$wym = 294$

Open-File Report 81-101

Open-File Report 81-101

UNITED STATES DEPARTMENT OF THE INTERIOR GEOLOGICAL SURVEY

 \mathbf{t} .

Ŷ.

PRELIMINARY RESULTS OF GRAVITY INVESTIGATIONS OF THE CALICO HILLS, NEVADA TEST SITE, NYE COUNTY, NEVADA

> Open-File Report 81-101 1981

Prepared by the U.S. Geological Survey

for the

Nevada Operations Office U.S. Department of Energy (Memorandum of Understanding DE-AI08-78ET44802) Copies of this Open-file Report may be purchased from

Open-File Services Section Branch of Distribution U.S. Geological Survey *Box 25425, Federal Center* Denver, Colorado 80225

PREPAYMENT IS REQUIRED

 \mathcal{A}

Price information will be published *in the monthly listing "New Publications of the Geological Survey"*

> *FOR ADDITIONAL INFORMATION* CALL: Commercial: (303)234-5888 FTS: 234-5888

Open-file Report 81-101 Open-file Report 81-101

 \sim .

 \sim

Contractor

 \mathcal{L}_{max} , where \mathcal{L}_{max} and \mathcal{L}_{max}

 \mathcal{L}_{max} , and \mathcal{L}_{max}

 $\mathcal{A}=\{x_1,\ldots,x_n\}$, where $\mathcal{A}=\{x_1,\ldots,x_n\}$, \mathcal{A}

 \mathcal{L}_{max} , the same

Contractor Contractor

 $\label{eq:2.1} \mathcal{L}_{\text{max}} = \mathcal{L}_{\text{max}} + \mathcal{L}_{\text{max}} + \mathcal{L}_{\text{max}}$

 $\mathcal{L}^{\text{max}}_{\text{max}}$, where $\mathcal{L}^{\text{max}}_{\text{max}}$

 $\ddot{\cdot}$

 $\mathcal{L}^{\mathcal{A}}(\mathcal{A})$ and $\mathcal{L}^{\mathcal{A}}(\mathcal{A})$ and $\mathcal{L}^{\mathcal{A}}(\mathcal{A})$ and $\mathcal{L}^{\mathcal{A}}(\mathcal{A})$

 $\mathcal{S}^{\mathcal{A}}$ and $\mathcal{S}^{\mathcal{A}}$ are the set of the set of $\mathcal{S}^{\mathcal{A}}$

 $\label{eq:2.1} \frac{1}{2} \int_{\mathbb{R}^3} \frac{1}{\sqrt{2}} \left(\frac{1}{\sqrt{2}} \int_{\mathbb{R}^3} \frac{1}{\sqrt{2}} \left(\frac{1}{\sqrt{2}} \int_{\mathbb{R}^3} \frac{1}{\sqrt{2}} \right) \frac{1}{\sqrt{2}} \right) \, d\mathcal{H}^3 \, d\mathcal{H$

 \sim $^{-1}$

 \mathcal{A}_c , \mathcal{A}_c

 $\Delta \sim 10^4$

 $\sim 10^{11}$ km $^{-1}$

UNITED STATES DEPARTMENT OF THE INTERIOR GEOLOGICAL SURVEY

PRELIMINARY RESULTS OF GRAVITY INVESTIGATIONS OF THE CALICO HILLS, NEVADA TEST SITE, NYE COUNTY, NEVADA

By

D. B. Snyder and H. W. Oliver , Inc.

"This report is preliminary and has not been reviewed for conformity with U.S. Geological Survey editorial standards and stratigraphic nomenclature. Any use of trade names is for descriptive purposes only and does not imply endorsement by the USGS."

Contents

 \mathcal{F}

ILLUSTRATIONS

 $\ddot{}$

 \overline{a}

 $\mathcal{O}(\sqrt{2\pi})$.

 $\sim 10^6$

 $\hat{\mathcal{L}}_{\text{max}}$ and $\hat{\mathcal{L}}_{\text{max}}$ and $\hat{\mathcal{L}}_{\text{max}}$

 $\hat{\boldsymbol{\cdot} }$

 $\mathcal{L}_{\mathrm{max}}$

 $\bar{\beta}$ \sim $\frac{1}{4}$

 \sim \sim

 $\ddot{}$

ABSTRACT

A total of 211 recently established gravity stations supplement ¹²⁸ existing regional stations in providing data for a detailed gravity interpretation of the Calico Hills area of the Nevada Test Site. Reduction of these data to complete Bouguer anomalies combined with a regional correction of -0.8 mGal/km N.15^o E. reveals an elliptical gravity high centered over the exposed Paleozoic rocks. The anomaly has an amplitude of about 7.5 mGal and extent of about 8 km, elongate in an east-west direction. A log from a 770-m drill hole and data from detailed surface geologic mapping, a seismic refraction line, electrical soundings, and magnetic traverses all provide some control on the interpretation of density variations or rock structure within the study area.

Two nearly perpendicular gravity profiles interpreted by $2/\tau$ dimensional modelling suggest the Calico Hills gravity anomaly can be entirely attributed to a lateral density contrast of 0.50 g/cm³ between Paleozoic rocks (= 2.60 $g/cm³$) and the overlying Tertiary tuff (= 2.10 $g/cm³$). This boundary can be constrained vertically to within several hundred meters and thus defines the domical structure of the Paleozoic rocks in this region. The strong contrast of this boundary produces anomalies that mask any gravity anomalies of smaller amplitude that might arise from structure internal to the Paleozoic rocks. Therefore, the existence of a relatively high- or low-density intrusive body in those rocks cannot be confirmed or denied.

INTRODUCTION

Gravity stations were established in the Calico Hills of the Nevada Test Site intermittently from March to August 1979 to help evaluate the area as a possible nuclear waste repository site. Thermal alteration observed in local rocks, surface dikes and plugs, and the structural dome in the Calico Hills suggest the existence of an intrusive body in the subsurface (Maldonado and others, 1979). Moreover, a broad magnetic anomaly occurs over the Calico Hills similar to a magnetic anomaly in the Wahmonie area about 15 km to the southeast where Tertiary felsic intrusive rocks crop out. Detailed gravity studies enlarging upon the work of Healey and others (1980b) were conducted in an attempt to verify the existence of such an intrusive body in the subsurface of Calico Hills.

The Calico Hills, the southernmost topographic expression of Shoshone Mountain, border Jackass Flats in the southwestern part of the Nevada Test Site (fig. 1). The area is best reached on land by roads traversing Jackass Flats from Mercury, about 35 km southeast. Many of the gravity stations were established using vehicles along a dirt road that cuts through the Calico Hills, generally in a north-south direction.

As one approaches the Calico Hills from Jackass Flats, the southern limb of a structural dme is being ascended (fig. 2). A facade of colorful altered tuffs forming a south-facing wall gives the area its name. Immediately upon crossing the topographic high, one stands directly on Paleozoic rock. To the left and right, high-angle faults form the contacts between Paleozoic rocks and tuffs in place of the dome-related erosional contacts. As one continues northward, an ever thickening tuff sequence overlies the Paleozoic rocks and continues into the Timber Mountain region far to the north.

ACKNOWLEDGMENTS

Appreciation is expressed to D. R. Jefferis, H. M. Van Buren, and D. A. Ponce for planning and coordinating the field work; to R. B. Livermore, of Fennix & Scisson, Inc., for many gravity observations, terrain corrections, and elevation control and for data reduction; R. N. Harris and J. B. Spielman for assisting in the gravity observation, and to D. C. Hohbach and $E. G.$ Miller for providing a part of the elevation control; C. R. Karish and T. A. Sagara drafted the illustrations.

Figure 1.--Index map of the Nevada Test Site showing the location of 1.--Index map of the Nevada Test Site showing the location of
the Calico Hills. The UE25a-3 drill-hole site and the border of the gravity study area of figure 8 are shown.

 $\hat{\mathcal{L}}_1$

 \sim

U ALLUVIUM (QUATERNARY)

- **A** VOLCANIC ROCKS (TERTIARY)
- **ELEANA FORMATION** (DEVONIAN-MISSIBIPPIAN, LOWER PLATE)
- CARBONATE ROCKS (DEVONIAN, UPPER PLATE)

⁴UE ²**6-S** DRILL HOLE

YYY THRUST FAULT (SAWTEETH ON UPPER PLATE)

FAULT-DASHED WHERE APPROXIMATELY LOCATED

.... CONCEALED FAULT

SULK DENSITIES IN g/cm ⁸

Figure 2.--Generalized geologic map of the Calico Hills, showing location of the drill-hole UE25a-3, seismic-refraction profile (S), and two modelled gravity profiles *(A-A'* and B-B'). Bulk densities determined by removing any topographic correlation with local gravity anomalies using the cmbined elevationcorrection factor of Nettleton (1976, p. 89) are indicated at the location at which each was determined.

GEOLOGY OF THE CALICO HILLS

The Calico Hills lie on the southern flank of a structural dome (Maldonado and others, 1979 fig. 2) that exposes thrust plates of Devonian limestone and dolomite superposed on argillite of the Eleana Formation (Mississippian-Devonian). The Paleozoic rocks in turn are overlain by volcanic rocks of Tertiary age (about 8 to 14 million years old). Porphyritic rhyolitic plugs and an andesite dike of Tertiary age intrude the Eleana Formation. Intense alteration of the volcanic rocks has destroyed their pyroclastic texture; the carbonate rocks are locally altered to calcite, quartz, montmorillonite and in places metamorphosed to a white marble. Local alteration of the Eleana argillite has produced hornfels in places and irregularly bleached the rock throughout. A network of radially trending faults appears to be related to the domical structure. The dome, the radially trending faults and the extensive hydrothermal alteration all may have been caused by the intrusion of a magma body at depth.

In addition to the low-angle Mesozoic(?) thrust fault, the Calico Hills area is structurally influenced by its setting in the Basin and Range province and on the south edge of the Timber Mountain caldera complex. Near the Calico Hills, the caldera complex has deposited many layers of tuffaceous sediments that thicken and become more complex to the northwest. The generally northsouth trending high-angle faults found throughout the Basin and Range province are present in the Calico Hills. These high-angle faults produce no impressive fault scarps in this area but do offset most tuff units and appear to modify the symmetry of the domical structure (McKay and Williams, 1964).

Geophysical basement rocks in the Calico Hills belong to the Eleana Formation of Carboniferous age. At least 2,350 m of these clastic rocks have been divided into ten major lithologic units, denoted in ascending order A through J (Poole and others, 1961), that contain argillite, quartzite, limestone, and conglomerate in varyious combinations. Only units I and J are of direct interest to gravity modelling. Unit J is predominately argillite; the uppermost unit is quartzite (Poole and others, 1961). At Calico Hills, unit J has been redivided by Maldonaldo and others (1979) into argillite, altered-argillite, and calcareous altered-argillite intervals. Digitized density logs were obtained for these rock units penetrated by the drill hole UE25a-3. Typical values are 2.55 g/cm³ for the altered argillite intervals and 2.50 $g/cm³$ for the argillite interval. Some fractures in the argillite interval are coated with, and may have been sealed or filled by, secondary mineralization during hydrothermal alteration but opened during drilling. The unit I marbles exhibit considerable variation in density. A median value of 2.60 g/cm³ matches density values used in gravity studies in other regions of the Nevada Test Site where the Eleana Formation occurs (Wahl, 1969) and will represent unit I as well as lower stratigraphic subunits of the Eleana in modelling the Calico Hills.

The relatively thin (less than 200 m) Devonian carbonate rocks of the upper thrust plate directly overlying the Eleana in places will be considered part of the Eleana and assume its density values in the gravity interpretation.

Tertiary tuffaceous sediments unconformably overlie the Eleana Formation in most of the Calico Hills area. Near the northern edge of the study area, nonwelded to densely welded tuff units totaling many thousands of meters rest upon the Paleozoic rocks (Orkild and O'Connor, 1970). This large stack of tuffs includes all or parts of the following formations, in ascending order: Crater Flat Tuff, Rhyolitic tuffs and lavas of Calico Hills, Paintbrush Tuff, and Timber Mountain Tuff. To the south, only the Calico Hills and Paintbrush Tuffs appear. These tuffs vary in density from 1.08 g/cm³ to 2.43 g/cm³ depending on the degree of welding and amount of alteration (Carr, 1966; App. 2). Density determinations in the immediate vicinity of the Calico Hills utilizing
the effect of density upon topographically related gravity anomalies (fig. 2) the effect of density upon topographically related gravity anomalies (fig. 2) indicate a range of values of 2.17 + 0.14 g/cm³. Bulk tuff densities of 2.1 $g/cm³$ and 2.25 $g/cm³$ will be considered.

Quaternary alluvium covers the surface in both the shallow basin directly north of the Calico Hills proper and in Jackass Flats. A density value of 1.7 $+$ 0.1 g/cm³ determined by the topographic effect mentioned above is deemed appropriate for this stratigraphic unit (D. L. Hoover, oral commun., 1980).

GRAVITY DATA COLLECTION AND REDUCTION

Healey and others (1980b) assembled regional gravity coverage that had existed in the southwest quadrant of NTS for a number of years. Better interpretation of the local gravity field required the more detailed field work begun in March 1979. This work established (1) 95 newly surveyed stations on the southern slopes of the Calico Hills reached by jeep or on foot, (2) 22 helicopter stations on isolated peaks, and (3) a line of 61 closely spaced stations trending north-northwest from the road pass in the Calico Hills (profile A-A', fig. 2). The principal facts for these stations and the 128 stations of Healey and others (1980b) are given in Appendix 1 at the end of this report.

All 339 stations have been reduced by the standard U.S. Geological Survey method using a computer program of D. F. Barnes (written commun., 1978) and Plouff (1977). The gravity datum is IGSN 1971 (Morelli, 1974); free-air anomalies are referenced to GRS 1967 (International Association of Geodesy, 1971). Terrain corrections to a radial distance of 0.068 km from the station were made in the field; those from 0.068 km to 0.59 km were made using cylindrical ring templates (Hammer, 1939). The remaining correction for terrain variations out to 167 km were computer-generated using the U.S. Geological Survey modification of the Defense Mapping Agency digital terrain data. Reduction density is 2.67 g/m^3 . The prime base station, Mercury, tied to IGSN Network stations DOD 0363-2 in Las Vegas and DOD 4027-1 in Indian Springs, is located at the geophysics laboratory bench in the USGS buildings at Mercury, Nev. The observed gravity value there is 979,518.80 mGals. All field measurements for these stations were made using LaCoste and Romberg gravity meters G17 and G177. These meters were calibrated on the Mount Hamilton or Mount Charleston gravity calibration loops before and after each field session (Barnes and others, 1969; Ponce and Oliver, unpub. data).

The use of residual gravity values in this study removes large scale effects of lateral density variations in Paleozoic rocks and even deeper structures. Diment and others (1960) calculated a regional gravity field using only stations on Paleozoic rock outcrops; in the Calico Hills region a

7

linear regional gradient from this study is -0.88 mGal/kmn at N.15 $^{\circ}$ E. (fig. 3). In addition, isostatic corrections were calculated using differing compensation models, producing linear gradients about -0.25 mGal/km less. Datum shifts and differing reduction methods further increase the possible inaccuracy of the regional gradient by $+0.05$ mGal/km, but the concept of direct contact with the local Paleozoic rocks makes the gradient of Diment and others the most attractive in a study area as limited as the Calico Hills. The uncertainty in this regional gradient of $+0.30$ mGal/km causes an uncertainty of several degrees in the slope of the Paleozoic surface on the north and south flanks of the Calico Hills dome.

OTHER INVESTIGATIONS AT CALICO HILLS

Drill-Hole exploration

The UE25a-3 exploratory drill hole was completed on October 10, 1978, at Nevada State Coordinates N. 234,498, E. 183,772 m (fig. 2). The drill penetrated 771.2 m of Eleana Formation without finding an intrusive crystalline body that it was expected to verify (Maldonaldo and others, 1979, p. 1). The hole penetrated 422.5 m of the argillite subunit of the Eleana Formation unit J, 298.1 m of the lower subunit of unit J, and 50.6 m of unit I. Because these stratigraphic Eleana subunits of Poole and others (1961) differ from the zones of hydrothermal alterations, the following intervals were subsequently determined on the basis of the presence or absence of thermal alteration (Maldonaldo and others, 1979, p. 5-12): argillite, 0-416.1 m; altered argillite, 416.1-676.9 m; and calcareous altered argillite, 676.9- 720.6 m.

The metamorphism of the rocks increases with depth. Although the argillite interval was not visibly bleached by thermal metamorphism, hydrothermal solutions migrated through fractures, depositing kaolinite. Heat drove off organic carbon, resulting in the bleaching of black argillite to lighter colors within the altered argillite and calcerious altered-argillite intervals. The marble present, representing the highest grade of metamorphism found in cored rocks, formed by thermal alteration or metamorphism of Eleana unit $I(?)$ carbonates at temperatures near 350^oC.

Heat-flow measurements in this drill hole indicate a high rate, about 3.2 HFU. Temperatures range from 14° C at the surface to 47° C at the bottom depth of 750 m. Deep upward migration of water is hypothesized for the Calico Hills region (Sass and others, 1980).

Seismic surveys

A seismic refraction line surveyed the northern half of the Calico Hills area (line S, fig. 2) using stations extending northeastward from the UE25a-3 drill hole across the alluvial basin north of the Calico Hills proper and into the wash to the northeast, Topopah Wash. Four main features appearing in the seismic interpretive cross section (fig. 4) have been incorporated into the gravity modelling. Stratigraphic units with velocities equal to or less than 3.3 km/s are considered alluvium or highly weathered rocks. The 3.7 km/s unit is interpreted as the southern edge of the tuff sequence forming Shoshone Mountain. The 4.0 to 4.5 km/s units correlate with the Eleana Formation argillites, the higher velocities associated with thermal alteration and

DISTANCE. KILOMETERS

Figure 3.--Removal of the regional gravity gradient of -0.88 mGal/Iam at N15°E from the complete Bouguer anomaly gravity profile A-A'. The resulting residual gravity profile is used for interpretation in this study.

9

 $\ddot{}$

Figure 4.--Interpretive cross section of the seismic refraction line at Calico Hills produced by Lee Pankratz of the U.S. Geological Survey (written comun., 1980). In this report, all units indicating velocities less than 3.5-km/s unit were considered alluvium or highly weathered rocks, $p = 1.7-1.8$ g/cm³, the 3,7-km/s unit was interpreted as tuff, $p=2.1-2.25$ g/cm³, the 4.0-4.5-km/s units as Eleana Unit J argillites, $p=2.5-2.55$ g/cm³, and the 5.3 km/s unit as Eleana Unit I, $p=2.6$ g/cm³. Vertical exaggeration, 2.85.

 $\overline{\mathbf{r}}$

related fracture filling. The 5.3 km/s stratigraphic layer is interpreted as the marbleized Eleana $I(?)$ unit, the deepest rocks characterized by the drill cores and seismic survey.

Electrical soundings

Hoover, Chornack, Nervick, and Broker (unpub. data) have established five magnetotelluric soundings, 23 Schlumberger vertical electrical soundings, and six dipole-dipole induced polarization traverses in the vicinity of the Calico Hills. A resistivity log for the depth interval of 620- to 760-m in UE25a-3 is also available (Maldonaldo and others, 1979, pl. 1).

Resistivities generally increase from values of about 50 ohm-meters in the unaltered Eleana Formation unit J to about 1000 ohm-meters in the deeper marbleized Unit I(?). South of the drill hole, high resistivities suggest the absence of unaltered Eleana unit J near the surface to the 1500 m depth explorable by the electrical methods. Below about 1 km, only the magnetotelluric data provide information. These data show a high resistivity body beneath the Calico Hills with a top at about 2.6 km depth and bottom around 100 km.

The dipole-dipole arrays indicate faulting within a conductive zone associated with Topopah Wash that is linear and strikes north-northeast. The zone has an apparent resistivity of 20 to 100 ohm-meters at 100 m depth, inferred to be due to unaltered Eleana unit I.

Magnetic surveys

Both aeromagnetic and ground magnetic data were collected along traverses in the Calico Hills area. The aeromagnetic map, figure 5, shows the sharp anomalies that originally turned attention to the Calico Hills. Anomalies from two sources may contribute to the broad anomaly over the Calico Hills (G. D. Bath, oral commun., 1980). The large circular anomalies are thought to be caused by the near-surface argillites of the Eleana Formation that have high magnetic content in this area. The broader anomaly probably arises from a deeper source. The ground traverse coincides with the detailed gravity survey and passes through the saddle between the two aeromagnetic maxima (fig. 5). The cross sections in figure 6 compare the ground magnetic traverse, the detailed gravity survey, and geologic outcroppings. In general, exposures of Paleozoic rocks correlate with magnetic highs, and alluvial pods and lenses match up with gravity lows.

Previous gravity measurements

No detailed gravity measurements had been made at Calico Hills prior to this study. Extensive measurements support weapon-testing north and east of this area (Diment and others, 1960); only regional reconnaissance, summarized by Healey and others (1980b), encompassed the present study area.

INTERPRETATION

General statement

Of primary concern to this study was to confirm or deny the existence of an intrusive body at Calico Hills. The extent and internal structure of the Eleana argillite are another major interest. Aeromagnetic measurements (fig. 5) reflecting a very distinct magnetic anomaly at Calico Hills first drew attention to the area as a possible location of a deep intrusive body. On the

 \sim \sim $\ddot{}$

 $\bar{\alpha}$

ALLUVIUM (QUATERNARY)

EL VOLCANIC ROCKS (TERTIARY)

ELEANA FORMATION (DEVONIAN-MISSIS81PPIAN. LOWER PLATE)

E CARBONATE ROCKS (DEVONIAN, UPPER PLATE)

 $+$ UE 25a-3 DRILL HOLE

I~rTHRUST FAULT (8AWTEETH ON UPPER PLATE)

FAULT-DASHED WHERE APPROXIMATELY LOCATED

.... CONCEALED FAULT

AEROMAGNETIC CONTOURS CONTOUR INTERVAL 18 10 **GAMMAS.** LOCAL MAXIMA AND **MINIMA** INDICATED.

'Figure S.--Aeromagnetic map of the Calico Hills and vicinity. The three sharp anomalies directly over the Calico Hills are prominent; a broader, lower amplitude anomaly is more subdued. The sharp anomalies are attributed to locally magnetic Eleana argillites (G. Bath, oral commun., 1980). Magnetic contours are residuals with respect to the 1975 IGRF corrected to November 1977 with a 5,000-gamna constant added. Survey was flown 400 ft above the surface (U.S. Geological Survey Openfile Map 79-587).

Figure 6.--High resolution north-south geologic cross section C-C' with coincident gravity and ground magnetic profile. Geology is taken from Orkild and O'Connor (1970), PMej represents Eleana
argillite, Ddn overthrust limestone. The gravity profile contains argillite, Ddn overthrust limestone. the residual Bouguer anomalies for reduction density of 2.67 $g/cm³$. The profile is both defined and located by the line of closely spaced stations in figure 8, centered at lat 36°52'30" N, long 116°18' 0" W. The total magnetic intensity near ground level was measured with a proton magnetometer (G. Bath, written commun., 1980). Magnetic highs tend to correlate with Paleozoic rock outcrops; gravity lows are associated with surficial lenses of alluvium.

other hand, surface geologic mapping and drill-hole logs indicate that Calico Hills is unusual within the southwestern Nevada Test Site in its surface exposure of Paleozoic argillites. These relatively dense rocks near or directly at the surface are another possible source for the Calico Hills gravity high

In some other areas of southern Nevada, the coincidence of distinct gravity and aeromagnetic highs is associated with known intrusive bodies, as at Wahmonie and Black Mountain (Healey and others, 1980a; 1980b; Spengler and others, 1979, p. 32-35). The gravity high at Calico Hills could possibly result from an intrusive body situated at depth greater than the drill-hole depth of about 800 m beneath the surface. However, intrusive masses do not necessarily produce gravity highs. For example, no local anomalies are associated with the Climax and Gold Meadows (Healey and Miller, 1963) stocks located in the northeastern part of the Nevada Test Site. The quartzite and carbonate country rocks that these stocks intrude themselves produce strong density contrasts with neighboring Cenozoic rocks, whereas the density contrast between the intrusive rock and quartzite is negligible (Healey and Miller, 1963, p. B64; Wahl, 1969, p. 8). Moreover, many felsic intrusions in eastern California are the source of gravity lows because the rocks they intrude, such as schist, amphibolite, and marble, have higher densities (2.8- 3.0 $g/cm³$) than the intrusions (2.7 $g/cm³$) (Oliver, 1977, 1980; Oliver and Robbins, 1980). Thus intrusive magma bodies are observed to have a wide variety of gravity signatures.

With this fact in mind, the interpretation of the gravity field at Calico Hills must begin by defining the gravity effects attributable to the Paleozoic/Tertiary rock interface. The significant density contrast (0.3-0.5 $g/cm³$) existing between these rocks makes this stage of interpretation critical. Once the effect of these relatively near surface structures is removed, source bodies at depth within the Paleozoic rocks can be explored. These deep bodies have low density contrast values (less than perhaps 0.2 $g/cm³$) and may be several kilometers below the surface. As a result, perturbations in the measured surface gravity field caused by their presence will be of rather low amplitude and difficult to define. This fact suggests the key question in interpreting the Calico Hills gravity field: Can the subsurface extent of the Paleozoic Eleana argillite be defined precisely enough to permit examination of its internal structure, including the existence of an intrusive body?

Contour-map interpretation

From a regional perspective, the Calico Hills gravity field is an elliptical anomaly of several milligals amplitude located on a northeasttrending shelf between high-gravity regions to the southeast near Wahmonie and Skull Mountain and low-gravity areas at Shoshone and Yucca Mountains (fig. 7). Over the study area, the gravity field, with a regional gravity gradient removed (fig. 8), shows an edge of the large gravity low at Shoshone Mountain to the north, the oval 7-mGal Calico Hills anomaly at the center, and, to the south, 1 to 3-mGal gradients within the northeast-trending band of figure 7. Removal of the regional gradient emphasizes the oval Calioo Hills anomaly (fig. 8).

116° 30' Figure 7.--Regional 2-m mGal complete Bouguer anomaly gravity contour map of the southwestern^{116°} 5' quadrangle of the NEvada Test Site; hachures indicate closed areas of lower gravity. Point A, the Calico Hills, appears as a small positive anomaly within a northeast-trending band that separates a gravity low to the northwest from a gravity high to the southeast near Skull Mountain. The Calico Hills anomaly is rather subdued and irregular in shape. Gravity contours are from D. L. Healey (written commun., 1979). Base from U.S. Geological. Survey, Death Valley 1:250,000 quadrangle, 1970.

Regionally, the Calico Hills anomaly is very subdued, especially at its east edge, and somewhat irregular in shape (fig. 7). Irregularly shaped anomalies are characteristic of shallow causative bodies, irregular in shape or distribution. An isolated sizable intrusive body, on the other hand, produces an anomaly noticeable mainly for its continuity and symmetry, although overlying rock units could make critical contributions to the extent and symmetry of the overall anomaly.

The Calico Hills residual anomaly (fig. 8) is roughly defined by the -23 mGal contour. The anomaly opens to the northeast and east beyond the -21 mGal contour in the vicinity of a ridge of tuffaceous rocks in the direction of the northeast-trending band noted earlier. To the north, the limb of the anomaly loses definition as residual anomaly values decrease to -46 mGals because of gradual increase in elevation of rocks having densities less than the 2.67 g/cm reduction density and possibly because of the deepening of the Paleozoic basement northward. In a three-dimensional perspective of the residual gravity map (fig. 9), the 1- to 5-mGal anomalies superimposed on the main anomaly appear clearly. These 0.5- to 1.0-km diameter anomalies do not correlate with changes in surface rock units, but many of them do correlate with local topographic features, and they have been used to estimate the bulk densities of those features relative to the 2.67 g/cm³ Bouguer reduction density (fig. 2). Some of the small anomalies may also be produced by local hydrothermal alteration causing density contrasts within the rock units.

In general, removing the regional gradient enhances the Calico Hills anomaly but does not simplify its outline. The proximity of the large gradient to the north with the generally small changes in gravity to the east further suggests complex near-surface density changes, perhaps masking the effects of more deeply seated causative bodies.

North-south profile interpretation

Gravity- profile modelling along both cross sections discussed in this study utilized a $2/\tau$ dimensional computer program (Cady, 1980) wherein source bodies are invariant in cross section but terminate a specifiable finite distance along strike in either direction. This $2^{1}\!/ _{Z}$ dimensional approach has the convenience and speed of 2-dimensional calculations yet retains much of the generality of 3-dimensional techniques.

The first interpretive cross section bears slightly west of north (A-A' in figs. 8 and 9). The profile begins over the alluvium in Jackass Flat, passes through a narrow gap in the Calico Hills proper, crosses over alluvium and ends to the north in the Shoshone Mountains. The gravity profile is based on stations at 500 m spacing in the southern third, 75 m spacing in the center, and highly variable spacing in the northern third.

Much of this profile appears in figure 10 along with an aeromagnetic profile and mapped surface geologic units. The 950-gamma magnetic high centered on the profile correlates best with a 1.3-nGal gravity feature rather than the broad 7-mGal high and occurs over an outcrop of Eleana argillite. A broad magnetic high with an amplitude of 150 gammas roughly correlates with the broad gravity feature and also with the overall surface expression of argillite. The 150-gamma correlations suggest that the argillite may be responsible for both geophysical anomalies.

GRAVITY STATION

GRAVITY CONTOURS **CONTOUR INTERVAL 18 1 mest. HACHURED CONTOURS** INDICATE AREAS OF LOW **OGAVITY** CLOSURE.

Figure 8.--One mGal contour residual gravity map of the Calico Hills study area. The Calico Hills gravity anomaly is roughly defined by the -23-mGal contour. The string of gravity highs just northwest of the large closed low northeast of the main anomaly are topographic effects due to an incorrect Bouguer reduction density. The large gravity low at the top of the figure corresponds with the south edge of Shoshone Mountain. The numerous small highs and lows superimpsoed on the main anomaly do not correlate with mapped geologic units and may. indicate density variations internal to the Paleozoic rocks.

Figure 9.--Three-dimensional perspective of the residual gravity field contoured in figure 8. The modelled profiles A-A' and B-B' are indicated both here and in figures 2 and 8. Drill hole UE25a-3 is indicated by the triangular symbol. The 1 to 5-mGal, 0.5 to l-km-diameter gravity anomalies superimposed on the Calico Hills main anomaly are rather noticeable in this perspective, as is the steep drop in gravity to the north.

Figure 10.--North-south geologic cross section A-A' with coincident residual gravity and aeromagnetic profiles. Tpt represents tuffs, PMe the Eleana argillite, and Qal alluvium. Gravity is from figure 8, aeromagnetic profile from figure 5. The large aeromagnetic spike correlates with the highest gravity values and Paleozoic outcrops. The main positive gravity anomaly matches a low-amplitude swell in the aeromagnetic profile and the structural dome.

Figure 11.--Modelled interpretation of gravity profile $A - A'$. The central dotted line indicates the A , "Best" interpretive model uses 1 7-4/-3 .
A, "Best" interpretive model uses 1 7-4/-3 interpretive model and respect or drill holeUE25a-3. X is calculated gravity profile A-A'. The central dotted line indicates to the represent interpretive model uses 1.7-g/cm3 bodies to represent alluvium. 2.1-g/cm3 being altered arguing tuff. The 2.50 altered argillite.
Tepresente en de A, "Best" interpretive model uses 1.7-g/cm³ bodies to represent alluvium, 2.1-g/cm³ bodies to represent Tertiary tuff. The 2.50-g/cm³, 2.55-g/cm³, and 2.60-g/cm³ bodies represent the argillite and remaining bulk represent Tertiary tuff. The 2.50-g/cm³ bodies to represent alluvium, 2.1-g/cm3 bodies to altered argillite, and remaining bulk of the Eleana Formation, respectively. The 2.25 σ /cm3 and intrusive rhyolite dike. B. a are 2.25 g/cm³ and a 3.0-g/cm³ pluton replaces the rhyolite dike. The smaller density contrast
between tuff and argillite required a thicker tuff section and a positive density contrast between

 $\frac{2}{3}$

 $\hat{\mathcal{L}}$

w)

 \mathbf{r}

The detailed profiles of figure 6 cover roughly the central third of the entire north-south cross section. Ground magnetics and gravity correlate nicely, matching 200 to 300-gamma and 1-mGal anomalies respectively. Surface geology generally indicates alluvium near gravity lows; Paleozoic rock crops out near gravity highs. The broad gravity high toward the left end of the profile is not readily explainable by any mapped rock units.

Modelling along profile A-A' started with the published geologic cross section for the area (McKay and Williams 1964: Orkild and O'Connor, 1970). Drill hole UE25a-3 provided geologic control for the top₃800 m near the center of the profile, including the density suite of 2.50 g/cm, 2.55 g/cm³, and 2.60 $q/cm³$ for the Eleana rocks. The seven bodies in the models (figs. 11A, B) are based on these constraints plus a seismic profile (fig. 4) that nearly parallels this profile between the 0 and 5 km values (fig. 2). This seismic profile indicated much complexity in velocity structure and was simplified when applied to gravity modelling. The shapes of all modelled bodies above 0.5 km elevation and between 0 and 5 km on the profile line were influenced by the results of the seismic interpretation. The $1.7-$ (1.8) g/cm³ body is noticeably thickened for gravity modelling. This body incorporates the 0.8 and 1.2- km/s bodies of the seismic profile, probably grouping fine alluvium with highly fractured and weathered bedrock. The $2.10-$ to $2.25-g/cm^3$ density range chosen for the tuff bodies represents median values of many density values determined for the many tuff and flow units represented by these bodies (Carr, 1966). The lower contact of the northern tuff body follows from some preliminary magnetic modelling in the area (D. Jefferis, oral cmun., 1978). The southern structure copies published cross sections (McKay and Williams, 1964) and is a logical extension of exposed rock units in the Calico Hills and Little Skull Mountain to the south. One fault is modelled at -3.0 km, as suggested by surface geologic mapping and electrical studies.

The seven body model (fig. 11A) produces a very reasonable fit between calculated and observed gravity curves. The match is especially striking between the -9 to -4 and -1 to 9 km values on the horizontal scale. The discrepancy between -4 and -1 km may result from approximating a structually complicated hillside with simply shaped bodies. The overall fit is rather good and cannot be much altered within the drill hole, seismic, and surface geologic constraints discussed.

The model bodies contributing most to the calculated profile are those representing tuffaceous rocks with $2.10-g/cm³$ modelling densities. The northern and southern bodies produce gravity effects of about -28 and -7 mGals, respectively. For comparison, the two alluvium bodies ($\rho = 1.7$ g/cm³) produce effects less than 5- mGal. Thus, the lower boundaries of the tuff bodies are very important, for slight adjustments in their location produce sizable shifts in the gravity profile. Unfortunately, there is no dependable control at appropriate depths near the edge of this cross section, and one must resort to geologic interpretations. Without control of these critical contacts, a number of equally close gravity profile matches could be generated by elevating the lower tuff boundaries while simultaneously injecting a highdensity body near the center of the cross section, or by thickening the tuff and adding low-density bodies within the Eleana Formation. The injecting of one such body, a rhyolite dike, is warranted by mapped surface expressions of such a dike near the center of the profile. In figure 11A, this rhyolite dike is represented by the 2.40- g/m^3 density body.

Another possible variation in modelling is indicated in figure 11B. Here the bulk densities of the tuffs is 2.25 $g/m³$. In addition, the dike-size body is enlarged and located deeper to represent the top of a pluton or some other large intrusive body. In order to achieve a match between calculated and observed gravity, the lower contact of the northern tuff body needed to be steepened and deepened. As with the dike, the modelling program used a least squares match of the gravity curves to determine the density of the intrusive body. In this case the pluton was assigned a density of 3.00 g/cm^3 . This value, undoubtedly too great, is intended to represent an extreme of modeling, as is figure 11B in general. A large dense intrusive is required to compensate the denser, though thickened, tuffs modelled in figure 11B with respect to those modelled in figure 11A.

For perspective, the several variations of spheres and upright rectangular prisms indicated in figures 12 and 13, respectively, produce maximum gravity effects of 1 to 11 mGals. The rectangular prism modelled in figure 11a had a maximum amplitude of -1.2 mGal; that in figure 11B, a maximum amplitude of $+2.2$ mGal.

East-west profile interpretation

The second interpretive cross section (fig. 14) shown as B-B' on figures 8 and 9, runs due east-west just south of lat 36°52'30" N. It extends from the edges of Yucca Mountain in the west, across Forty Mile Canyon, through the Calico Hills, to Kiwi Mesa. Because no planned magnetic or gravity lines were established in this direction, randomly spaced gravity stations define the observed gravity curve. The gravity map, figure 8, indicates no striking correlations between surface geology and gravity values. Notably, no clear gradient defines the Calico Hills anomaly on its eastern flank.

Modelling drew heavily upon the first interpretive profiles (fig. 11) for depth of contacts and densities. All nonalluvial bodies in the first model are included in this cross section. A higher density, 2.3 $g/cm³$, for the tuffs near Kiwi Mesa, 2.3 g/m^3 , was thought appropriate (D. Ponce, written commun., 1980). Body shapes resemble those of the geologic interpretations of McKay and Williams (1964) with three proposed faults represented at -3.75 , -2.5, and 3.3 km on the distance scale. A rhyolitic dike of 2.62 $g/cm³$ density was included in the modelling to approximate the observed gravity values better. This higher density, when compared with profile A-A' partly reflects the latitude in choosing the density of such an intrusive body.

The overall fit between observed and calculated gravity is reasonable. Terrain corrections based on an incorrect reduction density of 2.67 g/cm^3 undoubtedly contribute to the mismatch at -6 km. The tuff must also thicken west of this value and beyond the area of interest of this study. To the east, the nearly constant observed gravity suggests interesting structures.

Geologic interpretation of the surface contact between argillite and tuff mapped near the 3-km mark in the cross-section suggests at least a 0.5 km thickness of tuff, as do gravity interpretations of the west edge of the Wahmonie/Skull Mountain horst (D. A. Ponce, written commun., 1980). Such tuff thicknesses would produce about -10 mGal of computed gravity in profile $B-B'$ yet variation in the observed values is no greater than 3 mGals. The interpretation of profile B-B' in figure 14 assumed less than 0.2-km thickness of tuff to accommodate the small variation observed. An increase in density

25

Figure 12.--Resulting gravity anomalies for spheres of 1.07- and 1.68 km diameter compared with observed gravity profile A-A'. Density contrasts of 0.1 g/cm³ and 2.0 g/cm³ are indicated for the larger sphere. Cnly the large sphere and large contrast produce an anomaly of significant amplitude to affect the modelling, and that effect is too great. The short vertical line at -2 km again represents the drill hole.

Figure 13.--Resulting gravity anomalies for upright rectangular prisms of 0.1-g/cm³ density contrast and 1.22- and 3.66-km thicknesses and infinite extent. Anomalies are small, though noticeable when compared with the observed gravity profile A-A'.

Figure 14.--Modelled interpretation of gravity profile B-B'. Bodies with densities identical to those in Figure 11 represent the same geologic units. The 2.3 g/cm³ body represents somewhat denser volcanic rocks near Kiwi Mesa; the 2.62-g/cm3 body is another example of an intrusive dike, viewed toward the wide cross section of the 2-1/2-dimensional body. The calculated anomaly at -6 km is produced by the topographic effect of Forty Mile Wash. The vertical dotted line represents the \mathcal{L}^{\pm} drill hole. Symbols as in figure 11.

to the eastward at depth seems more appropriate, yet unproven. Thermal alteration associated with the Wahmonie intrusive body could either massively or locally increase density. Alternatively, greater coherence and fewer fractures away from the Calico Hills dome could lower effective porosity and increase density.

Conclusions

A 770-m drill hole log and complete surface geologic mapping provide unusally good geologic control for the Calico Hills study area. In addition, a seismic refraction line provided independent constraints on density variation within the tp kilometer of the central Calico Hills. This control plus geologic interpretation of the structures at the edges of the study' area were used to produce a modelling profile through the Calico Hills trending N 10⁰ W. Interpretation of this profile suggests that the Calico Hills gravity anomaly can be entirely attributed to topography on the surface between the Paleozoic Eleana argillite and overlying Tertiary tuff. The 0.35- to 0.50 $g/cm³$ contrast at this boundary produces gravity variations of several milligals for vertical changes of several hundred meters in the depth of the boundary. Spherical and prismatic bodies near 0 m elevation (sea level) produce 2- to 11- mGal anomalies depending on density contrasts and the extent of the bodies. Only with precise control of the depth of basement rocks beneath Jackass Flats and Shoshone Mountain such as a drill hole or seismic soundings could provide, can the existence of intrusive bodies be excluded or confirmed. The east-west-trending modelling profile produced similar conclusions with one exception. As before, no deep high-density body was required under the Eleana outcrops. To the east, beneath Kiwi Mesa, however, high-density rocks at depth must balance the 2.1- to 2.3-g/cm³ tuff mapped at the surface to produce the lack of gravity gradient in this region. Those high-density rocks may be an intrusive body, marbleized or hydrothermally altered rocks, or more coherent tuff and argillite than the highly fractured rock immediate to Calico Hills. This study does define the domical shape of the Paleozoic rock/tuff boundary but produces no convincing evidence of an intrusive body, although its existence cannot be ruled out at this time.

References cited

- Barnes, D. F., Oliver, H. W., and Robbins, S. L., 1969, Standardization of gravimeter calibrations in the Geological Survey: EOS, v. 50, p. 526- 527.
- Cady, John W., 1980, Calculation of gravity and magnetic anomalies of finitelength right-polygonal prisms: Geophysics, v. 45 p. 1507-1512.
- Carr, W. J., 1966, Geology and Test Potential of Timber Mountain Caldera Area, Nevada: U.S. Geological Survey Technical Letter NTS-174.
- Diment, W. H., Healey, D. L., and Roller, J. C., 1960, Gravity and seismic exploration at the Nevada Test Site: U.S. Geological Survey Professional Paper 400B, p. 156.
- Hammer, Sigmund, 1939, Terrain corrections for gravimeter stations: Geophysics, v. 4, p. 184-194.
- Healey, D. L., and Miller, C. H., 1963, Gravity survey of the Gold Meadows stock, Nevada Test Site, Nye County, Nevada: U.S. Geological Survey Professional Paper 475-B, p. B64-B66.
- Healey, D. L., Wahl, R. R., and Currey, F. E., 1980a, Bouguer gravity map of Nevada, Goldfield and Mariposa sheets: Map 68, Nevada Bureau of Mines and Geology, scale 1:250,000. (In press).
- Healey, D. L., Wahl, R. R., and Oliver, H. W., 1980b, Complete Bouguer gravity map of Nevada, Death Valley sheet: Map 69, Nevada Bureau of Mines and geology, scale 1:250,000, (In press).
- International Association of Geodesy, 1971, Geodetic reference system 1967: International Association of Geodesy, Special Publication No. 3, 116 p.
- Maldonado, Florian, Muller, D. C., and Morrison, J. N., 1979, Preliminary geologic and geophysical data of the UE25a-3 exploratory drill hole, Nevada Test Site, Nevada: U.S. Geological Survey report USGS-1543-6, 47 p.; available only from U.S. Department of Commerce, National Technical Information Service, Springfield, VA 22161.
- McKay, E. J., and Williams, W. P., 1964, Jackass Flats Quadrangle, Nye County, Nevada: U.S. Geological Survey Quadrangle Map GQ-368, scale 1:24,000.
- Morelli, C. (Ed.), 1974, The International Gravity Standardization Net 1971: International Association of Geodesy, Special Publication No. 4, 194 p.
- Nettleton, Lewis L., 1976, Gravity and magnetics in oil prospecting: New York, McGraw-Hill, 464 p.
- Oliver, H. W., 1977, Gravity and magnetic investigations of the Sierra Nevada batholith, California: Geological Society of America, v. 88, p. 445-461.
- Oliver, H. W., ed., 1980, Interpretation of the gravity map of California and its continental margin: California Division of Mines and Geology Bulletin 205, 9 ilus., 45 p.
- Oliver, H. W., and Robbins, S. L., 1980, Bouguer gravity map of California, Fresno sheet: California Division of Mines and Geology Map, scale 1:250,000, 15 p. text.
- Orkild, P. P., and O'Connor, J. T., 1970, Topopah Spring Quadrangle, Nye County, Nevada: U.S. Geoloigcal Survey Quadrangle Map GQ-849, scale 1:24,000.
- Poole, F. G., Hauser, F. N., and Orkild, P. P., 1961, Eleana Formation of Nevada Test Site and vicinity, Nye County, Nevada: U.S. Geological Survey Professional Paper 424-D, p. D104-D111.
- Plouff, Donald, 1977, Preliminary documentation for a FORTRAN program to complete gravity terrain corrections based on topography digitzed on a geographic grid: U.S. Geological Survey Open-file Report 77-535, 45 p.
- Sass, J. H., Lachenbruch, Arthur H., and Mase, C. W., 1980, Analysis of Thermal Data from Drill Holes UE25a-3 and UE25a-1, Calico Hills and Yucca Mounain, Nevada Test Site: U.S. Geoloigcal Survey Open-File Report 80- 826, 25 p.
- Spengler, R. W., Maldanaldo, Florian, Weir, J. E., Jr., Hanna, W. F., and Dixon, G. L., 1979, Inventory of granitic masses in the state of Nevada: U.S. Geological Survey Open-file Report 79-235, 264 p. text and 7 plates.
- Wahl, R. R., 1969, An analysis of gravity data in Area 12, Nevada Test Site: U.S. Geological Survey Open-file Report 09-Area-12-12, 23 p. and 3 plates.

Appendix 1. Principal facts of the 339 gravity stations used in this study.

Explanation of tables:

- STATION An alphanumeric combination of up to eight characters for station identification.
- LAT. and LONG. Latitude and longitude of the station in degrees and minutes to the nearest hundredth of a minute.
- ELEV Station elevation, in feet, to the nearest tenth of a foot.
- ACC $-$ One letter and three digits indicating elevation, location, and reading accuracies, see Appendix A. The remaining columns contain gravity values in milligals, to the nearest hundredth of a milligal.
- OB GRAV $-$ Observed gravity relative to the prime base Merc, $g = 979,518.80$ mGal.
- FA Free air anomaly.
- SB Simple Bouguer anomaly for a crustal density of 2.67 gcm⁻³.
- NEAR- Terrain corrections to a distance of 0.59 km using cylindrical ring templates (Hammer, 1939) and field measurements.
- TC Camputer calculated total terrain corrections out to 166.7 km using digitized terrain values.
- CBA Complete Bouguer anomaly for crustal density of 2.67 gcm⁻³.

 \overline{a}

 \rightarrow \rightarrow \rightarrow

 \bar{z}

 $\ddot{}$

amputer terrain corrections carried from station to 166.7 kilometers jensities are 2.67 and 2.50. Density of 2.67 used for values in columns labelled TC, (NEAR). (NEAR) represents inner zone hand corrections, C represents computer corrections
Station Lat. Long. Elev Ob C.

 $\mathcal{L}^{\mathcal{L}}$

omputer terrain corrections carried from station to 166.7 kilometers Densities are 2.67. Density of 2.67 used for values in columns labelled TC, (NEAR). (NEAR) represents inner zone hand corrections, iC represents computer corrections Station Lat. Long. Elev Ob Grav ACC FA SB NEAR TC CBA Sta

 $\frac{1}{\mu}$

 $\ddot{}$

LOCATION CODE (1st digit)

Code Station Location Method (vertical and horizontal) A) Survey marks (vertical) - Topo Maps (horizontal) 1) Base plate directly on bench mark B----------- a) USGS or USC&GS level-line bench mark. M----------- b) Level line bench marks other than a) such as USCE, BPR, CDH, private companies, etc. V----------- c) Vertical angle (VABM) bench marks. 2) Base plate near bench mark. N---------- a) USGS or USC&GS level-line bench mark. E----------- b) Level-line bench marks other than a) such as USCE, BPR H----------- c) Vertical angle (VABM) bench marks. P------- 3) Base plate on or near other reference marks (such as stakes, paint, etc.) that have been surveyed by the group doing the gravity survey or other known people. X ------- 4) On or near Section Corners, $1/4$ section marks, $1/8$ section markers, and other property boundary markers. D------- 5) Destroyed or not found bench or reference marks. B) Map locations (vertical and horizontal) F------- 1) Black spot elevations - field checked. G ------- 2) Brown spot elevations and elevations taken off original manuscripts . not field checked. W------- 3) Blue lake elevations. R------- 4) Lake or reservoir elevations determined from leveling to bench marks, and water level is determined from gauging stations. S **------** 5) Sea level elevations. C------- 6) Contour line interpolation. ^Q----- 7) River gradient interpolation. C) Air Photographs (vertical and horizontal) T------- 1) Elevations determined by U.S. Geological Survey Topographic Division by Kelsh plotter or least squares computer system. …------- 2) Elevations determined by other groups by Kelsh plotter or least squares computer system. L------- 3) Elevations determined by Laser methods. 3 **------** 4) Elevations determined by other methods. D) Altimetry (vertical) - Topo Maps (horizontal) A------- 1) Good control (Leap frog, double loop, two or more altimeters, etc.) Y------- 2) Poor control. E) Special sources z------- 1) Elevations determined by methods such as mobile elevation recorders - horizontal control from Topo Maps. I------- 2) Other special sources. F) Unknown Elevation Sources U------- 1) Elevation data sources unknown (this would include reference marks with unknown ties).

SUBJECTIVE NUMERICAL ACCURACY CODE FOR 2nd, 3rd & 4th DIGITS Elevation Accuracy Code (2nd digit) Relative to 1929 USC&GS mean sea level datum

 \bar{z}

large **U** 5 \pm .2 Data from loops with closure errors this 6 **+** .5 **H U n n U n n U** $7 + 1.0$ 8 $\frac{1}{2}$.0 **ⁿn U** a **H U** *U* **^N** \bullet ^an **U U** *U U U* **U** $9 + 5.0$ *** U U** *U* **U U** *^U***U** 0×5.0

J.

 $\ddot{}$

Totals

 $\label{eq:2.1} \frac{1}{\sqrt{2}}\int_{\mathbb{R}^3}\frac{1}{\sqrt{2}}\left(\frac{1}{\sqrt{2}}\right)^2\frac{1}{\sqrt{2}}\left(\frac{1}{\sqrt{2}}\right)^2\frac{1}{\sqrt{2}}\left(\frac{1}{\sqrt{2}}\right)^2\frac{1}{\sqrt{2}}\left(\frac{1}{\sqrt{2}}\right)^2.$ $\mathcal{L}^{\text{max}}_{\text{max}}$, where $\mathcal{L}^{\text{max}}_{\text{max}}$

 $\mathcal{L}^{\text{max}}_{\text{max}}$, where $\mathcal{L}^{\text{max}}_{\text{max}}$

 $\mathcal{L}(\mathcal{L}^{\mathcal{L}})$ and $\mathcal{L}(\mathcal{L}^{\mathcal{L}})$ and $\mathcal{L}(\mathcal{L}^{\mathcal{L}})$

 $\label{eq:2.1} \mathcal{L}(\mathcal{L}^{\text{max}}_{\mathcal{L}^{\text{max}}_{\mathcal{L}^{\text{max}}_{\mathcal{L}^{\text{max}}_{\mathcal{L}^{\text{max}}_{\mathcal{L}^{\text{max}}_{\mathcal{L}^{\text{max}}_{\mathcal{L}^{\text{max}}_{\mathcal{L}^{\text{max}}_{\mathcal{L}^{\text{max}}_{\mathcal{L}^{\text{max}}_{\mathcal{L}^{\text{max}}_{\mathcal{L}^{\text{max}}_{\mathcal{L}^{\text{max}}_{\mathcal{L}^{\text{max}}_{\mathcal{L}^{\text{$

 $\mathcal{L}^{\mathcal{L}}(\mathcal{L}^{\mathcal{L}})$ and $\mathcal{L}^{\mathcal{L}}(\mathcal{L}^{\mathcal{L}})$ and $\mathcal{L}^{\mathcal{L}}(\mathcal{L}^{\mathcal{L}})$

 \sim $\hat{\mathcal{A}}$

l, $\label{eq:1} \frac{1}{\sqrt{2\pi}}\int_{0}^{\infty}\frac{1}{\sqrt{2\pi}}\left(\frac{1}{\sqrt{2\pi}}\right)^{2}d\theta.$

 $\bar{\mathcal{A}}$