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MOL.20020227.0170

# **SUMMARY DATA REPORT**

**of** 

### **DOWNHOLE VELOCITY SURVEYS**

at the

**Proposed Waste Handling Building** 

 $(ATAC+MENT B)$  empbs/27/01

to accompany

**SCIENTIFIC NOTEBOOK: SN-M&O-SCI-030-V1** 

Date Completed: 10/17/01 emj 10/22/01

# Yucca Mountain Project

ENGINEERING GEOPHYSICS . BLAST DESIGN & PHENOMENOLOGY P.O.BOX 540 . MURPHYS . CALIFORNIA . 95247 209-728-3705

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### **0** Downhole Seismic Velocity Surveys at the Waste Handling Building Site Yucca Mountain Project, Nevada

#### Introduction

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This report presents shear- and compression-wave velocities obtained in 15 boreholes at the site of the proposed Waste Handling Building (WHB) at the Yucca Mountain Project (YMP), located on the Nevada Test Site (NTS). The report also describes the procedures and equipment used to acquire the raw data by means of the standard downhole technique. The downhole velocities, supplemented by corresponding information from Spectral Analysis of Surface Waves and from surveys using the OYO P-S Suspension Logger, are intended to be used in an earthquake-hazard analysis of the WHB site

All procedures and relevant calibration information are completely documented in Scientific Notebook SN-M&O-SCI-030-VI which was compiled as the field work progressed and as the data were interpreted.

The field work for these investigations was performed in two phases because the drilling had not progressed far enough to complete all the planned velocity surveys during the first mobilization to the site. I was assisted by Mr. Mark Dober of URS Corp. (Oakland) during the first phase in October 2000, and by Mr. Chris-Goetz, also of URS Corp. (Santa Ana), during the second phase of field work in November 2000.

The scope of work was as follows:



All of the holes were cased with 4-inch PVC pipe grouted in place. The water inside the casing in each hole, with the exception of RF-13, was pumped out prior to performing the downhole velocity surveys; RF-13 had about 160 ft of water in it at the time of the survey. Although the downhole geophone package is designed to operate under water, the holes were pumped dry to avoid any chance of tubes waves interfering with the shear-wave signals. A tube wave is a pressure pulse that propagates nearly unattenuated down the fluid in a borehole at a velocity close to that of the shear wave in the surrounding material.

#### Procedures and Equipment

We used the conventional downhole survey method in which travel times of signals from an impulsive source of energy at the surface are measured to progressively greater depths in the borehole, The corresponding plot of travel time vs. depth is then converted to velocity vs. depth by computing slopes of the interpreted major straight-line segments of the plotted data.

A vehicle-on-a-beam traction source located close  $(\approx 10 \text{ ft})$  to the collar of the hole was used to generate shear waves. The source consists of a 7-ft-long 6x6 wood beam with cleats on the bottom and steel end caps, the beam being held in firm contact with the ground by driving the front wheels of a vehicle on to it. Horizontal blows to the ends of the beam with a sledge hammer generate shear-wave pulses. Shear-wave signals with 'positive' and 'negative' polarities are generated by striking each end of the beam in sequence. The 'positive' and 'negative" shear waves are, ideally, mirror images of each other and, viewed as a pair, greatly assist in identifying the shear wave. The sledge hammer has an impact sensor attached to the handle near the head that starts the timing process in the recording instrument upon impact with the beam. Where site conditions permitted, the beam was aligned at an azimuth of 340° [true], which I understand is the approximate strike of the bedding.

Compression-wave signals were generated by vertical hammer blows to a steel striker plate on the ground, also located close to the collar of the hole.

The downhole sensor for the velocity surveys was a Model BHG-3 geophone manufactured by Geostuff of Saratoga, California. The sensor package contains an orthogonal array of three geophones, mounted on a gearhead-motor assembly, and a fluxgate compass. A manuallyactivated servo circuit links the compass to the motor for the geophone array so that the horizontal geophones maintain a constant, pre-determined azimuth at each measurement point This allows one of the two horizontal geophone elements to stay aligned with the shear-wave source as the sensor package is raised or lowered in the hole. This capability ensures that the detected shear-wave signal is always the optimum. The BHG-3 is mechanically locked in the hole casing with a motor-actuated clamping spring operated from the control box at the surface.

All signals were recorded with a Geometrics R24 'Strataview' digital seismograph. The R24 was configured to record 24 1024-sample channels of data, sampling at a rate of 125 or 250 microseconds, depending on the hole depth.

In a downhole survey the data are collected channel-by-channel until 24 channels have been recorded or until the survey is complete, at which time the signals are transferred to a hard disk drive and printed in the R24's internal printer for field examination. The raw signals on disk can be filtered and processed on playback and prior to printing. The data are ultimately stored in SEG-2 binary format on 3-1/2-inch disks.

Our procedure was to lower the BHG-3 to the bottom of the hole and expand the clamping spring so that the BHG-3 could be pulled up the hole but would still hold its position. Travel times were then measured from the bottom up at 5-ft intervals to a depth of 100 **ft,** and at **3-ft** intervals from there to the surface. Exceptions to this procedure were holes 14 and 19, in which measurements were made at 10-ft intervals below 480 and 490 **ft,** respectively. In some of the deep holes data were acquired in two phases, first from 400 **ft** up to the surface and then from 400 ft downwards. Because of the weight of the cable attached to the BHG-3, it is difficult for one person to pull more than 400 ft of cable up the hole between each measurement point

Carefui attention was given to the polarity of the shear-wave signals. Prior to beginning a survey, one of the two horizontal geophones inside the BHG-3 was aligned with the shearwave beam. The system was always configured so that at each measurement point the first blow to the beam would produce a shear wave for which the initial deflection of the trace on the seismograph record was upwards, *i.e.,* towards the top of the paper record. At each measurement point the same end of the beam *(e.g.,* on the passenger side of the vehicle) was always struck first and the resulting signal recorded on an odd-numbered channel. The signal from the blow to the other end of the beam was then recorded on the adjacent even-numbered channel. A compression-wave signal was recorded next (while still at the same depth) and a switch box directed that signal to a separate block of channels allocated to compression waves. Because each measurement point consists of two shear-wave signals and one compression-wave signal, a total of **8** measurement points comprise a 24-channel record; channels I through 16 are shear waves and channels 17 through 24 are compression waves. The ability to keep careful track of the polarity of the shear-wave signals proved to be extraordinarily important in the interpretation of the data acquired in these surveys.

It should be noted that the R24 seismograph has the capability of 'stacking' or linearly adding multiple signals. More than one hammer blow for each signal is almost always required at the deeper measurement points in order to increase the signal-to-noise ratio or to clearly develop the shear wave amidst interfering arrivals. As many as 5 blows to each end of the plank were sometimes required.

It was initially thought that some degree of shear-wave anisotropy might be present at the WHB site, *i.e.*, that the shear-wave velocity might be a function of the direction of polarization of the signals. For this reason we performed an experiment in borehole RF-25 in which two orthogonal shear-wave sources (340 $^{\circ}$  and 250 $^{\circ}$ ) were set up at equal distances from the collar of the hole. Prior to starting the survey the 'longitudinal' geophone in the BHG-3 was aligned with the 340°-source and, consequently, the 'transverse' geophone was aligned with the second source. The first blow to each shear-wave beam produced an initial upward deflection of the seismograph trace. At each measurement point shear waves from both

sources were recorded before the BHG-3 was raised to the next measurement point. A special switchbox was used which routed the signals from one shear-wave source to channels 1 through 8, signals from the second source were directed to channels 9 through 16, and the compression-wave signals were sent to channels 21 through 24; channels 17 through 20 were not used. Four measurement points comprised a complete record. In this manner a direct visual comparison of arrival times from the two shear-wave sources could be made at each measurement point. Significant anisotropy was not present; differences in the shear-wave travel times from the two sources at each depth were generally less than one millisecond, and no further experiments of this nature were performed.

#### Results and Analysis

Reduced copies of complete sets of the shear- and compression-wave signals are attached to this report. The source offset and the frequency of the low-pass filters used for playback are shown on each set of records. Although shear- and compression-wave signals are recorded on the same seismograph record in the field, they have been played back and presented here as separate paper records. It should be noted that the travel times that were plotted to obtain velocities were read from full-size copies of the recordings.

As stated previously, the horizontal geophones in the BHG-3 maintained a constant azimuth during each survey and, therefore, the polarity of the recorded shear-wave signal was always consistent with the direction of the hammer blow. A blow to the first end of the beam always resulted in an initial upward deflection of the seismograph trace. As noted on the attached copies of the records, some filtering was applied to all of the signals during playback from the hard disk. The signals acquired in the velocity surveys were transferred to the hard drive without any filtering; filtering was used only during playback to the R24's internal printer. The filtering was accomplished by using the digital filters built into the R24's operating system. The R24 filters used for processing these records were either 100-, 200- or 400-Hz low-pass.

The degree of filtering varied from hole to hole and depended upon the amount of interference present For example, it appears that shear-wave signals arriving at the BHG-3 transducer in portions of RF-15 may have traveled by different paths, i.e., multiple and overlapping signals were present. In this specific instance the least amount of filtering was used (400 Hz) to allow me to track the 'predominant' shear wave from the top of the hole to the bottom. A contrasting example is hole RF-22 where more filtering (100 Hz) was judged necessary because of weak signals near the bottom of the hole and because of noise from nearby drilling operations. Only one filter setting was used for the shear waves in any given borehole, and 400-Hz low-pass filtering was used for all the compression-waves signals.

The time of the onset or 'first break' of the shear wave was not used as the travel time. The onset of the pulse may be reasonably clear at shallow depths, but soon becomes vague or completely obscured as the depth increases. For this reason I prefer to pick the time of the first peak (and of the first trough of the corresponding reverse-polarity blow) as the travel time of the shear wave. In most boreholes at the WHB, even the first peak becomes too ambiguous at depth to time accurately and reliably, and it was necessary to use the maximum

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trough (which immediately follows the first peak) and its corresponding maximum peak as the travel time. The first peak was used only for holes RF-25 and -28. The times picked as travel times are indicated on the attached copies of the shear-wave signals by the small dots. In every case the average. time of a peak and its corresponding trough were used. The time of the first peak of the compression wave was always used as its travel time, but no dots are shown on the attached copies.

Because the energy source, whether compression or shear, is offset horizontally from the collar of the hole, it is necessary to correct the travel times at shallow depths for this offset. The correction is intended to convert the actual travel time along the slant path from source to receiver to the equivalent time required to travel vertically from the surface down to the, receiver. In the past this has always been done by simply multiplying the travel time by the cosine of the angle  $(\Theta)$  between the slant path and vertical. During the analysis of the data acquired in this program it was realized that this method of correcting the shallow travel times is valid only if the onset of a signal is used. If this simple version of the cosine correction is applied to a later portion of the signal, such as the first peak or the maximum peak, then the tines wini be over-corrected and the computed near-surface velocities will be too low. It is also necessary to correct for any time shift of the signals caused by low-pass filtering before making the cosine correction. Although the R24 uses digital filters, they are not 'zero-phase' filters and will cause a small time shift. The proper method of correcting for the offset is as follows:

#### $t_{\text{vertical}} = [t_{\text{peak}} - \Delta t - \Delta t_{\text{o}}] \times \cos\Theta] + \Delta t + \Delta t_{\text{o}}$

where  $t_{vertical}$  is the time that is plotted against the depth to the measurement point,  $t_{break}$  is the time of the selected peak, At is the time difference between the onset of the signal and the selected peak (and corresponding trough) for signals recorded at shallow depths, and  $\Delta t_o$  is the time shift of the onset of the signal due to filtering. This shift is easily measured at shallow depths by comparing the times of onset of unfiltered and filtered signals in each hole. A complete discussion of the filter-delay correction is given on page 109 of the Scientific Notebook. In practice, the value of  $cos\Theta$  is computed from the source offset, s, and the depth to the transducer, d, so that the above equation becomes:

 $t_{\text{vertical}} = [(t_{\text{peak}} - \Delta t - \Delta t_{\text{o}}) \times (d/(d^2 + s^2)^{1/2})] + \Delta t + \Delta t_{\text{o}}$ 

The cosine correction is required to a depth equal to approximately **10** times the offset distance. The cosine correction shown above was applied to the travel times in these surveys until the difference between the raw and the corrected time was only 0.1 millisecond.

Small-scale plots of shear- and compression-wave times *vs.* depth are attached to this report, and a 3-1/2-inch disk with an Excel file of these data accompanies the report. It must be emphasized that the determination of the reported velocities was performed on large-scale hand-drawn plots which are attached to the Scientific Notebook describing these surveys. The data points in these were plotted at a scale of 20 ft to the inch and 10 milliseconds to the inch.

The final values of shear- and compression-wave velocities are shown in the attached tables. Careful examination and comparison of the tabulated values of shear- and compression-wave velocities will reveal that the depths to the interfaces between layers based on the respective velocities are not always coincident, *i.e.,* the shear and compression wave boundaries may not agree. This may be due to the possibility that the shear and compression signals do not always follow the same path from source to receiver.

#### **Discussion**

In general the quality of the raw data was good to excellent. It will be noted on the largescale time *vs.* depth plots, and also on the smaller scale plots attached to this report, that there is some scatter of the data points about their respective trend lines over some portions of some boreholes. The scatter is most probably attributable to the complex geology at the WHB site. The scatter is not due to insufficient signal amplitudes, to noise from extraneous sources, nor to some systemic flaw in the recording instrumentation or procedures. In many cases an examination of the shear-wave records suggests that signals are arriving at the transducer from more than one pathway. This may also be happening in the case of compression waves, but this cannot be discerned by looking at the recorded signals. The degree of scatter is not large in absolute terms, typically a departure of only a millisecond or two from the overall trend line of the data points; in extreme cases the departure was as high as 4 msec. However, because of the relatively high velocities present at the site, even small variations in travel times can appear to be significant.

There is always some judgement involved when assigning travel-time data points to a layer. Often the change of slope of the data points is not especially pronounced when crossing the boundary between one velocity zone and the one immediately below, and the exact depth of the boundary becomes somewhat uncertain due to the inherent scatter in the data. Least squares calculations of slopes were applied in many cases, but it was also necessary to use judgement in many cases where it appeared that scatter precluded meaningful statistical calculations.

Some comments about the data acquired in specific boreholes will illustrate the above comments: The shear-wave signals at a depth of 335 ft in Hole RF-14 show a jump in arrival time of 13 msec, *i.e.,* a shear wave can be followed down to that depth but then simply fades away, and another shear wave at that depth can be seen to begin 13 msec later and to continue on to the bottom. I believe this is due to multiple travel paths that happen to coincide near 335 ft.

The best overall signal quality was observed in Hole RF-19, despite the fact that this was the deepest hole surveyed in this report. Only two hammer blows were required to produce a clear and unambiguous shear wave at a depth of 640 **ft,** and multiple or overlapping shear waves were not evident. The scatter in arrival times was typically less than 1 msec.

There is an apparent low-shear-wave-velocity zone in RF-22 from a depth of 175 ft to 192 **ft,** but there is no corresponding decrease in the compression-wave velocity over this depth

range. This anomaly suggests that the low shear-wave velocity may be an artifact of different travel paths which, in turn, is probably a manifestation of the complex geology associated with the graben in which the borehole is located.

Values of shear-wave velocity below 338 **ft** in RF-29 that result from a literal interpretation of the shear-wave travel times are simply not plausible. Specifically, the shear-wave velocity appears to be about 7000 ft/sec between 338 and 360 **ft,** but the compression-wave velocity is only 6040 ft/sec in this same zone. Despite the fact that the OYO P-S logger survey in this hole showed similar erratic behavior of the shear-wave velocity, the most reasonable approach is to use an 'eyeball' best fit through the data points from 338 **ft** to the bottom of the hole, which results in a continuation of the 3800 ft/sec velocity that starts at a depth of 230 ft.

Finally, the digital format of the raw data allowed some preliminary values of *in-situ* shearwave damping to be computed from spectral ratios of the dowrnhole signals. Although these damping calculations were not part of the original scope of work for the velocity surveys, it seemed appropriate to determine some initial values of damping for subsequent comparison with the corresponding values obtained from laboratory tests on core samples of the rocks. The damping calculations were not Quality Assured, but are completely described on pages 75 through 83 in the Scientific Notebook for these velocity surveys.

Briefly, the calculation of *in-situ* shear-wave damping is based on the spectral decay of the shear-wave signals as they propagate down through the subsurface materials. The damping calculations are concerned only with frictional losses in amplitude as the signal travels downward and not with geometric spreading, and these frictional losses are manifested by a progressive loss of the higher frequencies with increasing distance. By computing a sequence of spectral ratios between the downward traveling shear wave and a reference signal at shallow depth, low-strain (<10"5%) *in-situ* values of the fraction of critical damping can be determined.

I believe that these calculations of damping are meaningful in the sense that they are influenced by the characteristics of the entire rock mass through which the shear wave travels, including porosity and fractures. Laboratory tests on rock cores may be biased towards relatively low values of damping; only the more durable samples of rock that were competent enough to survive the drilling process will remain sufficiently intact to be suitable for laboratory testing.

Shear-wave damping calculations were performed using data acquired in holes RF-16 and -19. Spectral ratios were calculated at 100-ft depth intervals using the signals at a depth of 40 ft as a reference. In RF-16 the computed fraction of critical damping was 6.4% from 40 **ft** to 290 **ft,** and 1.4% below that. In RF-19 the shear damping was calculated to be 5.1 % from 40 ft to 280 t, and 2.3% from 280 **ft** to the bottom at 640 ft. The overall values of damping from a depth of 40 ft to the bottom of each hole, *i.e.*, ignoring any changes in the subsurface velocities, were 4.9% for RF-16 and 3.9% for RF-19. For comparison, low-strain values  $\left($  <10<sup>-3</sup>%) of shear damping obtained in laboratory tests of cores from depths of 81 and 127 ft in RF- 16 were approximately 1% or less.

#### Yucca Mountain Project Final Downhole Velocity Values

### SHEAR WAVES











### **SHEAR WAVES**















### SHEAR WAVES







Note: These tables were corrected for filter delay times in October of 2001.

### Yucca Mountain Project Final Downhole Velocities

# COMPRESSION WAVES











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### COMPRESSION WAVES















# **COMPRESSION WAVES**







Note: These tables were corrected for filter delay times in October 2001.

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Figure 10 Downhole Travel Time RF-23

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Figure 12. Downhole Travel Time RF-25

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YUCCA MOUNTAIN PROJECT Downhole Velocity Records **SHEAR WAVES** 

Borehole RF- 13

Playback Filter **100** Hz Low Pass<br>Time Scale 10 Milliseconds/Line Hz Low Pass Source Offset 9.3 feet








Borehole RF-14

 $200$ Playback Filter 2.00 Hz Low Pass<br>Time Scale 10 Milliseconds/Line<br>Source Offset 10.8 feet Hz Low Pass [All depths in feet]

















Borehole RF-16

Playback Filter  $\frac{200}{10}$  Hz Low Pass<br>Time Scale  $\frac{10}{10.7}$  Milliseconds/Line<br>Source Offset  $\frac{10.7}{10^{10}}$  feet







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Borehole RF-18

Playback Filter 200 Hz Low Pass<br>Time Scale 10 Milliseconds/Line<br>Source Offset 10.3 feet Hz Low Pass [All depths in feet]



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YUCCA MOUNTAIN PROJECT Downhole Velocity Records<br>SHEAR WAVES

Borehole RF-19

Playback Filter <u>400</u> Hz Low Pass<br>Time Scale 10 Milliseconds/Line<br>Source Offset 11.5 feet<br>[All depths in feet] Hz Low Pass



















Borehole RF-22

Playback Filter 100 Hz Low Pass<br>Time Scale 10 Milliseconds/Line<br>Source Offset 10.3 feet [All depths in feel]









Borehole RF-24 Playback Filter  $\frac{100}{10}$  Hz Low Pass<br>Time Scale  $\frac{10}{10}$  Milliseconds/Line<br>Source Offset  $\frac{9.8}{100}$  feet Hz Low Pass












# YUCCA MOUNTAIN PROJECT<br>Downhole Velocity Records<br>SHEAR WAVES

Borehole RF-26

Playback Filter  $\frac{200}{5}$  Hz Low Pass<br>Time Scale  $\frac{5}{5}$  Milliseconds/Line<br>Source Offset  $\frac{10.2}{100}$  feet

















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#### YUCCA MOUNTAIN PROJECT Downhole Velocity Records **COMPRESSION WAVES**

Borehole RF-14

Playback Filter 400<br>Time Scale 5 Millis<br>Source Offset 2.2, feet Hz Low Pass Milliseconds/Line (All depths in teat)





Hz Low Pass Milliseconds/Line





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## YUCCA MOUNTAIN PROJECT Downhole Velocity Records **COMPRESSION WAVES**















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# YUCCA MOUNTAIN PROJECT

Downhole Velocity Records **COMPRESSION WAVES** 

Borehole RF-19

Playback Filter <u>400</u> Hz Low Pass<br>Time Scale 5 Milliseconds/Line<br>Source Offset 11.5 feet<br>[All depths in teet] Hz Low Pass





### YUCCA MOUNTAIN PROJECT Downhole Velocity Records **COMPRESSION WAVES**

Borehole RF-20

Playback Filter 400 Hz Low Pass<br>Time Scale 5 Milliseconds/Line<br>Source Offset 10.3 feet [All depths in teel]












## YUCCA MOUNTAIN PROJECT Downhole Velocity Records **COMPRESSION WAVES**

Borehole RF-22







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## YUCCA MOUNTAIN PROJECT Downhole Velocity Records

**COMPRESSION WAVES** 

Borehole RF-23

Playback Filter 400 - Hz Low Pass TV LOW Pass<br>
5 Milliseconds/Line<br>
7.0 feet<br>
[All depths in feet] Time Scale Source Offset



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## YUCCA MOUNTAIN PROJECT Downhole Velocity Records

COMPRESSION WAVES

Borehole RF-24

Playback Filter 400 Hz Low Pass<br>Time Scale 5 Milliseconds/Line<br>Source Offsat 4.6 Text<br>Midepthe In text











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Borehole RF-26 Playback Filter 400 Hz Low Pass<br>Time Scale 5 Milliseconds/Line<br>Source Offset 10.2 feet<br>[All depths in lead]



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