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AN UPDATE ON THE SRP BURIAL GROUND
AREA WATER BALANCE AND HYDROLOGY

INTRODUCTION AND SUMMARY

A water budget for the burial ground area prepared by Hubbard and Emslie¹ concluded that about 15 inches, almost one-third of the average annual precipitation, normally infiltrates the land surface and recharges the groundwater. Also, evapotranspiration was estimated to average 30 inches annually, and runoff from the land surface was estimated as 1 to 3 inches.

More information has become available recently from lysimeter studies, climatic stations, groundwater studies, and stream discharge measurements. These additional data generally support the conclusions above with some modifications.

The type of vegetation cover on the land surface affects the site hydrology and water budget components of evapotranspiration and groundwater recharge. The lysimeter studies indicate that about

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12 inches more water is lost annually to the atmosphere by evapotranspiration with deep-rooted pine trees present than in areas where bare soil or shallow-rooted grass cover occur. Therefore, recharge in the burial ground area may differ from that with similar soils in forested areas of the Savannah River Plant.

Study of the hydrologic properties of soils in the burial ground area indicates that infiltration rates for the soils generally are relatively high, exceeding one inch per hour. Runoff, as overland flow tends to occur only with intense rainfall events of 1 inch or more. The soil-water characteristic curves are representative of relatively coarse-textured soils.

RESULTS

The values expressed here are "best estimates" based on the lysimeter studies, evaporation pan climatic data, and studies by other investigators. Estimates of the water budget, expressed as annual normals representative of the burial ground site, are:

rainfall	-----	48"
recharge	-----	16"
runoff	-----	2"
evapotranspiration	-----	30"

In forested areas near the burial ground, evapotranspiration is estimated to be about 40 inches annually, and therefore recharge to the water table is about 6 inches. About one-half of the recharge at the burial ground normally flows below the "tan clay."

Infiltration rates for burial ground soils are generally 1 inch per hour or greater. Surface runoff occurs infrequently, in intense rainstorms, and is unlikely to exceed 3 inches annually because of the relatively low topography, permeable soils, and vegetative cover over much of the burial ground. Problems of erosion from surface runoff tend to be limited to the exposed bare soils with a substantial clay content now located about the working area of active trenches.

DISCUSSION

1. Water Budget Estimates From Wasteform Lysimeter Data

Hubbard and Emslie¹ observed that an analysis of records of water volumes collected in Defense Waste Lysimeter studies compared well with other analysis techniques used for water budget estimates.

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The lysimeters, filled with soils from the burial ground trenches, are used for studies of leaching and waste migration. The volumes of percolate water pumped from the lysimeter sumps are measured regularly. Subtraction of the percolate volumes from the measured rainfall provided an estimate of the amounts of water lost to the atmosphere by evapotranspiration. Two additional years of lysimeter and rainfall data provided a larger sample of hydrologic processes. Records of the lysimeter studies from 1980 through 1984 provided by Emslie were used in most of this analysis.

The percolate water, collected from sumps 10 feet below ground surface, represents rainfall which infiltrated the soils of the tank lysimeters and flowed downward. The cylindrical lysimeters used in the study are exposed to normal rainfall and pumped regularly. The diameters are six or ten feet. There is no runoff because the rims of the lysimeters extend above the surface. The percolate volume thus represents the sum of runoff and groundwater recharge. Periods of three months, generally corresponding to seasons of the year, were used in the water budget analysis. January through March is called Winter in this analysis, April through June is called Spring, etc. Evapotranspiration was calculated as the difference between rainfall and the percolate volume during each three month period.

The analysis of the lysimeter data is shown in Table 1. Most of the annual recharge occurs in Winter and Spring periods. This is caused by the relatively high rainfall and the low amounts of evapotranspiration characteristic of these cooler seasons. Recharge is generally lower in Summer because of high evapotranspiration. Recharge is lower in Fall because of lower rainfall in that season.

Table 2 presents a climatic monthly evapotranspiration estimate typical of the region, based on data recorded at Blackville, South Carolina. The 34 inches estimated annual total is distributed as 7.4, 9.9, 10.8, and 6.2 inches in Winter, Spring, Summer, and Fall. The pan coefficients used with the Blackville, South Carolina data were developed for agricultural watersheds in the coastal plains near Tifton, Georgia, a region with somewhat similar topography, soils, and climate. Reported measurements of evapotranspiration in this region are few. Table 2 likely represents a better seasonal distribution of evapotranspiration amounts than does Table 1, and the higher total in Table 2 is more characteristic of areas of agricultural land.

The estimates of recharge and evapotranspiration in Table 1 are based on the water which reaches the lysimeter sump. The exact nature of the water movement as saturated and unsaturated flow in the soils is not well understood at this time. The time lag in its

movement from the surface to the lysimeter sumps reflects storage within the soil pores as well as transport processes. Some of the percolate collected as recharge may have infiltrated months before reaching the sump. Yet, after large rainfall events, water collects in the sumps in a few days. A simple tracer study has been started in a lysimeter to provide better understanding of processes involved.

2. The Time Variability of Recharge, Rainfall, and Evapotranspiration

Table 1 also indicates the relative variability for years and seasons. The coefficient of variability (C.V.) is the percentage ratio of the standard deviation to the mean. The C.V. value is greatest for recharge, 23%. Rainfall and evapotranspiration both have a C.V. of 16%. Seasonal variability is much greater than annual variability. Recharge has C.V. values exceeding 60% for Spring, Summer, and Fall seasons.

The effect on the amount of recharge during the year, related to the time distribution of rainfall, is shown in Table 1 for 1983 and 1984, years of relatively equal annual rainfall. The greater than normal Winter and Spring rainfall in 1984 resulted in 22 inches of recharge for that period. This was more than the total recharge of 1983 or of any other year in which rainfall amounts were more evenly distributed. But the period of low rainfall in Fall 1984 resulted in less recharge and less evapotranspiration than in 1983. Some lysimeters contained no percolate water at all in Fall 1984.

The elevation of the water table below the burial ground appears to have a time lag of about six months with respect to periods of excess rainfall. A relationship between recharge and the response of water table well C-21 is shown in Figure 1. To construct the figure, the assumptions were made that the yearly recharge is 16 inches and that recharge of 4 inches each quarter would maintain the water table. Recharge exceeding that amount would cause the water table to rise, and recharge less than 4 inches would cause it to fall. The cumulative net recharge from 1980 through Spring 1985 and the relative water table elevation are plotted in the figure to show an association.

3. Problems with Using the Lysimeters for Water Budget Estimates

The hydraulic performance of the Defense Waste Lysimeters was examined by Emslie.² The statistical analysis pointed out problems with lysimeter data associated with leaks in the systems and pump failure. Therefore, only unsaturated six- and ten-foot lysimeters at the site were used for water budget estimation. The six-foot lysimeters used in updating 1983 and 1984 data are

lysimeters #1, #6, #8, #11, #14, #18, #32, and #38. Ten-foot lysimeters #10 and #35 also provided representative data for the burial ground study. None of the lysimeters has vegetative cover as dense as the grass cover at 643-G.

4. The Effect of Vegetation Type on Evapotranspiration and Recharge

Some of the unsaturated ten-foot diameter lysimeters support the growth of the pine trees 10 to 14 feet high. Examination of the percolate volumes from these lysimeters suggests that the deep-rooted pines transpired at least 12 inches more water than the sparse cover of bahia grass and herbacious plants in most lysimeters.

Figure 2 shows the cumulative curves of 1984 rainfall and percolate volumes for lysimeters with and without pines present. The 1983 data also show a difference of about 12 inches. The increased transpiration from areas where more deeply rooted trees are present is expected to result in considerably less groundwater recharge.

5. The Results of Other Studies of Water Budgets in the SRP Region

McQueen Branch watershed, located a short distance northeast of the burial grounds, was studied by Parizek and Root.³ A weir was installed for stream discharge measurement. Surface runoff events were separated from base flow of the stream by hydrograph analysis. Although the McQueen Branch watershed differs from the burial ground area in some respects, estimates (Table 3) of 30 inches of evapotranspiration, 3 inches of surface runoff, and most of the remainder annually as recharge, are generally in agreement with those made by Hubbard and Emslie.¹

The McQueen Branch study provides additional information about groundwater movement in the Barnwell Formation above and the McBean Formation below the "tan clay." An estimated annual flow of 7 inches above the "tan clay" and 8 inches below the "tan clay" may be inferred. This is in reasonable agreement with the estimate for the burial ground area reported by Hubbard and Emslie.¹

Denehy and McMahon⁴ have completed a lysimeter study at the Chem-Nuclear waste disposal site in Barnwell, South Carolina, to be published by the U. S. Geological Survey. These data, taken from July 1983 through June 1984, showed that almost all of the recharge occurred in Winter and Spring 1984. The measurements show almost no runoff, 15 inches of recharge, and 30 inches of evapotranspiration annually. V. Ischamura of Chem-Nuclear, has begun a new water balance study at the Barnwell site (personal communication).

6. An Attempt to Analyze Hydrogeologic Control of Stream Discharges

The hydrogeologic stratigraphy described by Hubbard and Emslie¹ includes the Barnwell Formation, overlying and meeting an aquitard, the "tan clay," at an elevation 200 feet above sea level. The McBean Formation, the next layer in the sequence of formations, is similarly bounded 130 feet above sea level by another aquitard, the "green clay". The Congaree Formation is the next groundwater zone below the "green clay".

Table 4 presents a comparison of discharges of streams near the burial ground. Measurements reported by the U. S. Geological Survey⁵ are compared with those made of McQueen Branch by Parizek and Root³ from April 1983 through March 1984. The discharges and rainfall amounts are expressed in inches. The stream flow reflects surface runoff and discharges from groundwater zones.

The U. S. Geological Survey stream gage for Four Mile Creek at Road 4 is assumed to be located just above the "tan clay." Values representing the "natural behavior" of the creek were calculated by subtracting discharges at two H Area gages from the Road 4 gage. McQueen Branch includes flows from below the "tan clay" in the McBean Formation, according to Parizek and Root.³ Table 4 shows that McQueen Branch has a greater annual streamflow yield than Four Mile Creek.

Similarly, Upper Three Runs at Road A, where the gage is assumed to be below the "green clay", yields greater streamflow than at the Road C gage above the aquitard. This indicates base flows from Congaree Formation groundwater discharge. Further analyses of gaged streams may provide more quantitative understanding of groundwater-surface water interaction near the burial ground site.

7. Runoff Studies in the Burial Ground Area

Runoff estimates were made using daily precipitation and soils information, by a method used by the U. S. Department of Agriculture Research Service and described by Lane.⁶ The daily precipitation data were taken from 200-F Area, and soils information descriptive of the site was used to determine a curve number (CN) representing the hydrologic soil group present. A CN of 72 was selected. This refers to Soil Group A, characterized by bare soils with infiltration rates exceeding 0.30 inches per hour, which would be a conservative representation of the active burial ground. A CN of 90 also was selected for comparison. This refers to Soil Group D, characterized by clay soils with very low infiltration rates.

The runoff calculations were made by using the equation

$$Q = \frac{(P - 0.2S)^2}{(P + 0.8S)}$$

where S = a retention parameter (in.) = $\frac{1000}{CN} - 10$

P = daily rainfall (in.), and
 Q = daily runoff (in.)

Results of the runoff calculations for 1984, shown in Table 5, provide an estimated annual runoff of 3 inches. Runoff from clay soils, however, with a yield of 15 inches, could produce considerable surface erosion. The extensive gully development of F-Area Effluent Stream may have been caused by similarly increased runoff from the parking lots prior to 1978.

Table 5 also shows the runoff estimated for Four Mile Creek, using cumulative flow frequency curves for Four Mile Creek described by Hubbard and Emslie.¹ Total surface runoff in that well-vegetated watershed was 1.5 inches in 1984.

In the study of McQueen Branch, Parizek and Root³ analyzed stream gage hydrographs to separate surface runoff from base flow. About 2 inches average runoff occurs annually in that watershed, as shown in Table 3.

8. Infiltration Studies

A single-ring infiltrometer was constructed from a section of Shelby tube, rubber stopper rings, and a section of glass tubing inserted into the Shelby tube. The apparatus was implanted into the ground and filled with water. The rate of movement into the ground was observed as infiltration took place.

In three scouting measurements made at the burial ground site, the lowest rate observed was about 3 inches per hour. In another measurement the rate was 8 inches per hour, and in the tritium control lysimeter 3 inches of ponded water infiltrated the surface in 20 minutes. Further infiltration measurements should involve more sophisticated tools, such as rainfall simulators like those used by the U. S. Department of Agriculture, or other devices that can be fabricated.

The thesis by Gruber⁷ presents data on soil-water movement in an area of natural soil just west of the burial ground. Several days were required for ponded water in this area to reach a steady rate of infiltration, 0.9 inches per hour. This rate may simply reflect the hydraulic conductivity of a least-permeable layer of the soil horizons 10 feet below the land surface.

Most of the burial ground area soils have high rates of infiltration. However, clay soils brought to the surface in burial ground activities can be considerably less permeable. The contrast between Class A and Class D soils, described in the discussion on runoff and in Table 5, is considerable. At the Chem-Nuclear burial trench site in Barnwell, trench caps are made from clays excavated from the trenches. Runoff amounts in those areas can be fairly large and are being measured by diversion into a surveyed basin (V. Ischamura, personal communication).

9. Hydrologic Properties of Burial Ground Soils

The behavior of water in the earth materials below the burial ground involves saturated flow, unsaturated flow, and storage in the soil pores. The heterogeneous nature of the soils and geologic formations must be considered in any analysis. The work by Gruber^{7,8}, shows laboratory analyses of cores including a trench core from the site. Figure 3 presents the means, taken from curves presented in the thesis,⁷ of the fraction of soil volume filled with water in relation to soil moisture tensions.

The mean moisture content of saturated soils is slightly less than 50%. The moisture tension represented by 25 mm Hg would be unsaturation at field capacity, and perhaps that at 75 mm Hg would be representative of wilting point. Because hydraulic conductivity is a function of soil moisture tensions, these data may be useful for modeling. The difference in void space between saturation and field capacity appears to be about 10 per cent. Saturated flow of large amounts of rainfall infiltrating the soils may be rapid through large voids in the earth materials.

RECOMMENDATIONS

1. Water balance studies utilizing the lysimeters should be continued, particularly if the leaking lysimeters identified can be easily repaired. Rainfall measurements should be supplemented by those from the rain gage installed at the Tank 24 Saltstone Lysimeter station. The format for analyzing hydraulic data from the lysimeters, set up on the SRP computer system by R. Emslie, should be maintained.

2. The method of forecasting water table elevations in the burial ground area should be examined further, using a network of wells in recharge areas. The method, described in section 2 and Figure 1, may be useful with respect to areas with different soils and vegetative cover.
3. The tracer study noted in Section 1 should provide some insight into water movement through earth materials at the burial ground and should be repeated after analysis of this trial to determine the flow paths from surface to sump.
4. With additional development in areas near the burial ground site, McQueen Branch watershed should be gaged regularly, and the data should be related to rainfall, surface flow, and shallow groundwater flow.
5. The U. S. Department of Agriculture Research Service developed a CREAMS model to predict runoff, erosion, and chemical transport. This model should be investigated for potential applications at SRP.

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TABLE 1

EVAPOTRANSPIRATION ESTIMATE BY LYSIMETER WATER BALANCE*

Water Budget Components (Inches)									
1980-84 Means, (%Annual), (Standard Deviations)									
Time Periods	Rainfall P			Infiltration SR+Rx			Evapotranspiration ET		
Jan. - Mar.	14.4	(31%)	(3.7)	7.5	(38%)	(2.7)	6.9	(25%)	(1.1)
Apr. - Jun.	11.9	(25%)	(4.5)	6.4	(33%)	(4.3)	5.5	(20%)	(1.6)
Jul. - Sep.	13.2	(28%)	(3.4)	2.9	(15%)	(2.0)	10.4	(38%)	(1.8)
Oct. - Dec.	7.5	(16%)	(2.5)	2.8	(14%)	(1.8)	4.7	(17%)	(1.7)
Annual Total	47.0"		(7.6)	19.6"		(4.5)	27.4"		(4.5)

Period	1980			1981			1982		
	P	SR+Rx	ET	P	SR+Rx	ET	P	SR+Rx	ET
Jan. - Mar.	17.3	10.0	7.3	8.5	3.3	5.2	13.4	6.7	6.7
Apr. - Jun.	7.1	4.0	3.1	12.2	5.6	6.6	10.7	3.3	7.4
Jul. - Sep.	9.3	1.3	8.0	11.1	1.4	9.7	18.4	5.8	9.7
Oct. - Dec.	4.7	1.6	3.1	7.9	4.9	3.0	8.6	2.4	6.2
Annual	38.4	16.9	21.5	39.7	15.2	24.5	51.1	18.2	32.9

Period	1983			1984			1985		
	P	SR+Rx	ET	P	SR+Rx	ET	P	SR+Rx	ET
Jan. - Mar.	17.6	9.5	8.1	15.2	8.2	7.0	11.4	6.0	5.4
Apr. - Jun.	10.4	5.1	5.3	19.2	13.9	5.3	7.4	4.6	2.8
Jul. - Sep.	13.4	1.9	11.5	14.0	4.0	10.0			
Oct. - Dec.	10.9	4.4	6.5	5.3	0.8	4.5			
Annual	52.3	20.9	31.4	53.7	26.9	26.8			

* a. SRP unsaturated wasteform lysimeters; data provided by R. Emslie. Analysis by plotting cumulative percolate of nondeviant lysimeters.

b. Precipitation (P) as rainfall was recorded at the burial ground. Runoff and Groundwater Recharge (SR + Rx) infiltrated the lysimeters, percolated 10 feet, and was pumped from the lysimeter sumps. Evapotranspiration (ET) was taken as the difference, P - (SR + Rx).

TABLE 2

EVAPOTRANSPIRATION ESTIMATE
BY EVAPORATION PAN METHOD*

Month	Pan Coeff.	Actual Pan Evaporation (inches)		Estimated ET (inches)		
		1983	1984	1983	1984	Mean
January	1.02	1.40	1.93	1.43	1.97	1.70
February	0.83	2.38	3.05	1.98	2.53	2.26
March	0.65	4.28	5.20	2.78	3.38	3.08
April	0.52	5.14	5.10	2.67	2.65	2.66
May	0.46	7.18	6.96	3.30	3.20	3.25
June	0.52	7.42	7.90	3.86	4.11	3.98
July	0.60	8.20	7.02	4.92	4.21	4.56
August	0.56	7.51	6.15	4.21	3.44	3.82
September	0.47	5.27	4.92	2.48	2.31	2.40
October	0.55	3.94	4.20	2.17	2.31	2.24
November	0.81	2.72	2.73	2.20	2.21	2.20
December	1.02	1.38	2.00	1.41	2.04	1.72
Total		56.82	57.16	33.41	34.36	33.87

* Pan evaporation data were obtained from the South Carolina Edisto Agricultural and Educational Center at Blackville. Pan Coefficients were developed from watershed studies at the U. S. Department of Agriculture Agricultural Research Service Experiment Station at Tifton, Georgia.

TABLE 3

WATER BUDGET FOR McQUEEN BRANCH WATERSHED*

Component	Percentage of Rainfall	Average Year (Inches)
Precipitation		47.72
Evapotranspiration	63.9	30.53
Surface Runoff - H Area	2.5	1.19
Surface Runoff - Basin	4.3	2.05
Base flow	22.6	10.80
Groundwater storage	6.6	3.15

* From reference 3.

TABLE 4

RAINFALL AND STREAMFLOW FOR FOUR SITES NEAR THE BURIAL GROUND
FOR WATER YEAR 1983-1984*

Month	Rainfall (inches)	Streamflow (inches per month)			
		McQueen Branch	Four Mile	Upper Three Runs at Road C	Upper Three Runs at Road A
April 1983	4.71	1.48	1.96	1.64	2.10
May	2.33	1.07	0.94	1.23	1.25
June	3.66	0.79	0.50	1.13	1.11
July	2.94	0.65	0.18	0.97	0.94
August	4.68	1.05	0.24	0.91	0.89
September	4.62	0.84	0.70	1.19	1.15
October	1.64	0.76	0.08	1.03	1.00
November	5.16	1.26	0.28	1.31	1.43
December	2.89	1.08	0.79	1.43	1.57
January 1984	2.73	1.16	1.73	1.37	1.52
February	5.15	1.50	2.08	1.53	1.68
March	6.77	2.16	2.31	1.66	2.00
Total	47.28	13.80	11.79	15.40	16.64

* Rainfall and McQueen Branch streamflow are from reference 3, Table 5. Streamflow data for Four Mile Creek and Upper Three Runs Creek recorded by the U. S. Geological Survey were converted from cubic feet per second to area inches per month. The drainage areas and estimated elevations above sea level are:

McQueen Branch at weir: 4.4 square miles; 170 feet msl.
 Four Mile Creek at Road 4: 5.9 square miles; 205 feet msl.
 Upper Three Runs at Road C: 176 square miles; 130 feet msl.
 Upper Three Runs at Road A: 203 square miles; 120 feet msl.

TABLE 5

COMPARATIVE ESTIMATES OF SURFACE RUNOFF IN 1984:
 COMPARING THE USDA METHOD OF ANALYSIS WITH
 THE CUMULATIVE FLOW FREQUENCY METHOD*

Month	Precipitation at 200F Area, in.	Runoff by SCS Method		Runoff by Cumulative Flow Freq. Analysis, in. (# days)
		Case A, in. (# days)	Case D, in. (# days)	
Jan.	3.54	0.00	0.61 (4)	0.04 (2)
Feb.	5.34	0.65 (3)	2.68 (3)	0.24 (4)
Mar.	6.05	0.21 (1)	1.68 (7)	0.08 (2)
Apr.	7.11	0.37 (3)	2.27 (6)	0.32 (5)
May	10.73	1.42 (5)	5.29 (7)	0.39 (2)
Jun.	1.82	0.00	0.10 (2)	0.05 (2)
Jul.	6.46	0.04 (1)	1.35 (9)	0.17 (4)
Aug.	3.53	0.47 (1)	1.49 (1)	0.20 (2)
Sep.	1.06	0.00	0.20 (1)	0.00
Oct.	0.40	0.00	0.02 (1)	0.00
Nov.	0.97	0.00	0.03 (1)	0.00
Dec.	1.16	0.00	0.06 (2)	0.00
Total	48.17	3.16 (14)	15.78 (44)	1.49 (23)

* Runoff Case A represents a bare soil with high infiltration rates, even when wetted. Runoff Case D represents a bare soil with very slow infiltration rates, such as a clay soil (see reference 6). Cumulative flow frequency analysis of Four Mile Creek discharges is described in text. (# days) refers to the number of days on which runoff occurred.

FIGURE 1
GROUNDWATER RECHARGE AND WATER TABLE ELEVATION 1980-1985

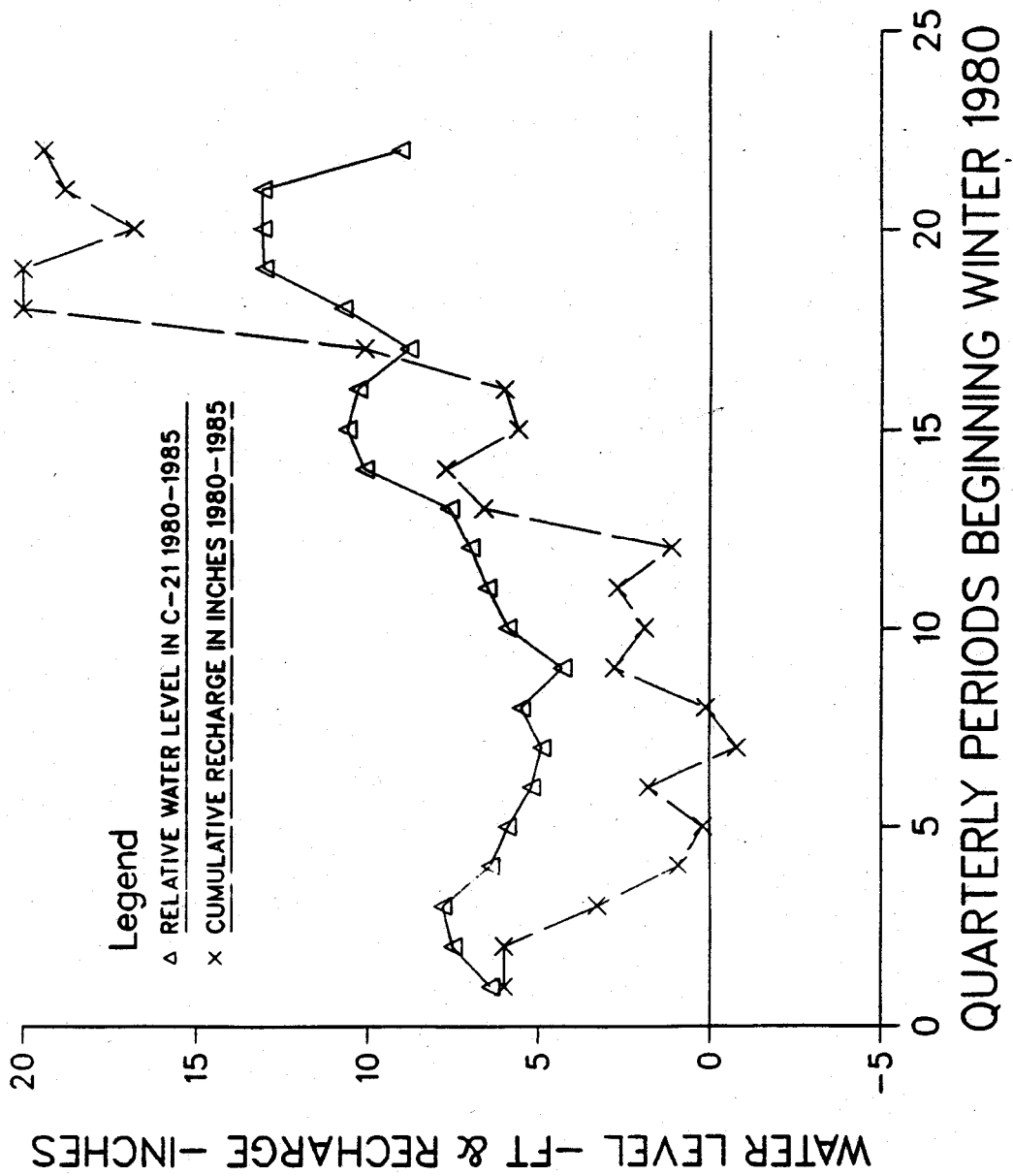


FIGURE 2
1984 RAINFALL AND PERCOLATE IN TEN-FOOT DW LYSIMETERS

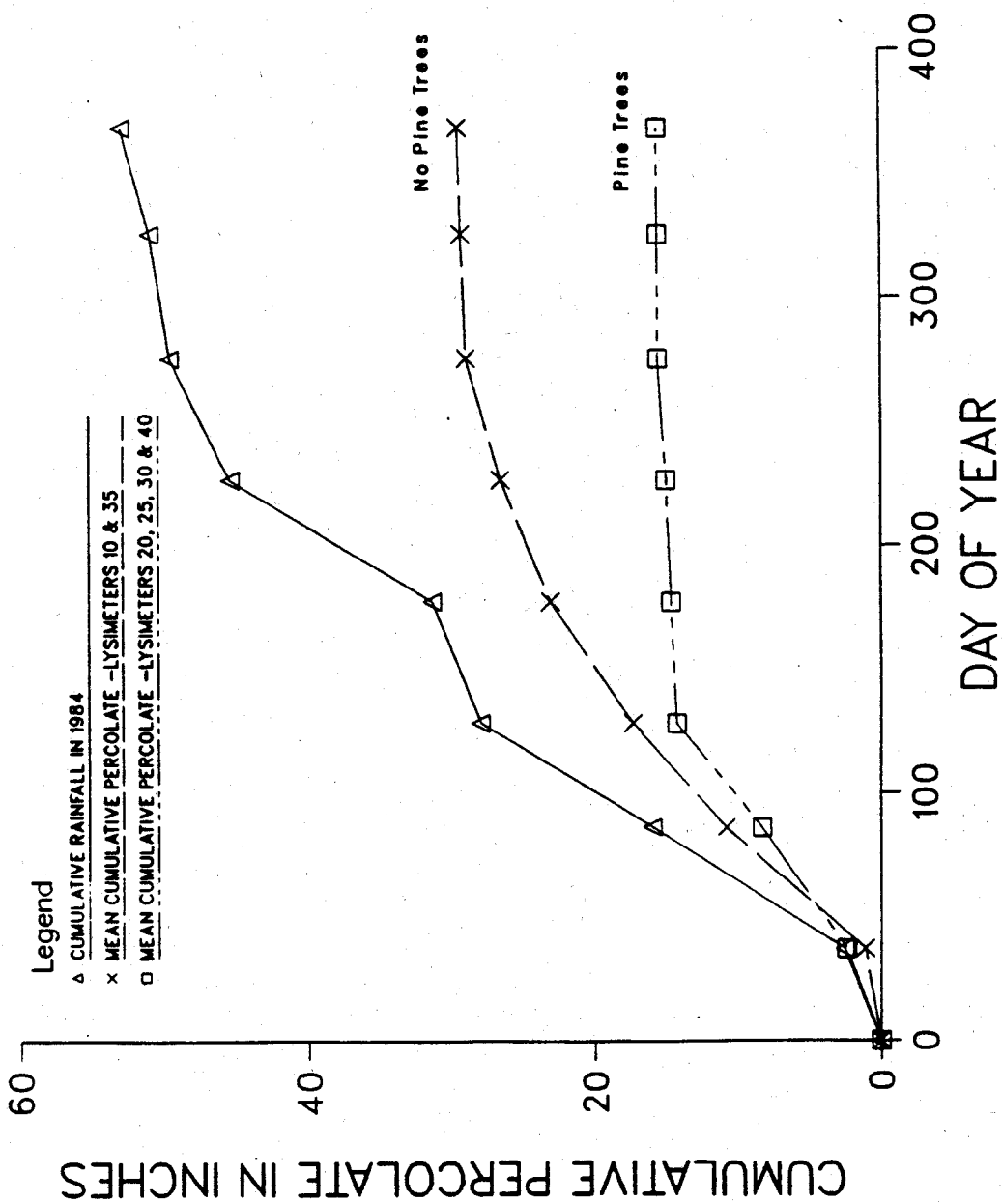


FIGURE 3

MEAN SOIL-WATER CHARACTERISTIC CURVE FOR TRENCH SOIL CORES

