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NUMERICAL SIMULATION OF GROUNDWATER FLOW FOR WATER RIGHTS ADMINISTRATION IN THE LEA COUNTY UNDERGROUND WATER BASIN NEW MEXICO

by

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ABSTRACT

Since 1931, the New Mexico Office of the State Engineer (OSE) has administered the ground waters of the Lea County Underground Water Basin, located within Lea County, eastern New Mexico. Water rights administrative criteria adopted by the OSE in 1952 have expired, and the Hydrology Bureau was requested to develop a groundwater flow model of the basin to be used in evaluating water rights applications. The Hydrology Bureau was further requested by Lea County to develop a model that could be used in local planning efforts, and in particular to evaluate the effects of groundwater pumping in Texas on water levels in New Mexico. To meet these objectives, a two-dimensional (one layer aquifer system), finite-difference numerical model of the High Plains groundwater flow system was developed. This report documents the approach used to model the aquifer, and the data base and the methods used to calibrate the model.

The groundwater flow model code, was calibrated for the Lea County Underground Water Basin in two phases. In the first phase, a steady-state (pre-development) model of 1948 groundwater flow conditions was used to constrain the relationship between the model parameters; recharge, hydraulic conductivity, evapotranspiration, and natural discharge. In the second phase, or transient calibration, historical changes in water levels from 1948 to 1996 were used to evaluate the specific yield. A trial and error procedure was used to calibrate aquifer parameters to field data. In this procedure parameters are estimated within a plausible range of values and a simulation run is performed. Simulated water levels are compared to observed water levels to check if an acceptable match is obtained. If the difference between observed and simulated water levels is unacceptable, aquifer parameters are adjusted and another simulation is performed. The process continued until an acceptable match was obtained between simulated and observed water levels. The steady-state calibration process was terminated when 67% of simulated water levels were within 10 feet of the observed, pre-developed water levels in wells where there are observational data. Transient calibration was terminated when 65% of simulated drawdowns are within 10 feet of the observed water level declines at the end of the simulation period.

Sensitivity analysis was performed to study the effects of changes in estimated aquifer parameters (hydraulic conductivity, recharge, and specific yield) on simulated water levels. Aquifer parameters tested were varied while others remained unchanged from the standard value. Aquifer response to variations in hydraulic conductivity and recharge rates was tested for pre-developed conditions. The response of the aquifer to variation in specific yield was tested for transient conditions. The model is most sensitive to an increase in hydraulic conductivity or a decrease in areal recharge. Slight increases in hydraulic conductivity (10% increase) or slight decreases in areal recharge (5% decrease) resulted in many dry cells. However, decreases in hydraulic conductivity, increases in recharge rates, or changes in specific yield caused simulated

water levels to change by 5 to 10 feet but did not have a significant effect on model results.

The effect of continued current withdrawals on water levels in the aquifer were projected to the year 2040. Future projections are important to determine how long the current saturated thickness of the Ogallala Formation would sustain current production rates. The current saturated thickness in the basin varies from one place to another and ranges from 25 feet in some areas to approximately 255 feet in the central part of the basin. Under current withdrawal rates, water level declines of 0.5 to 2.5 and 2 to 4 feet per year are simulated in New Mexico and Texas respectively. In Texas, in Yoakum County where the current saturated thickness is less than 50 feet, some areas will be de-watered as early as 2010. The de-watered area will increase as pumping continues to the year 2040. In New Mexico, a few areas located along the New Mexico/Texas state line are de-watered by the year 2040. In the vicinity of Hobbs, for the year 2040, the remaining saturated thickness ranged between 50 to100 feet to the north to less than 50 feet south of the city. In the vicinity of Lovington, for the year 2040, the remaining saturated thickness ranged from 50 to150 feet. A greater remaining saturated thickness is observed north of Knowles in New Mexico (100 to 200 feet). In general, groundwater withdrawals in New Mexico are sustainable for 40 years whereas groundwater withdrawals in Texas, especially in Yoakum County will result in aquifer de-watering. However, these projections are based on the estimated withdrawal rates used in historical simulations for current conditions. These rates are approximately 40% of the permitted water rights in New Mexico.

The effect of Texas pumping on New Mexico water levels was examined by performing a historical simulation with no pumping in New Mexico. Simulated drawdowns indicated that Texas pumping induced 10 to 20 feet of drawdown in New Mexico. The greatest drawdowns are along the state line (more than 20 feet) and decrease as we move further from the state line into New Mexico. A 10 foot drawdown is estimated in the vicinity of Hobbs, Knowels, and Nadine in New Mexico, due to Texas withdrawals.

A calibrated model is a valuable tool for water rights administration and an aid to planning efforts. As a tool for water rights administration, the model can provide an estimate of the amount of water remaining for appropriation and calculate the effects of proposed wells upon existing water rights. As an aid to local planning efforts, the model can be used for calculating regional effects and evaluating different predictive scenarios of future groundwater pumpage in both New Mexico and Texas.

I. INTRODUCTION

The principal source of water for agricultural, municipal, and industrial uses in Lea County is the High Plains aquifer. The area is a subdivision of the High Plains aquifer system that underlies 174,000 square miles in parts of Colorado, Kansas, Nebraska, New Mexico, Oklahoma, South Dakota, Texas, and Wyoming. The High Plains aquifer underlies the northern two-thirds of the county, as well as adjoining areas in western Texas. It includes parts of Lea County, southeastern New Mexico, and parts of the contiguous Cochran, Yoakum, Gaines, and Andrews counties in western Texas. The model simulates groundwater flow in the unconfined Ogallala Formation and related sediments of Tertiary and Quaternary age which locally comprise the High Plains aquifer.

Since the drilling of the first wells in the early decades of this century, it was recognized that progressive development of the aquifer will result in groundwater mining. To address this, areas of Lea County overlying the High Plains aquifer were declared an underground water basin by the Office of the State Engineer (OSE) for the purposes of ensuring the orderly development of the limited resource and protection of existing water rights. By declaring the basin, groundwater appropriation was made subject to a regulatory, permitting process. In 1952 the OSE adopted 40-year administrative criteria permitting limited mining of the aquifer, reserving the bottom one-third of the aquifer for use beyond 1992. Under this criteria, pumping effects and the life expectancy of the aquifer were computed by the Theis non-equilibrium formula using uniform values for transmissivity (T=100,000 gpd/ft) and specific yield (S=0.2) for the entire area.

Anticipating the need to update the criteria, the OSE's Water Rights Division requested the Hydrology Bureau to develop a model for the purposes of evaluating pending and future water rights applications. At the same time, both the OSE and Lea County identified the need to develop a groundwater flow model that could be used in local planning efforts, and in particular to evaluate the effects of groundwater pumping in Texas on water levels in New Mexico. An existing U. S. Geological Survey (USGS) numerical groundwater model (McAda, 1984) was developed for the Lea County Underground Water Basin. This model was evaluated and information from that study was utilized in the development of an updated model suitable for basin administration.

This report documents the development of a two-dimensional, finite-difference model of the High Plains aquifer in Lea County and the adjacent area in west Texas. The report summarizes the relevant work of previous investigators on the High Plains aquifer, but the emphasis is on the data base used to calibrate the model. The report also documents the ability of the model to reproduce measured pre-development (1948) water levels and subsequent (1948-1996) changes in water table elevation.

II. LEA COUNTY UNDERGROUND WATER BASIN

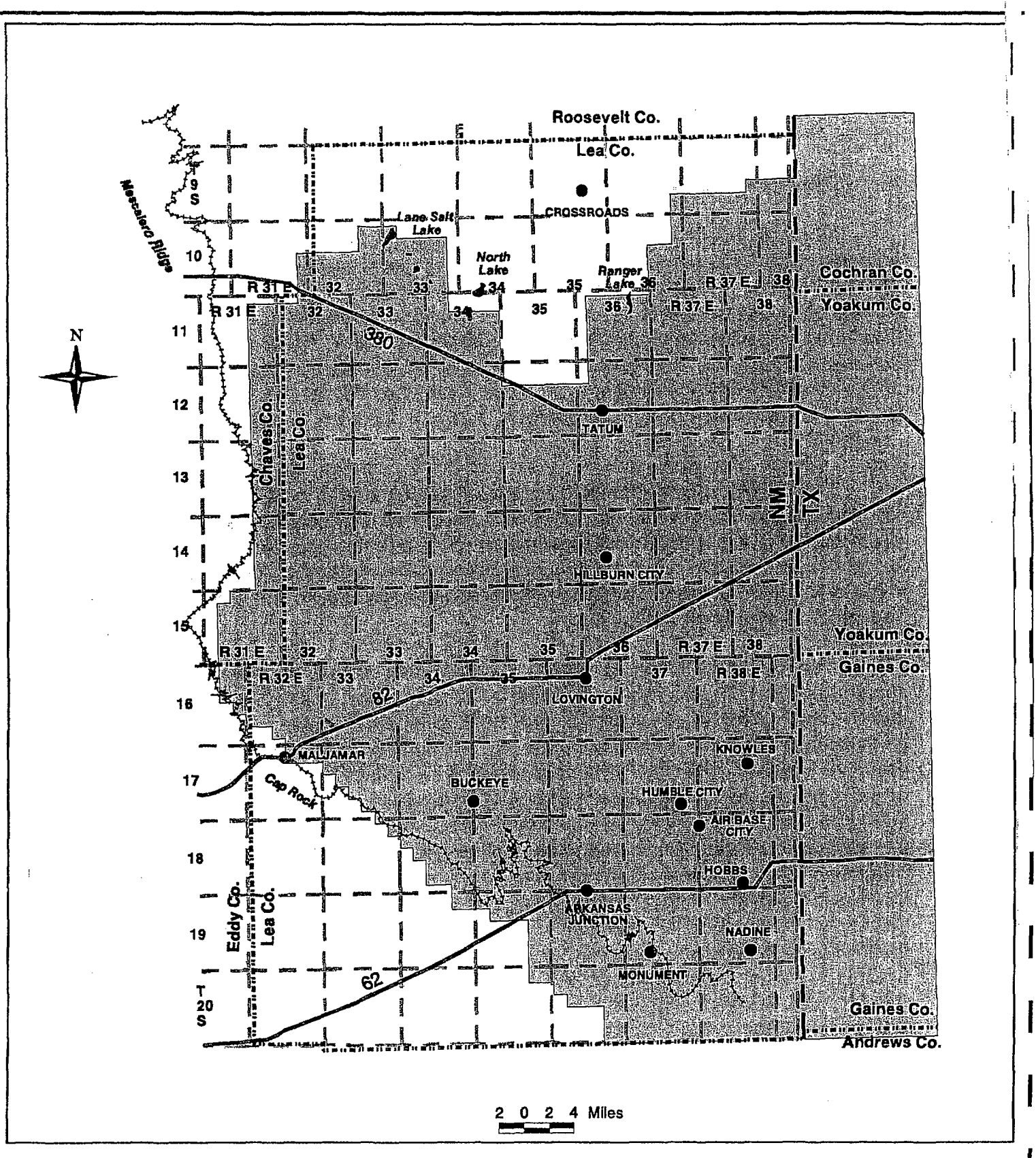
A. Study Area Description.

The Lea County Underground Water Basin is a declared basin encompassing approximately 2900 square miles within northern Lea County, southeastern New Mexico (Figure 1). The study area lies within the Southern High Plains of eastern New Mexico and western Texas. It extends from the eastern boundary of Chaves and Eddy Counties to approximately nine miles into Texas in parts of Yoakum and Gaines Counties to the east. The northern and southern boundaries of the study area includes portions of Cochran and Andrews Counties in western Texas. In this report, references to Lea County include portions of Gaines, Yoakum, Chaves, Cochran, and Andrews Counties.

The Ogallala Formation is the principal aquifer of the High Plains in Lea County. Unconsolidated sand, silt, clay, and gravel are the primary sediments of this Formation. Calcium carbonate cements the Ogallala sediments to form a caliche cap rock near the top of the formation. This caliche cap rock extends over most of the area. It is also found within the formation but generally decreases with depth until it becomes negligible at depths of 35 to 50 feet.

The Ogallala Formation overlies a pre-Ogallala erosional surface developed on Cretaceous and Triassic rocks. Triassic rocks composed of fine-to coarse-grained sand, clay, and shale are known to yield small amounts of water in local areas. Cretaceous rocks are composed of the Edwards Limestone and Paluxy Sandstone and are also known to yield small amounts of water. The Ogallala Formation is the major water yielding material in eastern New Mexico aquifers. This formation is truncated against the underlying Cretaceous and Triassic rocks in the northern and western part of Lea County and ranges in saturated thickness from zero to approximately 255 feet. Saturated thicknesses of 150 to 255 feet are found in the central part of the basin between Hobbs and Tatum in New Mexico and extends east into southern Gaines County in Texas. Smaller saturated thicknesses (50 to100 feet) are found in areas north-west of Lovington. Saturated thicknesses less than 50 feet are found in Cochran County in Texas, south of Monument, and at the north central boundary near the six lakes in New Mexico. Detailed descriptions of the sedimentation processes, sediments, and depositional environment are given by Nye, 1930; Nicholson and Clebsch, 1961; Ash, 1963; Havens, 1966; Cronin, 1969.

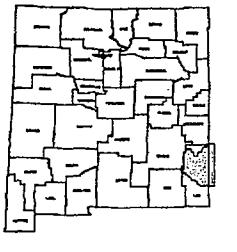
The Lea County area is a farming/ranching community with two primary population centers, Lovington and Hobbs. Groundwater is the primary and only source of water in the region. In New Mexico, the major agricultural pumping centers are located along the state line, in the vicinity of Hobbs, Lovington, and in the area



Legend

- City
- County Line
- State Line
- Mescalero Ridge
- US Highway
- Township/Range Line
- Lake
- Underground Water Basin

Figure 1: Underground Water Basin



between Hobbs and Lovington. Domestic and commercial pumping centers are distributed all over Lea County, New Mexico. In Texas, agricultural and commercial pumping centers are located in almost every cell and covers the entire area. Pumping centers and withdrawal rates are discussed in detail later in this report.

There are six perennial lakes near the northern boundary of the area that receive water from runoff and groundwater flow. These lakes are Lane Salt Lake, Ranger Lake, and four small lakes as shown in Figure 1. There are also many small shallow playa lakes scattered over the area. Most of these playas are perennial, losing collected water to infiltration and evapotranspiration although some retain water throughout the year.

B. Hydrological Setting and Aquifer Characteristics.

1. Water Table.

Under pre-developed conditions, groundwater in the High Plains aquifer is unconfined and flows southeastly, paralleling regional topographic and depositional gradients. The flow gradient varies form one place to another and averages 10 to12 feet per mile. Localized groundwater production has created numerous cones-of-depressions which caused reversed gradients around extensively irrigated areas. In New Mexico, major cones of depression are located along the state line extending from Hobbs in the south to Tatum in the north. In Texas, major cones of depression are concentrated in Gaines and Yoakum Counties. Larger cones-of-depressions are observed in Texas due to more concentrated pumping in the area.

Depth to the water table in the study area varies depending on surface and bedrock topography, proximity to discharge and recharge areas, and proximity to agricultural and urban well fields. Water table depths range from less than 20 feet near the four lakes and Monument, to more than 250 feet along Mescalero Ridge.

2. Hydraulic Conductivity.

Two fundamental characteristics of an aquifer system are its hydraulic conductivity (k), a measure of the aquifer's ability to transmit water, and the thickness of the water bearing material (b). The product of these two is aquifer transmissivity (T), a parameter that is usually estimated using several methods such as, aquifer tests, specific capacity test, grain size analysis from cores, laboratory column, and permeameter tests.

Aquifer tests designed to estimate transmissivity in the study area were concentrated around agricultural fields. They are site specific and reflect the highly variable nature of the aquifer due to grain size heterogeneity of the lithologic materials and subsequent interstitial cementing. Moulder and Frazor (1957) reported hydraulic conductivity values ranging between 21 to 25 feet per day from aquifer tests conducted near Amarillo, Texas. A wider range of hydraulic conductivity values (13 to 23 feet per day) was obtained by Cronin and Wells (1960) from long term aquifer tests near Plainview, Texas.

Several investigators have tried to estimate conductivity values for the Ogallala Formation in Texas and New Mexico using several methods. Mantei and others (1966-67) estimated a hydraulic conductivity of 112 feet per day in the Portales Valley based on laboratory tests and an electric analog model. Knowles and others (1984) mapped aquifer hydraulic conductivity and transmissivity for the High Plains in Texas using drillers logs and test cores. They reported a hydraulic conductivity range between 3 and 262 feet per day for the area. Myers (1969) obtained a conductivity range between 3 and 230 feet per day from 19 wells in 6 counties in Texas. He concluded that an average value of 57 feet per day is a reasonable estimate.

Estimated hydraulic conductivity for the Ogallala Formation in Lea County, New Mexico, showed lower values than the Ogallala Formation in Curry County, although the aquifers in both counties are part of the High Plains aquifer. Havens (1966) obtained an average conductivity value of 43 feet per day in Lea County, New Mexico for a recompacted sand sample collected from a Lovington well. Theis (1934), based on data from tests conducted near Hobbs by Nye (1932), recommended a conductivity value between 41 and 48 feet per day. Theis (1932) also performed laboratory analyses on samples collected in and near Lea County and estimated a hydraulic conductivity range between 2 to 17 feet per day for the Ogallala Formation.

Table 1 summarizes ranges and methods used to determine hydraulic conductivity for the Ogallala Formation. In general site-specific values of hydraulic conductivity ranged from 2 to 200 feet per day. As mentioned earlier, they reflect the highly variable nature of the aquifer due to grain size heterogeneity of the lithologic materials and subsequent interstitial cementing.

3. Specific Yield.

The specific yield of an unconfined aquifer is the volume of water that can be drained by gravity from the saturated sediments per unit drop in the water table per unit area. Specific yield may be determined by aquifer tests, laboratory core testing, and calculating the amount of water pumped divided by the volume of sediments de-watered.

STUDY AREA	HYDRAULIC CONDUCTIVITY VALUE RANGE (ft/d)	AVERAGE VALUE (ft/d)	METHOD	NUMBER OF TESTS	DATA SOURCE
Lea County, NM	41- 48		Aquifer Tests		Theis, 1934
Lea County, NM	2 - 17		Laboratory Analysis		Theis, 1934
Curry County, NM			Specific Capacity		Howard, 1954
Amarillo, Texas	21 - 25		Aquifer Tests		Moulder and Frazor, 1957
Plainview, Texas	13 - 23		Aquifer Tests		Cronin and Wells, 1960
Portales Valley, NM	112				Mantei, et. al., 1967
Lea County, NM	43		Estimated		Havens, 1966
Lovington, Texas	13		Laboratory Analysis		Havens, 1966
Portales Valley, NM	53 to 214				Galloway and Wright, 1967
Southern High Plains, Texas	< 200		Aquifer Tests		Myers, 1969
High Plains, Texas	10 - 230	56	Aquifer Tests	19	Myers, 1969
High Plains, Texas	3 - 262	30	Drillers Logs, Cores	41 Test Holes	Knowles, et. al., 1984
USGS Lea County GW Model	3 - 262		Model Calibration		McAda, 1984

Table 1: Tertiary Ogallala Hydraulic Conductivity Values from Various Sources (Musharrafieh and Logan, in press).

Knowles and others (1984) analyzed cores from 41 test holes to determine specific yield of the aquifer material. They found that finer sediments have low specific yield and that specific yield appeared to be evenly distributed throughout the saturated thickness of the Ogallala Formation. Alexander and others (1943) computed an average specific yield of 0.14 near Plainview and Hereford, Texas, by dividing the net amount of water pumped from 1938 to 1943 by the volume of de-watered sediments. Specific yields of 0.09 to 0.16 and 0.11 to 0.14 were obtained from tests conducted near Amarillo, Texas (Moulder and Frazor, 1957), and Plainview, Texas (Cronin and Wells, 1960) respectively. In Lea County, Havens (1966) used the same methodology used by Alexander and others (1943) to compute a specific yield of 0.24 from data collected from 1930 to 1955. Based on these and other studies, McAda (1984) in his model calibration, used a specific yield range between 0.12 and 0.25 for the Ogallala Formation in Lea County. Table 2 summarizes ranges and methods used to determine specific yield for the Ogallala Formation.

4. Groundwater Recharge and Discharge.

Semi-arid climatic conditions in the Southern High Plains region are characterized by low average annual precipitation and high average annual evaporation rates. Mean annual precipitation in Lea County ranges from 13 to 17 inches and potential evaporation (evaporation from a free water surface) ranges from 39 to 52 inches per year. Summers are usually low in humidity with clear, warm to hot, sunny days and cool nights. Most precipitation occurs from May through October corresponding to the highest seasonal evaporation rates. Winter temperatures are moderate with some severe cold spells. The average snowfall is light and the snows usually melt within a few days after occurrence. Light to moderate winds prevail most of the year except for January to May where strong winds are common.

Aquifer areal recharge in the Southern High Plains occurs by direct infiltration of areal precipitation, infiltration from playas and arroyos, and possibly irrigation return flows (Nativ and Riggio, 1990; Dvoracek and Black, 1973). The actual amount of recharge through infiltration of direct precipitation and through the playas varies depending on the amount and distribution of precipitation, soil type, and whether the underlying material restricts downward movement of water. In Lea County, it is estimated that only 0.25 to 0.5 inches of precipitation per year will infiltrate and recharge the aquifer (Theis, 1934; Theis, 1937; White and others, 1946; Cronin, 1961; Havens, 1966). The contributions of the playas to recharge is still not well understood. Nativ and Riggio (1990) found that areal recharge estimates varied by two orders of magnitude due to different interpretations of the importance of playa lake contributions to aquifer recharge. Playas are characteristic of the topography of the Southern High Plains with drainage areas capturing from eighty to ninety percent of rainfall (Dvoracek and Black, 1973). Early investigators believed that most water collected in a playa

STUDY AREA	SPECIFIC YIELD VALUE RANGE	AVERAGE	METHOD	NUMBER OF TESTS	DATA SOURCE
Lea Co., NM	0.28		Samples		Theis, 1934
Plainview, Hereford, Texas	0.14-0.15		Calculated, Q/de-watered Sediments		Alexander, et. al., 1943
Southern High Plains	0.15-0.20 Undisturbed	0.29 Compacted	Laboratory Analysis	8	Barnes, et. al., 1949
Amarillo, Texas	0.09-0.16		Aquifer Tests		Moulder and Frazor, 1957
Plainview, Texas	0.11-0.14		Aquifer Tests		Cronin and Wells, 1960
Portales Valley, NM	0.20 Storage Coeff.				Mantei, et. al., 1966
Lovington, NM	0.38		Laboratory	1, Repacked Sample	Havens, 1966
Lea County, NM		0.24	Calculated Q/de-watered Sediments		Havens, 1966
USGS Lea County GW Model	0.12-0.25		Model Calibration		McAda, 1984
Portales Valley, NM	0.20 Storage Coeff.		Reported		Galloway and Wright, 1967
Southern High Plains, Texas	0.072-0.195			19	Myers, 1969
Southern High Plains, Texas			Drillers Logs		Knowles, et. al., 1984
Southern High Plains, Texas			Cores	41 Test Holes	Knowles, et. al., 1984

Table 2: Specific Yield Values from Various Sources (Musharrafieh and Logan, in press).

evaporates and very little will infiltrate (Dvoracek and Black, 1973). Later studies indicated that infiltration through playas is important and higher recharge rates were observed near playas. Subsequent studies in the 1990's have shown that playas have different recharge profiles depending on their sedimentologic characteristics and moisture content at the time of a runoff event. In general, playas are focused recharge points where infiltration rates are initially very high and decrease exponentially with time. Mullican and others, (1994) modeled several recharge scenarios and found that focused (playa) recharge gave similar results to areal recharge. More information on aquifer recharge can be found in Dvoracek and Black, 1973; Kier and others, 1984; Stone, 1984; Wood and Osterkamp, 1984; Stone and McGurk, 1985; Scanlon and others, 1994; Zartman and others, 1994.

Recharge can also occur through ephemeral stream channels, arroyos, and irrigation return flow. A caliche layer occur in arroyos in this area, and it was originally thought that this layer would prevent infiltration (Cronin, 1964). Later investigators have decided that recharge through the caliche layer is possible through fractures and root tubes (Allison and others, 1985). Irrigation return flow is also a source of recharge in some areas. The amount of irrigation return flow reaching the aquifer depends on the depth to the water table and the loading rate. Stone's (1984, 1990) chloride tests indicated that 10.32 inches/year of irrigation returns to the Ogallala Formation, or 344 times more than natural recharge reaches the aquifer. Knowles and others (1984) found that the rate of water movement in the vadose zone is approximately 1 foot/year and that only in areas with a shallow water table irrigation return may be a source of recharge.

Methods to determine recharge rates and details on recharge investigations using these methods are described by Nativ and Riggio (1990); Stone and McGurk (1985); Stone (1984); Klemt (1981); Barnes (1949). In spite of all these studies, total recharge to an area is commonly unknown, difficult to estimate, and frequently determined through model calibration efforts.

Natural groundwater discharge in the study area occurs through evapo-transpiration when the water table is close to the land surface and through surface evaporation from lakes and playas. In general, surface discharge in the study area is of limited extent near the community of Monument, and is associated with the shallow depth to the water table. Natural discharge through subsurface flow into Texas and discharge through artificial means, or groundwater extraction constitute a major source of water loss from the Lea County Underground Water Basin. These topics will be further discussed in the modeling section.

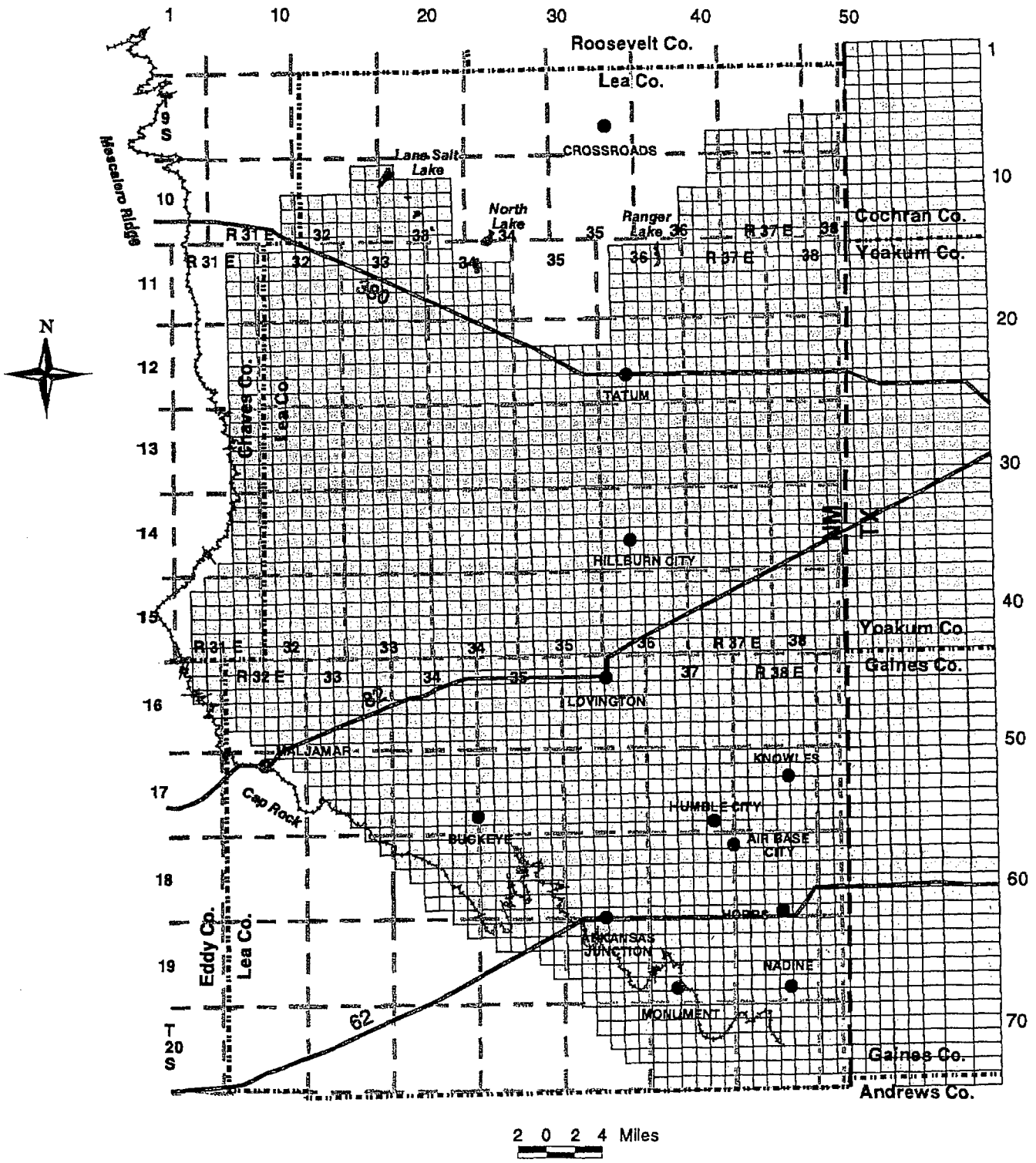
III. MODEL DESCRIPTION AND DEVELOPMENT

The modeled area covers 2916 square miles of which 2400 square miles are in New Mexico and 516 square miles are in western Texas (the model boundary extends 9 miles into Texas). In New Mexico, the modeled area includes the underground water basin declared by Office of the State Engineer. The 516 square miles into Texas are introduced to minimize boundary effects on the area of interest and to evaluate the effect of agricultural and municipal supply wells in Texas on water levels in Lea County, New Mexico.

The three-dimensional finite-difference model code developed by McDonald and Harbaugh (1988) was used to simulate a two-dimensional flow in the single layer (Ogallala Formation) aquifer in the Lea County Underground Water Basin. Model development began by discretizing the study area into a finite difference grid to facilitate a numerical solution. The area is divided into 74 rows and 56 columns as shown in Figure 2. Except for the last three columns in Texas, the grid size is one square mile. Column 54 is one mile long and 1.5 miles wide and columns 55 and 56 are one mile long and 2 miles wide. The purpose of the model is to predict average aquifer response to current and future pumping scenarios in Lea County, New Mexico, and adjacent areas in Texas. A one square mile model grid in the area of interest is adequate.

Following model discretization, boundary conditions are specified under pre-developed conditions. The types of boundaries that could be simulated depend on the proximity of the physical system to adjacent water sources. Boundaries could be "no flow", "constant head", and "specified flow". A "no flow" boundary is specified if the flow domain is adjacent to an impermeable geological formation that prevents water from entering or leaving the system. A "constant head" boundary is specified if the basin is hydraulically connected to a large water body. This type of boundary allows unlimited amount of water to enter or leave the system maintaining a constant hydraulic head at that boundary. A "specified flux" boundary allows a constant rate of water to enter or leave the flow domain independent of the hydraulic head at the boundary. A "no flow" boundary is a special case of a "specified flow" boundary where the amount of water entering or leaving the flow domain is zero.

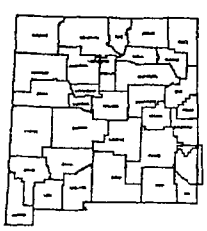
Boundary conditions used in the model for pre-development steady-state, simulations are presented in Figure 3. A "no flow" boundary is simulated along the Mescalero Ridge at the western edge of the modeled area due to the feathering out of the saturated Ogallala Formation. The north-central boundary is simulated as a "specified flow" boundary due to the presence of six perennial lakes in the area that recharge or discharge the aquifer. A "constant head" boundary is specified at the artificial eastern boundary (in Texas) to simulate natural subsurface flow out of the study area. A "constant head" boundary is also simulated along the northern side of

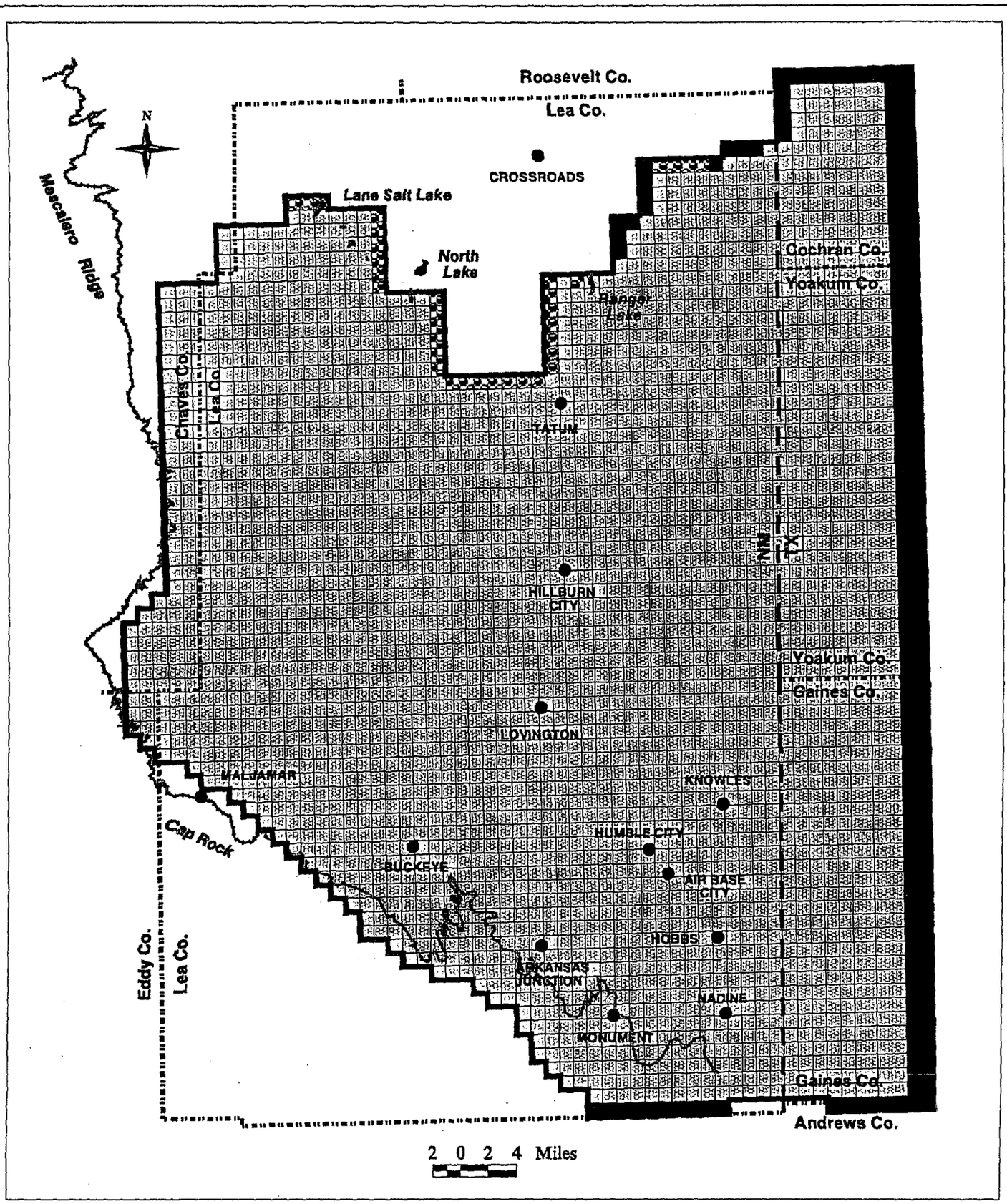


Legend

- City
- County Line
- State Line
- Mescalero Ridge
- US Highway
- Township/Range Line
- Lake
- Model Grid

Figure 2: Model Grid





Legend

- Model Boundary
- Variable Head Cell
- Constant Head Cell
- Specified Flow Cell
- State Line
- County Line

Figure 3 : Pre-Development Boundary Condition Simulated by the Model

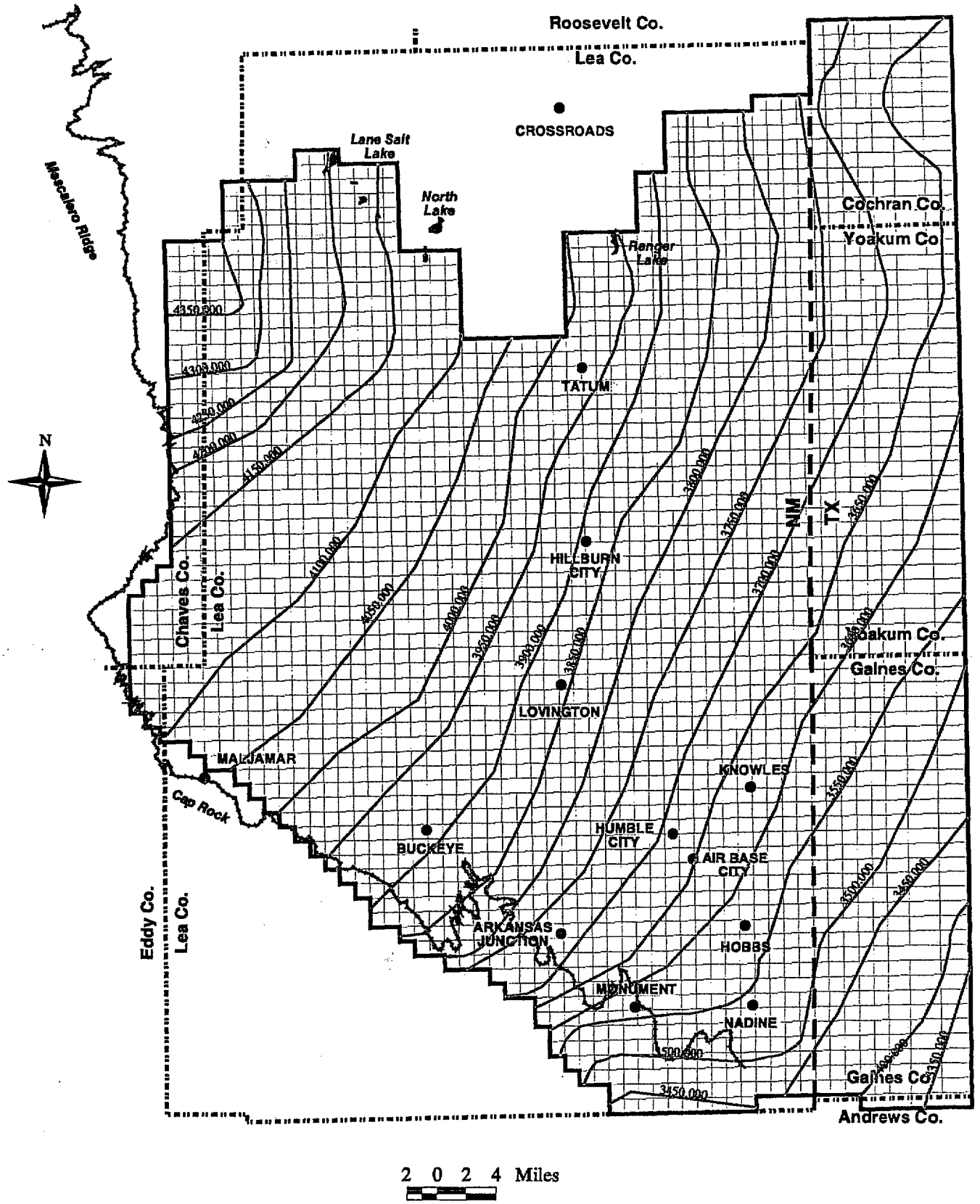
the basin where permeable Cretaceous rocks are present which permits water to flow from/to the modeled area. The southern boundary of the basin is simulated as a "constant head" and a "no flow " boundary. This was based on observed pre-developed water level contours in the area. The lower boundary of the Ogallala Formation is simulated as a " no flow " boundary since little leakage occur between this formation and the underlying impermeable material (Dutton and Simpkins, 1986; Nativ 1988; Fallin, 1989). The upper boundary, or the land surface, is simulated as "specified flux " since recharge and evapotranspiration occur through this boundary. Recharge and evapotranspiration from the Ogallala Formation are discussed in details later in the report.

After specifying boundary conditions, observed pre-development water levels and the elevation of the underlying impermeable material (Figures 4-5) were compiled and entered into MODFLOW files (Ash, 1963). Aquifer hydraulic conductivity, recharge rates, and evapotranspiration rates were estimated and also compiled. Estimates of specific yield are not necessary for steady-state simulations.

Historic pumping in both Lea County and contiguous areas in Texas were incorporated into the transient simulation. Irrigation wells (pumping centers) in Lea County were assumed to be located in close proximity to the irrigated lands they serve. The model assumes that every cell containing irrigated land has a well located in the center of the cell. Hence the term pumping center is a more appropriate description. Maps prepared in 1960, 1978 and 1991 were used to locate and quantify the amount of New Mexico irrigated acreage in each model cell. The 1960 maps were prepared by the OSE's Hydrographic Survey Bureau (Garvey 1961). The 1978 maps were prepared by the USDA-Soil Conservation Service (USDA-SCS,1978). The 1960 and 1978 maps were prepared from ground surveys. The 1991 maps were prepared for the OSE by a private contractor, and are based on aerial photography.

The 1991 map was used to approximate 1983-1996 conditions with respect to pumping center locations and initial acreage estimates; the 1978 map was used to approximate 1963 to 1983 pumping center locations and initial acreage estimates; and the 1960 distribution of irrigated acreage was used to approximate pre-1963 pumping center locations and initial acreage estimates.

Since the initial acreage estimates that were made using the maps had to extrapolate 48 years of acreage values (1948-1996) from only three years of data, additional refinements to the acreage estimates were made using annual estimates of irrigated acreage prepared by the OSE and others (Wilson, 1986; Wilson, 1992; Wilson and Lucero,1997; publications of the NMSU Agricultural Experiment Station). The annual surveys were not used directly because experience has shown that the acreage maps prepared from intensive field surveys are more accurate than the annual surveys. For example, the 1991 aerial survey found approximately 20% less irrigated acreage



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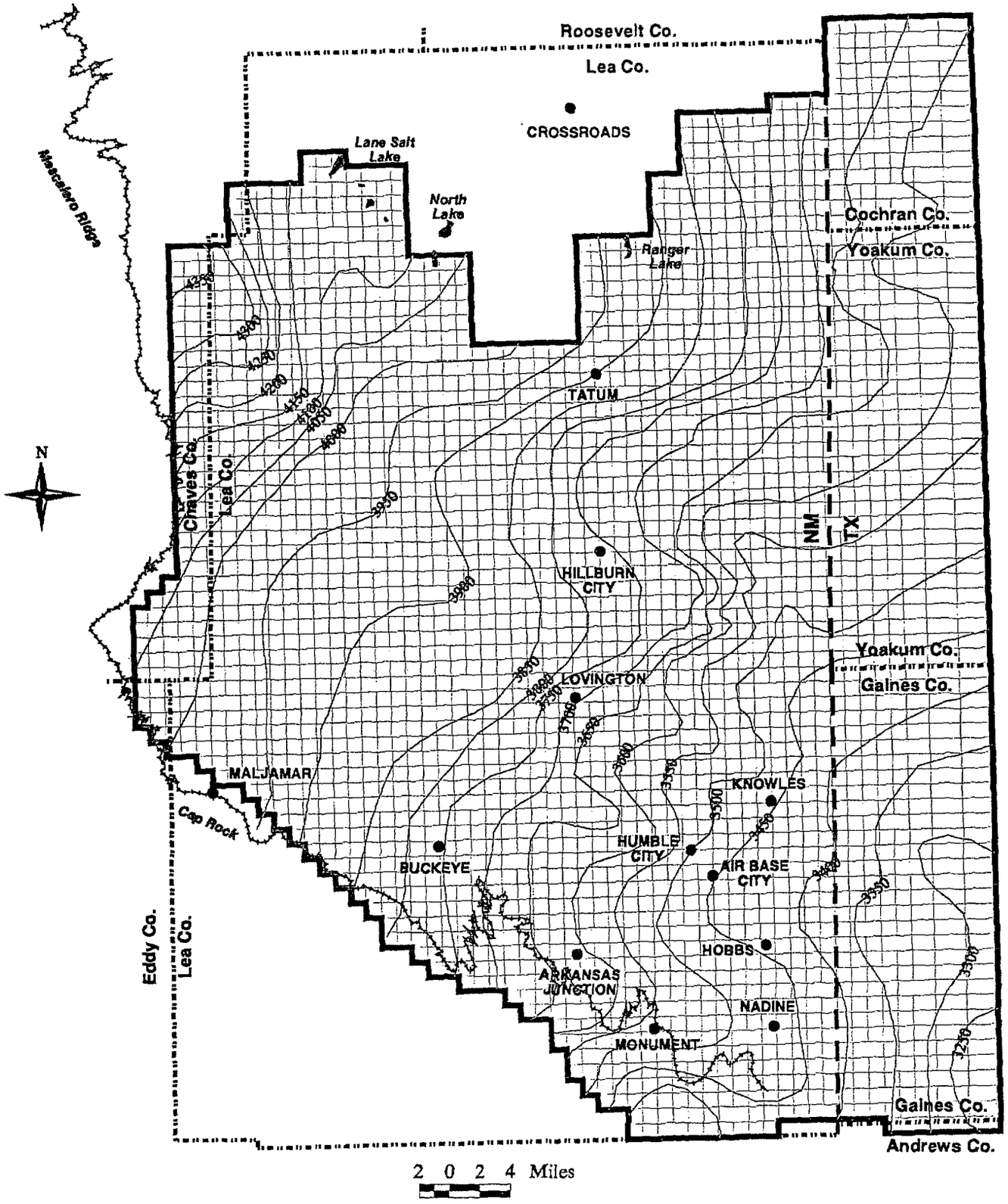
Model Boundary
State Line
County Line

Figure 4: Observed Pre-Development Water Table Elevation (ft)



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Model Boundary

State Line

County Line

Figure 5: Bottom Elevation (ft)

than the amount reported in the annual survey. However, the annual surveys were found useful in tracking annual percent changes in irrigated acreage. Therefore, the annual acreage estimates were used to proportionately increase or decrease the acreages obtained from the 1960, 1978 and 1991 maps. Consecutive years with relatively similar reported irrigated acreages in Lea County were combined into a single stress period. For the purposes of model calibration, pumping from wells from 1948 to 1996 were compiled into 7 stress periods (Figure 6 and Table 3).

The vast majority of Lea County irrigation diversions are not metered. Estimates of irrigation diversions were made for each stress period by multiplying the irrigated acreage estimate for each model cell for each stress period by a farm delivery requirement (FDR) value. The FDR is calculated from the average crop irrigation requirement (CIR) values. An average CIR value of 2.14 AF/acre for the period of simulation was calculated by Brian Wilson of the Water Use and reports Bureau. Based upon a farm irrigation efficiency of 65% (Wilson, 1997), a farm delivery requirement (FDR) of 3.29 AF/acre is estimated.

Recharge studies in Texas found that, with few exceptions, most tested fields that had been irrigated 25 years or more showed evidence of deep percolation only to depths of 20 to 30 feet (Klemt, 1981). The initial transient calibration run was made assuming that the estimated irrigation diversion (FDR) of 3.29 AF/acre is depleted with no return flow to the aquifer.

Municipal and commercial well locations in New Mexico were obtained from the well permits on file with the OSE. Pumping rates for Lea County municipal and commercial/industrial (M&I) wells were obtained from the Office of the State Engineer meter reading records and the annual reports of the basin administrator. Meter readings were available for most wells from the 1960's to the end of the transient calibration period. Prior to this period, M&I well diversions were relatively small and were assumed based upon available information. M&I withdrawals from the basin are assumed to be entirely depleted from the aquifer with no return flow to the aquifer.

Texas pumping center locations were obtained from the Texas Southern High Plains model GWSIM-III (Knowles, 1981). Pumping centers in the modeled area are shown in Figure 7. Historic well discharge estimates for all uses for the portion of the model in Texas were obtained from input files used in the Texas Water Development Board's (TWDB) Texas Southern High Plains model GWSIM-III (Knowles, 1981; Peckham, personal communication, 1995). The TWDB and OSE model grids are not compatible. Therefore, it was reasonable to create an overlay of the OSE model grid on the Texas model grid to convert TWDB model cell values to OSE model values. The TWDB model included pumpage estimates only for the period 1960 to 1996. Pumpages for the period 1948 to 1960 were estimated by prorating estimates of the number of irrigation wells and acreage published by Cronin (1964) to the 1960

Stress Period	Total Pumping	Texas Pumping	New Mexico Pumping
1948-1953	222138	103497	118641
1953-1958	299891	148959	150932
1958-1963	316662	154506	162156
1963-1978	235611	140460	95151
1978-1983	326030	158438	167591
1983-1993	250786	118511	132274
1993-1996	252790	130162	122628

Table 3 Pumping Rates for Each Stress Period Simulated in the Model (afy).

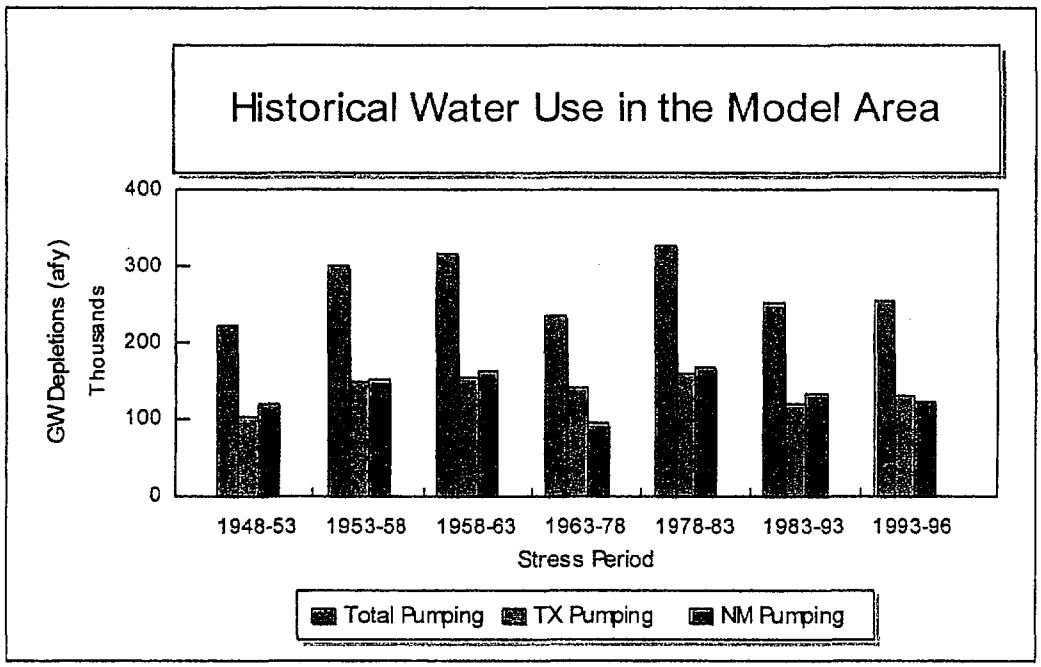
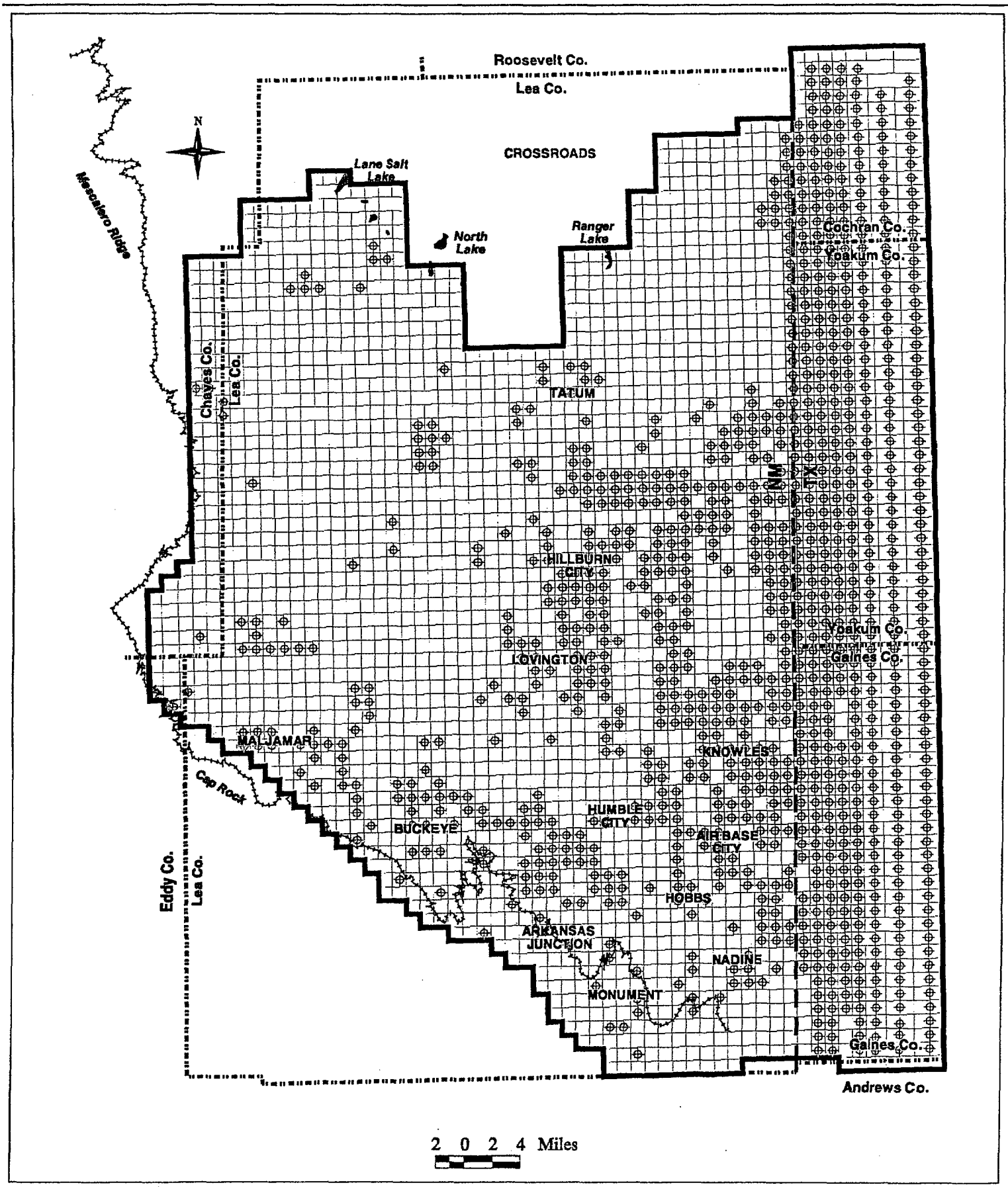


Figure 6. Total Pumping Simulated in the Model.



Legend

Model Boundary

State Line

County Line

Figure 7: Location of Pumping Centers Simulated in the Model

pumpage estimated model values. Pumpage estimates for the model area for all uses are provided in Table 3 and Figure 6.

Initial model runs resulted in simulated drawdowns exceeding observed drawdowns in the basin, especially in New Mexico. Since a stress period is the length of time in which pumping rates are assumed to be constant, the most likely source for this error is in using a fixed, average estimate of irrigated acreage for each stress period. Errors are also associated with extrapolating 48 years (1948-1996) of irrigation acreages from only three reliable field surveys (1960, 1978, 1991). A second possible source of error is in the assumption of no return flow from irrigation anywhere in Lea County. Another possible source of error is in the assumption that irrigators uniformally apply an average of 3.29 af/acre every year.

Realizing the uncertainty associated with the above methodology for estimating well depletions, and that additional data to better refine estimates of acreages and application rates are not available, depletion rates were treated as a lumped parameter and selectively reduced in some cells to obtain a better match between simulated and observed drawdowns in the basin. Table 3 gives the final aquifer depletions for each stress period simulated by the model. The average annual ground water depletion rate for Lea County for the 1948-1996 period is approximately 136,000 acre feet.

Since these errors do not influence the steady-state calibration, this phase of the calibration will be more useful in obtaining an acceptable distribution of hydraulic conductivity and recharge compared to the transient calibration.

For transient simulations, boundary conditions were also specified. Boundary cells represented as "no flow" and "specified flow" cells under steady-state conditions were represented the same under transient conditions. Boundary cells to the north, west, and south of the modeled area that were represented as "constant head" cells under steady-state conditions are represented as "specified flux" in subsequent transient simulations. Specified flows in these cells were obtained from the simulated steady-state flows. Constant head boundary on the eastern edge of the model in Texas presented a problem as to how it should be represented in transient simulations. It is not possible to represent the boundary as a "constant head" because water levels are changing due to pumping along this boundary in Texas. It is also not possible to represent it them as "no flow" because it is clear that extensive pumping in Texas will induce flow across the boundary. A possible representation of this boundary is as a "specified flux" boundary. The amount of specified flow through this boundary is assumed because of lack of field data in Texas. Since pumping wells exit along this boundary a specified flux is added to each cell so that the total withdrawals from these cells ranged between 75% (in stress period four) to about 90% (in the other stress periods) of simulated steady-state outflows.

To test whether the representation or location of the eastern boundary has an effect on simulated water levels in New Mexico, the area of interest, we first changed this boundary to a "no flow" boundary and compared simulated water levels (in New Mexico) under both conditions, "no flow" and "specified flow". The difference in simulated water levels was less than 1 foot in New Mexico, except for few cells along the state line where the difference was approximately 3 feet. We then moved the eastern boundary an additional 34 miles into Texas to determine if boundary location has an effect on simulated water levels in the area of interest. Aquifer parameters for the added 34 miles into Texas were assumed to be the same as those of the 9 miles in the standard simulations. Pumping centers were assumed to extend only an additional 8 miles into Texas with pumping rates the same as those in the standard simulation. The eastern boundary is simulated as a "constant head". Pumping centers are far enough to cause any change in water levels at that boundary. Here again, there is less than 1 foot difference in simulated water levels between this model and the standard model. It was concluded that boundary effect on the area of interest is minimal and the representation of the boundary as a "specified flux" boundary in the standard simulation is acceptable.

VI. MODEL CALIBRATION AND SIMULATIONS

As stated earlier, steady-state simulations were performed to adjust aquifer hydraulic conductivity values, areal recharge, and evapotranspiration rates within their plausible range to obtain a good match between simulated and observed pre-developed water levels. Also adjusted during these simulations are specified flows through model boundaries. A trial and error cycling process was used to estimate these parameters. A cycle starts by assuming a hydraulic conductivity, evapo-transpiration, and areal recharge distribution and a steady-state simulation is performed. Simulated steady-state water levels are compared with observed pre-developed water levels. If the difference between simulated and observed water levels is such that 60% or more of simulated water levels are within 10 feet of the observed and, simulated conditions are acceptable, cycling is terminated and steady-state calibration is considered satisfactory. Otherwise, spatially variable hydraulic conductivity, recharge rates, and evapotranspiration rates are adjusted within their plausible ranges and another cycle starts. The process continues until the difference between observed and simulated water levels is acceptable.

Following steady-state calibration, transient simulations were begun to estimate aquifer specific yield. During transient simulations, specific yield is varied within a plausible range to obtain an acceptable match between the rate of observed and simulated drawdowns for each stress period. Estimated hydraulic conductivity, recharge rates, and evapotranspiration rates are not varied in transient simulations. However, if varying specific yield within its plausible range did not yield an acceptable

match between simulated and observed drawdown rates, then hydraulic conductivity, recharge rates, and evapotranspiration rates are modified and another steady-state calibration is performed followed by transient calibration.

Calibrated model parameters are shown in Figures 8-11. Figure 8 shows areas where evapotranspiration from the aquifer occurs. A maximum evapotranspiration rate of 3.5 inches per year is used in the model. Actual evapotranspiration varies depending on aquifer material, extinction depth, or depth from the ground surface to the water table at which ET ceases, and the maximum evapotranspiration rate. An extinction depth of 25 feet was selected for the model.

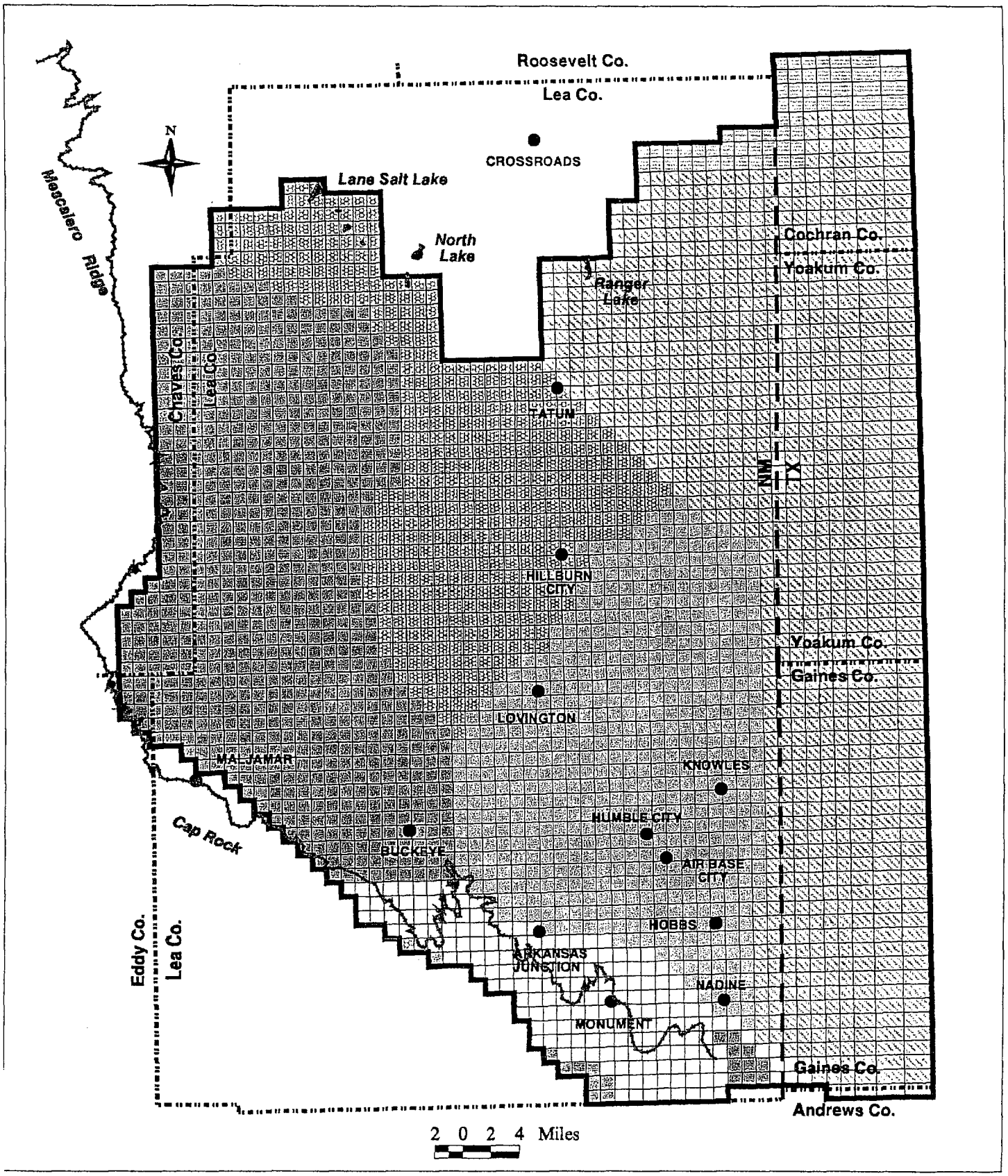
Figure 9 shows the distribution of areal recharge simulated in the model. Similar to evapotranspiration, areal recharge varied from one location to the other and ranges from 0.10 to about 0.50 inches/year. The amount and distribution of recharge in the model depended on annual precipitation and density of the playas.

Figures 10 and 11 shows hydraulic conductivity and specific yield distributions used in the model. Hydraulic conductivity values ranged from 4 to 120 feet/day. Specific yield values ranged from 12 to 24%. These parameters are within the ranges reported by various investigators.

A. Simulated Steady-State water Levels.

Simulated steady-state and observed pre-development (Ash, 1963) water levels are shown in Figure 12. The model reproduced observed pre-development water levels to within 10 feet except in a few locations. The difference between observed and simulated steady-state water levels is illustrated using the frequency distribution shown in Table 4 and Figures 13 and 14. An equal distribution of over and under estimates of water levels is obtained. For selected wells, or wells in which reliable historical observed water levels are available, 95% of simulated water levels are within 20 feet of the observed, whereas when the whole model grid (note that values for the whole model grid were obtained by extrapolating observed water levels in selected wells) is considered, only 83% of simulated water levels are within 20 feet.

The simulated water budget for steady-state calibration is given in Table 5. Areal recharge is the most important component of water inflow to the aquifer. Subsurface recharge through constant head cells and specified flows also contributes to recharge but to a lesser extent. Discharge from the aquifer occurs through subsurface flow to Texas and evapotranspiration. Evapotranspiration constitutes only a small part of the total discharge from the aquifer. The majority of the discharge is through the constant head cells representing outflow to Texas.

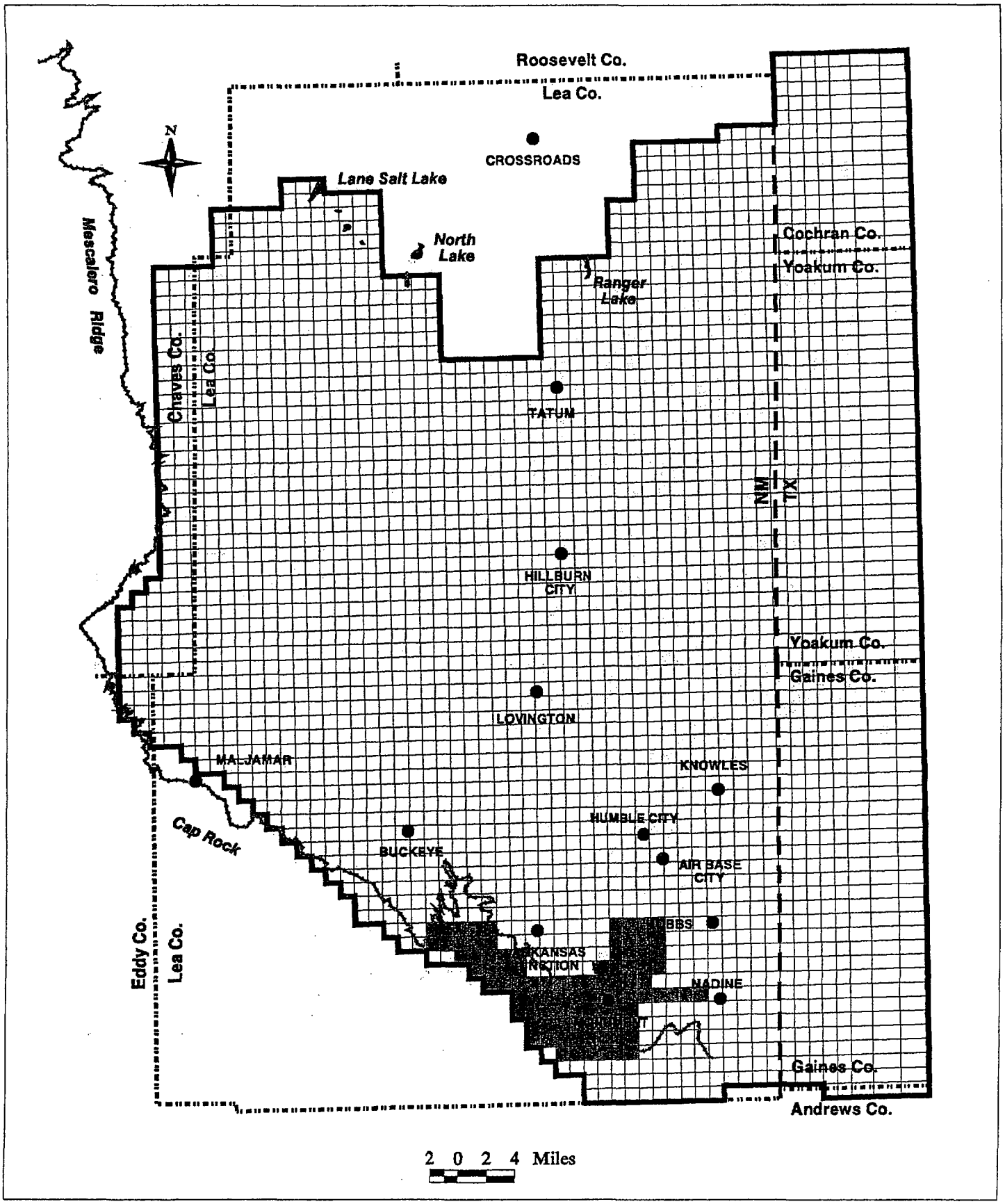


Legend

0.1	0.38
0.29	0.44
0.33	0.49

Model Boundary
State Line
County Line

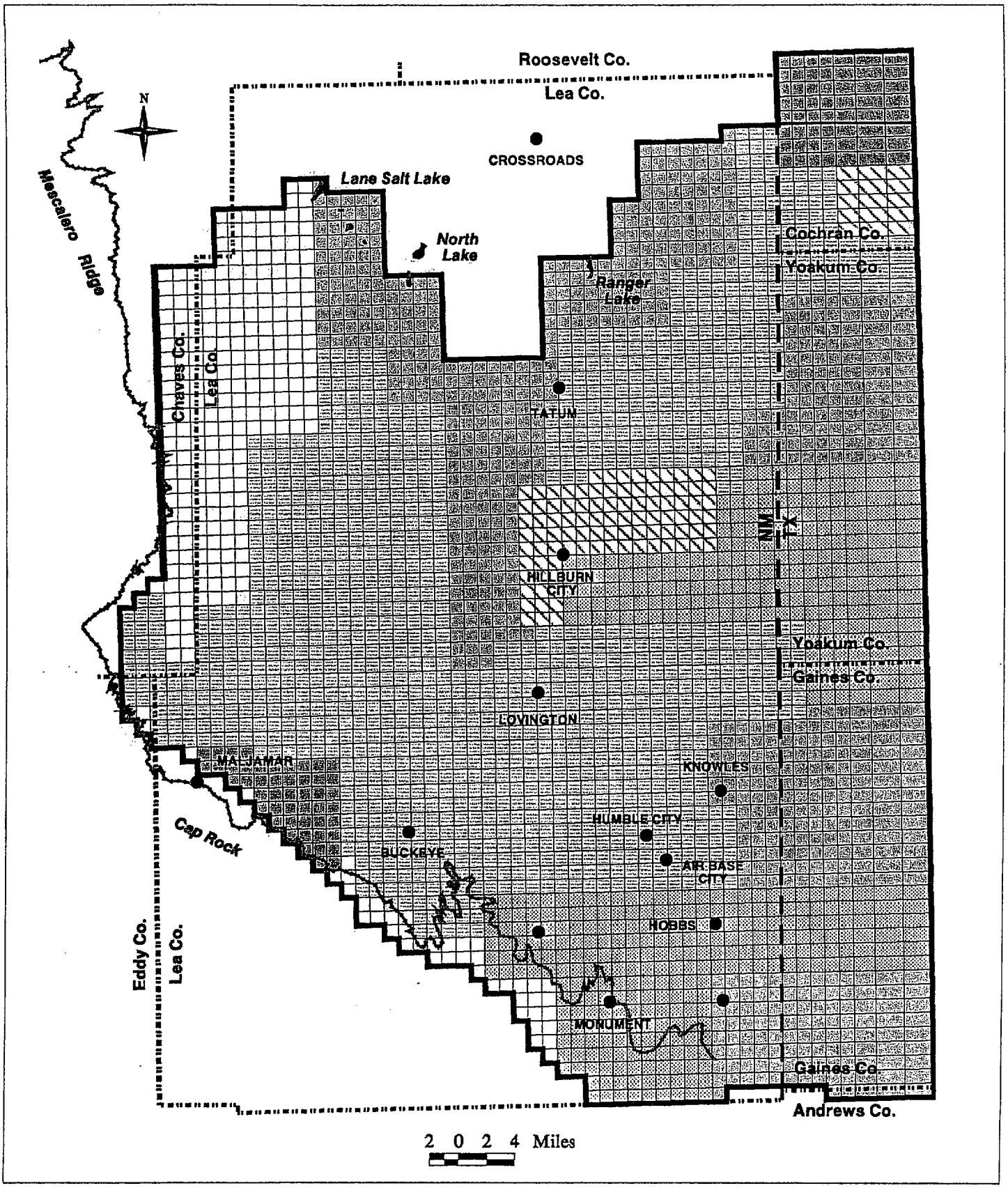
Figure 9 : Areal Recharge Distribution (in/year)



Legend

- ET Simulated Cell
- Model Boundary
- State Line
- County Line

Figure 8: Areas where Evapotranspiration is Simulated in the Model



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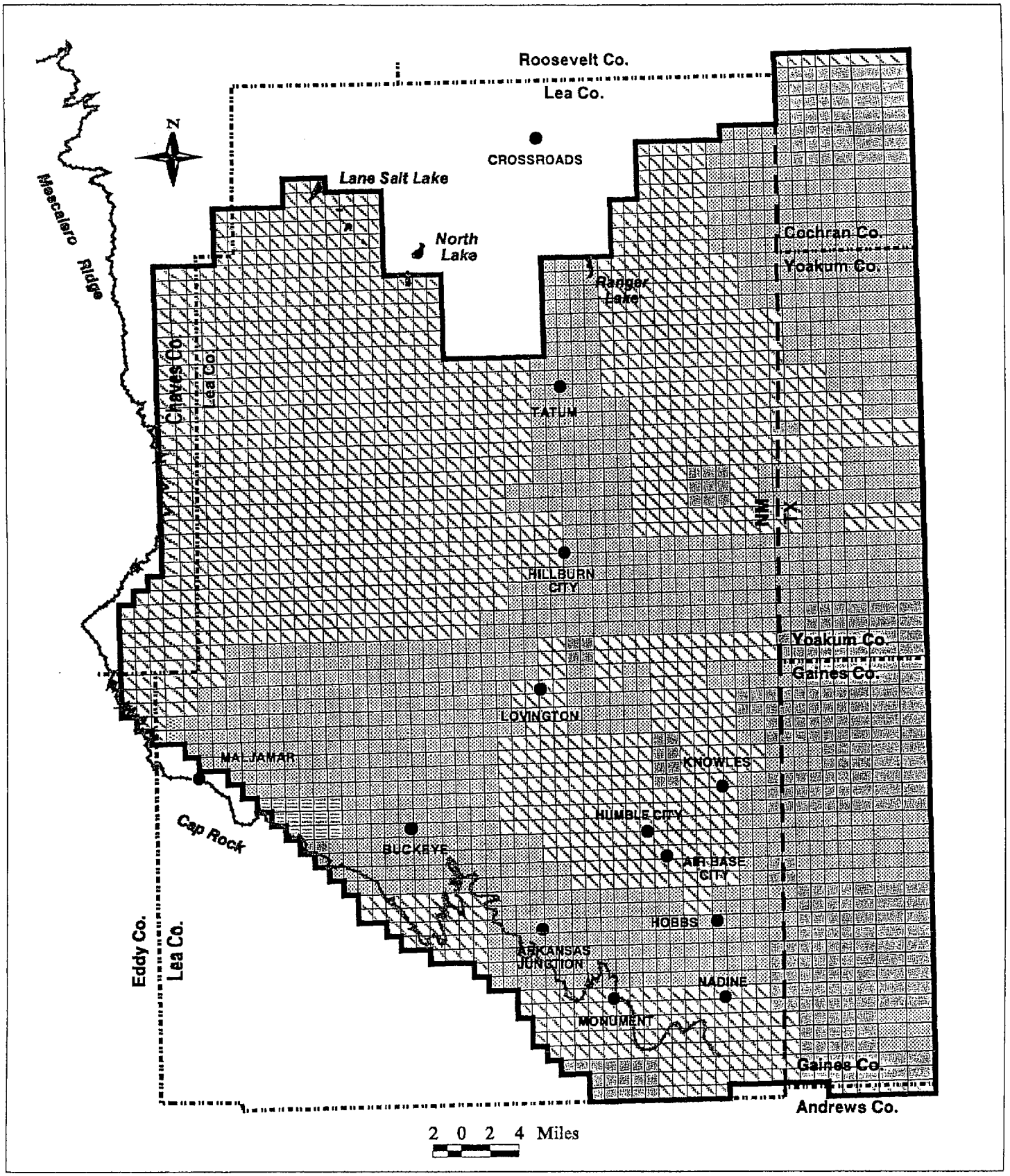
- 1-20
- 21-40
- 41-60
- 61-80
- 81-100
- >100
- Model Line
- State Line
- County Line

Figure 10 : Hydraulic Conductivity Distribution (ft/day)

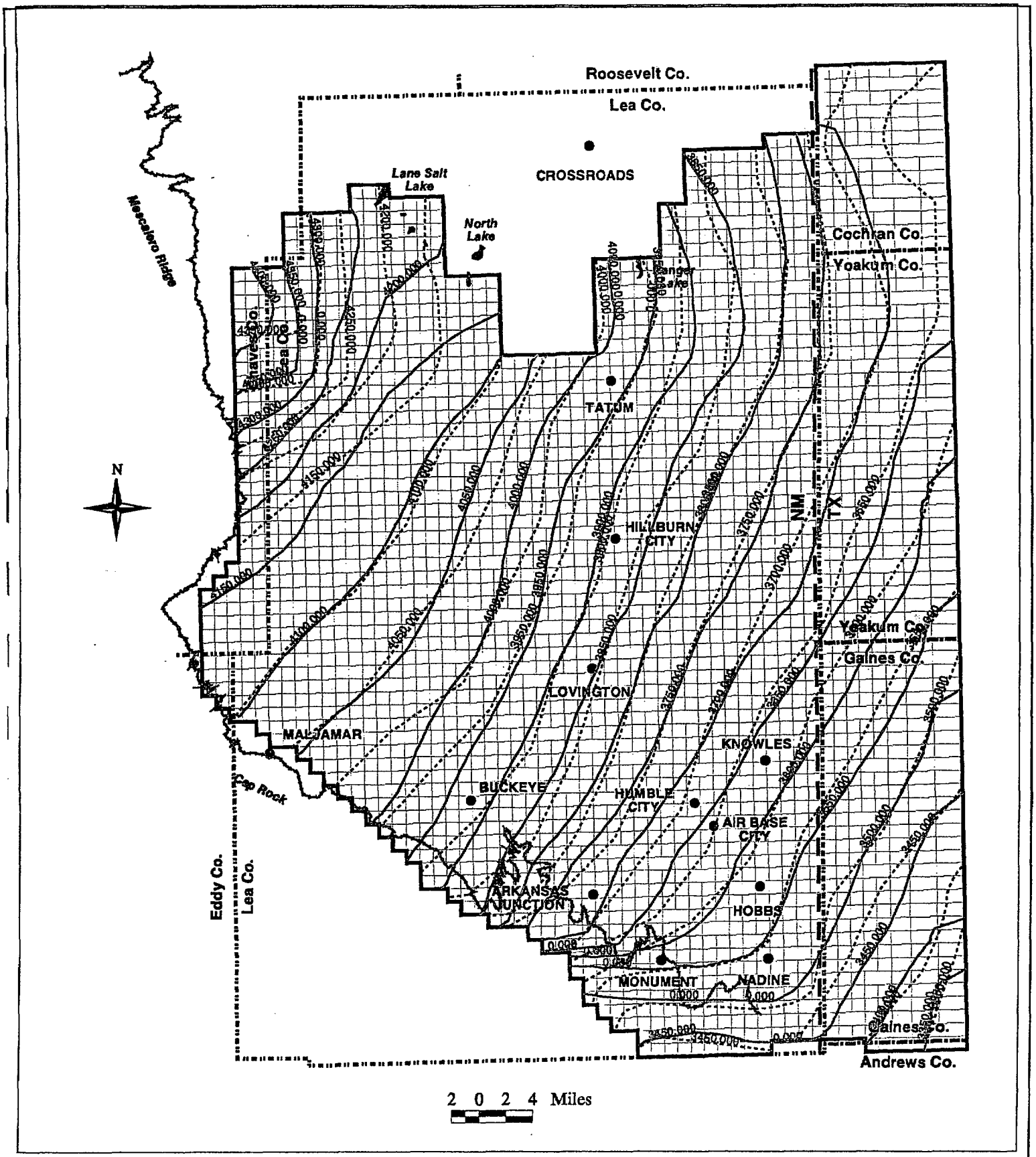


Legend

0.12
0.18
0.2
0.23
0.24

Model Boundary
State Line
County Line

Figure 11 : Specific Yield Distribution



Legend

- Observed Water Level
- Simulated Water Level
- Model Boundary
- State Line
- County Line

Figure 12: Simulated vs. Observed Steady-State Water Levels (ft)

Range of Difference between Computed and measured water Levels (ft)	For All Model Grid			For Measured Water Levels in Selected Wells		
	Under Pred.	Over Pred.	% of Total	Under Pred.	Over Pred.	% of Total
0-10	805	758	55.0	52	45	66.9
10-20	388	409	28.0	32	9	28.3
20-30	156	155	11.0	3	2	3.4
30-40	53	65	4.0	0	1	< 1.0
40-50	5	35	1.0	1	0	<1.0
50-60	1	14	1.0	0	0	
60-70	0	6	0	0	0	
70-80	0	1	0	0	0	
80-90	0	0	0	0	0	
Mean Error		-1.3		2.9		
Mean Absolute Error		11.5		8.7		
Mean Square Error		15.6		11		

Table 4. Difference Between Steady-State Simulated and Observed Water Levels.

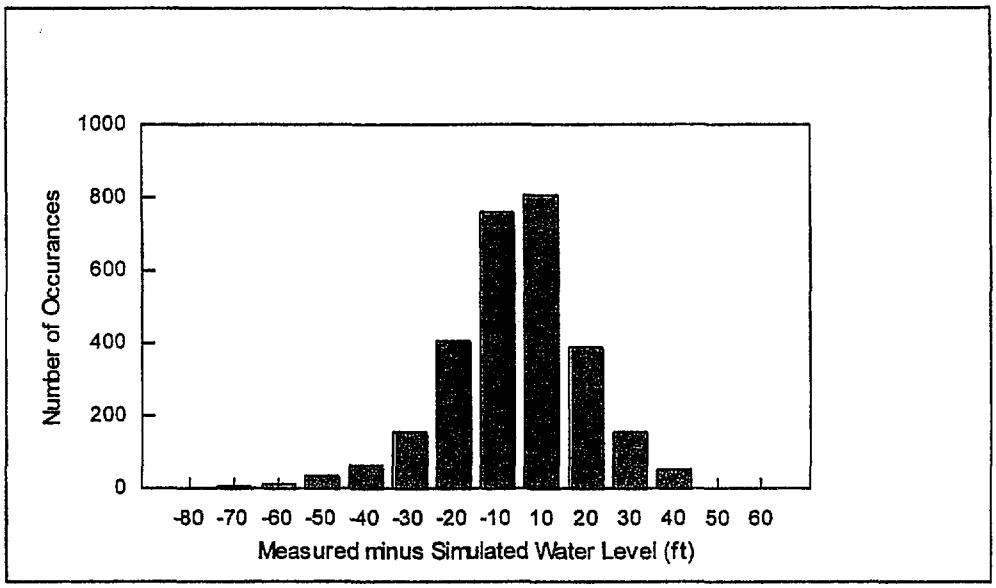


Figure 13. Simulated Steady-State Water Levels minus Measured Water Levels for the Entire Model Grid.

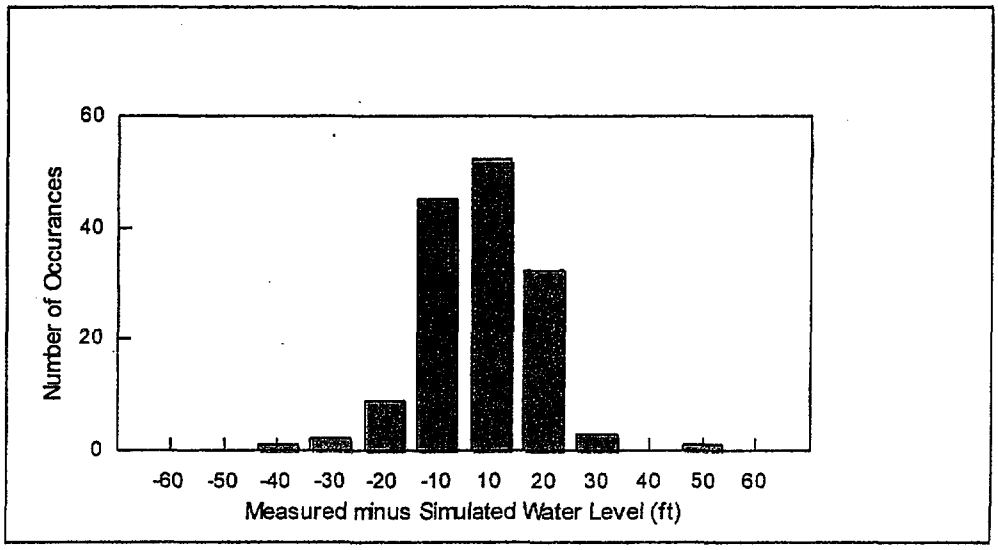


Figure 14. Simulated Steady-State Water Levels minus Measured Water Levels for Selected Wells.

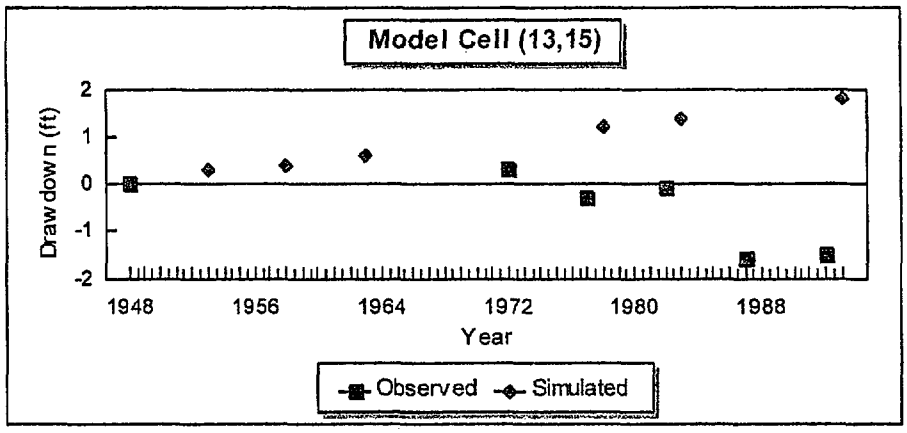
Source	Inflow to the Aquifer	Outflow from the Aquifer
Change in Aquifer Storage	0	0
Flow Through Constant Head Cells	2237	59183
Specified Flow	1320	652
Areal Recharge	57594	0
Evapotranspiration	0	1316
Total	61152	61152

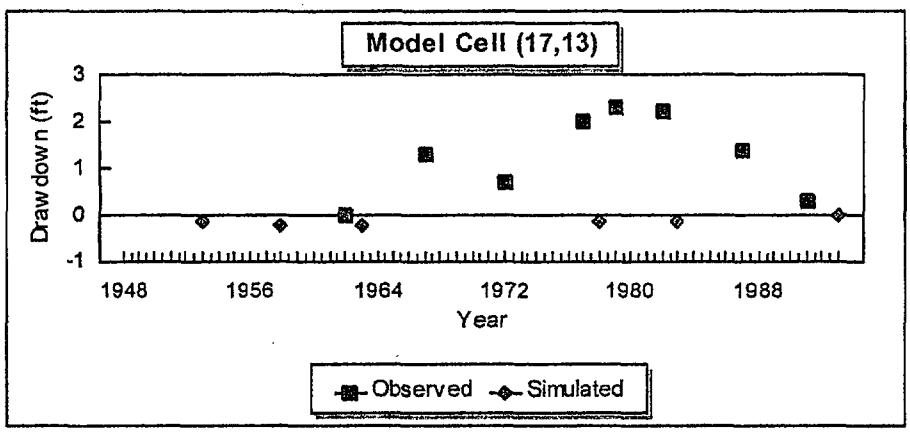
Table 5. Water Budget for the Steady-State Simulation (afy)

B. Simulated Transient Water Levels.

Historical drawdowns simulated by the model are shown in Figures 15 and 16. Figure 15 compares simulated with observed drawdowns for 48 selected wells, while Figure 16 shows total observed and simulated drawdowns at the end of the simulation period. Figure 17 shows the 1996 remaining saturated thickness in the basin. The frequency distribution for the difference between observed and simulated drawdowns at the end of the simulation period for 84 wells is shown in Figure 18 and Table 6. For the transient calibration, 65% percent of the differences between observed and simulated drawdowns are within 10 feet, 95% are within 20 feet, and 5% are greater than 20 feet.

It should be noted that in some wells (cells) shown in Figure 15, observed data show water table recovery during the last two stress periods while simulated results indicate continuous drawdowns in these wells. This deviation is also evident in the skewed error distribution shown in Figure 18 and is likely due to errors in estimated pumping and recharge rates used in the model. Simulated pumping rates are assumed constant for each stress period and continuous for the entire simulation. The recharge rate is also assumed constant for all years and stress periods and was not varied. These assumptions have probably resulted in the over-estimation of aquifer depletion rates and simulated drawdowns in some wells.





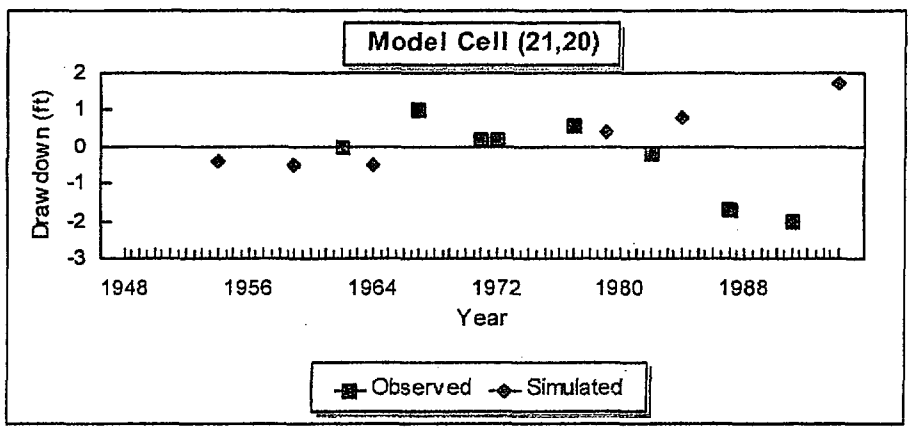
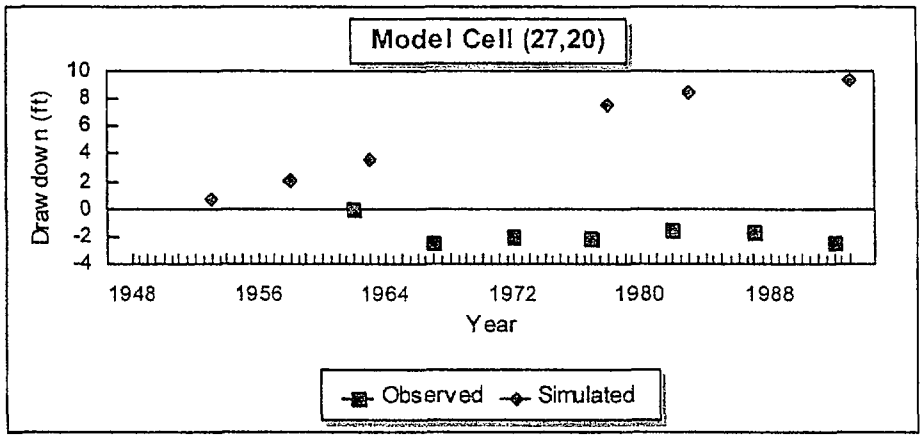
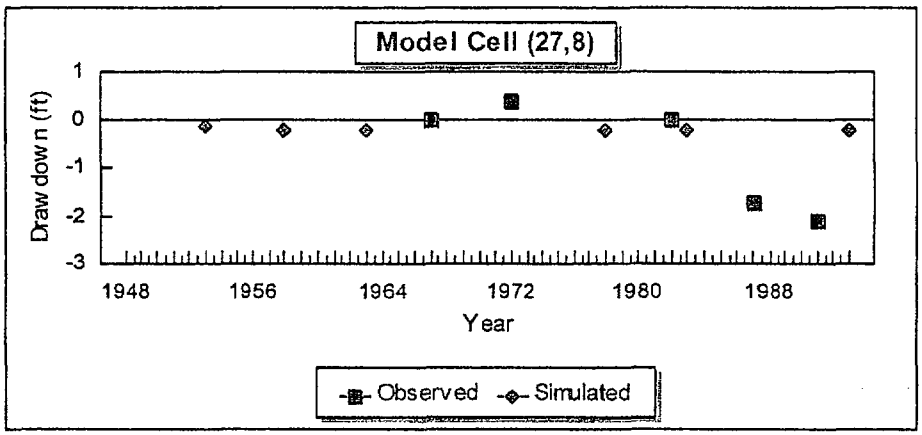
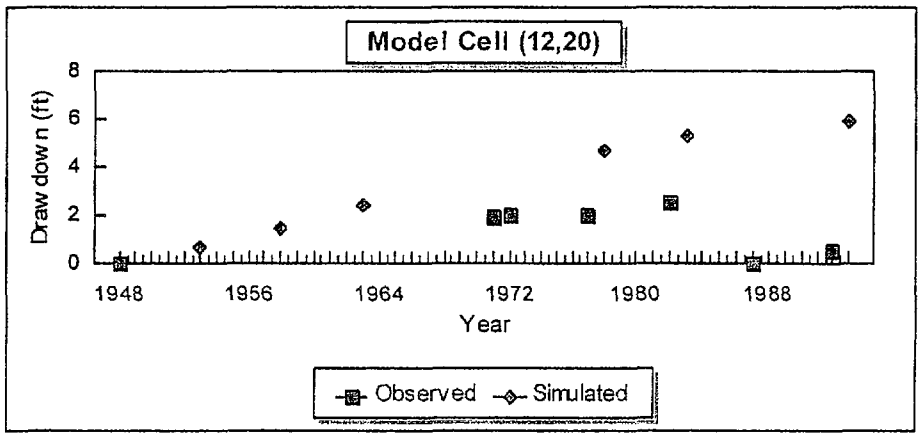


Figure 15. Simulated versus Observed Drawdowns in Selected Wells. (Negative Drawdowns Indicate a Rise in Water Level)







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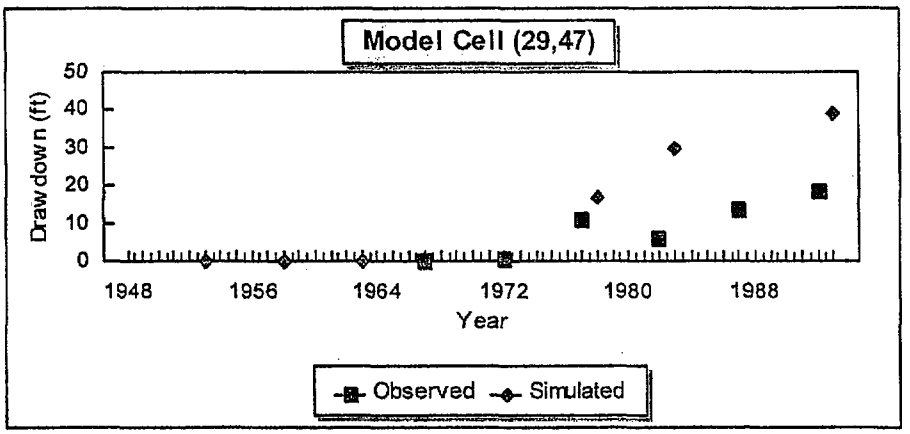

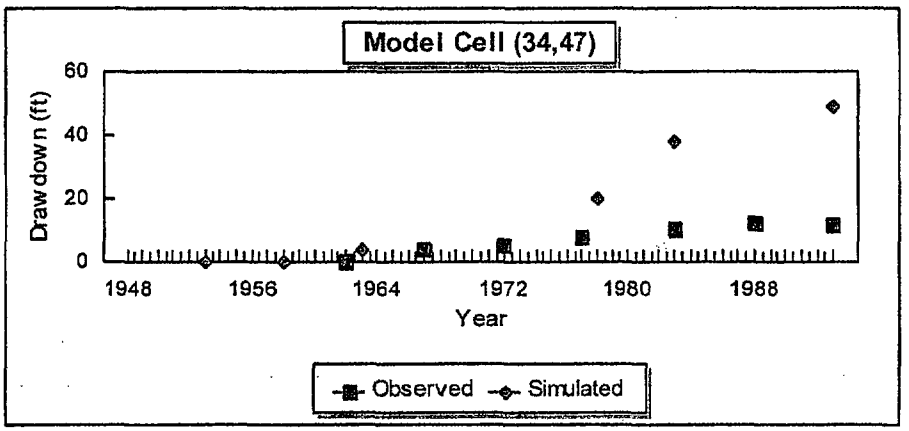

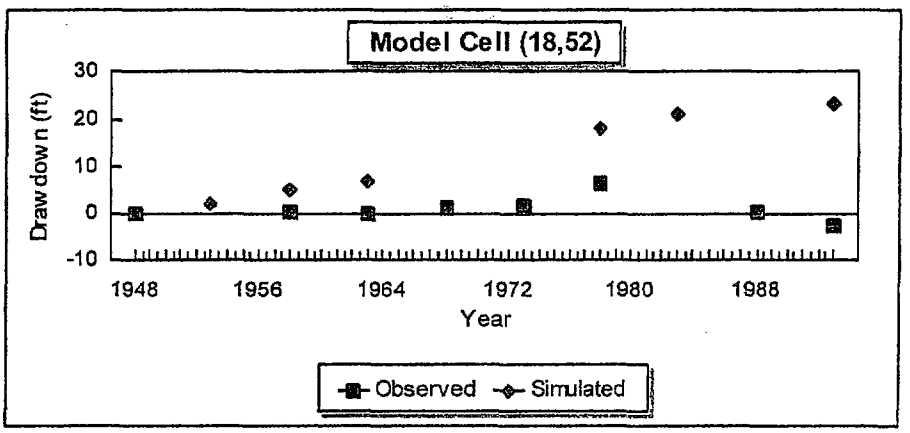

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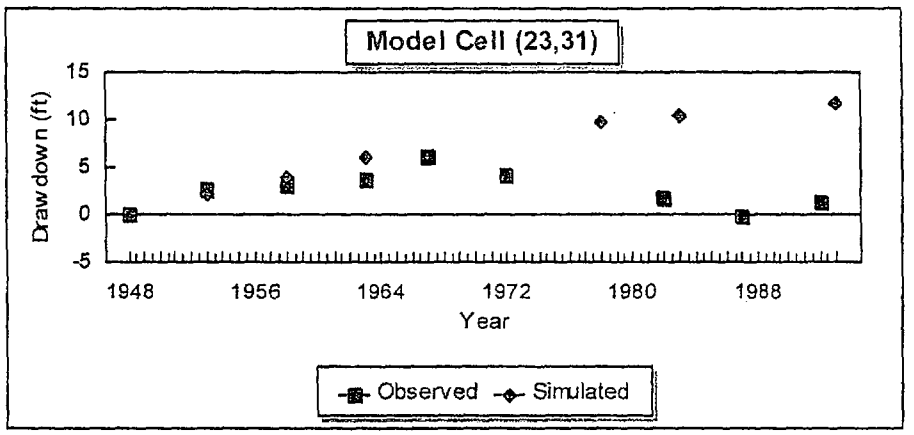

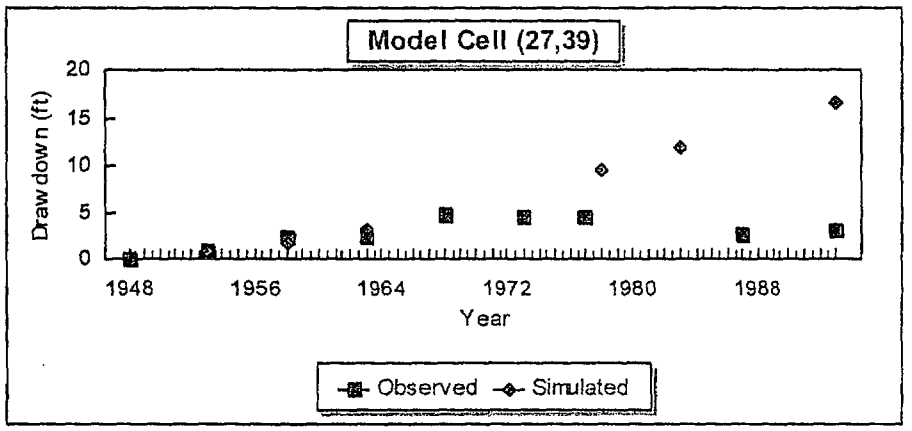

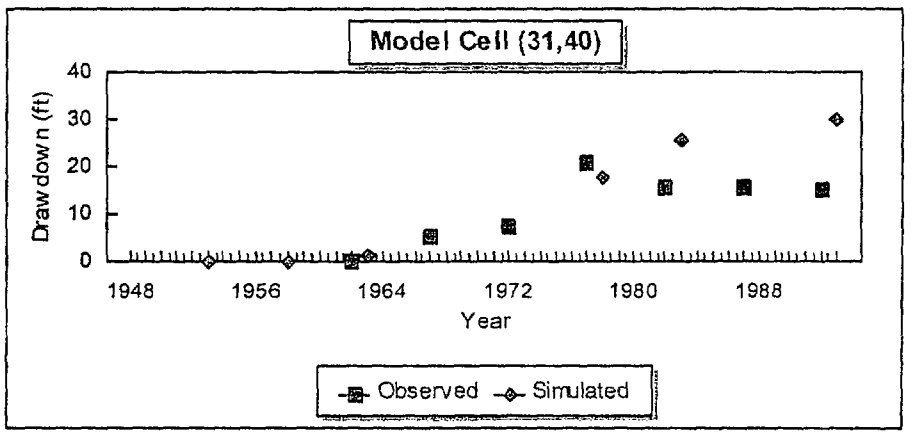

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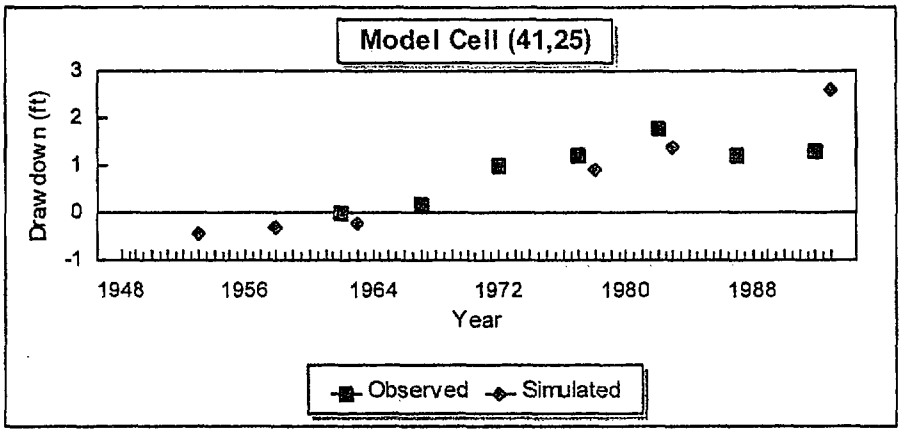

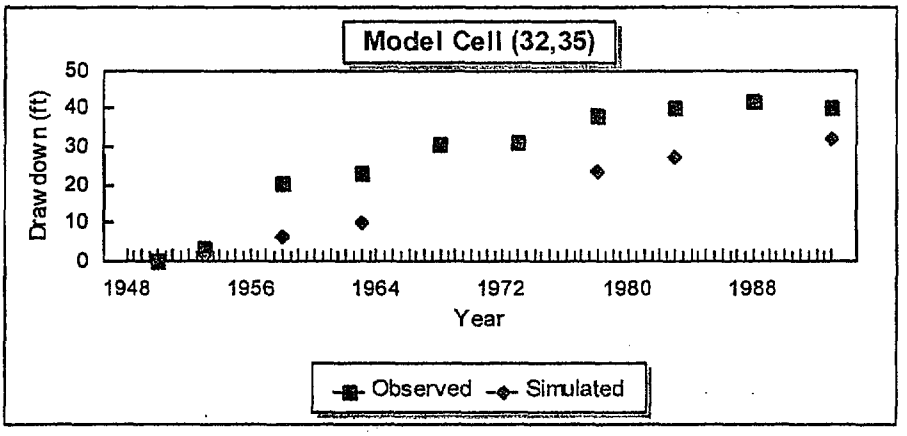

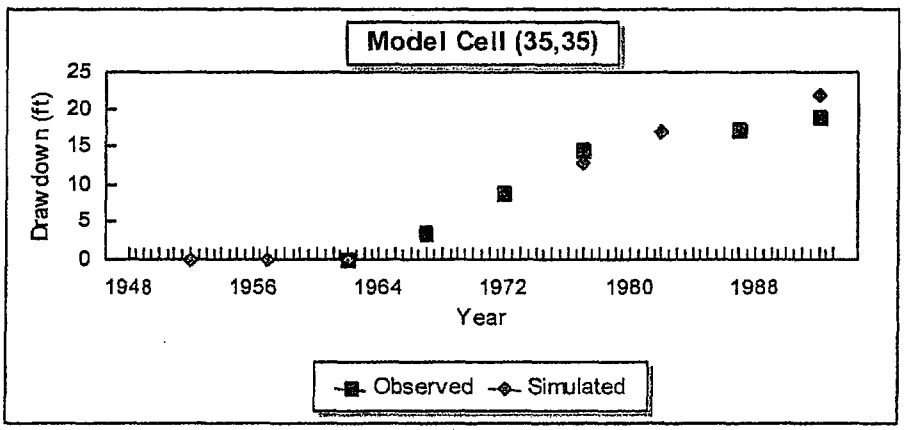

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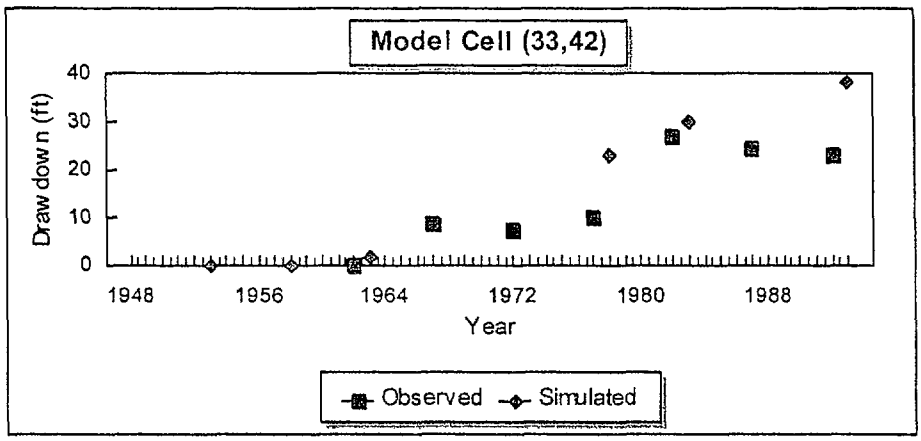

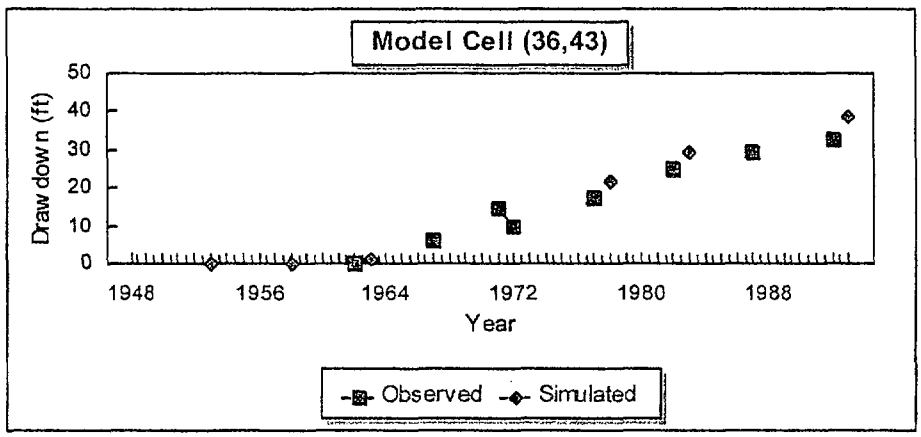

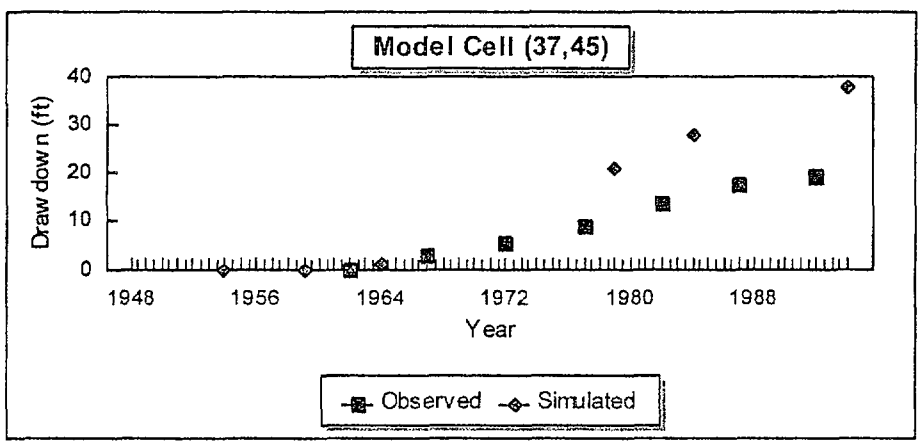

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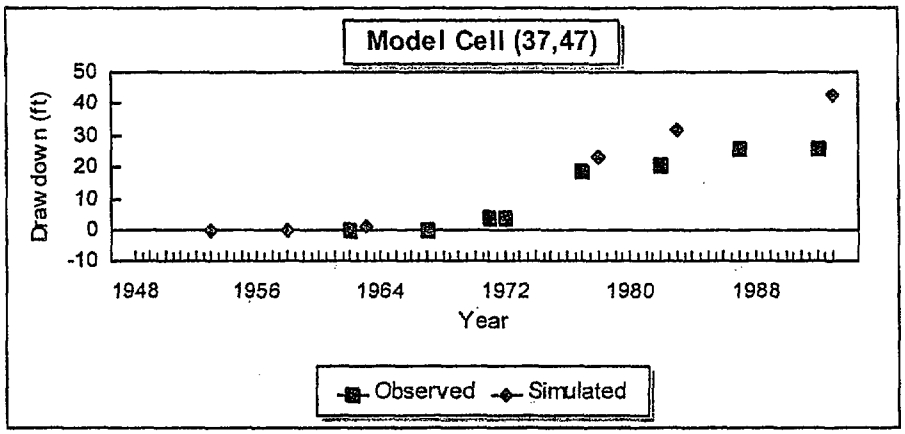

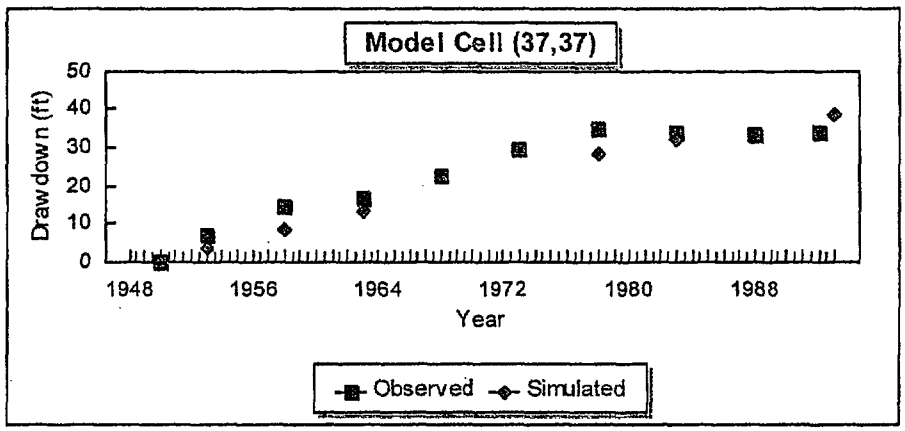

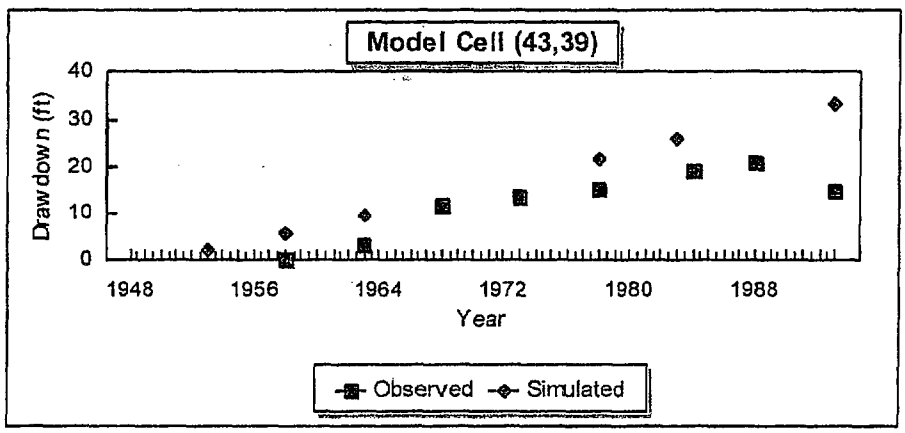

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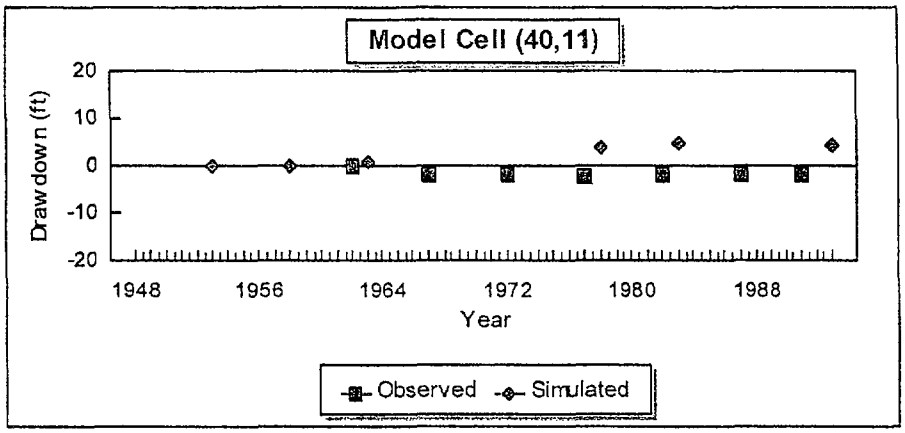

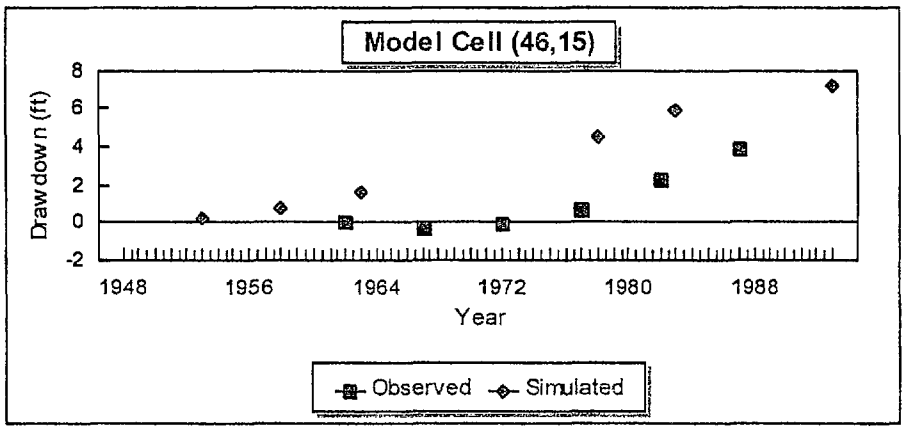

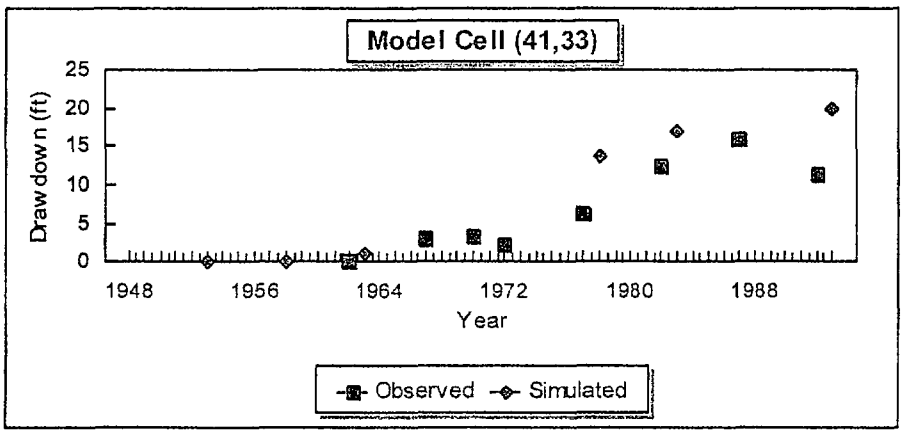

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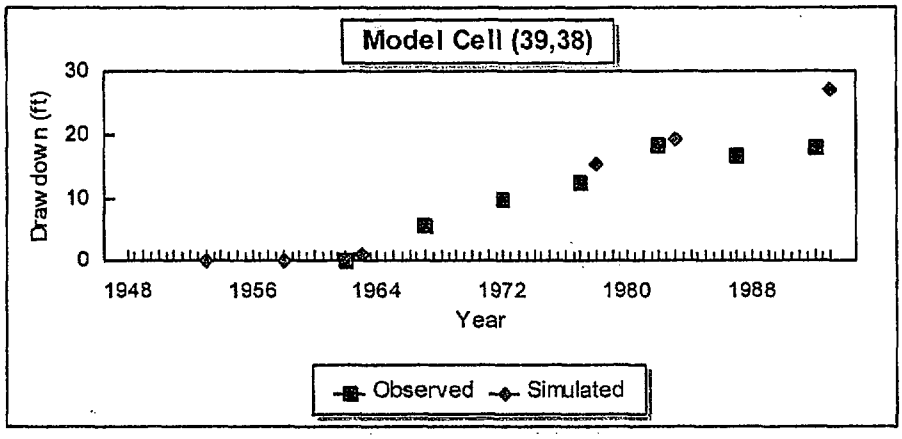

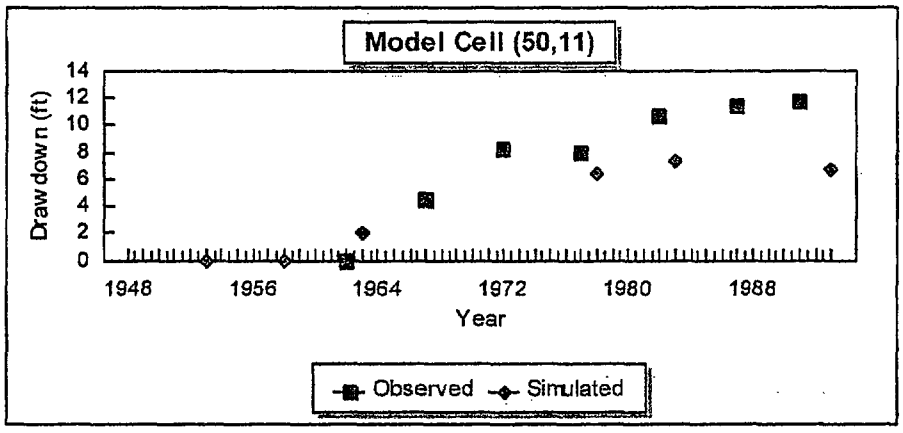

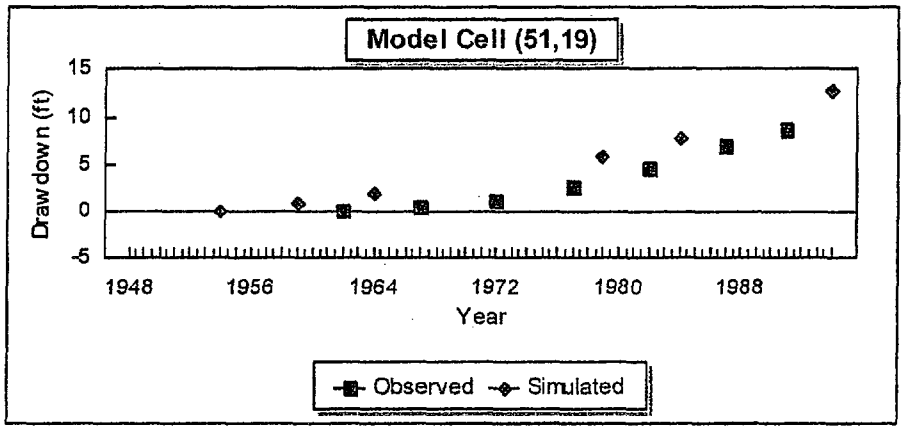

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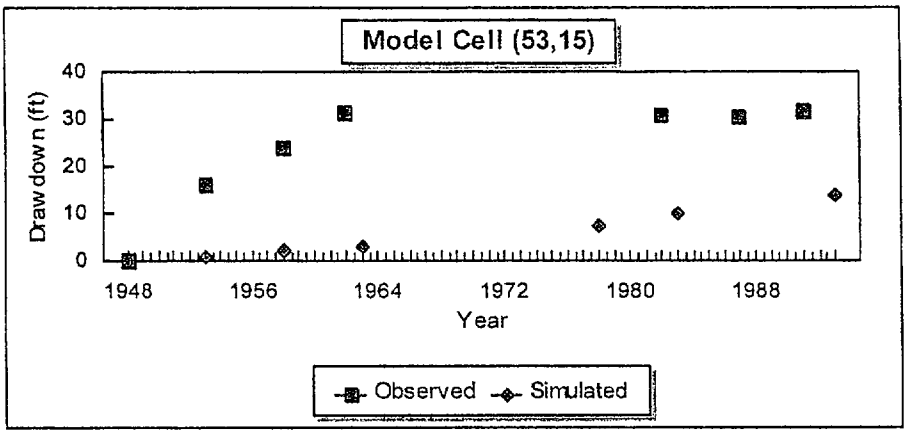

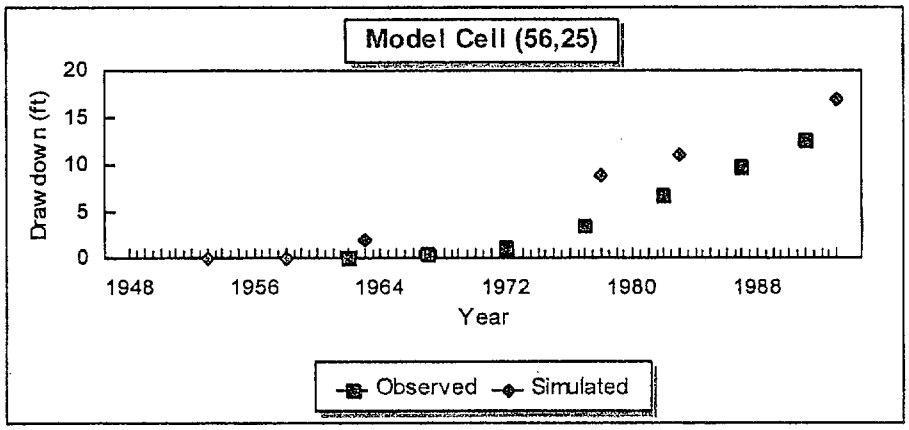

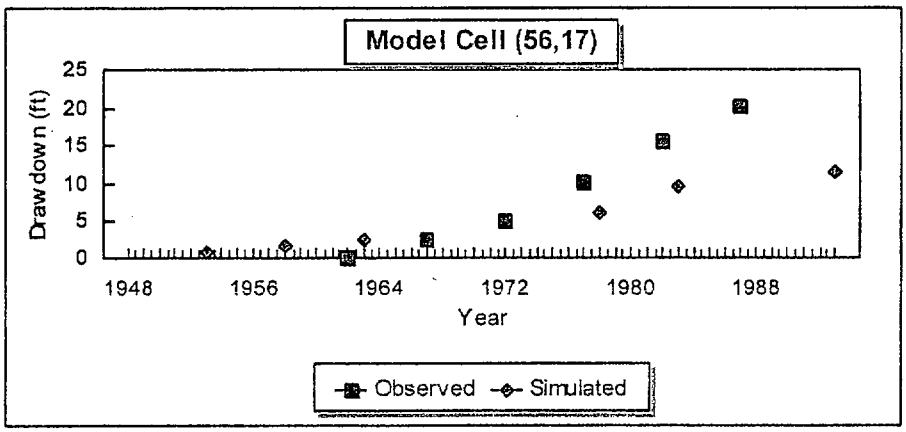

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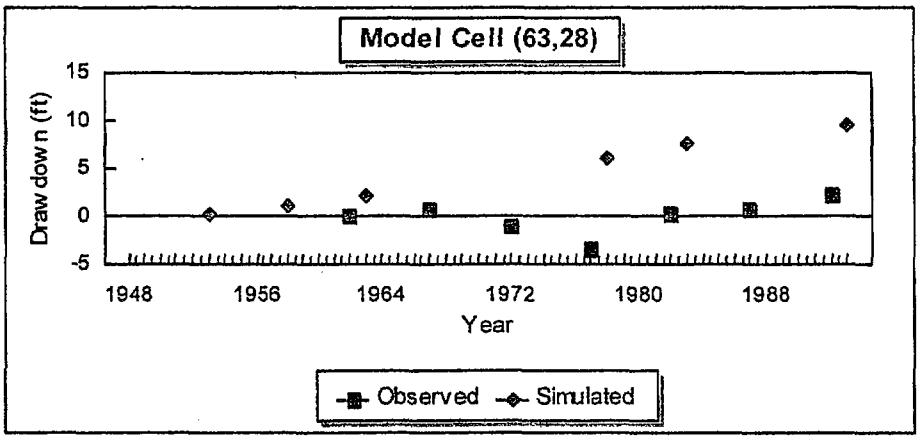

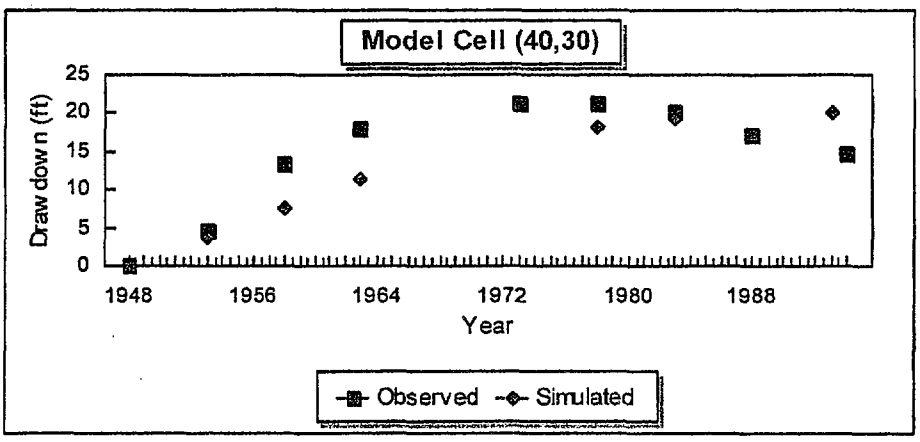

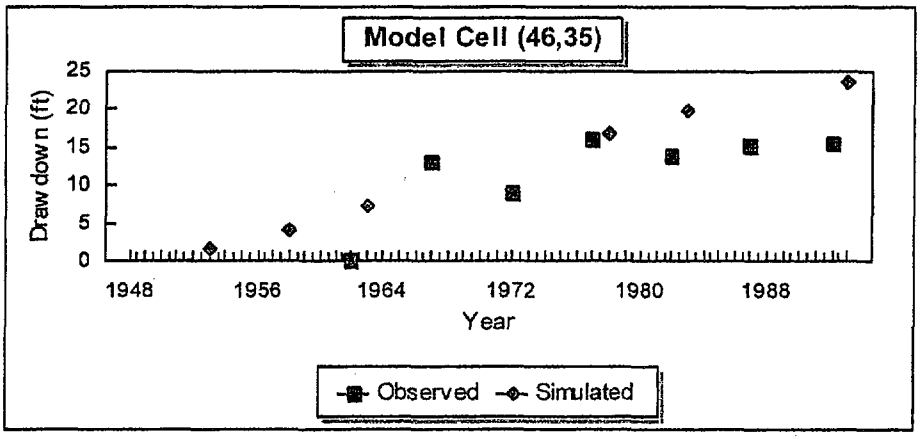

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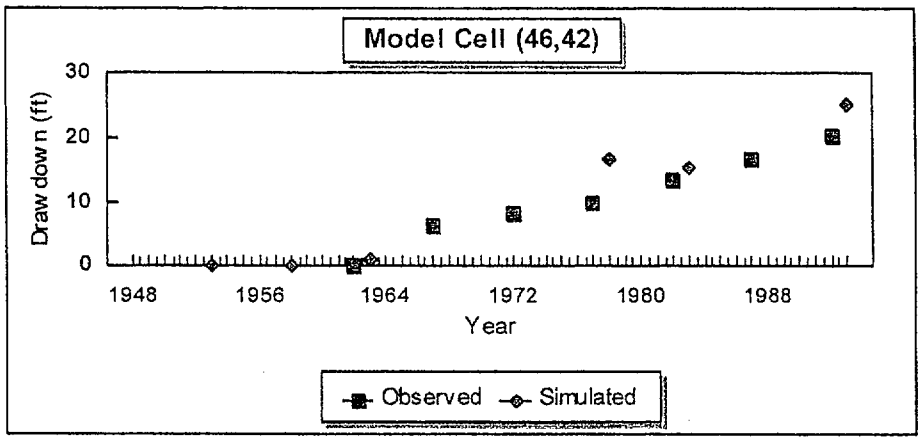

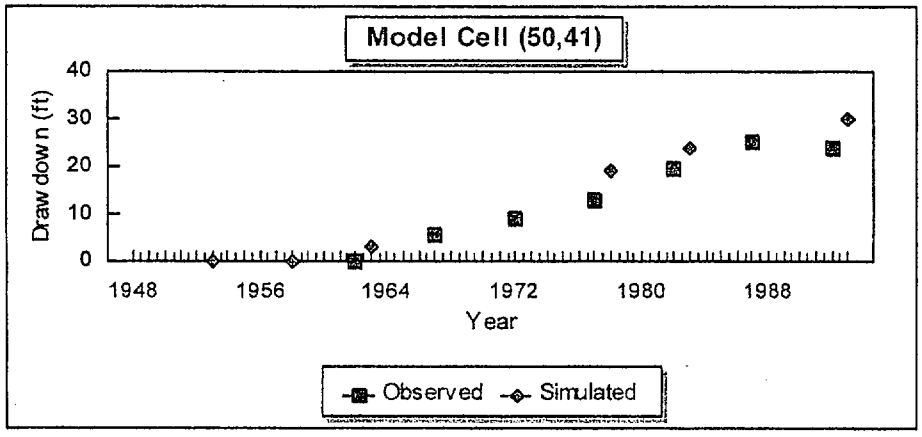

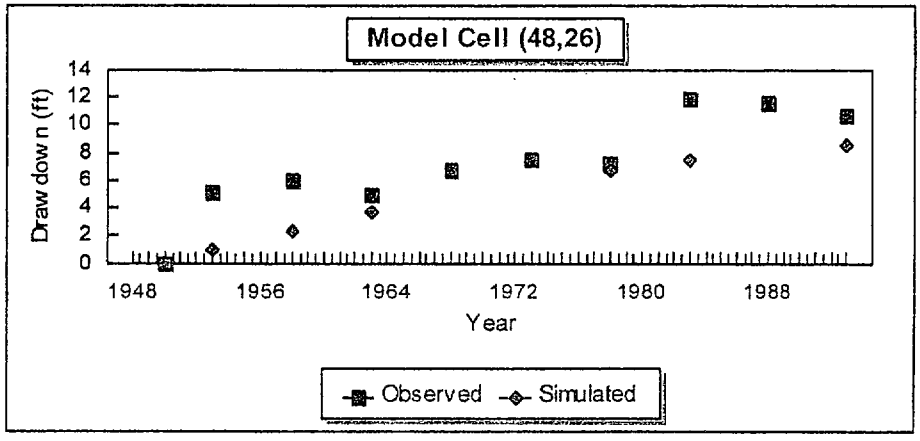

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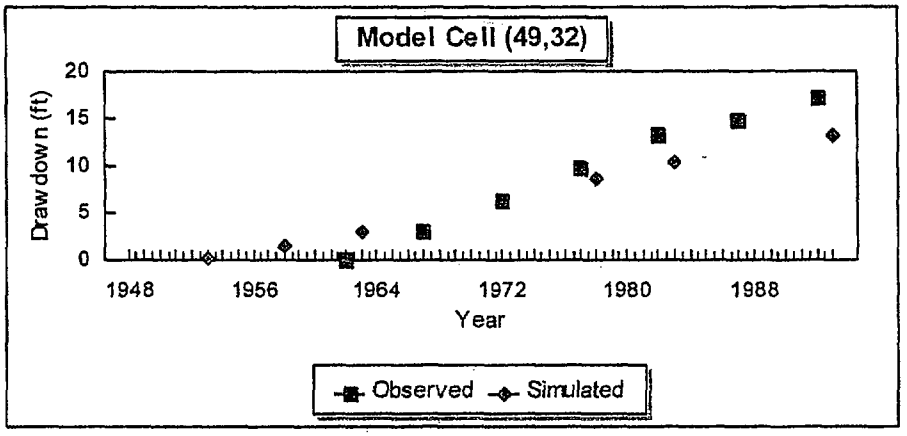

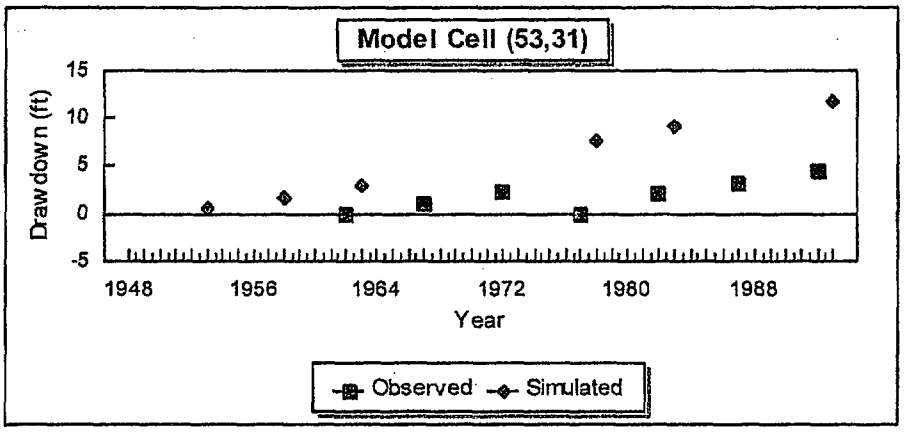

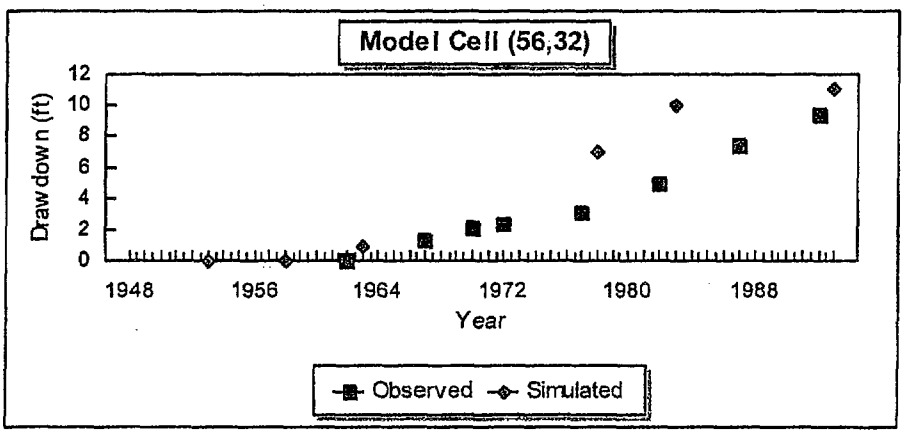

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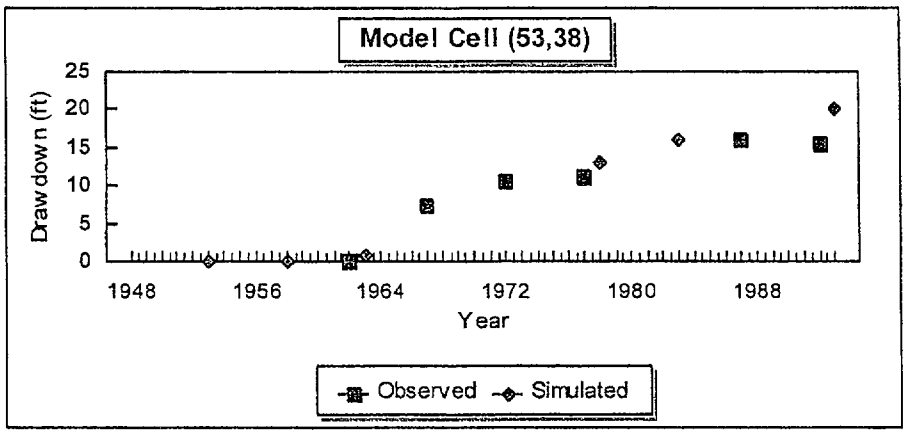

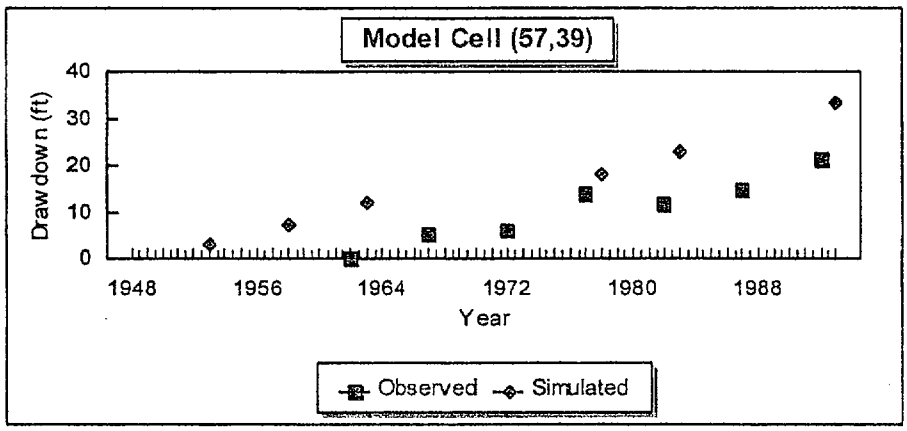

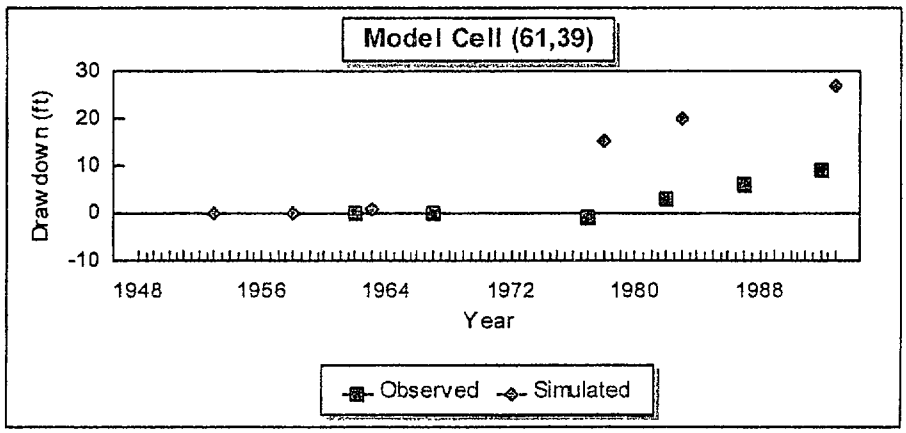

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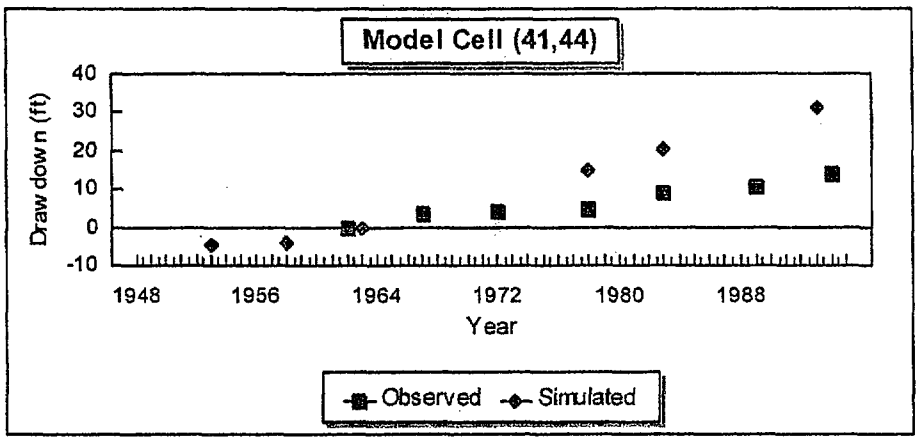

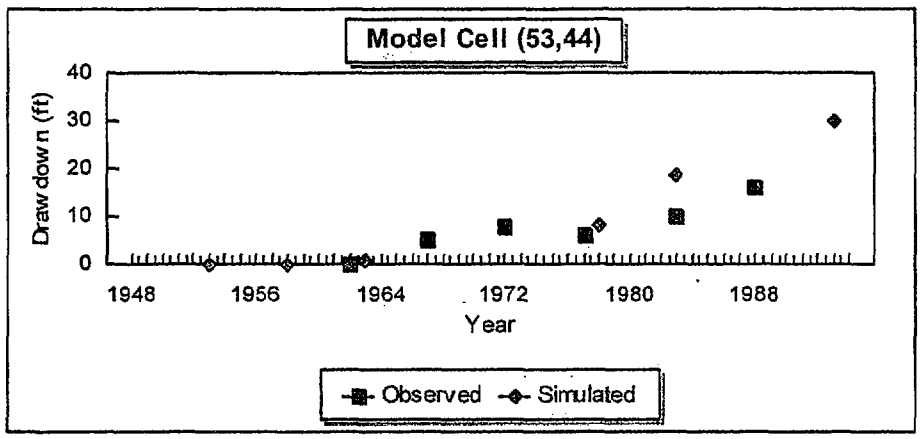

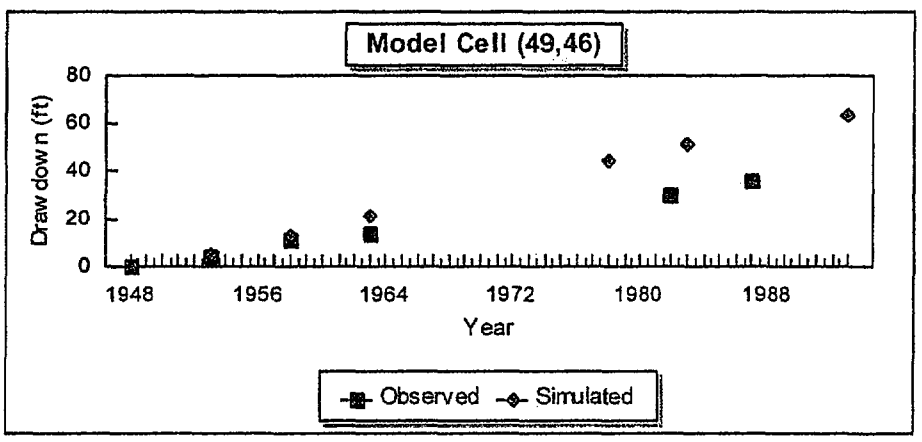

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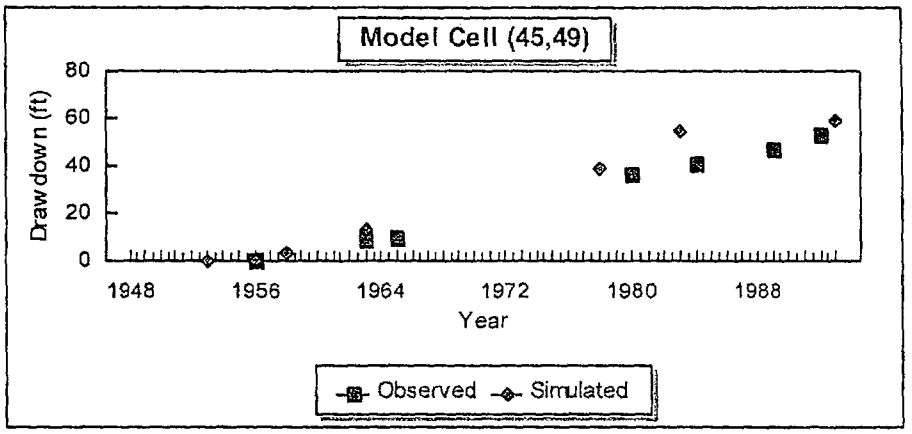

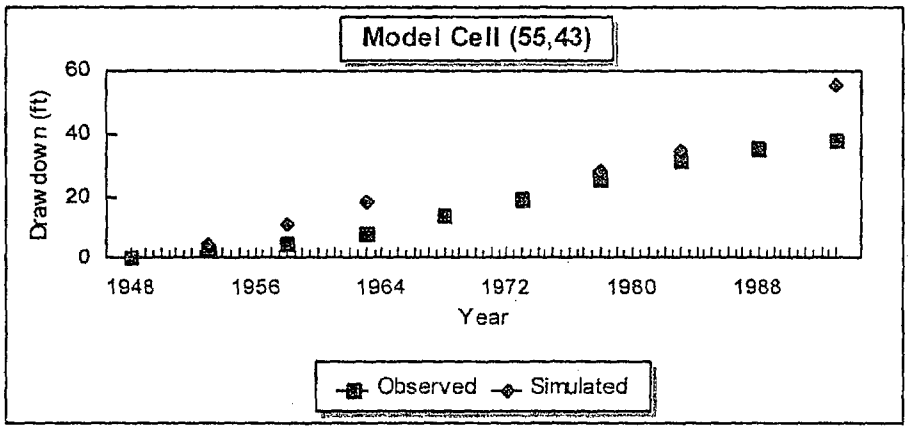

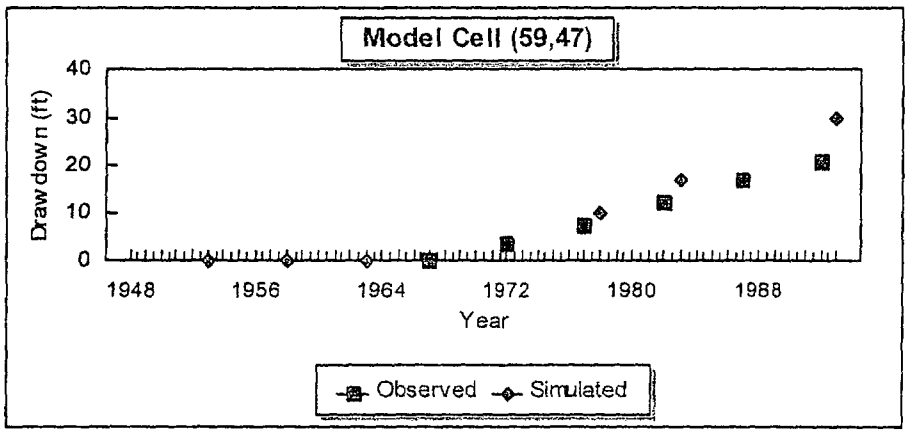

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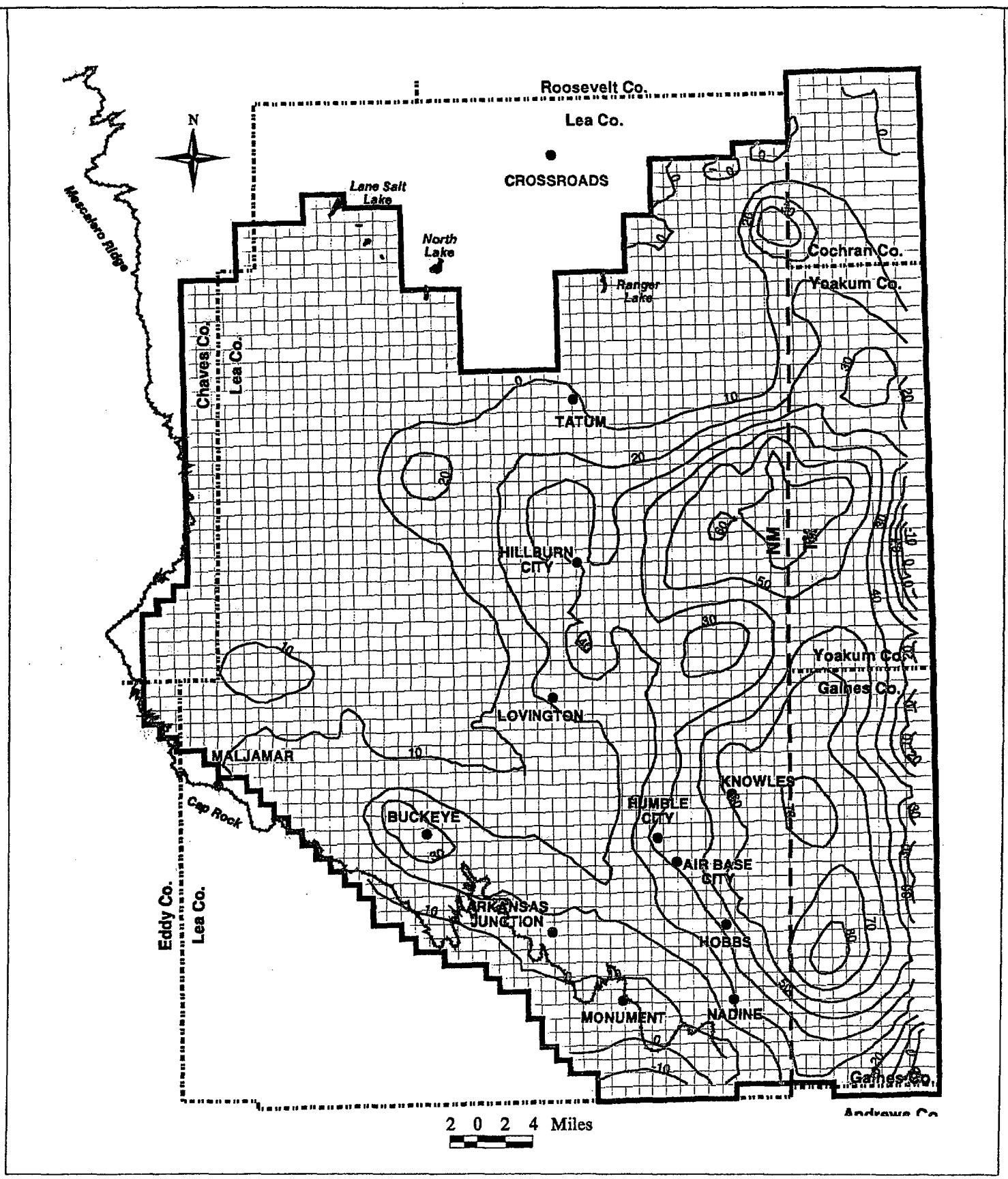
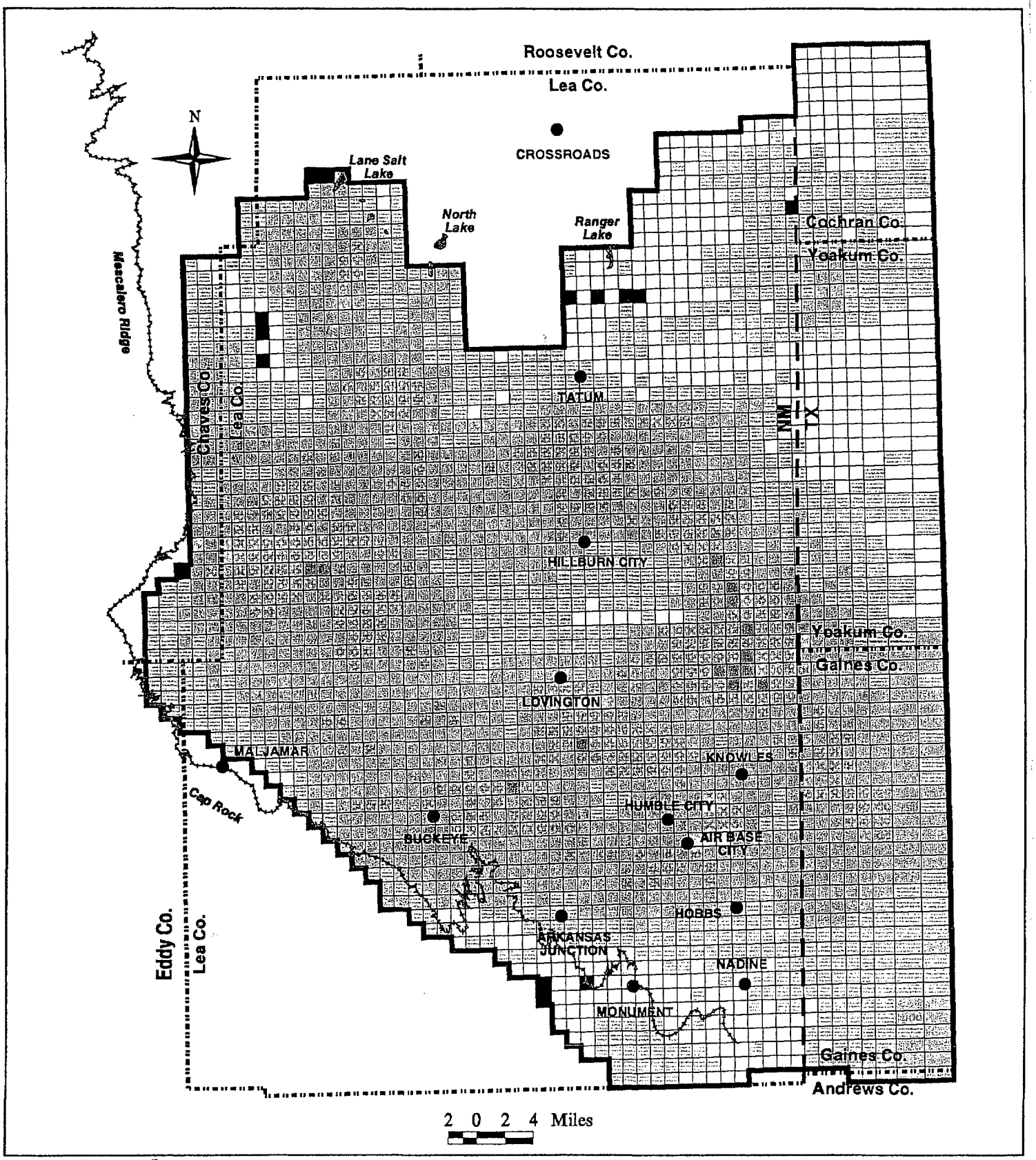


Legend

- Model Boundary
- 1996 Simulated Draw Down
- State Line
- County Line

Figure 16: Drawdown at the End of the Simulated Period (ft) (1948-1996)



Legend

- 0
- 1-49
- 50-99
- 100-149
- 150-199
- 200-249
- Model Boundary
- State Line
- County Line

Figure 17: Saturated Thickness at the End of the Simulated Period (ft)

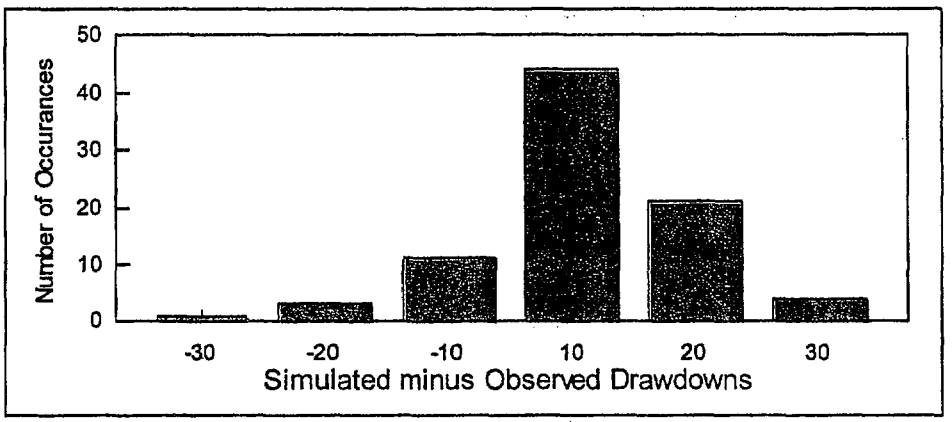


Figure 18. Simulated minus Observed Drawdowns for the Selected Wells.

Range of Difference between Simulated and measured Drawdowns	Under-Predicted	Over-Predicted	% Total
0-10	11	44	65
10-20	3	21	29
20-30	1	4	6

Table 6. Difference Between Simulated and Observed Drawdowns for the Selected Wells.

The largest drawdowns simulated by the model are near the Texas/ New Mexico state line as shown in Figure 16. The model was able to reproduced the general shape of the 1978 observed cones of depressions (Hart and McAda, 1985) in New Mexico. Simulated drawdowns are larger in Texas than in New Mexico due to more pumping on the Texas side. Drawdowns of 70 to 80 feet are simulated in Gains County, Texas while only 50 to 60 feet are simulated in the vicinity of Hobbs, New Mexico. Near Lovington, New Mexico, simulated drawdowns are 30 to 40 feet. Industrial and commercial wells near Buckeye caused 30 feet of drawdown in that area.

The water budget for the transient simulation is shown in Table 7. Most of the aquifer discharge comes from storage. Natural recharge contribution to aquifer discharge is small compared to aquifer storage (the recharge rate in the transient model is slightly higher than in the steady-state model because constant head cells in the steady-state model did not receive areal recharge but when converted to variable

head cells in the transient simulation these cells received areal recharge). Natural discharge through evapotranspiration decreased as water levels declined in the aquifer due to pumping.

Source	Inflow to the Aquifer	Outflow from the Aquifer
Storage[1]	192644	
Pumping wells and Specified Flows	3436	257654
Areal Recharge	62788	0
Evapotranspiration	0	1205
Total	258868	258860

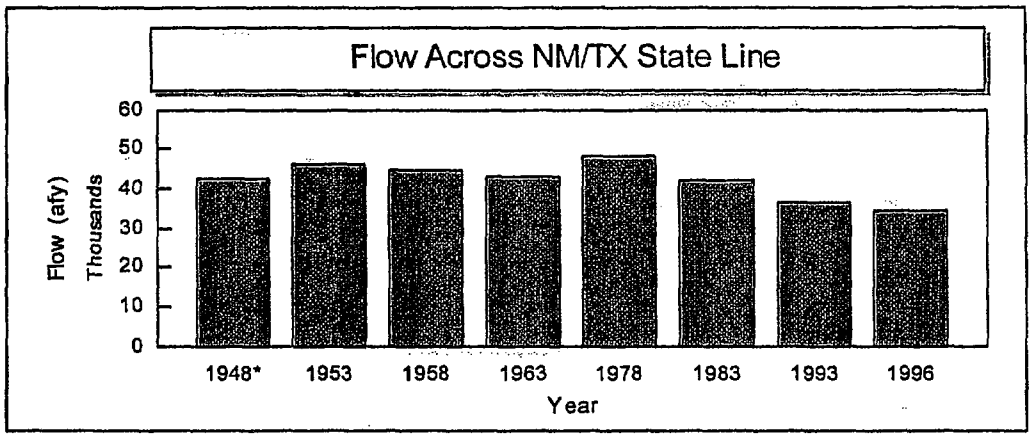
1 Inflow to the aquifer from storage means the net amount of water removed from storage.

Table 7. Water Budget for the Transient Simulation at the End of the Simulation (afy).

C. Flow Across the State Line and Effects of Texas Pumping.

The flow across the state line was computed to determine the effects of pumping in Texas on groundwater flow from New Mexico to Texas (Figure 19). State-line flows were calculated for steady-state conditions (1948) and for each stress period of the transient model. The results show that pumping in Texas caused an initial increase in groundwater flow across the state line. As pumping continued on both sides of the state line, the rate of groundwater flow across the entire Lea County section of the state line decreased to less than the1948 level. However, in the vicinity of Hobbs, the rate of groundwater flow across the state line is higher than under pre-development conditions due to the greater concentration of active wells in Texas.

The effect of pumping in Texas on groundwater levels in Lea County, New Mexico, was examined by performing a simulation with pumping in Texas only. A historical simulation was also performed to determine the effect of pumping in New Mexico on water levels in New Mexico. Figures 20 and 21 shows the 1948-1993 drawdowns in New Mexico resulting from New Mexico and Texas pumping respectively. Texas pumping resulted in 10 feet of drawdown near Hobbs and about 20 feet near the New Mexico/Texas state line. Pumping in New Mexico caused 20 to 50 feet of drawdowns in the area.



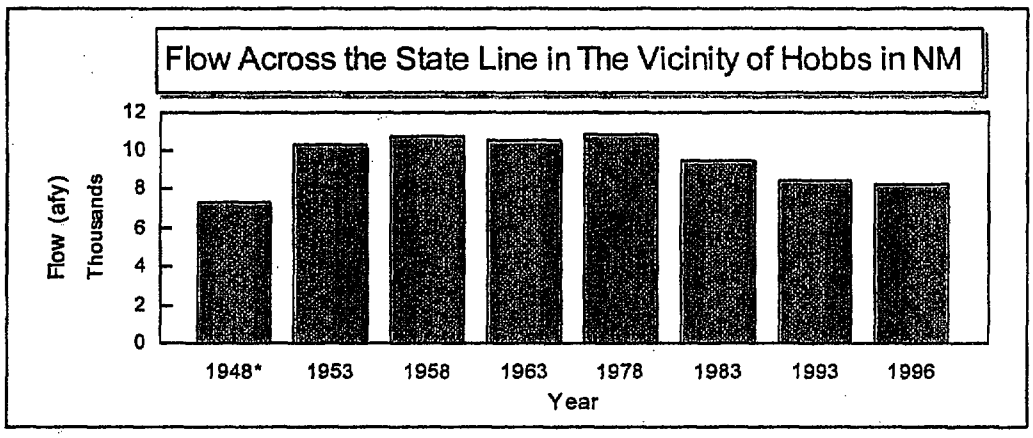
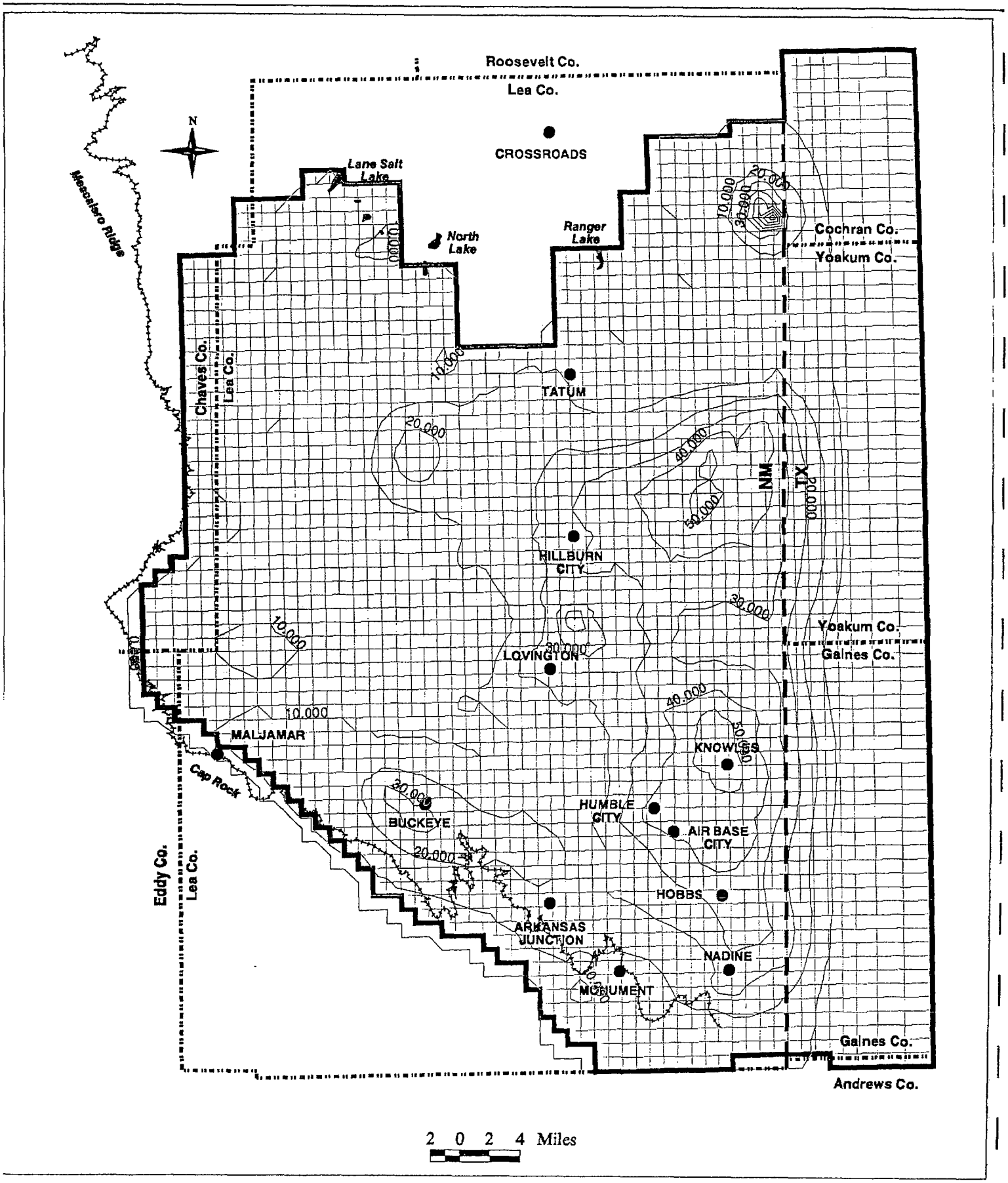


Figure 19. Flow Across the State Line in Lea County for Steady-State and Transient Conditions.
* Indicates Flow Across the State Line Under Steady-State Conditions.



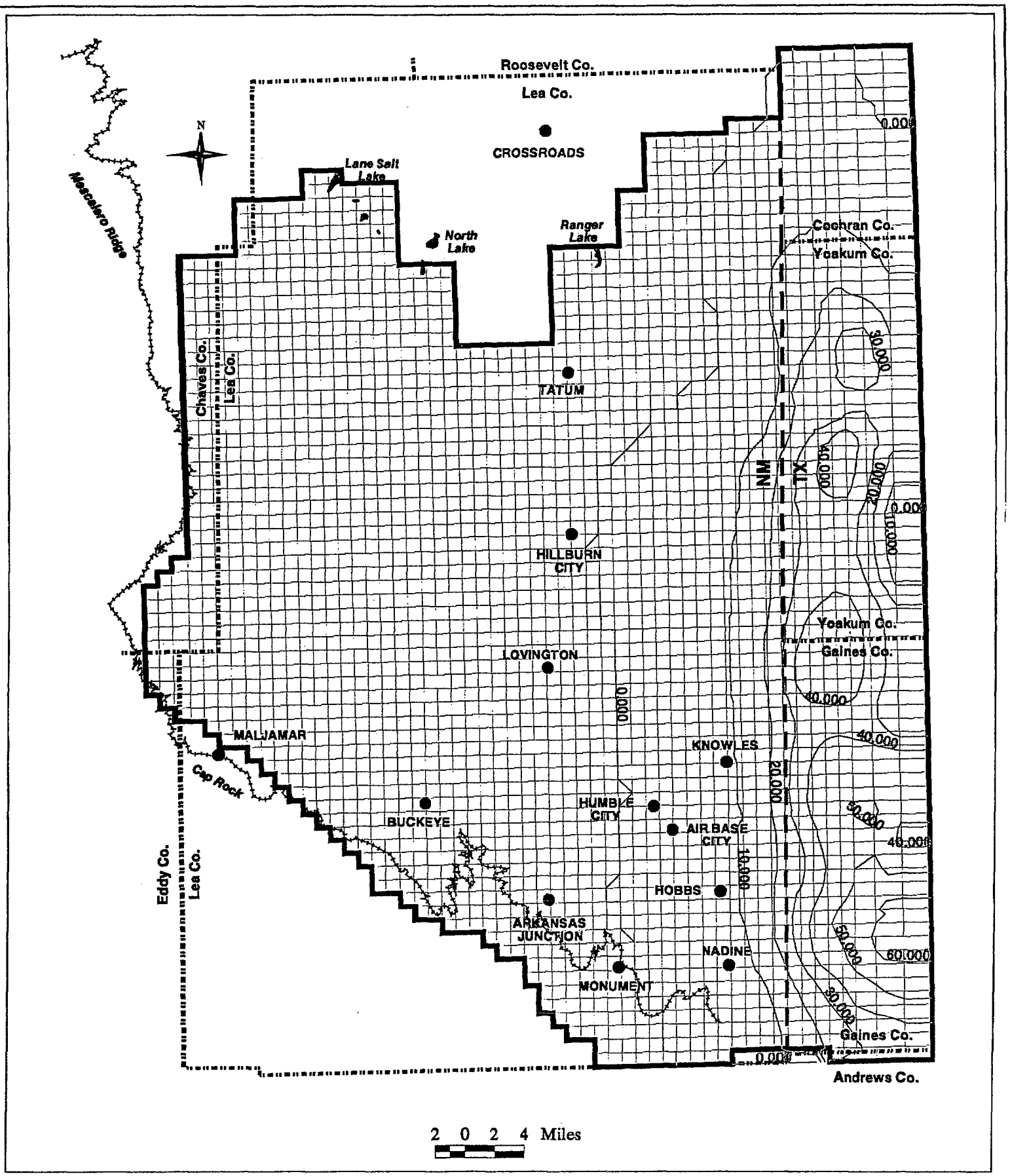
Legend

Model Boundary

State Line

County Line

Figure 20: Simulated Drawdown due to Historical New Mexico Pumping (ft)



Legend
Model Boundary
State Line
County Line

Figure 21: Simulated Drawdown due to Historical Texas Pumping (ft)

V. PROJECTED GROUNDWATER PRODUCTION

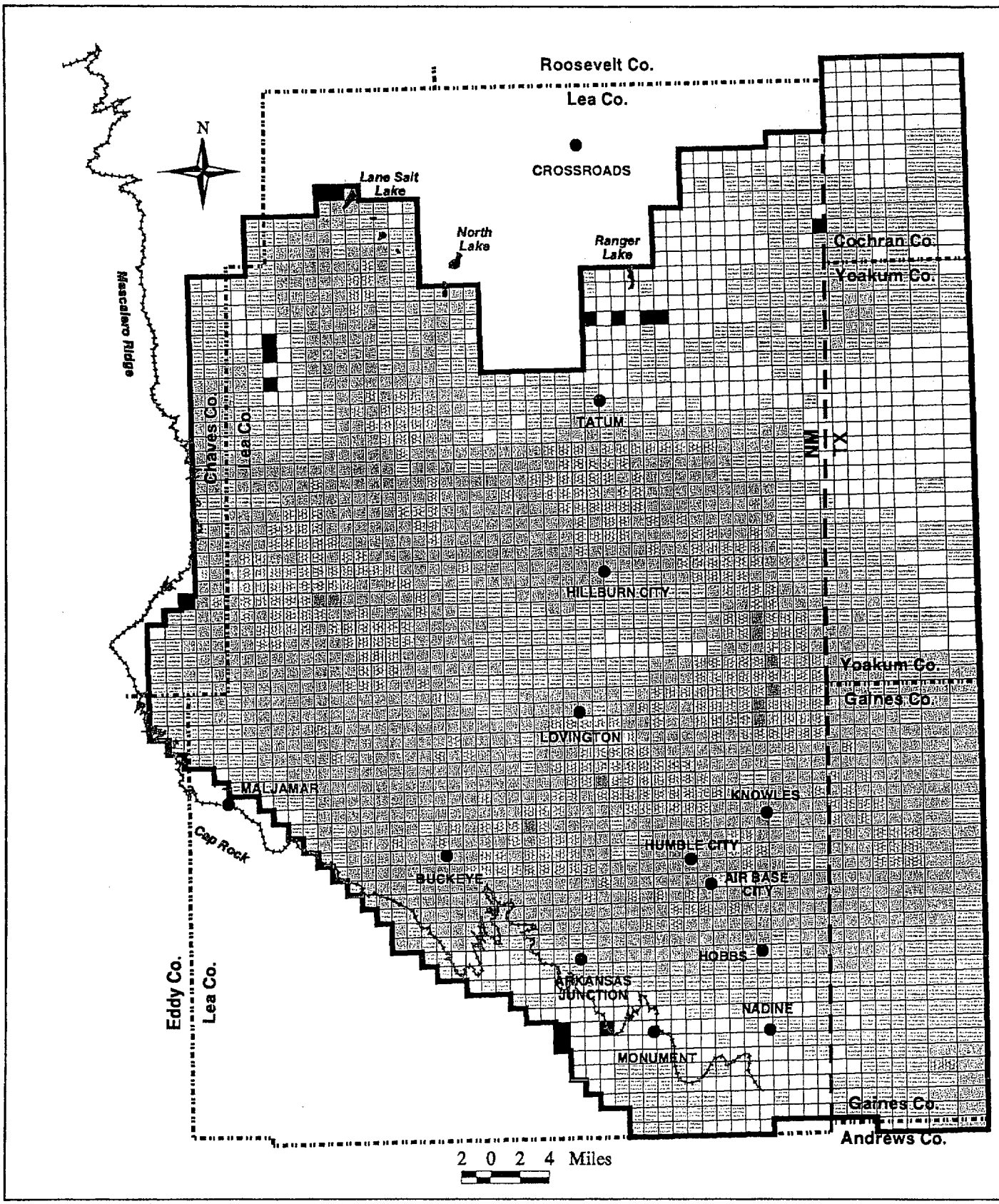
Projected groundwater withdrawals for the year 2040 were simulated using the 1993-1996 estimated pumping rates. Pumping rates from the last stress period were chosen for future projections because they are close to the long term average (1948-1996) of 272000 afy and best represent current conditions. Simulations through the year 2040 indicated that water level declines in the basin will continue, with total de-watering of some areas near the New Mexico/Texas state line. A few areas (cells) in Yoakum County in Texas, where the 1996 saturated thickness is less than 50 feet, could be de-watered as early as 2010. De-watered areas will increase by the year 2040 to include parts of New Mexico near the New Mexico/Texas state line. In other parts of New Mexico, the aquifer will sustain the current pumping rates to the year 2040 but the saturated thickness will decrease by 50 to 100 feet especially near Hobbs, Lovington, and between Lovington and Tatum. The greatest simulated drawdowns in the year 2040 are along the New Mexico/Texas state line due to the combined effect of pumping in New Mexico and Texas. In these areas the remaining saturated thickness decreases to as low as 5 feet. As we move away from the state line into New Mexico less drawdowns are simulated and the average remaining 2040 saturated thickness is more than 200 feet in some areas (Figures 22 to 24).

Table 8 shows water budgets for the years 2000, 2020, and 2040. Aquifer storage contributes about 70% of the extracted water from the basin causing an average water level decline of 0.5 to 2.5 feet/year. Areal recharge and subsurface flow, the only major source of aquifer recharge in the area, are small compared to aquifer discharge through pumping and contribute only about 30% of the extracted water from the basin. Current depletion rates are not indefinitely sustainable and would eventually de-water the basin. It should be noted that these projections are based on the present irrigation of approximately 51,000 acres. According to OSE record there are approximately 120,000 acres with permitted irrigation rights (annual report of basin administration). The annual rate of water level decline could increase if additional permitted acreages are brought back into irrigation.

Year	2000	2020	2040
Inflow to the Aquifer			
Storage[1]	191120	187162	172052
Specified Flows	3436	3436	3436
Areal Recharge	62788	62297	61467
Evapotranspiration	0	0	0
Total	257344	252895	236955
Outflow From the Aquifer			
Pumping wells and Specified Flows	257654	252246	237169
Areal Recharge	0	0	0
Evapotranspiration	1155	789	534
Total	258809	253035	237703

1 Inflow to the aquifer from storage means the net amount of water removed from storage.

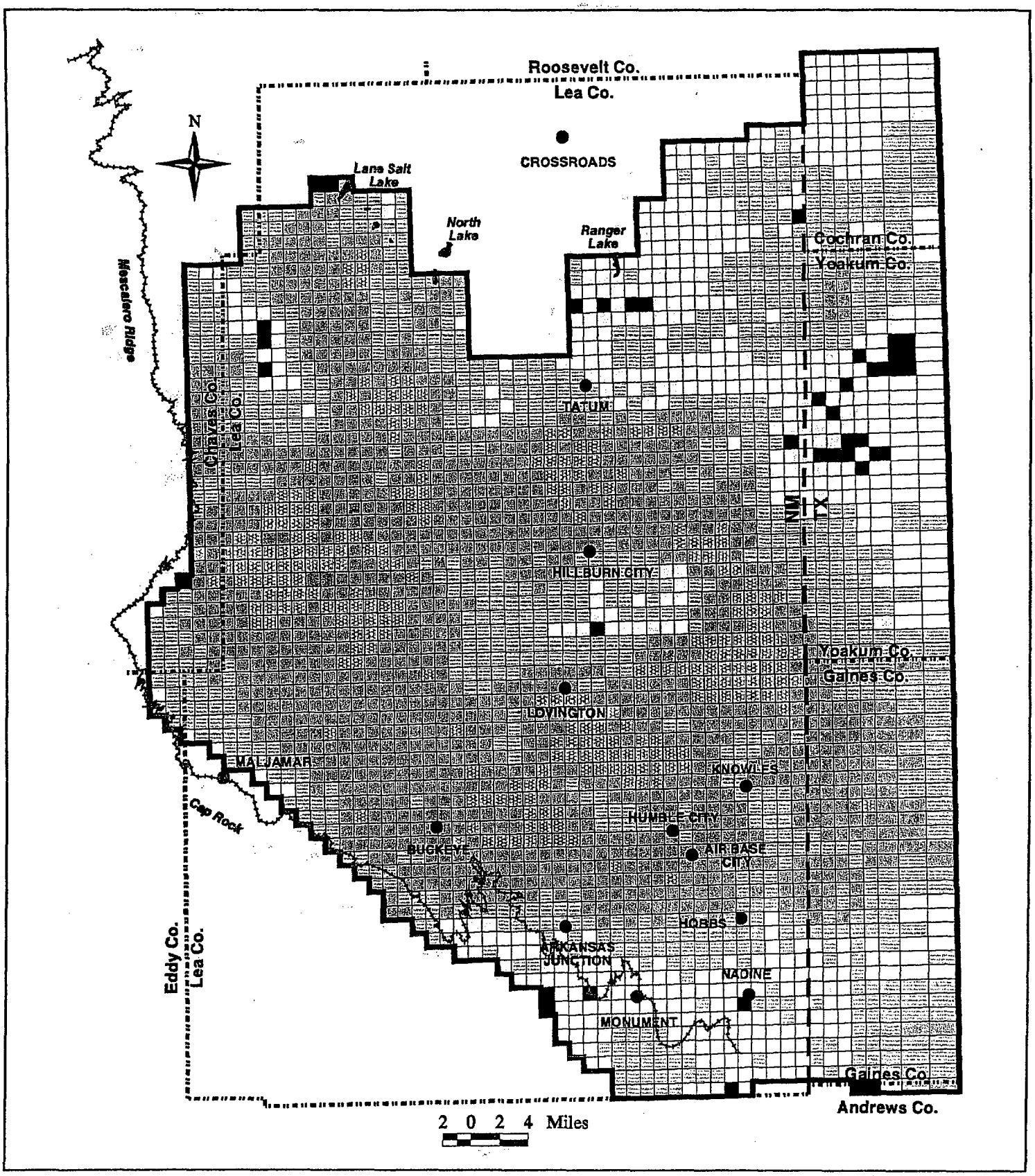
Table 8. Water Budget for Projected Withdrawals (afy).



Legend

- 0
- 1-49
- 50-99
- 100-149
- 150-199
- 200-249
- Model Boundary
- State Line
- County Line

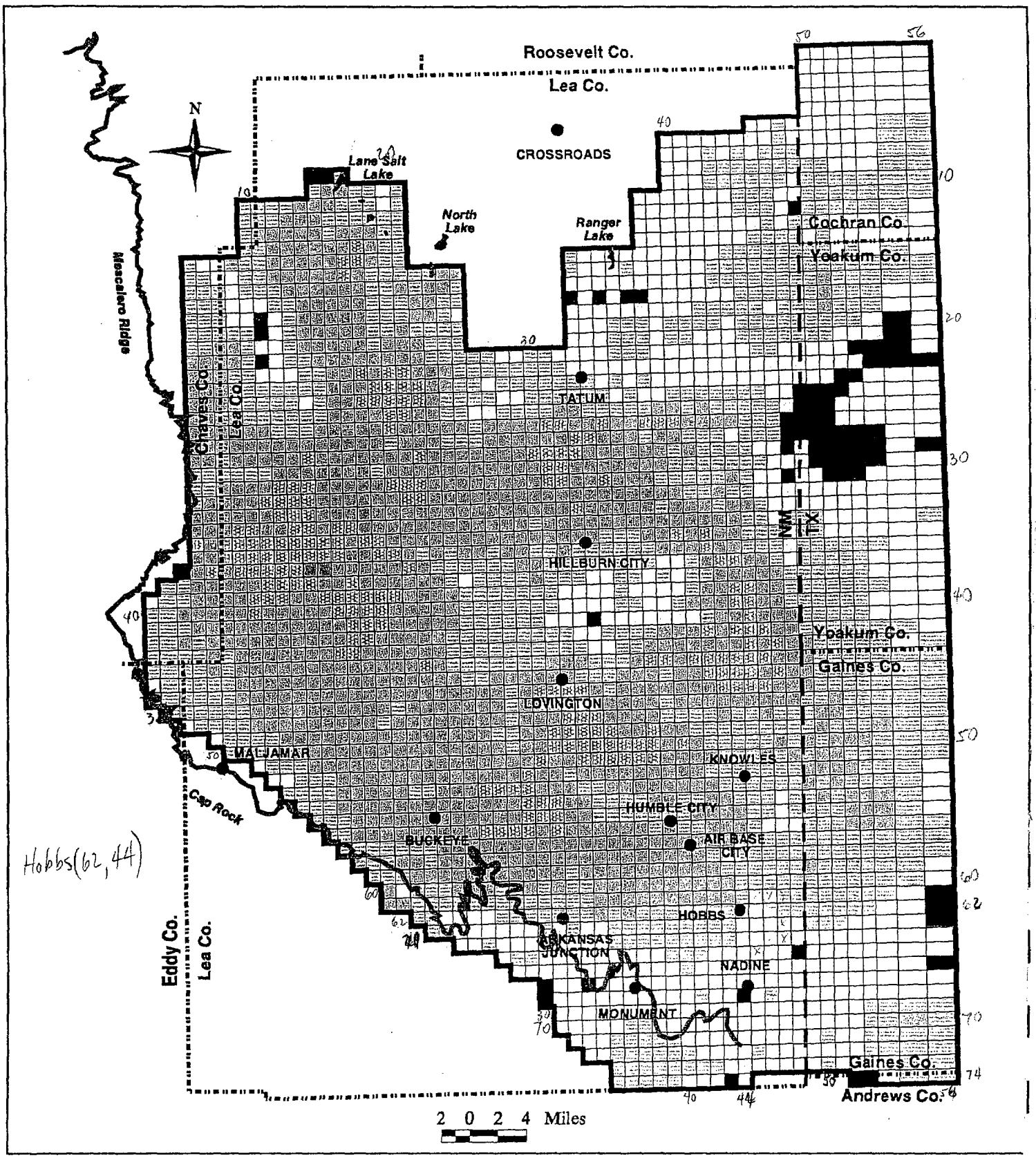
Figure 22: Projected Remaining Saturated Thickness in the Year 2000



Legend

- 0
- 1-49
- 50-99
- 100-149
- 150-199
- 200-249
- Model Boundary
- State Line
- County Line

Figure 23: Projected Remaining Saturated Thickness in the Year 2020 (ft)



Legend

- 0
- 1-49
- 50-99
- 100-149
- 150-199
- 200-249
- Model Boundary
- State Line
- County Line

Figure 24: Projected Remaining Saturated Thickness in the Year 2040

VI. SENSITIVITY ANALYSIS

The model's ability to predict the response of the system to applied stresses depends on how well the mathematical representation of the model represents the actual physical system. Major assumptions were made to characterize the physical system. These assumptions include, hydraulic conductivity, areal recharge, evaptranspiration, and specific yield distributions. These characteristics are known within a plausible range, and the system responds differently to variations in any of these parameters. To quantify the effect of variation of these characteristics on model results sensitivity analysis was performed by varying the values of these parameters within a specified range.

The response of the system to changes in a parameter is tested by varying that parameter while keeping others constant. Steady-state simulations were performed to predict system response to changes in recharge and hydraulic conductivity. Transient simulations were performed to predict system response to changes in specific yield. It is observed that the model is most sensitive to increases in hydraulic conductivity and decreases in areal recharge.

A. Steady-State Simulations.

1. Hydraulic Conductivity.

The effect of increasing and decreasing hydraulic conductivity by 10, 15, and 20% on simulated steady-state water levels are shown in Table 9. Shown are the mean error (ME), mean square error (MSE), and the root mean square error (RMSE) for the differences between simulated and observed steady-state water levels for the modeled area, as a result of changes in hydraulic conductivity. Increasing hydraulic conductivity by only 10% resulted in dry cells and model non convergence. The model is very sensitive to a slight increase in hydraulic conductivity. Decreasing hydraulic conductivity increased the difference between observed and simulated steady-state water levels. This is apparent by the increase in the ME, MSE, and the RMSE indicating that the difference between simulated and observed water levels has increased.

2. Recharge.

Decreasing recharge by 5, 15, and 20% had the same effect on simulated water levels as increasing hydraulic conductivity (Table 10). That is, the model did not converge and is very sensitive to a slight decrease in recharge. Increasing recharge increased the difference between simulated and observed water levels similar to

decreasing hydraulic conductivity.

	Standard	Decreasing K		
		10%	15%	20%
ME	-1.3	-0.7	-6.8	-10.2
MSE	11.5	7.40	10.4	13
RMSE	15.6	10.0	12.1	14.6

Table 9. Statistical Comparison Between Standard and Varied K for the Whole Model Grid.

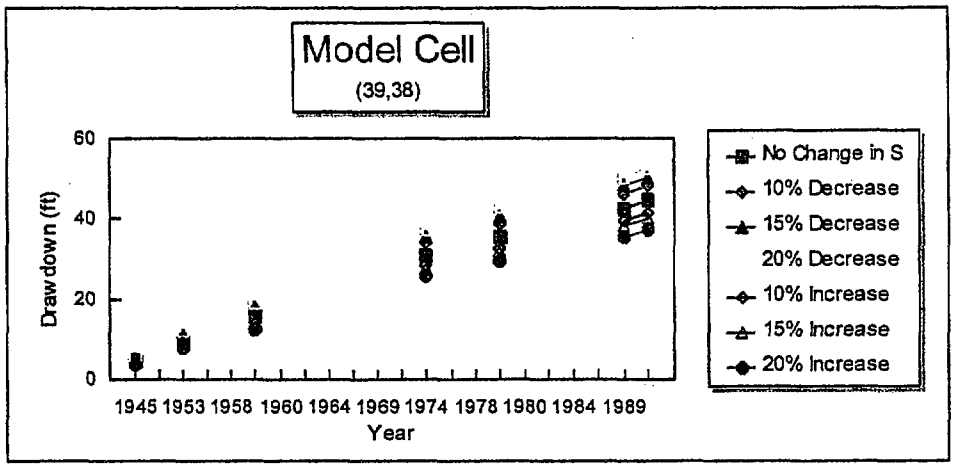
	Standard	Increasing R		
		5%	15%	20%
ME	-1.3	-1.0	-6.6	-9.1
MSE	11.5	7.4	10.0	11.8
RMSE	15.6	10.0	11.8	13.4

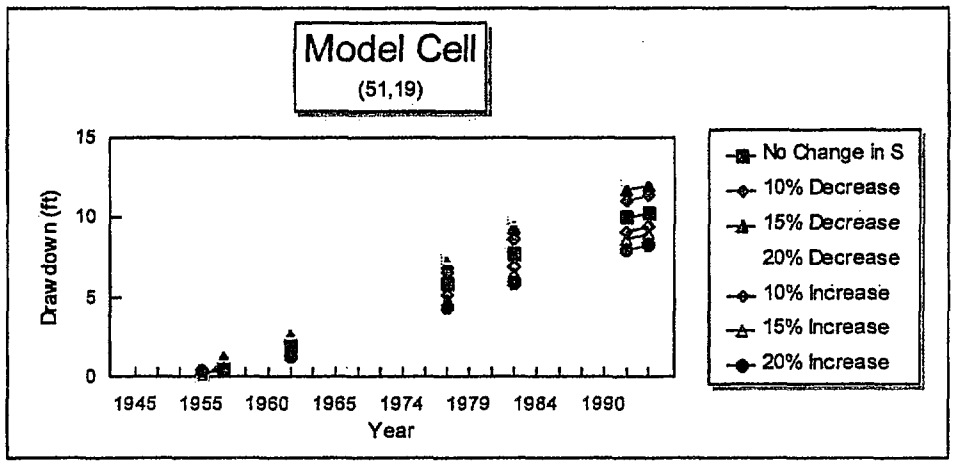
Table 10. Statistical Comparison Between Standard and Varied Recharge for the Whole Model Grid.

B. Transient Simulations.

1. Specific Yield.

The sensitivity of the model to changes in specific yield values was tested using 10,15, and 20% increases and decreases in specific yield. Sensitivity to specific yield was tested only for transient conditions. Figure 25 shows the difference in simulated water levels as a result of changes in specific yield. Increasing specific yield up to 20% decreased simulated drawdowns by about 10 feet in some cells. Decreasing specific yield up to 20% increased drawdowns in some cells by about 10 feet.





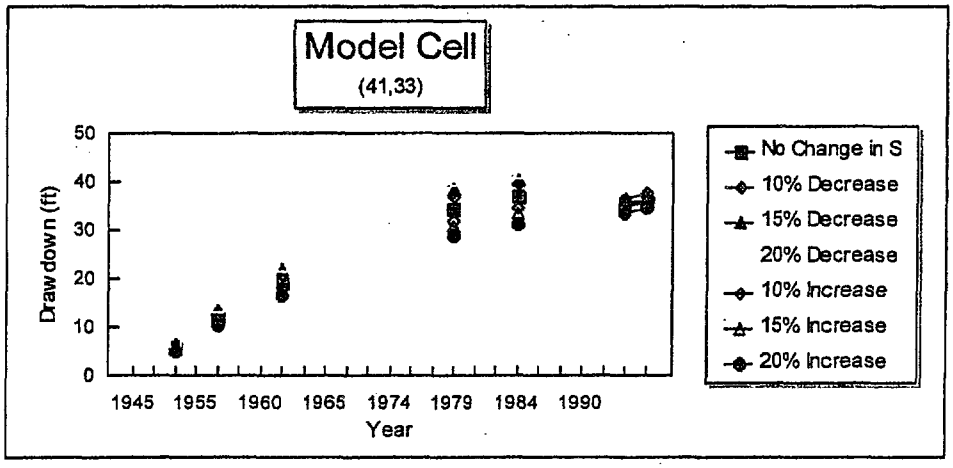
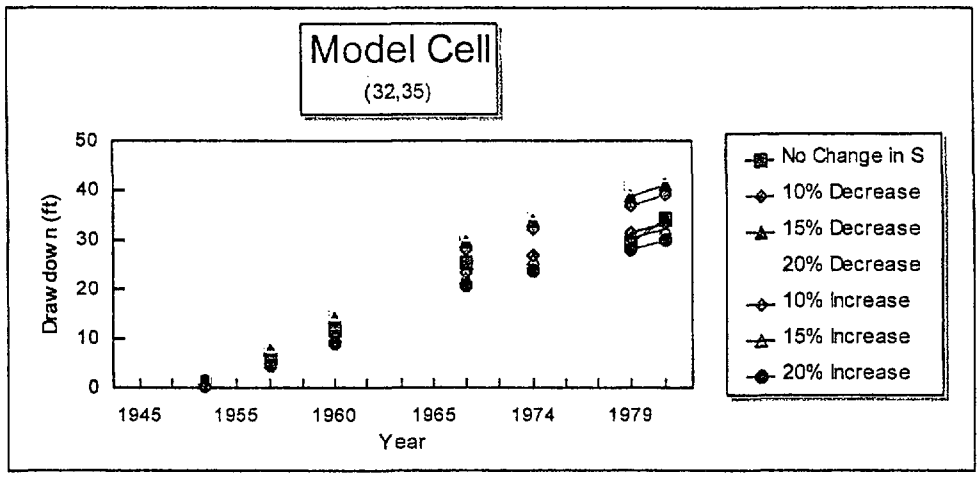
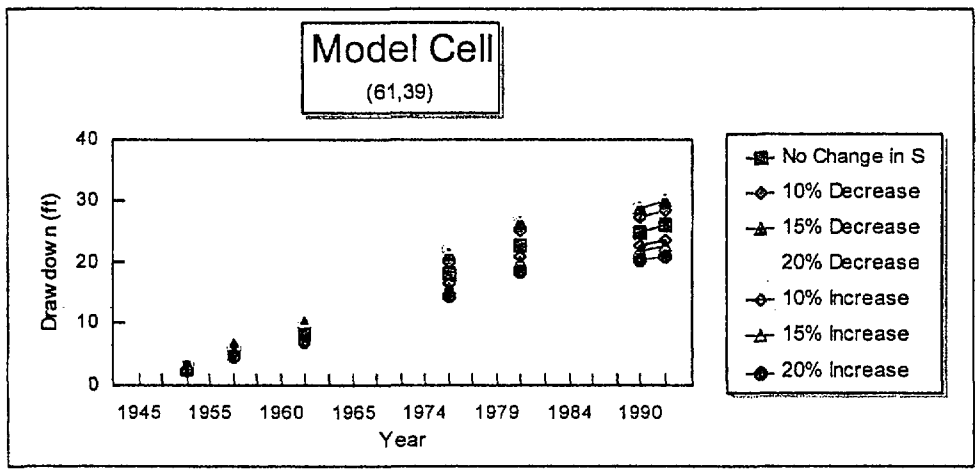


Figure 25. Effect of Changes in Specific Yield on Drawdown in Selected Wells.





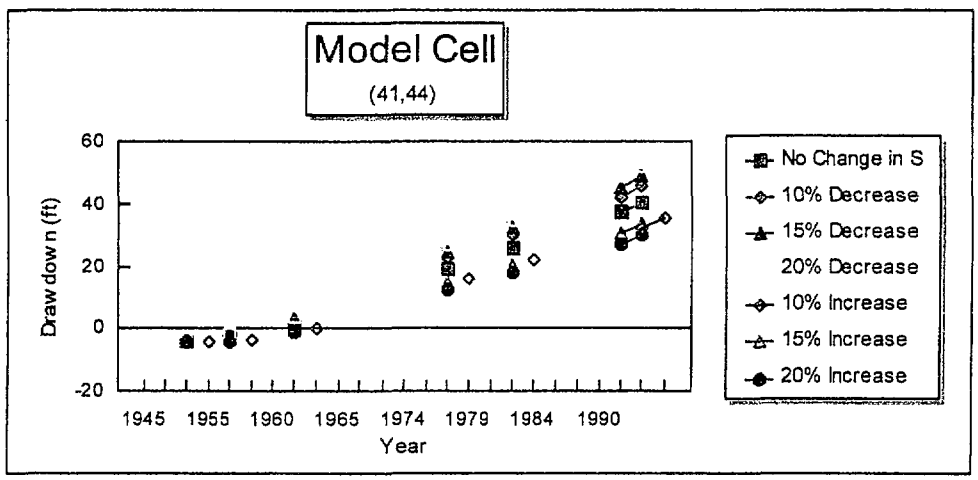
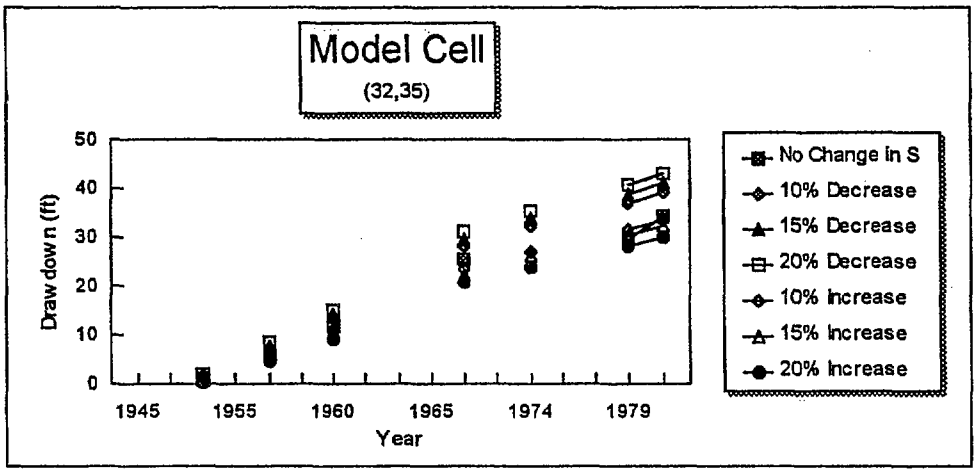
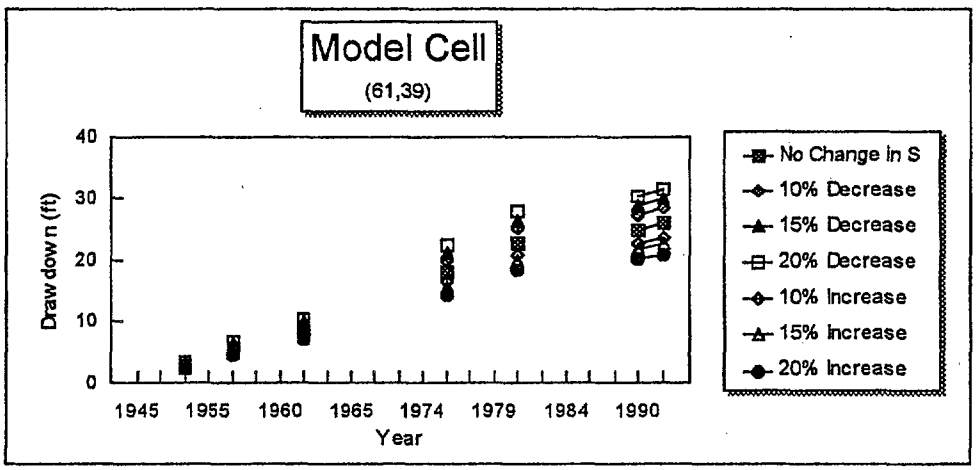


Figure 25 (Cont.)





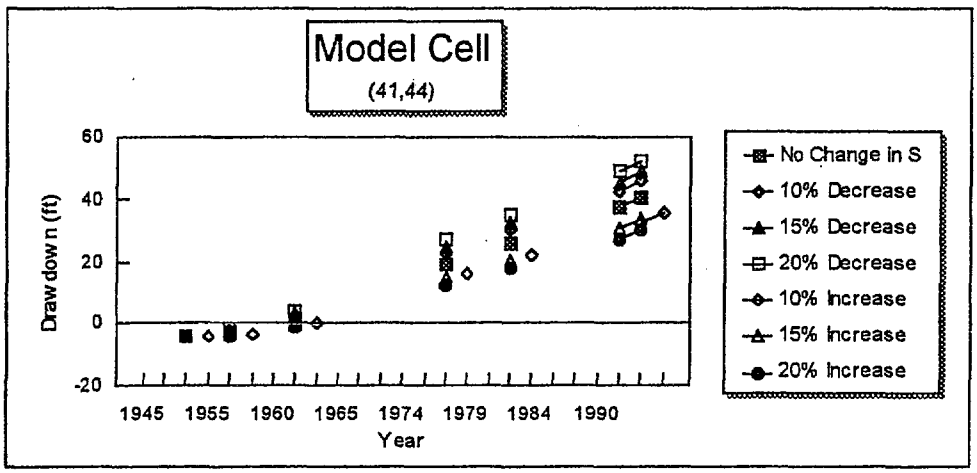


Figure 25 (Cont.)

VII. SUMMARY AND CONCLUSION

A digital groundwater flow model was developed for the Lea County Underground Water Basin for the administration of water rights. Aquifer hydraulic conductivity, transmissivity, and specific yield were estimated within a plausible range reported by previous investigators. Estimated parameters resulted in an acceptable difference between simulated and observed water levels. Steady-state and transient calibrations were performed and results were compared to available field observations. For steady-state conditions, 67% of the simulated water levels are within 10 feet of the observed pre-developed water levels in wells with available observational data. For the transient calibration, 65% of the simulated water levels are within 10 feet of the observed water levels in wells with available observational data at the end of the calibration period.

Sensitivity analysis was performed to study the effects of changes in estimated aquifer parameters (hydraulic conductivity, recharge, and specific yield) on simulated water levels. The model is most sensitive to an increase in hydraulic conductivity or a decrease in areal recharge. Hydraulic conductivity was varied by 10, 15, and 20%. Similarly, areal recharge was varied by the same amount. Increases in hydraulic conductivity or decreases in areal recharge resulted in many dry cells indicating that the model is extremely sensitive to these parameters. Decreases in hydraulic conductivity or increases in recharge caused simulated steady-state water levels to change by 5 to10 feet. Variations in specific yield caused simulated transient water levels to change by 5 to10.

A projected simulation was performed to determine the effect of current withdrawal rates on the sustainability of the aquifer for 40 years. Withdrawal rates from the last stress period were used for future projections because they are close to the long term average (1948-1996) rate used in the model. The exact pumping rates from the basin are not known and there is a high degree of uncertainty about future water uses in this region in both states. Projected simulations indicated that some areas in Yoakum County, Texas, will be de-watered by the year 2010 or even earlier. The current saturated thickness in these areas is less than 50 feet and is not expected to sustain the current pumping rates for a long time. The de-watered area in Texas will increase by the year 2040 to include some parts of New Mexico along the state line. No de-watered areas were simulated in other parts of New Mexico but the saturated thickness decreased at a rate 0.5 to 2.5 feet per year. The rate of groundwater decline in Texas ranges from 2 to 4 feet per year. In general, the current rate of groundwater depletion in New Mexico is sustainable for 40 years.

The effect of Texas pumping on New Mexico water levels was examined by performing a historical simulation with only pumping in Texas. Simulated drawdowns indicated that Texas pumping caused up to 20 feet drawdowns in New Mexico. The

highest drawdowns are along the state line (more than 20 feet) and decreases to 5 feet 3 to 4 miles away from the state line into New Mexico. About 10 feet drawdown is observed in the vicinity of Hobbs.

A calibrated model is a valuable tool for water rights administration and as an aid to planning efforts. As a tool for water rights administration, the model can provide an estimate of the amount of water remaining for appropriation and to calculate the effects of proposed wells upon existing water rights. As an aid to local planning efforts, the model can be used as a tool for calculating regional effects and evaluating different predictive scenarios of future groundwater pumpage in both New Mexico and Texas.

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X. APPENDICIES

APPENDIX A.

STEADY-STATE CELL-BY-CELL FLOW TERMS FOR CONSTANT HEAD CELLS.

Layer	Row	Column	Flow (cu ft/day)
1	1	51	23667.0
1	1	52	-1319.53
1	1	53	19157.9
1	1	54	-13891.9
1	1	55	17181.6
1	2	50	23667.0
1	2	56	-18672.8
1	3	50	22944.9
1	3	56	-26151.6
1	4	50	409.122
1	4	56	-59558.8
1	5	50	-18327.9
1	5	56	-57535.8
1	6	46	3386.91
1	6	47	-3767.84
1	6	48	8245.33
1	6	56	-108614.
1	7	41	2956.08
1	7	42	14840.2
1	7	43	42232.4
1	7	44	-37426.3
1	7	45	-16224.2
1	7	56	-46779.6
1	8	40	2956.08
1	8	56	-55368.8
1	9	40	1124.48
1	9	56	-37699.1
1	10	40	29775.8
1	10	56	-58831.4
1	11	39	2196.17
1	11	56	-57020.3
1	12	38	6633.31
1	12	56	-44573.5
1	13	38	30168.3
1	13	56	-48415.9
1	14	56	-24289.2
1	15	56	-32254.8
1	16	56	-42370.0
1	17	56	-49136.3
1	18	56	-49891.1
1	19	56	-39768.4
1	20	56	-31329.0
1	21	56	-26196.3
1	22	56	-25353.2
1	23	56	-31089.9
1	24	56	-47079.9
1	25	56	-40540.6
1	26	56	-34611.5
1	27	56	-28526.6
1	28	56	-28256.4
1	29	56	-20127.8
1	30	56	-92402.7
1	31	56	-81966.0
1	32	56	-84223.9
1	33	56	-129866.
1	34	56	-184551.
1	35	56	-98776.6
1	36	56	-92250.8
1	37	56	-93665.3
1	38	56	-83201.8

Layer	Row	Column	Flow (cu ft/day)
1	39	56	-124436.
1	40	56	-101209.
1	41	56	-83545.1
1	42	56	-206250.
1	43	56	-173604.
1	44	56	-174228.
1	45	56	-143990.
1	46	56	-219089.
1	47	56	-109209.
1	48	56	-92243.5
1	49	56	-125094.
1	50	56	-104914.
1	51	56	-106301.
1	52	56	-87362.3
1	53	56	-65486.2
1	54	56	-121757.
1	55	56	-104526.
1	56	56	-87325.0
1	57	56	-140935.
1	58	56	-111169.
1	59	56	-74609.0
1	60	56	-203444.
1	61	56	-162613.
1	62	56	-138891.
1	63	56	-144620.
1	64	56	-118264.
1	65	56	-94493.4
1	66	56	-105176.
1	67	56	-78626.8
1	68	56	-108160.
1	69	56	-93829.8
1	70	56	-172679.
1	71	56	-96922.8
1	72	56	-92820.4
1	73	56	-81227.9
1	74	35	-75411.2
1	74	36	-94038.6
1	74	37	-91198.0
1	74	38	-100327.
1	74	39	-113166.
1	74	40	-64788.9
1	74	41	-67778.1
1	74	42	-39155.7
1	74	43	-15567.8
1	74	44	-2419.73
1	74	52	-34515.0
1	74	53	-23976.1
1	74	54	-11869.2
1	74	55	-34263.6

APPENDIX B

A. MODFLOW Basic File

```
Conversion from ModelCad to MODFLOW
Transient model of leaco
     1      74     56      7      4
11 16 0 0 13 0 0 18 0 0 0 12 17
     0      1
     1      1(25I3)              -1
0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0
0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 1
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0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0
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1 1 1 1 1 1
0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0
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  5478    15    1.
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  3652    10    1.
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B. MODFLOW BCF File

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3.4550e+03	3.4750e+03	3.5000e+03	3.5000e+03	3.4750e+03	3.4600e+03	3.4500e+03	3.4500e+03	3.4250e+03	3.4000e+03
3.3750e+03	3.3500e+03	3.3250e+03	3.3000e+03	3.2000e+03	3.2000e+03				
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2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03
2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	3.5400e+03
3.5150e+03	3.5050e+03	3.4900e+03	3.4400e+03	3.4150e+03	3.4000e+03	3.4000e+03	3.4000e+03	3.4000e+03	3.4400e+03
3.4600e+03	3.4750e+03	3.4750e+03	3.5000e+03	3.4600e+03	3.4500e+03	3.4500e+03	3.4500e+03	3.4000e+03	3.3750e+03
3.3750e+03	3.3500e+03	3.3250e+03	3.3000e+03	3.2750e+03	3.2250e+03				
2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03
2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03
2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03
2.9500e+03	2.9500e+03	3.4750e+03	3.4400e+03	3.4050e+03	3.4000e+03	3.3800e+03	3.3800e+03	3.3800e+03	3.4000e+03
3.4200e+03	3.4250e+03	3.4500e+03	3.4500e+03	3.4250e+03	3.4500e+03	3.4500e+03	3.4500e+03	3.4000e+03	3.4000e+03
3.3750e+03	3.3500e+03	3.3250e+03	3.3000e+03	3.2500e+03	3.2250e+03				
2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03
2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03
2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03
2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	3.4000e+03	3.3800e+03	3.3700e+03	3.3700e+03	3.3700e+03	3.4000e+03
3.4000e+03	3.4000e+03	3.4250e+03	3.4350e+03	3.4450e+03	3.4450e+03	3.4450e+03	3.4500e+03	3.4000e+03	3.3750e+03
3.3750e+03	3.3500e+03	3.3250e+03	3.3000e+03	3.2500e+03	3.2250e+03				
2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03
2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03
2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03
2.9500e+03	2.9500e+03	2.9500e+03	2.9500e+03	3.3750e+03	3.3750e+03	3.3700e+03	3.3700e+03	3.3700e+03	3.4000e+03
3.4000e+03	3.4200e+03	3.4400e+03	3.4250e+03	3.4400e+03	3.4400e+03	3.4400e+03	3.4500e+03	3.4000e+03	3.3750e+03
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C. MODFLOW ET File

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D. MODFLOW Recharge File

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-1
-1
-1

E. MODFLOW Output Control File

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F. MODFLOW SIP File

```
  1590     10     1
 0000.1   00.1   0.90   0    0    0
```

G. MODFLOW Well File (Last Stress Period)

1663

1	1	51	23658.6
1	1	52	-1367.9
1	1	53	19115.3
1	1	54	-14029.8
1	1	55	17087.6
1	2	50	23658.6
1	3	50	22900.6
1	4	50	290.6
1	4	56	-59734.1
1	5	50	-19106.6
1	6	46	3106.8
1	6	47	-3893.5
1	6	48	7956.6
1	7	41	2838.5
1	7	42	14585.6
1	7	43	41809.5
1	7	44	-37925.9
1	7	45	-17131.8
1	8	40	2838.5
1	9	40	905.0
1	10	40	29344.2
1	11	39	2000.1
1	12	38	6522.0
1	13	38	29659.7
1	74	35	-94478.8
1	74	36	-68098.2
1	74	37	-67375.8
1	74	38	-73458.6
1	74	39	-82944.0
1	74	40	-80292.3
1	74	41	-90692.3
1	74	42	-57931.9
1	74	43	-29991.7
1	74	44	-30083.9
1	74	52	-58061.7
1	74	53	-37602.8
1	74	54	-38830.2
1	74	55	-49832.6
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1	22	27	5500.0
1	22	28	5500.0
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1	22	32	5500.0
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1	17	33	15500.0
1	18	33	25500.0
1	19	33	25500.0
1	20	33	25500.0
1	21	33	25500.0
1	9	16	-65000.0
1	10	21	-2300.0
1	11	21	-7300.0
1	12	21	-7300.0
1	13	21	-7300.0
1	14	21	-7300.0
1	15	21	-7300.0
1	16	25	-6000.0
1	17	25	-2000.0
1	18	25	-6000.0
1	19	25	-6000.0
1	20	25	-6000.0

1	21	25	-10000.0
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1	9	17	-5000.0
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1	2	51	-704.0
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1	2	54	-704.0
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1	3	56	-26151.6
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1	5	56	-57535.8
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1	6	56	-65168.4
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1	7	55	-1635.0
1	7	56	-46779.6
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1	8	54	-704.0
1	8	55	-1683.0
1	8	56	-55368.8
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1	11	55	-1802.0
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1	12	49	-70288.0
1	12	50	-4070.0
1	12	51	-4070.0
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1	12	55	-1802.0
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1	13	53	-1910.0
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1	13	55	-10145.0
1	13	56	-48415.9
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1	16	50	-11219.0
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