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Problem Statement:

Preliminary site selection performed jointly by the U.S. Department of Energy (DOE) and the Contractor has identified a 2,300 acre withdrawal area in the Crescent Flat area just northeast of Crescent Junction, Utah, as a possible site for a final disposal cell for the Moab uranium mill tailings. The proposed disposal cell would cover approximately 300 acres. Based on the preliminary site-selection process, the suitability of the Crescent Junction Disposal Site Is being evaluated from several technical aspects, including geomorphic, geologic, hydrologic, seismic, geochemical, and geotechnical. The objective of this calculation set is to present results of the rippability investigation based on seismic refraction activities at the Crescent Junction Disposal Site.

This calculation will be used in the *Remedial Action Plan and Site Design for Stabilization of Moab Title I Uranium Mill Tailings at the Crescent Junction, Utah, Disposal Site* (RAP), and summarized in the appropriate sections of the Remedial Action Selection (RAS) Report for the Moab site.

Method of Solution:

A refraction seismic survey was conducted along 10 seismic lines centered on existing boreholes at the Crescent Junction Site to assist in evaluation of suitability of the site for disposal of the Moab tailings. The purposes of the seismic surveys were to determine the seismic velocities of weathered and unweathered Mancos Shale deposits that underlie the site and relate those velocities to the rippability of the subsurface materials. The refraction seismic method.is routinely used for rippability investigations. Data collection and analysis methods for this project were performed In accordance with the *Standard Guide for Using the Seismic Refraction Method for Subsurface Investigation* ASTM Designation: D 5777-00.

The Final Report of the *Crescent Junction Disposal Site Seismic Rippability Investigation* Is in Appendix A and a review of this report is In Appendix B.

Assumptions:

N/A

Calculation:

N/A

Discussion:

Seismic velocities and the thickness of layers underlying the proposed disposal cell to a depth of approximately 60 feet, produced by means of a refraction seismic survey, have been provided in the report. This Information will be used to determine optimal and economic depths of excavation for construction of the disposal cell at the Crescent Junction Site. The suitability of selecting equipment based on the reported velocities should be based on the excavators experience with ripping machinery where seismic velocities are known. Data in this report will be used in making these determinations during the conceptual design phase of the disposal cell.

Conclusion and Recommendations:

Use of this information in the conceptual design of the disposal cell at the Crescent Junction will be incorporated into appropriate sections of the RAS and other design documentation.

Computer Source:

N/A

References:

See Cover Sheet.

Appendix A

Seismic Rippability Investigation Report

Groundwater, Engineering, Environmental & Mining

Final Report

Seismic.Rippability Investigation

Crescent Junction Disposal Site

for

S. M. Stoller Corporation Grand Junction; Colorado

November 21, 2005['] 4/2/*cs*

.'2473 North Leah Lane *Prescott, Arizona **86301 USA 928ý778-6320** (Telephone and Fax) jim@hasgeo.com (E-Mail) http://www.hasgeo.com (Homepage)

TABLE OF CONTENTS

Table

Figures

Borehole and Seismic Line Locations Map Borehole 202 NW to SE Seismic Line Elevation Section Borehole 202 SW to NE Seismic Line Elevation Section Borehole 204 NW to SE Seismic Line Elevation Section Borehole 204 SW to NE Seismic Line Elevation Section Borehole 206 NW to SE Seismic Line Elevation Section Borehole 206 SW to NE Seismic Line Elevation Section Borehole 207 NW to SE Seismic Line Elevation Section Borehole 207 SW to NE Seismic Line Elevation Section Borehole 208 NW to SE Seismic Line Elevation Section Borehole 208 SW to NE Seismic Line Elevation Section Borehole 202 NW to SE Seismic Line Depth Section Borehole 202 SW to NE Seismic Line Depth Section Borehole 204 NW to SE Seismic Line Depth Section Borehole 204 SW to NE Seismic Line Depth Section Borehole 206 NW to SE Seismic Line Depth Section Borehole 206 SW to NE Seismic Line Depth Section Borehole 207 NW to SE Seismic Line Depth Section Borehole 207 SW to NE Seismic Line Depth Section Borehole 208 NW to SE Seismic Line Depth Section Borehole 208 SW to NE Seismic Line Depth Section Caterpillar D8 Tractor Ripping Chart **Caterpillar D9 Tractor Ripping Chart** Caterpillar D10 Tractor Ripping Chart Caterpillar D11 Tractor Ripping Chart

CD

Velocity Analysis Tables Time-Distance (T-D) Plots Modeled elevations and depths beneath each source point and geophone **Observer Reports**

S. M. Stoller Corporation November 21, 2005

Crescent Junction Seismic Rippability Investigation Page i

INTRODUCTION

Refraction seismic surveys were conducted for **S.** M. Stoller Corporation along ten seismic lines centered on existing boreholes at the proposed Crescent Junction Disposaf Site to assist in the evaluation of the suitability of the site as a final repository for the Moab uranium mill tailings The purposes of the seismic surveys were to determine the seismic velocities of weathered and unweathered Mancos Shale deposits that underlie the site and relate those velocities to the rippability of the subsurface materials.

The refraction seismic method is routinely used for rippability investigations. Caterpillar Inc. has prepared charts that relate seismic velocities to different sized rippers. For typical refraction seismic rippability investigations a seismic.velocity versus depth or elevation profile is generated along each survey line and then the velocities are related to the Caterpillar charts so that a proper ripper can be selected by the construction contractor. Two types of refraction seismic surveys may be conducted to ascertain the rippability estimates of the subsurface: two-dimensional (2D) tomography and delay-time, The 2D tomography method offers a more detailed and gradational section of the subsurface seismic velocity, but takes a little more time in the field and thus is slightly more expensive. The delay-time method offers only a layered and averaged velocity section, but may be more familiar to construction contractors since it has been in use for a much longer time than 2D tomography. Stoller selected the delay-time method for this project.

This seismic survey was a joint effort between Bird Seismic Service, Inc. of Globe, Arizona and Hasbrouck Geophysics, Inc. of Prescott, Arizona. Bird Seismic acquired the seismic data using the survey design prepared by Hasbrouck Geophysics, while I lasbrouck Geophysics processed and interpreted all the data and prepared the final report. This final report will be reviewed by Mr. H. David MacLean of Grand Junction, Colorado. Ken Bernstein is president of Bird Seismic Services, Inc. and may be reached at ken@bridseismic.com or 928-719-1848. Jim Hasbrouck is president of Hasbrouck Geophysics, Inc. and may be contacted at jim@hasgeo.com or 928-778- 6320. Dave MacLean is available at 107770.3066@compuserve.com or 970-242-1649.

DATA ACQUISITION

Seismic surveys essentially consist of recording seismic waves that have been generated by artificial sources, observing the arrival times of these waves, and producing cross-sections of variations in subsurface seismic wave velocities that can then be related to geology. The source of seismic energy for surface surveys is primarily dependent upon the target depths and local geology, and for relatively shallow surveys is generally either a sledgehammer or weight-drop system. The seismic waves are detected by geophones in surface surveys. A geophone consists of a coil suspended by springs with magnets build into the case. A seismic wave moves the case and the magnets while the coil remains relatively stationary because of its inertia. The relative movement of the magnetic field with respect to the coil generates a voltage across the coil that is proportional to the velocity of the seismic wave. The electrical voltages produced by the geophones are transmitted back to a recording instrument (seismograph) via cables. In refraction seismic surveys it is necessary, according to Snell's Law, that velocities increase with depth so that the refracted seismic waves can be detected on the surface. For refraction seismic surveys in

S. M. Stoller Corporation Crescent Junction Seismic Rippability Investigation November 21, 2005 Page I

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most sedimentary environments it is typical that velocities increase with depth (i.e.. there arc no velocity reversals) ard it is assumed that this is the case for the Crescent Junction Disposal Site.

According to Stoller, the depth to weathered bedrock in the project area is assumed to vary from two to approximately 25 feet. Unweathered bedrock may be deeper than 50 feet thus the refraction seismic survey is designed to investigate to depths somewhat greater than 50 feet using the standard rule-of-thumb for refraction seismic surveys that the first geophone to "see" a refraction from a layer will be at a distance of three to five times the expected depth. For example, if an investigation depth of 60 feet is desired then the first geophone to see a refraction, if present, from that depth will be at 180 to 300 feet along the line of geophones, or spread, with the larger distance applicable to areas with generally slower velocities. In order to accurately map the deeper horizons, several geophones must be beyond the initial geophone that records the deeper refraction thus a geophone interval of 10 feet with 30 feet far offsets (resulting in a total spread of 500 feet) is used for this project.

The refraction seismic data for this project were acquired with a 48-channel Bison *9048* seismograph with 21-bits of dynamic range, 250 milliseconds (ms) record lengths, and 0.25 ms sample intervals with Mark Products 10-lz geophones implanted approximately three inches into the ground at intervals of **10** feet along each line. The seismic source was an Elastic Wave Generator (EWG) accelerated weight-drop mounted on the back of a 4x4 pickup and consisted or a 207-pound weight that was lifted hydraulically against large springs and then released resulting in a force much greater than the weight **itself.** For each seismic line data from a minimum of eleven source points were acquired (seven within the spread nominally between geophones 6 and 7, 12 and 13, **18** and **19,** 24 and 25, 30 and **31,** 36 and 37, and 42 and 43. and offeach end at distances of **10** and 30 feet). The geophone distances were initially measured with either a tape or takeout intervals on the geophone spread cable and after completion of data acquisition every fourth geophone and each offset source point was surveyed to at least centimeter accuracy by a contractor to Stoller. Because the surface topography change was minor and the seismic lines were relatively straight it was only necessary to-survey the coordinates and elevations of every fourth geophone and then interpolate values for the intermediate geophones.

The seismic data were stacked nominally four to six times (depending upon offiet and noise) at each source point to increase the signal-to-noise ratio. Stacking, or signal enhancement, involved repeated source impacts at the same point into the same set of geophones. For each source point, the stacked data were recorded into the same seismic data file and theoretically the seismic signal arrived at the same time from each impact and thus was enhanced, while noise was random and tended to be reduced or canceled. After recording the data on the hard disk of the seismograph, the seismic records were copied to a personal computer at the end of each.field day. These data were e-mailed nightly from the field to Mr. Hlasbrouck and copies of the'Observer Reports (field notes) were faxed at the same time. The quality-of the seismic data ranged from very good to excellent depending primarily upon offset, and identifiable first breaks (first arrivals of seismic energy) were present along all the lines.

DATA PROCESSING

The refraction seismic data were processed using the SiPwin (version 2.77) set of computer programs from Rimrock Geophysics Inc., Lakewood, Colorado. The general processing flow consisted of the initial selection, or "picking", of the seismic first breaks (first arrival of seismic energy) with the SIPIK program, creation of data files for input into the interpretation program with the *SllPIN* program. and interpretation of the data using modeling and iterative ray-tracing techniques with the *SIPT2* program. A first break was selected as the initial downward variation of the seismic signal from a horizontal line and was generally accurate to a time between 0.5 and $1¹$ ms. To enhance the accuracy of first break picks, the seismic record was zoomed to a time that only encompassed the breaks themselves (i.e., only the portion of the seismic record where the first breaks were visible) The *SIPT2* program uses the delay-time method to obtain a firstapproximation depth model, which is then trimmed by a series of ray-tracing and model adjustment iterations to minimize any discrepancies between the picked arrival times and corresponding times traced through a 2½-dimensional model. Arrival times at two geophones, separated by some variable XY-distance, are used in refractor velocity analyses and time-depth calculations.:I Using the principle of migration and iterative ray-tracing within the *SIP72* program. forward and reverse seismic rays emerge from essentially the same point on the reflector, thus requiring the reflector to be plane over only a very small distance. The ray-tracing procedure tests and corrects the estimated migrated position of points representing the locations of ray entry and emergence from the refracting horizon and takes into account the dip of the refracting horizon at those emergence points, therefore enabling accurate representation of steeply dipping horizons.

For any refraction seismic data analysis, it is important to determine accurate velocities. The *SII'72* program employs several routines for selection of the proper velocities. For the direct arrivals through the first layer, the velocity is computed by dividing the distances from each source point to each geophone by the corresponding arrival times. These individual velocities are averaged for each source point and a weighted average is computed. For layers beneath the first layer, velocities are computed by two methods:' **1)** Regression, in which a straight line is fit **by** least squares to the arrival times representing the velocity layer and average velocities are computed by taking the reciprocals of the weighted average of the slopes of the regression lines. and 2) the ltobson-Overton method wherein velocities are computed if there are reciprocal arrivals from two opposing source points at two or more geophones. Final velocities used in the *SIPT2* inversion process are computed by taking an average of the two methods. As quality control measures, time versus distance (T-D) plots (which represent velocities) are inspected along each seismic line relative to reciprocal times, irregularity and parallelism as per ASTM D5777. The refraction seismic data for this project adequately met the requirements of each of these tests.

Included within this report are a borehole and seismic line location map. and elevation and depth versus distance refraction seismic sections for each line with annotated average velocities for each layer. Also included is a CD with output from the **SIP72** program that includes velocity analysis tables, T-D plots indicating the picked arrival times, and modeled elevations and depths beneath each source point and geophone. Note that the distances in the modeled results have been corrected for horizontal foreshortening (i.e., corrections are made to obtain true horizontal

positions). The modeled results arc used to construct the elevation and depth sections. and are in Microsoft Excel format for future client use if desired.

RESULTS

According to Stoller. the geology of the project area consists of essentially three layers. The near surface is alluvial overburden composed of unconsolidated silt, clay, and sandstone fragments. Beneath the alluvium is weathered bedrock, or weathered Mancos Shale, composed of fractured. chemically weathered, siltstone, silty sandstone or clayey siltstone of variable thickness. The weathered layer is often highly fractured with calcite and gypsum fracture coatings. The Mancos Shale is present beneath the weathered layer and is increasingly competent with depth. Although the Mancos Shale appears to be a great shale mass, it is not one homogeneous unit. According to available lithologic logs for the borcholes within the project area, the Mancos Shale scenes to be described as consisting not only of shale but also some sandstone layers and what has been termed a silty claystone. The lithologic logs generally indicate variations in the composition of the unweathered Mancos Shale near its top with increasing shale constituents with depth.

Interpretation of the refraction seismic data indicates three layers, representing alluvial overburden, weathered Mancos Shale and competent Mancos Shale. Table **I** indicates the range in velocities and depths for each line. The first layer velocities range from about 1160 to *1330* feet per second and are consistent with typical unsaturated alluvial overburden values. The second layer velocities range from about 4060 to 5220 feet per second and represent typical values for weathered material such as the Mancos Shale. The variation in 'velocity values for the interpreted weathered Mancos Shale is probably related to the degree of fracturing and the amount of calcite and/or gypsum coating of the fractures. The higher velocities may have less fracturing or the fractures may be coated with an increased amount .of calcite and/or gy'psum. It is not possible from the seismic results to determine which scenario exists. The third layer, or competent Mancos Shale bedrock, Velocities range from about 9000 to as high as 10000 feet per second with the majority of the velocity values in a range from about 9000 to 9400 feet per second. Velocity variations of the interpreted Mancos Shale bedrock are considered relatively minor and probably related to slight changes in composition of the bedrock or some amount of fracturing. Velocity variations are present along the intersecting seismic lines at each borehole, but generally than about *5%* which is reasonable given that the velocity values are averages and the subsurface geology is variable (as evidenced by changes in the lithologic logs between boreholes).

The thickness of the overburden layer (or the depth to the top of layer 2 which is interpreted as weathered Mancos Shale) ranges from about 4¹/₂ to 18 feet, while the depth to the top of layer 3 (or interpreted unweathered Mancos Shale bedrock) varies from about 24 to 60 feet. The tie point depths between intersecting lines at each borehole are generally less than about 5% which is considered reasonable and quite acceptable for seismic surveys. Depth values at intersecting points from lines oriented in different directions often vary because of anisotropy within the subsurface geological formations. Anisotropy is defined as a variation of a physical property (e.g', velocity) depending upon the direction in which it is measured. In general, surface refraction seismic data have shown a 10% to 15% variation between the actual depths to velocity

layer anomalies, as Verified primarily by geophysical borehole logging, and the depth predicted by the models.

Table 1. Buildikn't be interpreted verbenies and depths					
Line	Layer 1 velocity	Layer 2 velocity	Layer 3 velocity	Depth to top of	Depth to top of
	f(t/s)	f(k/s)	(Ms)	layer $2(f)$	layer $3(f)$
202NW-SE	1230	4218	10005.	$4.5 - 15.1$	$34.3 - 58.7$
202SW-NE	1305	4305	9353	$11.3 - 17.1$	$31.5 - 53.0$
204NW-SE	1334	4674	9035	$8.0 - 14.9$	$40.4 - 61.1$
204SW-NE	1206	4705	9399	$10.1 - 18.1$	$40.5 - 59.3$
206NW-SE	1305	5221	9380	$9.9 - 16.0$	$29.5 - 46.8$
206SW-NE	1281	5169	9479	$7.7 - 15.2$	$25.4 - 47.5$
207NW-SE	1159	4195	9011	$7.1 - 14.3$	$26.7 - 49.1$
207SW-NE	1228	4061	9021	$5.9 - 15.0$	$28.9 - 45.9$
208NW-SE	1260	4430	- 9676	$11.0 - 14.5$	$33.3 - 48.8$
208SW-NE	1191	4633	9805	$9.0 - 13.6$	$23.4 - 48.4$

Table 1 : Summary of interpreted velocities and depths

Inspection **of** either the elevation or depth sections indicates that the subsurface is far from planar. with some areas showing signs of possible incised bedrock channels (e.g., particularly possibly both lines at borehole 208). Because both the first and last approximately 30 to 50 feet, or more, of the sections have less forward and reverse raypath coverage (refer to the T-D plots), results in those areas should be viewed with some caution. Nevertheless, subsurface depth variations are present along each of the seismic lines.

According to Caterpillar's ripping charts, shale is considered rippable at seismic velocities ranging up to about 6000 to 10200 feet per second for tractor models D8 to Di1, respectively. Rippable velocities are slightly different if the subsurface material is composed more of a siltstone (up to about 6500 to 9900 for a D8 to D11 tractor, respectively). Referencing the ripping charts from Caterpillar, it is reasonable to assume that all of the interpreted layer 2 or weathered Mancos Shale can be ripped with a tractor as small as a D8 (note that the Caterpillar ripping charts are not available for tractors smaller than a D8). If it is necessary to rip the interpreted competent Mancos Shale bedrock, with velocities interpreted to **be** greater than 9000 feet per second, it will be necessary to employ a D₁₁ tractor.

Although the seismic survey covered only a very small portion of the proposed Crescent Junction Disposal Site it is reasonable to assume that excavation in the proposed site will be impacted by the variable weathered and unweathered bedrock depths. Although the author of this report is not aware of the design depth of the, proposed disposal site excavation, if it is say 40 feet then there will be areas encountered with much higher velocity material at depth Which will require either larger rippers or other means of excavation. For example, if material is ripped along the borehole 207 SW to **NE** seismic line to a depth of 40 feet materials with average velocities of around 4000 and 9000 feet per second will both be encountered. Obviously, a **D8** tractor would not be able to rip to a depth of 40 feet along the entire length of this line.

LIMITATIONS OF **INVESTIGATION**

Although a refraction seismic investigation is the most cost-effective way to determine rippability of material in a project area (versus sporadic boreholes that offer only localized information), it must be realized that according to Caterpillar ripping is still more art than science. and much will depend upon operator skill and experience. Caterpillar states in their I landbook that tooth penetration is often the key to ripping success, regardless of seismic velocity. Low seismic velocities in sedimentary rocks can indicate probable rippability. However, if the fractures and bedding joints do not allow tooth penetration then the material may not be ripped effectively. Pre-blasting or "popping" may induce sufficient fracturing to permit tooth entry.

This survey was conducted with state-of-the-art instrumentation operated by experienced geophysicists, the data were processed by an experienced and licensed geophysicist with a commercial software package utilized on projects with similar objectives, and the results were interpreted by an experienced and licensed geophysicist. However, no warranty, either expressed or implied, is made as to the usability of the results of this survey. Additionally, the ripper' performance charts developed by Caterpillar are intended for estimating purposes only and neither Caterpillar Inc. nor Hasbrouck Geophysics. Inc. warrant that the tractors will perform as estimated.

REFERENCE

Caterpillar Performance Handbook, Edition 30. October 1999. Use *of Seismic Velocit, Charts,* pp. 1-71 to 1-78.

S. M. Stoller Corporation November 21. 2005

Crescent Junction Seismic Rippability Investigation Page 6

Seismic Rippability Investigation **Crescent Junction Disposal Site Borehole and Seismic Line Locations**

Hasbrouck Geophysics, Inc.

Borehole&LineLocations.grf

Seismic Rippability Investigation

Hasbrouck Geophysics, Inc.

Borehole202NSElev2D grf

Seismic Rippability Investigation Crescent Junction Disposal Site Borehole 204 NW to SE Seismic Line Elevation Section

Hasbrouck Geophysics, Inc.

Seismic Rippability Investigation Crescent Junction Disposal Site,' Borehole 204 SW to **NE** Seismic Line *Elevation Section*

Hasbrouck Geophysics, Inc. 5orehde204EWElev2D grf

Borehole206NSElev2D.at

Seismic Rippability Investigation **Crescent Junction Disposal Site** Borehole 206 SW to NE Seismic Line **Elevation Section**

Hasbrouck Geophysics, Inc.

Borehole206EWElev2D.grf

Borehole207NSElev2D grf

Borehole207EWElev2D.grf

Seismic Rippability Investigation Crescent Junction Disposal Site Borehole 208 NW to SE Seismic Line **Elevation Section**

Hasbrouck Geophysics, Inc.

Seismic Rippability Investigation **Crescent Junction Disposal Site Borehole 208 SW to NE Seismic Line Elevation Section**

Hasbrouck Geophysics, Inc.

Borehole208EWElev2D grf

Hasbrouck Geophysics, Inc. **Box 2022 Box 2022EWDepth2D gift**

Borehole204EWDepth2D grf

Borehole206NSDepth2D.gr

Hasbrouck Geophysics, Inc. **Borehole206EWDepth2D.grf**

Borehole207EWDepth2D grf

Hasbrouck Geophysics, Inc. **Boxes and Contact Contact**

Caterpillar D8 Ripping Chart

D8R

Hasbrouck Geophysics, Inc.

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Caterpillar D9 Ripping Chart

Hasbrouck Geophysics, Inc.

Caterpillar D10 Ripping Chart

D10R

- . Multi or Single Shank No. 10 Ripper
- · Estimated by Seismic Wave Velocities

Hasbrouck Geophysics, Inc.
Caterpillar D11 Ripping Chart

D11R

. Multi or Single Shank No. 11 Ripper

Hasbrouck Geophysics, Inc.

Appendix B

Review of Seismic Rippability Investigation Report

The Word Works Inc., dba

H. David Mac Lean, P. Geoph

November 23, 2005

Mr. Mark Kautsky S.M. Stoller Corp. 2597 B % Road Grand Junction CO 81503

Subject: Review of Report; Seismic Rippability Investigations at the Crescent Junction Disposal Site, for S.M. Stoller Corp. Grand Junction CO by Hasbrouck Geophysics.

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202 North Ave., PMB 252 Grand Junction Co 81501 Phone/fax: 970-242-1649

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H. David Mac Lean, **R** Geoph.

Review of Report "Crescent Junction November 23, 2005 Disposal Site Rippability Investigations" by Ilasbrouck Geophysics

In its simplest form, the delay time method involves measurement of the time of arrival ofa seismic wave at two geophone locations separated by a distance D. The description of the method and the-procedures employed to accomplish the measurements as set forth in the above reference report are in accordance with standard industry practice. It is a limitation of the method that in order to measure the seismic velocity of successively deeper units, the seismic velocity must increase sequentially with depth, as stated in the report. This is usually the case in the most frequently encountered field situations, but low velocity, or reversal situations are encountered on occasion. Low velocity reversals do not appear to occur at the Crescent Junction site.

The lithologic section and the depth of investigation were specified by Stoller. The near surface section was determined by careful logging of core and cuttings from boreholes located throughout the planned repository. Selected borcholes formed the centers of the seismic refraction spreads as shown in the Borehole and Seismic Line Location Map included with the subject report. The refraction surveys **were** intended to extend the layer thickness information for approximately 250 **fi** in four orthogonal directions from the borehole.

Field Data Acquisition

The equipment referenced in the subject report was inspected during a visit to the field operations on October 29. The equipment was found to **be** as specified, and to be in good working order. Field conditions were less thanoptimal. Heavy rains had turned the area into a quagmire; nevertheless, the field crew was able to bring equipment into the area and proceeded with the survey with only minimal interruption caused by the adverse road and access conditions.

Field work for the survey was conducted October 29 and 30,2005. It was observed that all field activities were conducted in a professional and workmanlike manner. Prior to commencement of operations, the field crew.was briefed on health and safety issues by a Stoller representative, and a Health and Safety Plan was provided to the crew. The briefing was attended by this reviewer. Requirements of the plan, including clothing specifications were carefully observed by all field personnel.. In accordance with plan requirements, any soil that became even slightly contaminated in the course of the field activities was removed from the site and disposed of in accordance with applicable procedures and regulations.

This reviewer participated in one day of the field operations and noted that they were conducted as described in the report. Field work was conduct by a crew provided by Bird Geophysics in accordance with survey design specifications developed by Hlasbrouck Geophysics. The crew was obviously well trained and performed all assigned tasks with competence and in a professional manner.

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H. David Mac Lean, P. Geoph.

Review of Report "Crescent Junction November 23, 2005 Disposal Site Rippability Investigations" by Hasbrouck Geophysics

Upon locating the center of the refraction spreads referenced in an expanded version of the "Borehole and Seismic Line Location Map" included in the report, two orthogonal survey lines were extended in a NE-SW direction and a NW-SE direction. The lines were run by chain and compass; markers were placed at 10 foot intervals for a distance of 270 ft in the four directions from the center point. Every fourth marker from the center was identified so that its position could be surveyed later to the required accuracy. Since the terrain was open and unencumbered by vegetation, the entire line could be viewed from vantage points along the line. Lines were visually determined to be straight along the length of the chained interval.

Geophones were placed at each 10 foot marker. All geophones in a linear string were connected to the Bison 48 channel seismograph with Mark Products geophone cables.

Seismic signals were generated by the accelerated weight drop hammer mentioned in the report. A 200 **lb** metal bar is raised against compressing springs, and is thus accelerated downward to strike an aluminum plate placed on the ground at the shot point. The hammering operation started at one end of the line, **and** continued at various points along .the spread asstated in the report. This is standard "slooting" procedure for seismic refraction surveys. The multiple shot points allow numerous depth and velocity determinations at various points along the spreads, and permit aiveraging and compensation for anisotropy and dip, since the seismic ray path can be observed in opposite directions. This procedure enables production of a much more detailed and represcntati **ve** velocity-depth section than would be possible if only.a single shot point was employed.

Several (up to six) "shots" or hammer blows were taken at each shot point, allowing the seismic signals at each geophone location to be stacked. This procedure increases the signal to noise ratio. As pointed out in the report, seismic waves that arrive at a geophone at the same time following a hammer blow are additive to the signal; random noise or seismic signals for which'the sirike instant is incorrect arc destructive and will not augment or enhance the initial seismic signal. On completion of the stacking activity, a seismogram was printed in the field for inspection and quality assurance purposes.

On the completion of the field survey day, digital data sets were forwarded to I lasbrouck Geophysics for processing and analysis. The data were processed by Hasbrouck Geophysics using the SIPwin software from Rimrock Geophysics. This processing software is state-of-the-technology for Refraction Seismic Data Processing. Given the software capabilities and the field procedures employed, Hasbrouck Geophysics was able to calculate seismic velocities over very short refractor distances. Velocities were calculated using both the regression and Hobson-Overton methods. This processing combination adequately deals with the effects anisotropy, and the distortions introduced by dipping layers. The resulting depth and velocity calculations were then employed to produce the very detailed velocity/depth sections included with the subject report.

H. David Mac Lean, **R** Geoph.

Review of Report "Crescent Junction November 23, 2005 Disposal Site Rippability Investigations" by Hasbrouck Geophysics

Analvsis

The velocities for the 3 layers discussed in the report, i.e., alluvium (layer 1) weathered Mancos Shale (layer 2) and Mancos Shale (layer 3) are well within thc range expected for these materials. In unconsolidated material such as Layer I, seismic velocities are often close to acoustic velocities in air (approximately **II** 00fps). Considerable variation in the measured velocity of Layer 2, (the weathered shale or regolith) can be cxpecied depending on the amount of sand or silt inclusioning and the degree of consolidation within local areas. As pointed out in the report, weathering will not be complete through the entire geologic section and the lithologic material is not uniform. As expected, the seismic velocities increase as a function of depth.

The velocity function for all three layers underlying the planned repository is well illustrated by the time-distance (T-D) plot for one of the survey lines at borehole 204. A copy of the T-D plot is attached hereto. The figure provides a visual indication of the seismic velocity for the various layers. Generally, the flatter the curve on the T-D plot, the higher the velocity. A segment of the T-D plot that is continuous over a measurable interval indicates an identifiable layer. A simple estimate of the velocity associated with (this interval can be made by dividing the distance interval D by the difference in arrival time (T) on the **"-D** plot. Of course the actual final determination of the depth associated with this interval involves a considerably more complex calculation, as has been discussed peripherally in the report..

Limitations

The purpose of measuring the seismic velocities of the layers underlying the proposed mill tailings repository.was to estimate the rippability of the underlying lithologic units. **I** lasbrouck Geophysics has developed depth and velocity sections for all of the surveyed lines that show the lithologic layers to depths of 50 to 60 feet and the measured seismic velocities within these layers to the accuracy that is achievable with the equipment and methodology employed. **I** lowever, the relationship of these measured seismic velocities to rippability of a particular unit is empirical, not an engineering certainty. Caterpillar. Inc. and others involved with heavy equipment operations have observed an apparent relationship and have published charts and graphs showing the ripping capabilities of certain tractor models for various geologic material with a range of seismic velocities.- **I lowever, there are many other factors that contribute to rippability, such as the degree** and orientation of.fracturing. Although the rippability charts published by Caterpillar Inc. reprsent that material with a seismic velocity in a certain range is usually within the ripping capability of certain tractor types, it is not an engineering certainty that this is the case.

> 202 North Ave.. PMB 252 Page 4of **7** Grand Junction, CO 81501. Phone/fax: 970-242-1649

H. David Mac Lean, P. Geoph..

Review of Report "Crescent Junction November 23, 2005 Disposal Site Rippability Investigations" by Hasbrouck Geophysics

Accordingly, any decision to employ a certain type of equipment based on the velocities provided in the subject report must be taken on the basis of the excavation contractor's own knowledge, and not on statements or implied statements in the report. The velocities and layer thicknesses provided in the report are valid within the accuracy of the seismic refraction method, and are reproducible by similar surveys. Nevertheless, the relationship of these in-situ measured velocities and the suitability of a specific tractor model for ripping a geologic unit with these velocities is strictly empirical and may vary from that presented in the rippability charts provided by Caterpillar Inc.

Conclusions

The subject report provides seismic velocities and the thickness of layers underlying the proposed tailings repository to a depth of about 60 **ft** or more. **The** sections showing these depths and velocities provided in the report were produced by means of a refraction seismic survey that was conducted in a professional and workmanlike manner, employing equipment that was suitable to the task. The measured interval velocities, unit thicknesses and variations to be expected are accurate to within the limitations of the current state of refraction seismic technology. The statement in the report that measured velocities arc accurate to within **10** per cent is probably overly pessimistic; the accuracy of the measurements is probably much closer to 5% or less. General experience suggests that the unit thicknesses stated in the report are accurate to within 10% or better.

As stated in the report, the suitability of selecting equipment based on the reported velocities is based entirely onthe experience of Caterpillar Inc. Nothing in the subject report should be construed as an endorsement of the suitability of a particular tractor model for ripping and excavating applications at the Crescent Junction repository. This decision must be taken on the basis of the excavator's own experience with ripping machinery in applications where seismic velocities are known.

Respectfully Submitted,

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11. David Mac Lean, P. Geoph.

11 DM/hdm

Enclosures:

Borehole 204 NW to SE Time-Distance plot

202 North Ave., PMB 252 Page 5of 7 Grand Junction, CO 8 *1* **501b.** 1hone'fax: 970.-242-1649

1. David Mac Lean, P. Geoph.

Review of Report "Crescent Junction November 23, 2005 H. **David**

Disposal Site Rippability Investigations" by Hashrouck Geophysics

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November 23, 2005

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H. David Mac Lean, **R** Geoph.

Rcview of Report "Crescent Junction Disposal Site Rippability Investigations" by Hasbrouck Geophysics

November 23, 2005

Statement of Qualifications

11. David Mac Lean is a Registered Professional Geophysicist in the State of Californa, Registration No. 440 and in the Province of Alberta, Canada, Registration no.M 15724. Mr. **Mac** Lean has been a practicing geophysicist for over 35 years.

.Mr. Mac Lean has gained experience with the seismic refraction method while engaged in aggregate mapping activities in the Beaufort Sea, and in laying out seismic surveys for oil exploration in Alberta. Canada.

Mr. Mac Lean is an emeritus member of the Society of Exploration Geophysicists, the Canadian Exploration Geophysical Society, The Australian Society of Exploration Gcophsicist, the Society for Mining and Metallurgy of the American Institute Mining Engineering and other technical and professional societies dedicated to the advancement of geophysics. Mr. MacLean is a frequent attendee at conventions, trade shows and seminars dedicated to geophysical technologies.

202 North Ave., PMB 252 Page 7of 7 Grand Junction, CO 8150). PHionelfax: 970-242-1649

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Problem Statement:

Preliminary site selection performed jointly by the U.S. Department of Energy (DOE) and the Contractor has identified a 2,300 acre withdrawal area in the Crescent Flat area just northeast of Crescent Junction, Utah, as a possible site for a final disposal cell for the Moab uranium mill tailings. The proposed disposal cell would cover approximately 300 acres. Based on the preliminary site-selection process, the suitability of the Crescent Junction Disposal Site Is being evaluated from several technical aspects, including geomorphic, geologic, hydrologic, seismic, geochemical, and geotechnical. The objective of this calculation set is to present results of the rippability investigation based on seismic refraction activities at the Crescent Junction Disposal Site.

This calculation will be used in the *Remedial Action Plan and Site Design for Stabilization of Moab Title I Uranium Mill Tailings at the Crescent Junction, Utah, Disposal Site* (RAP), and summarized in the appropriate sections of the Remedial Action Selection (RAS) Report for the Moab site.

Method of Solution:

A refraction seismic survey was conducted along 10 seismic lines centered on existing boreholes at the Crescent Junction Site to assist In evaluation of suitability of the site for disposal of the Moab tailings. The purposes of the seismic surveys were to determine the seismic velocities of weathered and unweathered Mancos Shale deposits that underlie the site and relate those velocities to the rippability of the subsurface materials. The refraction seismic method is routinely used for rippability Investigations. Data collection and analysis methods for this project were performed in accordance with the *Standard Guide for Using the Seismic Refraction Method for Subsurface Investigation* ASTM Designation: D 5777-00.

The Final Report of the *Crescent Junction Disposal Site Seismic Rippabli/ty Investigation* Is in Appendix A and a review of this report Is In Appendix B.

Assumptions:

N/A

Calculation:

N/A

Discussion:

Seismic velocities and the thickness of layers underlying the proposed disposal cell to a depth of approximately 60 feet, produced by means of a refraction seismic survey, have been provided in the report. This information will be used to determine optimal and economic depths of excavation for construction of the disposal cell at the Crescent Junction Site. The suitability of selecting equipment based on the reported velocities should be based on the excavators experience with ripping machinery where seismic velocities are known. Data In this report will be used In making these determinations during the conceptual design phase of the disposal cell.

Conclusion and Recommendations:

Use of this Information in the conceptual design of the disposal cell at the Crescent Junction will be Incorporated Into appropriate sections of the RAS and other design documentation.

2 Computer Source:

N/A

References:

See Cover Sheet.

U.S. Department of Energy
May 2006

Seismic Rippability Investigation
Doc. No. X0173000
Page 3

Appendix A

Seismic Rippability Investigation Report

X7• Hasbrouck Geophysics, Inc. *Groundwater, Engineering, Environmental & Mining*

Final Report

Seismic Rippability Investigation

Crescent Junction Disposal Site.

for

S. M. Stoller Corporation Grand Junction; Colorado

November 21, **2005**

2473 North Leah Lane Prescott, Arizona **86301 USA 928-778-6320** (Telephone and Fax) j im@hasgeo.com (E-Mail). http://www.hasgeo.com (Homepage)

TABLE OF CONTENTS

Table

-5

Figures

Borehole and Seismic Line Locations Map Borehole 202 NW to SE Seismic Line Elevation Section Borehole 202 SW to NE Seismic Line Elevation Section Borehole 204 NW to SE Seismic Line Elevation Section Borehole 204 SW to NE Seismic Line Elevation Section Borehole 206 NW to SE Seismic Line Elevation Section Borehole 206 SW to NE Seismic Line Elevation Section Borehole 207 NW to SE Seismic Line Elevation Section Borehole 207 SW to NE Seismic Line Elevation Section Borehole 208 NW to SE Seismic Line Elevation Section Borehole 208 SW to NE Seismic Line Elevation Section Borehole 202 NW to SE Seismic Line Depth Section Borehole 202 SW to NE Seismic Line Depth Section Borehole 204 NW to SE Seismic Line Depth Section Borehole 204 SW to NE Seismic Line Depth Section Borehole 206 NW to SE Seismic Line Depth Section Borehole 206 SW to NE Seismic Line Depth Section Borehole 207 NW to SE Scismic Line Depth Section Borehole 207 SW to NE Seismic Line Depth Section Borehole 208 NW to SE Seismic Line Depth Section Borehole 208 SW to NE Seismic Line Depth Section **Caterpillar D8 Tractor Ripping Chart Caterpillar D9 Tractor Ripping Chart** Caterpillar D10 Tractor Ripping Chart Caterpillar D11 Tractor Ripping Chart

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Velocity Analysis Tables Time-Distance (T-D) Plots Modeled elevations and depths beneath each source point and geophone **Observer Reports**

S. M. Stoller Corporation November 21, 2005

Crescent Junction Seismic Rippability Investigation Page i

INTRODUCTION

Refraction seismic surveys were conducted for *S. M. Stoller Corporation along ten seismic lines* centered on existing borcholes at the proposed Crescent Junction Disposaf Site to assist in the evaluation of the suitability of the site as a final repository for the Moab uranium mill tailings The purposes of the seismic surveys were to determine the seismic velocities of weathered and unweathered Mancos Shale deposits that underlie the site and relate those velocities to the rippability of the subsurface materials.

The refraction seismic method is routinely used for rippaliility investigations. Caterpillar Inc. has prepared charts that relate seismic velocities to different sized rippers. For typical refraction seismic rippability investigations a seismic velocity versus depth or elevation profile is generated along each survey line and.then the velocities are related to the Caterpillar charts so that a proper ripper can be selected by the construction contractor. Two types of refraction seismic surveys may be conducted to ascertain the rippability estimates of the subsurface: two-dimensional (2D) tomography and delay-time. The 2D tomography method offers a more detailed and gradational section of the subsurface seismic velocity, but takes a little more time in the field and thus is slightly more expensive. The delay-time method offers only a layered and averaged velocity section, but may be more familiar to construction contractors since it has been in use for a much longer time than 2D tomography. Stoller selected the delay-time method for this project.

This seismic survey was a joint effort between Bird Seismic Service, Inc. of Globe, Arizona and Hasbrouck Geophysics, Inc. of Prescott, Arizona Bird Seismic acquired the seismic data using the survey design prepared by Hasbrouck Geophysics, while **I** lasbrouck Geophysics processed and interpreted all the data and prepared the final report. This final report will be reviewed by Mr. H. David MacLean of Grand Junction, Colorado. Ken Bernstein is president of Bird Seismic Services, Inc. and may be reached at ken@bridseismic.com or 928-719-1848. Jim Hasbrouck is president of Hasbrouck Geophysics, Inc. and may be contacted at jim@hasgeo.com or 928-778-6320. Dave MacLean is available at 107770.3066@compuserve.com or 970-242-1649.

DATA ACQUISITION

Seismic surveys essentially consist of recording seismic waves that have been generated by artificial sources, observing the arrival times of these waves, and producing cross-sections of variations in subsurface seismic wave velocities that can then be related to'geology. The source of seismic energy for surface surveys is primarily dependent upon the target depths and local geology, and for relatively shallow surveys is generally either a sledgehammer or weight-drop system. The seismic waves are detected by geophones in surface surveys. A geophone consists of a coil suspended by springs with magnets build into the case. A seismic wave moves the case and the magnets while the coil remains relatively stationary because of its inertia. The relative movement of the magnetic field with respect to the coil generates a voltage across the coil that is proportional to the velocity of the seismic wave. The electrical voltages produced by the geophones are transmitted back to a recording instrument (seismograph) via cables. In refraction seismic surveys it is necessary, according to Snell's Law, that velocities increase with depth so that the refracted seismic waves can be detected on the surface. For refraction seismic surveys in

S. M. Stoller Corporation Crescent Junction Seismic Rippability Investigation **November 21, 2005** Page I

most sedimentary environments it is typical that velocities increase with depth (i.e., there are no velocity reversals) and it is assumed that this is the case for the Crescent Junction l)isposal Site.

According to Stoller, the depth to weathered bedrock in the project area iý assumed to vary firom two to approximately 25 feet. Unweathered bedrock may be deeper than 50 feet thus the refraction seismic survey is designed to investigate to depths somewhat greater than 50 feet using the standard rule-of-thumb for refraction seismic surveys that the first geophone to "see" a refraction from a layer will be at a distance of three to five times the expected depth. For example, if an investigation depth of 60 feet is desired then the first geophone to see a refraction, if present, from that depth will be at 180 to 300 feet along the line of geophones, or spread, with the larger distance applicable to areas with generally slower velocities. In order to accurately map the deeper horizons, several geophones must be beyond the initial geophone that records the deeper refraction thus a geophone interval of 10 **fiet** with 30 feet far offsets (resulting in a total spread of 500 feet) is used for this project.

The refraction seismic data for this project were acquired with a 48-channel Bison *9048* seismograph with 21 -bits **of** dynamic range, 250 milliseconds (ins) record lengths, and 0.25 mns sample intervails with Mark Products **10-1** lz geophones implanted approximately three inches into the ground at intervals of 10 feet along each line. The seismic source was an Elastic Wave Generator (EWG) accelerated weight-drop mounted on the back of a 4x4 pickup and consisted of a 207-pound weight that was lifted hydraulically against large springs and then released resulting in a force much greater than the weight itself. For each seismic line data from a minimum of eleven source points were acquired (seven within the spread nominally between geophones 6 and 7, 12 and 13,'18 and **19,** 24 and 25, 30 and 31, 36 and 37, and 42 and 43, and ollfeach end at distances of 10 and 30 feet). The geophone distances were initially measured with either a tape or takeout intervals on the geophone spread cable and after completion of data acquisition every fourth geophone and each offset source point was surveyed to at least centimeter accuracy by a contractor to Stoller. Because the surtace topography change was minor and the seismic lines were relatively straight it was orily necessary to survey the coordinates and elevations of every fourth geophone and then interpolate values for the intermediate geophones.

The seismic data were stacked nominally four to six times (depending upon offset and noise) at each source point to increase the signal-to-noise ratio. Stacking, or signal enhancement, involved repeated source impacts at the same point into the same set of geophones. For each source point. the stacked data were recorded into the same seismic data file and theoretically the seismic signal arrived at the same time from each impact and thus was enhanced, while noise was random and tended to be reduced or canceled. Afler recording the data on the hard disk of the seismograph, the seismic records were copied to a personal computer at the end of each field day. These data were e-mailed nightly from the field to Mr. Hasbrouck and copies of the Observer Reports (field notes) Were faxed at the same time. The quality of the seismic data ranged from very good to excellent depending primarily upon offset, and identifiable first breaks (first arrivals of seismic energy) were present along all the lines.

S. M. Stoller Corporation Crescent Junction Seismic Rippability Investigation **the November 21, 2005 1 Page 2 1 Page 2 1 Page 2 1 Page 2 1 Page 2 Page 2 1 Page 2 Pag**

DATA PROCESSING

The refraction seismic data were processed using the *SIPwin* (version 2.77) set of computer programs from Rimrock Geophysics Inc., Lakewood, Colorado. The general processing flow consisted of the initial selection, or "picking", of the seismic first breaks (first arrival of seismic energy) with the *SIPIK* program, creation of data files for input into the interpretation program with the *SIPIN* program, and interpretation of the data using modeling and iterative ray-tracing techniques with the *SIP72* program. A first break was selected as the initial dowmward variation of the seismic signal from a horizontal line and was generally accurate to a time between 0.5 and $1²$ ms. To enhance the accuracy of first break picks, the seismic record was zoomed to a time that only encompassed the breaks themselves (i.e., only the portion of the seismic record where the first breaks were visible) The SIPT2-program uses the delay-time method to obtain a firstapproximation depth model, which is then trimmed by a series of ray-tracing and model adjustment iterations to minimize any discrepancies between the picked arrival times and corresponding times traced through a 2½-dimensional model. Arrival times at two geophones, separated by some variable XY-distance, are used in refractor velocity analyses and time-depth calculations. Using the principle of migration and iterative ray-tracing within the SIPT2 program. forward and reverse seismic rays emerge from essentially the same point on the reflector, thus requiring the reflector to be plane over only a very small distance. The ray-tracing procedure tests and corrects the estimated migrated position of points representing the locations of ray entry and emergence from the refracting horizon and takes into account the dip of the refracting horizon at those emergence points, therefore enabling accurate representation of steeply dipping horizons.

-For any refraction 'sisniiic data analysis, it is important to determine accurate velocities. -The *SIPT2* program employs several routines for selection of the proper velocities. For the direct arrivals through the first layer, the velocity is computed by dividing the distances from each source point to each geophone by the corresponding arrival times. These individual velocities are averaged for each source point and a weightcd average is computed. For layers beneath the first layer, velocities are computed by two methods: **1)** Regression, in which a straight line is fit by least squares to the arrival times representing the velocity layer and average velocities are computed by taking the reciprocals of the weighted average of the slopes of the regression lines. and 2) the Hobson-Overton method wherein velocities are computed if there are reciprocal arrivals from two opposing source points at two or more geophones. Final velocities used in the *SM172* inversion process are computed by taking an average of the two methods. As quality control measures, time versus distance (T-D) plots (which represent velocities) are inspected along each seismic line relative to reciprocal times, irregularity and parallelism as per ASTM 1)5777. The refraction seismic data fbr this project adequately met **tile** requirements ofeach of these tests.

Included within this report are a borchole and seismic line location map, and elevation and depth versus distance refraction seismic sections for each line with annotated average velocities for each layer. Also included is a **CD)** with output from the *S8I'72* program that includes velocity analysis tables, T-D plots indicating the picked arrival times, and modeled elevations and depths beneath each source point and geophone. Note that the distances in the modeled results have been .corrected for horizontal foreshortening (i.e., corrections are made to obtain true horizontal

S. M. Stoller Corporation Crescent Junction Seismic Rippability Investigation t November 21, 2005 **Page 3**

positions). The modeled results are used to construct the elevation and depth sections, and are in Microsoft-Excel format for future client use if desired.

RESULTS

According to Stoller, tihe geology of the project area consists of essentially three layers. The near surlace is alluvial overburden composed of unconsolidated silt, clay, and sandstone fragments. Beneath the alluvium is weathered bedrock, or weathered Mancos Shale, composed of fractured, chemically weathered, siltstone, silty sandstone or clayey siltstone of variable thickness. The weathered layer is often highly fractured with calcite and gypsum fracture coatings. The Mancos Shale is present beneath the weathered layer and is increasingly competent with depth. Although the Mancos Shale appears to be a great shale mass, it is not one homogeneous unit. According to available lithologic logs for the boreholes within the project area, the Mancos Shale scents to be described as consisting not only of shale but also some sandstone layers and what has been termed a silty claystone. The lithologic logs generally indicate variations in the composition of the unweathered Mancos Shale near its top with increasing shale constituents with depth.

Interpretation of the refraction seismic data indicates three layers, representing alluvial overburden. weathered Mancos Shale and competent Mancos Shale. Table **I** indicates the range in velocities and depths for each line. The first layer velocities range from about 1160 to 1330 feet per second and are consistent with typical unsaturated alluvial overburden values. The second layer velocities range from about 4060 to 5220 feet per second and represent typical values for weathered material such as thc Mancos Shale. The variation in velocity values for the interpreted weathered Mancos Shale is probably related to the degree of fracturing and the amount of calcite and/or gypsum coating of the fractures. The higher velocities may have less fracturing or the fractures may be coated with an increased amount of calcite and/or gypsum. It is not possible from the seismic results to determine which scenario exists. The third layer, or competent Mancos Shale bedrock,, velocities range from about 9000 to as high as 10000 feet per second with the majority of the velocity values in a range from about 9000 to 9400 feet per second. Velocity variations of the interpreted Mancos Shale bedrock are considered relatively minor and probably related to slight changes in composition of the bedrock or some amount of fracturing. Velocity variations are present along the intersecting seismic lines at each borehole, but generally than about 5% which is reasonable given that the velocity values are averages and the subsurface geology is variable (as evidenced by changes in the lithologic logs between boreholes).

The thickness of the overburden layer (or the depth to the top of layer 2 which is interpreted as weathered Mancos Shale) ranges from about **4Y2** to 18 feet, while the depth to the top of layer 3 (or interpreted unweathered Mancos Shale bedrock) varies from about 24 to 60 feet. The tie point depths between intersecting lines at each borehole are generally less than about 5% which is considered reasonable and quite acceptable for seismic surveys. Depth values at intersecting points from lines oriented in different directions oflen vary because of anisotropy within the subsurface geological formations. Anisotropy is defined as a Variation of a physical property (e.g., velocity) depending upon the direction in which it is measured. In general, surface refraction seismic data have shown a **10%** to 15% variation between the actual depths **to** velocity

S. M. Stoller Corporation Crescent Junction Seismic Rippability Investigation) November 21, 2005 Page 4

layer anomalies, as verified primarily by geophysical borehole logging, and the depth predicted by the models.

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Line	Layer 1 velocity (h/s)	Layer 2 velocity (ft/s)	Layer 3 velocity (Ns)	Depth to top of layer $2(f)$	Depth to top of layer $3(f)$
202NW-SE	1230	4218	10005	$4.5 - 15.1$	$34.3 - 58.7$
202SW-NE	1305	4305	9353	$11.3 - 17.1$	$31.5 - 53.0$
204NW-SE	1334	4674	9035	$8.0 - 14.9$	$40.4 - 61.1$
204SW-NE.	1206	4705	9399	$10.1 - 18.1$	$40.5 - 59.3$
206NW-SE	1305	5221	9380	$9.9 - 16.0$	$29.5 - 46.8$
206SW-NE	1281	5169	9479	$7.7 - 15.2$	$25.4 - 47.5$
207NW-SE	1159	$4195 -$	9011.	$7.1 - 14.3$	$26.7 - 49.1$
207SW-NE	1228	4061	9021	$5.9 - 15.0$	$28.9 - 45.9$
208NW-SE	1260	4430	9676	$11.0 - 14.5$	$33.3 - 48.8$
208SW-NE	1191	4633	9805	$9.0 - 13.6$	$23.4 - 48.4$

Table 1: Summary of interpreted velocities and denths

Inspection of either the elevation or depth sections indicates that the subsurface is far from planar, with some areas showing signs of possible incised bedrock channels (e.g., particularly possibly both lines at borehole 208). Because both the first and last approximately 30 to 50 feet, or more, of the sections have less forward and reverse raypath coverage (refer to the T-D plots), results in those areas should be viewed with some caution. Nevertheless, subsurface depth variations are present along each of the seismic lines.

According to Caterpillar's ripping charts, shale is considered rippable at seismic velocities ranging up to about 6000 to 10200 feet per second for tractor models D8 to D11, respectively. Rippable velocities are slightly different if the subsurface material is composed more of a siltstone (up to about 6500 to 9900 for a D8 to D11 tractor, respectively). Referencing the ripping charts from Caterpillar, it is reasonable to assume that all of the interpreted layer 2 or weathered Mancos Shale can be ripped with a tractor as small as a D8 (note that the Caterpillar ripping charts are not available for tractors smaller than a D8). If it is necessary to rip the interpreted competent Mancos Shale bedrock, with velocities interpreted to be greater than 9000 feet per second, it will be necessary to employ a D11 tractor.

Although the seismic survey covered only a very small portion of the proposed Crescent Junction Disposal Site it is reasonable to assume that excavation in the proposed site will be impacted by the variable weathered and unweathered bedrock depths. Although the author of this report is not aware of the design depth of the proposed disposal site excavation, if it is say 40 feet then there will be areas encountered with much higher velocity material at depth which will require either larger rippers or other means of excavation. For example, if material is ripped along the borehole 207 SW to NE seismic line to a depth of 40 feet materials with average velocities of around 4000 and 9000 feet per second will both be encountered. Obviously, a D8 tractor would not be able to rip to a depth of 40 feet along the entire length of this line.

S. M. Stoller Corporation November 21, 2005

Crescent Junction Seismic Rippability Investigation Page 5

LIMITATIONS OF INVESTIGATION

Although a refraction seismic investigation is the most cost-eflictive way to determine rippability of material in a project area (versus sporadic' boreholes that offer only localized information), it must be realized that according to Caterpillar ripping is still more art than science, and much will depend upon operator skill and experience. Caterpillar states in their Handbook that tooth penetration is often the key to ripping success, regardless of seismic velocity. Low seismic. velocities in sedimentary rocks can indicate probable rippability. Ilowever, if the fractures and bedding joints do not allow tooth penetration then the material may not be ripped effectively. Pre-blasting or "popping" may induce sufficient fracturing to permit tooth entry.

This survey was conducted with state-of-the-art instrumentation operated by experienced geophysicists, the data were processed by an experienced and licensed geophysicist with a commercial software package utilized on projects with similar objectives, and the results were interpreted by an experienced and licensed geophysicist. However, no warranty, either expressed or implied, is made as to the usability of the results of this survey. Additionally, the ripper* performance charts developed by Caterpillar are intended for estimating purposes only and neither Caterpillar Inc. nor Hasbrouck Geophysics, Inc. warrant that the tractors will perform as estimated.

REFERENCE

Caterpillar Performance I landbook, Edition **30,** October 1999, *Use qf.Seismic Velocity ('liarts,* pp. 1-71 to 1-78.

S. **M.** Stoller Corporation November 21. 2005

Crescent Junction Seismic Rippability Investigation Page 6

Seismic Rippability Investigation **Crescent Junction Disposal Site Borehole and Seismic Line Locations**

Hasbrouck Geophysics, Inc.

Borehole&LineLocations.grf

Seismic Rippability Investigation

Hasbrouck Geophysics, Inc.

Borehole202NSElev2D.grf

Seismic Rippability Investigation **Crescent Junction Disposal Site**

Hasbrouck Geophysics, Inc.

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Seismic Rippability Investigation **Crescent Junction Disposal Site** Borehole 204 NW to **SE** Seismic **Line** *Elevation Section*

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Seismic Rippability Investigation **Crescent Junction Disposal Site** Borehole **204** SW to NE Seismic Line *Elevation. Section*

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Seismic Rippability Investigation **Crescent Junction Disposal Site Borehole 206 NW to SE Seismic Line.** *Elevation Section*

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Seismic Rippability Investigation Crescent Junction Disposal Site Borehole 206 SW to NE Seismic Line Elevation Section

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Seismic Rippability Investigation Crescent Junction Disposal Site Borehole 207 SW to NE Seismic Line

Hasbrouck Geophysics, Inc.

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Seismic Rippability Investigation Crescent Junction Disposal Site" Borehole **208** NW to SE **Seismic** Line

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*Elevation. Section**

Seismic Rippability Investigation Crescent Junction Disposal Site Borehole 208 SW to NE Seismic Line **Elevation Section**

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Caterpillar D8 Ripping Chart

Hasbrouck Geophysics, Inc.

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Caterpillar D9 Ripping Chart

Hasbrouck Geophysics, Inc.

Caterpillar D10 Ripping Chart

D10R

. Multi or Single Shank No. 10 Ripper.

· Estimated by Selsmic Wave Velocities 3 **Seismic Velocity** Meters Per Second × 1000 $12⁷$ $13₁₃$ $10 - 11$ $\frac{15}{25}$ $\mathbf{1}$ $\overline{\mathbf{c}}$ $\mathbf{3}$ $\ddot{4}$ 5 6 $\overline{\mathbf{z}}$ $\pmb{8}$ 9 14 Feet Per Second × 1000 **TOPSOIL CLAY GLACIAL TILL IGNEOUS ROCKS GRANTE BASALT TRAP ROCK** SEDIMENTARY ROCKS **SHALE** <u>.</u> **SANDSTONE SILTSTONE CLAYSTONE CONGLOMERATE BRECCIA CALICHE** LIMESTONE METAMORPHIC ROCKS **SCIRST SLATE MINERALS & ORES COAL IRON ORE** NON-RIPPABLE **ANYWW** MARGINAL | **RIPPABLE**

Hasbrouck Geophysics, Inc.

Caterpillar D11 Ripping Chart

D11R

- . Multi or Single Shank No. 11 Ripper
- **. Estimated by Seismic Wave Velocities**

Hasbrouck Geophysics, Inc.

Appendix B

Review of Seismic Rippability Investigation Report

The Word Works Inc.. dba

H. David Mac Lean, P. Geoph

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Subject: Review of Report; *Seismic Rippaibility Investigations at the Crescent Junction Disposal Site. at the Crescent Junction Disposal Site. for* S.M. Stoller Corp. Grand Junction CO
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QW 202 North Ave., PMB 252 Grand Junction Co 81501 Phonelfax :970-242-1649

H. David Mac Lean, P. Geoph. November *23.* ²⁰⁰⁵

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Review of Report "Crescent Junction Disposal Site Rippability Investigations" by Hasbrouck Geophysics

In its simplest form, the delay time method involves measurement of the time of arrival of a seismic wave at two geophone locations separated by a distance D. The description of the method and the procedures employed to accomplish the measurements as set forth in the above reference report are in accordance with standard industry practice. It is a limitation of the method that in order to measure the seismic velocity of successively deeper units, the seismic velocity must increase sequentially with depth, as stated in the report. This is usually the case in the most frequently encountered field situations, but low velocity, or reversal situations are encountered on occasion. Low velocity reversals do not appear to occur at the Crescent Junction site.

The lithologic section and the depth of investigation were specified by Stoller. The near surface section was determined by careful logging of core and cuttings from boreholes located throughout the planned repository. Selected borcholes formed the centers of the seismic refraction spreads as shown in the Borehole and Seismic Line Location Map included with the subject report. The refraction surveys were intended to extend the layer thickness information for approximately 250 **ft** in four orthogonal directions from the borehole.

Field Data Acquisition

The equipment referenced in the subject report was inspected during a visit to the field operations on October 29. The equipment was found to be as specified, and to be in good working order. Field conditions were less than optimal. Heavy rains had turned the area into a quagmire; nevertheless, the field crew was able to bring equipment into the area and proceeded with the survey with only minimal interruption caused by the adverse road and access conditions.

Field work for the survey was conducted October 29 and 30, 2005. It was observed that all field activities were conducted in a professional and workmanlike manner. Prior to commencement of operations, the field crew was briefed on health and safety issues by a Stoller representative, and a Health and Safety Plan was provided to the crew. The briefing was attended by this reviewer. Requirements of the plan, including clothing specifications were carefully observed by all field personnel.. In accordance with plan requirements, any soil that became even slightly contaminated in the course of the field activities was removed from the site and disposed of in accordance with applicable procedures and regulations.

This reviewer participated in one day of the field operations and noted that they were conducted as described in the report. Field work was conduct by a crew provided by Bird Geophysics in accordance with survey design specifications developed by Hasbrouck Geophysics. The crew was obviously well trained and performed all assigned tasks with competence and in a professional manner.

> 202 North Ave., PMB 252 \mathbb{R}^2 Page 2of 7 Grand Junction, CO 81501: Phone/fax. 970-242-1649

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Review of Report "Crescent Junction November 23, 2005 Disposal Site Rippability lnvestigations" by Hasbrouck Geophysics

Upon locating the center of thc refraction spreads refcrenced in an expanded version of the "Borehole and Seismic Line Location Map" included in the report. two orthogonal survey lines were extcnded in a NE-SW direction and a NW-SE direction. The lines werc run by chain and compass; markers were placed at 10 foot intervals for a distance of 270 ft in the four directions from the center point. Every fourth marker from the center was identified so that its position could be surveyed later to the required accuracy. Since the terrain was open and unencumbered by vegetation, the entire line could be viewed from vantage points along the line.. Lines were visually determined to be straight along the length of the chained interval.

Gcophones were placed at each **10** foot marker. All geophones in a linear string were connected to the Bison 48 channel seismograph with Mark Products geophone cables.

Seismic signals were generated by the accelerated weight drop hammer mentioned in the report. A 200 **lb** metal bar is raised against compressing springs, and is thus accelerated downward to strike an aluminum plate placed on the ground at the shot point. The hammering operation started at one end of the line, and continued at various points along the spread as stated in the report. This is standard "shooting" procedure for seismic refraction surveys. The multiple shot points allow numerous depth and velocity dctcrminations at various points along the spreads, and permit averaging and compensation for anisotropy and dip, since the seismic ray path can be observed in opposite directions. This procedure enables production of a much more detailed and representative velocity-depth section than would be possible if only a single shot point was employed.

Several (up to six) "shots" or hammer blows were taken at each shot point, allowing the seismic signals at each geophone location to be stacked. This procedure increases the signal to noise ratio. -As pointed out in the report, seismic waves that arrive at a geophone at the same time following a hammer blow are additive to the signal; random noise or seismic signals for which the strike instant is'incorrect are destructive and will not augment or enhance.the initial seismic signal. On completion of the stacking activity, a **se** ismogram was'printed in the field for inspection and quality assurance'purposes.

On the completion of the field survey day, digital data sets Were forwarded to Ilasbrouck Geophysics for processing and analysis. The data were processed by Hasbrouck Geophysics using the SIP win software from Rimrock Geophysics. This processing. software is state-of-the-technology for'Refraction Seismic Data Processing. Given the software capabilities and the field procedures employed, Hasbrouck Geophysics was able to calculate seismic velocities over very short refractor distances. Velocities were calculated using both the regression and Hobson-Overton methods. This processing combination adequately deals with the effects anisotropy, and the distortions introduced by.dipping layers. The resulting depth and velocity calculations were then employed to produce the very detailed velocity/depth sections included with the subject report.

> 202 North Ave., PMB 252. Page 3of 7 Grand Junction, **CO** 8150 **1.** Phone/fax: 970-242-1649

.H. David Mac Lean, P. Geoph.
Review of Report "Crescent Junction November 23, 2005

Disposal Site Rippability Investigations" by Hasbrouck Geophysics

Analysis

The velocities for the 3 layers discussed in the'report. i.e., alluvium (layer 1) weathered Maneos Shale (layer 2) and Mancos Shale (layer 3) are well within the range expected for these materials. In unconsolidated material such as Layer i, seismic velocities are often close to acoustic velocities in air (approximately *II* 00fps). Considerable variation in the measured velocity of Layer 2, (the weathered shale or rcgolith) can be expected depending on the amount of sand or silt inclusioning and the degree of consolidation within local areas. As pointed out in the report, weathering will not be complete through the entire geologic section and the lithologic material is not uniform. As expected, the seismic velocities increase as a function of depth.

The velocity function for all three layers underlying the planned repository is well illustrated by'the time-distance (T-D) plot for one of the survey lines at borehole 204. A copy of the T-D plot is attached hereto. The figure provides a visual indication of the seismic velocity for the various layers. Generally, the flatter the curve on the T-D plot, the higher the velocity. A segment of the T-D plot that is continuous over a measurable interval indicates an identifiable layer. A simple estimate of the velocity associated with this interval can be made **by** dividing the distance interval D by the difference in arrival time (T) on the T-D plot. Of course the actual final determination of the depth associated with this interval involves a considerably more complex calculation, as has been discussed peripherally in the report.

Limitations

The purpose of measuring the seismic velocities of the layers underlying the proposed mill tailings repository was to estimate the rippability of the underlying lithologic units. **I** lasbrouck Geophysics has developed depth and velocity sections for all of the surveyed lines that show the lithologic layers to depths of 50 to 60 feet and the measured seismic velocities within these layers to the accuracy that is achievable with the equipment and methodology employed. However, the relationship of these measured seismic velocities to rippability of a particular unit is empirical, not an engineering certainty. Caterpillar Inc. and others involved with heavy equipment operations have observed an apparent relationship and have published charts and graphs showing the ripping capabilities of certain tractor models for various geologic material with a range of seismic velocities. However, there are many other factors that contribute to rippability, such as the degree and orientation of fracturing. Although the rippability charts published by Caterpillar Inc. represent that material with a seismic velocity in a certain range is usually within the ripping capability of certain tractor types, it is not an engineering certainty that this is the case.

> 202_ North Avc_ PMB 252 Page 4of **7** Grand Junction, CO 81501. Phone/fax: 970-242-1649

H. David Mac Lean, P.Geoph..

Review of Report "Crescent Junction November 23, 2005 Disposal Site Rippability Investigations" by Hasbrouck Geophysics

Accordingly, any decision to employ a certain type of equipment based on the velocities provided in the subject report must be taken on the basis of the excavation contractor's own knowledge, and not on statements or implied statements in the report. The velocities and layer thicknesses provided in the report are valid within the accuracy of the seismic refraction method, and are reproducible by similar surveys. Nevertheless, the relationship of these in-situ measured velocities and the suitability of a specific tractor model for ripping a geologic unit with these velocities is strictly empirical and may vany from that presented in the rippability charts provided by Caterpillar Inc.

Conclusions

The subject report provides seismic velocities and the thickness of layers underlying the proposed tailings repository to a depth of about 60 **fR** or more. The sections showing these depths and velocities provided in the report were produced by means of a refraction .seismic survey that was conducted in a professional and workmanlike manner, employing equipment that was suitable to the task. The measured interval velocities, unit thicknesses and variations to be expected are accurate to within the limitations of the current state of refraction seismic technology. The statement in the report that measured velocities arc accurate to within 10 per cent is probably overly pessimistic; the accuracy of the measuremeats is probably much closer to 5% or less. General experience suggests that the unit thicknesses stated in the report are accurate to within **10%** or better.

As stated in the report, the suitability of selecting equipment based on the reported velocities is based entirely on the experience of Caterpillar Inc. Nothing in the subject report should be construed as an endorsement of the suitability of a particular tractor model for ripping and excavating applications at the Crescent Junction repository. This decision must be taken on the basis of the excavator's own experience with ripping machinery in applications where seismic velocities are known.

Respectfully Submitted,

I1. David Mac Lean, P. Geoph.

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Enclosures:

Borehole 204 NW to SE Time-Distance plot

202 North Ave., PMB 252 Page *5of 7* Grand Junction, CO 81501, Phone/fax: 970-242-1649

R. David **Mac** Lean, P. Geoph.

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. Caterpillar Performance Handbook, Ed 30, October 1999. *Use of Seismic Velocity Charts* pp 1-71 - 1-78

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Review of Report "Crescent Junction Disposal Site Rippability Investigations" by Hasbrouck Geophysics

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202 North Ave. PMB 252 Grand Junction, CO 81501. Phone/fax: 970-242-1649

H. David Mac Lean, P. Geoph.

Review of Report "Crescent Junction Disposal Site Rippability Investigations" by Hasbrouck Geophysics

November 23, 2005

Statement of Qualifications

I1 David Mac Lean is a Registered Professional Geophysicist in the State of Cal iforna, Registration No. 440 and in the Province of Alberta, Canada, Registration no.M 15724. Mr. Mac Lean has been a practicing geophysicist for over 35 years.

Mr. Mac Lean has gained experience with the seismic refraction method while engaged in aggregate mapping activities in the Beaufort Sea, and in laying out seismic surveys for oil exploration in Alberta. Canada.

,Mr. MactLan is an emeritus member of the Society of Exploration Geophysicists. the Canadian Exploration Geophysical Society, The Australian Society of Exploration Gcophsicist, the Society for Mining and Metallurgy of the American Institute Mining Engineering and other technical and professional societies dedicated to the advancement of geophysics. Mr. MacLean is a frequent attendee at conventions, trade shows and seminars dedicated to geophysical technologies.

202 North Ave., PMB 252 Page 7of 7 Grand Jimction, CO 81501. Phone/fax: 970-242-1649

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Contents

Appendix A Material Placement in the Disposal Cell

Appendix B Geochemical Attenuation and Performance Assessment Modeling

U.S. Department of Energy-Grand Junction, Colorado

Crescent Junction Disposal Site Geochemical Characterization of Mancos Shale and Reaction-Transport Modeling of Contaminant Attenuation

Executive Summary

The purpose of this work was to develop a robust one-dimensional reaction-transport model that could be used by site personnel to help evaluate the attenuation of contaminants in ground water that may migrate from a proposed uranium mill tailings disposal cell at Crescent Junction, Utah. Although it is unlikely that contaminated water from the disposal cell will penetrate the thick Mancos Shale, contaminant attenuation by bedrock affords increased protection for deep ground water systems. Data for the model were derived from laboratory characterization of the bedrock beneath the proposed disposal cell. The disposal cell will contain tailings from the Moab, Utah, (Atlas) uranium mill and is underlain by Mancos Shale Formation. Samples of Mancos Shale were collected from 10 cores that were evenly distributed in the area of the proposed disposal cell. Five samples were collected from each of the 10 cores for a total of 50 samples, at depths of 40 to 300 ft.

The chemistry of the water-soluble fractions of the samples was determined and used to help estimate the proportions and types of water-soluble minerals. Based on these results it is estimated that the watersoluble mineral assemblage is dominated by nahcolite (NaHCO₃) with major amounts of Ca-Na exchange and gypsum (CaSO₄'2H₂O). Halite, sylvite, and dolomite occur in lesser amounts and calcite is present. Gypsum and dolomite are more dominant in the shallow samples, while nahcolite and halite are more concentrated in the deeper samples. The mineral abundances were 'used as initial conditions for the ground-water reaction-transport model.

Cation exchange capacity (CEC) was measured on 20 core samples of Mancos Shale. **CEC** ranged from 0.54 to 36.29 meg/100 g with a mean of 11.23 meg/100 g. These values are consistent with shale composed of kaolinite, illite, and smectitic interlayered clay minerals as determined by x-ray diffraction analysis. In the reaction-transport modeling, the mean CEC value is used to specify the number of cation exchange sites. Cation exchange was used to model retardation of NH₄.

X-ray diffraction (XRD) analysis on 10 samples of Mancos Shale was used to'further identify minerals present in the core samples. The bulk mineralogy is dominated by quartz, with lesser amounts of dolomite and calcite, small amounts of feldspar, and traces of gypsum. The clay mineral fraction is dominated by mixed-layer (mostly illite/smectite) clays, illite, and kaolinite, with illite layers dominating the mixed-layer clays. The mineralogy was used to help estimate initial mineral composition for the reaction-transport model.

Particle surface area was determined by BET analysis on 10 Mancos Shale samples. Surface areas range from 8.81 to 13.22 m²/g with a mean of 11.02 m²/g. Surface area measured on a powdered split is consistent with surface areas measured on the **1-** to 2-mm fractions, indicating that the BET method is probably accounting for intraparticle surface area. Surface area was used as an input to the surfacecomplexation adsorption algorithm in the reaction-transport model.

Distribution ratios (R_d) are a measure of the partitioning of a contaminant between the ground water and the solid fraction of the aquifer. The higher the R_d, the more partitioning to the solids and the more retardation of the contaminant. The R_d value can vary significantly with solution chemistry. For example, Davis et al. (2004) showed that R_d values for uranium in a sample of alluvium varied by more than a factor of 10 depending on dissolved carbonate concentration and pH value. For the reaction-transport modeling, a surface-complexation approach is used instead of using constant R_d values to simulate retardation of U. Using the surface-complexation approach, the R_d value changes as chemical conditions (especially pH and

 $pCO₂$) change in the aquifer. The results of the R_d determinations were complicated by the analytical imprecision associated with relatively small concentration differences between tests containing sediment and controls without sediment. However, the data indicate that R_d values for uranium adsorption on Mancos Shale are relatively low, ranging from essentially 0 to 0.84 mL/g. The main use of the R_d values was to test and calibrate the surface-complexation module used for U transport in the reaction-transport modeling.

Two sequential batch-leaching tests were conducted to observe chemical evolution of tailings leachate as it interacted with Mancos Shale. One test simulated vertical transport and the other horizontal transport through the Mancos Shale. Chemical changes that occurred in the tailings leachate as it reacted with progressively more Mancos Shale in the horizontal simulation include slightly increasing values of pH, Cl, K, Na, S04, and U. Similar to the simulation of horizontal transport, pH values in the vertical simulation progressively increased with increased reaction. Progressive increases also occurred in the Cl, Na, SO₄, and U concentrations, but decreasing trends in alkalinity and possibly NH₄ occurred. The results of the sequential batch tests were used to help calibrate the reaction-transport models.

The laboratory data discussed above were used to develop a one dimensional reaction-transport model of tailings fluid interaction with the Mancos Shale to simulate transport of contaminants beneath the Crescent Junction disposal cell. Two simulations were run, one simulation representing horizontal flow through shallow Mancos Shale and the other representing vertically downward flow through progressively deeper Mancos Shale. The only difference in input parameters between the horizontal and vertical simulations is the initial abundances of minerals, the deeper Mancos Shale having more nahcolite and less gypsum than the shallow Mancos Shale. The simulations were calculated with the transport algorithm embedded in the PHREEQC geochemical computer program.

In both the horizontal and vertical transport simulations, NH4 is retarded for about 3.5 pore volumes, after which concentrations rise rapidly to the influent concentration. Retardation of NH4 was caused by the exchange of dissolved NH4 cations for solid-phase exchangeable Ca, K, Mg, and Na. Most of the exchange involves Na ion because Na dominates the cations in Mancos Shale ground water at the Crescent Junction site. Effluent U concentrations in the horizontal simulation are retarded for slightly less than one pore volume during which the U concentration is less than about 0.075 mg/L. After one pore volume, U concentrations increase rapidly and reach the influent value (4 mg/L) after about 3 pore volumes. In the vertical simulation, U concentrations are never less than 2.4 mg/L; however, U concentrations remain less than the influent until about 4 pore volumes. In the horizontal simulation, K_d values were initially 0.26 mL/g and decreased to 0.19 mL/g after about 2 pore volumes. In the vertical simulation, K_d values were negligible initially and increased to 0.15 mL/g after about 4 pore volumes. The differences in K_d values and dissolved U concentrations between the two simulations result from variation in the ionic composition of the solutions resulting from equilibration with minerals and cation exchange sites. A critical factor was the presence of a higher concentration of gypsum in the horizontal simulation than in the vertical simulation. Calcium released from dissolution of gypsum in the horizontal simulation caused precipitation of calcite, which resulted in decreased pH. These conditions led to increased partitioning of U to the specific adsorption sites on the Mancos Shale and increased retardation in the horizontal simulation.

In summary, the transport results suggest that $NH₄$ migration is retarded by several pore volumes. Uranium is retarded by about **1** pore volume, but only if Ca is released from gypsum dissolution. To evaluate the effects of this analysis on contaminant transport beneath the proposed Crescent Junction disposal cell, it is necessary to know the flux of contaminated water from the cell. The effective porosity of the Mancos Shale must also be known; attenuation would be much less in a system dominated by fracture flow than one dominated **by** porous media flow. Evaluation of the flow regime was beyond the scope of this investigation; thus, results of the model simulations are provided in terms of pore Volumes. To maximize the benefit of these results in the field setting, project personnel will need to couple these results with results of hydrologic investigations to convert units of pore volume to more useful units of travel time and distance. Alternatively, a sensitivity analysis that uses reasonable bounds for the hydrologic parameters may be appropriate to assess the impact of chemical attenuation at the Crescent Junction site.

Much of the work presented in this paper on U adsorption uses an approach recently developed by Dr. James Davis and co-workers at the U.S. Geological Survey. Davis et al. were funded by the U.S. Nuclear
Regulatory Commission specifically to advance the state-of-the-art in geochemical methods used for the transport of U at uranium milling sites. Their model of surface complexation adsorption is well suited to the work presented here.

Introduction

This work was conducted to characterize the geochemistry of the bedrock beneath and adjacent to the proposed disposal cell location.at Crescent Junction, Utah. The disposal cell will be constructed to contain tailings from the former Moab, Utah, (Atlas) uranium-ore processing mill and Will be underlain by the Mancos Shale Formation. The purpose of this work is to provide data to help evaluate the potential for ground water contamination and transport of tailings constituents at the disposal site. The basis for this work is provided in Section 4.5 of the "Work Plan for Characterization of Crescent Junction Disposal Site" (Appendix A) and was modified based on discussions with Moab Uranium Mill Tailings Remedial Action (UMTRA) Project personnel. The scope, which includes laboratory investigations and reaction-transport modeling, is defined in Appendix A.

The geochemical approach involves collecting site-specific data and using thete data to model geochemical interactions between tailings pore fluid and the Mancos Shale. Results of the laboratory investigations provide the following information about the Mancos Shale: (1) abundance and mineralogy of water-soluble minerals; (2) mineralogy of water-insoluble minerals, including clays; (3) cation exchange capacity (CEC); (4) surface area; and (5) chemical distribution ratios (R_d) . These data were used to construct a onedimensional reaction-transport model of tailings water transport through the Mancos Shale. The model includes equations governing aqueous speciation, mineral dissolution, mineral precipitation, mixing with other ground water, cation exchange, and adsorption. In addition, two sequential batch-leaching tests were conducted, and the results were used to help calibrate the geochemical model. Numerous analyses of tailings pore fluids have been made previously, and no additional analyses were made for this study.

Samples of Mancos Shale were collected from 10 cores that were evenly distributed in the area of the proposed disposal cell (see Appendix A for locations). Samples were obtained from the Prairie Canyon and Blue Gate Shale Members of the Mancos Formation, the geologic units that directly underlie the proposed disposal cell (Table 1). Five samples were collected from each of the 10 cores for a total of 50 samples. The disposal cell excavation will extend approximately 20 feet (ft) below the ground surface; thus, the uppermost sample of each core was collected from a depth of 40 ft. The remaining four samples per core were collected at equal intervals below 40 ft; with the deepest samples obtained at 300 ft. Evaluation of these samples provided a reasonable set of data to assess lateral and vertical distributions of geochemical properties in the Mancos Shale underlying the location of the proposed disposal cell.

Table 1. Geologic Members of Mancos Shale in Samples Used in this Studf

PC = Prairie Canyon Member. BG = Blue Gate Shale Member.

The water-soluble chemistry of all 50 samples was measured. Samples from the 40-ft depth (and some from the 105-ft depth) were used for additional analyses because these samples are representative of rock directly beneath the proposed tailings disposal cell and are most likely to contact migrating tailings fluids. All 40-ft-deep samples were analyzed for surface area, mineralogy by X-ray diffraction (XRD), CEC, and uranium (U) distribution ratios. In addition, all 105-ft-deep samples were measured for U, uranium R_d and **CEC** values.

This report is divided into 10 sections. Section 0 describes the general methods used for acquisition of samples and analyses. Sections 0 through 0 present methods and results for each characterization activity, and Section 0 provides the laboratory results of two sequential batch-leaching tests. Section 0 develops each module of a reaction-transport model and presents reaction-transport simulations for the sequential batch-leaching tests and the Crescent Junction Site. Section 0 provides a summary and conclusions.

Much of the work presented in this paper on U adsorption uses an approach recently developed by Dr. James Davis and co-workers at the U.S. Geological Survey. Davis et al. present a sophisticated method to evaluate U adsorption using detailed data collected at a former uranium milling site at Naturita, Colorado, a site administered by the U.S. Department of Energy Office of Legacy Management. The U.S. Nuclear Regulatory Commission funded the recent work by Davis and others specifically to advance the state-ofthe-art in geochemical methods used for the transport of U at uranium milling sites.

General Methods

Sample Collection and Preparation

Five samples were collected from each of the 10 cores. The samples were collected at depths of 40, 105, 170, 235, and 300 ft. Sample numbers are in the format CJ-201-40, where CJ is Crescent Junction, 201 is the boring number, and 40 is the depth measured in feet. Samples were 2 inches in diameter and contained approximately 2.5 linear inches of core. Samples were air dried and stored in plastic containers until used for the tests. Dried samples were lightly crushed with a hammer or pestle, sieved, and split using a riffle splitter as needed for the tests. Bulk sample weights ranged from 238 grams (g) to 615 **g.** Moisture contents range from 0.76 percent to 3.01 percent and have no obvious relationship to sample depth or lithology. Laboratory notes are presented in Appendix B and a complete set of raw data and Excel calculations are presented in Appendix C.

Analytical Methods

The laboratory portion of the work was conducted in the DOE Environmental Sciences Laboratory at Grand Junction, Colorado. A subcontract was procured with Dr. William Hood, Grand Junction, Colorado, to conduct XRD analysis. Micromeritics Analytical Services, Norcross, Georgia, measured particle surface areas using the standard Brunauer, Emmett, and Teller (BET) method. Table 2 presents analytical methods used for water chemistry measurements.

Table 2. Analytical Methods

Analysis of Water-Soluble Extracts

Introduction

This portion of the work scope was designed to identify and estimate abundances of minerals present in the water-soluble fraction of the Mancos Shale. The water-soluble fraction is the most reactive portion of the aquifer and will likely have a significant effect on chemical evolution of the ground water system. Appendix C contains a complete set of chemical results for the water-soluble extractions.

Methods

Samples were crushed, air dried, and sieved to -10 +18 mesh (1 to 2 millimeters [mm] nominal diameter). This size fraction was selected because it **(1)** is likely to have sufficient material to accomplish the tests, (2) is uniform enough so that surface area is relatively constant and easy to determine, (3) is efficient to work with in the laboratory (testing apparatus can be relatively small), and (4) should be reasonably representative of the bulk rock encountered in the field.

The leaching procedure was modified from procedure CB (BT-1) (STO 210). Two grams of the sieved sample was placed in a plastic centrifuge tube with 100 milliliters (mL) of deionized water at room temperature. The tube was agitated end over end for 4 hours. The solids were separated from the water by centrifuging and decanting to produce a clear solution. The solutions were analyzed for pH, oxidationreduction potential (ORP), specific conductance, alkalinity, ammonium (NH4), calcium (Ca), chloride (Cl), magnesium (Mg), nitrate (NO₃), potassium (K), sodium (Na), sulfate (SO₄), and U.

Results and Discussion

All 50 samples were analyzed by the water-extraction method. Analyses included major ions that constitute most water-soluble minerals as well as the mill tailings contaminants NH4 and U. Mass balance calculations made using the computer program NETPATH (Plummer et al. 1994) were used to estimate the mineralogy of the water-soluble fraction. Mineral identifications were aided by XRD analyses (Section 5), but XRD is limited in its ability to detect small amounts (typically less than 10 percent) of minerals and does not provide quantitative information on mineral abundance. This. section reports the concentrations of the constituents in the water-soluble fractions of the Mancos Shale samples and the mineral assemblages calculated using NETPATH.

Masses of Constituents in the Water-Soluble Extracts

The solid-phase concentration of each constituent (Table 3) was calculated from the measured concentration and the solids-to-solution ratio (2 **g** in 100 mL). The total water-soluble fraction ranged from 0.40 to 2.85 percent (by weight) of the sample, with a mean value of 0.72 percent (Table 3; 1 percent is equivalent to 1,000 micrograms per gram [µg/g]). Only two samples (CJ-203-40 and CJ-204-40) exceeded 1 percent, and both were from the 40-ft depth. Values of pH in the solutions after agitation ranged from 6.17 to 9.88, with a mean pH value of 9.50. Only five samples, all from the 40-ft depth, had pH values less than 9.00.

The water-soluble fractions are dominated by carbonate (alkalinity), Na, and SO₄, with lesser amounts of Ca, Cl, K, and Mg (see "Mean" in Table 3). Data in Table 3, show that Ca and SO₄ are more dominant in the shallow (40-ft) samples than in the samples obtained from deeper depths, suggesting the presence **of'** more gypsum (CaSO₄ *2H₂O) in the shallow horizon. Alkalinity and Na appear to increase in the deeper horizons, suggesting the presence of nahcolite (NaHCO $_3$) in the deeper horizons.

Concentrations of **N03-N** were less than the detection limit of 5.5 pg/g in all but three samples ***** The three samples with higher **N03-N** values were collected at the 40-ft depth. Ammonia-N (NH3-N) values ranged from less than the detection limit of 4 μ g/g to 36 μ g/g and had no obvious correlation with depth. Watersoluble U concentrations ranged from 0.006 to 0.519 μ g/g. Only one sample (CJ-208-40) had a U) concentration more than 0.1 pg/g, and that sample was from the 40-ft depth. These U values are low compared to an average total concentration of U reported for the Earth's crust (1.8 pg/g; Mason and Moore

1982) and an average value (20 µg/g) reported by Wedepohl (1974) for marine black shale such as the Mancos Shale. Part of the reason for the lower results is that the literature crustal and marine black shale averages were analyzed on total digestions of the rock rather than the water-soluble fractions.

Mineralogy of the Water-Soluble Fraction

The geochemical computer program NETPATH was used to calculate possible mineral assemblages dissolved from the Mancos Shale samples by the deionized Water. The most common use of NETPATH is to model mineral-phase reactions that could occur as ground water flows through an aquifer, based on the water chemistry data of two ground water samples collected along a flow path. In this study, we used NETPATH to help evaluate the mineralogy of the water-soluble fractions. To do this, deionized water was used as the initial solution composition and the analytical results of the water after reaction with solid samples was used as the final water composition. Water-soluble concentrations of NH₃ and U were also measured, but the identification of the mineral phases containing these constituents was not pursued because of low concentrations of these contaminants. Results of XRD analysis helped to identify mineral phases that are most likely to contain the water-soluble components (Section 0).

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To run NETPATH, the solution chemistry is entered into a database. From the database, dissolved species, mineral saturation indices, dissolved inorganic carbon concentration (from alkalinity), and electrical balance were calculated for each solution. The molal concentrations of the constituent elements were then imported to NETPATH for mass balance calculations. Combinations of mineral phases were specified, and the program calculated the amount of each phase that must precipitate or dissolve to meet the compositional constraints.

NETPATH solves a set of mass balance chemical equations and can determine a set of mineral assemblages to account for the chemical differences between two water compositions. For example, if water A has 10 millimoles (mmol) of Ca and evolves to water B that has 2 mmol of Ca, then 8 mmol of Ca was lost. Several chemical processes can account for the loss, including (1) precipitation of calcite $(CaCO₃)$, (2) precipitation of gypsum, (3) exchange of Ca for Na on an exchanger, and other processes. Clearly, each process will produce other changes in the water chemistry; for example, calcite precipitation will cause carbon (C) loss, gypsum precipitation will cause SO₄ loss, etc. Thus, NETPATH models for this example are also constrained by mass balances for C and SO4.

The user selects *constraints* and *permissible phases* for use With a simulation. *Constraints* are elements and *permissible phases* are minerals, exchange sites, or gases. Given the number of *constraints,* there must be enough *permissible phases* to afford a solution and there must be a phase for each *constraint* (for example if Ca is a *constraint,* there must be a phase that contains Ca).

Important points to remember about NETPATH models are (1) models are not unique, (2) models are not supported by thermodynamic principles (for example, a NETPATH model can have calcite precipitating, even though precipitation is impossible because of undersaturation), (3) the final water composition must have evolved from the initial composition (e.g., two water samples from the same stream tube in a flow system or water reacted with rock in a laboratory batch test such as was done in this study), and (4) the models are strongly a function of the *constraints* and *permissible phases* selected by the user. The user selects *constraints* and *permissible phases* by evaluating the data available about the site and applying geochemical principles.

Table 3. Results of Water Extractions Expressed as Micrograms of Constituent Per Gram of Rock

pH values in final solution; alkalinity as CaCO₃

U.S. Department of Energy
June 2006

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Seven elements (C, Ca, Cl, K, Mg, Na, and S) were used to constrain the NETPATH models. After trial and error using a wide range of phase assemblages in NETPATH, seven permissible phases were selected because they produced a reasonable set of results. The selection of permissible phases was aided by XRD (see Section 5), and hand-lens observations, and literature on soluble mineral phases that are likely to be present in the Mancos Shale. The minerals selected as permissible phases are calcite, gypsum, thenardite $(Na₂SO₄)$, nahcolite, halite (NaCI), sylvite (KCI), and dolomite $[CaMq(CO₃)₂]$. Cation exchange of Ca and Na was also permitted. Because eight phases (including exchange) were permitted and only seven constraints were used, each sample analysis produced two to three models that exactly match the mass balance constraints. Selection of the model for each sample was somewhat arbitrary, but the major phases are similar among all the models.

The NETPATH results are given in millimole per liter (mmol/L) of test solution (water-soluble extractions). The tests used 2 g of rock in 0.10 L of water. Assuming a porosity of 0.25 and a mineral density of 2500 grams per liter (g/L) for the Mancos Shale at the Crescent Junction Site, a conversion factor of 0.375 is used to convert the NETPATH results to input values for the geochemical modeling program PHREEQC used in Section 9 (Parkhurst and Appelo 1999):

where "gw" is ground water and "test soln" is test solution.

The results of the mineral abundances calculated by NETPATH are presented in units of mmol/L of ground water (Table 4). Although thenardite was included in the list of permissible phases, it was not present in any of the selected mineral assemblages (Table 4).

On the basis of the means presented in Table 4, the mineral assemblage is dominated by nahcolite with major amounts of Ca-Na exchange and gypsum. Halite, sylvite, and dolomite occur in lesser amounts. Potassium is contained only in sylvite, thus the concentration of sylvite is constrained by the K concentration. Similarly, the concentration of Mg dictates the concentration of dolomite. All other constraints are contained in more than one phase, and mass balance equations must be solved. Calcite constitutes to 24.23 percent of the soluble mineral mass, but in some simulations calcite had to precipitate to meet the mass balance constraints.

Gypsum and dolomite are more dominant in the samples from the 40-ft depth than in samples from deeper depths (Table 4). Nahcolite and halite are more concentrated in the samples from deeper depths than in the samples from 40-ft depths. Sylvite is concentrated more in the 40-ft and 105-ft samples. The results presented in Table 4'are used in Section 9.3 to designate initial mineral concentrations for a ground water reaction-transport model.

Table 4. Water-Soluble Mineralogy Estimates Based on Mass Balance Approach (expressed as millimole per literof *ground water; negative values indicate precipitation)* *Table 4 (continued). Water-Soluble Mineralogy Estimates Based on Mass Balance Approach (expressed as* millimole *per* liter *of ground water, negative values indicating precipitation)*

U.S. Department of Energy June 2006

Cation Exchange Capacity

Introduction

CEC is the ability of a solid substance to freely exchange one cation for another. For example, cations in smectite clay minerals are able to readily enter and exit the interlayer (exchangeable) sites. Cations are selective in their ability to occupy exchange sites, and selectivity is often influenced by the hydration state of the dissolved cation. **CEC** can be dependent on other solution variables such as pH values. **CEC** is commonly expressed in units of milliequivalents per 100 grams (meq/100 g). Typical values of **CEC** for pure clays are 3 to 15 meq/100 g for kaolinite, 80 to 150 meq/100 g for smectite, and 10 to 40 meq/100 g for illite (Grim 1953). Mancos Shale samples analyzed in this study contain mostly kaolinite, illite, and smectitic interlayered clay minerals (Section 5).

Because clay minerals are abundant, the Mancos Shale may have a large **CEC** that can cause significant changes to ground water chemistry. Knowledge of the **CEC** is required to develop a geochemical model of water-rock interactions. It is assumed that most of the **CEC** in the Crescent Junction samples is due to the clay minerals. In the reaction-transport modeling, presented in Section 9.5, the **CEC** is considered a fixed property of the rock, but selectivity of cations occupying the exchange sites is controlled by solution chemistry. The laboratory results presented in this section are used to specify the number of exchange sites in the models.

Methods

CEC was measured on 20 core samples of Mancos Shale collected from 40-ft and 105-ft depths. Samples were crushed, air dried, and sieved to -10 +18 mesh (1- to 2-mm nominal diameter) in the same manner as for the other tests. This size fraction was selected because it is easy to work with in small-scale **CEC** tests and has a reasonably constant surface area (Section 6). The results could be scaled to field conditions by normalizing to the surface area. **CEC** was also measured on a powdered Mancos Shale sample to provide information on maximum **CEC.**

Various methods have been used to measure CEC. Most methods rely on saturation of the exchange sites with a single composition of cation; subsequently the cation is removed from the fully saturated mineral and its mass is measured to determine CEC. No general method exists that can be reliably used for all claybearing samples (Bain and Smith 1987). Methods include saturation with Ba, Ca, K, Na, or NH4, (Jackson 1969; Bain and Smith 1987).

In this study, **CEC** was determined using the ammonium saturation method (Chapman 1965 as described in Bain and Smith 1987). This method was selected largely because the saturating solution (ammonium acetate) is highly buffered and maintains its near-neutral pH value throughout the test (Chapman **1965).** Another advantage is that NH4 concentrations are easily measured in the sodium-chloride extracts. Accuracy of the method was checked by measuring the **CEC** of a calcium-montmorillonite clay standard (sample number STx-1) from the Source Claycollection of the Clay Minerals Society. **CEC** values published by Borden and Giese (2001) for this specimen were 89 meq/1 00 g with a standard deviation of 2 meq/1 00 g. **CEC** values of nine repetitions in our laboratory ranged from 71.4 to 85.7 meq/1 00 g with a mean of 77.0 and a standard deviation of 6.6. Although slightly lower than the Borden and Giese (2001) results, our values are reasonably similar and probably within the range of analytical uncertainty.

For the ammonium saturation method, the clay sample was first saturated with $NH₄$ ions using 1 molar (M) ammonium acetate followed by exchange with sodium chloride (STO 210, method CB [CEC-1]). For a **CEC** measurement, an exact weight of clay ranging from 20.milligrams (mg) to 200 mg was combined with 20 mL of **1** M ammonium acetate and agitated end over end for 2 hours. The solids were separated from the liquid phase by centrifugation, and the saturation process with ammonium acetate was repeated five times. Isopropyl alcohol (20 mL) was then added to the solids, the mixture was agitated by hand, and centrifuged. Five additional washings with isopropyl alcohol were conducted, after which 20 mL of 100 g/L sodium chloride solution was added to the solid phases to initiate removal of the NH₄ cations from the exchange sites. The sodium chloride solution was agitated by hand seven times. The NH4 concentration in the resultant solution was measured spectrophotometrically and was used to calculate the **CEC.**

Results and Discussion

CEC measurements ranged from 0.54 to 36.29 meq/100 g with a mean of 11.23 meq/100 g (Table 5). The mean is within the range published by Grim (1953) for pure kaolinite (3 to 15 meq/100 g) and pure illite (10 to 40 meq/100 g) but is significantly less than pure smectite (80 to 150 meq/100 g). The values for these Mancos Shale samples are reasonable for the clay mineral compositions as determined by XRD analysis (Section 5).

One sample (CJ-205-40) was ground to a powder to compare CEC values for a finer grain size. The CEC measurements of the powder (10.29 and 10.43 meq/100 g for duplicate samples) were similar to the 1- to 2-mm grain size. Although limited to one analysis, this result provides confidence for applying the laboratory CEC values to the field in the reaction-transport models (Section **9.5).** The CECs are used in the reaction-transport models to simulate cation exchange of Ca, K, NH4, and Na. Retardation of NH4 in the model is assumed to be caused by the cation exchange with the Mancos Shale.

Table 5. Results of Cation Exchange Measurements

 ${}^{a}P$ = Powdered Sample, Dup = Duplicate

U.S. Department of Energy June 2006

A brief literature search was conducted to determine if the **CEC** values of Crescent Junction Mancos Shale samples are comparable to **CEC** measurements of Mancos Shale samples from other areas. Evangelou et al. (1984) collected samples of partially weathered and unweathered outcrop samples of Mancos Shale from the West Salt Creek Watershed near Grand Junction, Colorado, and analyzed the CEC using the calcium/barium exchange method. They reported **CEC** values ranging from 13.25 to 19.96 meq/100 g. These values are similar to the Crescent Junction values and suggest that the **CEC** for Mancos Shale may be relatively constant.

X-Ray Diffraction Analysis

Introduction

XRD was used to identify minerals present in the core samples. XRD is capable of defining the internal arrangement of atoms in a crystalline lattice, thus making it possible to positively identify minerals. Identification of mineralogy based on chemical methods (Section 3.3.2) is more ambiguous because often minerals have similar chemical compositions and, in some cases, identical compositions. Unlike chemical methods, XRD analysis is not able to detect small quantities of minerals (XRD is generally limited to -detection of approximately 10 percent, but detection is dependent on mineral crystallinity) and only semiquantitative estimates of mineral abundances are possible. XRD analyses were conducted at Mesa State College in Grand Junction, Colorado, by Dr. William Hood (Appendix D).

Chemical interaction between ground water and Mancos Shale is likely to occur mostly at the surfaces of clay minerals. Therefore, clay mineral chemistry is important to the transport of contaminants by the ground water. XRD is one of the best analytical tools to identify clay minerals. For this study, 10 core samples, all from the 40-ft depth, were analyzed by XRD.

For XRD analysis, the sample is subjected to an x-ray beam. The atomic lattice within the individual mineral crystals diffracts (reflects) the x-ray beam, and the angle of diffraction is measured. The angle of diffraction
and the intensity of the diffracted x-rays produce a "fingerprint" that can be used to identify the minerals (Jackson 1969).

Methods

A random powder mount was used for identification of major minerals (Appendix D). Samples were finely powdered with a mortar and pestle, and placed in a sample holder, and scanned from 4.2 to 50 degrees 20 using a Rigaku Miniflex x-ray diffractometer (Appendix D). The Jade computer program was used to' calculate spacings and to compare diffraction patterns of the samples with diffraction patterns of known minerals.

Splits of the samples were analyzed for clay mineralogy (Appendix **D).** Analysis of clay mineralogy is more complex than for the major minerals because the sample requires four separate preparation methods and an XRD analysis of each. To prepare the clay mineral splits, the clays were segregated into fine particles and not flocculated. The bulk sample was first placed in water overnight to remove readily soluble material and initiate disaggregation of the clays. Calgon (sodium metaphosphate) was added to further disperse clay minerals and the mixture was allowed to sit for 8 hours. Magnesium was added to the suspension to saturate the clay mineral interlayers (to produce a constant d spacing of this lattice plane). A small quantity of the suspension was smeared on a glass microscope slide and allowed to dry. The four methods used to treat the clay mineral separates are (1) air drying, (2) glycolating, (3) heating to 300 °C, and (4) heating to * 550 °C. More details on the preparation and analysis methods are available in Appendix **D.** XRD patterns from these four treatments were used to positively identify the clay minerals and to estimate the amount of smectite layers in interlayered illite/smectite clays.

Estimates of the mineral abundance of non-phyllosilicate minerals were made by a process including background removal, normalizing the peak intensities to a quartz standard, and summing the peak intensities. Estimates of clay mineral abundance used the glycolated diffractograms and employed a series of computer enhancements, including background removal and peak assessment. The ratio of illite to smectite layers in the mixed-layer clays was estimated from a comparison of the Mg-saturated and

U.S. Department of Energy Geochemical Attenuation and Performance Assessment Modeling **June 2006 Doc. No. X0173400** Page 13 glycolated diffractograms. More details on the methods used to estimate mineral abundances are presented in Appendix **D.** X-ray diffraction methods provide only a semi-quantitative estimate of the mineral abundances. According to Schultz (1964) based on his work with Pierre Shale (a Mancos Shale equivalent), the abundance estimates are adversely affected by sampling, sample preparation, machine response, and, most importantly, interpretation. Schultz (1964) also states that if a mineral makes up more than 15 percent of the sample, the precision of the abundance estimate is usually within about 10 percent. At lower mineral concentrations, the uncertainty increases.

Results and Discussion

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XRD analysis was conducted on the 10 samples of Mancos Shale collected from the 40-ft depth. The mineralogical composition of these samples is dominated by quartz, with lesser amounts of dolomite and calcite, small amounts of feldspar, and traces of gypsum (Table 6.). Orthoclase and plagioclase feldspar. occur in approximately equal amounts.

Table 6. Non-Phyllosilicate Mineral Abundance Estimates in Percent Based on XRD Analysis of Bulk Samples

The clay mineral fraction is dominated by mixed-layer (mostly illite/smectite) clays, illite, and kaolinite (Table 7.). The mixed-layer. clays are dominated by illite layers. A small amount of mixed-layer chlorite/vermiculite is probably present.

'Mostly mixed-layer illite/smectite with minor chlorite/vermiculite.

A brief literature survey was conducted to determine if the XRD results for the Mancos Shale at Crescent •.J Junction are typical of other Mancos Shale localities. Schultz (1997) reported that clays in the Mancos Shale from the Colorado Plateau region contain 50 to 60 percent mixed-layer illite/smectite, 12 to 15 percent poorly ordered kaolinite, and 30 to 35 percent illite; however, no source of data is provided in that publication. Nadeau and Reynolds (1981b) discuss clay mineral XRD results of 77 bentonite-shale paired samples collected from the Mancos Shale throughout the Western Interior of North America. They determined that the clays are primarily randomly mixed-layer illite/smectite with illite compositions ranging from 0 to 85 percent. They attribute the origin of the clays to volcanic ash. The illitic component increases in response to increased burial metamorphism, as does the ordering of the illite/smectite.

Nadeau and Reynolds (1981 a) discuss XRD clay mineralogy results of 690 samples of Mancos Shale collected frpm 154 sites in the four-state region of Colorado, Utah, Arizona, and New Mexico. These clays are dominated by mixed-layer illite/smectite, with 20 to 60 percent illite layers. The regional distribution of ordering in the mixed-layer illite/smectite is attributed to differences in burial metamorphism of the Mancos Shale. Evangelou et al. (1984) collected samples of partially weathered and unweathered outcrop samples of Mancos Shale from the West Salt Creek Watershed near Grand Junction, Colorado, and analyzed the clay fractions by XRD. These clay fractions contain mica, kaolin, smectite, and randomly interstratified mixed-layer clay, possibly mica/vermiculite. These studies generally indicate that the clays identified in the Crescent Junction samples (kaolinite, illite, and mixed-layer illite/smectite) are typical of the Mancos Shale throughout much of its depositional basin.

Surface Area

Introduction

Knowledge of the surface area of the Mancos Shale samples is needed to relate the results to the transport .of contaminants through the subsurface. Processes such as cation exchange and adsorption are directly related to the surface area that the ground water contacts. For example, for the same travel distance, interaction of dissolved contaminants with the rock will be less in a fracture-dominated matrix (small surface area) than in a porous-media (higher surface area) flow. It is beyond the scope of this study to determine the nature of the flow (fracture versus porous media) in the subsurface at the Crescent Junction Site. However, to properly use the data collected during this study in site models, these data will need to be normalized to surface area. Therefore, it is important to measure the surface area of the samples used in the tests. Surface area was determined for 10 core samples from the 40-ft depth.

Methods

Samples were crushed, air dried, and sieved to **-10** +18 mesh (1- to 2-mm nominal diameter) in the same manner as for other tests. Surface area was determined by the standard BET method. This method is the most widely used method for determining particle surface area. Samples were prepared by heating while simultaneously evacuating to remove impurities. The prepared samples were then cooled with liquid nitrogen and analyzed by measuring the volume of N_2 gas adsorbed at specific pressures (Micromeritics 2006; Jackson 1969). Multiple-point isotherm measurements were conducted. The multiple point tests are more accurate than single point tests. Micromeritics Analytical Services, Norcross, Georgia, conducted the surface area measurements (Appendix E).

Results and Discussion

Measured surface areas of the 10 samples from the 40-ft depth range from 8.81 to 13.22 square meters per gram (m^2/g) with a mean of 11.02 m^2/g (Table 8.). These values are reasonably comparable to surface areas measured by Davis and Curtis (2003) on alluvial aquifer samples obtained from the former Naturita, Colorado, uranium-ore processing site. Surface areas measured on the 16 Naturita samples range from 5.2 to 20.0 m²/g with a mean of 12.43 m²/g and standard deviation of 3.77 m²/g. Surface area was used as an input to the surface-complexation algorithm in the transport model presented in Section 0.
Surface area was also measured on a powdered split of Crescent Junction sample CJ-205-40. The surface area of the powdered sample (10.7 m²/g) is consistent with surface areas measured on the 1- to 2-mm fractions, indicating that the BET method is probably accounting for intraparticle surface area.

Table 8. Results of BET Surface Area Analysis of 1- to 2-mm Size Fraction

) Distribution Ratios

Introduction

Distribution ratios (R_d) are a measure of the partitioning of a contaminant between the ground water and the solid fraction of the aquifer. The higher the R_d value, the more partitioning to the solids and the more retardation of the contaminant. The R_d value is empirical and is simply the ratio of the measured concentration in the solids (milligrams per kilogram [mg/kg]) to the measured concentration in the ground water (milligrams per liter [mg/L]), and has units of milliliters per gram (mL/g). R_d values are often used to simulate retardation of contaminants in ground water models. When used in these models, an implicit assumption exists that chemical retardation occurs by the process of adsorption under equilibrium conditions. For the equilibrium assumption, the R_d value is often referred to as a distribution coefficient (K_d). Many models also assume that K_d does not vary with the concentration of the contaminant, but research has shown that it often does, and algorithms such as the Langmuir or Freudlich equations are used to produce a better fit to R_d data. To test if R_d values vary with contaminant concentrations, multiple points using various concentrations are measured, and plots of these "adsorption isotherms" are fitted with the various models. If the plot is reasonably linear, then models using a constant K_d value (also termed linear isotherm) will produce satisfactory results.

The R_d value can vary significantly with solution chemistry. For example, Davis et al. (2004) show that R_d values for U in a sample of alluvium vary by more than a factor of 10, depending on dissolved carbonate concentrations and pH values. Also, the R_d approach is only valid for contaminants that occur in trace concentrations.

For the reaction-transport modeling, a surface-complexation approach is employed instead of using R_d values to simulate retardation of U (Section 0). With the surface-complexation approach, the R_d value changes as chemical conditions (especially pH and $pCO₂$) change in the aquifer. The main use of the R_d values determined in this study is to test and calibrate the surface-complexation module used in the reaction-transport modeling.

U.S. Department of Energy June 2006

Methods

Uranium R_d values were determined using procedure CB (Rd-1) in STO 210. In summary, a sample of Mancos Shale was crushed and sieved to $-10 + 18$ mesh (1- to 2-mm nominal diameter). The sample (5 g for a single-point R_d value) was placed in a plastic centrifuge tube with 100 mL of pH value-adjusted synthetic pore fluid (SPF) that simulates tailings or evolved pore fluid. The tubes were agitated end over end for 24 hours and then centrifuged, decanted, and filtered through a 0.45-micrometer (um) filter. Filtrates were analyzed for electrical conductivity, pH value, alkalinity, and U concentration. Controls included SPF without solids and duplicates. The concentration of U in the solid phase was determined from the loss of U from the SPF solution. R_d values were determined from the calculated concentration of U in the solid phase and the measured U concentration in the solution.

Uranium R_d values were determined using three different SPF compositions designed to simulate ground water that could be present in the Mancos Shale after construction of the disposal cell. One solution (SPF-1) simulates pore fluids that currently exist in the tailings. Another solution (SPF-2) simulates tailings pore water mixed with 50 percent Mancos Shale ground water. **A** third solution (SPF-3) simulates water that results from the sequential batch-leaching testing (Section 0). Table 9 provides compositions of the SPF fluids. The SPF solutions were spiked with 1 mg/L U to enable measurement of the R_d values.

Table 9. Composition (milligrams per liter) of SPF Used in Uranium R_d Tests

"The value for **Cl** for SPF-2 should have been approximately 4541 mg/L. The low value (309.49) resulted from an error in designing the SPF-2 solution.

Single-point R_d measurements for U were determined for 20 samples, and 6-point isotherms were determined for five samples of SPF-1. For the five samples used for isotherms, single-point R_d values were also determined using SPF-2 and SPF-3 to evaluate sensitivity to solution chemistry.

Results and Discussion

R_d values were corrected for the amount of "labile" U contained in the Mancos Shale samples. The labile U fraction'is defined as U in the solid fraction capable of being released to the solution during a test (Davis et al. 2004). Release of labile U is usually a function of time with longer agitation periods resulting in higher concentrations of labile U. A constant value of 0.0436 µg/g for the labile U fraction was used in our study. This Value is the mean of eight measurements of the sequential batch-leaching tests (Section 0). In the sequential batch tests, U-free SPF was reacted for various time periods with Mancos Shale samples, and U concentrations were measured in the final solutions. For comparison, Davis et al. (2004) used a value of 0.21 pg/g for labile U in uncontaminated alluvial aquifer sediment from the Naturita Site.

Despite making corrections for the labile U, some tests produced negative R_d values (Table 11). Negative R_d values are a combination of (1) analytical uncertainty, (2) relatively low R_d values, and (3) variable contribution from labile U. Minimum and maximum R_d values presented in the tables in this section are

based on analytical imprecision of ± 2.5 percent in all U measurements. Many of the R_d determinations were fairly sensitive to the analytical imprecision.

Single-point determinations of R_d using SPF-1 (tailings fluid) ranged from negative values to 0.84 mL/g (Table 10). The maximum R_d value, accounting for 2.5 percent error in all measurements, is 1.9 mL/g. These values are reasonably similar to R_d values measured on alluvial aquifer material from other uranium-ore milling sites. For example, R_d values for samples collected at the DOE New Rifle, Colorado, Site, for alluvium and Wasatch Formation, range from negative values to 3.7 mL/g and 1.6 mL/g, respectively (DOE 1999).

^aTest conditions: pH 7.57, alkalinity 570 mg/L CaCO₃, 50 g/L. Minimum and maximum R_d values based on ±2.5 percent for U analyses on final solutions and SPF-1.

Single-point R_d values for samples from the 40-ft depth using SPF-2 (simulated tailings pore water mixed with 50 percent Mancos Shale ground water) range from -0.46 to 0.51 mL/g with one negative value (Table 11). Single-point R_d values for samples from the 40-ft depth using SPF-3 (simulated water that results from the sequential batch-leaching testing) ranged from 0.41 to 1.92 mL/g (Table 12). Maximum R_d values for SPF-2 and SPF-3 are 1.56 and 3.04 mL/g, respectively. The R_d values for SPF-3 are slightly higher than those of the SPF-1 and SPF-2 solutions, indicating increased partitioning to the solid phase. This might be expected for SPF-3, as it has a much lower dissolved carbon concentration and a higher pH value (Table 9).

U.S. Department of Energy June 2006

Table 11. Final Concentrations of U in Liquid (U_i) and Solid (U_s) Phases and R_d for Single-Point *Determination with SPF-2a*

^aTest conditions: pH 7.69, alkalinity 907 mg/L CaCO₃, 50 g/L. Minimum and maximum R_d values are based on ± 2.5 percent for U analyses on final solutions and SPF-2. D in sample ID number indicates duplicate analysis.

Table 12. Final Concentrations of U in Liquid (U_I) and Solid (U_s) Phases and R_d for Single-Point Determination with SPF-3'

Test conditions: pH 7.91, alkalinity 313 mg/L CaCO₃, 50 g/L. Minimum and maximum R_d values are based on ±2.5 percent for U analyses on final solutions and SPF-3.

Isotherm tests using mass-to-water ratios of 10, 20, 50, 100, 200, and 250 **g/L** were conducted on five Mancos Shale samples from the 40-ft depth using SPF-1. For a linear isotherm, the results should plot as a straight line on a log-log plot such as shown on Figure 1a. The lack of linearity of these plots is likely due in large part to analytical imprecision. The analytical uncertainty is greatest at the highest values of dissolved U (Figure **1** b). Better definition would result if the tests had spanned a larger range of U concentrations and used larger masses of solids. Although these data are scattered, it is apparent that the uranium R_d values of the Crescent Junction Mancos Shale samples are relatively low.

The same data shown on Figure 1a are plotted on Figure 2 at a larger scale, along with the Naturita uranium R_d results from Davis et al. (2004). Data from Crescent Junction form a relatively tight group near the dissolved U concentration of 10^{-5.50} mol/L. Many of the Crescent Junction sample uranium R_d values are lower than any of the Naturita results. The lowest Naturita R_d values were measured on solutions with high **C02** partial pressures and a near-neutral pH value of 6.88 (Davis1 on Figure 2). The Crescent Junction SPFs also have high CO₂ concentrations and near-neutral pH (Table 9), which accounts for their low R_d values. The mineralogy of the Mancos Shale (e.g., low in iron oxyhydroxide) compared to alluvial samples may also be partly responsible for the lower R_d values.

U.S. Department of Energy June 2006

Figure 1. a. Isotherm Plot of Uranium R_d Determinations Made on Crescent Junction *Mancos Shale Sample Using SPF-1,*

> *b. Isotherm Plot of Sample 202-40 Showing Effect of Analytical Imprecision on Calculated Uranium Rd Values; Maximum and Minimum Rd Values are Calculated Using* ±2.5% *Uncertainty on all Analytical Measurements.*

U.S. Department of Energy June 2006

Geochemical Attenuation and Performance Assessment Modeling Doe. No. X0173400 Page 20

a.

"b.

Davis1 = AGW-6, pH 6.88, 6.8 percent **C02,** 820 g/L Davis2 = AGW-5, pH 7.18, 1.57 percent **C02,** 250 g/L Davis3 = AGW-7, pH 7.58, 0.47 percent **C02,** 125 **g/L** Davis4.= AGW-3 (laboratory air), pH 7.94, 0.05 percent **C02,** 25 **g/L**

The uranium R_d values from our testing are more scattered than the results measured by Davis et al. (2004) on the Naturita samples. The scatter is largely due to the lower rock-to-water ratios used in our tests. Davis et al. (2004) used up to 820 g/L, whereas, we used a maximum of 250 g/L. Because of the low R_d values, a higher rock-to-water ratio would have reduced the scatter in our data. However, we believe that additional measurements at higher rock-to-water ratios are unnecessary because the results are not likely to change the conclusion that the parameters used in the Davis et al. (2004) model are a reasonable approximation for use at the Crescent Junction site.

Sequential Batch-Leaching Tests

Introduction

Two sequential batch-leaching tests were conducted to measure evolution of tailings leachate chemistry as it interacts with Mancos Shale. Results of these tests were used to help calibrate the reaction transport models presented in Section 0. In the sequential batch-leaching tests, the test fluid (SPF-1) was not spiked with U as it was for the H_d tests (Section 7.27). Rather, any U in the solution was released from the Mancos Shale samples. The evolution of major-ion chemistry and pH were monitored to determine the effects of these changes on U partitioning.

Methods

A 100-g sample of Mancos Shale was crushed, air dried, and sieved to -10 +18 mesh (1- to 2-mm nominal diameter). The sample was combined with 400 mL of SPF-1 in a 500-mL glass Erlenmeyer flask, agitated on an orbital shake table for 24 hours, and then centrifugated, decanted, and filtered (0.45 µm) to separate solids from the solution. A 50-mL split was retained and analyzed for pH, ORP, specific conductance, alkalinity, Ca, Cl, K, Mg, Na, NH₃, NO₃, SO₄, and U. The remaining solution was placed with approximately 87.5 **g** of the next Mancos Shale sample. The exact amount of the Mancos Shale sample was calculated so that the water-to-rock ratio remained constant. The procedure was repeated to monitor the changing solution composition as the fluid reacted with progressively more Mancos Shale.

Results and Discussion

Two sequential batch-leaching tests were conducted, one to simulate vertical transport and one to simulate horizontal transport through the Mancos Shale. To simulate horizontal transport, the 40-ft-depth samples from borings CJ-203, -204, and -206 were used, in that order. Chemical changes that occurred in the solution as it reacted with progressively more Mancos Shale include slightly increasing values of pH, Cl, K, Na, SO_4 , and U. Small decreases occurred in ORP and NH₃-N, while alkalinity, Ca, and Mg remained fairly stable (Table 13). The increase in U may be caused by longer reaction times with the sediment or by decreasing R_d values because of increasing pH values and nearly constant dissolved carbonate concentration.

To simulate vertical transport, SPF-1 (Table 9) was reacted sequentially with samples from boring CJ-205 from the 40-, 105-, 170-, 235- and 300-ft depths. Similar to the simulation of horizontal transport, pH values progressively increase with increased reaction. Progressive increases also occurred in the Cl, Na, SO4, and U concentrations. Decreasing trends occur in alkalinity and possibly NH3-N. Dissolved U concentrations increased despite lowered concentrations of dissolved carbonate (alkalinity). Some carbonate was apparently lost from the solution because of outgassing and/or mineral precipitation. The progressively higher U concentrations may simply be caused by desorption from fresh adsorption sites at each reaction step, or may be caused by variable pH and $pCO₂$ values.

**amV = millivolt.
^bCond. = electrical conductivity in microsiemens per centimeter (μS/cm).
^cAlk = alkalinity in milligrams per liter as CaCO_{3.}**

Another outcome of the sequential batch-leaching tests was an estimate of the amount of labile U in the Mancos Shale. The labile U concentration is the readily releasable portion and is needed to calculate uranium R_d values as discussed in Section 0. Labile U should be measured using a solution that strongly favors partitioning of U into the liquid phase, such as the strong carbonate solution used by Davis and Curtis (2003). SPF-1 has a relatively high carbonate concentration and should produce a reasonable estimate of labile U; however, our results should be considered as minimum values. The labile U concentrations range from 0.0120 to 0.0660 ug/g, with a mean of 0.0436 ug/g (Table 14).

Table 14. Labile U Concentrations

Reaction-Transport Modeling

Introduction

This task consists of developing a reaction transport model using the PHREEQC code (Parkhurst and Appelo 1999). Ion exchange is modeled using data from the **CEC** determinations (Section 0) and surface area measurements (Section 0). Initial conditions include concentrations of minerals as estimated from the soluble chemistry results (Section 0) and XRD results (Section 0). A surface-complexation model (SCM) for U was developed and calibrated against the uranium R_d measurements provided in Section 0.

The reaction-transport model includes one-dimension transport of tailings water through the Mancos Shale. Aqueous speciation reactions are typically fast with respect to ground water flow and were modeled at chemical equilibrium.. Water-to-rock interaction includes mineral precipitation and dissolution, adsorption,' and cation exchange. For the models used in this study, it is assumed that the system is oxidized; no reduced species were included. Results from the sequential batch reaction tests were used to test and calibrate the reaction-transport model. Input files for the sequential batch models are presented in Appendix F. Input files for the Crescent Junction models are presented in Appendix G, and the thermodynamic database is presented in Appendix H.

The reaction-transport model is adaptable to allow inclusion of such factors as (1) mixing with ground water, (2) reaction kinetics, and (3) changing oxidation-reduction state (e.g., due to biologic activity). Thus, sensitivity of the transport to various parameters can be readily estimated with additional model simulations. In this section, each "module" of the reaction-transport model is discussed separately, and input parameters used in the reaction-transport modeling are specified.

Aqueous Speciation Module

The aqueous speciation reactions used in this study are identical to those used by Davis and Curtis (2003), supplemented with reactions in the PHREEQC.dat thermodynamic database provided with the PHREEQC program(Parkhurst and Appelo 1999). All aqueous U species are from Davis and Curtis (2003). Table 15 and Table 16 provide lists of the non-U-bearing and U-bearing aqueous speciation reactions respectively, used in this study.

^aFrom Davis and Curtis (2003) and Parkhurst and Appelo (1999).

U.S. Department of Energy June 2006

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For this study, it was assumed that the reactions among all aqueous chemical species are at equilibrium.
This assumption is reasonable because reaction rates among dissolved aqueous species are relatively fast. Numerically, this is accomplished by solving the set of Log K expressions. For example, the numerical expression for the first reaction. in Table 15 is

$$
Log K = Log a_{(HCO^{-1})} - Log a_{(CO^{-2})} - Log a_{(H^{+})}
$$
 Equation 1

where a is the activity of the dissolved species. Activities are related to concentrations using activity coefficients calculated using the Debye-Hückle theory; for example:

$$
a_{(HCO_2^-)} = \eta m_{(HCO_2^-)}
$$

where γ is the Debye-Hückle activity coefficient and m $_{(HCO_2^-)}$ is the molality of the bicarbonate ion.

Mineral Precipitation/Dissolution and Gas-Phase Module

Four minerals (calcite, gypsum, halite, and nahcolite) were allowed to equilibrate with the aqueous solution during the reaction-transport simulations. Table 17 presents these mineral reactions and their associated equilibrium constants. Similar to the aqueous speciation reactions in the previous section, equilibrium was forced between the aqueous solution and each of these minerals. Numerically, this means that the following condition must hold for each mineral, using calcite as an example:

$$
Log K = Log a_{(CO_2^{-2})} - Log a_{(Ca^{2*})} - Log a_{(CaCO_3)}
$$
 Equation 3

where $a_{(CaCO3)}$ is the activity of calcite, which is assumed to be in the pure phase and, thus, has unit activity.

Table 17. Reactions and Logarithmic Equilibrium Constants for Minerals Allowed to Equilibrate During Reaction-Transport Simulationsa

From Wateq4f.dat database supplied with the PHREEQC program (Parkhurst and Appelo 1999).

Minerals present were allowed to dissolve; therefore, the initial concentration of a mineral is an important input parameter. The initial concentrations of minerals were estimated from the NETPATH results of the water-soluble chemistry results (Table 4). The mineral compositions of samples from the 40-ft depth were distinctly different from samples from deeper depths in that they contained more gypsum and less nahcolite. Because calcite was not completely dissolved during the water extraction tests, its composition was set at 1 weight percent $[0.75 \text{ mol/L}_{(gw)}]$, considered to be a reasonable value for the Mancos Shale based on hand-lens observations of core samples and identification by XRD analysis. Table 18 presents the estimated mineral abundances; these Values were used in the reaction-transport modeling.

Table 18. Estimated Mineral Abundances (in mole per liter of ground water) Used in Reaction-Transport Modeling

Mineral	Abundance (vertical transport)	Abundance (horizontal transport)
Calcite	0.75	0.75
Gypsum	0.02	0.50
Nahcolite	0.50	0.20
Halite	0.02	0.02

Though only the minerals listed in Table 17 were allowed to react during the transport simulations, other minerals were included in the thermodynamic database to track the saturation indices (SI). Saturation index . (SI) is defined as

 $SI = \frac{IAT}{Log K}$ Equation 4

where IAP is the ion activity product. If the SI is positive, the solution is oversaturated, and the mineral will tend to precipitate. If the SI is negative, the solution is undersaturated, and the mineral, if present, would tend to dissolve. By tracking the mineral SIs, the modeler is aware of additional mineral precipitation that may need to be included. Table 19 provides the minerals included only for information on the **SI.**

Table 19. Reactions and Logarithmic Equilibrium Constants for Minerals Included in Reaction-Transport Modeling but Not Allowed to Equilibrates

^a From Davis and Curtis (2003) and Parkhurst and Appelo (1999).

Several gas phases were also included in the thermodynamic database. The SI for a gas-phase reaction is the logarithm of its partial pressure (measured in atmospheres). Thus, by tracking the partial pressures the modeler can be aware of unusual situations, such as a partial pressure that exceeds 1 atmosphere, which could result in separation of agas phase. Table 20 provides the three gas phases included in the modeling. *-Table 20. Reactions and Logarithmic Equilibrium Constants for Gases Included in Reaction-Transport Modeling but not Allowed to Equilibratea*

8C02 from Davis and Curtis (2003); H20 and NH3 from Parkhurst and Appelo (1999). $aq = aqueous$

'Uranium Adsorption Module

Uranium adsorption is simulated using an SCM (Davis et al. 1978). In an SCM, adsorption is a function of the electrostatic potential and the species complexation configuration at a mineral surface; these parameters vary with solution composition. The SCM approach has been widely used to model adsorption of U(VI) to pure phase minerals, in particular, iron oxyhydroxide (Davis 2001, Hsi and Langmuir 1985, .Tripathi 1984, Morrison et al. 1995). Naturally occurring sediments are more complicated because of the presence of a mixture of many minerals. To simplify the use of the SCM approach in modeling field situations, Davis et al. (2004) developed an approach they called the Generalized Composite Model (GCM). In a GCM, the rock is treated as a composite of mineral phases, and surface-complexation constants for the composite are calibrated from laboratory measurements. As with the SCM, adsorption processes in a GCM vary with ground water composition. Thus, instead of treating the distribution of contaminant between solids and water as a constant (the so-called K_d approach), the K_d values vary throughout a transport simulation. To simplify computations, Davis et al. (2004) elected not to consider the effects of-surface charge potential on adsorption, and we also omit this effect. Thus, the adsorption module is solved numerically using equilibrium expressions in a manner similar to the aqueous-speciation module. Three types of adsorption sites (weak, strong, and very strong) were required to produce an acceptable fit to the laboratory data.

Davis and Curtis (2003) provide thermodynamic data and model parameters for a GCM. Because they used data different from that in the PHREEQC database and a different numerical processor, we started by entering their data and parameters into PHREEQC and checking the accuracy of the calculated results. The results shown on Figure 3 indicate a good match between our modeled results using PHREEQC and the laboratory results of Davis and Curtis (2003). The small differences between the PHREEQC and laboratory results for AGW-6 (Davis and Curtis 2003, label synthetic ground water as AGW) probably arise because the model is sensitive to solution parameters in this range; some of the aqueous thermodynamic data were not provided by Davis and Curtis 2003. However, the small deviation should not significantly affect the adsorption calculations for the Crescent Junction reaction-transport model.

The uranium **Rd** data reported in Section 0 were used to check the calibration of the **GCM** for U adsorption to Mancos Shale. By comparison with the data of Davis and Curtis (2003), the laboratory measurements from our study grouped tightly at relatively low values of R_d (Figure 4). Models of the Crescent Junction R_d measurements using the GCM, without any changes to the surface-complexation constants, are also presented on Figure 4 (SPF-1, SPF-2, and SPF-3). The R_d measurements generally plot above the models. As discussed in Section 0, many of the R_d measurements were negative because of variable amounts of labile U and analytical imprecision. Thus, the measured R_d values are maximum values, and the actual R_d values may be lower, as indicated by the SPF models on Figure 4. Therefore, as a first approximation, we elected to use the Davis and Curtis (2003) GCM as it is without attempting to modify the surfacecomplexation constants. With this approach, the uranium R_d values used for the SPF fluid migration are kept quite low, consistent with results of the R_d tests. Table 21 provides the GCM uranium adsorption surface complex reactions and associated equilibrium constants. Other input parameters needed for the GCM are proportions of weak, strong, and very strong sites = 0.9879, 0.012, and 0.0001, respectively; site density = 1.92 μ mol/m² (Davis and Kent 1990); surface area for Mancos Shale = 11.02 m²/g (mean value of 10 BET measurements, Section 0); and

rock-to-water ratio **=** 7500 g (based on 25 percent porosity and 2.5 g/mL rock density).

Compositions of solutions AGW-3, AGW-5, AGW-6, and AGW-7 are provided in Davis and Curtis (2003). AGW-3: pH 7.94, 0.05 percent CO₂ (lab air), 25 g/L
AGW-5: pH 7.22, 1.24 percent CO₂, 125 g/L AGW-6: pH 6.88, 6.8 percent CO₂, 820 g/L
AGW-7: pH 7.58, 0.47 percent CO₂, 125 g/L

Figure 4. Comparison of Uranium R_d Measurements (black dots) Made in This Study With Naturita Sediment Models (AGW plots) from Davis and Curtis (2003)

 -48

Table 21. Uranium Surface-Complexation Reactions and Logarithmic Equilibrium Constants Used in Reaction-Transport Modeling"

⁶From Davis et al. (2004); w, s, v **=** weak, strong, and very strong sites. Hfo represents an adsorption site.

Cation Exchange Module

Cation exchange was considered to be a prominent chemical mechanism for ground water transport in the Mancos Shale. The Mancos Shale contains a high proportion of clay minerals that include a significant fraction of smectite (Section 0). Smectites are layered silicate minerals with freely exchangeable cations within the interlayer space. The **CEC** was used in the reaction-transport modeling as a measure of the total amount of exchangeable sites present in the rock.

The mean **CEC** value of 11.23 meq/100 g (Section 0) was used for the modeling. The number of exchange sites in equivalents per liter (eq/L) of ground water (0.842 eq/L) was calculated as the product of the mass of rock per liter of ground water and the **CEC:**

Exchange Sites = 7500 $g_{(rock)}L_{(cwy)}$ \star 11.23 meq/100 $g_{(rock)}$ \star 1 eq/1000 meq = 0.842 eq/L Equation 5

Calculations of cation exchange by PHREEQC are accomplished by solving equilibrium expressions for each exchange reaction. For example, the exchange reaction and logarithmic equilibrium expression for Ca is

$$
Ca^{+2} + 2X = CaX_2
$$

$$
Log(0.8) = Log a_{(CaX_2)} - Log a_{(Ca^{+2})} - 2Log a_{(X^{+})}
$$

Log(0.8) = Log apcax.) -Log *a(ca.2) -2Log a X.)* Equation 6

where X is on the exchange site.

Table 22 presents the exchange reactions and associated logarithmic equilibrium constants used in the modeling. Some cations have a higher selectivity for exchange sites than other cations; this selectivity is accounted for in the equilibrium constants. Retardation of NH4 contamination from the tailings is likely to be largely a function of cation exchange.

Table 22. Exchange Reactions and Logarithmic Equilibrium Constants Used in Table 22. Exchange Reactions and Logarithmic Equilibrium Constants Used in
Reaction-Transport Modeling^e

> ^aFrom Parkhurst and Appelo (1999). X represents an exchange site.

Model Calibration Using Sequential Batch-Leaching Results

The sequential batch-leaching results (Section 0) were modeled to calibrate the release of background U from the Mancos Shale and to compare modeled major-ion concentrations with experimental results. Appendix F provides the PHREEQC input files for the sequential batch-leaching simulations.

The rock-to-water ratio used in the sequential batch-leaching tests was 250 g/L (100 g of rock to 400 mL of water). Because input parameters for the models are normalized to a liter of water, the parameters for the sequential batch-leaching models are different from the field models that used a rock-to-water ratio of 7,500 g/L. Table 23 provides the normalized parameters for the sequential batch-leaching models.

Table 23. Parameters Used in Sequential Batch-Leaching Models

The initial distributions of cations on the exchange sites were determined by assuming the rock sample was in equilibrium with the ground water sample collected at the Crescent Junction Site (ground water sample from Borehole 210 collected November 7, 2005). The composition of this water (Table 24) was equilibrated with the solid sample to fix the initial distribution of cations on the exchange sites.

Table 24. Composition of Borehole 210 Ground Water Collected November 7,2005

U.S. *Department of* Energy June 2006

Geochemical Attenuation and Performance Assessment Modeling Doe. No. X0173400 Page 30

To calibrate the distribution of U-bearing surface adsorption complexes, the U concentration in the Borehole 210 ground water sample was adjusted until the U concentration in the model matched reasonably well with the labile U released (Table 14) during the sequential batch-leaching tests. The concentration of U required to match well with the labile U release was 0.2 mg/L. To achieve the same concentrations of U-bearing adsorption sites for the Crescent Junction reaction-transport modeling, this U concentration was scaled to 6.7 µg/L to account for the 30-fold difference in rock-to-water ratio.

After initializing the cation exchange sites and adsorption sites, a sequential batch-leaching test was simulated by reacting SPF-1 fluid with appropriate mineral phases in a step-wise fashion. For the horizontal sequential batch-leaching scenario, three reaction steps were conducted using minerals representing samples collected from the 40-ft depth Mancos Shale. For the vertical scenario, SPF-1 was first reacted with the 40-ft mineral assemblage followed by four additional reaction steps with the mineral phases representing progressively deeper Mancos Shale.

As discussed in Section **0,** the results of the horizontal sequential batch-leaching tests showed several distinctive concentration trends. For example, concentrations of Na, $SO₄$, and U increased, while NH₄ decreased and Ca remained constant. These trends are simulated well with the model (Table 25). Modeled pH values were less than measured results and had a reverse trend (Table 25). Modeled alkalinity values (not listed in the table) were higher than the measured values. The inconsistencies in pH and alkalinity values are attributed largely to addition of bicarbonate ion to the solution from the dissolution of nahcolite in the model. As the bicarbonate component increases, calcite precipitates, causing pH values to decrease.

Comparison of modeled and measured results of the vertical sequential batch-leaching test shows trends similar to the horizontal scenario (Table 26). Unlike the horizontal model, pH values show an increasing trend after reaction step 3, and Ca has a decreasing trend. In both the horizontal and the vertical models, the observed trends result from interactions among several chemical processes that transfer mass between the solid and liquid phases, including mineral precipitation/dissolution, cation exchange, and specific adsorption for U. Considering the complexity of the system, the ability to simulate the measured results is quite good.

Table 25. Comparison of Measured and Modeled Results of Sequential Batch-Leaching Test for Horizontal Transporta

^aModeled results in parentheses.

Table 26. Comparison of Measured and Modeled Results of Sequential Batch-Leaching Test for Vertical Transport'

^aModeled results in parentheses.

U.S. Department of Energy June 2006

Transport Modeling

Two modeling scenarios of the Crescent Junction Site are presented: (1) horizontal flow through the upper Mancos Shale (depths of less than 40-ft below ground surface) and (2) vertical flow through deeper Mancos Shale (depths greater than 40-ft below ground surface). Input files for the two scenarios are provided in Appendix G. Parameters used for the transport models are similar to those used for the sequential batchleaching models (Section 0) but are scaled to account for the different rock-to-water ratio. Table 27 presents a summary of the parameter values. Modeling was conducted using the PHREEQC code (Parkhurst and Appelo 1999).

Table 27. Parameters Used in the Crescent Junction Reaction-Transport Models

Transport Model Setup

The horizontal flow simulation involves the flow of tailings pore fluid, represented by SPF-1 (Table 9), through shallow Mancos Shale, represented by the mineralogy of the 40-ft samples. For the vertical transport simulation, SPF-1 water flows through deeper Mancos Shale, represented by the mineralogy of samples deeper than 40 ft. For both simulations, the compositions of cation exchange sites and specific adsorption sites were set by equilibration with Borehole 210 (BH210) ground water. BH210 ground water was equilibrated with calcite and gypsum prior to equilibrating the exchange/adsorption sites. Sufficient U (6.7 µg/L) was added to BH210 ground water to match the calibration condition for composition of U-bearing surface-complexation sites determined from the sequential batch model (Section 0).

Transport was simulated in one dimension (similar to flow through a laboratory column or flow along a flow system streamline). The model domain consisted of 20 equally spaced cells and water flows through the domain by invoking 200 "shifts." Each shift transports water through a cell; thus, a total of 10 pore volumes (a pore volume being the volume of water within the fully saturated domain) were modeled. Pore volume is related to ground-water travel time or distance, if the ground water flow velocity is known. For simplicity, dispersion and diffusion were not included in the simulations although PHREEQC includes these capabilities.

Transport Model Results

For both simulations, NH₄ is retarded for about 3.5 pore volumes, after which concentrations increase rapidly to the influent concentration (Figure 5 and Figure 6). Retardation of NH₄ is caused by the exchange of dissolved NH₄ cations for solid-phase Ca, K, Mg, and Na. Most of the exchange involves Na ions because Na dominates the cations in BH210 ground water (Table 24). Thus, the cation exchange sites on the model Mancos Shale are initially dominated by Na. The compositional variation in the cation exchange

sites caused by the different mineral assemblages in the horizontal and vertical simulations did not significantly effect retardation of NH₄.

Effluent U concentrations in the horizontal simulation are retarded for slightly less than a pore volume during which the U concentration is less than about 0.075 mg/L (Figure 5). After 1 pore volume, U concentrations increase rapidly and reach the influent value (4.0 mg/L) after about 3 pore volumes. In the vertical simulation, U concentrations are never less than 2.4 mg/L (Figure 6). However, U concentrations in the vertical simulation remain less than the influent until about 4 pore volumes.

Uranium retardation is simulated by specific adsorption to the Mancos Shale. Specific adsorption is modeled using the U surface complexes listed in Table 21. Using this approach, the distribution coefficient (K_d) is not held constant but varies throughout a simulation. K_d values are not used explicitly in the simulations but were computed and are plotted on Figure 5 and Figure 6.

In the horizontal simulation, K_d values were initially 0.26 mL/g and decreased to 0.19 mL/g after about 2 pore volumes. In the vertical simulation, K_d values were negligible initially and increased to 0.15 mL/g after about 4 pore volumes. The differences in K_d values and dissolved U concentrations between the two simulations are products of the variation in the ionic composition of the solutions resulting from equilibration with minerals and cation exchange sites. A critical factor was the presence of a higher concentration of gypsum in the horizontal simulation than in the vertical simulation. Calcium released from dissolution of gypsum in the horizontal simulation caused precipitation of calcite, which resulted in decreased pH values. Initial Ca concentrations in the horizontal simulation were about 850 mg/L compared to only 9 mg/L in the vertical simulation (Figure 7 and Figure 8). These conditions led to increased partitioning of U to specific adsorption sites and increased retardation.

Values are scaled for plotting purposes using the multipliers shown in the legend. K_d values (mL/g) are calculated as the ratio of adsorbed concentration to dissolved concentration.

Figure 5. Effluent (Cell 20) Concentrations of U (mg/L), NH₄ (mg/L), and Uranium K_d (mL/g) and pH Values for the Horizontal Transport Simulation

 $C41$

Figure 6. Effluent (Cell 20) Concentrations of U (mg/L), NH₄ (mg/L), and Uranium K_d (mL/g) and pH Values for the Vertical Transport Simulation

Figure 8. Effluent (Cell 20) Concentrations of Major Ions for the Vertical Transport Simulation

Elevated initial SO₄ concentrations in the horizontal simulation are caused by gypsum dissolution, and lower initial dissolved carbon (IV) concentrations result from calcite precipitation. Chloride concentrations are sim to affect it.

In summary, the transport results suggest that NH₄ migration is retarded by several pore volumes. Uranium is retarded by about 1 pore volume but only if Ca is released from gypsum dissolution.

It is beyond the scope of this investigation to estimate ground water velocity beneath the proposed disposal
cell at the Crescent Junction Site; thus, results of the model simulations were provided in terms of pore
volume equivalent to 2 pore volumes after 1 year. After 2 years, the concentrations at the 10-ft and 5-ft distances are equivalent to 2 and 4 pore volumes, respectively, and so on. If the porosity of the Mancos Shale aquifer is significantly different than the value (0.25) used in the simulations, then model input would need to be refor

Model Limitations

The models presented in the previous section couple many of the chemical processes that are likely to affect transport of constituents at the Crescent Junction Site. Model limitations include both hydrologic and chemical

Ground water flow modeling with PHREEQC is limited to one dimension. Thus, models are constrained to a
single streamline. Although the modeling code is able to simulate dispersion and diffusion, these processes
were not in

 -51

ground water transport through the Mancos Shale is likely to be through fractures or other large-scale features. Such features are not explicitly considered in the models presented in this study. Rather, the flow
regime is implicit by the values used to assign concentrations of minerals, exchange sites, and adsorption sites. Because we based these values on the results of laboratory determinations of surface area and bulk density, we made an implicit assumption of pseudo porous-media flow. Adaptation of the model to fracture flow would be accomplished by decreasing the concentrations of sites and minerals (normalizing to a liter of ground water).

The ionic strength of the tailings pore fluids is higher than is commonly prescribed as a limit for use of the Debye-Hückle theory used in PHREEQC to calculate activity coefficients. The ionic strength limitation could be improved by employing empirical activity coefficient algorithms, such as the Pitzer equations (Mariner 2001). Unfortunately, these activity models exist for some of the major ions but are not available for U.

As a first approximation, oxidation-reduction (redox) processes Were ignored in the transport models. Although the PHREEQC code is capable of calculating redox equilibria, many redox processes are rate limited. Rate-limited processes can be incorporated in PHREEQC, but universally applicable kinetic data needed to apply these processes are currently unavailable. Another reason for omitting redox processes in the models is that no data are available to confirm that redox processes are occurring in the Mancos Shale at the Crescent Junction Site. Scattered occurrences of pyrite and carbonaceous material suggest that redox processes could be occurring. However, some data suggest relatively oxidized conditions as deep as 300 ft (the depth of the boreholes at the site). Ground water collected from six of these boreholes has redox potentials ranging from 234 to 442 millivolt (mv) with one exception of 15 mV (SEEPro database). The two values of dissolved oxygen that were determined on these same ground water samples have values (2.41 mg/L and 2.8 mg/L) consistent with oxidized conditions. Also, scattered occurrences of oxidized iron (limonite) staining are present in the upper portion of the Mancos Shale. If data collected later confirm the presence of redox processes, these processes are readily stimulated by the reaction-transport model.

Other chemical factors that limit the modeling are omission of solid solutions and rate-limited processes. These are omitted from the working model because of lack of data to support a viable analysis rather than lack of model capabilities.

Conclusions

The water-soluble mineral assemblage of the Mancos Shale at the' Crescent Junction site is likely dominated by nahcolite with major amounts of Ca-Na exchange and gypsum. Halite, sylvite, and dolomite occur in lesser amounts and calcite is present. Gypsum and dolomite are more dominant in the shallow .samples, while nahcolite and halite are more concentrated in the deeper samples. Cation exchange capacity of the Mancos Shale ranges from 0.54 to 36.29 meq/100 g with a mean of 11.23 meq/100 g. The bulk mineral assemblage in the Mancos dominated by quartz, with lesser amounts of dolomite and calcite, small amounts of feldspar, and traces of gypsum. The clay mineral fraction is dominated by mixed-layer (mostly illite/smectite) clays, illite, and kaolinite with illite layers dominating the mixed-layer clays. Particle surface area ranges from 8.81 to 13.22 m²/g with a mean of 11.02 m²/g. Distribution ratios for uranium adsorption on Mancos Shale are low, ranging from essentially 0 to 0.84 mL/g.

As tailings fluids reacted with progressively more Mancos Shale in sequential batch tests, slight increases were observed in pH, and concentrations of Na, K, Cl, SO₄, and U. Results of a reaction-transport model of sequential batch tests showed reasonably good agreement with observed concentrations for Na, Ca, SO4, NH4, and U; but the modeled pH values were slightly less than the observed values. Considering the complexity of the chemical interactions, these results were considered favorable; thus, this model was used in a one-dimensional simulation of contaminant transport beneath the Crescent Junction disposal cell. The transport results suggest that NH4 migration is retarded by several pore volumes. Uranium is retarded by about 1 pore volume, but only if Ca is released from gypsum dissolution. To evaluate the effects of this analysis on contaminant transport beneath the proposed Crescent Junction disposal cell, it is necessary to know the flux of contaminated water from the cell and the effective porosity of the Mancos Shale. Evaluation of the flow regime was beyond the scope of this investigation; thus, results of the model simulations are provided in terms of pore volumes. To maximize the benefit of these results in the field setting, project personnel will need to couple these results given in units of pore volume, with the results from hydrologic investigations to yield more useful units'of travel time and distance. Alternatively, a sensitivity analysis that

uses reasonable bounds for the hydrologic parameters may be appropriate to assess the impact of chemical attenuation at the Crescent Junction site.

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Appendix A

(,) Copy of "Work Plan Geochemical Analysis of Mancos Shale Cores from the Crescent Junction Disposal Cell Site"

DOE EM/GJ 1118-2006

Work Plan Geochemical Analysis of Mancos Shale Cores from the Crescent Junction Disposal Cell Site, Moab UMTRA Project

January 2006

Work Performed by S.M. Stoller Corporation under DOE Contract No. DE-AC01-02GJ79491 for the U.S. Department of Energy Office of Environmental Management, Grand Junction, Colorado

Contents

Tables

Table 1. Analytical M ethods **...** 2

K.)

Geochemical Analysis of Mancos Shale Cores, Crescent Junction, Moab UMTRA Project Doc. No. X0134000 Page ii

1.0 Introduction

This work plan details work that will be done to characterize the geochemistry of the bedrock beneath and adjacent to the proposed disposal cell at Crescent Junction, Utah. The disposal cell will contain tailings from the Moab, Utah, (Atlas) uranium mill and will be underlain by Mancos Shale Formation. The purpose of the work is to provide data that will help evaluate the potential for ground water contamination and transport by constituents in the tailings. The basis for this work is provided in Section 4.5 of Department of Energy (DOE) (2005) and has been modified based on discussions with Moab UMTRA Project personnel. The scope includes laboratory investigation and geochemical modeling, and is presented as nine individual tasks.

The geochemical approach involves collecting site-specific data that can be used to model geochemical interactions between tailings pore fluid and the Mancos Shale. Results from this work will provide the following information about the Mancos Shale:

- **9** Abundance and mineralogy of water soluble minerals
- Mineralogy of water insoluble minerals, including clay mineralogy
- **9** Cation exchange capacity
- Surface area
- Chemical distribution ratios (R_d)

The data will be used to construct a one-dimensional coupled hydrogeochemical model of **•,)** tailings water transport through the Mancos. The model will include equations governing aqueous speciation, mineral dissolution, mineral precipitation, mixing with other ground water, oxidation/reduction, cation exchange, and adsorption. In addition, a sequential batch-leaching test will be conducted and the results used to help calibrate the geochemical model. Numerous analyses of tailings pore fluids have been made previously and no additional analyses will be made for the current work scope.

Samples of Mancos Shale will come from 10 cores that were collected from borings evenly distributed in the area of the proposed disposal cell (see DOE 2005 for locations). Samples will primarily consist of the Blue Gate Shale Member of the Mancos Formation because this unit is the most likely to receive contaminated drainage from the disposal cell. Some samples may also be collected from the Prairie Canyon Member of the Mancos. Approximately five samples will be collected from each of the 10 cores for a total of 50 samples. The disposal cell will be excavated approximately 20 ft into the ground; thus, the uppermost sample will be collected from a depth of about 40 ft. The remaining four samples (per core) will be collected at equal intervals below 40 ft. These samples should provide a reasonable set of data to evaluate lateral and vertical distribution of geochemical properties in the Mancos.

2.0 Task Descriptions

This section provides details of the' nine individual laboratory and modeling tasks. The laboratory portion of the work will be conducted in the Applied Sciences and Technology (AST) testing

facility at the Grand Junction site. Equipment available in the AST facility that will be used for this work include:

- **0** Ion chromatograph
- **0** PH/ORP/conductivity probes and meters
- **S** Balances
- **0** Shaker tables
- Atomic absorption spectrometer
- Solids/liquid separation equipment
- Drying ovens
- Specialized spectrometers and other sample analysis equipment

A subcontract will be procured with Mesa State College to conduct X-ray diffraction (XRD) analysis. Particle surface areas using BET will be measured by a contract laboratory. Analytical methods are presented in Table 1.

Table 1. Analytical Methods

2.1 Task 1 - Preparation and Sample Collection

Task 1 includes procurement of chemicals and supplies, many of which are already available in the AST testing facility. Subcontracts with Mesa State College for XRD analysis and a contract laboratory for BET analysis Will be procured.

Five samples will be collected from each of the 10 cores. The samples will be collected at depths of 40 **ft,** 105 **ft,** 170 **ft,** 235 **ft,** and•300 ft. Each sample will contain approximately 3 linear inches of core. Samples will be air-dried and stored in plastic containers until used for the tests. A sample log will be prepared that contains sample numbers, bulk weights, and moisture contents. Dried samples will be lightly crushed with a hammer or pestle, sieved, and split using a riffle splitter as needed for the tests.

2.2 Task 2 - Analysis of Water Soluble Fractions

Task 2 is designed to identify and estimate abundance of minerals present in the water-soluble fraction of the Mancos. Mineral identification will be aided by the XRD work in Task 4; however, XRD is limited in its ability to detect small amounts (less than about 5 percent) of

mineral and does not provide information on mineral abundance. All 50 samples will be analyzed. Water-soluble minerals are likely to include thenardite $(Na₂SO₄)$, halite $(NaCl)$, gypsum ($CaSO₄$.2H₂O), and other Na-Ca-K-Mg-SO₄-CO₂-NO₃-Cl-bearing salts. Analyses will include major ions that make up these minerals as well as mill tailings contaminants NH_3 and U.

Samples will be crushed, air dried, and sieved to $-10 + 18$ mesh (1 to 2 mm nominal diameter). The size fraction was selected because: (1) it is likely to have sufficient material to accomplish the testing, (2) it is uniform enough so that surface area is relatively constant and easy to determine, and (3) it is efficient to work with in the laboratory (testing apparatus can be relatively small).

The leaching procedure is modified from DOE 2005b, procedure CB (BT-l). Two grams of the sieved sample is placed in a plastic centrifuge tube with 100 mL of deionized water at room temperature. The tube is agitated end-over-end for four hours. The solids are separated from the water by centrifuging and decanting, and/or filtering to produce a clear solution. The solution is then analyzed for pH, ORP, specific conductance, alkalinity, Ca, Na, Mg, K, SO_4 , Cl, NO₃, $NH₃$, and U.

The solid-phase concentration of each leachable major ion will be calculated from the measured concentration and the solid/solution masses used. These solid-phase concentrations will be cast in mole units and mineral stoichiometry will be used to estimate abundances of soluble minerals. The geochemical computer program NETPATH (Plummer et al., 1994) may be used to help determine possible mineral mixtures. In NETPATH, the user can define the initial solution by the composition of the test leachate and the final solution as pure water. Various combinations of mineral phases can be specified and the program calculates the amount of each phase that must precipitate to meet the compositional constraints. Results of'X-ray diffraction will help to identify possible mineral phases. Solid-phase concentrations of the contaminants NH₃ and U will also be calculated; however, concentrations are expected to be small and identification of the mineral phases containing these constituents will not likely be possible.

2.3 Task **3 -** Measurements of Cation Exchange Capacity **(CEC)**

The Mancos Shale contains abundant clay minerals. Many types of clay have large CEC that can cause significant changes to ground water chemistry. Knowledge of the CEC is required to develop a geochemical model of water rock interactions. Therefore, CEC will be measured on 20 core samples of Mancos.

Samples will be crushed, air dried, and sieved to **-10** +18 mesh (I to 2 mm nominal diameter) in the same manner as for other tests. The CEC will be determined using a standard method such as the Ca/Mg exchange method (Jackson, 1969) or the ammonium saturation method (Chapman, .1965 as described in Bain and Smith, 1987). The choice of method will be determined after a thorough review of the applicability to the Mancos cores. For the Ca/Mg exchange method, approximately 0.5 to 2 g (exact weight depends on expected CEC) of Mancos is saturated with Ca²⁺ using 0.5 N CaCl₂ solution. The Ca²⁺ ion is then replaced with Mg²⁺ using 0.5N MgCl₂, and the concentration of $Ca²⁺$ released is used to calculate CEC. For the ammonium saturation method, the clay sample is first saturated with ammonium ion using one molar ammonium acetate followed by exchange with sodium chloride. CEC will also be measured on one powdered Mancos sample to provide information on maximum CEC.

2.4 Task **4-** X-ray Diffraction Analysis

Chemical interaction between ground water and Mancos Shale is likely to occur mostly at the surfaces of clay minerals. Therefore, clay mineral chemistry is important to the transport of contaminants by the ground water. XRD is one of the best analytical tools to identify the mineralogy of the clay mineral fraction of the Mancos. Ten core samples will be selected for XRD analysis.

Mancos samples will be finely powdered with a mortar and pestle and fractions will be separated by suspension in Calgon solution to concentrate the clay minerals. These clay-mineral separates will be treated in four different ways prior to XRD analysis (Wilson, 1987): (1) air dried, (2) glycolated, (3) heated 300 **"C,** and (4) heated 550 **'C.** XRD patterns from these runs will be used to identify clay minerals such as illite (I), smectite (S), interlayered **1/S,** chlorite, and kaolinite.

X-ray diffraction (XRD) will also be used to determine the major minerals present in the core samples. For these analyses, the 10 core samples will be powdered and oriented randomly.

2.5 Task 5 - Determine Surface Area of Mancos Shale

Knowledge of the surface area of the Mancos Shale samples used in the testing is needed to relate the results to the transport of contaminants through the subsurface. Processes such as cation exchange and adsorption are directly related to the surface area that the ground water contacts. For example, for the same travel distance, interaction of dissolved contaminants with the rock will be much less in a fracture-dominated matrix than for porous-media flow. It is beyond the scope of this work plan to determine the nature of the flow (fracture-verses-porous media) in the subsurface at the Crescent Junction site; however, to properly use the data collected during this study in site models, they will need to be normalized to surface area. Therefore, it is important to measure the surface area of the samples used in the testing.

Surface area will be determined for a subset of 10 core samples. Samples will be crushed, air dried, and sieved to $-10 + 18$ mesh (1 to 2 mm nominal diameter) in the same manner as for other tests. Surface area will also be measured on a powdered Mancos sample to provide information on maximum surface area. Surface area will be determined by the standard Brunauer, Emmett, and Teller (BET) method. This method uses N_2 gas adsorption isotherms at -183 °C to measure surface area (Jackson 1969).

2.6 Task **6-** Determination **of** Distribution Ratios

Distribution ratios (R_d) provide a measure of the partitioning of a contaminant between the ground water and the solid constituents that comprise the aquifer. The higher the R_d , the more partitioning to the solids and the more retardation. The R_d value is an empirical value that is simply the ratio of the measured concentration in the solids (mg/kg) to the measured concentration in the ground water (mg/L) and has units of mL/g. The R_d values are often used to simulate retardation of contaminants in ground water models. When used in these models there is an implicit assumption that chemical retardation occurs by the process of adsorption under equilibrium conditions. For the equilibrium assumption, the R_d value is often referred to as a

 K_d value. Many models also assume that K_d does not vary with the concentration of the contaminant. However, research has shown that this variation sometimes exists and algorithms such as the Langmuir or Freudlich equations are used to produce a-better fit to the R_d data. To test if R_d varies with contaminant concentration, multiple points using various concentrations are measured and plots of these "adsorption isotherms" are fitted with the various models. If the plot is reasonably linear, then a K_d (also termed linear isotherm) will produce satisfactory results.

Rd values will be determined using DOE (2005b) procedure CB (Rd-1). In summary, a sample of Mancos Shale is crushed and sieved to -10 +18 mesh **(I** to 2 mm nominal diameter) as in Task 2. The sample $(5 g)$ is placed in a plastic centrifuge tube with 100 mL of synthetic pore fluid (SPF) that simulates tailings pore fluid. The R_d value can vary significantly with solution chemistry. For example, Davis et al. (2004) showed that R_d values for U in a sample of alluvium varied by more than a factor of 10 depending on dissolved carbonate concentration and pH value. **Rd** values will be determined using three SPF compositions that are designed to simulate ground water that could be present in the Mancos Shale after construction of the disposal cell. One solution (SPF-1) will simulate pore fluids that currently exist in the tailings. Another solution (SPF-2) will simulate tailing pore water mixed with 50% Mancos ground water. A third solution (SPF-3) will simulate water that results from the sequential batch-leaching test discussed in Task 7.

Single-point R_d measurements for U will be made on 20 samples and 6-point isotherms will be measured on five of these using SPF-1. For the five samples used for isotherms, single-point Rd values will also be determined using SPF-2 and SPF-3 to evaluate sensitivity to solution chemistry. The R_d approach is only valid for contaminants that occur in trace concentrations.
Due to the high concentrations present in the pore fluid, modeling the transport of NH₃ by adsorption is not valid; thus, R_d values for NH₃ will not be measured.

2.7 Task **7 -** Sequential Batch Leaching Test

Two sequential batch-leaching tests will be conducted to "measure evolution of tailings leachate", chemistry as it interacts with Mancos Shale. For this study, it is assumed that the major transport path is either vertically downward beneath the tailings pile or subhorizontal through the weathered zone. Results of this test will be used to help validate the coupled hydrogeochemical transport model presented in the next section.

For the vertical scenario, a 100-g sample of Mancos (40-ft depth) will be crushed, air dried, and sieved to -10 +18 mesh **(I** to 2 mm nominal diameter). The sample will be combined with 400 mL of SPF-1 in a 500-mL glass Erlenmeyer flask. The mixture will be agitated on an orbital shake table for 24 hours. At that time, the solids will be separated from the liquid by centrifugation, decantation, and/or filtering. A 50-mL split will be retained and analyzed for pH, ORP, specific conductance, alkalinity, Ca, Na, Mg, K, SO₄, Cl, NO₃, NH₃, and U. The remaining solution will be placed with approximately 87.5 g of Mancos sample from 105-ft depth. The exact amount of Mancos sample will be calculated so the water-to-rock ratio remains constant. The procedure will be repeated three times with progressively deeper samples.

For the horizontal scenario, a 3-step leach will be conducted using the 40-ft deep samples from cores 203,204, and 206.

2.8 Task **8 -** Coupled Hydrogeochemical Transport Modeling

This task consists of developing a coupled hydrogeochemical transport model using the PHREEQC code (Parkhurst and Appelo, 1999). Ion exchange will be modeled using data from the CEC tests in Task 3 and surface area measurements in Task 5. Initial conditions will include concentrations of minerals as estimated from the results of Tasks 2 and 4. A surface complexation model for U, based on R_d measurements (Task 6) will be incorporated.

The model will include one-dimension transport of tailings water through the Mancos Shale. Aqueous speciation reactions are typically fast with respect to ground water flow and will be modeled at chemical equilibrium. Water-rock interaction will include mineral precipitation and dissolution, adsorption, and cation exchange. The model will be adaptable to all6w inclusion of such factors as: **(1)** mixing with ground water, (2) reaction kinetics, and (3) changing redox state (e.g., due to biologic activity). Thus, sensitivity of the transport to various parameters can be readily estimated with additional model simulations. Two model simulations Will be conducted to simulate the vertical and horizontal scenarios described in Section 2.7. The model can be used to simulate other transport paths with minor changes.

2.9 Task **9 -** Reporting

Results of the testing and modeling will be presented as Moab Calculation Sets.

3.0 Schedule

24.0 References

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Appendix B

Laboratory Notes

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Task 4 $CJO1-O1-O6$ $X \mathcal{L} \Delta$ 1/17/06 Place aux approx 150 split of the following sumples in Simple $LCI - 201 - 40$ $202 - 40$ $203 - 40$ $204 - 40$ $205 - 40$ $206 - 40$ $307 - 40$ $208 - 40$ $309 - 40$ $210 - 40$ Deliver to D. Bill Hood for XRD analysis

 $CJ01-01-07$ Crescent Jets. Task5 - BET Surface aux. Prépare the following Binifiles for Alument to Meromereties for $1/17$ lo 6 Particle piz Semple LJ 201-40 L 2 ∂ 02--40 ノーブ $203 - 40$ $1 - 2 204 - 40$ $1 - 3 205 - 40$ 1-2-(pourdered) $205 - 40P$ $2/$.706-40 /- ブ $207 - 40$ /-み $208 - 40$ $1 - 3 209 - 40$ $1 - 3 \partial 10 - 40$ ノーブ tra Kdex Sind to Micromenties analistical Irvices Que Micromerítics DS Narcross, GA 300937877 $1 - 770 - 662 - 3630$ Multipunt surpacience analizus = Nigas Request Cat#OOS-01 all Micromenities re analysis - 1-2mm enc pac compueste to pewdered sample.
Confirmed samples run as pullmetted with pewdered) wij puone conversation
with analyst that perfermed trits. He surpreted in meccule was smiller 2/2/06

Sarah Morris From: Stan Morrison Sent: Thursday, February 02, 2006 9:35 AM To: Sarah Morris Subject: FW: Micromeritics Analytical Services - Analysis Results Sarah, Surface Area results. I've included an Excel sheet.("Results_Surface Area") with the results tabulated (in the format your original data table). Of interest, the powdered sample was very similar to the coarser sample. Could you contact them and confirm that the samples were not powdered prior to analysis. Thanks, **Stan -** Original Message--- From: michael.poston@micromeritics.com [mailto:michael.poston@micromeritics.com] **Q** Sent: Thursday, February 02, 2006 7:49 AM To: Stan Morrison Cc: mark.talarico@micromeritics.com Subject: Micromeritics Analytical Services - Analysis Results Dear Stan Morrison, *•J* Thank you for submitting your samples to Micromeritics Analytical Services. We appreciate the opportunity to serve you. Attached below, please find the results for your samples and a copy of the submission form (for future submissions). If you have any questions, comments or concerns please contact us at 770-662-3630. Sincerely, Michael Poston - Lab Analyst Micromeritics Analytical Services **1** Micromeritics Dr. Norcross, GA 30093 770.662.3630 www.particletesting.com *This e-mail is intended for the use of the addressee(s) only and may contain privileged, confidential, or proprietary information that is exempt from disclosure under law. If you have received this message in error, please inform us promptly by reply e-mail, then delete the e-mail and destroy any printed copy., Thank you.*

2/2/2006

Sarah Morris

rom: Stan Morrison Thursday, February 02, 2006 1:02 PM Sarah Morris -oubject: RE: Micromeritics Info

Sarah,

Thanks. That satisfied my concern. Please put the response in the project file.

Stan

I just called and talked to the person that did the surface area analysis. He confirmed that the samples were analyzed as instructed, that the samples were not powdered prior.

When I told him your concern that the surface area of the 1-2 mm fraction was comparable to the powdered split, his response was that "it is possible that the N2 molecule was smaller than the interstitial space of the 1-2 mm fraction and was seeing "inside" the particles." He said that he was probably not the one to converse with about that he was probably not the one to converse with about that possibility. If you wanted to call and discuss, the best person would be Greg Fields, their business manager who would probably be knowledgeable, but if he didn't, would know who to put you in contact with. Their number is 770-662- 3630.

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Sarah

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(start 315-290 mL = 25 mL retained unoverds) 425 MC setuded in Aolids. CJ 205 e 170 $+2$ H Go g/m 24 ml residual SPFI CT 206 $e40 + 24$ 60 g in 240 m Denduce SPFI 0900 Return to arbitral straker

22-210 \mathcal{D} $CJ01-01-19$ CJ Task 7 Remode from attital shaper. Aurile Vol & Aolids between Lea 250 ml nalge 0745 2/13/2006 Place in centrique essodern x20" Decant and filter through 0.45 um filter Vol = 215 ml (star 240-25 ml = 25 ml retained) Herizontil fest is complete! Retail all for a undeds) Herefortil fat is complete | Retail all for analyses |
Vertifial test continues | Retain 50mL | 2050 | 76 | arralyses CT 30507 35 -2/71 4/259 in 165 ml residuel SPF1 Retigento creital shaker. Millolo ogoo Remiere from orbital shaker Durice vol: solies tielween sea 125ml pulge Place interntique essaprem à 20" Decant and piller through 0.45 un filter Vol= 145 ml (start 165me-145 - 20 ml retimed madido) Return Home CT2050235 for anneysis $452056300 - 2/7$ $35g/m$ /00 ml residuel SPF Return to orbited shaker Measure SPFI prep 2/9/06 pt, ore, det conf 3/15/05 0830 Remione from arrital philiar decant and filter thru 0.45 cm/ filter. Refain all for analysis

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Appendix **C**

Complete Chemical Results and Computations

- (1) *Loss on Drying*
- *(2) Size Fraction Analysis*
- *(3). X-Ray Diffraction Results*
	- *(4) Surface Area Results*
- *(5) Cation Exchange Capacity Results*

(6) Water Soluble Minerals - *Raw Data*

- *(7) Water Soluble Minerals NETPATH Results*
- *(8) Water Soluble Minerals PHREEQC Input*

(9) Distribution Ratios: Single-Point, SPF1

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- (10) *Distribution Ratios: Single-Point, SPF2*
- (11) *Distribution Ratios: Single-Point, SPF3 (12) Distribution Ratios: Multiple-Point Isotherms*

(13) Results of Sequential Batch Tests

Page **1** of 2

Page 2 of 2

Page **I** of 2

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Page 1 of I

Page **1** of I

Page **I** of I

Page 1 of I

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Single Point Kd's SPF 3 Mean Labile U (from Sequential Batch Tests)= 0.0436 ug/g Sample Mass UU U U Cs **CS** Cs Rd Rd Rd Log **U** Log Rd (g) **) (ug/L) (ug/L)** (ugOL) (ug/g) (ug/g) **(ug/g) mug mug mug** -(mol/L) mUgI |Meas |Max |Min |Meas. |Min |Max |Meas. |Min |Max | |Meas. |Meas. |Meas 2.5% Error CJ-202-40 5 742.2 760.76 723.65 1.4236 0.6469 2.2003 1.92 0.85 3.04 -5.50606 0.282867 CJ-205-40 5 756.7 775.62 737.78 1.1336 0.3496 1.9176 1.50 0.45 2.60 -5.49765 0.175536 CJ-207-40 5 783.4 802.99 763.82 0.5996 -0.1977 1.3969 0.77 -0.25 1.83 -5.48259 -0.11612 CJ-208-40 5 797 816.93 777.08 0.3276 0.4765 1.1317 0.41 -0.58 1.46 -5.47512 0.38611
CJ-210-40 5 793.9 813.75 774.05 0.3896 0.4130 1.1922 0.49 -0.51 1.54 -5.47681 0.30915 CJ-210-5 1793.9 813.75 774.05 0.3896 0.4130 1.1922 0.49 0.51 1.54 -5.47681 0.30915 SPF3¹

Page 1 of **1**

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SPF1 (Final) |7.85 |191.9 |29600 |510 |1018 |738 |18272 | | |196 |5300 |310 |129

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Appendix D

Dr. William C. Hood, 2006

) X-Ray Diffraction Analysis Determination of the Mineralogy of 10 Mancos Shale Samples

Prepared for S.M. Stoller Corporation, February 3, 2006 Grand Junction, Colorado
X-ray-determination of the mineralogy of 10 Mancos Shale samples

A report prepared by:

William C. Hood, Phi) AAPG Certified Professional Geologist No. 2185 February 3, 2006 $\ddot{}$

X-ray determination of the mineralogy of **10** Mancos Shale samples

Introduction

At the request of Stan Morrison, ten samples of Mancos Shale from the proposed Atlas Mill tailings disposal site near Crescent Junction were examined by x-ray diffraction techniques. The purpose of the examination was to identify the minerals present in the shale, with special emphasis on the clay mineralogy.

Suminary and Conclusions

The mineralogy of the ten samples is quite similar, as might be expected in a marine environment such as the Mancos Shale. The non-phyllosilicate minerals identified include quartz, calcite, dolomite, orthoclase, plagioclase and gypsum. These minerals are estimated to comprise from a third to one-half of the total sample. The clay minerals identified, in order of abundance, are mixed-layer illite/smectite (a little less than 38 percent, discrete illite (about 33 percent), kiolinite (about 28 percent) and mixed-layer chlorite/vermiculite (perhaps 2 or 3 percent). The mixed-layer illite/sniectite contains at least 60 percent illitic layers, but the actual amount could not be determined with accuracy. Kaolinite crystallites are the largest clay minerals in the samples at an average grain size of 297 A, followed by illite at **152** A and mixed-layer illite/smectite at 67 A.

Methods

The samples were supplied in pulverized form. A portion of each of the **10** original samples was ground for several minutes using a porcelain mortar and pestle. A five-gram sample of this material was removed for clay analysis. The technique used to separate clay minerals is described in Appendix **1.** About two grams of the remaining material was ground for several more minutes to obtain a powdered sample for determining the bulk mineralogy

The powdered bulk samples were loaded into sample holders and scanned from 4.2 to 50 degrees two-theta using the Rigaku Miniflex x-ray diftiractometer at Mesa State College.

Clay samples were scanned over the same interval so that direct comparisons of the clay patterns to the bulk samples could be made. Four scans were made of each clay sample: untreated, glycol solvated, heat treated to 300° C and heat treated to 550° C. Glycol solvation is accomplished by putting the samples into a desiccator that contains ethylene glycol and heating the desiccator in an oven at 60° C overnight. The purpose of this treatment is to expand the mineral smectite or sniectite layers within mixed-layer clays to approximately 17 Angstroms d-spacing, facilitating the identification of clays containing an expandable smectitic component. The two heat treatments were accomplished by placing-the samples into a muffle furnace and heating to the'appropriate temperature, as indicated by a thermocouple thermometer. **The300"** C heat treatment collapses the expandable layers to 9.2 Angstroms but does not destroy chlorite, chlorite layers in mixed-layer clays or kaolinite. The 550° C heat treatment destroys kaolinite, facilitating its distinction from chlorite.

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Results

Bulk samples. The minerals identified in the bulk samples are, in order of general abundance, quartz, dolomite, calcite, plagioclase, orthoclase and gypsum. The estimated abundances of these minerals are given in Table 1 and the method of estimating the abundances is explained in Appendix 2, which also includes diffraction patterns for quartz, calcite, dolomite, albite and gypsum standards.

Table I. 'Estimated abundances (in percent) of non-phyllosilicate minerals in bulk Mancos Shale samples.

Minerals were identified using the diffraction peaks listed in Appendix 3. The x-ray patterns for the bulk samples themselves are in Appendix 4. For each sample, two copies of the x-ray diffraction patterns are included, one with d-spacings for most of the peaks and one with most of the peaks identified as to which mineral they result from.

To verify the presence of the carbonate minerals, calcite and dolomite, a small amount of .sample CJ-205-40 was placed in a small container, mixed with water and then several drops of dilute hydrochloric acid were added. The sample effervesced vigorously, confirming the presence of carbonate minerals.. After the effervescence stopped, the sample was washed twice to remove the dissolved constituents, dried and x-rayed. The peaks associated with calcite disappeared and the dolomite peak was considerably reduced in intensity, confirming the presence of both minerals (Figure 1).

Clay separates.

Identification of clay minerals.

Examination of a typical x-ray diffraction pattern reveals the presence of four phyllosilicate and two non-phyllosilicate minerals in these samples. The two non-phyllosilicates are quartz and dolomite. The phyllosilicate minerals identified are kaolinite, illite, mixed-layer illite/smectite and mixed-layer chlorite/vermiculite. The following paragraphs explain how these minerals were identified.

Kaolinite is recognized by peaks at approximately 7.14 Angstroms, 3.57 Angstroms and 2.38 Angstroms. These correspond to the first, second and third order diffractions from the (001) plane. These spacings do not change when the sample is glycol-solvated or heated to 300° C. They disappear upon heating the sample to 550° C, indicating that the kaolinite has been destroyed. Kaolinite is present in all of the samples

Illite is recognized by peaks at approximately **10** Angstroms, 5.0 Angstroms, 3.3 Angstroms, 2.5 Angstroms and 2.0 Angstroms. These represent the first, second, third, fourth and fifit order diffractions from the (001) plane. Peak intensities are strongest for the odd orders, with the third order peak being the strongest, as is typical for a dioctahedral phiyllosilicate. In the case of the CJ sample set, the peak at 3.3 Angstroms **(3rd** order) is quite strong because it is combined with the most intense peak for quartz. The fourth order peak at 2.5 Angstroms is very weak, hardly rising above background in most of the samples. Glycol solvation shifts the peak of expandable minerals, making the **10** Angstrom peak quite obvious. Heat treatments bring about slight peak shifts and significant increases in intensity. These changes are not so much due to changes in the illite itself, but are mainly a result of the collapse of mixed-layer -clays to **9.9** Angstroms or slightly less and the diffraction from that material being added to that of the discrete illite.

Mixed-layer illite/smectite is recognized by a broad peak between about 11.2 and 11.9 Angstroms in the untreated sample. The peak is broad and there is enough "chatter" in the background that it is difficult to pick the center of the peak with any certainty. Another complication is that the peak overlaps the peaks from illite and chlorite. The additive effects of these peaks can cause the apparent peak position to shift. In the glycol-solvated sample, the peak shifts to higher d-spacing (lower. angle 2-theta or left on the diffraction pattern), indicating expansion of smectitic layers. In these samples, the shift does not usually result in a well-developed peak, but rather a significant drop in intensity near the 10 Angstrom illite peak and a smeared-out low hump between 14 and 17 Angstroms. Heating to 300' C collapses the expandable layers to approximately 9.9 Angstroms, resulting in a marked increase in intensity of the peak at approximately **10** Angstroms. This peak is asymmetric toward low angles, suggesting that not all of the layers are totally'collapsed. Heating to 550" collapses these layers and sharpens the peak-at approximately 10 Angstroms.

The presence of small amounts of a chlorite component in the CJ samples is recognized by a low, broad peak about 14.3 Angstroms and a sharper, third order peak, at 4.8 Angstroms, which appears as a shoulder or small peak on the high angle side of the second-order illite peak. The 14,3 Angstrom peak is small and sometimes is just a shoulder on the mixed-layer

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illite smectite peak. It persists through glycol solvation and low heat and does not seem to change position. although interference from mixed-layer illite/smectite makes the peak position indeterminate. The peak remains after heating to 550° C but shifts to lower dspacings, typically about 13.5 angstroms. The shift indicates that the material is not a pure chlorite, but rather a mixed-layer material. To summarize its characteristics, it does not seem to expand upon glycol salvation and collapses somewhat with high heat. Based on the material presented in Moore and Reynolds (1997) the mineral is probably a mixed-layer chlorite/vermiculite.

Figure 2 shows the results of the various treatments for sample CJ201-40 combined into one display to show how the patterns change with vairious treatments. Stacked patterns for all ten of the samples are contained, in Appendix 5. The individual x-ray patterns are located in Appendix 6.

Estimating the relative amounts of clay minerals.

To estimate the relative amounts of the various clay minerals, the patterns of the glycolsolvated samples were used. By expanding the smectitic layers in mixed-layer clay, this treatment separates peaks of clays with expandable components from the 10 Angstrom peak of illite. Although the glycol-solvated pattern allows the separation of discrete iltite, it does not allow the separation of the mixed-layer illite/smectite from the mixed-layer chlorite/vermiculite.

The first step is to draw a smooth background through the low-angle portion of the x-ray pattern. Figure 3 shows what this looks like. Once the background curve is established, the Jade computer program was used to eliminate the background. Appendix 7 contains the patterns after this step. The next step was to use the peak-painting module within Jade to determine the area under the peak for each of the first-order peaks of mixed-layer clay, discrete illite and kaolinite. Figure 4 shows the painted peaks of a typical sample and Table 2 gives the'results of these measurements.

Table 2. Areas under the eaks, as calculated by the Jade computer program.

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Finally, the areas under the peaks of the three minerals were added and the relative amounts of the components calculated as a percentage of the total. These values are given in Table 3.

Table 3. Estimated abundance of individual clays, derived from areas under the peaks.

Note. Mixed-layer includes both mixed-layer illite/smectite plus a much smaller amount of mixed-layer chlorite/vermiculite. Mixed-layer chlorite/vermiculite mnakes up perhaps 5 percent of the mixed-layer component.

The values in Table 3 should be taken only as semi-quantitative approximations. To obtain more quantitative values, standards of similar grain size and crystallinity would have to be mixed in various amounts and x-rayed under the same conditions as the Mancos Shale samples. Such standards were not available. The values that were obtained in this study are, however, useful for inter-sample comparisons.

Little published work exists on the clay mineralogy of the Mancos Shale in the Colorado Plateau region. L. G. Schultz, cited in Eckel. (p. 276) reports that the clays consist of 50 to 60 percent mixed-layer illite/smectite, 12 to 15 percent poorly ordered kaolinite, and 30 to 35 percent discrete illite. No information was given about the method used to determine these amounts. The **CJ** samples contain more kaolinite and less mixed-layer illite/smectite than reported by Schultz.

Amount of illite in mixed-layer illite/smectite

The amount of illitic and smectitic components in the mixed-layer illite/smectite is very difficult to determine in this sample set due to interference among the peaks of the various minerals. The usual procedure for determining the amount of illite and smectite in a mixedlayered illite/smectite is to use the glycol solvated pattern and determine the peak positions for the (001)/(002) and (002)/(003) peaks. Unfortunately in the **CJ** sample set, the $(001)/(002)$ peak (between 10.3 and 9.0° 2-theta) lies on the flank of a much stronger illite peak and cannot be resolved. Likewise, the (002)/(003) peak (between 15.8 and 17.4' 2 theta) lies on the flank of the second order illite peak and just appears as a broadening of the illite peak and not as a distinct peak itself. However, this fact by itself suggests that the material is richer in the illite component than the smectite component. Materials with large amounts of smectite have peaks at the low angle end of the range, whereas peaks with a large

9

amount of illite component have peaks at the high end of the range, close to the peak for pure illite

In order to get an independent approximation as to the amount of illite layers in the mixedlayer illite/smectite, I examined the position of the (001)/(001) peak in the Mg-saturated but otherwise untreated sample. Pure smectite would have a peak at 14 Angstroms and pure illite a peak at 10 Angstroms. The assumption this method makes is that there is complete randomness in the mixing of illitic and smectitic components and that the peak position varies in a straight-line manner from 14 to 10 Angstroms. Table 4 presents the results of this analysis. The values are rounded to the nearest 10 percent because of the difficulty in picking the peak position.

Table 4. Estimated amounts of illitic layers in mixed-layer illite/smectite.

The average amount of illitic layers in the material is 60 percent. This should probably be considered to be a minimum estimation, because the (002)/(003) peak of the glycol solvated samples suggest a value that would probably be more on the order of 80 percent illitic layers.

Crystallite size.

The Jade computer program contains an algorithm for determining crystallite size of crystallites smaller than 1 micron. This information is produces as a by-product of measuring the area under a peak and became available by determining the relative amounts of the various clay minerals. In general, kaolinite has the largest grains with an average size of 297 A, followed by illite at 152 A and the mixed-layer material at 67 A (Table 5).

Table 5. Crvstallite size of clay minerals in the CJ sample set.

References

- Eckel, E. B. (1997) Minerals of Colorado: Friends of Mineralogy Edition, Fulcrum Publishers.
- Moore, D. M., and R. C. Reynolds (1997) X-ray diffraction and the identification and analysis of clay minerals, 2nd ed.: Oxford University Press, New York, 378 p.

Nadeau, P. H. and R. C. Reynolds (1981) Burial and contact metamorphism in the Mancos Shale: Clays and Clay Minerals, 29, 249-259.

Appendix 1.

Method used to separate and prepare clay minerals for x-ray diffraction

Appendix 1. Method used to separate and prepare clay minerals for x-ray diffraction

I. Place bulk sample in water overnight to remove any easily soluble material and to partially disaggregate clay. Remove as much of the water as practical without losing the sample.

2. Add fresh water and ¼ tsp. of sodium metaphosphate (Calgon). Agitate vigorously to get the material into suspension.

3. Allow to sit undisturbed for 8 hrs. With a turkey baster or some other large pipette, remove the upper **10** cm. of the suspension. Place this in a separate container. This suspension contains clay (less than 2 micron particles).

4. Re-suspend the bulk material to collect additional clay. If the original suspension yielded little clay or if it flocculated, add additional Calgon. Otherwise, do not add Calgon.

5. Add ¼ tsp. of magnesium chloride or magnesium sulfate to the clay suspension. This does two things; it flocculates the clay and begins the magnesium saturation process. When the flocculated particles have settled to the bottom of the container, remove as much water as you can with the pipette.

6. After re-suspending the bulk sample. repeat step 3. Add the clay suspension to the container with the flocculated clay. Repeat step 5.

7. Repeat steps 4 and 5, but do not add more Calgon to the bulk sample. **If** there is insufficient clay to do the required analysis, steps $4 \& 5$ can be repeated as necessary.

8. When sufficient clay has been collected and step 5 has been completed on the accumulated clay, the clay should by now be Mg saturated. At this point, centrifuge the sample to concentrate the clay and remove as much of the water as possible.

9. Add distilled water and get the flocculated clay into suspension. Centrifuge and remove the water.

10. Repeat step 9 until you are confident that you have removed the excess Mg. This usually takes 3 washings, but it depends on the size of the centrifuge tube. When the excess Mg salt is gone, it will take much longer to clear the sample because the finest fraction of clay will start staying in suspension.

1I. Add enough water to make a rather thick suspension of clay and agitate vigorously. Allow to sit for a couple of hours so that clumps of clay will settle out. With a pipette, collect the clay from the top centimeter (to obtain only -2 micron material and not the larger floes), put it onto a glass slide and allow to dry.

If the sample contains a lot of smectite or if it too thick, it may curl up off of the slide. There are two things that can be done if this happens. Make another slide using less material or frost the glass slide lightly and remake the sample.

Appendix 2.

Estimating the abundance of non-phyllosilicate minerals in Mancos Shale

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Appendix 2. Estimating the abundance of non-phyllosilicate minerals in Mancos Shale

The amounts of non-phyllosilicate minerals in the Mancos Shale samples were estimated using the following technique. First, diffraction patterns of pure samples of quartz, calcite, dolomite and albite were obtained under the same conditions as the bulk samples of Mancos Shale. Because the intensities of the most intense peak of the minerals were not identical, the next step was to obtain a normalization factor so that each of the other minerals could be compared to quartz. This was done by first removing the background from the patterns and determining the peak intensity, then dividing the peak intensity of each of the minerals into the peak intensity ofquartz. A pure sample of orthoclase was not available, so the intensity of the main peak of this mineral was assumed to be the same as that of albite. This most likely gives an incorrect normalization factor, but because this mineral is present in amounts just slightly above the detection limit, it makes little difference in the interpretation.

The second step is to normalize the peak intensities to quartz and sum the various peak intensities. This gives an approximation of what the quartz peak intensity would be if only quartz plus phyllosilicate minerals were present in the samples. The ratio of intensities of the normalized peaks to the pure quartz standard is a measure of the amount of non-phyllosilicate minerals present in the sample. This assumes that the presence of phyllosilicate minerals suppresses the quartz peak in an amount directly proportional to their abundance. The total amount of non-phyllosilicate minerals is esfimated from 33 to 52 percent in the 10 Mancos Shale samples.

[he final step is to compare the intensities of the various normalized peaks to the total intensity values and multiply that by the percent of non-phyllosilicate minerals obtained in the preceding step. This gives an approximation of the abundance of each mineral in the bulk sample.

The mineral gypsum was not included in the calculation just described because it was detected in only.4 samples. By comparison with the values for the other minerals, the amount of gypsum is estimated to be 1 percent or less.

The data and results of calculations are shown on the next page. Diffraction patterns for quartz, calcite, dolomite, albite and gypsum follow the calculations page.

Method for calculating abundance of non-phyllosilicate minerals in bulk samples

Normalize calcite, dolomite and feldspar standards to quartz (background removed)

Measured peak heights (in counts/second) after background removed

Peak heights (in counts/second) normalized to quartz

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Estimated Percentages of Non-phyllosilicate minerals

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<c \hood> Thursday, Jan 26, 2006 @10.26a

<GYPSUM.RAW> Gypsum standard ____ _____ ______ ____

 $. d = 761$ 2000- **1500.** I; **d%3 ⁰⁶** Intensity(Counts) 1000 $d = 428$ 1000 . de 2000 . de 3000 . de $. d = 379$ ÷. **4=³⁰⁰** $d = 280$ 500 $d = 184$ **d=8 03 d=3** 47f **d--288 =22** $d = 1.90$ $d = 288$ $d = 603$ $d = 347$ $d = 2.22$ $d=2.14$ $d=2.08$ $d=1.99$ $\left| \begin{array}{ccc} | & | & | \end{array} \right|$ $\left| \begin{array}{ccc} | & | & | & | \end{array} \right|$ $\left| \begin{array}{ccc} d=2.49 & d=2.49 \\ d=2.59 & d=2.45 \\ d=2.45 & 40 \\ 0 & 0 & 0 \end{array} \right|$ **Welch Brick West Art Water Real WAMAMA** *Rith Walls* la u" $\mathbf{0}$ **10** 20 **30** 40 **50**

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Appendix 3.

X-ray diffraction peaks used to identify minerals present in bulk samples

Appendix 3. X-ray diffraction peaks used to identify minerals present in bulk samples.

Notes: Not all of the peaks listed in the ASTM data file of x-ray diffraction peaks were observed for these minerals. Except for quartz, most were present in too small amounts for any peaks except for those of high intensity to appear

Not all peaks appear at their ideal position. In mixtures of several minerals such as these samples, interference of peaks from different minerals can cause the resulting peak to shift. Also, it is difficult to determine exactly the position of the top of a peak when the peaks are of low intensity such as most of the peaks that appear on these samples.

Appendix 4.

X-ray diffraction patterns of bulk samples

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Appendix 4. X-ray diffraction patterns of bulk samples.

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Two copies of each diffraction pattern are presented. The first of each pair is annotated with the d-spacings of most diffraction peaks. A few were not annotated because the closeness of other peaks would make the annotations unreadable.

The second of each pair is annotated with letters showing the identities of the most intense peaks.. Not all of the minor peaks were annotated, although they were accounted for.

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Appendix 5.

Stacked x-ray diffraction patterns of clay minerals in CJ sample set

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CJ202HH RAW> 0J202-40 H550 <CJ202G RAW> CJ202-40 giveol
<CJ202C RAW> CJ 202-40 Clay

CJ206HH RAW> CJ206-40 H550 **J206HL RAW - CJ206-40 H300** <CJ206G RAW> CJ206-40 glycol
<CJ206C.RAW> CJ 206-40 Clay

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<CJ207HH RAW> CJ207-40 H550 CAZUTHE RAW - 0.207-40 H300 <CJ207G RAW> CJ207-40 glycol
<CJ207C RAW> CJ 207-40 Clay

CJ2010HH RAW> CJ2010-40 H550 J2010HLRAW - CJ2010-401+300 <CJ2010G.RAW> CJ2010-40 glycol
<CJ2010C.RAW> CJ 2010-40 Clay

Appendix 6.

X-ray diffraction patterns of clay minerals in **CJ** sample set

Patterns are arranged in the sequence: untreated, glycol solvated, heated to **3300"** C and heated to **500'** C for each sample.

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Appendix 7.

X-ray diffraction patterns of glycol solvated clays with background removed

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Appendix E

Copies of Analytical Reports of BET Surface Area Analysis

Prepared by Micromeritics Analytical Services Norcross, Georgia

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TriStar 3000 V6.05.01 A Unit 1 Port 1 Correspondent Serial #: 1098 Page 1

Sample: CJ-201-40 06-0210 3/3 Operator: MJP Submitter: S.M.Stoller File: C:\..\01JAN\06-0210.SMP

Started: 2/1/2006 12:27:55PM Completed: 2/1/2006 3:23:59PM Report Time: 2/1/2006 3:24:00PM Warm Free Space: 5.1716 cm³ Measured Equilibration Interval: 10 s Sample Density. 1.000 **g/cm ³**

Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 4.2757 g Cold Free Space: 14.2717 cm³ Measured Low Pressure Dose: None Automatic Degas: Yes

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TriStar 3000 V6.05.01 A Contract Unit 1 Port 1 Contract 1098 Contract 2008 Page 3

Sample: CJ-201-40 06-0210 313 Operator: MJP Submitter: S.M.Stoller File: C:\...\01JAN\06-0210.SMP

Started: 2/1/2006 12:27:55PM Completed: 2/1/2006 3:23:59PM Report Time: 2/1/2006 3:24:00PM Warm Free Space: 5.1716 cm³ Measured Equilibration Interval: **10** s Sample Density: 1.000 g/cm³

Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 4.2757 g Cold Free Space: 14.2717 cm3 Measured Low Pressure Dose: None Automatic Degas: Yes

Sample Prep: Stage **I**

Soak Temperature (°C) 100

Ramp Rate (°C/min) 10

Soak Time (min) 960

BET Surface Area Report

BET Surface Area: 10.6561 **±** 0.0383 m2/g Slope: 0.405995 **±** 0.001443 g/cm 3 STP Y-Intercept 0.002523 **±** 0.000277 g/cm 3 STP C: 161.941253 Qm: 2.4479 cm³/g STP Correlation Coefficient 0.9999432 Molecular Cross-Sectional Area: 0.1620 nm²

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TriStar 3000 V6.05.01 A **Unit 1 Port 1** Serial #: 1098 Page 5

Sample: CJ-201-40 06-0210 **3/3** Operator: MJP Submitter: S.M.Stoller File: C:\..\01JAN\06-0210.SMP

Started: 2/1/2006 12:27:55PM Completed: 2/1/2006 3:23:59PM Report Time: 2/1/2006 3:24:00PM Warm Free Space: 5.1716 cm³ Measured Equilibration Interval: 10 s Sample Density. 1.000 g/cm³

Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 4.2757 g Cold Free Space: 14.2717 **cm³**Measured Low Pressure Dose: None Automatic Degas: Yes

Sample Prep: Stage **1**

Soak Temperature (°C) 100

Ramp Rate (°C/min) 10

Soak Time (min) 960

Summary Report

Surface Area Single point surface area at $P/Po = 0.301305895$: 10.4413 m²/g

BET Surface Area: 10.6561 m²/g

TriStar 3000 V6.05.01 A Unit **1** Port 2 Serial #: 1098 Page 1

Sample: CJ-202-40 06-0211 38/38 Operator: MJP Submitter: S.M.Stoller File: C:\...\01JANW06-0211 .SMP

Started: **2/11/2006** 12:27:55PM Completed: 2/1/2006 3:23:59PM Report Time: 2/11/2006 3:42:29PM Warm Free Space: 6.0917 cm³ Measured Equilibration Interval: 10 s Sample Density: 1.000 g/cm³

Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 4.8504 g Cold Free Space: 17.8637 cm³ Measured Low Pressure Dose: None Automatic Degas: Yes

Sample Prep: Stage **I**

Soak Temperature (°C) 100

Ramp Rate (°C/min) 10

Soak Time (min) 960

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TriStar 3000 V6.05.01 A Unit I Port2 Serial #: 1098 Page 3

Sample: CJ-202-40 06-0211 38/38 Operator: MJP Submitter: S.M.Stoller File: C:\..\01JAN\06-0211.SMP

Started: 2/1/2006 12:27:55PM Completed: 2/1/2006 3:23:59PM Report Time: 2/1/2006 3:42:29PM Warm Free Space: 6.0917 cm³ Measured **Equilibration Interval: 10** Sample Density: 1.000 g/cm³

Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 4.8504 g Cold Free Space: 17.8637 **cm³**Measured Low Pressure Dose: None Automatic Degas: Yes

Sample Prep: Stage **1**

Soak Temperature (°C) 100

Ramp Rate (°C/min) 10

Soak Time (min) 960

BET Surface Area Report

BET Surface Area: 12.2117 ± 0.0440 m²/g Slope: 0.354268 **±** 0.001265 g/cm 3 STP Y-Intercept 0.002211 ± 0.000225 g/cm³ STP C: 161.265813 Qm: 2.8052 cm³/g STP Correlation Coefficient 0.9999490 Molecular Cross-Sectional Area: 0.1620 nm²

TriStar 3000 V6.05.01 A Unit I Port 2 Serial #: 1098 Page 5

Sample: CJ-202-40 06-0211 38/38 Operator: MJP Submitter: S.M.Stoller File: C:\..\01JAN\06-0211.SMP

Started: 2/1/2006 12:27:55PM Completed: 2/1/2006 3:23:59PM Report Time: 2/1/2006 3:42:29PM Warm Free Space: 6.0917 cm3 Measured Equilibration Interval: 10s Sample Density. 1.000 **g/cm ³**

Analysis Adsorptive: .N2 Analysis Bath Temp.: 77.300 K Sample Mass: 4.8504 g Cold Free Space: 17.8637 cm³ Measured Low Pressure Dose: None Automatic Degas: Yes

Sample Prep: Stage 1

Soak Temperature (°C)

Ramp Rate (°C/min) 10

Soak Time (min) 960

Summary Report

Surface Area Single point surface area at P/Po = 0.276325488: 11.9531 m²/g

100

BET Surface Area: 12.2117 m²/g

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TriStar 3000 V6.05.01 A Unit **I** Port **3** Serial #: 1098 Page 1

Sample: CJ-203-40 06-0212 815/815 Operator: MJP Submitter: S.M.Stoller File: C:\...\01JAN\06-0212.SMP

Started: 2/1/2006 12:27:55PM Completed: 2/1/2006 3:23:59PM Report Time: .2/1/2006 3:24:02PM Warm Free Space: 5.8383 cm³ Measured Equilibration Interval: 10 s Sample Density: 1.000 g/cm³

Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 4.9651 g Cold Free Space: 16.8406 **cm³**Measured Low Pressure Dose: None Automatic Degas: Yes

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TriStar 3000 **V6.05.01** A Unit **I** Port **3** Serial **#: 1098** Page **3**

Sample: CJ-203-40 06-0212 815/815 Operator: MJP Submitter: S.M.Stoller File: C:\...\01JAN\06-0212.SMP

Started: 2/1/2006 12:27:55PM Completed: 2/1/2006 3:23:59PM Report Time: 2/1/2006 3:24:02PM Warm Free Space: 5.8383 cm³ Measured Equilibration Interval: **los** Sample Density 1.000 **g/cm ³**

Analysis Adsorptive: **N;** Analysis Bath Temp.: 77 .300 K Sample Mass: 4.9651 g Cold Free Space: 16 .8406 **cm³**Measured Low Pressure Dose: None Automatic Degas: Ye

Sample Prep: Stage 1

Soak Temperature (°C) 100

Ramp Rate (°C/min) 10

Soak Time (min) 960.

BET Surface Area Report

BET Surface Area: 8.8087 ± 0.0164 m²/g **.** Slope: 0A89798 **±** 0.000906 **g/cm ³**STP Y-Intercept 0.004396 **±** 0.000173 g/cm³STP C: 112.409201 Qm: 2.0235 cm3/g STP Correlation Coefficient 0.9999846 Molecular Cross-Sectional Area: 0.1620 nm2

TriStar 3000 V6.05.01 A Unit 1 Port 3 Serial #: 1098 Page 5

Sample: CJ-203-40 06-0212 815/815 Operator: MJP Submitter: S.M.Stoller File: C:\..\01JAN\06-0212.SMP

Started: 2/1/2006 12:27:55PM Completed: 2/1/2006 3:23:59PM Report Time: 2/1/2006 3:24:02PM Warm Free Space: 5.8383 cm³ Measured Equilibration Interval: 10 Sample Density. **1.000** g/cm ³

Analysis Adsorptive: **N2** Analysis Bath Temp.: 77 .300 K Sample Mass: 4.9651 g Cold Free Space: 16.8406 cm³ Measured Low Pressure Dose: None Automatic Degas: Yes

Sample Prep: Stage **1**

Ramp Rate (°C/min) 10

Soak Time (min) 960

Summary Report

Surface Area Single point surface area at P/Po = 0.299973650: 8.6142 m²/g

Soak Temperature (°C) 100

BET Surface Area: 8.8087 m²/g

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TriStar 3000 V6.05.01 A Unit 2 Port **I** Serial **#:** 1449 Page 1

Sample: CJ-204-40 06-0213 **GI/G1** Operator: MJP Submitter: S.M.Stoller File: C.\..\01JAN\06-0213.SMP

Started: 2/1/2006 12:24:05PI Completed: 2/1/2006 3:16:44PM Report Time: 2/1/2006 3:43:34PM Warm Free Space: 5.6492 **cm³**Measured **Equilibration Interval: 10 s** Sample Density: **1.000 g/cm ³**

Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 4.4734 g Cold Free Space: 16.0508 **cm³**Measured Low Pressure Dose: None Automatic Degas: Yes

Sample Prep: Stage 1

Soak Temperature (°C) 100

Ramp Rate (°C/min) 10

Soak Time (min) 960

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TriStar 3000 V6.05.01 A Unit 2 Port 1 Serial #: 1449 Page 3

Sample: CJ-204-40 06-0213 **GI/G1** Operator: MJP Submitter: S.M.Stoller File: C:\..\01JAN\06-0213.SMP

Started: 2/1/2006 12:24:05PM Completed: 2/1/2006 3:16:44PM Report Time: 2/1/2006 3:43:34PM Warm Free Space: 5.6492 cm³ Measured Equilibration Interval: 10 s Sample Density: 1.000 g/cm³

Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 4.4734 g Cold Free Space: 16.0508 cm³ Measured Low Pressure Dose: None Automatic Degas: Yes

Sample Prep: Stage **I**

Soak Temperature (°C) 100

Ramp Rate (°C/min) 10

-Soak Time (min) 960

BET Surface Area Report

BET Surface Area: 12.9162 ± 0.0501 m²/g Slope: 0.335315 **±** 0.001288 g/cm 3 STP Y-Intercept 0.001719 ± 0.000229 g/cm3 STP C: 196.026540 Qm: 2.9671 cm³/g STP Correlation Coefficient 0.9999410 Molecular Cross-Sectional Area: 0.1620 nm'

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TriStar 3000 **V6.05.01** A Unit 2 Port **1** Serial #: 1449 Page 5 Sample: CJ-204-40 06-0213 **GI/G1** Operator: MJP Submitter: S.M.Stoller File: C:\..\01JAN\06-0213.SMP Started: **2/1/2006** 12:24:05PM Completed: 2/1/2006 3:16:44PM Report Time: 2/1/2006 3:43:34PM Warm Free Space: 5.6492 **cm³**Measured Equilibration Interval: **los** noration interval: 15 5
Sample Density: 1.000 g/cm³ Sample Prep: Stage **.** 1 Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 4.4734 g Cold Free Space: 16.0508 cm³Measured Low Pressure Dose: None Automatic Degas: Yes Soak Temperature (°C) 100 Ramp Rate (°C/min) 10 Soak Time (min) $.960$ Summary Report

Surface Area Single point surface area at P/Po **=** 0.276424207: 12.6720 m²/g

BET Surface Area: 12.9162 m²/g

TriStar 3000 V6.05.01 A Unit 2 Port 2 Correspondent Serial #: 1449 Page 1

Sample: CJ-205-40 06-0214 G2/G2 Operator: MJP Submitter: S.M.Stoller File: C:\..\01JAN\06-0214.SMP

Started: 2/1/2006 12:24:05PM Completed: 211/2006 3:16:44PM Report Time: 211/2006 3:16:49PM Warm Free Space: 5.6027 **cm³**Measured Equilibration Interval: 10s Sample Density: 1.000 g/cm³

Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 4.2657 g Cold Free Space: 15.9013 cm3 Measured Low Pressure Dose: None Automatic Degas: Yes

Sample Prep: Stage **1**

Soak Temperature (°C) 100

Ramp Rate (°C/min) 10

Soak Time (min) 960

TriStar 3000 V6.05.01 A UUnit 2 Port 2 Serial **#:** 1449 Page **3**

Sample: CJ-205-40 06-0214 G2/G2 Operator: MJP Submitter: S.M.Stoller File: **C:** ...\01JAN\06-0214.SMP

Started: **2/11/2006** 12:24:05PM Completed: 2/11/2006 3:16:44PM Report Time: 2/11/2006 3:16:49PM Warm Free Space: 5.6027 cm³ Measured Equilibration Interval: **los** Sample Density. 1.000 g/cm ³

Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 4.2657 g Cold Free Space: 15.9013 **cm³**Measured Low Pressure Dose: None Automatic Degas: Yes

Sample Prep: Stage 1

Soak Temperature (*C) 100

Ramp Rate (°C/min) 10

Soak Time (min) 960

BET Surface Area Report

BET Surface Area: 9.9511 **±** 0.0246 m² /g Slope: 0A34049 **±** 0.001061 g/cm 3 STP Y-Intercept 0.003410 **±** 0.000203 g/cm 3 STP **C:** 128.300420 Qm: 2.2859 cm3/g STP Correlation Coefficient 0.9999731 Molecular Cross-Sectional Area: 0.1620 nm ²

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TriStar 3000 \6.05.01 A Unit 2 Port 2 Serial #: 1449 Page 5

Sample: CJ-205-40 06-0214 G21G2 Operator: MJP Submitter: S.M.Stoller File: **C:** ...\01JAN'06-0214.SMP

Started: 2/1/2006 12:24:05PM Completed: 2/1/2006 3:16:44PM Report Time: 211/2006 3:16:49PM Warm Free Space: 5.6027 cm³ Measured Equilibration Interval: 10s Sample Density: 1.000 g/cm³

Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 4.2657 g Cold Free Space: 15.9013 cm³ Measured Low Pressure Dose: None Automatic Degas: Yes

Sample Prep: Stage 1

Soak Temperature (°C) 100

Ramp Rate (°C/min) 10

Soak Time (min) 960

Summary Report

Surface Area Single point surface area at P/Po = $0.300337231: 9.7341 m^2/g$

BET Surface Area: 9.9511 m²/g

TriStar 3000 V6.05.01 A Unit 2 Port **3** Serial #: 1449 Page **1** Sample: CJ-205-40-P 06-0215 G3/G3 Operator: MJP Submitter: S.M.Stoller File: **C.\ ..\1** JAND6-0215.SMP Started: 2/1/2006 12:24:05PM Completed: 2/1/2006 3:16:44PM Report Time: 2/1/2006 3:16:50PM Warm Free Space: 5.8945 cm³ Measured Equilibration Interval: 10 sample Density: 1.000 g/cm³ .Sample Prep: Stage ¹ Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 4.4765 g Cold Free Space: 17.0075 **cm³**Measured Low Pressure Dose: None Automatic Degas: Yes .Soak Temperature (°C) **100** Ramp Rate (°C/min) 10 Soak Time (min) 960 Isotherm Tabular Report

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TriStar 3000 V6.05.01 A Unit 2 Port 3 Serial #: 1449 Page 3

Sample: CJ-205-40-P 06-0215 G3/G3 Operator: MJP Submitter: S.M.Stoller File: C:\..\01JAN\06-0215.SMP

Started: 2/1/2006 12:24:05PM Completed: 2/1/2006 3:16:44PM Report Time: 2/1/2006 3:16:50PM Warm Free Space: 5.8945 cm³ Measure Equilibration Interval: 10 Sample Density-1.000 **g/cm ³**

Analysis Adsorptive: **N2** Analysis Bath Temp.: 77.300 K Sample Mass: 4.d 4765 **g** Cold Free Space: 17,0075 cm³ Measure Low Pressure Dose: Non Automatic Degas: Ye

Sample Prep: Stage **I**

Soak Temperature (°C) 100

Ramp Rate (°C/min) 10

Soak Time (min) **960**

BET Surface Area Report

BET Surface Area: 10.6955 **±** 0.0222 m2/g Slope: 0.403460 **±** 0.000830 **g/cm ³**STP Y-Intercept 0.003553 ± 0.000159 g/cm 3 STP C: 114.541452 Qm: 2.4569 cm3/g STP Correlation Coefficient 0.9999810 Molecular Cross-Sectional Area: 0.1620 nm ² Relative

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BET Surface Area: 10.6955 m²/g

TriStar 3000 V6.05.01 A **A CONTEX 1 Unit 1 Port 1** CONTEX 1098 **Page 1** Page 1

Sample: CJ-206-40 06-0216 F3/F3 Operator: MJP Submitter: S.M.Stoller File: C:\...\O1JAN\06-0216.SMP

Started: 2/1/2006 4:22:41PM Completed: 2/1/2006 8:12:53PM Report Time: 2/1/2006 8:12:54PM Warm Free Space: 5.3716 cm³ Measured Equilibration Interval: 10s Sample Density. 1.000 g/cm ³

Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 4.8010 g Cold Free Space: 15.0468 cm³ Measured Low Pressure Dose: None Automatic Degas: Yes

Soak Temperature (°C) 100

Ramp Rate (°C/min) 10

Soak Time (min) 960

TriStar 3000 V6.05.01 A **Unit 1 Port 1** Serial #: 1098 **Page 3**

Sample: CJ-206-40 06-0216 F3/F3 Operator: MJP Submitter: S.M.Stoller File: **C:** ...\01JAN\06-0216.SMP

Started: 2/1/2006 4:22:41PM Completed: 2/1/2006 8:12:53PM Report Time: 2/1/2006 8:12:54PM Warm Free Space: 5.3716 cm³ Measured Equilibration Interval: 10s Sample Density: 1.000 g/cm³

Analysis Adsorptive: **N;** Analysis Bath Temp.: 77.300 K Sample Mass: 4.8010 Cold Free Space: 15.0468 cm³ Measured Low Pressure Dose: None Automatic Degas: Yes

Sample Prep: Stage Soak Temperature (°C) 1 100

Ramp Rate (°C/min) 10

Soak Time (min) 960

BET Surface Area Report

BET Surface Area: 9.4615 ± 0.0195 m²/g Slope: 0.456311 ± 0.000930 g/cm³ STP Y-Intercept 0.003784 **±** 0.000178 g/cm 3 STP C: 121.597783 Qm: 2.1735 cm³/g STP Correlation Coefficient 0.9999813 Molecular Cross-Sectional Area: 0.1620 nm ²

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TriStar 3000 V6.05.01 A Unit **1** Port **1** Serial #: 1098 Page 5 Sample: CJ-206-40 06-0216 F3/F3 Operator: MJP Submitter: S.M.Stoller File: C:\..\01JAN\06-0216.SMP Started: 2/112006 4:22:41PM Completed: 2/112006 8:12:53PM Report Time: 211/2006 8:12:54PM Warm Free Space: 5.3716 **cm ³**Measured Equilibration Interval: 10s Sample Density. 1.000 g/cm ³ Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 4.8010 g Cold Free Space: 15.0468 **cm³**Measured Low Pressure Dose: None Automatic Degas: Yes Sample Prep: Stage 1 Soak Temperature (°C) 100 Ramp Rate (°C/min) 10 Soak Time (min) 960 Summary Report Surface Area

Single point surface area at P/Po = 0.300121926: 9.2542 m²/g

BET Surface Area: 9.4615 m²/g

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TriStar 3000 V6.05.01 A Unit **1** Port 2 Serial #: 1098 Page 1

Sample: CJ-207-40 06-0217 **N1/N1** Operator: MJP Submitter: S.M.Stoller File: C:\..\01JAN\06-0217.SMP Started: 2/1/2006 4:22:41PM Completed: 2/11/2006 8:12:53PM Report Time: 2/2/2006 7:50:36AM Warm Free Space: 5.5902 cm³ Measured Equilibration Interval: 10s Sample Density: 1.000 g/cm³ Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 4.3543 g Cold Free Space: 15.9894 cm3 Measured Low Pressure Dose: None Automatic Degas: Yes Sample Prep: Stage 1 Soak Temperature (°C) 100 Ramp Rate (°C/min) 10 **1** Soak Time (min) 960

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ERVICES TriStar 3000 V6.05.01 A Unit **1** Port 2 Serial #: 1098 Page 3 Sample: CJ-207-40 06-0217 **N1/N1** Operator: MJP Submitter: S.M.Stoller File: C:\...\01JAN\06-0217.SMP Started: 2/1/2006 4:22:41 PM Completed: 2/1/2006 8:12:53PM Report Time: 2/2/2006 7:50:36AM Warm Free Space: 5.5902 cm³ Measured Equilibration Interval: **los** nbratton interval: 1555
Sample Density: 1.000 g/cm³ Sample Prep: Stage Soak Temperature **(°C)** 100 Analysis Adsorptive: N2 Analysis Bath Temp.: 77 .300 K Sample Mass: 4. 3543 g Cold Free Space: 15 .9894 **cm³**Measured Low Pressure Dose: None **Automatic Degas: Yes** Ramp Rate (°C/min) 10 Soak Time (min) 960 BET Surface Area Report BET Surface Area: 13.1245 ± 0.0495 m²/g Slope: 0.330055 **i** 0.001233 **g/cm ³**STP Y-Intercept 0.001631 **±** 0.000219 **g/cm ³**STP C: 203.384811 Qm: 3.0149 cm3/g STP Correlation Coefficient 0.9999442 Molecular Cross-Sectional Area: 0.1620 nm². Relative Pressure (PIPo) 0.051121487 0.075556860 0.098313891 0.123338473 Quantity $1/[Q(Po/P - 1)]$ Adsorbed (cm³ /g STP) 2.8779 3.0562 3.1924 3.3267 0.018721 0.026743 .0.034154 0.042292

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0.148133515 0.173619345 0.199105163 0.224676473 0.250356073 0.276132543 **3A562** 3.5836 3.7082 3.8333 3.9600 4.0890 0.050313 .0.058627 0.067041 0.075597 0.084335 0.093290

TriStar 3000 V6.05.01 A Control Unit 1 Port 2 Control Serial #: 1098 Page 5

Sample: CJ-207-40 06-0217 **Ni/Nl** Operator: MJP Submitter: S.M.Stoller File: C:\..\01JAN\06-0217.SMP

Started: 211/2006 4:22:41PM Completed: 2/1/2006 8:12:53PM Report Time: 2/2/2006 7:50:36AM Warm Free Space: 5.5902 cm3 Measured Equilibration Interval: 10s Sample Density. 1.000 **g/cm ³**

Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 4.3543 g Cold Free Space: 15.9894 cm³ Measured Low Pressure Dose: None Automatic Degas: Yes

Sample Prep: Stage

Soak Temperature (°C) 100

Ramp Rate ("C/min) 10

Soak Time (min) 960

Summary Report

Surface Area Single point surface area at P/Po = 0.276132543: 12.8851 m²/g

BET Surface Area: 13.1245 m²/g

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TriStar 3000 V6.05.01 A Unit **1** Port **3** Serial #: 1098 Page **1**

Sample: CJ-208-40 06-0218 N2/N2 Operator: MJP Submitter: S.M.Stoller File: C:\..\01JAN\06-0218.SMP

Started: 2/1/2006 4:22:41PM Completed: 2/1/2006 8:12:53PM Report Time: 2/1/2006 8:12:55PM Warm Free Space: 5.7202 cm³ Measured Equilibration Interval: 10s Sample Density:. 1.000 **g/cm ³**

Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 4.3662 g Cold Free Space: 16.4086 **cm³**Measured Low Pressure Dose: None Automatic Degas: Yes

Sample Prep: Stage **I**

Soak Temperature (°C) 100

Ramp Rate (°C/min) 10

Soak Time (min) 960

Isotherm Tabular Report

TriStar 3000 V6.05.01 A Unit 1 Port 3 Serial #: 1098 Page 3

Sample: CJ-208-40 06-0218 N2/N2 Operator: MJP Submitter: S.M.Stoller File: C:\..\01JAN\06-0218.SMP

Started: 2/1/2006 4:22:41PM Completed: 2/1/2006 8:12:53PM Report Time: 2/1/2006 8 12:55PM Warm Free Space: 5.7202 cm³ Measured Equilibration Interval: 10 s. Sample Density. 1.000 g/cm ³

Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 4.3662 g Cold Free Space: 16.4086 **cm³**Measured Low Pressure Dose: None Automatic Degas: Yes

Sample Prep: Stage **1**

Soak Temperature (°C) Ramp Rate (°C/min)

100 10

10

Soak Time (min) 960

BET Surface Area Report

BET Surface Area: 9.3897 + 0.0250 *m*² Ig Slope: 0.460156:t 0.001212 **9/cm ³**STP Y-Intercept 0.003458 ± 0.000232 g/cm³ STP C: 134.056409 Qm: 2.1570 cm³/g STP Correlation Coefficient 0.9999688 Molecular Cross-Sectional Area: 0.1620 nm ²

TriStar 3000 V6.05.01 A Unit 1 Port 3 Serial #: 1098 Page 5 Sample: CJ-208-40 06-0218 N2/N2 Operator: MJP Submitter: S.M.Stoller File: C:\..\01JAN\06-0218.SMP Started: 2/1/2006 4:22:41PM Analysis Adsorptive: N2 Completed: 2/1/2006 8:12:53PM Analysis Bath Temp.: 77.300 K Report Time: 2/1/2006 8:12:55PM Sample Mass: 4.3662 g Warm Free Space: 5.7202 **cm³**Measured Cold Free Space: 16.4086 cm³ Measured Equilibration Interval: 10 s Low Pressure Dose: None Sample Density. 1.000 **g/cm ³** Automatic Degas: Yes $\ddot{}$ Soak Temperature (°C) Sample Prep: Stage Ramp Rate (°C/min) Soak Time (min) 1* 100 10 960

Summary Report

Surface Area Single point surface area at P/Po = $0.300589849: 9.1880$ m²/g

BET Surface Area: 9.3897 m²/

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TriStar 3000 V6.05.01 A Unit 2 Port 1 Serial #: 1449 Page 1

Sample: CJ-209-40 06-0219 N5/N5 Operator: MJP Submitter: S.M.Stoller File: C:\..\01JAN\06-0219.SMP

Started: 2/11/2006 4:23:29PM Completed: 2/1/2006 7:40:08PM Report Time: 2/1/2006 7:40:06PM Warm Free Space: 5.4357 cm³ Measured Equilibration Interval: 10s Sample Density. 1.000 **g/cm ³**

Analysis Adsorptive: N; Analysis Bath Temp.: 77.300 K Sample Mass: 5.1155 Cold Free Space: 15 .2828 **cm ³**Measured Low Pressure Dose: Nor Automatic Degas: Ye

Sample Prep: Stage Soak Temperature (°C) 1 10

Ramp Rate (°C/min) 10

Soak Time (min) 960

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TriStar 3000 V6.05.01 A Unit 2 Port 1 Serial #: 1449 Page 3

Sample: CJ-209-40 06-0219 **N5/N5** Operator: MJP Submitter: S.M.Stoller File: C:\..\01JAN\06-0219.SMP

Started: 2/1/2006 4:23:29PM Completed: 2/1/2006 7:40:08PM Report Time: 2/1/2006 7:40:06PM Warm Free Space: 5.4357 cm³ Measured Equilibration Interval: 10 s Sample Density. 1.000 g/cm³

Analysis Adsorptive: **N2** Analysis Bath Temp.: 77 .300 K Sample Mass: **5.1** 1155g Cold Free Space: 15.2828 cm³ Measured Low Pressure Dose: None Automatic Degas: Yes

Sample Prep: Stage Soak Temperature (°C)

1 100

Ramp Rate (°C/min) 10

Soak Time (min) 960

BET Surface Area Report

BET Surface Area: 10A560 **±** 0.0226 m2/g Slope: 0.412643 **±** 0.000883 g/cm3 STP Y-Intercept 0.003694 **±** 0.000169 g/cm³STP C: 112.707884 Qm: 2.4019 cm³/g STP Correlation Coefficient 0.9999794 Molecular Cross-Sectional Area: 0.1620 nm ²

TriStar 3000 V6.05.01 A Unit 2 Port 1 Serial #: 1449 Page 5 Sample: CJ-209-40 06-0219 N5/N5 Operator: MJP Submitter: S.M.Stoller File: C:\..\01JAN\06-0219.SMP Started: 2/1/2006 4:23:29PM Completed: 2/1/2006 7:40:08PM Report Time: 2/1/2006 7:40:06PM Warm Free Space: 5.4357 **cm³**Measured Equilibration Interval: 10s Sample Density: 1.000 g/cm³ Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 5.1155 g Cold Free Space: 15.2828 cm³ Measured Low Pressure Dose: None Automatic Degas: Yes Sample Prep: Stage **1** Soak Temperature (°C) 100 Ramp Rate (°C/min) 10 Soak Time (min) 960 Summary Report

Surface Area

Single point surface area at P/Po = 0.300100534: 10.2099 m²/g

BET Surface Area: 10.4560 m²/g

TriStar 3000 V6.05.01 A Unit 2 Port 2 Serial #: 1449 Page **1**

Sample: CJ-210-40 06-0220 N6/N6 Operator: MJP Submitter: S.M.Stoller File: C:\..\01JAN\06-0220.SMP

100

Started: **2/11/2006** 4:23:29PM Completed: 2/11/2006 7:40:08PM Report Time: 2/2/2006 7:51:21AM Warm Free Space: 5.6495 **cm³**Measured Equilibration Interval: 10s Sample Density: 1.000 g/cm³

Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 5.1428 g Cold Free Space: 16.1759 **cm ³**Measured Low Pressure Dose: None Automatic Degas: Yes

10

Soak Time (min) 960

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TriStar 3000 V6.05.01 A Unit 2 Port 2 Serial #: 1449 Page 3

Sample: CJ-210-40 06-0220 N6/N6 Operator: MJP Submitter: S.M.Stoller File: C:\..\01JAN\06-0220.SMP

Started: 2/1/2006 4:23:29PM Completed: 2/1/2006 7:40:08PM Report.Time: 2/2/2006 7:51:21AM Warm Free Space: 5.6495 **cm³**Measured Sample Density. 1.000 g/cm³

Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 5.1428 g Cold Free Space: 16.1759 cm³ Measured Low Pressure Dose: None Automatic Degas: Yes

Sample Prep: Stage $\mathbf 1$

Soak Temperature (°C) 100

Ramp Rate (°C/min) 10

Soak Time (min) 960

BET Surface Area Report

BET Surface Area: 13.2247 **±** 0.0591 m² /g Slope: 0.327644 **±** 0.001449 g/cm³STP Y-Intercept 0.001529 **±** 0.000258 g/cm 3 STP C: 215.310851 Qm: 3.0379 cm³/g STP Correlation Coefficient 0.9999217 Molecular Cross-Sectional Area: 0.1620 nm ²

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TriStar 3000 V6.05.01 A Unit 2 Port 2 Serial **#:** 1449 Page **5** Sample: CJ-210-40 06-0220 N6/N6 Operator: MJP Submitter: S.M.Stoller File: C:\...\01JAN\06-0220.SMP Started: 2/112006 4:23:29PM Completed: 2/1/2006 7:40:08PM Report Time: 2/2/2006 7:51:21AM Warm Free Space: 5.6495 cm³ Measured Equilibration Interval: 10 s Sample Density. 1.000 g/cm³ Analysis Adsorptive: N2 Analysis Bath Temp.: 77.300 K Sample Mass: 5.1428 g Cold Free Space: 16.1759 **cm³**Measured Low Pressure Dose: None Automatic Degas: Yes Sample Prep: Stage " **¹** Soak Temperature (°C) 100 Ramp Rate (°C/min) 10 Soak Time (min) 960

Summary Report

Surface Area Single point surface area at P/Po = 0.276623808: 12.9795 m²/g

BET Surface Area: 13.2247 m²/g

Appendix F

PHREEQC Input Files for Sequential Batch Reaction Models

List of Files

(1) Transport_seq_batch_horizontal
(2) Transport_seq_batch_vertical $Transport_seq_batch_vertical$

Transport_seq_batch_horizontal
TITLE Crescent Junction Transport Model - Horizontal Transport in Sequential Batch
Test Test # Includes Mineral PPT, adsorption, and IX # Transport of Tailings Water through Mancos shale # Transport of Tarrings mater embagn mantos Sh
Database "Phreeqc_Davis_Naturita_GC_Short.txt" # GC (Generalized composite) Adsorption Model (to match Davis and Curtis Figure 5.6) # React in 3 separate batch steps using "fresh" Mancos (from 40-ft depth) each step # use BH210 gw only to set the IX and adsorption sites **#** Added U to BH210 to get the right amount adsorbed to show the measured U increase SOLUTION 0 SPF-1 (Tailings Fluid). Using final control values from batch tests. temp 25 pн 7.85 pH 7.85
pe 7
redox pe redox pe
units mg/L units mg
Alkalinity Alkalinity 510.0 as CaCO3
Amm 2000 Amm 2000
Ca 196 Ca 196
Na 530 5300 Mg 310 $K = \begin{bmatrix} 129 \\ 5(6) & 182 \end{bmatrix}$ s(6) 18272 as S04 Cl 1018 SOLUTION **1** Groundwater from **CJ** BH 210. Initial solution for exchange sites # Groundwater chemistry from Borehole 210 (CJ) # sampled 11/7/05. This sample is considered represenative of Mancos groundwater # This composition sets the initial conditions for the exchange sites units mg/ temp $25.$ temp 25.0
pH 7.23 pe 12.5
Alkalinity 634 as CaCO3 Alkalinity 63
4 Amm **#** Amm 0 Ca 180
Na 120 Na 12000
Mg 140 140
58 K 58 S(6) • 1700 as S04 Cl 23000 U defined by $\mathsf{O.2}$ END EXCHANGE **1** equilibrate with solution 1 # used 400 mL-fluid with 100g rock, ie. R/W = 250g/L
Mean measured CEC=11.23 meq/L
Exchange Sites = 250 g/L x 11.23 meq/100g x 1 eq/1000 meq = 0.028 eq/l # Exchange Sites = 250 g/L x 11.23 meg/100g x 1 eq/1000 meg = 0.028 eq/L
x 0.028 SURFACE **1** # Site Density= (1.92 umoles hydroxyl sites)/mA2 (assumed bidentate bonds). **#** from Davis and Kent (1990) and **#** used in Davis and Curtis (2003) Naturita GC Model # Surface area (Mean of 10 values from 1-2 mm fraction) = 11.02 mA2/g (BET Measurement) **#** Used 250'g rock/L gw **#** Of Total Sites: very strong Sites = 0.01 %,'Strong Sites = 1.2 % (Davis and "
Curtis) # Total.Sites = 5290 umol sites/L gw (from 250*11.02*1.92) # Very Strong sites = 0.0001*5290 =,0.53 umoles/L (5.3e-7 mol/L) # Strong Sites = 0.012 * 5290 = 63.5 umoles/L (6.35e-5 mol/L) # weak sites = 0.9879 * 5290 = 5226 umoles/L (5.226e-3 mol/L) -equilibrate with solution **¹** Hfo_wOH 5.226e-3 11.02 250.0 Page **1**

Transport_seq_batch_horizonta¹ Hfo-sOH 6.35e-5 11.02 250.0 Hfo_vOH 5.3e-7 11.02 250.0 Hfo_vOH
-no_edl End Use Solution 0 use Exchange 1 Use Surface 1 Equilibrium-Phases 1 **#** The total no. of moles is based on water soluble and XRD tests of Mancos Shale # For verticle transport through deeper Mancos (in mol/Lgw):Gypsum 6.67e-4, Nahcolite 0.017 # For horizontal transport through shallower Mancos (in mol/Lgw):Gypsum 0.017, Nahcolite 6.67e-3 # Calcite **I** wgt % = 0.025 mol/Lgw for all depths. calcite was not completely dissolved in tests. # Halite 6.67e-4 mol/Lgw for all depths Calcite 0.0 0.025 Carcrice 0.0
Gypsum 0.0 0.017 Gypsum 0.0 0.017
Nahcolite 0.0 6.67e-3
Halite 0.0 6.67e-4 Halite 0.0 PRINT -reset true SELECTED-OUTPUT _outPUI
-file transport_seq_batch_horizontal.se -Tite
-totals Na Amm K Ca Cl U S(6) -molalities NaX AmmHX KX CaX2 Save Solution **0-** END use Solution 0 Use Exchange 1 use Surface 1 Equilibriunu.Phases **I .#** The total no. of moles is based on water soluble and XRD tests of Mancos shale **#** For verticle transport through deeper Mancos (in mol/Lgw):Gypsum 6.67e-4, Nahcolite 0.017 # For horizontal transport through shallower Mancos (in mol/Lgw):Gypsum 0.017, Nahcolite 6.67e-3 # Calcite 1 wgt % = 0.025 mol/Lgw for all depths. Calcite was not completely dissolved in tests. aissoived in tests.
Halite 6.67e-4 mol/Lqw for all depths c 0.07e-4 mor/Lyweron
Calcite 0.0 0.025 Carcrice 0.0
Gypsum 0.0 0.017 Nahcolite **0.0** 6.67e-3 Halite 0.0 PRINT -reset true SELECTED_OUTPUT -file transport-seq-batch-horizontal .sel -totals Na Amm K Ca **Cl** U S(6) -molalities NaX AmmHX KX CaX2 Save solution 0 END .Use solution **0..** Use Exchange 1 Use surface 1 EquilibriumPhases **1 #** The total no. of moles is based on water soluble and XRD tests of Mancos Shale **#** For verticle transport through deeper Mancos (in mol/Lgw):Gypsum 6.67e-4, * rui vertitie
Nahcolite 0.017 # For horizontal transport through shallower Mancos (in mol/Lgw):Gypsum 0.017, Page 2

with a substant that a control of the second and control the state and control

Transport_seq_batch_horizontal (\) Nahcolite 6.67e-3

wancoffice 0.07e-3
Calcite 1 wgt % = 0.025 mol/Lgw for all depths. Calcite was not completely
dissolved in tests. # Halite 6.67e-4 mol/Lgw for all depths calcite 0.0 0.025 Gypsum 0.0 0.017 Nahcolite **0.0** 6.67e-3 Halite 0.0 6.67e-4

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true transport_seq_batch_horizontal.se Na Amm K Ca CT U S(6) **NaX** AmmHX KX CaX2

Page 3

Transport_seq_batch_verticle
TITLE Crescent Junction Transport Model - Verticle Transport in Sequential Batch () Test
In **#** Includes Mineral PPT, adsorption, and Ix # Incrudes Anneral Pri, duscription, and IX
Transport of Tailings Water through Mancos Shale # Transport of Tarrings water through mantos shi
Database "Phreeqc_Davis_Naturita_GC_Short.txt # Database "Phreeqc_Davis_Naturita_GC_Short.txt"
GC (Generalized Composite) Adsorption Model (to match Davis and Curtis Figure 5.6)
React in 5 separate batch steps using "fresh" Mancos (from 40-ft depth) at first step **#** then, "fresh" Mancos from deeper (lower gypsum) for remaining steps # Use BH210 gw only to set the IX and adsorption sites **#** Added U to BH210 to get the right amount adsorbed to show the measured U increase SOLUTION 0 SPF-1 (Tailings Fluid). using final control values from batch tests. temp 25 temp 25
pH 6.63 pe redox pe
units mg/L units mg
Alkalinity Alkalinity 510.0 as CaC03 Amm 2000
Ca 196 Ca 196
2530 Na Na .5300 Mg 310
K 129 K 129
S(6) 18272 S(6) 18272 as S04
Cl 1018 Cl 1018 SOLUTION 1 Groundwater from **CJ** BH 210. initial solution for exchange sites *#* Groundwater chemistry from Borehole 210 (CJ) **#** sampled 11/7/05. This sample is considered represenative of Mancos groundwater # This composition sets the initial conditions for the exchange sites This composition sets the initial conditions for the exchange sites
units mg/L
temp 25.0 temp **25.0** pH 7.23 pe_1 , 12. Alkalinity 634 as CaCo3 **# Amm 0.** Ca 180 Na 12000 Mg 140
K 58 K 58 $\begin{array}{cc} S(6) & 1700 \text{ as } S04 \\ C1 & 23000 \end{array}$ Cl 23000 \cup 0.2 END **EXCHANGE** ¹ .. equilibrate with solution **^I** # Used 400 mL fluid with 100g rock, ie. R/W = 250g/L
Mean measured CEC=11.23 meq/L
Exchange Sites = 250 g/L x 11.23 meq/100g x 1 eq/1000 meq = 0.028 eq/l $x = 250$ g/c $x = 11.25$ and $y = 0.028$ SURFACE 1 **#** Site Density = (1.92 umoles hydroxyl.sites)/mA2 (assumed bidentate bonds) **#** from Davis and Kent (1990) and **#** used in. Davis and Curtis (2003) Naturita GC Model # Surface area (Mean of 10 values from 1-2 mm fraction) = 11.02 m \triangle 2/g (BET Measurement)
Used 250 g rock/L gw # Of Total Sites: Very Strong Sites = 0.01 %, Strong Sites = 1.2 % (Davis and Curtis) **#** Total Sites = 5290 umol sites/L gw (from 250*11.02*1.92) **#** Very Strong Sites = 0.0001*5290 = 0.53 umoles/L (5.3e-7 molQL) # Strong Sites **=** 0.012 * 5290 = 63.5 umoles/L (6.35e-5 mol/L) **#** weak sites = 0.9879 * 5290 = 5226 umoles/L (5.226e-3 mol/L) Page 1

Transport-seq-batch.verti cl e -equilibrate with solution 1 Hfo-wOH 5.226e-3 11.02 250.0 Hf_{0} 50H 6.35e-5 11.02 250.0 Hfo_s0H 6.35e-5 11.02 250.0
Hfo_v0H 5.3e-7 11.02 250.0 -no_ed1 End # Reaction with 40-ft depth sample Use solution 0 Use Exchange 1 use Surface 1 #PHASES **#** Fi xH+ **#** H+ **=** H+ 1oa_k Equilibrium_Phases 1
The total no. of m # The total no. of moles is based on water soluble and XRD tests of Mancos shale # For verticle transport through deeper Mancos (in mol/Lgw):Gypsum 6.67e-4, Nahcolite 0.017 # For horizontal transport through shallower Mancos (in mol/Lgw):Gypsum 0.017, Nahcolite 6.67e-3 # calcite 1 wgt % = 0.025 mol/Lgw for all depths. calcite was not completely dissolved in tests. # Halite 6.67e'4 mol/Lgw for all depths calcite 0.0 0.025 **Gypsum** 0.0 0.01 Nahcolite **0.0** 6.67e-3 Nahcolite 0.0 6.67e-3

Halite 0.0 6.67e-4

co2(g) -2.1 1.0 **#** . co2(g) -2.1 1.0 **#** FixH+ -7.73 C02(g) 10.0 PRINT -reset true) save Solution 0 END # Reaction with 105-ft depth sample. Use Solution 0 Use Exchange 1 Use Surface 1 Equilibrium_Phases 1 # The total no. of moles is based on water soluble and XRD tests of.Mancos Shale # For verticle transport through deeper Mancos (in mol/Lgw):Gypsum 6.67e-4, Nahcolite 0.017 # For horizontal transport through shallower Mancos (in mol/Lgw):Gypsum 0.017, Nahcolite 6.67e-3 # calcite 1 wgt % = 0.025 mol/Lgw for all depths. calcite was not completely dissolved in tests. # Halite 6.67e-4 mol/Lgw for all depths calcite **0.0** 0.025 Calcite 0.0 0.025
Gypsum 0.0 6.67e-4
Nahcolite 0.0 0.017 Nahcolite 0.0 0.017
Halite 0.0 6.67e-4 Halite PRINT -reset true Save Solution 0 END # Reaction with 170-ft sample use solution 0 Use Exchange 1 Use Surface **I1** Equilibrium-Phases **1** # The total no. of moles is based on water soluble and.XRD tests of Mancos Shale Page 2

Transport-seq-batch.verti cl e # For verticle transport through deeper Mancos (in mol/Lgw):Gypsum 6.67e-4,
Nahcolite 0.017 # For horizontal transport through shallower Mancos (in mol/Lqw):Gypsum 0.017, Nahcolite 6.67e-3 # Calcite 1 wgt $\frac{1}{x}$ = 0.025 mol/Lgw for all depths. Calcite was not completely dissolved in tests. # Halite 6.67e-4 mol/Lgw for all depths calcite 0.0 0.025 Gypsum 0.0 6.67e-4 Nahcolite **0.;0** 0.017 Nahcolite 0.0 0.017
Halite 0.0 6.67e-4 PRINT -reset true Save solution 0 END # Reaction with 235-ft sample use solution 0 Use Exchange 1 Use Surface **I** equilibrium Phases 1
The total no. of moles is based on water soluble and XRD tests of Mancos Shale # The Local no. Of moles is based on water soluble and AKD Lests of Mancos
For verticle transport through deeper Mancos (in mol/Lgw):Gypsum 6.67e-4 Nahcolite 0.017 # For horizontal transport through shallower Mancos (in mol/Lgw):Gypsum 0.017, Nahcolite 6.67e-3 # calcite 1 wgt **%=** 0.025 mol/Lgw for all depths..Calcite was not completely dissolved in tests. # Halite 6.67e-4 mol/Lgw for all depths calcite 0.0 0.025 Calcite 0.0 0.025

Gypsum 0.0 6.67e-4

Nahcolite 0.0 0.017 Nahcolite 0.0 0.017
Halite 0.0 6.67e-4 Halite PRINT -reset true Save Solution 0 END # Reaction with 300-ft sample Use solution 0 Use Exchange 1 Use Surface 1 Equilibrium-Phases 1 # The total no. of moles is based on water soluble and XRD tests of Mancos shale # For verticle transport through deeper Mancos (in mol/Lgw):Gypsum 6.67e-4, Nahcolite 0.017 # For horizontal transport through shallower Mancos (in mol/Lgw):Gypsum 0.017, Nahcolite 6.67e-3 # calcite 1 wgt % = 0.025 mol/Lgw for all depths. Calcite was not completely dissolved in tests. #.Halite 6.67e-4 mol/Lgw for all depths $\frac{1}{2}$ calcite 0.0 0.025 Calcite 0.0 0.025

Sypsum 0.0 6.67e-4

Nahcolite 0.0 0.017 Nahcolite 0.0
Halite 0.0 6.67e-4 **PRINT** -reset true Save Solution 0 END

Page 3

Appendix **G**

PHREEQC Input Files for Transport Models

List of Files

(1) Transport horizontal .(2) Transport vertical

Transport-hori zontal TITLE Crescent Junction Transport Model - Horizontal Transport Through Mancos shale Includes Mineral PPT, adsorption, and IX # Transport of Tailings Water through Mancos shale # Database "Phreeqc_Davis_Naturita_GC_Short.txt"
GC (Generalized Composite) Adsorption Model (to match Davis and Curtis Figure 5.6)
SOLUTION 0 SPF-1 (Tailings Fluid) $temp$ 25 temp 25
pH 6.63 pe 7 redox pe
units mg/L
Alkalinity 570.0 as CaC03 Amm 2181.82 ca 209.62 Na 5781.41
Ma 315.71 **Mg.** 315.71" $K = 112.23$
S(6) 17454. s(6) 17454.37 as S04 Cl 910.26 U 4.0 SOLUTION 1-40. Groundwater from **CJ** BH 210. Initial solution for column sampled 11/7/05. This sample is considered representative of Mancos groundwater
units mg/L units mg/L
temp 25.0 pH 7.23 pe 12.5 Alkalinity 634 as CaC03 **Amm** Ca 180 Na 12000 Mg 140 K 58 $\begin{array}{cc} \kappa^2 & 58 \\ 5(6) & 1700 \text{ as } 504 \end{array}$ Cl 23000 # uranium concentration was calibrated by seq batch test model to produce **#** the amount.of required adsorption. 0.2 /30 = 0.0067mg/L ^U0.0067 END # Equilibrate BH 210 Fluid with calcite and Gypsum USE solution 1 Equilibrium_Phases .Calcite 0.0 10.0 Gypsum 0.0 10.0 •save solution **I1** End # Equilibrate Soln 1 (BH 210) with exchange sites and Minerals EXCHANGE 1-20
 \leq equilibrate with solution 1
 $\# R/W = 7500$ g/L # Mean measured CEC=11.23 meq/100g
Exchange Sites = 7500 g/L x 11.23 meq/100g x 1 eq/1000 meq = 0.842 eq/L
0.842 Equilibrium Phases 1-20 # The total no. of moles is based on water soluble and XRD tests of Mancos shale # For. verticle transport through deeper Mancos (in mol/Lgw):Gypsum 0.02, Nahcolite **0.5 #** For horizontal transport through shallower Mancos (in mol/Lgw):Gypsum 0.5, Nahcolite 0.2 # Calcite 1 wgt % = 0.75 mol/Lgw for all depths. Calcite was not completely
Page 1

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Transport.hori zontal dissolved in tests. # Halite 0.02 mol/Lgw for all depths $Calcite$ 0.0 0.75 Calcite 0.0 0.75

Gypsum 0.0 0.5

Nahcolite 0.0 0.2 Nahcolite 0.0 0.2

Halite 0.0 0.02 Halite SURFACE 1-20 **#** site Density = (1.92 umoles hydroxyl sites)/mA2 (assumed bidentate bonds) **#** from Davis and Kent (1990) and **#** used in Davis and Curtis (2003) Naturita GC Model # Surface area (Mean of 10 values from 1-2 mm fraction) = 11.02 m^2/g (BET Measurement) # used 7500 **9** rock/Lgw # of Total sites: very Strong Sites = 0.01 %, Strong sites'= 1.2 % (Davis and Curtis) # Total Sites = 158,688 umol sites/L gw (from 7500*11.02*1.92) # very Strong sites = 0.0001*158688 = 15.87 umoles/L (SiM calculation) $\frac{1}{2}$ Strong sites = 0.012 $\frac{1}{2}$ 158688 = 1904 umoles/L (SJM Calculation) # weak sites = 0.9879 * 158688 = 156768 umoles/L (SJM calculation) .ccs
-equilibrate 1 Hfo_wOH 1.568e-1 11.02 7500.0 Hfo_w0H 1.568e-1 11.02 7500.0

Hfo_sOH 1.904e-3 11.02 7500.0

Hfo_v0H 1.587e-5 11.02 7500.0 1.587e-5 -no-edl PRINT -reset true ADVECTION **"** -cells 20 -shifts 20
-nunch 20 -punch 20 **#** -punch 10 20 -punch-frequency 1 $-$ print -print_frequency 50 USER-GRAPH \ldots -headings PV Na C NH4*5 Ca*10 Cl S04 U*1000 KD*25000 pH*1000 -headings PV TOTADS CS TOT("U") KD -chart-title "Crescent Junction Horizontal Transport Model" $\#$ -chart_title "Crescent Junction Horizontal Transport Mode
-axis_titles "PORE VOLUME" "Effluent Concentration (mg/L) -axis_scale x_axis 0 10.1 0.1 -axis-scale y-axis 0 25000.0 1000.0 -axis_scale y_axis o 250
-initial_solutions false -plot_concentration_vs time -piot_concentracion_vs cime
-start
5 TOTADS = MOL("Hfo_wOUO2OH") + MOL("Hfo_sOUO2CO3-") + MOL("Hfo_vOUO2OH") + MOL("Hfo_vOUO2CO3-" $6 \text{ CS} = \text{TOTADS}/7500$ 7 KD **=** (CS/TOT("U"))*1000 10 GRAPH X (STEP $NQ + 0.5$) / 20 20 GRAPH_Y TOT("Na")*23000, TOT("C")*12000, TOT("AMM")*18000*5, TOT("Ca")*40000*10, TOT("C1")*35900, ToT("s(6)")*96000, **TOT("U")*238000*1000,** KD*25000, 1000*(-LA"H+") **#** 20 GRAPHY TOT("U")*238000*1000, KD*25000 -end PRINT * -user.graph true. SELECTED_OUTPUT
file --file * **.** CJtransport.sel -totals Na Amm K Ca **Cl** NaX AmmHX KX CaX2 END

Page 2

 \mathcal{D}_1 in the \mathcal{D}_2 -considerable for the second \mathcal{D}_1 , and the second second second

Transport_vertical TITLE Crescent Junction Transport Model - Vertical Transport Through Mancos shale # Includes Mineral PPT, adsorption, and Ix # Transport of Tailings water through Mancos shale # Database "Phreeqc_Davis_Naturita_GC_Short.txt"
GC (Generalized Composite) Adsorption Model (# GC (Generalized Composite) Adsorption Model (to match Davis and Curtis Figure 5.6)
SOLUTION 0 SPF-1 (Tailings Fluid). From means in SOWP Table 9
temp 25 temp
PH $\frac{6.63}{7}$ pe 7
redox pe redox pe
units mg/L units mg
Alkalinity Alkalinity 570.0 as Caco3 Amm 2181.82
Ca 209.62 Ca 209.62
Na 5781.4 Na 5781.41
Mg 315.71 Mg 315.71
K 112.23 K^2 112.23
S(6) 17454. S(6) 17454.37 as S04 * **Cl** 910.26 U 4.0 SOLUTION 1-20 Groundwater from **CJ** BH 210. Initial solution for streamtube. # Groundwater chemistry from Borehole 210 (Cc) # sampled 11/7/05. This-sample is considered represenative of Mancos groundwater units mg/ $temp$ 25. pH 7.23 pH 7.23
pe 12.5
Alkalinity 634 a 634 as CaCO3 **#** AmM 0 الا التي تعدد التي ا
التي تعدد التي تعدد
 Na 12000 Mg 140
K 58 K 58 S(6) 1700 as S04
Cl 23000 **cl** 23000 # U value is 0.2 (used in model of Seq Batch) divided by 30 (W/R factor) $6.67e-3$ END #Equilibrate Tailings Fluid with Calcite USE solution 0 PHASES **Fix** 'H+ $H+ = H+$...
]og_k 0.0 Equilibrium_Phases calcite **0.0** 10.0 Fix_{H+} -6.63 NaOH 10.0 Save solution **0** End # Equilibrate BH 210 Fluid with calcite USE solution 1 Equilibrium_Phases Calcite 0.0 10.0
Fix_H+ -7.23 NaOH FixH+ -7.23 NaOH 10.0 **#** Gypsum 0.0 10.0 Save solution 1 End # Equilibrate soln 1 (BH 210) with exchange sites and Minerals EXCHANGE 1-20 equilibrate with solution 1 Page 1

;7,

R/W = 7500 *g/L* Mean measured CEC=11.23 meq/100g # Exchange Sites = 7500 g/L x 11.23 meq/100g x 1 eq/1000 meq **=** 0.842 eq/L $x \left(0.842 \right)$ EquilibriumPhases 1-20 # The total no. of moles is based on water soluble and XRD tests of Mancos Shale. # For verticle transport through deeper Mancos (in mol/Lgw):Gypsum 0.02, Nahcolite 0.5 # For horizontal transport through shallower Mancos (in mol/Lgw):Gypsum 0.5, Nahcolite 0.2 # calcite 1 wgt % **=** 0.75 mol/Lgw for all depths. calcite was not completely dissolved in tests. # Halite 0.02 mol/Lgw for all depths Calcite 0.0 0.75 Gypsum 0.0 0.02
Nahcolite 0.0 Nahcolite 0.0 0.5 Halite 0.0 SURFACE 1-20 # Site Density = $(1.92 \text{ umoles hydroxyl sites})/m^2$ (assumed bidentate bonds)
from Davis and Kent (1990) and **#** from Davis and Kent (1990) and **#** used in Davis and Curtis (2003) Naturita GC Model # Surface area (Mean of 10 values from 1-2 mm fraction) = 11.02 m $\triangle 2/g$ (BET Measurement): #.Used 7500 **9** rock/Lgw # Of Total Sites: Very Strong Sites = 0.01 %, Strong Sites = 1.2 % (Davis and Curtis)
Total # Total Sites **=** 158,688 umol sites/L gw (from 7500*11.02*1.92) $*$ very Strong Sites = 0.0001*158688 = 15.87 umoles/L (SJM Calculation) $\#$ Strong Sites = 0.012 $*$ 158688 = 1904.umoles/L (SJM calculation) # weak Sites **=** 0.9879 * 158688 156768 umoles/L (SJM calculation) .ccs
-equilibrate 1 Hfo_woH 1.568e-1 11.02 7500.0 Hfo_w0H 1.568e-1 11.02 7500.0
Hfo_sOH 1.904e-3 11.02 7500.0
Hfo_v0H 1.587e-5 11.02 7500.0 1.587e-5 -no-edl PRINT -reset true ADVECTION -cells 20 -shifts 20
-punch 20 -punch 20 **#** -punch 10 20 -punch.frequency 1 -print 20 -print
-print_frequency 50 USER-GRAPH -headings PV Na C NH4*5 Ca*10 cl S04 U*1000 KD*25000 pH*1000 **#** -headings PV TOTADS CS TOT('!U"),KD -chart-title "Crescent Junction Vertical Transport Model" -chart_tritles "PORE VOLUME" "Effluent Concentration (mg/L)
-axis_titles "PORE VOLUME" "Effluent Concentration (mg/L) -axis_tities Fore Vocumi
-axis_scale x_axis 0 10 1 0.1 -axis-scale x-axis 0 25000.0 1000. -axis_scale_y_axis_0_2500
-initial_solutions false -plot_concentration_vs time -start 5 TOTADS **=** MOL("HfowOUO2OH'.') **+** MOL("HfosOU02CO3-") + MOL("Hfo_..vOUO2OH") + MOL("Hfo.vOUO2CO3-") $\overline{6}$ CS = TOTADS/7500 $KD = (CS/TOT("U"))*1000$ **10** GRAPHX (STEP-NO + 0.5) / 20 20 GRAPH_Y TOT("Na")*23000, TOT("C")*12000, TOT("Amm")*18000*5 TOT("Ca")*40000*10, TOT("Cl")*35900, TOT("S(6)")*96000, TOT("U")*238000*1000, KD*25000, 1000*(-LA"H+") Page 2

 $\mathcal{P}(\mathcal{E}_\alpha,\mathcal{E}_\beta)$. The same of the set of the s

Transport-vertical

Transport_vertical
20 GRAPH_Y TOT("U")*238000*1000, KD*25000
end

PRINT -user_graph true

#

ransport.sel
mm K Ca Cl
MaX AmmHX KX CaX2

Page 3

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Appendix H

Thermodynamic Database Used for PHREEQC Models

Phreeqc_Davis_Naturita_GC_Short

Phreeqc_Davis_Naturita_GC_Short $Ca+2 + CO3-2 + H+ = CAHCO3+$ log_k 11.435 $\text{Ca+2 + } \text{SO4-2} = \text{CaSO4}$ log_k 2.3 $Mg+2 + H2O = MgOH + + H+$ log_k -11.44 Mg+2 + CO3-2 = MgCO3
 log_k 2.98 Mg+2 + H+ + CO3-2 = MgHCO3+
log_k 11.399 Mg+2 + SO4-2 = MgSO4

log_k 2.37 $\text{Na+} + \text{H2O} = \text{NaOH} + \text{H+}$
 $\log_k -14.18$ $Na+ + CO₂-2 = NaCO₂$ log_k 1.27 $Na+ + HCO3- = NaHCO3$ log_k -0.25 $Na+ + SO4-2 = NaSO4$ log_k 0.7 $K+ + H20 = KOH + H+$
 $log_k -14.46$ $K+ + SO4-2 = KSO4$ log_k 0.85 #URANIUM SPECIES FROM TABLE 2-1 DAVIS AND CURTIS NRC REPORT 2003 #U020H+ #U02(OH)2 $\frac{1002+2 + 2 + 20}{9} = \frac{1002(0H)}{11.5}$
10g_k #U02(OH)3- $UO2+2 + 3H2O = UO2(OH)3 - 3H+$ log_k -20.00 #UO2(OH)4-2 $UOZ+2 + 4H2O = UOZ(OH)4-2 + 4H+$ log_k -33.0 #(UO2)20H+3 $2002+2 + H20 = (002)20H+3 + H+$
 log_k -2.70 $\frac{1}{2}$ 597 #(UO2)2(OH)2+2 $2002+2 + 2H20 = (U02)2(OH)2+2 + 2H+$ log_k -5.62 #(UO2)3(OH)4+2 Page 3

Phreeqc_Davis_Naturita_GC_Short
3002+2 + 4H2O = (002)3(OH)4+2 + 4H+ log_k -11.90 598 #(UO2)3(OH)5+ $3002+2$ + 5H2O = (002)3(OH)5+ + 5H+
log_k -15.55 #(UO2)3(OH)7- $3002+2 + 7H20 = (002)3(OH)7 - + 7H +$ log_k -31.00 #(UO2)4(OH)7+ $4002+2 + 7H20 = (U02)4(OH)7+ + 7H+$ log_k -21.90 #U02C03 603 $UO2+2 + CO3-2 = UO2CO3$ 9.67 log_k #U02(C03)2-2 604 $1002+2 + 2003-2 = 102(003)2-2$ 16.94 log_k #U02(C03)3-4 605 $1002+2 + 3003-2 = 102(003)3-4$ 21.60 log_k #(UO2)3(CO3)6-6 648
3UO2+2 + 6CO3-2 = (UO2)3(CO3)6-6 54.0 log_k . #(UO2)2CO3(OH)3- $2\overline{u}02+2 + \overline{c}03-2 + 3\overline{h}20 = (\overline{u}02)2\overline{c}03(\overline{0}01)3- + 3\overline{h}4$ ∵log_k -0.86 #(UO2)3CO3(OH)3+ $3002+2 + 1$
 $3002+2 + 203-2 + 3H20 = (002)3C03(OH)3+ 3H+$
 log_k
 0.66 log_k #CUO2)11(CO3)6(OH)12-2
11UO2+2 + 6CO3-2 + 12H2O = CUO2)11(CO3)6(OH)12-2 + 12H+
1og k 10g 36.43 #U02N03+ 646 $UO2+2 + NO3- = UO2NO3+$ log_k 0.3 #U02C1+ 611 $\text{UO2+2} + \text{Cl} = \text{UO2Cl} +$ log_k 0.17 de Ita_h 1.9 kcal # #U02C12 644 $U02+2 + 2C1 - = U02C12$]og_k -1.1 de t a_h 3.6 kcal $#$. #U02S04 612 $\sqrt{02+2}$ + SO4-2 = $\sqrt{025}$ O4 \log_k 3.15 delta_h 4.7 kcal #U02(SO4)2-2 613

Page 4

Phreeqc_Davis_Naturita_GC_Short
U02:2 **+ 2SO4-2 = UO2(SO4)2-2** log_k 4.14 delta_h 8.4 kca PHASES Calcite CaCO3 = C03-2 **+** Ca+2 .log._k -8.48 .og__^
delt<u>a</u>_h -2.297 kca $-$ analytic -171.9065 -0.077993 2839.319 71.595 Aragonite CaCO3 = CO3-2 + Ca+2
log_k -8.336
delta_h -2.589 kcal -analytic -171.9773 -0.077993 2903.293 71.595 Dolomite CaMg(CO3)2 = Ca+2 + Mg+2 + 2 CO3-
log_k -17.09 delta_h -9.436 kca Gypsum CaSO4:2H20 = Ca+2 **+** S04-2 **+** 2 H20 log_k -4.58 del ta_h -0.109 kcal $-$ analytic 68.2401 0.0 $-$ 3221.51 $-$ 25.0627 Anhydrite CaSO4 = Ca+2 **+** S04-2 log.k -4.36 del tah -1.710 kcal uerta_n -1.710 KCa1
-analytic 197.52 0.0 -8669.8 -69.83! Halite 64 NaCl = Na+ **+ Cl-**log.k **1.582** del tah 0.918 kcal Thenardite 65 Na2SO4 = 2Na+ **+** S04-2 log_k -0.179
delta_h -0.572 kcal Nahcolite 58 NaHCO3 = Na+.+ HCO3- logq-k -0.548 del tah 3.720 kcal Trona $\sim 10^{-1}$ 59 NaHCO3:Na2CO3:2H20 2H20 **+** log1 k -0.795 3Na+ **+** C03-2 **+** HCO3 delta_h -18.0 kca^l 60 Natron Na2CO3:IOH20 = 2Na+ **+** cO3-2 + 10H20 log_k -1.311 log_k -1.
delta_h 15.745 kcal Thermonatrite 61 Na2CO3:H20 = 2Na+ **+ C03-2 +** H20 Page 5 \bigcap

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Phreeqc_Davis_Naturita_GC_Short 0.125 1og_k de ¹ $tan h$ -2.802 kca¹ # Uranium Mineral FROM TABLE 2-2 DAVIS AND CURTIS NRC REPORT 2003 602 Schoepite $U_02(OH)2:H2O + 2H+ = U_02+2 + 3H2O$ 4.93 \log_k Rutherfordine 606 $\frac{1002003}{\log k}$ = $\frac{1002+2}{-14.49}$ 600 Gummite $UO3 + 2H + = UO2 + 2 + H2O$ 10.403 log_k delta_h -23.015 kcal 599 $U₀₃(gamma)$ $\begin{array}{rcl} 1003 + 2H + & = & 002 + 2 + H20 \\ \log k & & 7.719 \end{array}$ delta_h -19.315 kcal Na4U02(C03)3 571 $\frac{100}{30002}$ (co3)3 = 4Na+ + UO2+2 + 3CO3-2
1og_k -16.290 B-U02(OH)2 601 1002 (OH)2 + 2H+ = UO2+2 + 2H2O
1og_k 5.544 delta_h -13.730 kcal # CO2 gas FROM TABLE 2-1 DAVIS AND CURTIS NRC REPORT 2003 $CO2(q)$ $\cos = \cos$ log_k -1.472
delta_h -4.776 kcal $\pmb{\sharp}$ 108.3865 0.01985076 -6919.53 -40.45154 -analytic 669365.0 $H2O(g)$ $H2O = H2O$ 1.51 log_k -44.03 delta_h kJ Stumm and Morgan, from NBS and Robie, Hemmingway, and Fischer (1978) # Gases from LLNL.dat... $Amm(g)$ $Amm = Amm$ 1.7966 loq_k EXCHANGE_MASTER_SPECIES x $X \sim 10^{-11}$ Page 6

Phreeqc_Davis_Naturita_GC_Short EXCHANGE_SPECIES \overline{x} = x - 1 oq $_k$ 0.0 $Na + X - = Nax$
 log_k 0.0
-gamma 4.0 0.075 $K + + X - = KX$ log_k 0.7
-gamma 3.5 0.015 $delta_h$ -4.3 # Jardine & Sparks, 1984 $# 11111$ \ddot{r} $H + X^- = HX$
 log_k 1.0 $-q\overline{a}$ mma 9.0 0.0 $AmmH+ + X- = AmmHX$ 0.6 1og_k 0.0 $-g\overline{a}$ mma 2.5 $d\tilde{e}$ lta_h -2.4 # Laudelout et al., 1968 $Ca+2 + 2x- = Cax2$
 log_k 0.8 $-g\overline{a}$ mma 5.0 0.165 $d\tilde{e}$ lta_h 7.2 # Van Bladel & Gheyl, 1980 $Mg+2 + 2X - = MgX2$
 log_k 0.6 $-$ gamma 5.5
delta_h 7.4 0.2 # Laudelout et al., 1968 # Surface Species FROM TABLE 5.4 DAVIS AND CURTIS NRC REPORT 2003
SURFACE MASTER SPECIES
HEO V. HEO YOH Hfo_v Hfo_vOH Hfo_S Hfo_sOH Hfo_w Hfo_wOH SURFACE_SPECIES # Very strong binding site Hfo_ss # Strong binding site Hfo_s
weak binding site Hfo_s $HfO_VOH = HfO_VOH$ log_k} 0.0 $Hf_{0,SOH} = Hf_{0,SOH}$ log_k 0.0 Hf_0 woh = Hf_0 woh : 1 oq \pm k 0.0 # Uranium Surface Species FROM TABLE 5.4 DAVIS AND CURTIS NRC REPORT 2003 # Uranyl Hfo_wOH + UO2+2 + H2O = Hfo_wOUO2OH + 2H+
log_k -6.74 Hf_0 _sOH + UO2+2 + CO2 + H2O = H f_0 _sOUO2CO3- + 3H+ $log_k - 8.00$ Page 7

Hfo_vOH + UO2+2 + H2O = Hfo_vOUO2OH + 2H+
log_k -2.06

END

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