grain sizes and sorting conditions. Interpolating from this graph for median grain sizes in the Shallow Aquifer and using the curve for average material, suggests that the specific yield is approximately 80 percent of the porosity of the Shallow Aquifer.

Permeability or hydraulic conductivity of sands with a D_{10} grain size between 0.1 and 3.0 mm can be estimated using the Hazen approximation (Reference 2.4S.12-18). This formula was based on empirical studies for the design of sand filters for drinking water. The formula was developed for use in well-sorted sand and application to poorer-sorted materials would result in over-prediction of permeability. Figure 2.4S.12-27 includes the grain size derived hydraulic conductivity with aquifer pumping test and slug test derived hydraulic conductivity. Comparison of the boxplots suggests that the grain size derived hydraulic conductivity is within the range of regional hydraulic conductivity values and the STP aquifer test ranges. Comparison of geometric means indicates the grain size derived hydraulic conductivity is similar to the STP aquifer test results.

The hydraulic conductivity of the clay materials was measured in the STP 1 $\&$ 2 subsurface investigation (Reference 2.4S.12-9). Table 2.4S.12-13 summarizes the results of these tests. The geometric mean hydraulic conductivity of the clay samples is 0.004 gpd/ft² (1.72 x 10⁻⁷) cm/sec). The clay samples were collected to a maximum depth of 39 ft below ground surface. The uniform depositional history and effects of consolidation and loading on clay hydraulic conductivity suggest that it would be a conservative assumption to apply these hydraulic conductivity values to deeper clays at the site.

2.4S.12.2.4.3 Representative Properties of Hydrogeologic Units

A simplified conceptual model of the STP site was developed to apply site parameters to the estimation of groundwater flow and contaminant transport. Figure 2.4S.12-29 presents a simplified hydrostratigraphic section of the site. The units presented on the section were used as a framework to relate measured or estimated properties to the groundwater system. A summary of important properties related to groundwater flow and contaminant transport is presented in Table 2.4S.12-14. The values for bulk density, total porosity, and effective porosity for the Deep Aquifer were taken from tests performed in the Lower Shallow Aquifer. The similarity of depositional environments and the observed grain size distributions suggest that an assumption of equivalence between the units is reasonable.

To assign representative values, the properties were divided into spatially and temporally variable data. Spatially variable data includes unit thickness, hydraulic conductivity, bulk density, porosity, and effective porosity. Representative values for the spatially variable data were assigned either an arithmetic mean (unit thickness, bulk density, porosity, and effective porosity) or a geometric mean (hydraulic conductivity) of the referenced data set. Temporally variable data are the hydraulic gradient measurements; the maximum value from each data set is assigned as the representative value.

2.4S.12.2.5 Hydrogeochemical Characteristics

Regional hydrogeochemical data were obtained from Reference 2.4S.12-6 and are presented in Table 2.4S.12-15. The data set includes ten wells in the Deep Aquifer and seven wells in the Shallow Aquifer. The analytical data was compared to EPA Primary and Secondary Drinking

Water Standards (Reference 2.4S.12-19) and exceedances are identified on the table. The principal exceedances were for total dissolved solids and chloride (Secondary Drinking Water Standards). Examination of data suggests that the highest concentrations of total dissolved solids and chlorides are present in the Shallow Aquifer.

STP site-specific hydrogeochemical data are presented in Table 2.4S.12-16, which includes seven samples from the Deep Aquifer and 23 samples from the Shallow Aquifer. The analytical data were compared to EPA Primary and Secondary Drinking Water Standards and the exceedances are identified in the table. The principal exceedances were for total dissolved solids and chloride. The data indicate that the highest concentrations of total dissolved solids and chloride are present in the Shallow Aquifer.

The hydrogeochemical data can also be used as an indicator of flow patterns in the groundwater system. Variations in chemical composition can be used to define hydrochemical facies in the groundwater system. The hydrochemical facies are classified by the dominant cations and anions in the groundwater sample. These facies may be shown graphically on a trilinear diagram (Reference 2.4S.12-20). A trilinear diagram showing the regional and STP sitespecific data is presented on Figure 2.4S.12-30. The predominant groundwater type for the Deep Aquifer regional groundwater data is sodium-bicarbonate, while for the Shallow Aquifer regional data the groundwater type varies from sodium-bicarbonate to sodium-chloride. The predominant STP site-specific groundwater type is sodium-bicarbonate in the Deep Aquifer, sodium-chloride in the Upper Shallow Aquifer, and sodium-bicarbonate in the Lower Shallow Aquifer. An exception to the Lower Shallow Aquifer hydrochemical facies pattern is observed at observation wells OW-332L and OW-930L, where the water type is sodium-chloride. This facies change may indicate the proximity of a zone of vertical interconnection between the Upper and Lower Shallow Aquifers. This observation would be consistent with the findings of aquifer pumping test WW-4 (Reference 2.4S.12-7), which indicates a localized hydraulic connection between the Upper and Lower Shallow Aquifers. The conclusion that this is a localized connection is based on the absence of a hydraulic connection at the other three aquifer pumping test sites. The source of this interconnection may be either a natural feature, such as an incised channel or scour feature, or a man-made feature such as an excavation backfilled with pervious material or a leaking well seal. The manmade sources of interconnection are less probable, since the depth to the Lower Shallow Aquifer is on the order of 60 ft below ground surface, which would be below most site excavations, and leaky well seals also typically exhibit elevated pH associated with the impacts of cement grout, which is not observed at either of the wells.

Comparison of historical and more recent hydrogeochemical data indicates a general temporal consistency in groundwater chemistry for the individual aquifers present in the site area. This suggests that there are no long-term variations in groundwater chemistry occurring the site area.

2.4S.12.3 Subsurface Pathways

This section presents an evaluation of subsurface pathways for offsite exposure resulting from a liquid effluent release at STP $3 \& 4$. The section focuses on advective groundwater flow. Discussion of sorption and radioactive decay effects on offsite exposure is presented in Subsection 2.4S.13.

2.4S.12.3.1 Exposure Point and Pathway Evaluation

Figure 2.4S.12-31 presents the Blessing SE U.S. Geological Survey 7.5 minute quadrangle map of the site area (Reference 2.4S.12-21). This map shows onsite and offsite surface features considered in the evaluation. Review of regional groundwater use data presented in Subsection 2.4S.12.2.1 indicates that there is a credible Shallow Aquifer groundwater user exposure point in the vicinity of the STP site at Well 2004120846. This would be the most likely exposure point for the Shallow Aquifer groundwater. A second exposure pathway is via surface water, where the Shallow Aquifer discharges to local creeks or the Colorado River. The most likely exposure point for the Deep Aquifer would be the onsite groundwater production wells.

Off-site migration pathways were evaluated for the following hydrogeologic units:

- -Upper Shallow Aquifer
- -Lower Shallow Aquifer
- -Deep Aquifer

The Upper Shallow Aquifer is the most likely hydrogeologic unit to be impacted by an accidental liquid effluent release onsite. Due to the shallow depth of this unit, a conservative release scenario would be a direct injection of liquid effluent into the Upper and Lower Shallow Aquifer. The Upper Shallow Aquifer has a flow direction toward the southeast, as discussed in Subsection 2.4S.12.2.2. Examination of Figure 2.4S.12-31 indicates that a potential Upper Shallow Aquifer groundwater discharge area would be the unnamed tributary, located to the east of the STP 1 & 2 Essential Cooling Pond (ECP), which flows into Kelly Lake, approximately 7300 ft from STP 3. A second possible discharge area for both the Upper and Lower Shallow Aquifer is at Well 2004120846, which is an 80 ft deep livestock well, located east of the site boundary approximately 9000 ft from STP 3. This pathway assumes the well discharges to stock watering containers and that the groundwater is consumed by livestock, which would be an indirect human exposure pathway. Information from Appendix 2.4S.12-A3 indicates this well is estimated to produce 200,000 gallons per year or approximately 0.4 gpm. A third possible discharge area for both Shallow Aquifer units would be the Colorado River, approximately 17,800 ft from STP 3.

The Lower Shallow Aquifer is isolated over much of the site by the Lower Shallow Aquifer Confining Layer. However, aquifer pumping test data (Subsection 2.4S.12.2.4.1) and hydrogeochemical data (Subsection 2.4S.12.2.5) suggest that leakage through the less permeable confining layer is occurring. Additionally, excavations for the foundations of some of the deeper structures are projected to enter the Lower Shallow Aquifer. Subsection 2.4S.12.2.2 indicates that a consistent downward vertical hydraulic gradient exists between the Upper and Lower Shallow Aquifer, which would provide the driving force for movement of groundwater from the Upper to the Lower Shallow Aquifer in the leakage areas. A conservative effluent release scenario would be a direct effluent release into the Lower Shallow Aquifer. Subsection 2.4S.12.2.2 indicates the Lower Shallow Aquifer has an east to southeast flow direction. Due to the depth to the top of the aquifer and the downward vertical hydraulic gradient in the Lower Shallow Aquifer, it is unlikely that discharge would occur into the unnamed tributary to the east of the STP $1 \& 2$ ECP. Likely discharge points are Well

2004120846, as discussed above, or the Colorado River alluvium, where the river channel has incised into the Lower Shallow Aquifer, approximately 17,800 ft from STP 3 & 4.

The Deep Aquifer is the least likely hydrogeologic unit to be impacted by an accidental liquid effluent release. The Deep Aquifer is separated from the Shallow Aquifer by a 100 ft to 150 ft thick clay and silt layer. Recent potentiometric surface maps for the Deep Aquifer (Subsection 2.4S.12.2.2) indicate that groundwater flow in the plant area is moving toward the production wells at the site, thus precluding the potential for offsite migration should the effluent pass through the clay layer. The additional groundwater needs for operation of STP $3 \& 4$ will further depress the potentiometric surface in the Deep Aquifer. The combined effects of horizontal flushing by flow in the Shallow Aquifer, radionuclide sorption as the effluent passes through the 100+ ft thick clay layer, and groundwater capture by the site production wells suggest that there is no credible offsite release pathway for the Deep Aquifer.

2.4S.12.3.2 Advective Transport

Advective transport assumes that a accidental liquid effluent release travels at the same velocity as groundwater flow. The groundwater flow velocity or average linear velocity is estimated from the following equation (Reference 2.4S.12-17):

$$
v = \frac{Ki}{n_e}
$$

where:

 $v =$ average linear velocity (ft/day)

 $K =$ hydraulic conductivity (ft/day)

 $i =$ hydraulic gradient (ft/ft)

 n_e = effective porosity (decimal)

The travel time from the effluent source to the receptor would be:

$$
T = \frac{D}{\nu}
$$

where:

 T = travel time (day)

 $D =$ distance from source to receptor (ft)

 $v =$ average linear groundwater velocity (ft/day)

Table 2.4S.12-17 presents average linear velocity and travel time estimates for the Upper Shallow Aquifer using information from Table 2.4S.12-14. The table includes ranges of groundwater velocities and travel times for the extremes (high and low) of the data set. The average linear velocity in the Upper Shallow Aquifer is estimated to be 0.2 ft/d and in the Lower Shallow Aquifer to be 0.09 ft/d. In the Upper Shallow Aquifer, travel time to the unnamed tributary east of the ECP would be 100 years, to Well 2004120846 it would be 123 years, and to the Colorado River it would be 244 years. In the Lower Shallow Aquifer, travel time to Well 2004120846 would be 274 years and to the Colorado River it is estimated to be 541 years.

2.4S.12.3.3 Plant Groundwater Use and Effects

Groundwater is projected to be the main source of water for STP 3 & 4 plant construction and operation. During construction, groundwater use requirements will vary and will be used for the following activities: onsite personnel consumption and use; manufacturing of concrete, concrete curing, and clean-up; dust control; addition of moisture and placement of engineered backfill; and piping hydro tests and flushing. Preliminary estimates indicate that up to 1200 gpm of groundwater will be required during construction.

STP is currently permitted to use up to 3000 acre-ft per year of groundwater from their existing production wells. STP currently uses about 1300 acre-ft per year for plant operations. Therefore, approximately 1700 acre-ft per year (1050 gpm) of groundwater could be available for construction use. Water demand could be met by increasing the yield of the existing wells or by installing new wells with the objective that total STP use would not exceed the 3000 acreft per year permitted amount. A detailed evaluation of groundwater availability and estimates of aquifer drawdown, requirements for permitting of new wells and yields, water conservation measures, and the identification of alternative sources, if practicable, will be addressed as part of the detailed engineering for STP 3 & 4.

Operation of STP 3 $\&$ 4 is predicted to require a typical groundwater consumption of 1077 gpm or 1738 acre-ft per year, whereas the peak groundwater consumption for STP 3 $\&$ 4 is expected to be as great as 3935 gpm, when required (i.e., outages). The projected combined STP plant normal groundwater consumption for STP 1 $& 2$ and STP 3 $& 4$ is expected to be between 2938 and 3038 acre-ft per year, which is approximately equal to the current STP permitted use of 3000 acre-ft per year. Peak demand for outages could be met by increasing the permitted groundwater allotment for short-term uses or by obtaining water from other sources such as the MCR or the Colorado River.

Based on these estimates, additional groundwater wells will be required to satisfy site demands. As with STP 1 $\&$ 2, it is expected that no sustained pumping will be permitted within 4000 ft of the plant safety-related facility areas in order to minimize the potential for regional subsidence resulting from lowering of the Deep Aquifer zone potentiometric head. Based on this requirement, the location of the additional groundwater wells required for expanded plant operations would most likely be located in the northwestern and northeastern sections of the STP site and/or in the southeastern and southwestern site areas adjacent to the MCR.

As stated in Subsection 2.4S.12.2.2, comparison of a regional potentiometric surface map for the Deep Aquifer in Matagorda County in 1967 (Figure 2.4S.12-15) and that of a potentiometric surface map for the Gulf Coast Aquifer from data collected between 2001 and 2005 (Figure 2.4S.12-16) suggests that groundwater elevations have increased in some parts of Matagorda County. In 1967, groundwater elevations above mean sea level were primarily located in the northern portion of the county. In the 2001-2005 potentiometric surface map, groundwater

elevations in the northern and central portions of the county were above mean sea level. Therefore, the regional impacts of groundwater production on the aquifer groundwater levels appear to be decreasing, thus minimizing impact to the regional aquifer as the result of STP plant expansion with the construction and operation of STP $3 \& 4$. Some additional aquifer drawdown would be expected near the STP site boundaries as the result of installing and operating new groundwater wells. Based on Figure 2.4S.12-18, it can be expected that the lowering of the potentiometric head in the Deep Aquifer at the existing STP production would expand over most of the northern portion of the site due to the installation of the new site production wells. The decrease in head would be expected to extend beyond the site boundaries but the impact would be less than that beneath the site.

As part of the detailed engineering for the STP 3 $\&$ 4, the impact of the groundwater pumping in the Deep Aquifer will be evaluated to the current site conditions and that of nearby, offsite groundwater users (Figure 2.4S.12-13). Permitting of new wells and yields, plant water conservation methods, and the identification of alternative sources or recycling, if practicable, will be addressed as part of the detailed engineering for STP 3 & 4.

2.4S.12.4 Monitoring or Safeguard Requirements

Groundwater level monitoring in the STP $3 \& 4$ area is currently being implemented through the use of the groundwater observation wells installed in 2006 for the site subsurface investigation and through the periodic review of water levels from selected wells in the vicinity of the site.

Some of the existing STP 3 $\&$ 4 area observation wells will be taken out of service prior to construction activities due to anticipated earth moving and construction requirements. Prior to construction activities, the observation well monitoring network will be evaluated in the detailed design to determine groundwater data gaps and needs created by the abandonment of existing wells.

As part of the detailed design for STP 3 $\&$ 4, the current STP groundwater monitoring programs will be evaluated with respect to the addition of STP $3 \& 4$ to determine if any modification of the existing programs is required to adequately monitor plant effects on the groundwater. Considerations to revise the site groundwater monitoring program will include the following components:

- - Deep Aquifer - Periodic water level measurements in deep observation wells and geochemical sampling and analysis of production wells would detect changes in the Deep Aquifer that may impact groundwater supply availability or the accident release analysis.
- - Shallow Aquifer - Periodic water level measurements in the Upper and Lower zone observation wells and collection of geochemical samples and analysis will be performed in selected observation wells. The water level monitoring program objective is to detect changes in flow patterns in the Shallow Aquifer that might impact accident analysis and would track temporal trends in groundwater levels that might impact structural stability. The geochemical monitoring would detect changes in groundwater geochemistry that would be deleterious to plant structures and subsurface components.
- - Subsidence Monitoring - The current plant subsidence monitoring program will be expanded to include STP 3 & 4.
- - Operational Accident Monitoring - In the unlikely event of an operational accident, site observation wells in the Shallow and Deep Aquifers and onsite groundwater production wells in the Deep Aquifer would be sampled for radionuclides associated with the plant. Additional monitoring locations may be added if onsite monitoring indicates the potential for offsite exposure.

Groundwater level measurements in the Deep and Shallow Aquifers would be collected starting during construction and after plant startup Selection of observation wells to be included in the program will be made prior to the start of operation based on well condition, position relative to plant site and other observation wells (provide optimal spatial distribution for potentiometric map preparation and vertical hydraulic gradient assessment), and long-term viability of the observation well (likelihood well will survive construction).

Geochemical sampling and analysis in the Deep and Shallow Aquifers would be performed during construction and after startup. Analysis will include field parameters (pH, temperature, specific conductance, oxidation-reduction potential, and dissolved oxygen), major cations, major anions, total dissolved solids, and silica. Sampling would be performed in site production wells, any new production wells installed to support STP $3 \& 4$ operation, and selected observation wells in the Shallow Aquifer. Observation wells would be selected during detailed design.

Additional near-surface subsidence monuments would be installed around STP 3 $\&$ 4 structures. The onsite subsidence monitoring frequency would increase during construction and after startup.

Operational accident monitoring would be triggered in the unlikely event of a release of liquid effluent from the plant. Quarterly groundwater samples would be collected from site production wells and downgradient Shallow Aquifer observation wells. Selection of downgradient observation wells would be based on flow directions determined from the most recent groundwater level measurements.

Safeguards will be used to minimize the potential of adverse impacts to the groundwater by construction and operation of the new units. These safeguards would include the use of lined containment structures around storage tanks (where appropriate), hazardous materials storage areas, emergency cleanup procedures to capture and remove surface contaminants, and other measures deemed necessary to prevent or minimize adverse impacts to the groundwater beneath the STP 3 & 4 site.

2.4S.12.5 Site Characteristics for Subsurface Hydrostatic Loading

Subsurface hydrostatic loading estimates for structures at STP $3 \& 4$ were evaluated using two approaches. First, by conservatively assuming the maximum groundwater level of 61 cm (2 ft) below ground surface as specified in DCD Table 2.0-1 for the ABWR. The existing plant grade at the site is approximately 30 ft MSL and the finished plant grade in the Power Block area is anticipated to be between approximately 32 and 36.5 ft MSL, thus a grade elevation of 35 ft MSL would result in a maximum groundwater elevation of 33 ft MSL. The second approach

uses the maximum observed groundwater level elevation (December 2006 - June 2007), within the STP 3 & 4 power block area; elevation 25.85 ft MSL from observation well OW-332U on April 27, 2007. The maximum hydrostatic loading is estimated using the following formula:

 $\rho_w = z_w \times \gamma_w$

where:

 ρ_w = hydrostatic pressure (psf)

 z_w = depth below groundwater level (ft)

 γ_w = unit weight of water (62.4 pcf)

Figure 2.4S.12-32 presents a graph of building elevation versus hydrostatic pressure. Two lines are provided on the graph, one representing the upper bound condition using the DCD maximum groundwater level and the second using the maximum observed groundwater level in the power block area.

Excavations for the construction of STP 3 $\&$ 4 are preliminarily planned to depths of about 90 ft below existing grade. The reactor building mat is expected to be placed at a depth of approximately 85 ft with the control building at a depth of approximately 75 ft, the UHS at a depth of approximately 40 ft, and the turbine building at a depth of approximately 30 ft. Perimeter dewatering will be required to a depth of at least 35 ft with deeper excavation dewatering to a depth of at least 100 ft. To minimize excess dewatering, the UHS would utilize a separate perimeter dewatering system. The excavation design may require the use of slope stability structures. The actual excavation design will be refined as part of the detailed design.

During excavation and construction of STP $3 \& 4$, the hydrostatic loading on the excavation and structures will be controlled by a temporary construction dewatering system. Typical dewatering systems for this type of cut and fill excavation would consist of a combination of perimeter dewatering wells and open pumping from sumps within the excavation. The perimeter dewatering wells would control lateral inflow and assist in removing water stored within the excavation. The open pumping system would control precipitation run-off, assist in water storage removal, and removal of any inflow to the excavation.

To prevent uplift of foundation soils, groundwater levels will be maintained a minimum of 5 ft below the bottom of the deepest excavation. The STP 3 & 4 excavation is deeper than the excavation for STP 1 $\&$ 2 (Reference 2.4S.12-9). The hydrogeologic conditions encountered beneath the proposed STP 3 & 4 are, in general, similar to that beneath STP 1 & 2. A longterm, steady state dewatering flow rate is estimated to be between 1800 and 4200 gpm. The range in pumping rates is dependent on the hydraulic conductivity used in the analysis (low range or geometric mean of the pumping test hydraulic conductivity values) and on the excavation plan. Because the excavation required for the construction of STP $3 \& 4$ is estimated to be deeper than that for STP 1 & 2, the flow rates estimated for STP 3 & 4 are considered to be within reason in comparison to actual flow rates measured at STP 1 & 2 (between 1300 gpm to 2900 gpm). Alternatives that could reduce the amount of water to be removed include various types of cut-off walls. The cut-off walls could include a slurry wall, grout curtain, or freeze-wall. The slurry wall and grout curtain are permanent features, while

the freezewall can be temporary. Some dewatering would still have to be performed to remove storage, precipitation run-off, and vertical inflow. Methods to mitigate the subsidence beneath existing structures include cut-off walls, injection wells, and infiltration trenches. The dewatering system design will be refined as part of the detailed design.

Another concern is rewatering after completion of excavation and backfill around structures. Groundwater levels will be raised in a controlled manner to prevent rapid hydrostatic pressure build-up or damage to the subsurface backfill materials. Prior to the start of excavation, a dewatering and rewatering plan will be prepared to document the construction dewatering system design and groundwater control criteria.

In summary, based on the water level elevations collected to date, the groundwater depth in both power block areas is below the maximum groundwater level of 61 cm (2 ft) below ground surface as specified in DCD Table 2.0-1 for the ABWR. Based on this observation, a permanent dewatering system is not anticipated to be a design feature for the STP 3 & 4 facility. Post-construction groundwater conditions are anticipated to have some localized changes resulting from excavation and backfilling, however, based on observations of STP 1 & 2 postconstruction groundwater conditions, the effects would be minimal and may include localized communication between the Upper and Lower Shallow Aquifers and an increased cone of depression in the Deep Aquifer resulting from increased groundwater use for STP $3 \& 4$. The groundwater supply wells to be installed for STP 3 $\&$ 4 are not a safety-related source of water because the UHS has a 30-day supply of water, which is sufficient for plant shutdown without a supplementary water source.

2.4S.12.6 References

- 2.4S.12-1 "Physiographic Map of Texas," Bureau of Economic Geology, The University of Texas at Austin, 1996. Available at http://www.lib.utexas.edu/geo/maps.html.
- 2.4S.12-2 "Ground Water Atlas of the United States: Oklahoma, Texas, HA 730-E," Ryder, P. D., United States Geological Survey, 1996. Available at http://capp.water.usgs.gov/gwa/ch_e/index.html.
- 2.4S.12-3 "Major Aquifers of Texas," 2000. Available at http://www.twdb.state.tx.us/mapping/index.asp.
- 2.4S.12-4 "Region VI Sole Source Aquifer Map," United States Environmental Protection Agency. Available at http://www.epa.gov/region6/6wq/swp/ssa/maps.htm, accessed March 2, 2007.
- 2.4S.12-5 "Stratigraphic and Hydrogeologic Framework of Part of the Coastal Plain of Texas, Report 236," Baker, E.T., Texas Department of Water Resources, 1979. Available at http://www.twdb.state.tx.us/publications/reports/GroundWaterReports/GWRepor ts/Individual%20Report%20htm%20files/Report%20236.htm.
- 2.4S.12-6 "Ground-Water Resources of Matagorda County, Texas, Texas Water Development Board Report 91," Hammond Jr., W., Cooperative study by the Texas Water Development Board, Lower Colorado River Authority, and Matagorda County Commissioners Court, 1969. Available at http://www.twdb.state.tx.us/publications/reports/GroundWaterReports/GWRepor ts/Individual%20Report%20htm%20files/Report%2091.htm.
- 2.4S.12-7 "Basic Data Report Ground Water South Texas Project, submitted to Brown and Root, Inc.," Woodward-Clyde Consultants, August 29, 1975.
- 2.4S.12-8 "Deep Aquifer Ground-Water Evaluation and Pump Test Results, report Y180XR042-WL, prepared for Brown & Root, Inc.," Woodward-Clyde Consultants, July 9, 1975.
- 2.4S.12-9 "STPEGS Updated Final Safety Analysis Report Units 1 and 2," Revision 13.
- 2.4S.12-10 "United States Environmental Protection Agency Safe Drinking Water Information System (SDWIS)." Available at http://oaspub.epa.gov/enviro/sdw_query_v2.get_list?wsys_name=&fac_search= fac_beginning&fac_county=MATAGORDA&pop_serv=500&pop_serv=3300&po p_serv=10000&pop_serv=100000&pop_serv=100001&sys_status=active&pop_ serv=&wsys_id=&fac_state=TX&last_fac_name=&page=1&query_results=&tota l_rows_found=, accessed March 15, 2007.
- 2.4S.12-11 "Texas Water Development Board Water Information Integration and Dissemination (WIID) System." Available at http://wiid.twdb.state.tx.us/, accessed March 15, 2007.
- 2.4S.12-12 "Coastal Plain/Coastal Bend Groundwater Conservation Districts Database." Available at http://www.gis.aecom.com/cbcpgcd/, accessed March 20, 2007.
- 2.4S.12-13 "Texas Water Development Board historical water use information." Available at: http://www.twdb.state.tx.us/wushistorical/, accessed March 12, 2007.
- 2.4S.12-14 "Texas Water Development Board water use projections." Available at http://www.twdb.state.tx.us/DATA/db07/defaultReadOnly.asp, accessed March 12, 2007.
- 2.4S.12-15 "Hydrogeochemistry, Salinity Distribution, and Trace Constituents: Implications for Salinity Sources, Geochemical Evolution, and Flow Systems Characterization, Gulf Coast Aquifer, Texas" Chapter 5 in 'Aquifers of the Gulf Coast of Texas,' Texas Water Development Board Report 365," Chowdhury, A., Boghici, R., and Hopkins, J., 2006. Available at http://www.twdb.state.tx.us/publications/reports/GroundWaterReports/GWRepor ts/R365/AGCindex.htm.
- 2.4S.12-16 "Texas Water Development Board Well Database Information." Available at http://www.twdb.state.tx.us/publications/reports/GroundWaterReports/GWDatab aseReports/Database%20in%20ASCII/Matagorda/wlevels.txt, accessed March 14, 2007.
- 2.4S.12-17 "Applied Hydrogeology, second edition," Fetter, C.W., 1988.
- 2.4S.12-18 "Hydrogeology," Davis, S.N, and DeWiest, R.J.M., 1966.
- 2.4S.12-19 "Drinking Water Contaminants," United States Environmental Protection Agency. Available at http://www.epa.gov/safewater/contaminants/index.html, accessed April 19, 2007.
- 2.4S.12-20 "A Graphic Procedure in the Geochemical Interpretation of Water Analyses 'Transactions, American Geophysical Union', 25, 914-923," Piper, A.M., 2006.
- 2.4S.12-21 "Blessing SE, Texas, 28096-G1-TF-024, DMA 6741 1 SE-SERIES V882," United States Geological Survey, 1995. Available at http://www.tnris.org/datadownload/quad.jsp?Quad=Blessing%20SE&num=2809 6-G1.
- 2.4S.12-22 "Climate Division: Temperature-Precipitation-Drought Data," National Climatic Data Center (NCDC), 2007. Available at http://www.ncdc.noaa.gov/ oa/climate/onlineprod/drought/ftppage.html, accessed April 13, 2007.

Table 2.4S.12-2 STP Production Well Information **Table 2.4S.12-2 STP Production Well Information**

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Table 2.4S.12-4 Listing of U.S. Environmental Protection Agency (EPA) Safe Drinking Water Information System (SDWIS) Community and Non-Community Groundwater Systems in Matagorda County, Texas

Table 2.4S.12-4 Listing of U.S. Environmental Protection Agency (EPA) Safe Drinking Water Information System (SDWIS) Community and Non-Community Groundwater Systems in Matagorda County, Texas (Continued)

Table 2.4S.12-4 Listing of U.S. Environmental Protection Agency (EPA) Safe Drinking Water Information System (SDWIS) Community and Non-Community Groundwater Systems in Matagorda County, Texas (Continued)

Transient Non-Community Water Systems: Water Systems that do not consistently serve the same people (e.g. rest stops, campgrounds, gas stations)

Source: Reference 2.4S.12-10

Table 2.4S.12-5 Matagorda County Historical Water Use

Source: Reference 2.4S.12-13

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Table 2.4S.12-6 Matagorda County Projected Water Use

Source: Source 2.4S.12-14

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Table 2.4S.12-9 Regional Aquifer Properties from Aquifer Pumping Tests **Table 2.4S.12-9 Regional Aquifer Properties from Aquifer Pumping Tests**

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Data Source:

1- Reference 2.4S.12-8

2- pumping test data interpretation

3- Reference 2.4S.12-7

ND = Not Determined

Table 2.4S.12-11 STP Slug Test Results

 $P = Poor$ curve match or questionable data

Test Methods:

KGS = Kansas Geological Survey

B-R = Bouwer and Rice

ND = No data - data not recovered from data logger

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Table 2.4S.12-13 Hydraulic Conductivity of Clay

Source: Reference 2.4S.12-9, Section 2.5.4.2.6.1

Table 2.4S.12-14 Representative Properties of Hydrogeologic Units

Table 2.4S.12-15 Regional Hydrogeochemical Data **Table 2.4S.12-15 Regional Hydrogeochemical Data**

BDL = Below analytical detection limit
ND = Not Determined
National Secondary Drinking Water Standard Exceeded
National Primary Drinking Water Standard Exceeded
Source: Reference 2.4S.12-6

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BDL = Below analytical detection limit
ND = Not Determined
National Secondary Drinking Water Standard Exceeded BDL = Below analytical detection limit
ND = Not Determined
National Secondary Drinking Water Standard Exceeded

Table 2.4S.12-17 Estimated Average Linear Velocity and Travel Time

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Figure 2.45.12-3 Physiographic Map of Texas (modified from Reference 2.45.12-1)

Figure 2.45.12-4 Aquifers of Texas (modified from Reference 2.45.12-2)

[1) Present only in the subsurface

[2] Called Catahoula Tuff west of Lavaca County [3] Not recognized at surface east of Live Oak County

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Texas coastal uplands aquifer system

Coastal lowlands aquifer system-Dot patterned area indicates freshwater

Confining unit

Hydrogeologic unit-See figure 2.4S.12-5 D

Figure 2.4S.12-6 Generalized Cross Section through the Coastal Lowlands/Coastal Uplands Aquifer Systems (modified from Reference 2.4S.12-2)

Groundwater

Figure 2.45.12-7 Major Aquifers of Texas (Reference 2.45.12-3)

REGION 6 SOLE SOURCE AQUIFERS

Figure 2.4S.12-8 Sole Source Aquifers in EPA Region VI (Reference 2.4S.12-4)

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Final Safety Analysis Report

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15 Sept 2007

Final Safety Analysis Report

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Figure 2.4S.12-12 Well Locations in Matagorda County and Adjacent areas from the TWDB Database (See Plate I of Appendix 2.4S.12-A for clarity)

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Figure 2.4S.12-13 Coastal Plains Groundwater Conservation District Well Locations (See Plate II of Appendix 2.4S.12-A for clarity)

Groundwater

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Groundwater

Figure 2.4S.12-15 Potentiometric Surface in the Deep Aquifer in Matagorda County in 1967 (modified from Reference
2.4S.12-6)

Figure 2.45.12-16 Regional Potentiometric Surface Map including water level measurements from 2001 to 2005 (Reference 2.45.12-15)

4478

MCR

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November 1, 2005 Upper Shallow Aquifer

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Figure 2.4S.12-19 Quarterly Potentiometric Surface Maps in the STP 3 & 4 Areas

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Figure 2.4S.12-20 Hydrogeologic Cross-Section A-A'

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Figure 2.4S.12-21 Conceptual Hydrogeologic Cross-Section B-B'

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Figure 2.4S.12-22 Regional Hydrographs for Deep Aquifer

Lower Shallow Aquifer

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Figure 2.4S.12-23 Hydrographs of Selected Wells at the STP Site

Upper Shallow Aquifer

Elevation (ftmsl)

Groundwater
Lower Shallow Aquifer Slug Test Hydraulic Conductivity

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1000
Scale in fee

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Contour Interval = 100 gpd/sq. ft.

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OW-930L

Figure 2.4S.12-27 Summary of Hydraulic Conductivity from Aquifer Pumping Tests, Slug Tests, and Grain Size

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Groundwater

Unit	Hydrogeologic Zone	Ground Surface	Thickness	Geologic Materials
Shallow Aquifer	Upper Shallow Aquifer Confining Layer		$10 - 30$ ft	Clay and Silt
	Upper Shallow Aquifer		20 - 30 ft	Silty Sand and Poorly Graded Sand
	Lower Shallow Aquifer Confining Layer		$15 - 25$ ft	Clay and Silt
	Lower Shallow Aquifer		$25 - 50$ ft	Silty Sand and Poorly Graded Sand with thin Clay and Silt Layers
Deep Aquifer Confining Layer			$100 - 150$ ft	Silty Clay and Silt with thin Sand Layers
Deep Aquifer			>500 ft	Sand with thin Clay and Silt Layers

Figure 2.4S.12-29 Simplified Hydrostratigraphic Section

Figure 2.4S.12-30 Trilinear Diagram of Hydrogeochemical Data

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Figure 2.4S.12-31 Blessing SE 7.5 minute Topographic Map (modified from
Reference 2.4S.12-21)

Maximum Hydrostatic Pressure

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