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Fracture and Matrix Hydrologic Characteristics of Tuffaceous Materials from Yucca Mountain, Nye County, Nevada

R. R. Peters, E. A. Klavetter, I. J. Hall, S. C. Blair, P. R. Heller, G. W. Gee

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FRACTURE AND MATRIX HYDROLOGIC CHARACTERISTICS OF TUFFACEOUS MATERIALS FROM YUCCA MOUNTAIN, NYE COUNTY, NEVADA

Sandia National Laboratories

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ABSTRACT

The geological formations in the unsaturated zone at Yucca Mountain, on and adjacent to the Nevada Test Site (NTS), are currently being studied for consideration as the host for a radioactive-waste repository; the U.S. Department of Energy is carrying out these studies through the Nevada Nuclear Waste Storage Investigations project. The formations are composed of tuffaceous (tuff) materials that must be evaluated to estimate the rate at which radionuclides would migrate to the accessible environment. According to the available evidence, the flux of water in the unsaturated zone beneath the Yucca Mountain site is low; quantifying such low flow rates through direct measurements is difficult. To help provide data that can be used to assess unsaturated flow, Pacific Northwest Laboratory (PNL), under contract to Sandia National Laboratories (SNL), performed hydrologic tests on tuffaceous samples from 48 different locations in Yucca Mountain.

The primary purpose of this document is to provide a compilation of the testing procedures used and the hydrologic data obtained. The results of the testing indicate the following:

- There are wide variations in water-retention characteristics for the tuffaceous materials tested. Nonwelded, nonzeolitized samples tended to drain readily, while welded samples and nonwelded, zeolitized samples stayed near saturation at relatively high suction head values (10 to 50 m). The water retention characteristic curves (pressure vs saturation), in general, were replicated by mercury-intrusion tests on similar samples.
- 2. The measured saturated hydraulic conductivities for welded tuff samples were low, ranging from 10^{-10} to 10^{-14} m/s. Most of the nonwelded, zeolitized samples exhibited low conductivities similar to those of welded samples. The nonwelded, nonzeolitized samples exhibited conductivities ranging from 10^{-6} to 10^{-10} m/s, values that compare with silts and clay-type soils.
- 3. The fracture saturated conductivity was significantly higher than the matrix conductivity on all samples tested and flow through all fractured and unfractured samples was reduced at elevated effective pressure. The degree of welding of the matrix and the fracture surface characteristics influenced the response of fractured samples to elevated pressure.

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This report contains the entire set of psychrometer measurements of desaturation curves for tuffs from Yucca Mountain as well as a substantial number of saturated conductivity measurements. It is, to the best of the authors' knowledge, the most complete set of published, unsaturated, hydrologic data for any hard rock.

ACKNOWLEDGMENTS

The authors wish to acknowledge the efforts of other individuals who were instrumental to this study. B. M. Schwartz of Sandia National Laboratories (SNL) contributed to document and procedures review. J. H. Gauthier (SNL) provided valuable support for much of the graphics. C. E. Haines, of Pacific Northwest Laboratories (PNL), provided technical assistance in parameter measurement.

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INTRODUCTION

The geological formations in the unsaturated zone at Yucca Mountain, on and adjacent to the Nevada Test Site (NTS), are currently being studied for consideration as the host for a radioactive-waste repository; the U.S. Department of Energy is carrying out these studies through the Nevada Nuclear Waste Storage Investigations project. Because water flow through unsaturated rock is the principal mechanism for the transport of soluble radionuclides and other contaminants from a repository to the biosphere at an arid site, characterization of the hydrologic system within Yucca Mountain is an essential portion of the evaluation of this location. At Yucca Mountain, rates of flow in the unsaturated zone have been shown to be low, but must be quantified or bounded if the site is to be characterized for possible use as a waste repository. However, the quantification of low flow rates deep within any arid formation is difficult, and few direct measurements of unsaturated flow rates have been made (Winograd, 1981). An alternative to direct measurement of the low flow rates in Yucca Mountain is to estimate the flow rate indirectly by using models that account for the dependence of hydraulic properties on water content. Unsaturated-flow codes, combined with the hydrologic properties of tuffaceous (tuff) materials at Yucca Mountain under unsaturated conditions, can be used to predict pressure-head profiles and water content as a function of depth and water flux. These predicted pressure profiles can be compared with pressure profiles and water content determined by field measurements to determine the flow rate within Yucca Mountain. Modeling flow in the unsaturated zone is currently receiving considerable attention. A good introduction to this topic is provided by several authors including Narasimhan (1982).

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In a previous contract, Pacific Northwest Laboratory (PNL) provided Sandia National Laboratories (SNL) with data on the relationship between water content and pressure head for tuffs tested in the pressure-head range of 0 to -30.000 m.^(a) However, because of limitations of the test apparatus, little water-retention data were collected in the range of -10 to -1000 m of pressure head (which is equivalent to negative suction head). To determine the hydraulic properties in the range of -10 to -1000 m, tests were run on 19 samples of tuff taken from drill hole USW GU-3 and 29 samples taken from drill hole USW G-4 on the NTS. Direct measurement of unsaturated conductivities of tuff within a reasonable time period is currently impossible because of limitations of test equipment. However, theoretical models of unsaturated flow (e.g., Mualem, 1976) allow one to calculate the unsaturated conductivity if water-retention data and saturated-conductivity data are available. Four samples of unfractured tuff from drill hole USW GU-3 were tested at elevated confining pressures to determine the saturated conductivity, and five fractured samples taken from drill hole USW G-4 were tested at elevated confining pressures to determine saturated conductivity. This report discusses the methods used to characterize the hydraulic properties of the tuffs, documents the results of those tests, and provides analysis of those results.

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⁽a) Gee, G. W., P. R. Heller, and M. B. Dodson. 1982. "Laboratory Report on the Unsaturated Flow Characteristics of Core Samples from Nevada Test Site Well-USW GU-3 (NTS)." Letter Report, October 1982, Pacific Northwest Laboratory, Richland, Washington.

MATERIALS AND METHODS

Tuffaceous core materials for testing on this project were taken from drill holes USW GU-3 and USW G-4 in Yucca Mountain. Table 1 presents a summary of the functional stratigraphies found in these two drill holes. The functional stratigraphies were developed by Ortiz (1984) and are based on logs of the drill holes and mineralogical work. Table 2 presents the PNL code, drill hole depth, and functional unit for core samples tested from drill holes USW GU-3 and USW G-4, respectively.

Three types of tests were performed by PNL. They include 1) waterretention tests, 2) unconfined, saturated hydraulic conductivity tests, and 3) confined, saturated hydraulic conductivity tests. Mercury-intrusion tests were performed by Micromeritics^(a) so that water-retention curves determined by the mercury intrusion measurements could be compared with those determined by the psychrometric measurements made by PNL.

All 48 samples selected by SNL were tested for water-retention characteristics. The tests were run on cylinders 1.4 x 1.2 cm (diameter x length) that were subcored from the original core samples. The water solution used in testing the hydraulic properties was of low ionic strength and had a composition similar to that of ground water sampled near the test well. Table 3 lists the chemical composition of the test solution.

Unconfined saturated-conductivity tests were also run on the 1.4 x 1.2 cm cylinders. Saturated conductivities at elevated confining pressures were measured on the fractured and unfractured samples listed in Table 4.

(a) Micromeritics Instr. Corp., Norcross, Georgia 30093
 Tests done on Auto-Pore 9200 Porosimeter to 60,000 psi.

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TABLE 1. Description of Units

<u>Unit</u>	GU-3 Depth(a) (ft)	G-4 Depth (ft)	Hydrologic Unit	Description
I-A	0-343	22-118	Tiva Canyon welded unit	Moderately to densely welded, devitrified ash-flow tuff in the Tiva Canyon Member of the Paintbrush Tuff.
I-B	343-430	118-243	Paintbrush nonwelded unit	Partially welded to non- welded, vitric and occas- ionally devitrified tuffs of the Tiva Canyon, Yucca Mountain, Pah Canyon, and Topopah Spring Members of the Paintbrush Tuff.
II-L	430-690	243-670	Topopah Spring welded unit	Moderately to densely welded, devitrified zone of the Topopah Spring Member of the Paintbrush Tuff that con- tains more than approximately 10% by volume of vugs.
II-NL	690-1187	670-1293	Topopah Spring welded unit	Moderately to densely welded, devitrified zone of the Topopah Spring Member of the Paintbrush Tuff that con- tains less than approximately 10% by volume of vugs. This is the potential repository unit.
III	1187-1269	1293-1345	Basal Vitro- phyre of the Topopah Spring welded unit	Basal Vitrophyre of the Topopah Spring Member.
IV- <u>A</u> -v(b)	1269-1507	13451360	Vitric Calico Hills non- welded unit	Nonwelded ashflows, bedded and reworked tuffs, vitric and primarily nonzeolitized Topopah Spring Member and/or the Calico Hills.

⁽a) Sample depths reported in feet are primary data supplied by Sandia National Laboratories.

⁽b) The lower case "v" or "z" in a unit number (e.g., IV-A-v) indicates whether the unit is vitric or water-induced zeolitization has occurred.

TABLE 1. Description of Units (continued)

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<u>Unit</u>	GU-3 Depth(a) (ft)	G-4 Depth (ft)	Hydrologic Unit	Description
IV-B-v	1507-1560	Missing	Vitric Calico Hills nonwelded unit	Basal, bedded and reworked zone of the vitric tuffs and tuffaceous sandstones of the Calico Hills.
IV-C-v	1560-1601	Missing	Vitric Calico Hills non- welded unit	Upper vitric zone of the Prow Pass Member of the Crater Flat Tuff.
IV-A-z	Missing	1360-1706	Zeolitized Calico Hills nonwelded unit	Nonwelded ashflows, bedded and reworked tuffs, primarily zeolitized, from the Topopah Spring Member and/or the Calico Hills.
IV-B-z	Missing	1706-1761	Zeolitized Calico Hills nonwelded unit	Basal, bedded and reworked zone of the zeolitized tuffs and tuffaceous sandstones of the Calico Hills.
IV-C-z	Missing	1761-1792	Zeolitized Calico Hills nonwelded unit	Upper zeolitized zone of the Prow Pass Member of the Crater Flat Tuff.
V	1601-1746	1792-1960	Prow Pass welded unit	Moderately welded, devitri- fied zone of the Prow Pass Member of the Crater Flat Tuff.
VI	1746-2069	1960-2250	Crater Flat nonwelded unit	Zeolitized portions of the lower Prow Pass Member and of the upper Bullfrog Member of the Crater Flat Tuff.
VII	2069-2508	2250-2698	Bullfrog welded unit	Welded devitrified zone of the Bullfrog Member of the Crater Flat Tuff.

Drill Hole USW G-4	PNL Sample Code	Depth Below Surface, ft	Unit
	G4-1	43	I-A
	G4-2	124	I-B
	G4-3	208	I-B
	G4-4	247	II-L
	G4-5	864 (dark section)	II-NL
	G4-24	864 (light section)	II-NL
	G4-6	1158	II-NL
	G4-1F	1215	II-NL
	G4-7	1256	II-NL
	G4-2F	1278	II-NL
	G4-8	1299	111
	G4-9	1324	111
	G4-3F	1359	IV-A-v
	G4-10	1405	IV-A-z
	G4-11	1548	IV-A-z
	G4-4F	1551	IV-A-z
	G4-12	1686	IV-A-z
	G4-13	1728	IV-B-Z
	G4-14	1737	IV-B-z
	G4-15	1769	IV-C-z
	G4-16	1778	IV-C-z
	G4-5F	1778	IV-C-z
	G4-17	1787	IV-C-z
	G4-18	1899	v
	G4-19	2006	vi
	G4-20	2101	VI
	G4-20 G4-21	2401	VII
		2401	VII
	G4-22		ATT
	G4-23	(a)	

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TABLE 2. Information Concerning Samples Taken fromDrill Holes USW G-4 and USW GU-3

Drill Hole USW GU-3	PNL Sample	Depth Below Surface, ft	Unit
	GU3-1	82	T A
			I-A
	GU3-2	120	I-A
	GU3-3	155	I-A
	GU3-4	257	I-A
	GU3-5	316	I-A
	GU36	374	I-B
	GU3-7	378	I-B
	GU3-8	397	I-B
	GU3-9	1132	II-NL
	GU3-10	1197	111
	GU3-11	1245	111
	GU3-12	1311	IV-A-v
	GU3-13	1331	IV-A-V
	GU3-14	1440	IV-A-v
	GU3-15	1499	IV-A-v
	GU3-16	1555	IV-B-V
	GU3-17	1628	v
	GU3-18	1680	v
	GU3-19	1730	V

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(a) Sandia Sample Code 1010A (commercial clinoptilolite).

TABLE 3. Constituent Concentration of J-13 Well Water^(a) Used in Saturation and Conductivity Tests

<u>Constituent</u>	_mg/L
A1	<0.1
Si	31.35
Fe	0.05
Mg	1.98
Ca	12.93
Sr	0.04
Li	0.04
Na	48.00
K	4.69
C03	71.70
HCO3	0.0
C1	3.7
SOA	9.3
NO3	1.5
PO4	0.02

(a) pH = 8.10

TABLE 4. Samples Tested for Conductivity at Elevated Confining Pressure

Unfractured Samples	Fractured Samples			
GU32	G4-1F			
GU33	G4-2F			
GU3-11	G4-3F			
GU3-15	G4-4F			
	G4-5F			

Unfractured samples tested at elevated confining pressure were wafers 5.4 \times 1.9 cm (diameter x length).

The criteria for selecting the fractured core samples were

- That the fracture be natural, or at least not be obviously caused by the drilling process.
- 2. That the orientation of the fracture allow testing.

The second criterion indicated that it would be advantageous to look for core containing a single fracture oriented approximately parallel to the core axis. A fracture with this orientation is a member of the vertical fracture set, which is thought to be the predominant fracture set at Yucca Mountain (Scott, 1983). Therefore, all the fractured core selected for testing contained a single fracture oriented approximately parallel to the core axis; the fracture divided the sample into two discrete pieces. Fractured samples were right circular cylinders 6 x 8 or 6 x 15 cm (diameter x length). Photographs of fractured core samples from depths of 1215 (G4-1F) and 1778 (G4-5F) ft in drill hole USW G-4 are presented in Figure 1. Table 5 gives a brief description of each fractured sample tested.

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FIGURE 1. Photographs of Two Fractured Core Samples from Drill Hole USW G-4. Fracture surfaces for sample G4-5F mate extremely well while fracture surfaces for G4-1F mate poorly.

Sample	Sample Depth, ft	Description	Length,	Diameter,	Crack Width, 	Fracture Surface	Fracture Mating	Comments
G4-1F	1215	Densely welded tuff, Brown color	15.29	6.08	6.05	Rough, but planar surface Oriented parallel to core axis	Роог	Asperities on surface oriented parallel to core axis Fracture has poor fit and largest aperture Nissing chips along fracture surface
G4-2 F	1278	Densely welded tuff, Brown color	7.67	6.10	5.80	Smooth, curved surface Subparallel to axis	Fair	Obvious voids along surface, flow channels visible, one large chip loose at end of core
G4-3F	1359	Densely welded tuff, Dark grey	15.28	6.07	6.05	Smooth, planar surface Parallel to core axis	Good	
G4-4P	1551	Moderately consoli- dated, Whitish matrix with angular black fragments	7.72	5.99	5.89	Undulating surface Subparallel to core axis	Fair	Lithic fragments, (~0.5-0.8 cm dia.) present throught rock
G4-5F	1778	Moderately consoli- dated tuff, Pink in color	15.27	6.10	4.95	Planar surfac e Parallel to core axis	Excellent	Fracture surfaces mate extremely well

TABLE 5. Characterization of Fractured Tuff Cores from Hole USW G-4

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This section describes the testing methods used to obtain data concerning the functional relationship between sample saturation and pressure head. This relationship is one of the two "characteristic curves" required by unsaturated hydrology models. The data, when combined with the saturated hydraulic conductivity, can be used to estimate the functional relationship between sample unsaturated conductivity and pressure head--the second characteristic curve required by the models. (The calculation of the unsaturated conductivity is discussed in the Results section.)

Two methods were used to obtain the water-retention curve. The first is a direct method using a thermocouple psychrometer. The second is a somewhat indirect one using the results of mercury-intrusion tests.

All samples were vacuum saturated before testing with the thermocouple psychrometer. The samples were placed in a vacuum chamber, which was then filled with J-13 well water to just cover the samples. The chamber was pumped for 45 minutes and the samples were allowed to saturate inside the closed chamber for 24 hours before testing. Grain density and bulk density were measured for each sample to determine porosity. Samples were tested using a thermocouple psychrometer, yielding data for water retention as a function of pressure.

Psychrometer Tests

A commercial thermocouple psychrometer^(a) was used to measure water potential (used to estimate the suction head, h,) of the matrix water in the range from 10 to 10,000 m (Figure 2). The samples were wiped clean of any free water, placed in a preweighed bottle, weighed, and transferred to a small

-12-

⁽a) Decagon, Inc., Pullman, Washington.

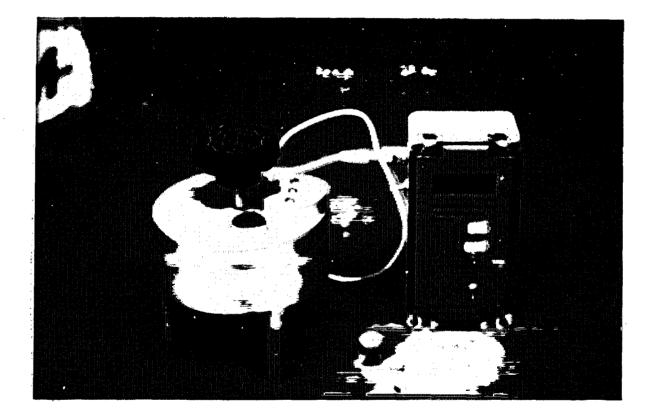


FIGURE 2. Thermocouple Psychrometer Sample Changer, Small Chamber, and Direct Readout Unit cup that fit into a vapor-tight chamber in the thermocouple psychrometer. After approximately 1 hour, during which time the samples and chamber were allowed to equilibrate thermally, the relative humidity was measured with a miniature thermocouple psychrometer. The samples were removed, dried slightly (for about 30 seconds) in a microwave oven and allowed to cool; the process was then repeated. Moisture content was determined gravimetrically, and the water potential was determined from the measured humidity.

Microwave drying (Gee and Dodson, 1981) was used to desaturate the samples in a stepwise fashion because it saved time and dried the samples more uniformly than drying in a conventional electric oven. Some welded samples did not dry completely using this technique, apparently because water was trapped in the pores. Other samples were incompletely saturated and dried too quickly. These latter samples were then resaturated and the entire process repeated. In all cases, heating in a conventional electric oven was used to determine the final oven-dry weight. Samples were dried at 105°C for a minimum of 24 hr before weighing.

Humidity readings were taken until the sample reached a predesignated suction head, usually between 1000 and 3000 m. The data points corresponding to suction heads greater than 3000 m were obtained by allowing the samples to air dry and weighing the samples. An Assmann psychrometer^(a) was used to measure the relative humidity of the air and, hence the sample. For all humidity measurements three subsamples for each sample depth (e.g., G4-1) in USW G-4 and two subsamples for each sample depth in USW GU-3 were tested to determine the homogeneity of subsamples as well as the consistency of tests, and to gather more data in the pressure range where the sample water content changed quickly.

(a)Weather Measure Corp., Sacramento, California.

A standard salt solution (0.5 M NaCl) with a water potential of 233 m at 25°C was used for calibrating the psychrometer. The psychrometer was considered to be functioning correctly when the measured value of water potential of the salt solution was within 10 m of the standard value. This test result also indicated that thermal equilibrium was being maintained in the sample chamber, since a sample temperature control of approximately 0.01°C is required to attain reproducibility of 10 m (Brown and Van Haveren, 1972). The inside walls of the sample cup were coated with paraffin to prevent water from adsorbing on the cup walls.

Estimation of Suction Head From Psychrometric Measurements

The psychrometer measures the wet-bulb depression (humidity) in a closed chamber directly above the sample using a miniature thermocouple connected to a microvolt sensing unit. The equilibrium vapor pressure and relative humidity can be related directly to the water potential of the water in the sample (Campbell, 1977). The relationship between the total water potential and relative humidity can be written as

$$\Psi = \frac{-RT}{M} \ln (RH/100)$$
 (1)

where ψ = total water potential

- R = universal gas constant
- T = Kelvin temperature
- M = molecular weight of water
- RH = relative humidity (percent).

Numerous references (e.g., Hillel, 1981; Papendick and Campbell 1980) detail the relationship of total water potential to individual component potentials and to sample suction. The total water potential of the water in the sample is the sum of the individual water-potential components as follows:

$$\psi = \psi_{\text{os}} + \psi_{\text{m}} + \psi_{\text{g}} + \psi_{\text{p}} + \psi_{\text{ov}}$$
(2)

where ψ_{os} , ψ_{m} , ψ_{g} , ψ_{p} , ψ_{ov} are osmotic, matrix, gravitational, pressure and overburden potential components, respectively.

For an unsaturated tuffaceous rock sample at atmospheric pressure, the major components of the total potential are ψ_{os} and ψ_{m} . Because the J-13 well water used in the tests was low in salts, it was assumed that ψ_{os} was negligible; hence, ψ is approximately equal in magnitude to the matrix potential, ψ_{m} . The value of ψ or ψ_{m} can be expressed either in terms of pressure units (Pascals or bars) or head units (m) (Campbell and Van Schilfgaarde, 1981). The suction head, h, is equal to $-\psi_{m}$ and represents the water head created by capillary and surface adsorptive forces of the matrix material. The water potential readings (negative bars) from the psychrometer were converted to suction head values (m) by multiplying by -10.2.

Mercury-Intrusion Tests

Mercury intrusion tests were performed on samples from USW G-4. Sample size was approximately 1.2 x 2.0 cm (diameter x length). The Micromeritics AutoPore 9200 testing apparatus was used with standard testing procedures. The sample and chamber were evacuated before testing began. The pressure range during the mercury-intrusion testing was 1.5 to 60,000 psi. The approximate range of pore sizes corresponding to this pressure range is 120 to

-16-

0.003 microns based on capillary bundle theory (e.g., Hillel, 1982). The primary result of these tests is the measured relationship between pressure and intruded mercury volume. The intruded mercury volume data is converted to mercury-saturaton data by dividing each data point by the maximum value of intruded mercury volume. The manner in which the mercury-saturation data points are used is explained below.

The adjustment for the differences in the properties of mercury and water was based on capillary bundle theory (e.g., see Hillel, 1982) with the suction head for water calculated as:

$$h_{w} = \frac{P_{Hg}}{\rho_{w}g} \frac{\sigma_{w} \cos(\gamma_{w})}{\sigma_{Hg} \cos(\gamma_{Hg})}$$
(3)

where

P = pressure

 σ = surface tension between the fluid and tuff

 γ = contact angle between the fluid and tuff.

 $\rho_{...}$ = density of water

g = acceleration of gravity

The "w" and "Hg" subscripts refer to water and mercury, respectively. The specific values used for sigma and gamma were estimates based on information contained in a variety of sources (e.g., Hillel, 1982) and are listed below.

$$\sigma_{W} = 72 \text{ dynes/cm} \qquad \gamma_{W} = 15^{\circ}$$

$$\sigma_{Hg} = 484 \text{ dynes/cm} \qquad \gamma_{Hg} = 130^{\circ}.$$

The water saturation corresponding to this adjusted pressure is 1 minus the mercury saturation determined in the intrusion tests. Thus, at zero pressure the mercury saturation of the sample is zero and the corresponding water saturation is 100%. At the maximum recorded pressure, corresponding to an h_w of about 1.0 x 10⁴ m of water, the mercury content of the sample is a maximum, the mercury saturation is assumed to be 100%, and the corresponding water saturation is 0%.

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TESTING METHODS--SATURATED HYDRAULIC CONDUCTIVITY

The saturated hydraulic conductivity testing consists of

- Unconfined matrix-conductivity testing of the small samples used in the water-retention tests and of larger, core samples.
- 2. Confined, matrix-conductivity testing of core samples.
- 3. Confined, conductivity testing of fractured core samples.

All samples were saturated in the manner described previously.

Unconfined Matrix Testing

Unconfined, saturated hydraulic conductivity was measured by a constant head method. This method measures conductivity by determining the time necessary for a given amount of liquid at room temperature (20 to 22°C) to pass through the core under a fixed pressure gradient. All 29 samples from drill hole USW G-4 were tested using the Ruska permeameter; ^(a) for about half of the samples more than one subsample was tested. Cross-sectional samples from drill hole USW GU-3 and six samples from USW G-4 were tested using a permeameter designed to accommodate the core samples.

Ruska Permeameter Tests

Tuff matrix samples were cut to size $(1.4 \times 1.2 \text{ cm})$, vacuum saturated, and inserted into the core holder. The core holder and buret above the sample were filled with test liquid to a level slightly above the upper index mark. A 2-bar (20.4-m) pressure gradient was applied on the test liquid from the top of the buret. The time necessary for the liquid level in the buret to drop

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⁽a) Model No. 1013-801-00, Ruska Instr. Corp., Houston, Texas.

from the upper to the lower index mark was determined and the saturated hydraulic conductivity calculated using Darcy's law for laminar flow in a right circular cylinder:

$$K = (Q/At)(L/\Delta H)$$
(4)

where K = saturated hydraulic conductivity (m/s)

Q = volume of water passing through the sample (m³)

- A = cross-sectional area of the sample (m^2)
- t = time (seconds)
- L = length of sample (m)
- ΔH = hydraulic head difference (m).

Note: At 20°C, with water's density at 1 g/cm³, the relationship between the permeability unit of "darcy" and the conductivity unit of "m/s" is 1 darcy = 1.02×10^{-5} m/s.

Core Permeameter Tests

Because the cores were not uniform, cross sections were taken from selected core samples to represent the heterogeneity of the entire core cross section. Cross sections in the form of wafers 6 x 2 cm (diameter x length) were taken from six of the core samples from drill hole USW G-4, (G4-5, G4-11, G4-13, G4-14, G4-17, and G4-18). The dimensions of the USW GU-3 samples were 6 x 5 cm. These samples were positioned in a specially built, rigid-wall, plastic permeameter and sealed in place with a silicone rubber compound to prevent water flow between the sample and the permeameter wall. The samples

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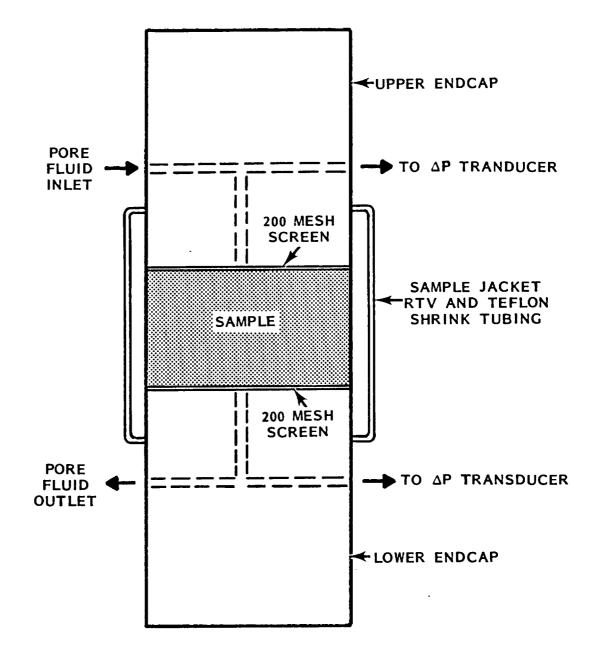
were then vacuum saturated and allowed to soak a minimum of 24 hr before testing. For these constant-head tests, elevated pressures up to 3 bars (30.6 m equivalent head) were applied to the permeant solution (J-13 well water). The effluent from each sample and the time required to collect it were recorded, and Equation 4 was used to determine the conductivity of the sample.

Confined Matrix Testing

This section describes conductivity tests performed on relatively thin (1.9 cm) samples of unfractured tuff under saturated conditions and confining stresses. The purpose of these measurements was to measure hydraulic conductivity under conditions that simulated the overburden stress encountered at depth. The sample was placed between two metal endcaps and coated with silicon rubber and shrink tubing as shown in Figure 3. The sample stack was then assembled into the base plate of the pressure vessel and vacuum saturated with J-13 well water. The base plate was then raised into a pressure vessel. The high-pressure apparatus is shown pictorially in Figure 4 and is described in detail by Blair and Stottlemyre (1981).

To determine the flow through unfractured tuff, the confining fluid pressure was raised to 100 bars and upstream pore pressure maintained at 50 bars. A high-pressure positive-displacement pump was used to supply pore fluid to the sample at a constant rate and a gas-backed accumulator was used to maintain a constant pressure on the downstream pore fluid reservoir. A differential pressure transducer was used to monitor pressure drop across the sample. The conductivity test apparatus is shown schematically in Figure 5a. The fluid supply rate was adjusted until steady-state flow was obtained. This

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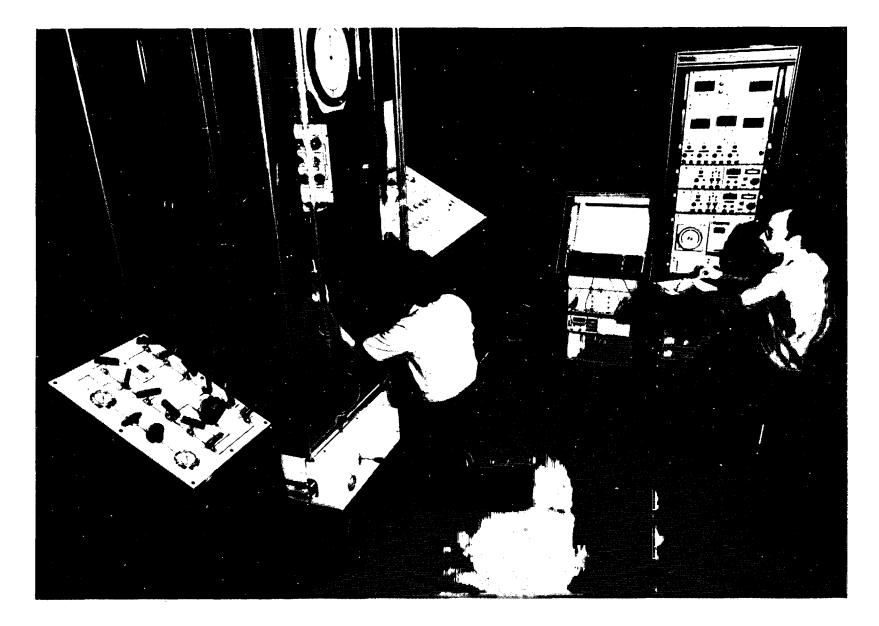


FIGURE 4. High-Pressure Testing Apparatus

was determined by monitoring upstream and downstream pressure as well as pressure drop across the sample. Steady-state flow was assumed when all pressure values (confining fluid, pore fluid inlet, etc) remained constant over a reasonable time period (on the order of minutes to hours). The exact length of time required to reach steady-state is dependent on many factors with the major one being the sample's saturated conductivity. Five different parameters were digitally recorded during the tests. These parameters, along with their estimated accuracy, are listed in Table 6. An analog chart recorder was also used to monitor various parameters throughout the testing.

Saturated conductivity measurements at increasing effective pressures (stresses) were conducted by increasing the confining fluid pressure and then reducing the fluid supply rate to attain a desired effective pressure. The saturated conductivity was calculated using Darcy's law for a laminar flow (Equation 4).

The effective pressure is a measure of the average confining stress on the sample and is calculated as

$$P_{eff} = P_{c} - \frac{[P_{i} + P_{o}]}{2}$$
(5)

where

P_c = confining fluid pressure
P_i = pore fluid pressure at sample inlet
P_o = pore fluid pressure at sample outlet

The effective pressure applied to a sample in the testing apparatus is approximately equivalent to the in situ confining stress.

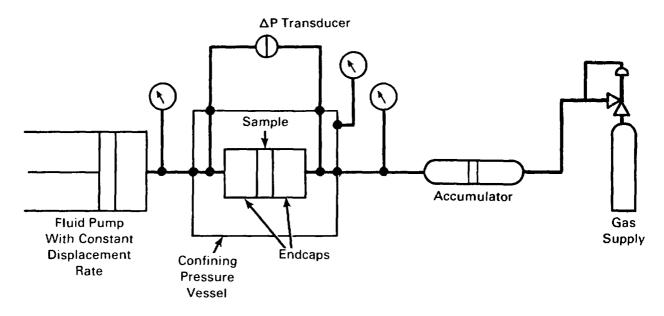
- 23-

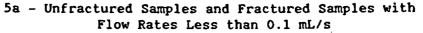
Confined Fracture Testing

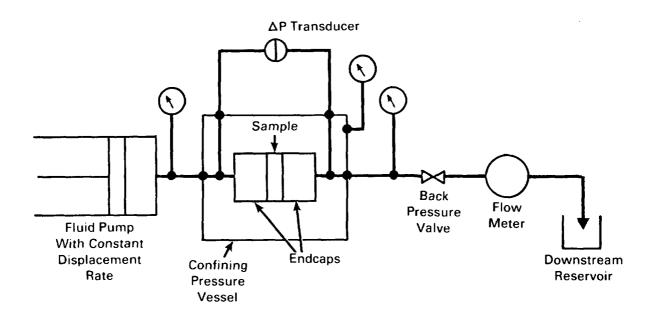
This section describes the saturated hydraulic conductivity testing performed on large (6 x 8 or 6 x 15 cm long) cylindrical samples of fractured tuff. Dimensions of each sample were measured, the fracture surfaces were examined, and the characteristics were documented (see Table 5). The sample was prepared in the same manner as that used for the saturated hydraulic conductivity testing of matrix samples under confining stress (discussed in the previous section).

To test fractured samples, the confining pressure and pore pressure were initially raised to 35 and 30 bars, respectively. The pore pressure was kept at approximately 30 bars throughout all tests, while the confining pressure was varied over a range of 35 to 150 bars. A pressure differential was introduced across the sample and flow through the sample was measured using one of two methods. For samples accommodating high flow rates (greater than 0.1 mL/s), flow was measured using a flow meter downstream of the sample, and back pressure was maintained through the use of a metering valve. This configuration is shown schematically in Figure 5b. For samples that were less permeable, the system was configured as shown in Figure 5a, and flow rate was calculated from piston displacement in the pore-fluid supply pump. In both methods, a differential pressure transducer was used to measure pressure drop across the sample. Data for confining pressure, upstream and downstream fluid pressure, differential pressure, pump displacement, flow rate, and temperature were recorded digitally. The accuracy of measurement of all parameters. except temperature, is listed in Table 6 (Blair and Stottlemyre, 1981). Sample temperature was that of the ambient conditions (22°C).

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5b - Fractured Samples with Flow Rates Greater than 0.1 mL/s

FIGURE 5. Conductivity Test Apparatus

Parameter	Accuracy (Better Than)		
Fluid pump displacement	4.6 x 10^{-5} mL		
Upstream pore fluid pressure	1%		
Downstream pore fluid pressure	1%		
Differential pressure across sample	1.4%		
Confining fluid pressure	0.25%		

TABLE 6. Accuracy of Parameters Digitally Recorded During Experiment

For all tests, data for confining pressure were recorded manually in a notebook, and data for pore pressure were recorded on an analog recorder. A detailed procedure for the laboratory testing is provided in Appendix A in Figure A.1. For all samples, saturated conductivity was calculated only for intervals where flow was judged to be steady state.

The cubic law approximation for steady-state laminar flow between parallel plates was used to reduce the laboratory fracture-flow data. This approximation for flow in a crack is given by Gale (1975) as

$$Q = \frac{e^3 w}{12 \mu L} \Delta P$$
(6)

where

e = effective parallel-plate hydraulic fracture aperture (m)

$$\mu$$
 = fluid viscosity $\left(\frac{Ns}{m^2} = 10 \text{ poise}\right)$
w = fracture width (m).

The effect of surface roughness was not included in this analysis. The ratio of physical aperture to effective parallel-plate hydraulic aperture (e) appears to generally be in the range of 2 to 7 (Barton, et al., 1983). Laboratory data were used to solve for e and then a fracture conductivity, K_{f} , was calculated by using the approximation (Gale, 1975):

$$\kappa_{f} = \frac{e^{2}\rho_{g}}{12\mu}$$
(7)

where

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$$\rho$$
 = fluid density (kg/m³)

All measurements were conducted at 22°C, and a fluid viscosity of μ = 0.9548 x 10^{-2} poise was used in all calculations.

RESULTS

In this section, we discuss the data obtained by testing the water-retention characteristics, the hydraulic conductivities of the tuff matrix, and the hydraulic conductivity of the matrix and fractures at elevated confining pressures. These data are listed in tabular form in Appendix A.

RESULTS OF WATER RETENTION TESTING

Psychrometer Water Retention Data

Table A.1 in Appendix A presents the water-retention characteristics at various suction heads in the range from 0 to about 10,000 m. Appendix B contains plots based on these data. Three individual subsamples were tested for each USW G-4 depth (e.g., depth of 43' denoted as sample G4-1) and two individual subsamples were tested for each USW GU-3 depth. The relative humidity was converted to water potential using Equation 1. Table A.2 (Appendix A) summarizes the densities and hydraulic conductivities. The total porosity (Tables A.1 and A.2) was calculated using the bulk density data (Table A.1) and the grain density (Table A.2).

The total porosity and maximum measured water content ("vol/vol" column) listed in Table A.1 for each subsample are approximately the same value (i.e., the subsample was almost completely saturated--see G4-2a). There are two possible causes for those cases where the maximum water content and porosity differ greatly (e.g., G4-12a or G4-9a).

 The porosity determined from the subsample's bulk density and a single grain-density measurement for that sample depth is inaccurate because tuff variability within that sample.

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2) When the subsample's maximum water content is much less than the porosity (e.g., G4-9a), the sample may contain many very small and/or disconnected pores that could not be saturated by a vacuum saturation technique within a reasonable amount of time. It can be assumed that these pores will not contribute significantly to water flow.

For these reasons, the maximum volumetric water content rather than the porosity was used as a basis on which the relative saturation was computed. The data plotted in Appendix B are the volumetric water content divided by the volumetric water content when the sample is "saturated." This quantity will be referred to as the saturation.

Comparison with Mercury-Intrusion Tests

The results of mercury-intrusion tests have also been used to estimate the shape of the saturation curve. Mercury-intrusion tests were run on almost all of the G-4 samples. Appendix C contains 23 figures showing a comparison of the saturation data derived from the psychrometer tests and its curve fit (discussed in the next section) with the adjusted mercury-intrusion data. Table A.3 summarizes the ancillary data on porosity, bulk density, and grain density supplied by Micromeritics.

The plots in Appendix C indicate that the saturation curves derived from the mercury-intrusion data are in qualitative agreement with the psychrometer-derived saturation curves at suction-head values less than about 2×10^3 m of water. Above this suction head it becomes very difficult to push mercury into the extremely small pores, and so the water saturation, derived from the mercury-intrusion tests, at high suction heads drops to

-29

zero. In some cases there are significant differences between the mercury-intrusion results and the psychrometer test results at all suction heads (e.g., G4-3 and G4-22). The conclusions drawn from the figures in Appendix C are

- The mercury-intrusion data and the psychrometer data generally support each other over the pressure range where both tests are valid. Thus, the confidence in the results obtained from both testing methods is increased.
- 2. In a few cases there is significant disagreement both qualitatively and quantitatively. It should be noted, however, that the two fluids used in the testing procedures are fundamentally different, with mercury being a nonwetting fluid and water being a wetting fluid. The assumptions made to convert the mercury intrusion data to water saturation data are quite simplistic and may miss important effects (due to sample structure or mineralogy) that may be present in some samples and not in others. The psychrometer data are a much more direct measurement of the desired information and so they should be more reliable.

Fitting the Water-Retention Data

Many different functions have been suggested to be used to fit water retention data. A discussion of these may be found in a paper by Van Genuchten (1978). Several different functions were tried; functions suggested by Haverkamp (1977) and Van Genuchten appeared the most promising.

The Van Genuchten curve was used to fit the saturation data because it gave as good a fit as other methods and it yields an analytical expression when the relative hydraulic conductivity is calculated by the method of Mualem

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(1976) (McKeon et al., 1983). A discussion of a comparison of Van Genuchten and Haverkamp fits of the data may be found in Appendix D. The conclusion of this comparison is that when the curve is well defined by the data points, the two curve fits provide nearly identical results.

The equation for the saturation (or equivalently, moisture content) is:

$$S = (S_{s} - S_{r}) \left[\frac{1}{1 + |\alpha h|} \right]^{\lambda} + S_{r}$$
(8)

The equation for the unsaturated conductivity using the Van Genuchten fit and the method of Mualem (1976) is:

$$K(h) = K_{s} \begin{pmatrix} \frac{\left|1 - |\alpha h|^{\beta-1} \cdot \left(1 + |\alpha h|^{\beta}\right)^{-\lambda}\right|^{2}}{\lambda/2} & \alpha \\ (1 + |\alpha h|^{\beta}) & \alpha \\ suction head - \frac{\beta_{c}}{\ell_{g}} \end{pmatrix}$$
(9)

The values of K_s , α , β , and S_r for each sample are listed in Table A.2. Estimates of the parameters were calculated using the SOILGEN computer codes described by McKeon et al. (1983).

The method of Mualem for calculating unsaturated conductivities has been shown to work well for a variety of soils, but the authors have not found any evidence of any methods which have proven validity over the large suction head ranges and low unsaturated conductivities characteristic of the tuff samples tested. There is currently no direct way to measure low unsaturated conductivities ($<10^{-10}$ m/s) for liquid flow in a reasonable time period. Therefore, this calculational method was determined to be acceptable to meet current needs.

Discussion of Water-Retention Results by Unit

The water retention data and curve fits have been collected for each rock unit in Figures 6-14 (for definition of the units see Table 1). The figure

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for each unit has the data points for all the samples belonging to that unit plotted in the upper half of the figure with the corresponding fitted curves plotted in the lower half of the figure. These plots, together with Table 2, may be examined to gain understanding of the homogeneity of a unit and the similarities and differences between the units. The accuracy of the psychrometer data is of the order of 10 m of suction head; a fact which should be considered during examination of Figures 6-14. The information in Figures 6-14 also was used to attempt to pick representative samples for each unit for use in future hydrologic computations. The representative sample was chosen on the basis of the following criteria which are listed in order of decreasing importance.

- 1. Saturation curve shape.
- 2. Saturated conductivity.
- 3. Porosity and other bulk properties.

As information resulting from upcoming tests becomes available an attempt will be made to apply statistical methods to this larger data set to more rigorously quantify the definition of a representative sample, the variation in parameter values, etc. This rigorous definition of property variability will be necessary for calculations required as a part of the licensing process.

Figure 6 shows a consistent pattern for Unit I-A (the welded, devitrified Tiva Canyon). A representative sample for the samples tested in this unit is G4-1.

Unit I-B (Figure 7) looks self-consistent except for G4-2 which, according to the tables, has a relatively low porosity (0.27) when compared to the other samples (0.40 to 0.65). X-ray diffraction tests of G4-2 show that it is a

-32--

devitrified, partially welded sample. Thus, its saturation curve is dissimilar to the rest of the samples which are vitric, partially welded to nonwelded tuffs. A representative sample for the samples tested in this unit is GU3-7.

The saturation curves of Units II-L and II-NL were combined in Figure 8 because the matrix of the two units should be very similar in chemical composition and bulk properties. The curves appear fairly self-consistent. X-ray diffraction testing of samples G4-5 and G4-24 show no significant differences in the major constituents of the samples. A representative sample for the samples tested in this unit is G4-6.

The saturation curves (Figure 9) for Unit III are fairly self-consistent for a unit consisting of a vitrophyre that may be locally altered. The plot of the data points shows much more consistency than the plot of the fitted curves. The porosities for this unit are very low (0.02 to 0.11) and it was difficult to saturate the samples. G4-8 has a typical or average curve shape and saturated hydraulic conductivity; however, its porosity is the highest. There appears to be no sample that can qualify as representative or average for the samples tested in this unit. GU3-10 or GU3-11 could be used in calculations.

In Figure 10, Units IV-A-v through IV-C-v were combined into Unit IV-v which is a nonwelded, vitric unit with occasional zeolitization. The major difference between Unit IV-v and Unit IV-z is the prevalence of zeolitization. The variability of the unit is also seen in the plot of the saturation curves with some of the curves appearing to be similar to those of Unit I-B (a vitric, nonzeolitized unit). Other curves tend to resemble those of Unit IV-z (a zeolitized unit). Inspection of the Unit IV-v curves together

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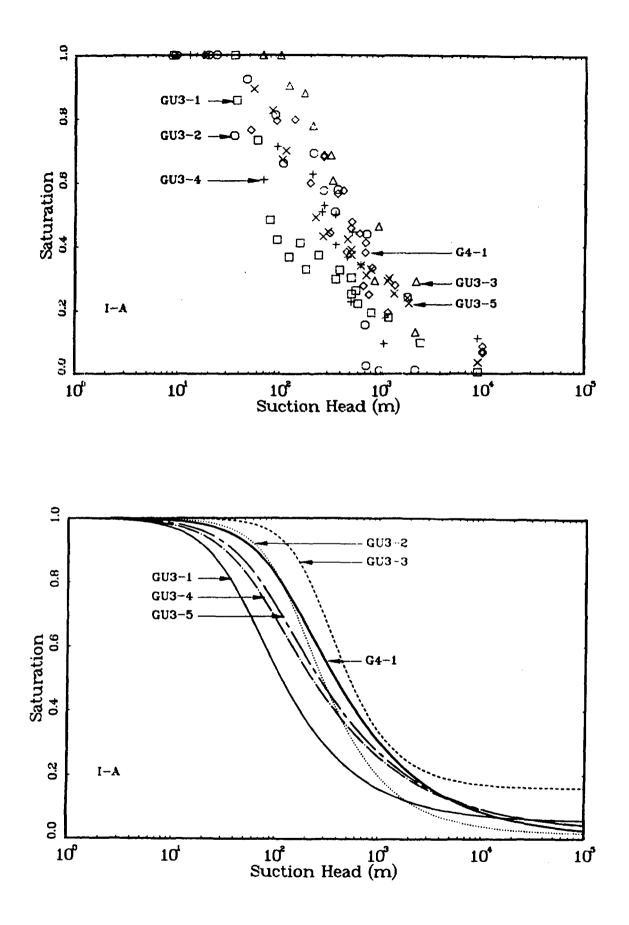


Figure 6 Tiva Canyon Welded Unit I-A

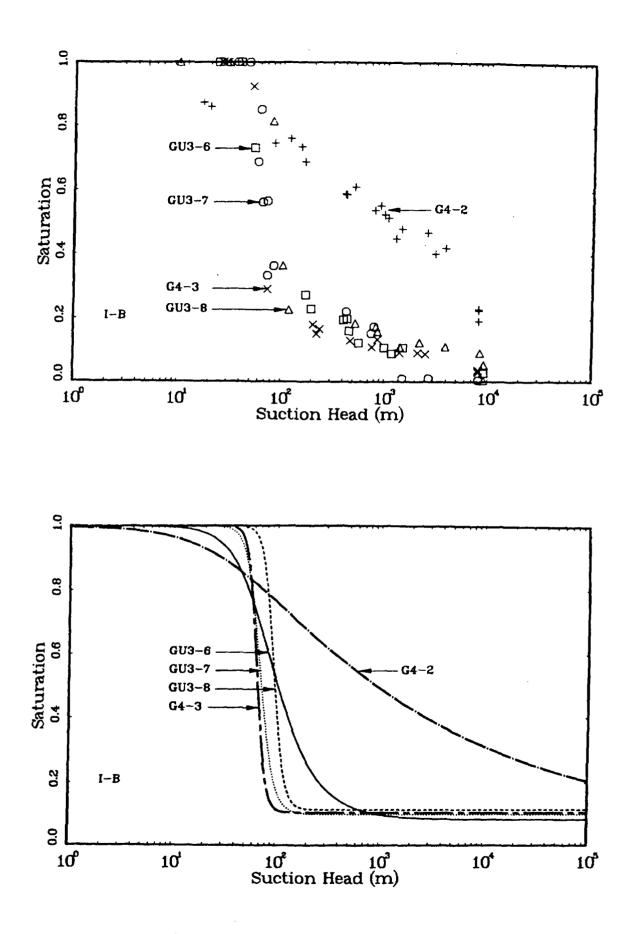


Figure 7 Paintbrush Nonwelded Unit I-B

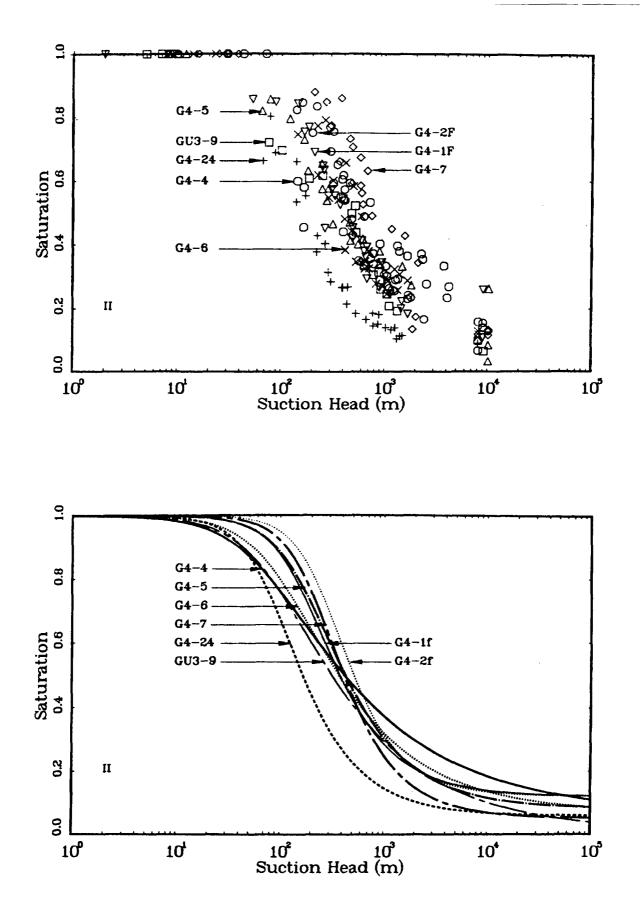


Figure 8 Topopah Spring Welded Unit II

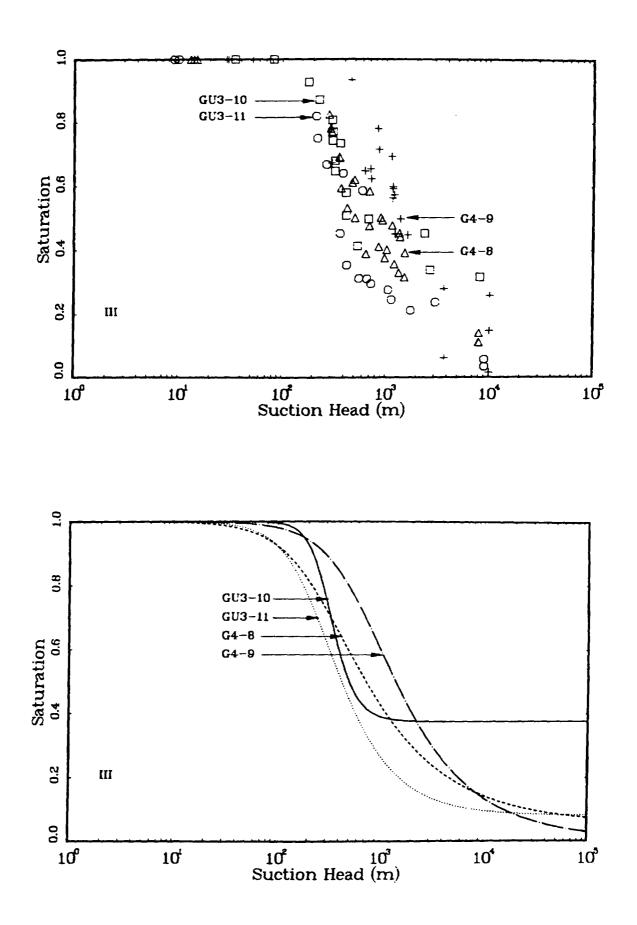


Figure 9 Basal Vitrophyre of Topopah Spring Welded Unit III

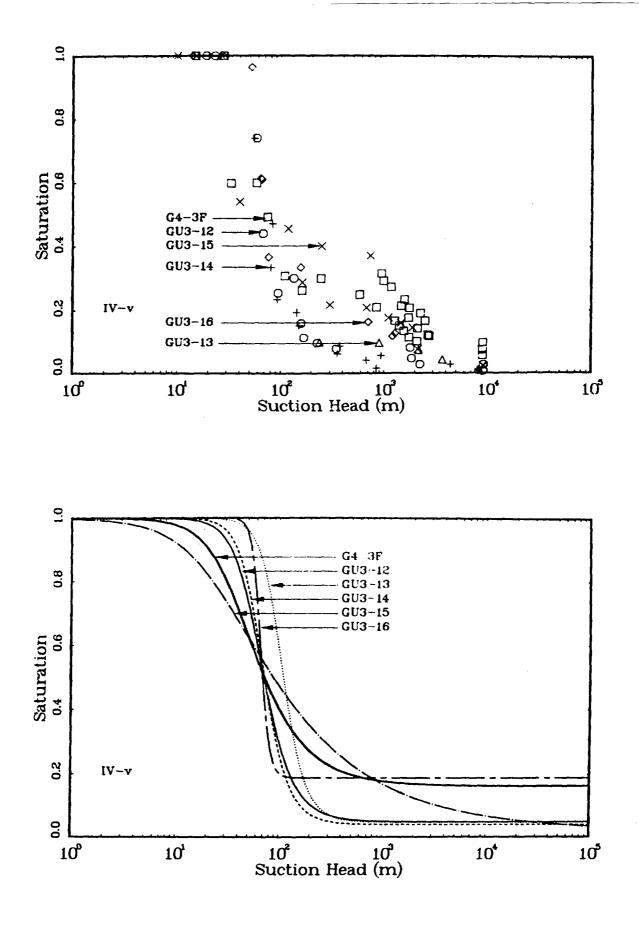


Figure 10 Vitric Calico Hills Nonwelded Unit IV-v

with Table A.2 reveals that the saturated hydraulic conductivities appear to generally correlate with the curve shape. Curves with gentle slopes (small betas) have small conductivities (e.g., G4-3F), while those with steep slopes (large betas) have high conductivities (e.g., GU3-14). This qualitative correlation appears to hold for the nonwelded units. A representative sample for the samples tested in this unit is GU3-14, which was chosen primarily on the basis of the relatively large amount of data in the region where the saturation decreases rapidly. The values of the other parameters describing GU3-14 appear consistent with those of the other Unit IV-v samples.

Unit IV-z curves in Figure 11 appear to be fairly self-consistent. That G4-10 is somewhat anomalous is not entirely unexpected, because it comes from the upper boundary of this unit. A representative sample for the samples tested in this unit is G4-11.

Samples were taken from Units V, VI, and VII for completeness. There are some portions of Yucca Mountain where Unit V occurs above the water table. In Figure 12 we see that samples from Unit V are fairly self-consistent. These samples were taken from locations below the water table, and their data appear somewhat different from those for samples taken from welded, devitrified units above the water table (I-A and II). In particular, they generally have a steeper slope and a larger conductivity than those of samples from Units I-A and II. G4-18 appears to be representative for the samples tested in this unit.

Data from Unit VI (Figure 13) are typical of those from a nonwelded, zeolitized unit, although the saturated hydraulic conductivities for this unit are somewhat larger than those measured for samples from Unit IV-z.

Unit VII is a welded devitrified unit below the water table. The saturation curves and the saturated hydraulic conductivities are presented in

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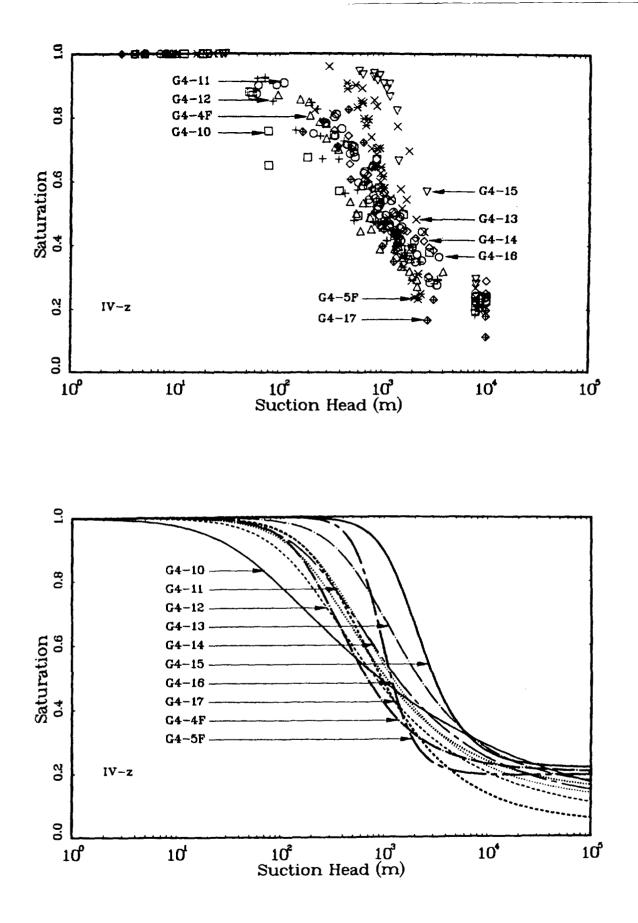


Figure 11 Zeolitized Calico Hills Nonwelded Unit IV-z

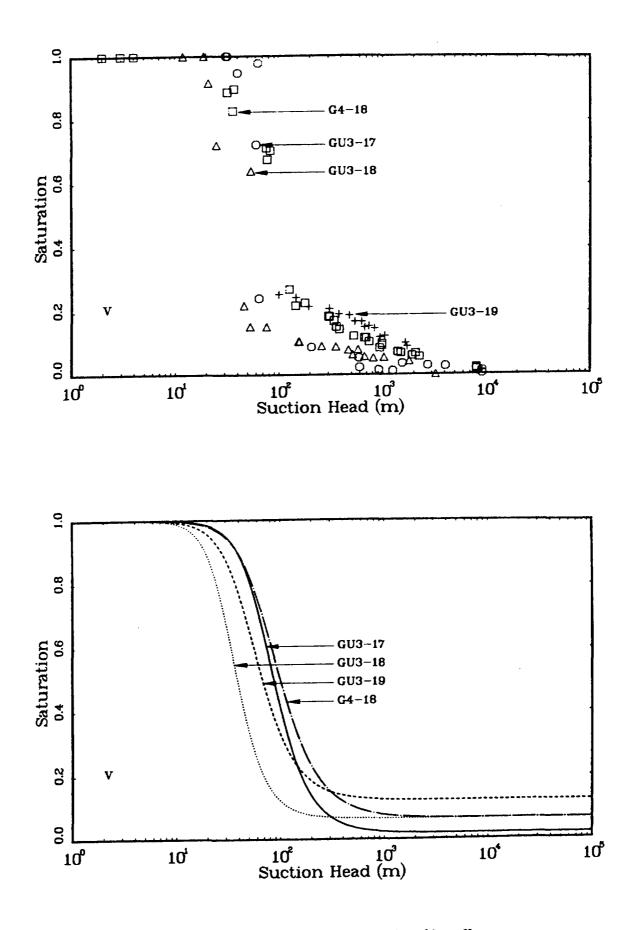


Figure 12 Prow Pass Welded Unit V

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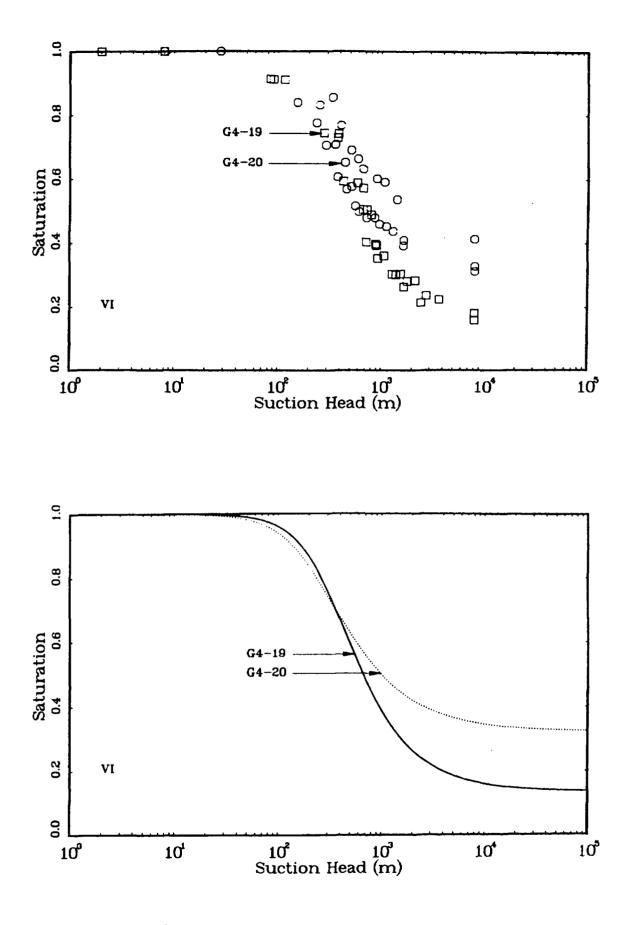


Figure 13 Crater Flat Nonwelded Unit VI

Figure 14 for this unit and visually appear different than those seen for Units I-A and II. The data visually appear to be similar to those for the Unit V samples, which were also taken from a location below the water table.

RESULTS OF SATURATED HYDRAULIC CONDUCTIVITY TESTING

This section contains the results of unconfined and confined saturated hydraulic conductivity testing of the tuff matrix and confined saturated conductivity testing of five fractured core samples.

Unconfined Matrix-Testing Results

Table A.4 in Appendix A gives the conductivities of all samples tested in the Ruska permeameter and in the core permeameter (referred to as "Large Disk" in Table A.4). It should be noted in Table A.4 that there is no statistically significant difference between the conductivity values of the small (Ruska permeameter) and large disk samples except possible for samples G4-13 and G4-17.

The following visual observations of the samples are relevant to the saturated hydraulic conductivity data.

From visual observations of the cross-sectional samples (G4-5, G4-17, G4-11, G4-13, G4-14, and G4-18), it appeared that in samples G4-5 and G4-17 the flow occurred through microsized cracks in the core samples. The texture of samples G4-11, G4-13, G4-14, and G4-18 was very coarse, and several different types of material were cemented together. Observations confirmed that water flow was through the matrix of the material. However, it could not be determined whether the water flowed through a single type of material or through all of the matrix within a sample.

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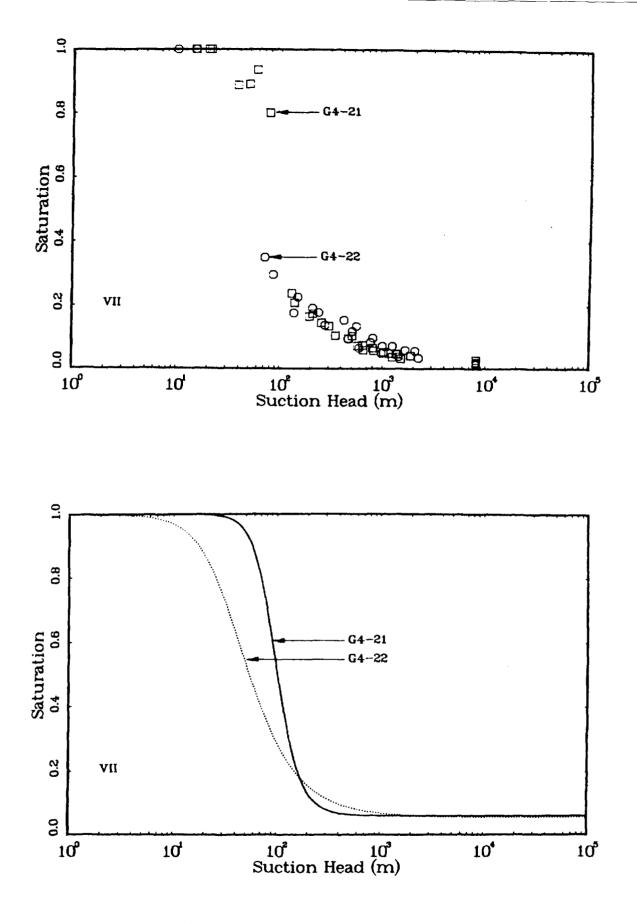


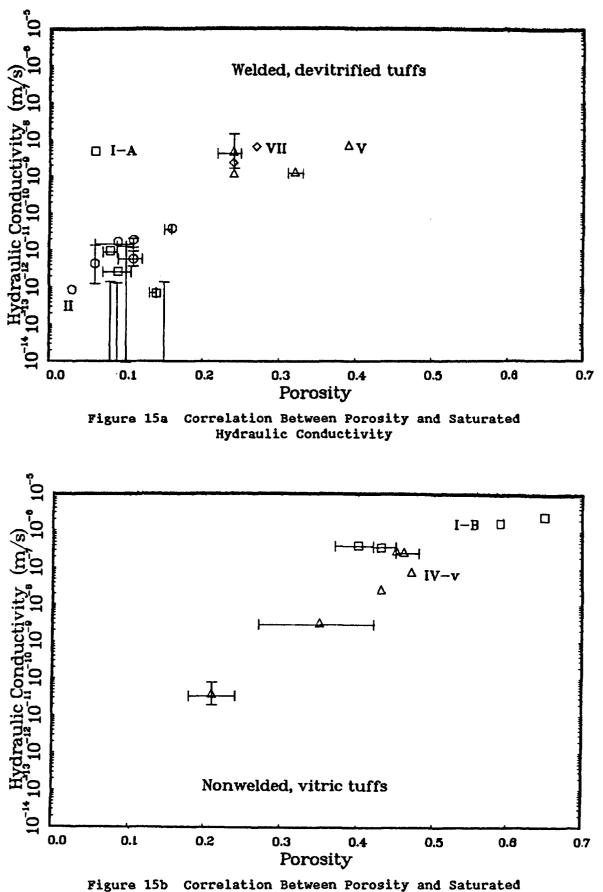
Figure 14 Bullfrog Welded Unit VII

Sample G4-3 was very fragile and loosely compacted and crumbled easily, especially when saturated. When a slight pressure head was applied to determine the conductivity, several samples cracked before a measurement could be taken. The subsample tested had no apparent cracks at the end of the run.

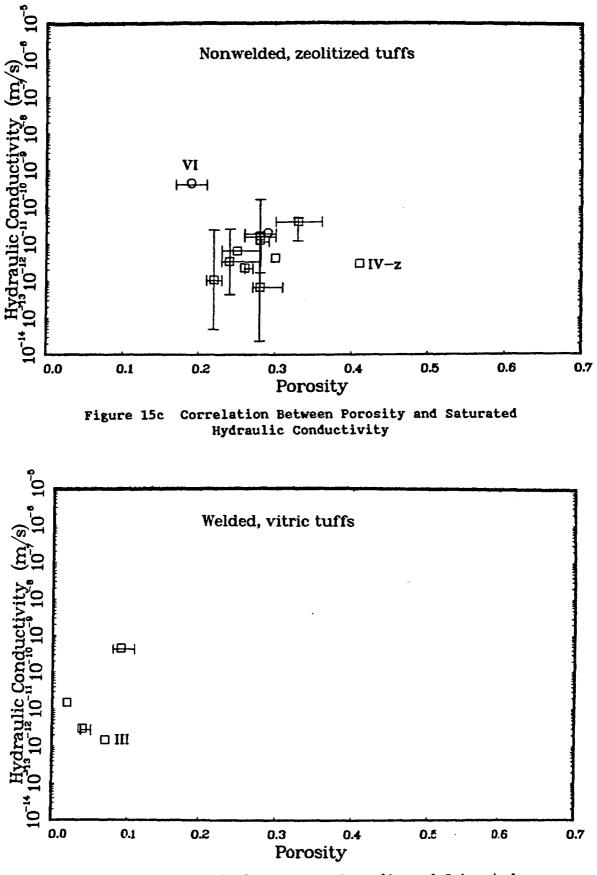
Samples G4-1, G4-1F, G4-2F, G4-4, G4-5, G4-6, G4-7, G4-8, G4-9, G4-19, G4-20, G4-21, G4-22, G4-23, G4-24, GU3-1, GU3-2, GU3-3, GU3-4, GU3-5, GU3-9, GU3-10, and GU3-11 were all welded, small-grained samples of apparently uniform material. In general, flow probably would occur only through microsized cracks in these samples. The remaining samples appeared to be less welded, lighter in weight, and of nonuniform coarse material. Flow probably occurs through the matrix of these samples.

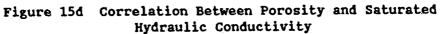
Figure 15 contains plots of each sample's average unconfined saturated hydraulic conductivity versus average porosity, with these data organized according to the four rock types found at Yucca Mountain. The type of symbol (e.g., a square) indicates the unit from which the saturated sample was taken. The bars indicate the measured range of porosity and saturated hydraulic conductivity for each sample depth. Bars that extend to the bottom of the plot indicate a test where the result indicated only an upper limit on the saturated conductivity (e.g., <1.3 x 10^{-11} m/s). These plots indicate the conductivities for the welded samples and the nonwelded, vitric samples appear to be positively correlated with porosity, while the nonwelded, zeolitized samples appear to have no correlation. The zeolitized samples all exhibit low saturated hydraulic conductivities, presumably due to the higher microporosity relative to the nonzeolitized samples.

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Confined Matrix-Testing Results

Conductivity measurements of unfractured samples at elevated effective pressures were made on four unfractured tuff core samples from drill hole USW GU-3 at the NTS. These experiments measured conductivity as a function of increasing effective pressure at ambient temperature in order to begin gathering data concerning the affect of confining stress on saturated conductivity. As a point of reference, the vertical stress at the water table under Yucca Hountain is about 130 bars. Higher confining stresses were investigated to gain further understanding of rock behavior. A discussion of the results for each sample follows.

GU3-2

Data for sample GU3-2 are shown in Figure E.1 (also see Table A.5, Appendix A). This sample displays a loss of conductivity in the effective pressure range of 50 to 350 bars that is typical of microcrack closure. At higher effective pressures the loss of conductivity for this sample is probably associated with inelastic sample deformation that was observed after testing was completed.

GU3-15

Data for sample GU3-15 are also shown in Figure E.1 (see Table A.6, Appendix A). This sample displays a loss of conductivity very similar to sample GU3-2 for the range 50 to 150 bars. However, when effective pressure was increased between 150 and 550 bars, conductivity for this sample decreased sharply, losing 90% of its conductivity. GU3-15 is a high porosity, low density, nonwelded sample and much of the conductivity reduction above 150 bars is due to inelastic sample compaction.

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GU3-11

Conductivity of sample GU3-11 was 10^{-13} m/s at initial conditions of pressure. As pressure was increased, conductivity of this sample decreased to below our detection limit (~10⁻¹³ m/s).

GU3-3

Sample GU3-3 was tested in a slightly different way from the above samples. For this sample, effective pressure was cycled during testing as shown in Figure E.2 (also see Table A.7, Appendix A). The effective pressure was cycled to determine the amount of hysteresis in conductivity loss with increasing effective pressure. Conductivity data for this sample are plotted versus effective pressure in Figure E.3. Sample GU3-3 showed a more rapid decrease in conductivity at low effective pressures (<150 bars) than either of the other samples. The major portion of the conductivity loss was irrecoverable when effective pressure was lowered to 50 bars. When the second pressure cycle was imposed (maximum $P_{eff} = 250$ bars), conductivity decreased linearly at a rapid rate, and 50% of the loss was unrecoverable. Subsequent pressure cycles showed conductivity to decrease with increasing pressure in a recoverable fashion.

The results of testing on this sample were unexpected, especially the rapid and irrecoverable loss of conductivity noted at pressures between 50 and 150 bars. Close inspection of this sample after testing revealed a well-developed, through-going crack. This rapid and unrecoverable loss of conductivity with increasing effective pressure is consistent with crack closure and deformation of asperities along the crack with increasing pressure (Johnson, 1983). Thus, the conductivity data for this sample should be interpreted as data for a fractured tuff sample.

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Laboratory conductivity data indicate that all four samples had very low conductivity, and data for GU3-2 and GU3-15 indicate that in the pressure range that simulates repository conditions ($P_{eff} = 0 - 150$ bars) conductivity decreased approximately 10%, probably because of the closure of microcracks in the samples. This decrease in conductivity is small compared to that due to other factors (e.g., variability within a unit or saturation).

Confined Fracture-Testing Results

Conductivities of five fractured tuff samples were measured at a series of elevated confining fluid pressures between 35 and 150 bars, with the average pore pressure held constant at 30 bars. Effective pressures ranged from 5 to 120 bars for these samples.

Each sample behaved differently during the flow testing, and a discussion of each sample's results is presented below.

Sample G4-1F

This sample is a highly welded tuff and the surfaces of the fracture are rough and display a poor fit when compared to the other samples (see Figure 1). Laboratory data and calculated parameters are listed in Table A.8. Figures E.4 through E.6 show that calculated aperture and relative conductivity for the sample decreased evenly with increasing pressure over the entire pressure range. The sample also displayed the least amount of change in relative conductivity.

Because the fracture surfaces were rough and poorly mated, it is likely that flow through this sample occurred mainly in channels with relatively large apertures and that the increased pressure may have reduced the size of

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these channels, but did not close them. When effective pressure was reduced the sample regained 74% of its original calculated fracture conductivity. The hysteresis evident in the plots can be attributed to 1) mechanical (frictional) bonding of the fracture during pressurization, and 2) asperity mating or breaking of asperities during pressurization.

Sample G4-2F

This sample is a highly welded tuff containing a fracture with a relatively smooth surface. Data and calculated parameters for the sample are listed in Table A.9 and are plotted in Figures E.7 through E.9. Computed aperture and fracture conductivity decreased as effective pressure was increased to 70 bars; they then remained nearly constant as effective pressure was increased to approximately 120 bars. The rock recovered at least 50% of its initial fracture conductivity (Figure E.9). Upon removal from the pressure vessel, a small thin chip about 1 cm² in size was found in place, but broken from the fracture surface. Hysteresis evident in the plots may be caused by plastic deformation along the fracture surface.

Sample G4-3F

This is a dense, strong tuff containing a very planar fracture with smooth, well-matched surfaces. The sample displayed the largest relative change in computed aperture and fracture conductivity, with most of the decrease occurring as P_{eff} was increased from 20 to 60 bars (see Table A.10 and Figures E.10 through E.12). Although the computed fracture conductivity decreased to 10% of the original value, it recovered to 83% of the original value when effective pressure was reduced to 10 bars. Hysteresis in plots for this sample is attributed to mechanisms mentioned previously for sample G4-1F.

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Sample G4-4F

This sample is a nonwelded tuff. Laboratory data and calculated parameters are listed in Table A.11 and plots of computed aperture, normalized aperture, and normalized fracture conductivity are shown in Figures E.13 through B.15. Fracture conductivity decreased with increasing pressure over the entire pressure range; however, conductivity decreased more rapidly as effective pressure was increased over the range of 10 to 60 bars. It is also interesting to note that the conductivity of this sample only returned to 36% of the original conductivity when pressure was reduced. This permanent loss of conductivity may be due in part to a plastic deformation along the fracture surface that improved mating between the fracture surfaces as pressure was increased. Also, the sample matrix contained several voids that were deformed as pressure was increased; however, this probably had little effect on fracture conductivity because the low conductivity of the matrix severely restricted flow through the internal voids.

Sample G4-5F

This is a nonwelded tuff sample with a planar fracture oriented parallel to the core axis. The surfaces of the fracture are extremely well matched (see Figure 1). Laboratory data and selected calculated parameters are listed in Table A.12. Computed aperture, normalized aperture, and normalized fracture conductivity are shown as a function of effective pressure in Figures E.16 through E.18, respectively. These figures show that, as effective pressure was increased from 10 to 20 bars, computed aperture and conductivity decreased rapidly. Conductivity of the fracture continued to decrease with increasing effective pressure to a value of 20% of the initial conductivity. Conductivity returned to 100% of the original value when effective pressure

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was lowered. Data for this sample exhibit much less hysteresis than other samples during the entire pressurization cycle, probably because of the extremely good mating between the two surfaces of the fracture.

Summary of Fracture Testing Results

Table 7 contains a summary of the data obtained from the G4-IF through G4-5F samples. The initial hydraulic aperture of the discrete fractures (at low effective confining pressures) ranges from 6 μ m for samples G4-2F and G4-5F to 67 μ m for sample G4-1F, with respective fracture conductivities of 3.0 x 10⁻⁵ to 3.8 x 10⁻³ m/s. It should be noted that the aperture calculated is an equivalent parallel-plate hydraulic aperture and may be significantly different from the physical, or real, effective fracture aperture, with the equivalent hydraulic aperture always being smaller because of the effects of the fracture surface roughness on flow through the fracture. The ratio of the effective physical aperture to the effects of the surface roughness more significant at lower apertures (Barton et al., 1983).

Fitting the Confined Fracture-Conductivity Data

Normalized conductivity as a function of effective confining pressure is shown in Figures E.6, E.9, E.12, E.15, and E.18 for the five fractured core samples. Analyses by Walsh (1981) indicate that the cube root of the fracture conductivity should be linearly related to the logarithm of the effective confining pressure. The following relationship suggested by Walsh was used to fit both the loading and unloading laboratory data:

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	Fracture	Aperture	Fracture Conductivity			
Sample <u>Code</u>	Computed Initial Aperture	Normalized Minimum Aperture	Recovered Aperture	Computed Fracture <u>Conductivity</u>	Normalized Minimum <u>Conductivity</u>	Recovered Fracture <u>Conductivity</u>
	e _o (microns)	(% e ₀)	(% e ₀)	K _o (10 ⁻⁵ m/s)	(% K _o)	(% K _o)
G4-1F	67	57%	86%	378	32%	74%
G4-2F	6	49%	85%	3.5	24%	74%
G4-3F	22	31%	91%	43	10%	83%
G4-4F	31	39%	60%	79	15%	36%
G4-5F	6	44%	100%	3.1	20%	100%

TABLE 7. Summary Data for Fractured Tuff Samples

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$$(K/K_{init})^{1/3} = a - b \cdot \ln(P_{eff})$$
 (10)

where

K = fracture saturated conductivity

- K = fracture conductivity at the initial confining
 pressure
- P = effective confining pressure, bars (see eq. 5)
 - a,b = determined constants, related to the fracture
 surface topography

Although at some high pressures the fracture saturated conductivity will reach a minimum value, it can be seen from the curves in Figures E.6, E.9, E.12, E.15 and E.18 that, even at 130 bars, the fracture conductivity is still decreasing for the samples. This is a result of both the rock strength and the surface topography.

The laboratory data were entered in a regression analysis routine which produced estimates of the parameters "a" and "b". The fitted curves are plotted in the Figures E.6, E.9, E.12, E.15, and E.18 showing the normalized conductivity for samples G4-1F thru G4-5F. These values are physically related to the topography of the fracture surface, with "b" proportional to the root-mean-square (RMS) value of the asperity height distribution and "a" inversely proportional to the normalizing conductivity K_{init}. The estimates of both parameters and the corresponding coefficient of determination, R^2 , are shown in Table 8 for both the loading and unloading curves. When the parameter "b" is multiplied by the initial conductivity to yield a relative RMS value of the asperity height distribution, it can be seen that sample G4-1F has the largest value and G4-5F the smallest. This is consistent with the visual appearance of the fracture surfaces, with sample G4-1F described previously as having the roughest surface and worst fracture mating and sample G4-5F as having a nearly planar fracture with the fracture surfaces extremely well matched.

With values of both saturated fracture conductivities as a function of effective confining pressure (corresponding to overburden stress) for the welded and nonwelded tuff of the functional units in Yucca Mountain, and estimates of the frequency and size of fractures in those units provided by other NNWSI groups, it should be possible to estimate an effective, saturated conductivity for the rock mass.

Parameter <u>a</u>	Parameter b	<u>_R</u> ² _	$\frac{K_{\text{init}} \times 10^3}{(\text{m/s})}$
1.204	0.084	0.94	3.78
1.572	0.196	0.94	0.035
1.906	0.296	0.97	0.43
1.529	0.200	0.97	0.79
1.244	0.131	0.97	0.030
1.028	0.070	0.93	3.78
1.034	0.079	0.94	0.035
1.349	0.190	0.97	0.43
0.962	0.093	0.97	0.79
1.241	0.140	0.97	0.030
	8 1.204 1.572 1.906 1.529 1.244 1.028 1.034 1.349 0.962	<u>a</u> <u>b</u> 1.204 0.084 1.572 0.196 1.906 0.296 1.529 0.200 1.244 0.131 1.028 0.070 1.034 0.079 1.349 0.190 0.962 0.093	$\begin{array}{c ccccccccccccccccccccccccccccccccccc$

Table 8. Regression Parameter Estimates for Confined-FractureConductivity Curve Fits

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SUMMARY

The testing and analysis of 48 samples taken from drill holes USW GU-3 and USW G-4 provide information required for hydrologic and radionuclide transport analyses to determine the suitability of the Yucca Mountain site as a location for a high-level nuclear waste repository. The results of these tests indicate:

Matrix Water Retention

1) The relationship between water content and suction head for each individual sample is unique for the specific core matrix material and suggests that the matrix properties could be described reasonably well by the smooth curves.

2) A comparison of psychometric and mercury intrusion data for 22 individual samples indicates that the two testing methods give results that are, for the most part, in good qualitative agreement. There appears to be no general correlation between the type of rock and degree of agreement between the two testing methods. The psychrometric method is a much more direct measurement of the desired information and so appears more reliable.

3) The data on water content versus suction head data for the limited number of samples taken from a particular rock type (functional unit) form a reasonably coherent group.

4) Comparison of Haverkamp and Van Genuchten curve-fits of data for selected samples indicates that when the curve is well defined by the data points the two curve fits yield identical results. The Van Genuchten curve fit was

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chosen because it is more convenient to use, as it yields an analytical expression when the unsaturated hydraulic conductivity is calculated by the method of Maulem (1976).

Matrix Saturated Hydraulic Conductivity

1) The nonwelded, vitric tuff samples had conductivities orders of magnitude higher than those of either the welded tuff samples or the nonwelded, zeolitic tuff samples.

2) As individual groups, the nonwelded, vitric tuff samples and the welded, devitrified samples appear to have a general correlation between the porosity of their samples and the hydraulic conductivity. The other two tuff types did not show any correlation.

3) The six core cross sections (diameter about 6.3 cm) tested for hydraulic conductivity generally had hydraulic conductivities equal to or lower than those for the cylindrical samples (diameter about 1.2 cm).

4) The reduction in conductivity, as confining pressure is increased to approximately the lithostatic load (maximum value of about 130 bars), is fairly small compared to the reduction due to other factors (e.g. saturation). For example, the reduction in saturated conductivity for the three samples for which we have data was less than 15% as the effective pressure was increased from 50 to 150 bars (see Tables A.5 to A.7). If the matrix saturation of sample GU3-2 is reduced from 100% to 90%, the calculated matrix conductivity will decrease by 75%.

The contrast in matrix properties between the unsaturated-zone hydrologic units may be seen in Table 9 and Figure 16, which contain information for samples "representative" of the individual units. Information contained in this paper indicates that the matrix data fall into three general groups:

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- 1. Densely welded tuffs (Units I-A, II, and III).
- 2. Vitric, nonwelded tuffs (Units I-B and IV-v).
- 3. Zeolitized, nonwelded tuffs (Unit IV-z).

The information concerning the representative samples (contained in Table 9 and Figure 16) supports this conclusion. The representative samples are suggested for "first-cut" hydrologic calculations concerning Yucca Mountain. As additional information becomes available, an attempt will be made to apply statistical methods to this larger data set to more rigorously quantify the definition of a representative sample, the variation in parameter values, etc.

Fracture Saturated Hydraulic Conductivity

1) The saturated conductivity of the fractures is several orders of magnitude higher than that of the matrix.

2) Flow through all fractured samples was substantially reduced at elevated pressures; however, the response of a particular sample to pressure was influenced by the fracture surface characteristics and the mechanical strength of the sample.

3) Fractured samples that were composed of strong rock regained 75 to 100% of initial conductivity when pressure was lowered to initial levels, while a weaker sample recovered only 36% of its initial conductivity indicating that irreversible (plastic) mating of the fracture surfaces had occurred.

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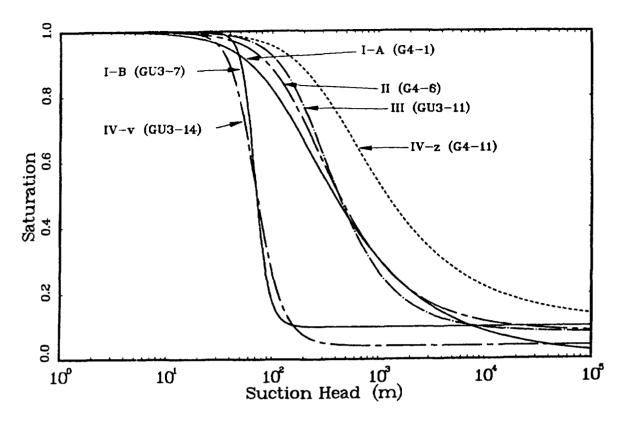


Figure 16. Representative Water-Retention Curves for the Unsaturated Zone Hydrologic Zone Units

Table 9.	Properties for Representative Samples Taken From
	the Unsaturated Zone Hydrologic Units

	Sample	Grain Density		Hydraulic Conductivity		Alpha	Beta
Unit	Code	(g/cm ³)	Porosity	(m/s)	s _r	(1/m)	
I-A	G4-1	2.49	0.08	9.7E-12	0.0020	0.821E-02	1.558
I-B	GU3-7	2.35	0.40	3.9E-07	0.1001	0.150E-01	6.872
II	G4-6	2.58	0.11	1.9E-11	0.0801	0.567E-02	1.798
III	GU3-11	2.38	0.07	1.5E-12	0.0804	0.441E-02	2.058
IV-v	GU3-14	2.37	0.46	2.7E-09	0.0405	0.160E-01	3.872
IV-z	G4-11	2.23	0.28	2.0E-11	0.1095	0.308E-02	1.602

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APPENDIX A

TEST RESULTS

This appendix contains the data gathered by PNL and is organized in the following manner:

- Table A.1 Porosity and Water Retention Data for Samples from Drill Holes USW G-4 and USW GU-3
- Table A.2 Properties of Samples Taken From USW G-4 and USW GU-3
- Table A.3 Summary of Mercury-Intrusion Data Supplied by Micromeritics
- Table A.4 Unconfined Saturated Hydraulic Conductivity Data for Samples Taken From USW G-4 and USW GU-3
- Figure A.1 Technique for Measuring Sample Saturated Conductivity as a Function of Elevated Confining Pressure
- Table A.5Data and Calculated Values for Confined, SaturatedConductivity Tests of Sample GU3-2
- Table A.6Data and Calculated Values for Confined, SaturatedConductivity Tests of Sample GU3-15
- Table A.7 Data and Calculated Values for Confined, Saturated Conductivity Tests of Sample GU3-3

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- Table A.8Laboratory Data and Calculated Parameters for Confined,Saturated Conductivity Tests of Sample G4-1F
- Table A.9 Laboratory Data and Calculated Parameters for Confined, Saturated Conductivity Tests of Sample G4-2F
- Table A.10 Laboratory Data and Calculated Parameters for Confined, Saturated Conductivity Tests of Sample G4-3F
- Table A.11 Laboratory Data and Calculated Parameters for Confined, Saturated Conductivity Tests of Sample G4-4F
- Table A.12 Laboratory Data and Calculated Parameters for Confined, Saturated Conductivity Tests of Sample G4-5F

						Water	Content
Sample	Depth (ft)	Unit	Bulk Density (g/cm ³)	Porosity	Head (m)	g/g	vol/vol
G4-1 a	43	A-I	2.26	.09	0 9 52 199 314 459 670 757 1183 10200	0.0322 0.0322 0.0246 0.0193 0.0143 0.0124 0.0090 0.0081 0.0063 0.0023	0.07 0.07 0.06 0.04 0.03 0.02 0.02 0.02 0.01 0.005
G4-16	43	I-A	2.31	.07	0 10 92 270 372 505 699 816 1387 10200	0.0242 0.0242 0.0192 0.0165 0.0137 0.0111 0.0100 0.0081 0.0068 0.0016	0.06 0.04 0.04 0.03 0.03 0.02 0.02 0.02 0.02 0.02
G4–1e	43	I-A	2.32	.07	0 18 138 272 423 515 612 699 10200	0.0217 0.0217 0.0173 0.0149 0.0125 0.0104 0.0096 0.0083 0.0019	0.05 0.05 0.04 0.03 0.03 0.02 0.02 0.02 0.02 0.02 0.02
G4-2a	124	I -8	1.72	.28	0 5 20 83 408 973 1436 3825 8109	0.1420 0.1420 0.1220 0.1061 0.0837 0.0746 0.0683 0.0598 0.0330	0.24 0.24 0.13 0.14 0.13 0.12 0.10 0.06
G4-2Þ	124	I-8	1.72	.28	0 12 151 418 785 2550 3060 8109	0.1434 0.1434 0.1055 0.0841 0.0773 0.0672 0.0579 0.0327	0.25 0.25 0.18 0.15 0.13 0.12 0.10 0.05
G4-2c	124	I-B	1.80	.25	0 5 17 118 165 500 890 1056 1265 8109	0.1199 0.1199 0.0913 0.0826 0.0733 0.0663 0.0618 0.0539 0.0233	0.22 0.22 0.19 0.16 0.15 0.13 0.12 0.11 0.10 0.04

TABLE A.1. Porosity and Water Retention Data for Samples from Drill Holes USW G-4 and USW GU-3 $\,$

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Sample	Depth (ft)	Unit	Bulk Density (g/cm ³)	Porosity	Head (m)	Water g/g	Content vol/vol
G4-3a	208	I-B	0.85	.65	0 26 51 237 467 2499 8109	0.5359 0.5359 0.4945 0.0901 0.0704 0.0481 0.0169	0.45 0.42 0.08 0.06 0.05 0.001
G4-3Þ	208	I-B	0.85	. 65	0 26 73 222 757 1409 8109	0.5446 0.5446 0.1591 0.0833 0.0606 0.0510 0.0196	0.46 0.46 0.14 0.07 0.05 0.04 0.02
G4–3c	208	I-B	0.85	. 65	0 29 204 856 2107 8109	0.4947 0.4947 0.0901 0.0668 0.0463 0.0189	0.42 0.42 0.08 0.06 0.04 0.02
G4—4a	247	11 <i>-</i> L	2.49	.031=	0 30 166 397 900 1377 1650 3743 4055 8109	0.0131 0.0076 0.0058 0.0053 0.0050 0.0050 0.0048 0.0044 0.0031 0.0013	0.033 0.033 0.019 0.014 0.013 0.012 0.012 0.012 0.011 0.003 0.003
G4-47	247	11-L	2.49	.031 ¤	0 43 318 377 713 1329 2251 2319 2550 8109	0.0118 0.0118 0.0070 0.0065 0.0063 0.0047 0.0044 0.0042 0.0033 0.0008	0.029 0.029 0.017 0.016 0.016 0.012 0.011 0.010 0.008 0.002
G4–4e	247	II-L	2.49	.031 *	0 71 144 165 561 1051 1341 4212 8109	0.0125 0.0125 0.0075 0.0057 0.0044 0.0038 0.0037 0.0034 0.0020	0.031 0.031 0.019 0.014 0.010 0.009 0.009 0.008 0.005
G4–5a	864 (dark)	II-NL	2.26	.11	0 10 65 181 275 466 589 1140 10200	0.0482 0.0482 0.0395 0.0305 0.0260 0.0200 0.0169 0.0127 0.0016	0.11 0.09 0.07 0.06 0.04 0.03 0.02 0.01

TABLE A.1. (cont)

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 $^{\rm B}$ - The porosity value listed is an estimate based on the grain density from a nearby sample; see Table A.2 for further information.

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						Water	Content
Sample	Depth	Unit	Bulk Density	Porosity	Head	g/g	vol/vol
	(ft)		(g/cm^3)		(m)		
G4-56	864	II-NL	2.30	.09	0	0.0362	0.08
					12 121	0.0362 0.0288	0.08 0.07
					245	0.0237	0.05
					287	0.0209	0.05
					456	0.0171	0.04
					612	0.0152	0.03
					887	0.0138	0.03
					1474	0.0121	0.02
					10200	0.0096	0.004
G4-5c	864	II-NL	2.38	.06	0	0.0493	0.12
					8	0.0493	0.12
					78	0.0422	0.10
					167	0.0360	0.09
					249 315	0.0284 0.0230	0.07 0.05
					553	0.0199	0.04
					882	0.0169	0.03
					1822	0.0136	0.03
					10200	0.0042	0.01
G4-24a	864	II-NL	2.21	. 16	0	0.0586	0.13
8*3-10	(light)	11-06	C.C	. 10	9	0.0586	0.13
					87	0.0404	0.09
					172	0.0325	0.07
					224	0.0252	0.06
					303 392	0.0168	0.04
					529	0.0158 0.0109	0.04 0.02
				•	667	0.0098	0.02
					780	0.0086	0.02
					1166	0.0079	0.02
					1326	0.0062	0,01
•					8109	0.0052	0.01
G4-24D	864	II-NL	2.23	. 15	0	0.0545	0.12
	•••				2	0.0545	0.12
					67	0.0362	0.08
					140	0.0292	0.07
					220 285	0.0207 0.0172	0.05 0.04
					394	0.0146	0.03
					433	0.0118	0.03
					771	0.0102	0.02
					867	0.0083	0.02
					1261	0.0077	0.02
					1492 8109	0.0063 0.0050	0.01 0.01
					0109	0.0000	0.01
G4-24c	864	II-NL	2.21	.16	0	0.0558	0.12
					1	0.0558	0.12
					78 140	0.0448 0.0369	0.09 0.08
					267	0.0226	0.05
					442	0.0151	0.03
					882	0.0102	0.02
					1032	0.0079	0.02
					1415	0.0064 0.0050	0.01 0.01
					8109	0.0000	0.01
G4-6a	1158	II-NL	2.30	.11	0	0.0399	0.09
			-		27	0.0399	0.09
					144	0.0298	0.07
					226 286	0.0247 0.0219	0.06 0.05
					353	0.0216	0.05
					414	0.0154	0.04
					528	0.0139	0.03
					642	0.0131	0.03
					828 040	0.0112	0.03
					949 1054	0.0109 0.0099	0.03 0.02
					8109	0.0044	0.01
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Sample	Depth (ft)	Unit	Bulk Density (g/cm ³)	Porosity	Head (m)	Water g/g	Content vol/vol
G4-65	1158	II-NL		. 11	14 224 418 497 592 645 702 1061 1076 1342 8109	0.0406 0.0406 0.0314 0.0267 0.0199 0.0142 0.0133 0.0120 0.0117 0.0106 0.0053	0.09 0.07 0.06 0.05 0.03 0.03 0.03 0.03 0.03 0.03 0.02 0.01
G4–6c	1158	II-NL	2.30	.11	0 23 266 316 408 476 696 901 1242 1394 1666 8109	0.0445 0.0445 0.0352 0.0268 0.0215 0.0198 0.0170 0.0153 0.0145 0.0138 0.0129 0.0054	0.10 0.08 0.06 0.05 0.05 0.04 0.04 0.04 0.03 0.03 0.03 0.03
G4-1Fa	1215	II-NL	2.28	. 12*	0 9 88 167 258 369 671 934 1145 1443 1681 9027	0.0395 0.0395 0.0335 0.0298 0.0251 0.0219 0.0208 0.0156 0.0137 0.0099 0.0080 0.0073 0.0043	0.09 0.08 0.07 0.06 0.05 0.05 0.04 0.03 0.02 0.02 0.02 0.02 0.02
G4-1FD	1215	II-NI	. 2.28	. 12*	0 8 53 185 251 307 487 598 630 676 1464 9027	0.0395 0.0395 0.0338 0.0304 0.0256 0.0227 0.0180 0.0164 0.0151 0.0117 0.0089 0.0044	0.09 0.09 0.07 0.06 0.05 0.04 0.04 0.03 0.03 0.02 0.01
G4–1Fc	1215	II-N	L 2.36	.09 =	0 2 144 209 265 489 652 893 9027	0.0324 0.0324 0.0273 0.0224 0.0147 0.0137 0.0118 0.0087 0.0085	0.08 0.08 0.07 0.05 0.04 0.03 0.03 0.02 0.01
G4-7a	1256	II-N	L 2.30	. 09	0 30 270 459 577 581 609 2020 10200	0.0341 0.0341 0.0289 0.0250 0.0230 0.0200 0.0180 0.0060 0.0040	0.08 0.08 0.07 0.06 0.05 0.05 0.04 0.01 0.01

* - The porosity value listed is an estimate based on the grain density from a nearby sample; see Table A.2 for further information.

TABLE A.1. (cont)

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						Water	Content
Sample	Depth (ft)	Unit	Bulk Density (g/cm ³)	Porosity	Head (m)	g/g	vol/vol
G4-7Ъ	1256	II-NL	2.28	. 10	0 38 379 486 675 747 1161 1564 2105 10200	0.0321 0.0321 0.0276 0.0227 0.0203 0.0158 0.0145 0.0138 0.0111 0.0043	0.07 0.06 0.05 0.05 0.04 0.03 0.03 0.03 0.03 0.03
G4–7c	1256	II⊸NL	2.27	. 10	0 16 209 302 362 477 596 1887 10200	0.0403 0.0403 0.0354 0.0267 0.0239 0.0227 0.0055 0.0052	0.09 0.09 0.08 0.07 0.06 0.05 0.05 0.05 0.01 0.01
G4–2Fa	1278	II-NL	2.35	.060	0 10 294 396 457 628 776 1277 1703 9027	0.0294 0.0294 0.0242 0.0226 0.0182 0.0175 0.0144 0.0110 0.0086 0.0072 0.0046	0.07 0.06 0.05 0.04 0.04 0.03 0.03 0.03 0.02 0.02 0.02
G4-2FЪ	1278	II-NL	2.34	.064	0 30 159 200 299 347 408 487 774 988 1249 1673 2427 9027	0.0310 0.0262 0.0233 0.0215 0.0202 0.0178 0.0134 0.0106 0.0096 0.0080 0.0072 0.0052 0.0044	0.07 0.07 0.06 0.05 0.05 0.04 0.03 0.03 0.02 0.02 0.02 0.02 0.01 0.01
G4–2Fc	1278	II-NL	2.36	.056	0 10 219 319 387 409 478 730 1028 1114 1799 9027	0.0307 0.0256 0.0232 0.0186 0.0166 0.0148 0.0110 0.0103 0.0082 0.0073 0.0043	0.07 0.06 0.06 0.04 0.04 0.04 0.04 0.03 0.02 0.02 0.02 0.02 0.01

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						Water	Content
Sample	Depth (ft)	Unit	Bulk Density (g/cm ³)	Porosity	Head (m)	g/g	vol/vol
G4-8a	1299	III	2.23	.11	0 13 292 370 424 505 646 979 1212 1342 1531 8109	0.0371 0.0290 0.0221 0.0198 0.0187 0.0145 0.0140 0.0133 0.0123 0.0118 0.0042	0.08 0.06 0.05 0.04 0.04 0.03 0.03 0.03 0.03 0.03 0.03
G4-8b	1299	III	2.28	.09	0 14 283 357 498 694 896 928 1155 1364 1382 8109	0.0246 0.0203 0.0170 0.0153 0.0144 0.0124 0.0122 0.0118 0.0112 0.0109 0.0035	0.06 0.05 0.04 0.03 0.03 0.03 0.03 0.03 0.03 0.03
G4–8e	1299	111	2.31	.08	0 15 300 355 473 696 856 1025 1539 8109	0.0320 0.0320 0.0247 0.0222 0.0196 0.0153 0.0132 0.0129 0.0126 0.0036	0.07 0.07 0.05 0.05 0.05 0.04 0.03 0.03 0.03 0.03
G4-9a	1324	111	2.28	.038	0 30 301 624 1168 1209 1391 1633 10200	0.0040 0.0040 0.0027 0.0026 0.0024 0.0023 0.0020 0.0018 0.0006	0.009 0.009 0.005 0.005 0.005 0.005 0.005 0.002 0.001
G4-95	1324	III	2.25	. 051	0 52 459 836 857 1132 1162 3700 10200	0.0046 0.0046 0.0043 0.0036 0.0033 0.0032 0.0026 0.0013 0.0012	0.010 0.010 0.010 0.007 0.007 0.006 0.003 0.003
G4-9c	1324	III	2.27	.042	0 29 298 711 724 1173 1234 3784 10200	0.0064 0.0050 0.0050 0.0042 0.0040 0.0038 0.0029 0.0004 0.0004	0.015 0.015 0.011 0.010 0.009 0.009 0.007 0.001 0.0002

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						Water Con	itent
Sample	Depth	Unit Bu	lk Density	Porosity	Head	g/g	vol/vol
•	(ft)		(g/cm^3)		(m)		
G4-3Fa	1359	IV-A-v	1.88	.21∎	0	0.0992	0.19
					15	0.0992	0.19
					33	0.0595	0.11
					245	0.0299	0.06
					578	0.0249	0.05
					1461	0.0214	0.04
					1730	0.0176	0.03
					2118	0.0143	0.03
					2673	0.0122	0.02
							0.02
					2740	0.0119	
					9027	0.0076	0.01
G4–3Fb	1359	IV-A-v	1.82	.24	0	0.1196	0.22
04-510		1		•=•	15	0.1196	0,22
					58	0.0718	0.13
					109	0.0370	0.07
					161	0.0316	0.06
					840	0.0252	0.05
					1264	0,0201	0.04
					1410	0.0183	0.03
					1558	0,0162	0.03
					1748	0.0138	0.03
					2098	0,0123	0.02
					9027	0.0069	0.01
					<u><u></u></u>		•••
G4–3Fc	1359	IV-A-v	1.97	.18*	0	0.0666	0.13
_					28	0.0666	0.13
					74	0.0329	0.07
					936	0.0211	0.04
					986	0.0196	0.04
					1164	0.0183	0.04
					1574	0.0157	0.03
					1762	0.0139	0.03
					2256	0.0128	0.03
					2498	0.0112	0.02
					9 027	0.0066	0.01
G4-10a	1405	IV-A-z	1.39	.41#	0	0,2723	0.38
04+108	1405	1	1. 19	• • •	12	0.2723	0.38
					51	0.2395	0.33
					78	0.2063	0.29
					187	0.1839	0.26
					711	0.1601	0.22
					750	0.1452	0.20
					784	0.1327	0.18
					1163	0.1256	0.17
					1786	0.1058	0.15
					8109	0.0580	0.08
					-		o
G4—10Ъ	1405	IV-A-v	1.39	. 41≢	0	0.2231	0.31
					20	0.2231	0.31
					53 .	0.1959	0.27
					79	0.1452	0.20
					381	0.1275	0.18
					577	0.1102	0.15
					992	0.1019	0.14
					1306	0.0963	0.13
					2867	0.0846	0.12
					2007 8109	0.0442	0.06
						-	
G4-10c	1405	IV-A-z	1.39	.41#	0	0.2453	0.34
			,		18	0.2453	0.34
					55	0.2131	0.30
					336	0.1829	0.25
					861	0.1643	0.23
					918	0.1426	0.20
					938	0.1313	0.18
					1586	0.1221	0.17
					8109	0.0590	0.08

* - The porosity value listed is an estimate based on the grain density from a nearby sample; see Table A.2 for further information.

TABLE	A.1.	(cont)
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Sample I	Depth	Unit Bu	lle Donation	Demosta	11		Content
oampie i	(ft)		lk Density (g/cm ³)	Porosity	Head (m)	g/g	vol/vol
G4-11a	1548	IV-A-z	1.63	.27	0 5	0.1708 0.1708	0.28 0.28
					338	0.1370	0.22
					523	0.1212	0,20
					664	0.1079	0.17
					1212	0.0924	0.15
					1471	0.0837	0.14
					2413	0.0759	0.12
					2820	0.0512	0.08
					10200	0.0412	0.07
G4-115	1548	IV-A-z	1.61	.28	0	0.1840	0.30
					8	0.1840	0.30
					399 536	0.1407	0.23
					665	0.1246	0.20
					1275	0.1104 0.0928	0.18
					1535	0.0845	0.15 0.14
					2044	0.0762	0.14
					2918	0.0526	0.08
					10200	0.0420	0.07
G4-11c	1548	IV-A-z	1.55	. 31	0	0.1974	0.31
					4	0.1974	0.31
					110	0.1794	0.28
					352	0.1601	0.25
					446	0.1413	0.22
					639	0.1159	0.18
	•				912 1461	0.1031	0.16
					3417	0.0897 0.0549	0.14
					10200	0.0441	0.09 0.07
G4-4Fa	1551	IV-A-z	1.51	• 36*	0	0.2307	0.35
					10	0.2307	0.35
					97	0.2007	0.30
					198	0.1859	0.28
					286 485	0.1695	0.26
					609	0.1241 0.1029	0.19
					945	0.0896	0,16
					1553	0.0837	0.14 0.13
					1962	0.0703	0.11
					2203	0.0629	0,10
					9027	0.0479	0.07
G4-4FD	1551	IV-A-z	1.62	.32*	0	0.1825	0.30
					8 150	0.1825	0.30
					159 245	0.1563	0.25
					245 344	0.1436	0.23
					544 645	0.1291 0.0974	0.21 0.16
					756	0.0974	0.18
					954	0.0849	0.14
					1387	0.0763	0.12
					1441	0.0724	0.12
					1877	0.0654	0.11
					3905	0.0582	0.09
					9027	0.0448	0.07

 $^{\tt \#}$ - The porosity value listed is an estimate based on the grain density from a nearby sample; see Table A.2 for further information.

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				Water	Content
Sample Depth (ft)	Unit Bulk Density (g/cm?)	Porosity	Head (m)	g/g	vol/vol
G4_4Fc 1551	IV-A-z 1,65	. 30*	0 5 192 284 370 559 774 1303 1527 1839 2182 9027	0.1737 0.1737 0.1473 0.1357 0.1216 0.0870 0.0785 0.0678 0.0582 0.0556 0.0511 0.0412	0.29 0.24 0.22 0.20 0.14 0.13 0.11 0.10 0.09 0.08 0.07
G4-12a 1686	IV-A-z 1.57	. 30	0 16 61 86 144 260 433 529 1095 1632 1685 3457 8109	0.2334 0.2334 0.1986 0.1774 0.1568 0.1316 0.1119 0.0972 0.0882 0.0840 0.0690 0.0428	0.37 0.34 0.31 0.28 0.25 0.21 0.18 0.15 0.14 0.13 0.11 0.07
G4-12b 1686	IV-A-z 1.57	. 30	0 27 72 202 248 374 570 813 1110 1682 1916 3188 8109	0.2143 0.2143 0.1976 0.1813 0.1591 0.1437 0.1232 0.1014 0.0890 0.0809 0.0765 0.0605 0.0428	0.34 0.31 0.28 0.25 0.23 0.19 0.16 0.14 0.13 0.12 0.09 0.07
G4-12c 1686	IV-A-z 1.57	. 30	0 24 72 232 421 782 993 1529 1740 2313 3448 8109	0.1901 0.1901 0.1759 0.1570 0.1381 0.1105 0.0101 0.0763 0.0703 0.0666 0.0557 0.0385	0.30 0.28 0.25 0.22 0.17 0.16 0.12 0.11 0.10 0.09 0.06
G4-13a 1728	IV-B-z 1.78	.23	0 23 300 624 817 870 1380 1788 8109	0.1665 0.1665 0.1599 0.1485 0.1421 0.1390 0.1287 0.1159 0.0400	0.30 0.28 0.26 0.25 0.25 0.25 0.23 0.21 0.07

 $^{\rm z}$ ~ The porosity value listed is an estimate based on the grain density from a nearby sample; see Table A.2 for further information.

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						Water C	ontent
Sample	Depth (ft)	Unit Bu	lk Density (g/cm ³)	Porosity	Head (m)	g/g	vol/vol
G4–13b	1728	IV-B-z	1.84	.21	0 16 219 770 862 948 1490 1698 2126 8109	0.1421 0.1421 0.1170 0.1050 0.0973 0.0891 0.0820 0.0773 0.0685 0.0381	0.26 0.26 0.22 0.19 0.18 0.16 0.15 0.14 0.13 0.07
G4–13c	1728	IV-B-z	1.83	.21	0 23 330 770 862 944 1186 1545 2536 10200	0.1534 0.1534 0.1244 0.1079 0.1005 0.0918 0.0840 0.0792 0.0680 0.0375	0.28 0.23 0.20 0.18 0.19 0.15 0.15 0.12 0.07
G4-14a	1737	IV-B-2	1.67	.28	0 8 442 714 1017 1459 2098 3137 10200	0.1478 0.1478 0.1056 0.0919 0.0828 0.0719 0.0626 0.0570 0.0429	0.25 0.25 0.18 0.15 0.14 0.12 0.10 0.10 0.01
G4–146	1737	IV-B-z	1.80	.23	0 9 464 776 959 1288 1591 2842 10200	0.1345 0.1345 0.1000 0.0867 0.0774 0.0673 0.0603 0.0530 0.0338	0.24 0.24 0.18 0.16 0.14 0.12 0.11 0.10 0.07
G4-14c	1737	IV-B-z	1.80	.23	0 8 342 477 790 1316 1729 2535 2856 10200	0.1416 0.1416 0.0928 0.0818 0.0698 0.0631 0.0587 0.0430 0.0335	0.25 0.25 0.19 0.17 0.15 0.13 0.11 0.11 0.08 0.06
G4-15a	1769	IV-C-z	1.77	.26	0 20 585 878 1078 1162 1365 1418 2677 8109	0.1406 0.1406 0.1329 0.1288 0.1251 0.1217 0.1156 0.0936 0.0801 0.0396	0.25 0.25 0.24 0.23 0.23 0.22 0.20 0.17 0.14 0.07
G4-15b	1769	IV-C-z	1.75	.26	0 30 635 807 1009 8109	0.1379 0.1379 0.1289 0.1279 0.1250 0.0409	0.24 0.24 0.23 0.22 0.22 0.07

TABLE	A.1.	(cont)
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						Water C	ontent
Sample	Depth (ft)	Unit Bu	lk Density (g/cm ³)	Porosity	Head (m)	g/g	vol/vol
G4-15c	1769	IV-C-2	1.74	.27	0 28 811 927 1149 8109	0.1391 0.1391 0.1305 0.1297 0.1259 0.0391	0.24 0.24 0.23 0.23 0.22 0.07
G4–16a	1778	IV-C-z	1.75	.24	0 7 60 280 500 792 844 889 1068 1290 1953 3558 8109	0.1480 0.1480 0.1293 0.1159 0.1057 0.0962 0.0810 0.0760 0.0734 0.0663 0.0583 0.0583 0.0541 0.0362	0.26 0.23 0.20 0.18 0.17 0.14 0.13 0.13 0.13 0.12 0.10 0.09 0.06
G4–16b	1778	IV-C-z	1.69	.27	0 4 213 478 578 732 933 1220 1413 2081 8109	0.1673 0.1673 0.1506 0.1255 0.1151 0.1001 0.0909 0.0828 0.0745 0.0709 0.0584 0.0383	0.28 0.25 0.21 0.19 0.17 0.15 0.14 0.13 0.12 0.10 0.06
G4–16c	1778	IV-C-z	1 . 78	. 23	0 93 344 521 654 811 1047 1164 1312 1514 2171 2416 8109	0.1499 0.1351 0.1114 0.1038 0.0940 0.0836 0.0716 0.0687 0.0654 0.0623 0.0559 0.0531 0.0337	0.27 0.24 0.20 0.18 0.17 0.15 0.13 0.12 0.12 0.12 0.11 0.10 0.09 0.06
G4–5Fa	1778	IV-C-z	1.67	.28	0 9 445 611 652 870 959 1047 1145 1315 1392 1574 2240 2382 9027	0.1711 0.1711 0.1552 0.1457 0.1363 0.1203 0.0994 0.0951 0.0834 0.0750 0.0679 0.0622 0.0534 0.0428 0.0361	0.29 0.29 0.26 0.24 0.23 0.20 0.17 0.17 0.15 0.12 0.11 0.10 0.09 0.07 0.06

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						Water Co	
Sample	Depth (ft)		lk Density (g/cm ³)	Porosity	Head (m)	g/ g	vol/vol
G4–5Fb	1778	IV-C-z	1.65	. 29	0 10 540 636 717 927 927 1029 1384 1909 2084 9027	0.1623 0.1623 0.1463 0.1372 0.1290 0.1144 0.1047 0.0943 0.0715 0.0642 0.0388 0.0317	0.27 0.24 0.23 0.21 0.19 0.17 0.16 0.12 0.11 0.06 0.05
G4–5Fc	1778	IV-C-z	1.67	. 28	0 4 562 686 1000 1081 1362 1600 1947 2254 9027	0.1686 0.1686 0.1500 0.1403 0.1308 0.1040 0.0788 0.0632 0.0563 0.0492 0.0395 0.0352	0.28 0.28 0.25 0.23 0.22 0.17 0.13 0.11 0.09 0.08 0.07 0.06
G4-17a	1787	IV-C-z	1.64	. 29	0 5 265 494 699 836 1028 1346 1346	0.1815 0.1815 0.1429 0.1278 0.1091 0.0915 0.0818 0.0723 0.0361	0.30 0.24 0.21 0.18 0.15 0.14 0.12 0.06
G4-17b	1787	IV-C-2	1.71	.26	0 3 464 648 838 939 1055 1464 3192 10200	0.1573 0.1573 0.1297 0.1136 0.1016 0.0837 0.0741 0.0673 0.0365 0.0282	0.27 0.27 0.22 0.19 0.17 0.14 0.13 0.12 0.06 0.05
G4-17c	1787	IV-C-z	1.63	. 30	0 3 168 365 490 903 997 1291 2806 10200	0.1887 0.1887 0.1425 0.1341 0.1147 0.0883 0.0752 0.0661 0.0315 0.0214	0.31 0.23 0.22 0.19 0.14 0.12 0.11 0.05 0.04
G4-18a	1899	V	1.94	.25	0 2 37 75 126 144 339 385 666 987 1407 8109	0.1224 0.1224 0.096 0.0872 0.0332 0.0270 0.0211 0.0178 0.0145 0.0115 0.0089 0.0026	0.24 0.21 0.17 0.06 0.05 0.04 0.03 0.03 0.02 0.02 0.02

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Sample	Depth (ft)	Unit	Bulk Density (g/cm ³)	Porosity	Head (m)	g/g	vol/vol
G4-18b	1899	v	(g/cm ²) 2.02	. 22	0	0.1049	0.21
					3	0.1049	0.21
					32	0.0929	0.19
	•				77 178	0.0709 0.0240	0.14 0.05
					301	0.0240	0.04
					694	0.0124	0.03
					989	0.0105	0.02
					1504	0.0073	0.02
					1935 8109	0.0064 0.0017	0.01 0.003
G4-18c	1899	V	1.94	.25	0	0.1218	0.24
					4 36	0.1218 0.1007	0.24 0.20
					82	0.0858	0.17
					307	0.0225	0.04
					359	0.0187	0.04
					527	0.0151	0.03
					741 935	0.0128 0.0104	0.02 0.02
					2084	0.0083	0.02
					2268	0.0070	0.01
					8109	0.0023	0.01
G4-19a	2006	VI	1,60	. 30	0	0.1996	0.32
		••			ž	0.1996	0.32
					117	0.1818	0.29
					386 585	0.1483	0.24
					665	0.1175 0.1011	0.19 0.16
					719	0.0808	0.13
					1072	0.0723	0.12
					1275	0.0609	0.10
					2150 3686	0.0569	0.09
					8109	0.0450 0.0364	0.07 0.06
G4-196	2006	VI	1.61	. 30	0	0.1959	0.32
04-190	2000	•1	1.01	• 30	8	0.1959	0.32
					91	0.1786	0.29
					379	0.1431	0.23
					666 798	0.1122 0.0960	0.18 0.15
					887	0.0781	0.13
					931	0.0694	0.11
					1550	0.0599	0.10
					1791 2760	0.0554 0.0467	0.09
					8109	0.0356	0.08 0.06
				_		-	
G4-19c	2006	VI	1.69	.26	0	0.1729	0.29
					2 85	0.1729 0.1579	0.29 0.27
					280	0.1288	0.22
					428	0.1027	0.17
					726	0.0875	0.15
					907 1393	0.0681 0.0524	0.12 0.09
					1669	0.0459	0.08
					2459	0.0374	0.06
					8109	0.0277	0.05
G4-20a	2101	VI	1.76	.21	0	0.1143	0.20
			· -		28	0.1143	0.20
					336	0.0978	0.17
					408 512	0.0879	0.15 0.14
					513 597	0.0790 0.0759	0.14
					666	0.0723	0.13
					903	0.0688	0.12
					1073	0.0676	0.12
					1422 8109	0.0614 0.0473	0.11 0.08
						0.0713	0.00

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						Water	Content
Sample	Depth	Unit	Bulk Density	Porosity	He ad	g/g	vol/vol
•	(it)		(g/cm ³)		(m)		
G4-20b	2101	VI	1.85	. 17	0	0.1173	0.22
					28	0.1173	0.22
					253 360	0.0976 0.0832	0.18 0.15
					447	0.0766	0.14
					514	0.0677	0.13
					606	0.0586	0.11
					720	0.0564	0.10
					1123	0.0532	0.10
					1669	0.0481	0.09
					8109	0.0387	0.07
G4-20c	2101	VI	1.82	. 19	0	0.1198	0.22
					8 154	0.1198 0.1006	0.22 0.18
					238	0.0931	0.18
					294	0.0847	0.15
					376	0.0728	0.13
					460	0.0682	0.12
					557	0.0620	0.11
					864	0.0576	0.10
					956	0.0552	0.10
					1311	0.0525	0.10
					1655	0.0471	0.09
					8109	0.0376	0.07
			a	0 h	•	0 1156	0.00
G4-21a	2401	VII	2.00	.24	0 20	0.1156	0.23
					38	0.1156 0.1025	0.23 0.21
					58 77	0.0925	0.19
					132	0.0276	0.06
					211	0.0202	0.04
					303	0.0157	0.03
					509	0.0121	0.02
					801	0.0076	0.02
					1023	0.0061	0.01
				•	1407	0.0053	0.01
					8109	0.0032	0.01
G4-215	2401	VII	1.97	, 25	0	0.1250	0.25
44-215	2401	***	• 31	.25	21	0, 1250	0.25
					49	0,1112	0.22
					141	0.0262	0.05
					256	0.0183	0.04
					351	0.0132	0.03
					641	0.0094	0.02
					820	0.0072	0.01
					1159	0.0061	0.01
					1252	0.0047	0.01
					1517	0.0040	0.01
					8109	0.0026	0.01
G4-21c	2401	VII	1.97	.25	0	0 1177	0.22
04-21C	270)	***	1+21	. 25	15	0.1177 0.1177	0.23 0.23
					58	0.1101	0.21
					196	0.0196	0.04
					469	0.0114	0.02
					573	0.0086	0.02
					650	0.0070	0.01
					1355	0.0056	0.01
					1892	0.0049	0.01
					8109	0.0033	0.01
61.22-	2807	VTT	1 00	24	~	A 1917	A A A
G4-22a	2407	VII	1,92	.27	0	0.1215	0.23
					10 71	0,1215	0.23
					151	0.0425 0.0276	0.08
					240	0.0217	0.05 0.04
					510	0.0217	0.04
					806	0.0119	0.03
					996	0.0088	0.02
					1664	0.0072	0.02
					8109	0.0016	0.01
						5,0010	0.01

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				. (conc)			
						Water	Content
Sample	Depth (ft)	Unit	Bulk Density (g/cm ³)	Porosity	Head (m)	g/g	vol/vol
G4–22b	2407	VII	1.93	.27	0 10 86 210 422 557	0.1202 0.1202 0.0356 0.0231 0.0186 0.0163	0.23 0.23 0.07 0.05 0.04 0.03
					765 1248 2080 8109	0.0101 0.0087 0.0069 0.0021	0.02 0.02 0.01 0.01
G4-22c	2407	VII	1.92	.27	0 15 138 277 464 592 979 1446 2256 8109	0.1305 0.1305 0.0231 0.0181 0.0124 0.0086 0.0064 0.0050 0.0045 0.0027	0.25 0.25 0.04 0.03 0.02 0.02 0.01 0.01 0.01 0.01
G4–23a	Commer Clinopt	cial ilolite	1.60	• 37	0 61 142 241 432 700 1116 1472 2829 8109	0.1824 0.1824 0.1481 0.1349 0.1261 0.1070 0.0870 0.0799 0.0740 0.0695 0.0565	0.29 0.29 0.24 0.22 0.20 0.17 0.14 0.13 0.12 0.11 0.09
G4-23b	Commer Clinop	cial tilolite	1.50	. 41	0 15 76 247 472 627 1088 1367 2191 3369 8109	0.1600 0.1600 0.1361 0.0973 0.0890 0.0818 0.0745 0.0631 0.0620 0.0585	0.24 0.24 0.20 0.16 0.15 0.13 0.12 0.11 0.09 0.09 0.09
G4-23c	Commer Clinop	cial tilolite	1.46	. 42	0 15 76 115 209 277 464 580 1239 2866 8109	0.2284 0.2284 0.2041 0.1649 0.1517 0.1412 0.1141 0.1048 0.0899 0.0827 0.0592	0.33 0.33 0.24 0.22 0.21 0.17 0.15 0.13 0.12 0.09
GU 3-1a	82	I-A	2.11	. 15	0 9 38 94 123 181 359 514 592 803 1200 9180	0.0594 0.0594 0.0509 0.0251 0.0219 0.0196 0.0178 0.0151 0.0133 0.0116 0.0107 0.0003	0.13 0.13 0.05 0.05 0.04 0.04 0.03 0.03 0.03 0.02 0.02 0.02

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Sample	Depth (ft)	Unit	Bulk Density (g/cm ³)	Porosity	Head (m)	Water g/g	Content vol/vol
GU 3–16	82	I-A	2.11	. 15	0 36 80 158 244 392 512 567 2557 9180	0.0604 0.0443 0.0293 0.0226 0.0226 0.0198 0.0184 0.0160 0.0060 0.0003	0.13 0.09 0.06 0.05 0.05 0.04 0.04 0.04 0.04 0.03 0.01 0.001
GU 3-2a	121	I-A	2.16	. 13	0 24 47 89 214 371 714 1877 2265	0.0404 0.0373 0.0328 0.0280 0.0234 0.0178 0.0099 0.0005	0.09 0.09 0.08 0.07 0.06 0.05 0.04 0.02 0.01
GU 3–2b	121	I-A	2.14	. 14	0 20 36 107 271 349 701 721 969	0.0420 0.0420 0.0314 0.0278 0.0242 0.0214 0.0066 0.0011 0.0005	0.09 0.09 0.07 0.06 0.05 0.05 0.01 0.01 0.01
GU 3-3a	155	I-A	2.31	.071	0 68 121 211 331 859 2279	0.0135 0.0135 0.0122 0.0105 0.0082 0.0040 0.0018	0.03 0.03 0.02 0.02 0.02 0.01 0.01
GU 3-36	155	I-A	2.22	. 11	0 101 172 317 933 2294	0.0208 0.0208 0.0183 0.0143 0.0097 0.0061	0.05 0.05 0.04 0.03 0.02 0.01
GU 3-4a	257	I-A	2.33	.065	0 15 94 209 260 352 520 630 1115 9180	0.0255 0.0255 0.0182 0.0160 0.0130 0.0128 0.0114 0.0088 0.0046 0.0029	0.06 0.04 0.04 0.03 0.03 0.03 0.03 0.02 0.01 0.01
GU 3–4 Þ	257	I-A	2.33	.065	0 13 69 274 359 467 510 1085	0.0287 0.0287 0.0175 0.0152 0.0117 0.0106 0.0066 0.0028	0.07 0.07 0.04 0.03 0.02 0.02 0.02 0.03

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Sample	Depth (ft)	Unit	Bulk Density (g/cm ³)	Porosity	Head (m)	Water g/g	Content vol/vol
GU 3-5a	316	I-A	2.26	.089	0 84 114 225 268 512 630 714 1204 1358 9180	0.0298 0.0246 0.0209 0.0147 0.0129 0.0112 0.0102 0.0093 0.0090 0.0076 0.0011	0.07 0.06 0.05 0.03 0.03 0.03 0.03 0.02 0.02 0.02 0.02
GU 3-56	316	I-A	2.27	.085	0 19 55 301 466 512 793 1168 1871 1938 9180	0.0306 0.0273 0.0206 0.0137 0.0130 0.0120 0.0101 0.0090 0.0074 0.0069 0.0011	0.07 0.07 0.06 0.05 0.03 0.03 0.03 0.02 0.02 0.02 0.02 0.02
GU 3-6 a	374	I-B	1.00	.59	0 39 196 403 455 562 1178 9180	0.3807 0.3807 0.0880 0.0750 0.0616 0.0475 0.0342 0.0107	0.38 0.38 0.09 0.08 0.06 0.05 0.03 0.01
GU 3бъ	374	I-B	1.00	.59	0 24 53 171 435 995 1518 9180	0.3904 0.3904 0.2855 0.1074 0.0783 0.0428 0.0428 0.0113	0.39 0.29 0.11 0.08 0.04 0.03 0.01
GU 3-7 a	378	I-B	1.48	. 37	0 47 58 65 84 425 788 1511 8160	0.2560 0.2560 0.1763 0.1441 0.0932 0.0570 0.0450 0.0027 0.0015	0.38 0.26 0.21 0.14 0.08 0.07 0.004 0.002
GU 3-76	378	IB	1.36	. 42	0 36 61 72 73 745 2713	0.3209 0.3209 0.2731 0.1816 0.1073 0.0495 0.0038	0,44 0,44 0,37 0,25 0,15 0,07 0,01
GU 3-8 a	397	I-B	1.42	. 42	0 10 80 103 822 2189 3896 8402 9180	0.2998 0.2998 0.2441 0.1096 0.0526 0.0378 0.0339 0.0281 0.0166	0.43 0.43 0.35 0.16 0.07 0.05 0.05 0.04 0.02

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Sample	Depth (ft)	Unit	Bulk Density (g/cm ³)	Porosity	Head (m)	water g/g	Content vol/vol
GU 3-86	397	I-B	1.34	. 45	0 30 119 520 859 1425 9180	0.3239 0.3239 0.0740 0.0604 0.0518 0.0358 0.0014	0.43 0.43 0.10 0.08 0.07 0.05 0.002
GU 3–9a	1132	II-NL	2.33	. 082	0 5 101 186 395 475 527 631 877 921 1109 9180	0.0304 0.0304 0.0212 0.0185 0.0166 0.0152 0.0134 0.0103 0.0096 0.0080 0.0064 0.0020	0.07 0.05 0.04 0.04 0.03 0.02 0.02 0.02 0.02 0.02 0.02 0.02
GU 3–9Ъ	1132	II-NL	2.33	.082	0 76 250 514 640 778 900 1056 1331	0.0298 0.0298 0.0215 0.0184 0.0156 0.0100 0.0092 0.0087 0.0074 0.0058	0.07 0.05 0.04 0.02 0.02 0.02 0.02 0.02 0.02
GU 3-10a	1197	III	2.33	.026	0 35 179 302 306 361 413 683 2382	0.0110 0.0110 0.0102 0.0089 0.0085 0.0081 0.0064 0.0055 0.0050	0.03 0.03 0.02 0.02 0.02 0.02 0.02 0.02
GU 3–10Þ	1197	III	2.34	.021	0 82 227 306 322 323 416 537 2708 8190	0.0094 0.0094 0.0082 0.0070 0.0064 0.0061 0.0048 0.0039 0.0032 0.0032	0.02 0.02 0.02 0.02 0.02 0.01 0.01 0.01
GU 3-11a	1246	III	2.22	.066	0 9 218 267 363 423 554 724 1146 3055 9180	0.0121 0.0121 0.0091 0.0055 0.0043 0.0038 0.0038 0.0036 0.0030 0.0029 0.0007	0.03 0.02 0.02 0.01 0.01 0.01 0.01 0.01 0.01
GV 3-11b	1246	111	2.20	.074	0 10 212 385 593 661 1060 1765 9180	0.0201 0.0165 0.0129 0.0118 0.0063 0.0056 0.0043 0.0007	0.04 0.04 0.03 0.03 0.01 0.01 0.01 0.01

TABLE A.1. (cont)

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		_		Water Co	
Sample Depth (ft)	Unit Bulk Density (g/cm ³)	Porosity	Head (m)	g/g	vol/vol
GU3-12a 1311	IV-A-v 1.72	.27	0 23 67 134 168 346 1867 9180	0.1359 0.1359 0.0601 0.0411 0.0155 0.0107 0.0065 0.0038	0.23 0.23 0.10 0.07 0.03 0.02 0.01 0.01
GU3-12b 1311	IV-A-v 1.37	. 42	0 19 58 94 158 227 1800 2243 9180	0.2885 0.2885 0.2136 0.0739 0.0462 0.0281 0.0236 0.0083 0.0020	0.40 0.29 0.10 0.06 0.04 0.03 0.01 0.003
GU3-13a 1331	IV-A-v 1.29	. 45	0 26 239 2179 8190	0.1929 0.1929 0.0192 0.0145 0.0024	0.25 0.25 0.02 0.02 0.03
GU3-13b 1331	IV-A-v 1.29	. 45	0 28 899 3713 8190	0.2009 0.2009 0.0197 0.0087 0.0025	0.26 0.26 0.03 0.01 0.003
GU3-14a 1440	IV-A-v 1.30	. 45	0 20 55 80 92 143 373 940 4436 8190	0.3185 0.3185 0.2355 0.1069 0.0750 0.0620 0.0280 0.0180 0.0090 0.0040	0.41 0.31 0.14 0.08 0.04 0.02 0.01 0.005
GU3-145 1440	IV-A-v 1.24	. 48	0 30 83 151 355 674 844 8160	0.3850 0.3850 0.1818 0.0590 0.0240 0.0160 0.0060 0.0030	0.48 0.48 0.23 0.07 0.03 0.02 0.01 0.003
GU3-15a 1499	IV-A-v 1.33	.43	0 15 162 298 672 1102 1457 1895 9180	0.2777 0.2777 0.0803 0.0607 0.0582 0.0494 0.0447 0.0406 0.0070	0.37 0.37 0.11 0.08 0.08 0.07 0.06 0.05 0.01
GU 3-156 1499	IV-A-v 1.33	. 43	0 10 40 117 247 731 2137 9180	0.3073 0.3073 0.1664 0.1405 0.1236 0.1144 0.0238 0.0070	0.41 0.22 0.19 0.16 0.15 0.03 0.01

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			TABLE	A.1. (con	t)		
GU 3-16a	1555	IV-B-v	1.29	. 47	0 14 65 156 699 1209	0.2814 0.2814 0.1722 0.0944 0.0463 0.0336	0.36 0.36 0.38 0.12 0.06 0.04
GU 3-166	1555	IV-B-v	1.29	. 47	0 52 64 76 1280	0.2707 0.2607 0.1663 0.0996 0.0347	0.35 0.35 0.21 0.13 0.04
GU 3-17a	1628	V	1.57	• 39	0 31 63 207 592 1553 2740 4065 9180	0.2243 0.2243 0.2193 0.0200 0.0120 0.0080 0.0065 0.0060 0.0027	0.35 0.35 0.33 0.01 0.01 0.01 0.01 0.01 0.01
GU 3-176	1628	v	1.57	. 39	0 32 40 64 64 918 1254 9180	0.2392 0.2263 0.1729 0.0583 0.0054 0.0035 0.0028 0.0007	0.38 0.38 0.27 0.09 0.01 0.01 0.004 0.001
GU 3-18a	1680	V	1.73	. 33	0 12 25 46 53 156 260 470 516 1031 3284	0.1755 0.1755 0.1605 0.2266 0.0388 0.0272 0.0189 0.0161 0.0140 0.0112 0.0096 2.83E-5	0.30 0.28 0.22 0.07 0.05 0.03 0.03 0.02 0.02 0.02 4.90E-5
GU 3-186	1680	V	1.76	.31	0 19 54 76 158 356 580 670 813 1817	0.1392 0.1392 0.0890 0.0215 0.0148 0.0125 0.0112 0.0080 0.0073 0.0060	0.25 0.25 0.16 0.04 0.03 0.02 0.02 0.02 0.02 0.01 0.01
GU 3-19a	1730	v	1.91	.24	0 20 194 378 538 672 944 1734 9180	0.1170 0.0299 0.0255 0.0226 0.0199 0.0178 0.0140 0.0106 0.0013	0.22 0.22 0.06 0.05 0.04 0.04 0.03 0.03 0.03 0.02 0.002
GU 3–19Ъ	1730	V	1.92	.24	0 20 146 306 474 624 734 836 1049 1668 9180	0.1152 0.0283 0.0242 0.0219 0.0196 0.0179 0.0171 0.0144 0.0117 0.0015	0.22 0.22 0.05 0.04 0.04 0.03 0.03 0.03 0.03 0.02 0.003

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TABLE A.2. Properties for Samples Taken from USW G-4 and USW GU-3

Sample Code	Depth (ft)	Unit	Grain Density (g/cm ³)	Porosity	Sat. Hydraulic Conductivity (m/s)	S r	Alpha (1/m)	Beta
G4-1	43	I-A [a]	2.49	.08 [b]	9.7E-12	0.0020	0.821E-02	1.558
G4-2	124	I-8 [c]	2.40	.27	2.6E-11 - 2.9E-12	0.0400	0.305E-01	1.220
G4-3	208	I-B	2.45	. 65	2.4E-6	0.1053	0.158E-01	10.563
G4-4	247	II-L	2.57 [d]	.03	8.6E-13	0.0600	0.152E-01	1.400
G4-5	864 (dark)	II-NL	2.54	.09	2.2E-11 - 1.3E-11 [e]	0.0662	0.985E-02	1.561
G4-24	864 (lt.)	II-NL	2.47	. 16	3.9E-11	0.0578	0.119E-01	1.945
G4-6	1158	II-NL	2.58	.11	1.9E-11	0.0801	0.567E-02	1.798
G4-1F	1215	II—NL	2.58 [f]	.11	9.2E-12 - 3.8E-12	0.1199	0.602E-02	1.916
G4-7	1256	II-NL	2.53	. 10	<1.3E-11 [g]	0.0704	0.258E-02	1.907
G4-2F	1278	II—NL.	2.50	.06	1.5E-11 - 1.3E-12	0.1198	0.372E-02	2.116
G4-8	1299	III	2.50	.09	4.5E-10	0.0517	0.419E-02	1.622
G4-9	1324	III	2.37	.04	3.0E-12	0.0026	0.169E-02	1.708
G4-3F	1359	IV-A-v	2.39 [h]	.21	8.0E-11 - 2.0E-11	0.1638	0.265E-01	2.223
G4-10	1405	IV-A-z	2.35 [1]	.41	3.0E-12	0.0100	0.220E-01	1.236
G4-11	1548	IV-A-z	2.23	.28	2.0E-11 - 2.4E-14	0.1095	0.308E-02	1.602
G4-4F	1551	IV-A-z	2.37 [j]	. 33	5.1E-11 - 1.3E-11	0.2017	0.415E-02	1.894

Notes: a) Mineralogical analyses indicate probable vapor-phase alteration of matrix.

b) Average value listed, see Table A.1 for values.

- c) Mineralogical analyses indicate devitrified, nonwelded sample. d) Value estimated based on grain density test of G-4 sample from
- 251 feet of depth.e) If multiple tests were made, the range is indicated, see Table A.4.
- f) Value estimated based on grain density test of G-4 sample from 1241 feet of depth.
- g) Actual value below this limit.
- h) Value estimated based on grain density test of G-4 sample from 1373 feet of depth.
- i) Value estimated based on grain density test of G-4 sample from 1410 feet of depth.
- j) Value estimated based on grain density test of G-4 sample from 1550 feet of depth.

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Sample Code	Depth (ft)	Unit	Grain Density (g/cm ³)	Porosity	Sat. Hydraulic Conductivity (m/s)	S r	Alpha (1/m)	Beta
G4-12	1686	IV-A-z	2.25	. 30	4.24E-12	0.0600	0.600E-02	1.460
G4-13	1728	IV-B-z	2.32	.22	2.5E-11 - 4.7E-14	0.1500	0.158E-02	1.685
G4-14	1737	IV-B-z	2.33	.24	2.5E-11 - 4.6E-13	0.1000	0.370E-02	1.496
G4-15	1769	IV-C-z	2.38	.26	2.3E-12	0.2154	0.605E-03	2.487
G4-16	1778	IV-C-z	2.31	.25	6.5E-12	0.1330	0.425E-02	1.560
G4-5F	1778	IV-C-z	2.31	.28	2.3E-11 - 6.9E-12	0. 1939	0.120E-02	3.322
G4-17	1787	IV-C-z	2.32	.28	1.6E-10 - 1.7E-12	0.0370	0.286E-02	1.675
G4-18	1899	V	2.59	.24	1.4E-8 - 1.6E-9	0.0658	0.141E-01	2,639
G4-19	2006	VI	2.29	. 29	2.0E-11	0.1346	0.316E-02	2.019
G4-20	2101	VI	2.24	. 19	4.4E-10	0.3217	0.448E-02	1.872
G4-21	2401	IIV	2.62	.24	2.3E-9	0.0608	0.112E-01	4.148
G4-22	2407	VII	2.63	.27	6.3E-9	0.0559	0.293E-01	2.257
G4-23 Anaconda	1010A,	 Commercial	2.53 Clinoptilo	.40 lite	<1.3E-11	0. 1897	0.134E-01	1.407

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Sample Code	Depth (ft)	Unit	Grain Density (g/cm ³)	Porosity	Sat. Hydraulic Conductivity (m/s)	S r	Alpha (1/m)	Beta
GU 3-1	82	I-A	2.49	. 15	<1.5E-12	0.0535	0.231E-01	1.693
GU 3-2	121	I-A	2.49	. 14	7.0E-13	0.0139	0.701E-02	1.851
GU 3-3	155	I-A	2.49	.09	2.7E-12	0.1600	0.389E-02	2.130
GU 3-4	257	I-A	2.49	.06	4.8E-09 [k]	0.0200	0.168E-01	1.500
GU 3-5	316	I-A	2.48	.09	<1.5E-12	0.0180	0.137E-01	1.513
GU 3-6	374	I-B	2.41	.59	1.6E-06	0.0835	0.144E-01	2.528
GU 3-7	378	I –B	2.35	. 40	3.9E-07	0.1001	0.150E-01	6.872
GU 3-8	397	I-8	2.44	. 43	3.5E-07	0.1145	0.110E-01	8.878
GU 3-9	1132	II-NL	2.54	.08	<1.5E-12	0.0075	0.123E-01	1.488
GU 3-10	1197	III	2.39	.02	1.5E-11	0.3757	0.328E-02	4.118
GU 3-1 1	1246	III	2.38	.07	1.5E-12	0.0804	0.441E-02	2.058
GU 3-12	1311	IV-A-v	2.37	.35 [1]	3.2E-09	0.0497	0.172E-01	3.283
GU 3-13	1331	IV-A-v	2.36	. 45	2 .9 E-07	0.0479	0.103E-01	4.203
GU 3-14	1440	IV-A-v	2.37	. 46	2.7E-07	0.0405	0.160E-01	3.872
GU 3-15	1499	IV-A-v	2.35	. 43	2.6E-08	0.0200	0.440E-01	1.496
GU 3-16	1555	IV-B-v [m]	2.44	. 47	7.9E-8	0.1892	0.155E-01	10.140
GU 3-17	1628	v	2.58	• 39	6.9E-09	0.0180	0.144E-01	2.964
GU 3-18	1680	V j	2.57	. 32	1.3E-09	0.0665	0.314E-01	3.442
GU 3-19	17 30	۷	2.52	.24	1.2E-09 [k]	0.1239	0.211E-01	2.851

Notes: k) Flow pattern indicated cracks in sample. 1) Subsamples had porosities of .27 and .42. m) Minerological analysis indicates this sample is a nonzeolitized, devitrified 'inclusion' in this unit.

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	Bulk	Apparent (Skeletal)	Average Pore
	Density,	Density,	Diameter,
Sample	g/cm ³	g/cm ³	pm
G4-1	5.1755	5.9929	0.0156
G4-2	1.8715	2.6008	0.0259
G4-3	1.2409	2.1764	0.2472
G4-4		Data not available	
G4-5	2.3192	2.5902	0.0306
G4-24	1.8512	2.5552	0.1311
G4-6	2.3235	2.5922	0.0315
G4-7	2.3579	2.5681	0.0180
G4-8	2.2826	2.5104	0.0187
G4-8(b)	2.2724	2.5416	0.0153
G4-9		Data not available	
G4-10	1.6501	2.2863	0.0177
G4-11	1.6706	2.2648	0.0192
G4-12	1.7101	2.3781	0.0217
G4-13	1.7623	2.2773	0.0236
G4-14	1.8813	2.3443	0.0198
G4-15	1.8433	2.2440	0.0131
G4-16	1.6577	2.2527	0.0204
G4-17	1.6419	2.3662	0.0363
G4-18	1.9570	2.6050	0.0284
G4-19	1.5779	2.3005	0.0329
G4-20	1.8792	2.2441	0.0307
G4-21	2.5419	3.7244	0.0672
G4-22	1.9541	2.6303	0.1360
G4-23	1.5690	2.1422	0.0233

<u>TABLE A.3</u>. Summary of Mercury Intrusion Data Supplied by Micromeritics^(a)

(a) Micromeritics Instrument Corporation
 5680 Goshen Springs
 Norcross, Georgia 30093

(b) Two G4-8 subsamples were tested by Micromeritics.

TABLE A.4	Unconfined	Hydraulic (Conductiv	ity for
	Samples Tak	en from USW	IG-4 and	USW GU-3

Sample	Depth (ft)	Unit	Saturated Hydraulic Conductivity (m/s)					
			Ruska	Permeameter	r	Large Disc		
			Run #1	Run #2	Run #3	-		
G4-1	43	I-A	9.68E-12					
G4-2	124	I-B	2.63E-11	2.86E-12				
G4 - 3	208	I-B	2.35E-06					
G4-4	247	II-L	8.60E-13					
G4-5	864	II—NL	2.17E-11	1.31E-11	1.97E-11	2.17E-11		
	(dark)							
G4-24	864	II-NL	3.90E-11					
	(light)							
G4-6	1158	II-NL	1.86E-11					
G4–1F	1215	II-NL	3.77E-12	7.11E-12	9.19E-12			
G4-7	1256	II-NL	<1.31E-11[a]					
G4-2F	1278	II-NL	8.58E-12	1.34E-12	1.48E-11			
G4-8	1299	III	4.47E-10					
G4-9	1324	III	3.02E-12					
G4–3F	1359	IV-A-v	3.25E-11	1.99E-11	8.03E-11			
G4-10	1405	IV-A-z	2.99E-12					
G4-11	1548	IV-A-z	<1.31E-11[a]	5.9 E-12	1.97E-11	2.37E-14		
G4-4F	1551	IV-A-z	5.06E-11	1.88E-11	1.33E-11			
G4-12	1686	IV-A-z	4.24E-12					
G4-13	1728	IV-B-z	1.86E-11	2.45E-11	1.97E-11	4.69E-14		
G4-14	1737	IV-B-z	<1.31E-11[a]	4.59E-13	2.48E-11	1.59E-12		
G4-15	1769	IV-C-z	2.30E-12					
G4 - 16	1778	IV-C-z	6.47E-12					
G4–5F	1778	IV-C-z	6.89E-12	1.83E-11	2.25E-11			
G4-17	1789	IV-C-z	1.61E-10	1.97E-11	1.24E-10	1.68E-12		
G4-18	1899	V	2.34E-09	1.38E-08	1.58E-09	4.46E-09		
G4-19	2006	VI	2.03E-11					
G4-20	2101	VI	4.36E-10					
G4-21	2401	VII	2.31E-09					
G4-22	2407	VII	6.26E-09					
G4-23	Anaconda	1010A	<1.31E-11[a]					

[a] Indicates upper limit.

Sample	Depth (ft)	Unit	Saturated Hydraulic Conductivit (m/s)	, y
			Ruska Permeameter L	arge Disc
			Run #1 Run #2 Run #3	
GU 3-1	82	I-A	<	(1.5 E-12[a]
GU 3-2	121	I-A		7.0 E-13
GU 3-3	155	I-A		2.66E-12
GU 3-4	257	I-A		4.83E-09
GU 3–5	316	I-A		(1.5 E-12[a]
GU 36	374	I-B		1.63E-06
GU 3-7	378	I-B		3.90E-07
GU 3-8	379	I –B		3.52E-07
GU 3-9	1132	II-NL		(1.5 E-12[a]
GU 3-10	1197	III		1.46E-11
GU 3-11	1246	III		1.52E-12
GU 3-12	1311	IV-A-v		3.15E-09
GU 3-1 3	1331	IV-A-v		2.92E-07
GU 3-14	1440	IV-A-v		2.68E-07
GU 3-15	1499	IV-A-v		2.57E-08
GU 3-16	1555	IV-B-v		7.90E-08
GU 3-17	1628	V		6.92E-09
GU 3-18	1680	V		1.28E-09
GU 3-19	1730	V		1.18E-09

TABLE A.4. (cont)

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[a] Indicates upper limit.

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FIGURE A.1. Technique for Measuring Sample Saturated Conductivity as a Function of Elevated Confining Pressure

- 1. Sample preparation.
 - a. Measure sample length and diameter.
 - b. Jacket sample with end caps as per standard procedure for low temperature samples.
- Conductivity system preparation.
 Note: Concurrent with sample preparation.
 - a. Clean system as necessary to remove contaminants.
 - b. Configure permeability apparatus as per Figure 5a for normal permeability measurement or as per Figure 5b for low permeability measurement depending on anticipated sample permeability.
 - c. Fill and flush system with J-13 well water.
 - d. Pressurize accumulator with nitrogen gas to one half the pore fluid operating pressure specified by the test plan.
- 3. Sample Installation.
 - a. Assemble sample stack and test apparatus as required by test plan on base plug as per standard procedure.
 - b. Leak check sample stack and correct any leaks found.
 - c. Vacuum saturate sample stack with J-13 well water.
 - d. Install differential pressure transducer, calibrated for anticipated range.
 - e. Raise base plug with sample stack into pressure vessel and complete connections as per standard operating procedure.
- 4. Vessel pressurization.
 - a. Fill confinement pressure vessel as per standard operating procedure.
 - b. Pressurize vessel to initial confining pressure specified by test plan as per standard operating procedure.
 - c. Introduce pore fluid to sample and increase pore fluid pressure to specified value.

- 5. Test data collection.
 - a. Start data collection using DAS and X-Y recorder.
 - b. Adjust pore fluid pump speed and/or back pressure valve to obtain specified pore pressure and a reasonable flow or sample differential pressure.
 - c. Vary confining pressure as per test plan and collect data at each pressure specified. Adjust pump speed or back pressure valve as necessary to obtain differential pressures within calibrated range.
- 6. System shutdown.
 - a. Stop data collection.
 - b. Depressurize sample pore fluid to atmospheric pressure.
 - c. Depressurize pressure vessel as per standard operating procedure.
 - d. Disengage sample stack assembly as per standard operating procedure.

P-C (Bars)	<u>P-E (Bars)</u>	Flow (mL/Sec)	D.P. (PSI)	<u>K (Millidarcy)</u>
100	50	0.0043	172	0.031
200	150	0.0043	196	0.027
400	350	0.0043	217	0.024
700	650	0.0043	233	0.023
960	910	0.0043	265	0.020
700	650	0.0043	261	0.020
400	350	0.0043	255	0.021
200	150	0.0043	250	0.021
100	50	0.0043	235	0.022

<u>TABLE A.5</u> Data and Calculated Values for Confined, Saturated Conductivity Tests of Sample GU3-2

Column Heading Explanation - Tables A.5 - A.7

- P-C Confining fluid pressure
- P-B Effective pressure
- Flow Fluid flow rate through the sample
- D.P. Pressure drop across the sample
 - K Calculated permeability

Note: At 20°C with water's density at 1 g/cm³ the relationship between the permeability unit of "millidarcy" and the conductivity unit of "m/s" is: 1 millidarcy = 1.02×10^{-8} m/s

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<u>P-C (Bars)</u>	<u>P-E (Bars)</u>	Flow (mL/Sec)	<u>D.P. (PSI)</u>	<u>K (Millidarcy</u>)
100	50	0.0055	3.5	20.2
200	150	0.0055	4.0	17.7
300	250	0.0055	5.1	13.9
400	350	0.0055	11.0	6.4
600	550	0.0055	34.0	2.1
800	750	0.0055	72.0	1.0
1000	950	0.0011	41.0	0.3
800	750	0.0011	41.0	0.3
600	550	0.0011	39.0	0.4
400	350	0.0011	38.0	0.4
200	150	0.0011	29.0	0.5
100	50	0.0011	23.0	0.6

<u>TABLE A.6</u> Data and Calculated Values for Confined, Saturated Conductivity Tests of Sample GU3-15

(See Table A.5 for column headings explanation)

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<u>P-C (Bars)</u>	<u>P-E (Bars)</u>	Flow (mL/Sec)	<u>D.P. (PSI)</u>	<u>K (Millidarcy</u>)
100	50	0.049	129	0.384
100	50	0.034	104	0.333
200	150	0.034	243	0.054
200	150	0.018	154	0.120
100	50	0.018	115	0.160
100	50	0.031	199	0.155
200	150	0.018	187	0.098
300	250	0.018	386	0.050
300	250	0.011	237	0.046
300	250	0.007	192	0.041
300	250	0.007	201	0.395
200	150	0.007	172	0.046
100	50	0.007	120	0.066
300	250	0.007	223	0.035
300	250	0.005	147	0.034
400	350	0.005	206	0.023
400	350	0.003	180	0.020
300	250	0.003	156	0.023
200	150	0.003	130	0.028
300	250	0.003	151	0.024
500	450	0.003	292	0.012
500	450	0.002	241	0.011
500	450	0.002	215	0.010
300	250	0.002	180	0.014
100	50	0.002	80	0.028
100	50	0.004	111	0.044
100	50	0.002	62	0.037
300	250	0.002	112	0.020
500	450	0.002	202	0.011
500	450	0.001	119	0.009
600	550	0.001	22	0.005
600	550	0.001	532	0.001

TABLE A.7 Data and Calculated Values for Confined, Saturated Conductivity Tests of Sample GU3-3

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(See Table A.5 for column headings explanation)

0	1 P-C (BARS)	2 P+UP (BAR\$)	3 P+DN (BARS)	4 DIFF-P	5 FLOW (ML/SEC)	6 P-P (BARS)	7 P-EFF (BARS)	8 APTR (MICKONS)	9 NORM 4PTR	10 PERM (DARCY)	11 NURM Perm
1	35	28	27	0.367	0.385	27	8	67	1.00	378	1,00
5	40	29	28	0.389	0.377	29	11	65	0,97	359	0.95
3	50	28	27	0.441	0.376	28	22	62	0,93	329	0.87
4	60	29	59	0.474	0.404	29	31	62	0.93	329	0.87
5	70	23	53	0.369	0.241	23	47	57	0.85	275	0.73
6	70	30	59	0.560	0.388	50	41	58	0.87	287	0.76
7	80	28	27	0.639	0.392	59	52	56	0,84	264	0.70
8 9	100	33	32	0.703	0.376	35	68	53	0.80	241	0.64
10	120 140	32 34	31	0.86A	0.412	32	88	51	0.77	553	0.59
11	150	34	32	1.179	0.397	33	107	46	0.68	177	0.47
12	140	35	32	1.319	0.392	33	117	44	0.66	163	0.43
13	120	32	33 30	1.793	0.391	34	106	40	0,59	132	0.35
14	100	32	30	1,926	0.381	31	89	38	0.57	124	0.33
15	80	34	31	1.616	0.384 0.387	31 32	69	41	0.61	140	0.37
15	70	33	35	1.348	0.400	33	48 37	42	0.63	148	0.39
17	60	31	29	1.210	0.407	30	30	44 46	0,66	163	0.43
18	50	31	29	0.933	0.406	30	20	50	0.68 0.75	177 210	0.47
19	40	29	28	0.745	0.398	28	12	53	0.80	241	0.56
20	35	28	27	0.575	0.387	28		58	0,86	281	0.64
Column Heading Explanation - Tables A.8 - A.12P-CConfining fluid pressureP-UPUpstream pore fluid pressureP-DNDownstream pore fluid pressureDIFF-P(P-UP - P-DN) Pressure drop across the sample											
			FLO P-P				_	he sample average po	re press	sure	
			P-E	FF	(P-C -	P-P) Eff	ective co	onfining pre	ssure		
			APT	ĸ	Apertur	e calcula	ited using	; "cubic law	" (Equat	ion 6)	
			NOR	M APTR	R Normalized aperture						
			PER	Н	Calcula	ted perme	ability (see Equatio	n 7)		
			NOR	M PERM	Normali	zed perme	ability				

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<u>TABLE A.8</u> Laboratory Data and Calculated Parameters for Confined, Saturated Conductivity Tests of Sample G4-1F

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<u>TABLE A.9</u> Laboratory Data and Calculated Parameters for Confined, Saturated Conductivity Tests of Sample G4-2F

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0	1 P-C (BARS)	2 UP=P	3 DN-P	4 DIFF PRESS	5 FLOW (ML/SEC)	6 P-P (BARS)	7 P=E (BARS)	8 APTR (MICRONS)	9 NORM Aptr	10 PERM (DARCY)	11 NURM PERM
1	38	20	16	3	0.00465	18	20	6	1,00	3,51	1.00
2	51	24	20	3	0.00465	22	29	6	0,95	3,14	0,89
3	56	28	23	4	0,00465	56	30	6	0,85	2.75	0,75
4	60	27	21	5	0,00465	24	36	5	0,82	5,36	0,67
5	68	30	24	5	0,00465	27	41	5	0,79	2,19	0,62
6	67	59	21	4	0,00232	23	a 4	4	0,69	1.67	0,47
7	83	25	20	4	0.00232	23	60	4	0,67	1,58	0,45
8	102	27	21	5	0,00116	24	78	3	0,49	0,85	0.24
9	104	27	21	5	0,00116	24	80	3	0,52	0.94	0.27
10	101	53	18	4	0,00116	20	81	4	0,56	1.12	0,32
11	120	24	19	4	0,00116	21	99	4	0,57	1,13	0,32
12	140	27	53	3	0.00116	25	115	4	0,60	1,26	0,36
13	148	30	25	3	0,00116	28	120	4	0,57	1.14	0.33
14	152	35	59	4	0,00116	29	123	3	0,53	1.00	0,29
15	130	33	29	4	0.00116	31	99	4	0,56	1,10	0,31
16	121	31	25	4	0,00116	28	93	4	0,56	1,12	0.32
17	99	33	59	3	0,00116	31	68	4	0,58	1,18	0.33
18	80	35	31	3	0.00116	33	47	4	0,61	1.32	0,38
19	71	37	33	3	0.00116	35	36	4	0,61	1,30	0.37
20	74	25	55	2	0,00116	24	50	4	0,65	1.47	0,42
21	62	23	24	2	0,00116	26	36	4	0,67	1.60	0,45
22	49	29	59	5	0,00116	27	22	5	0,73	1.89	0,54
23	42	30	28	1	0,00116	29	13	6	0,86	2.59	0,74
24	42	27	21	5	0,00232	24	18	4	0,66	1.51	0.43
25	36	25	19	4	0,00232	22	14	5	0,70	1.73	0,49

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0	1 P-C (BARS)	2 UP-P (BAPS)	3 DN-P	4 DIFF Press	5 FLOW (ML/SEC)	6 P-P (BARS)	7 P-E (BARS)	8 APTR (MICRONS)	9 NORM Aptr	10 PERM (DARCY)	11 NORM PERM
*****	45	20	16	5.95	0.0870	18	27	22	0.91	35	54.0
ż	40	22	19	2.18	0.0866	21	19	55	1.00	43	1.00
3	50	17	13	3,54	0,0860	15	35	19	0.85	31 39	0.72 0.92
ā.	50	28	25	2,45	0,0866	27	23	22	0.96	39	0.92
Ś	50	27	24	2,45	0.0860	25	25	22 20	0.95 0.87	32	0.76
6	60	28	24	3.33	0.0570	26	34	-		29	0.69
7	60	27	22	3.74	0.0850	24	36	19 18	0.83 0.80	27	0.64
8	60	26	20	4.22	0.0860	23	37	16	0.73	23	0.53
9	70	28	21	5.64	0.0866	24	46 43	16	0.71	21	0.50
10	70	59	25	3,06	0.0430	27	51	15	0.67	19	0.44
11	80	31	27	3.67	0.0430	29 28	52	13	0,57	14	0.32
12	80	31	24	5,85	0,0430	25	55	13	0,56	14	0,32
13	80	29	55	5.98	0.0429	26	54	13	0.57	14	0.32
14	80	85	24	2.92 1.36	0.0100	30	50	13	0.57	14	0.32
15	80	32	59	1.30	0.0100	23	57	12	0.53	12	0.28
16	80	24 27	22 24	2.04	0.0100	25	75	11	0.50	11	0,25
17	100	30	26	2,99	0.0100	28	92	10	0.ដដ	8	0.19
18	120 140	22	15	6,26	0.0100	18	155	5	0.34	5	0.12
19	140	25	23	1.84	0 0056	24	116	10	0.42	8	0.18
20	140	24	20	3.47	0.0055	22	118	8	0.34	5	0.12
21 22	150	29	25	3.40	0.0055	27	123	8	0.34	5	0.12
23	150	27	22	4.01	0.0056	25	125	7	0.33	2	0.11
24	140	30	26	3.33	0.0054	28	112	8	0.34	2	0.12
25	140	23	13	4.28	0.0053	21	119		0.31	· 5	0,10
56	120	24	20	3,33	0.0053	55	98		0.34	5	0.14
27	100	27	23	2.72	0,0057	25	75		0.37	8	0,18
28	80	58	25	1.90	0.0056	27	53		0.42	10	0.24
29	70	30	27	1,16	0.0055	28	42				
30	70	31	29	1.36	0.0055	30	40			11	0.27
31	70	24	22	0,95	0.0052	23	47				
32	70	23	50	1.50	0.0055	21	49 43			10	-
33	70	59	25	2.24	0.0109	27 19	51			9	
34	70	21	17	2.72	0.0110	19	52	-	· · · · · · · · · · · · · · · · · · ·	9	0,21
35	70	20	17	2.99	0.0111	23	47			9	0.21
36	70	27	20	5,51	0.0216	25	35			12	
37	60	27	23	3.81	0.0217 0.0218	25				16	0.37
38	50	27	23	2.45	0.0217	56				26	0.61
39	40	27	25 18	1.16 3.20	0.0436	20					
40	40	22		2.18	0.0431	28				27	
41	40	29 29	26 24	4.22							
42	40	29	24	4.08	0.0871	26					
43	40 40	23	17	5.64				16			
44	35	24	21	2.86					0.91	36	0,83
45	22	£.1	c. 1	2.00			_				

TABLE A.10 Laboratory Data and Calculated Parameters for Confined, Saturated Conductivity Tests of Sample G4-3F

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0	1 P=C (BARS)	2 UP=P	3 DN=P	4 DIFF PRESS	5 FLOW (HL/SEC)	6 P-P (BARS)	7 P-E (BARS)	8 APTR (MICRONS)	9 NORM Aptr	10 PERM (DARCY)	11 NORM Perm
1	35	21	17	1.904	0.3687	19	16	31	1.00	79	1.00
2	35	24	50	2,108	0.3690	55	13	30	0,97	74	0.93
3	40	23	50	2,584	0.3692	22	18	28	0,90	65	0.82
4	50	20	16	2,788	0.3696	18	32	27	0.88	61	0.78
5	50	28	23	3,400	0,3687	25	25	25	0.82	54	0.68
6	50	22	17	3.740	0.3686	19	31	24	0,60	50	0.64
7	65	25	16	4,556	0.3691	19	46	23	0.75	44	0,56
8	56	23	17	^ _556	0.3687	20	36	23	0,75	44	0,56
9	55	22	16	4.692	0.3695	19	36	23	0.74	43	0,55
10	60	21	14	6,256	0.3699	17	43	21	0,67	36	0,45
11	60	26	21	3,604	0.2214	24	36	21	0.68	37	0.47
12	60	23	18	4.080	0,2195	21	39	50	0,65	34	0.43
13	60	21	17	3.196	0.1746	19	41	20	0.66	34	0,43
14	60	22	19	1.700	0.0922	20	<i>a</i> 0	20	0.65	34	0.43
15	70	24	21	1.836	0.0925	23	47	20	0.64	32	0.41
16	80	24	21	1,972	0.0923	23	57	19	0.65	31	0.39
17	100	53	50	2.312	0.0919	21	79	18	0.59	28	0.35
18	100	24	19	3.400	0.0925	22	78	16	0.52	21	0.27
19	120	23	18	4.284	0.0926	20	100	15	0,48	18	0.23
50	140	26	19	5,304	2560.0	22	118	14	0.45	16	0.20
51	140	24	17	6,120	0.0924	20	120	13	0.43	14	0.18
55	140	25	50	3.740	0.0553	22	118	13	0.42	14	0,18
53	150	25	53	4.284	0.0548	26	124	12	0.40	13	0.16
24	150	23	17	4.964	0.0552	20	130	12	0.39	12	0.15
25	140	24	18	4.828	0.0550	21	119	12	0,39	12	0.15
56	120	24	19	4.420	0.0550	22	9.9	12	0,40	13	0.16
52	100	24	19	4.012	0.0554	22	78	13	0.41	14	0.17
85	100	59	23	4.284	0.0556	56	74	12	0.41	13	0.16
59	100	55	16	4.012	0.0557	19	81	13	0.42	14	0.17
30	80	25	21	3.536	0.0559	23	57	13	0,43	15	0.19
31	8 Q	24	19	3,332	0.0553	21	59	13	0.44	15	0.19
32	70	27	55	3,128	0.0553	24	46	14	0.45	16	0.20
33	60	27	55	2.380	0.0557	25	35	15	0.49	19	0.24
34	60	21	17	2.720	0.0556	19	41	14	0.47	18	0.25
35	50	21	17	1,972	0.0554	19	31	16	0,53	22	0.28
36	40	21	15	1,428	0,0564	19	21	18	0,59	27	0,35
37	35	21	17	1.292	0.0550	19	16	18	0.60	29	0,36

TABLE A.11 Laboratory Data and Calculated Parameters for Confined, Saturated Conductivity Tests of Sample G4-4F

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TABLE A.12	Laboratory Data and Calculated Parameters for Confined,
	Saturated Conductivity Tests of Sample G4-5F

0	1 P=C (BARS)	2 UP=P	3 DN-P	4 DIFF -P (BAR\$)	5 FLOW (ML/SEC)	6 P=P (BARS)	7 P-E (BARS)	8 APTR (MICRONS)	9 NORM APTR	10 PERM (DARCY)	11 NORH Рени
1	37	30	26	3.214	0_001962	28	8,999	6,0	0,98	3.0	0.97
2	50	31	26	4.840	0,001156	29	21,274	4.4	0.72	1.6	0.52
3	38	29	24	4,180	0.001503	29 27	11.272	5,0	0.82	2.1	0,68
4	80	27	23	3.944	0.000558	25	55,316	3,7	0.60	1.1	0.37
5	80	27	23	4,284	0.000420	25	55,214	3,2	0,54	0,9	0,27
6	80	26	22	3.536	0.000793	24	55 622	4.3	0,70	1,5	0.50
7	120	26	22	3,672	0.000448	24	95 724	3,5	0.58	1.0	0.33
8	116	25	22	2.312	0.000206	23	92.574	3,1	0,52	0.8	0.27
9	115	24	22	1.836	0.000133	23	92,118	2,9	0.48	0,7	0,23
10	150	25	22	2.516	0.000159	23	126,778	2.8	0.46	0.7	0.21
11	150	25	22	2,992	0.000163	23	126 778	2.7	0.44	0.6	0,19
51	150	25	51	3.332	0.000000	23	126,710	0.0	0.00	0.0	0.00
13	120	25	21	3.536	0.000194	23	96 778	2.7	0.44	0,6	0,19
14	80	25	21	3,944	0.000305	23	56 846	3.0	0.49	0.8	0.24
15	54	23	21	1.632	0.000209	22	32,240	3,5	0,58	1,1	0.34
16	50	23	21	2.040	0.000615	22	28,138	4.7	0.78	1,9	0.61
17	50	29	25	3,196	0.000479	27	22,868	3.7	0,62	1,2	0.38
18	35	30	27	2.040	0.001306	28	6,678	6.1	1,00	3,1	1.00

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(See Table A.8 for explanation of column headings)

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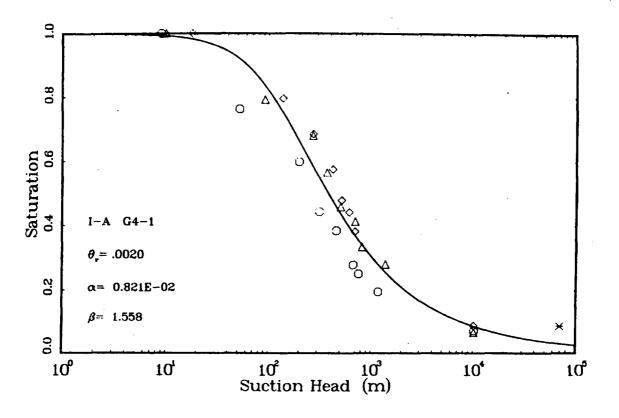
APPENDIX B SATURATION DATA PLOTS AND CURVE FITS

This appendix contains plots based on the psychrometer data contained in Table A.1. The saturation data points are, as discussed in the main body of the report, the water content divided by the maximum water content. The subsample data are keyed with all subsample "a" (e.g., G4-1a) points plotted as circles, subsample "b" points plotted as triangles, and subsample "c" (G-4 samples only) points plotted as diamonds.

These data were fit using the Van Genuchten function. The fitted curve is plotted and the parameters are listed.

The "stars" indicate data points added to stabilize the curve fit routine at a reasonable result.

The plots are in the order listed in Table 2.





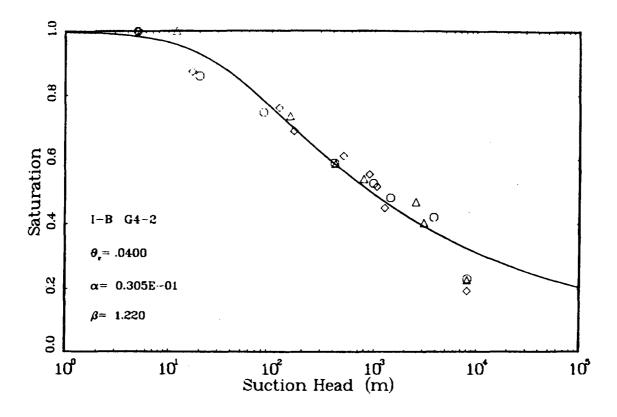
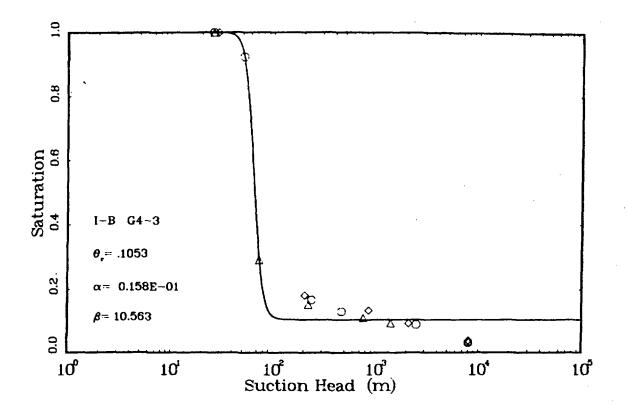


Figure B.2





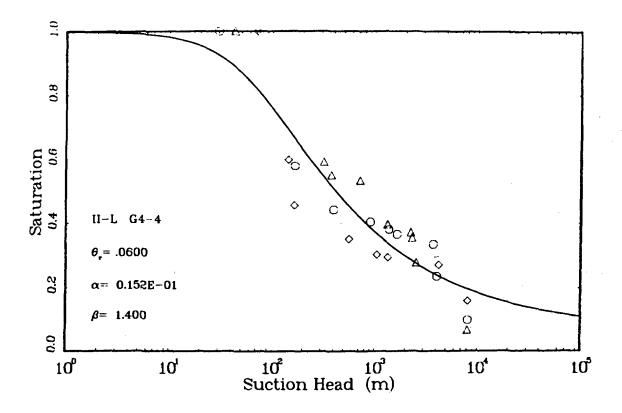


Figure B.4

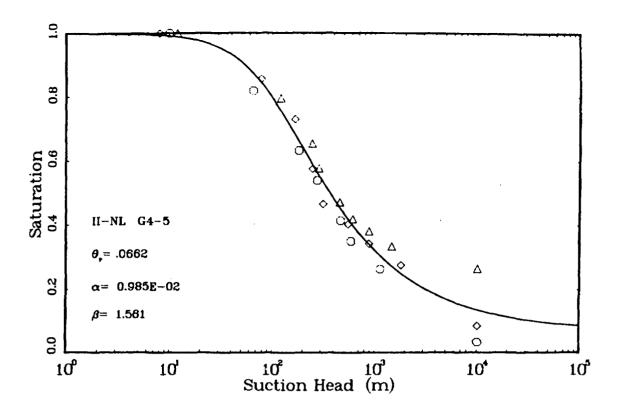


Figure B.5

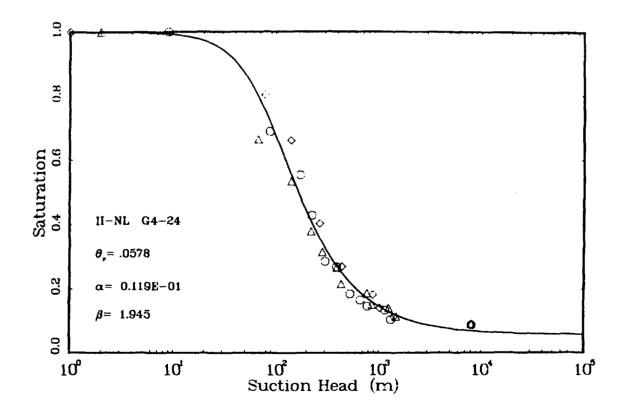


Figure B.6



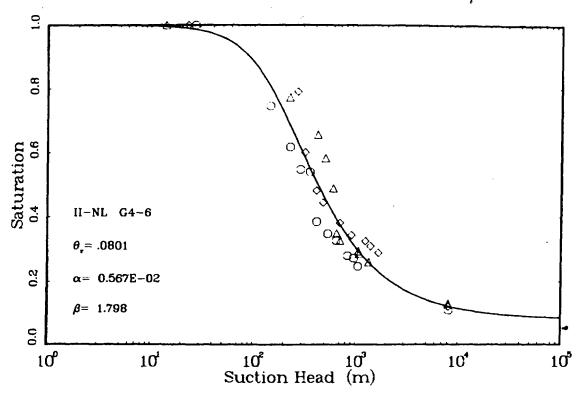


Figure B.7

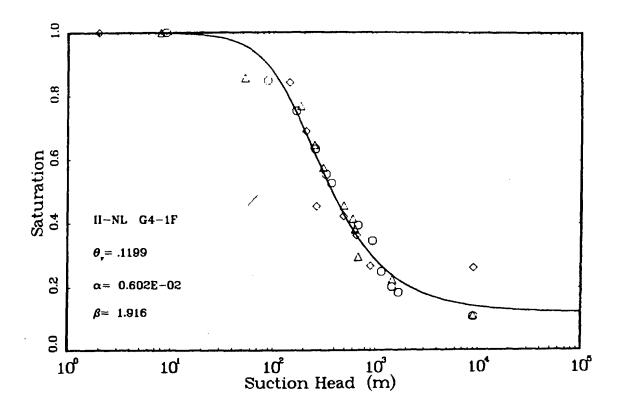


Figure B.8

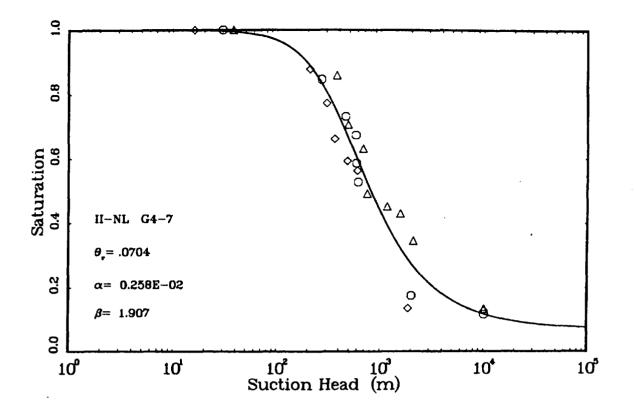


Figure B.9

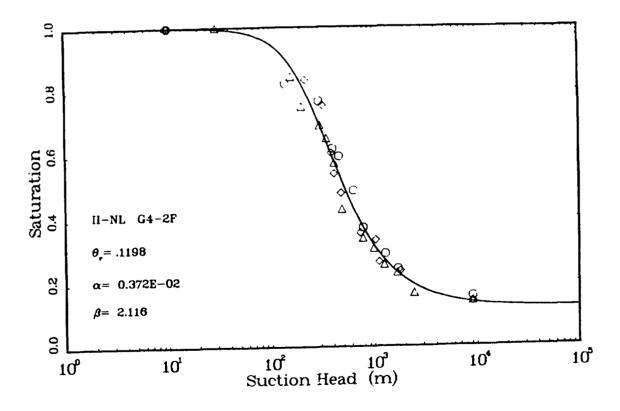


Figure B.10

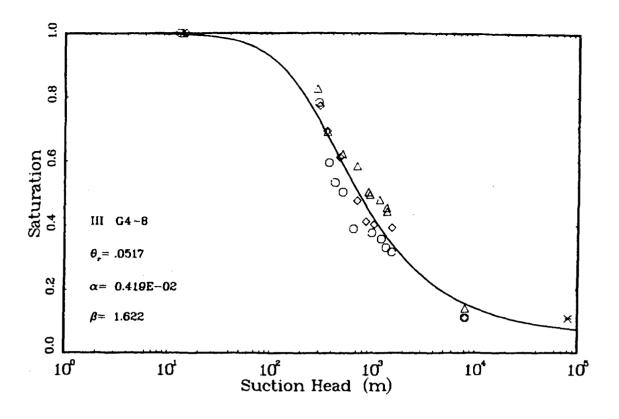


Figure B.11

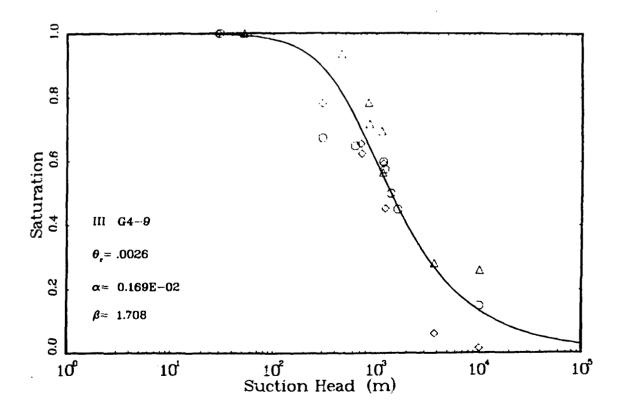


Figure B.12

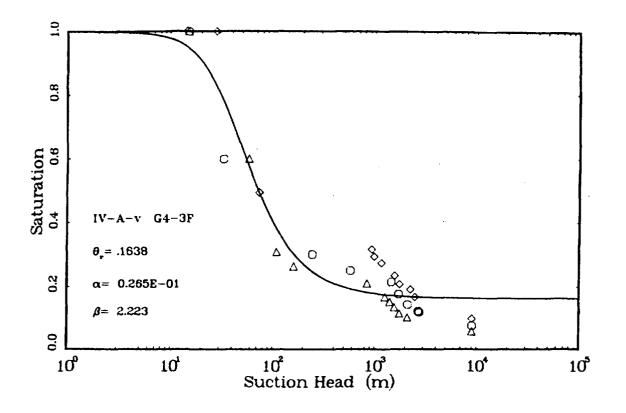


Figure B.13

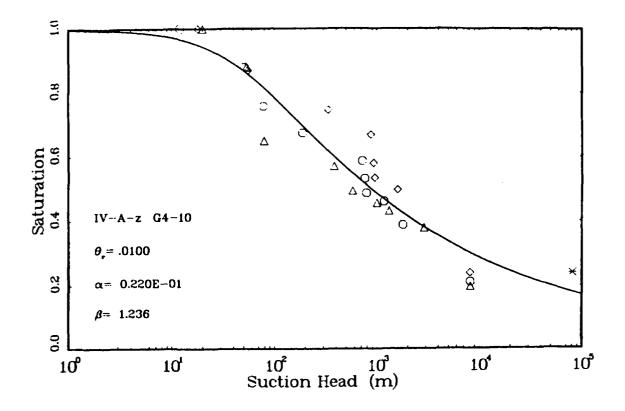


Figure B.14

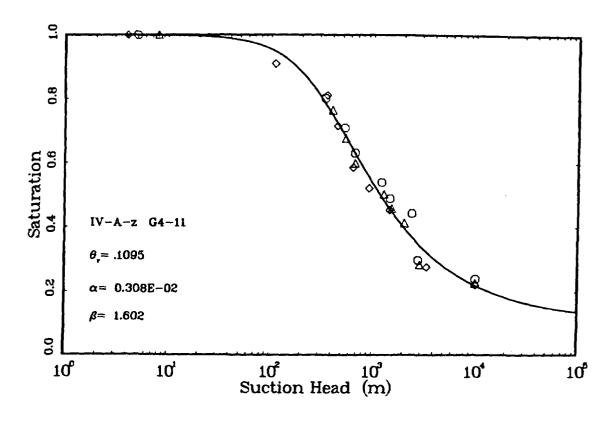


Figure B.15

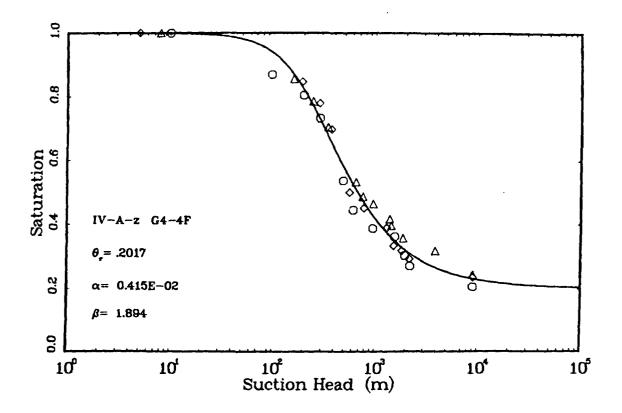


Figure B.16

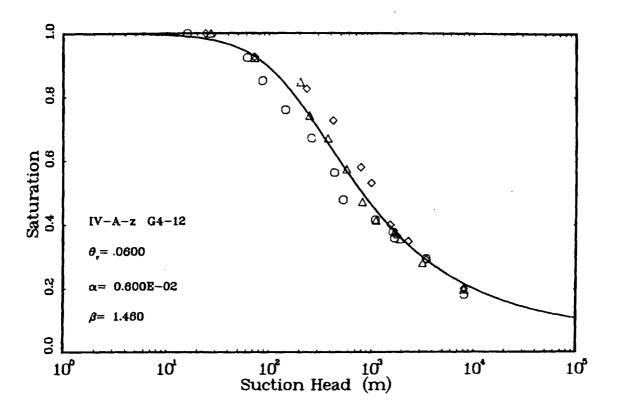


Figure B.17

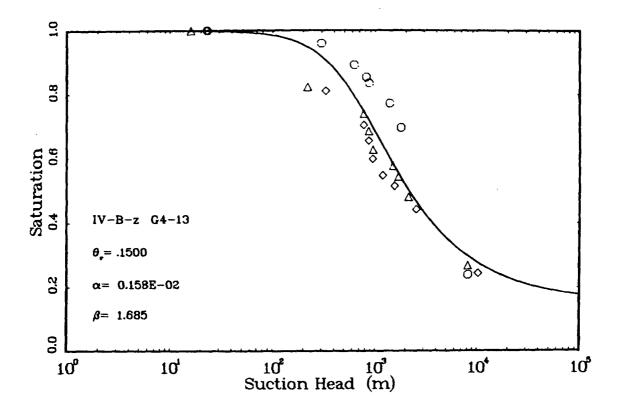


Figure B.18

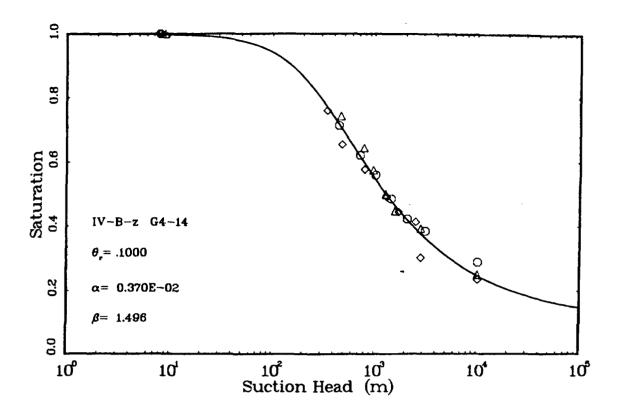
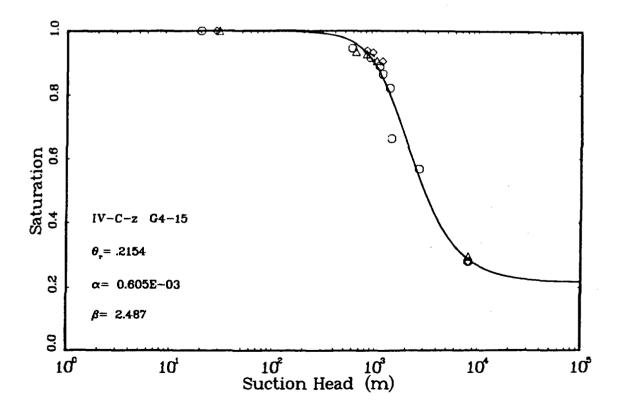


Figure B.19





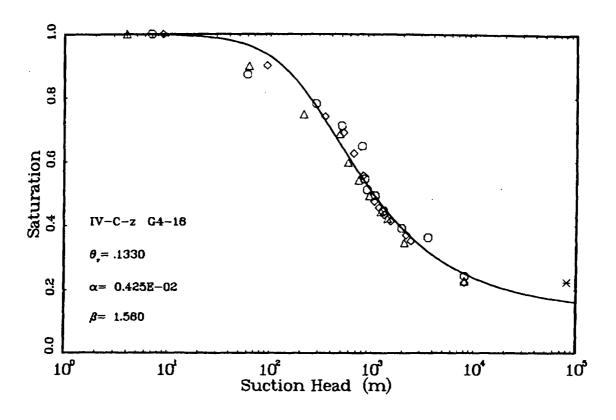


Figure B.21

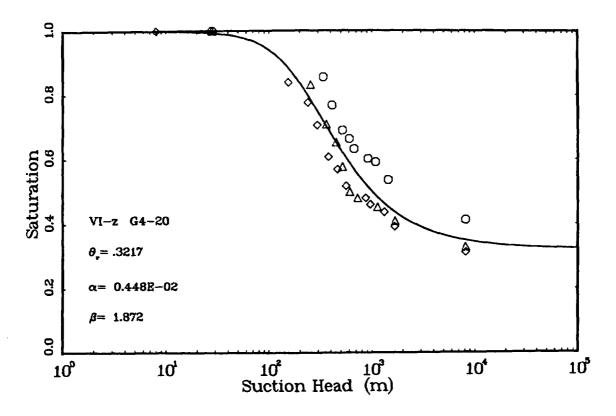
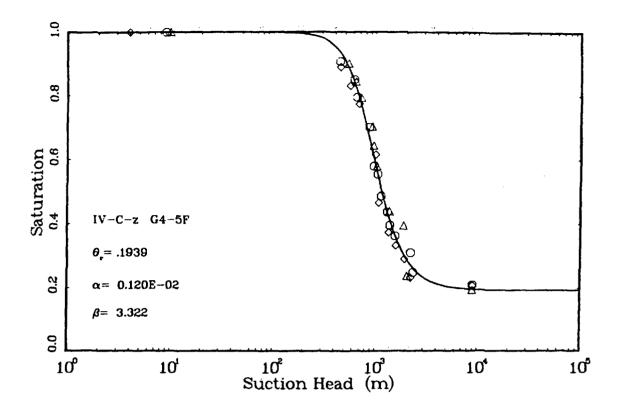
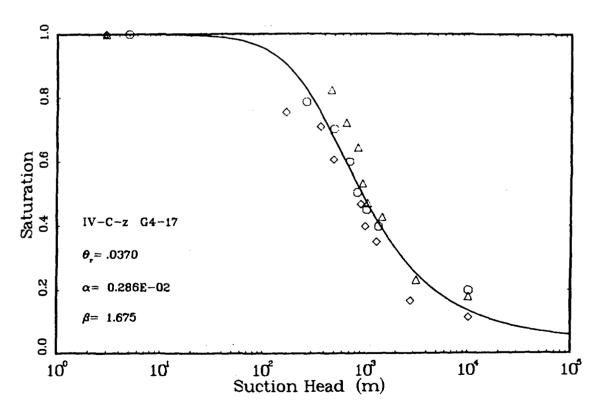


Figure B.22









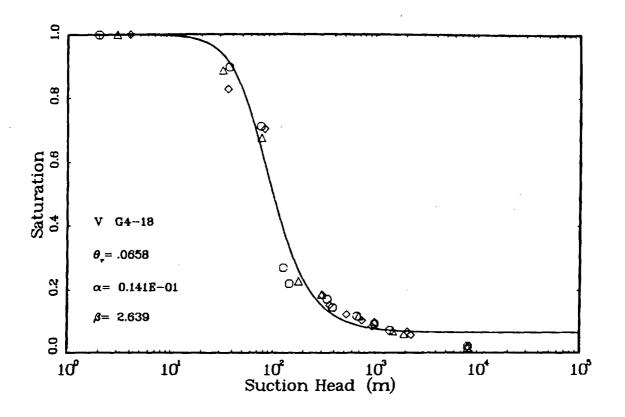


Figure B.25

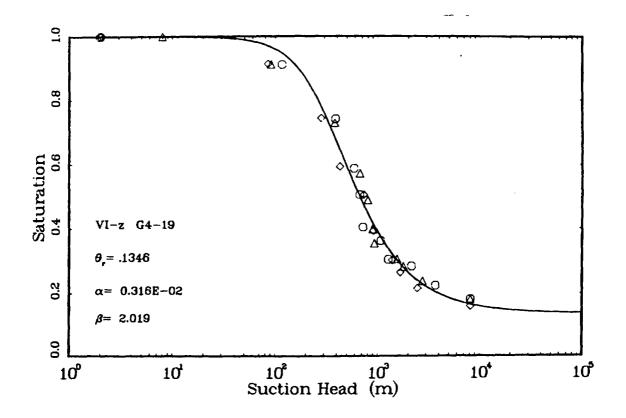


Figure B.26

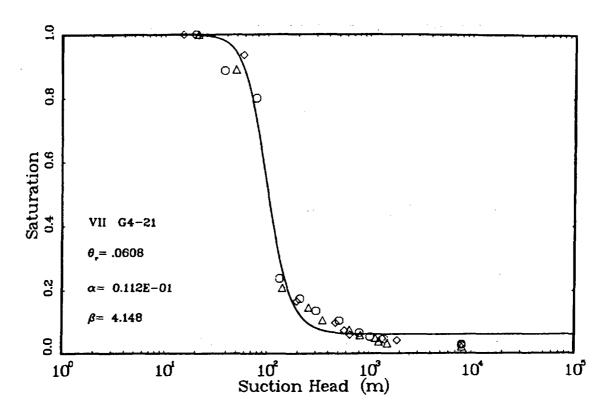


Figure B.27

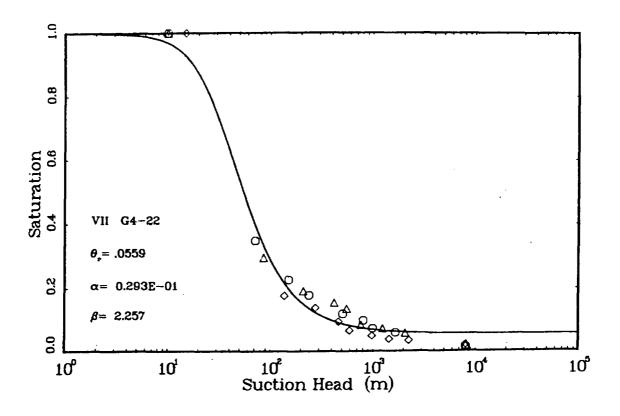
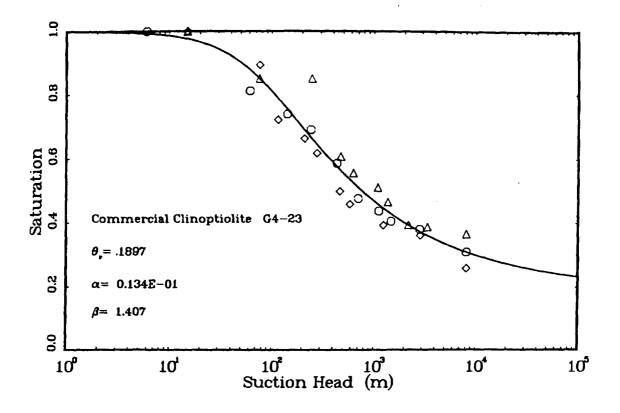
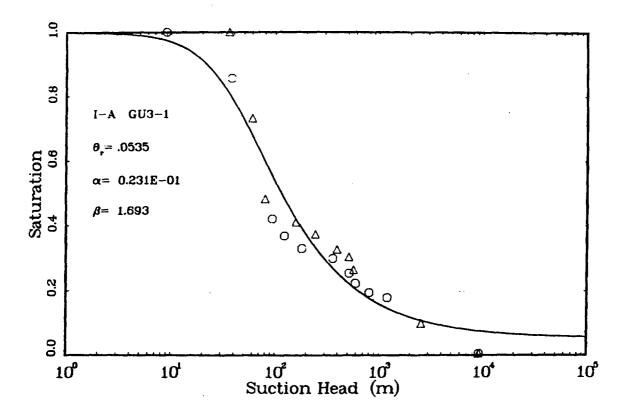
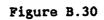


Figure B.28









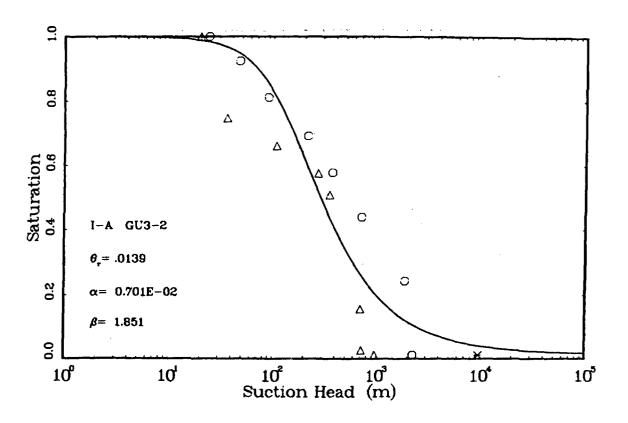


Figure B.31

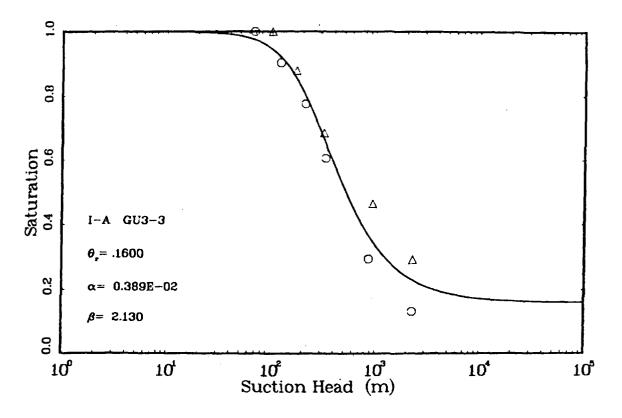
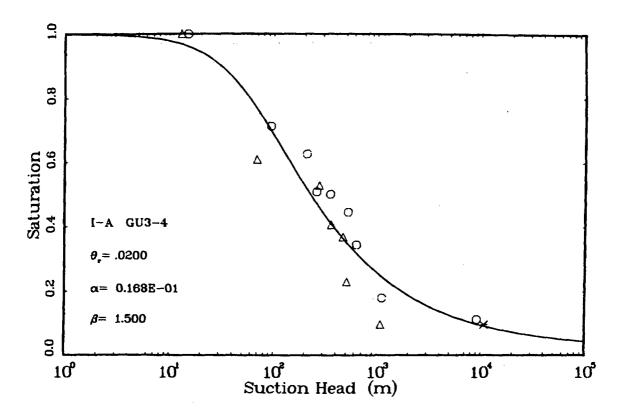


Figure B.32





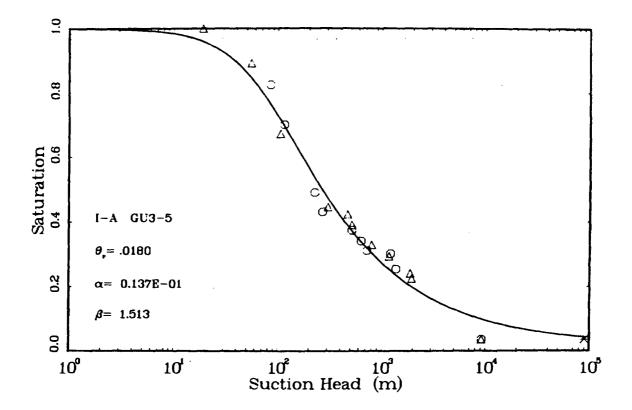
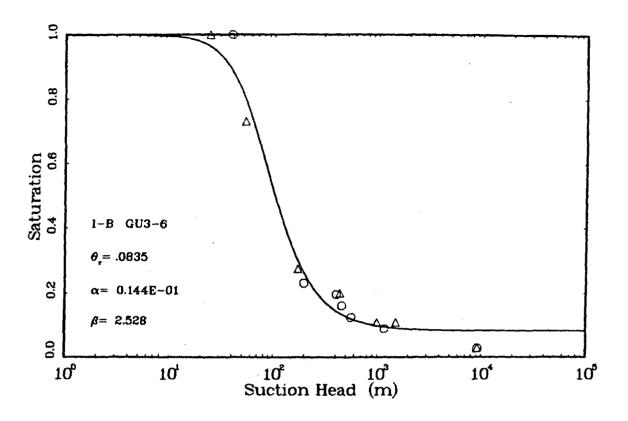


Figure B.34



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Figure B.35

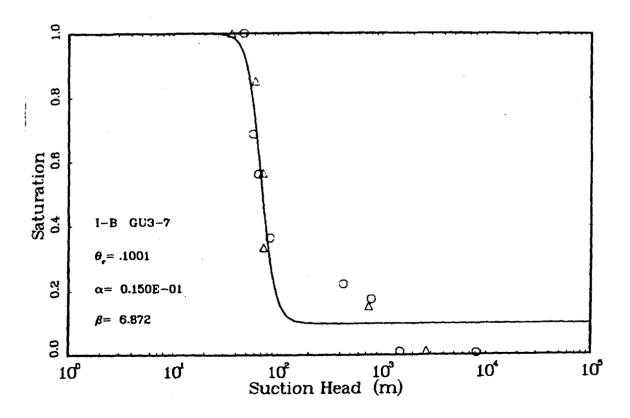
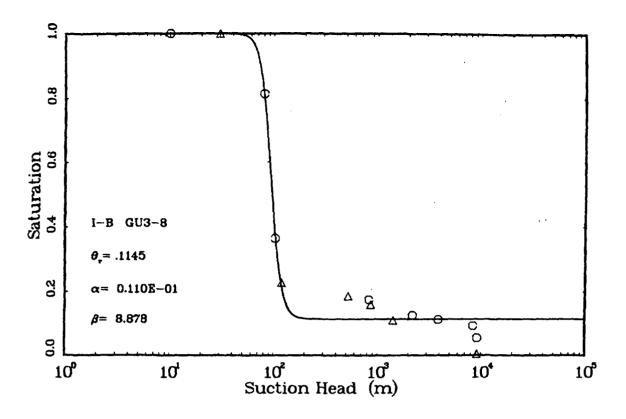


Figure B.36





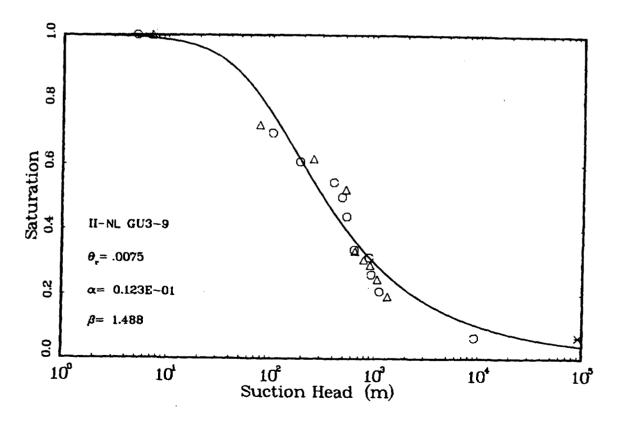


Figure B.38

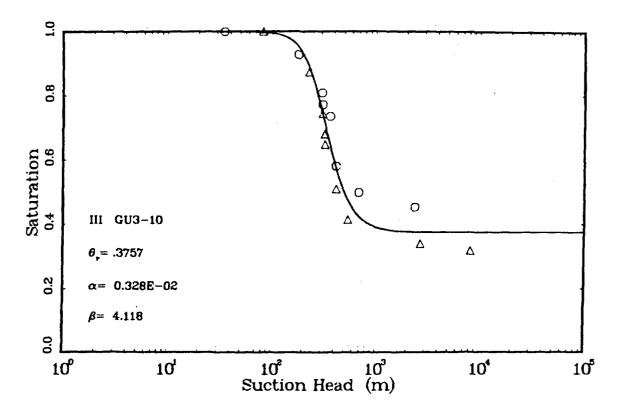


Figure B.39

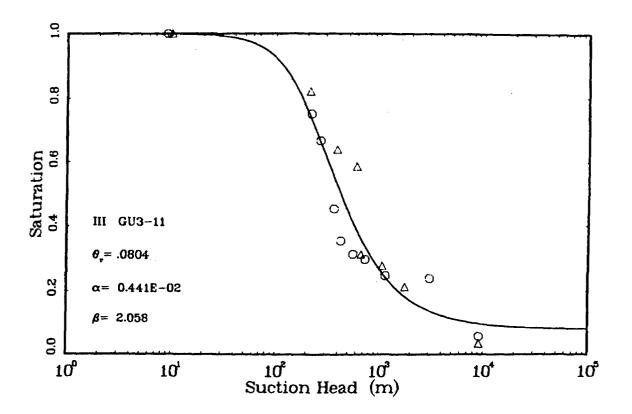
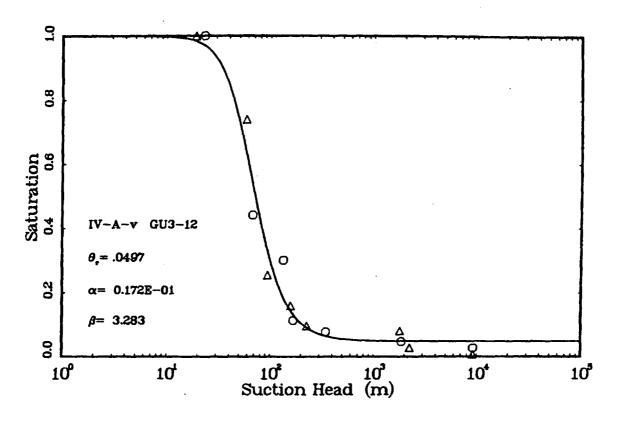


Figure B.40





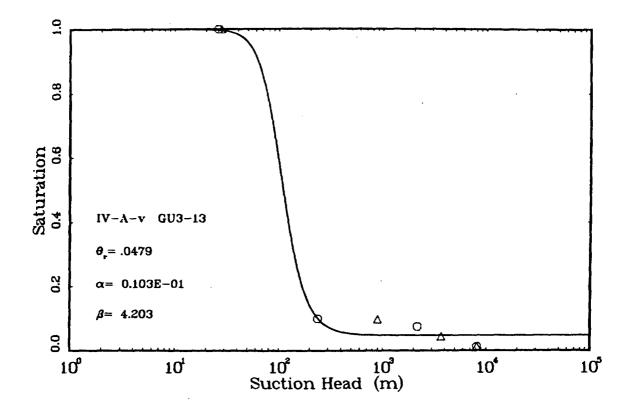


Figure B.42

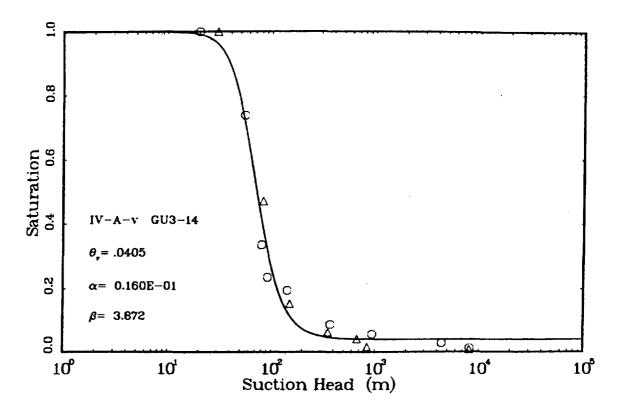


Figure B.43

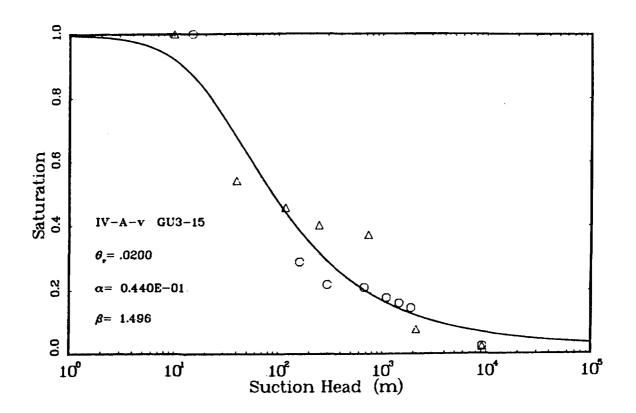
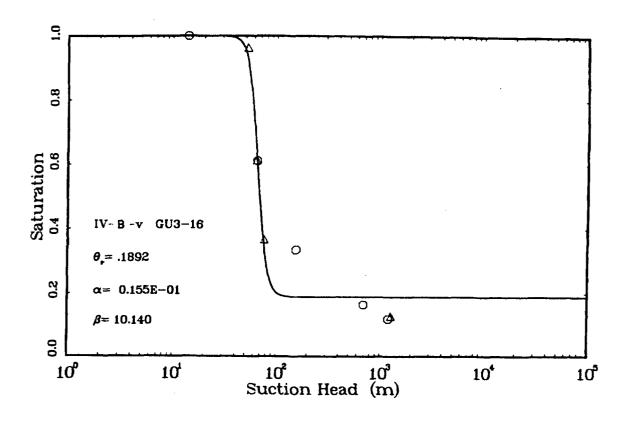


Figure B.44





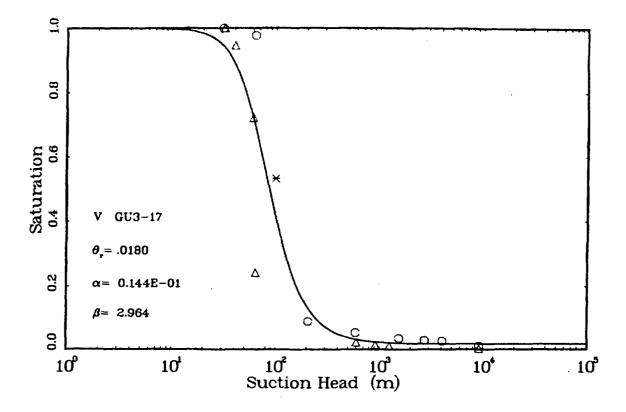


Figure B.46

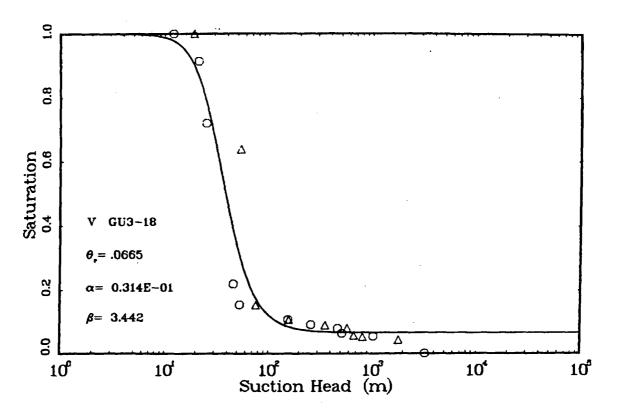
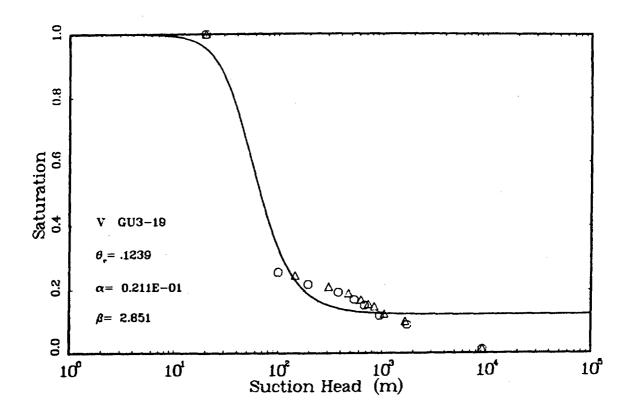
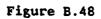


Figure B.47





APPENDIX C

COMPARISON OF SATURATION CURVES AS DETERMINED BY PSYCHROMETER DATA AND MERCURY-INTRUSION DATA

This appendix contains plots of the psychrometer data, the Van Genuchten curve fit of the psychrometer data, and the adjusted mercury-intrusion data. The manner in which these data were taken and analyzed is discussed in the text, particularly in the section TEST METHODS--WATER RETENTION and the first three subsections of RESULTS OF WATER RETENTION TESTING.

The root-mean-square (RMS) difference between the mercury-intrusion data points and the curve fit of the pyschrometer data in the suction head range of 10 to 3000 m (corresponding roughly to a pore size range of 50 to 15,000 angstroms) was calculated and the results are listed in Table C.1. The values listed in Table C.1 indicate the quantitative agreement between the mercury-intrusion data and the psychrometer data. The degree of agreement between the two testing methods indicated by this table agrees with that reached from a visual inspection of the figures. For example, the RMS difference for sample G4-5 is.05, which indicates good agreement; inspection of Figure C.4 confirms this conclusion. Sample G4-3 has a RMS difference of .37, which indicates poor agreement; inspection of Figure C.3 confirms this conclusion.

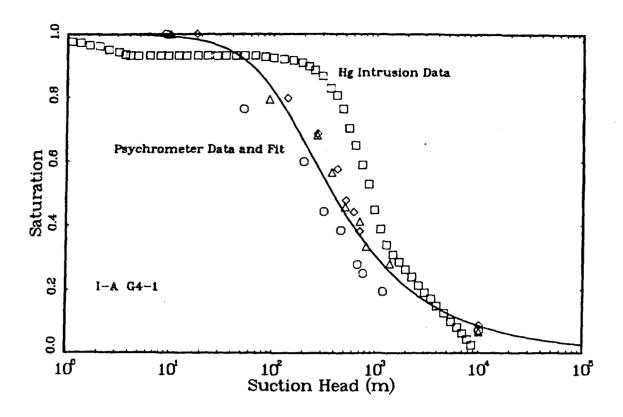
C-1

TABLE C.1 RMS Difference* Between the Mercury Intrusion Data and the Van Genuchten Fit of the Psychrometer Data

	Depth		RMS
Sample	<u>(ft)</u>	<u>Unit</u>	Difference
G4-1	43	I-A	0.174
G4-2	124	I-B	0.139
G4-3	208	I-B	0.369
G4-5	864	II-NL	0.098
G4-24	864	II-NL	0.425
G4-6	1158	II-NL	0.079
G4-7	1256	II-NL	0.079
G4-8	1299	III	0.056
G4-8	1299	III	0.070
G4-10	1405	IV-A-z	0.088
G4-11	1548	IV-A-z	0.097
G4-12	1686	IV-A-z	0.113
G4-13	1728	IV-B-z	0.275
G4-14	1737	IV-B-z	0.103
G4-15	1769	IV-C-Z	0.111
G4-16	1778	IV-C-z	0.128
G4-17	1789	IV-C-z	0.208
G4-18	1899	V	0.190
G4-19	2006	VI	0.134
G4-20	2101	VI	0.194
G4-21	2401	VII	0.236
G4-22	2407	VII	0.111
G4-23	Anaconda	1010A	0.138

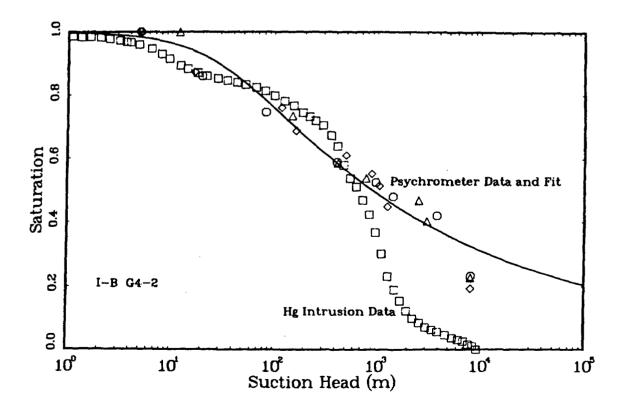
*RMS Difference =
$$\begin{pmatrix} 3000m \\ \sum (Y_p(h) - Y_{Hg}(h))^2 \\ \frac{h=10m}{N} \end{pmatrix}^{\frac{1}{2}}$$

C-2



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Figure C.1





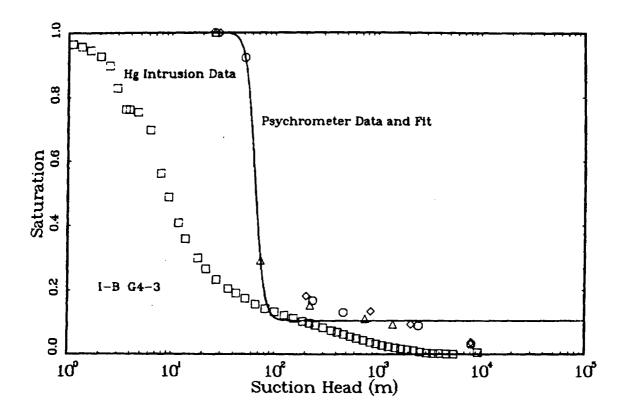
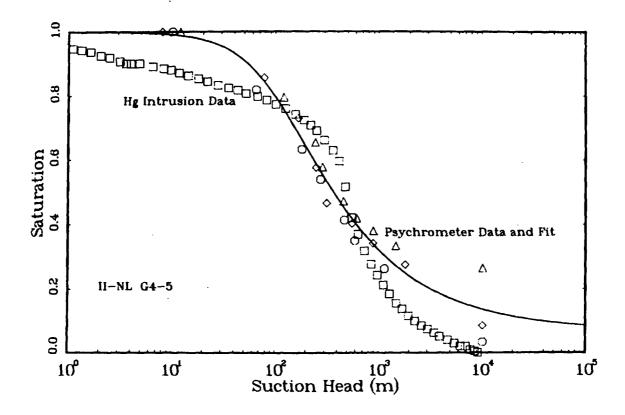
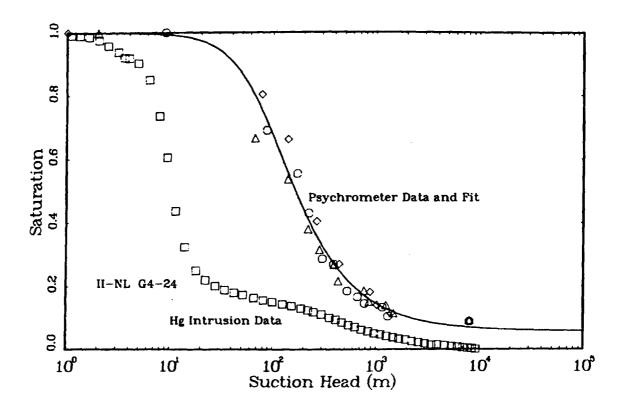


Figure C.3







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Figure C.5

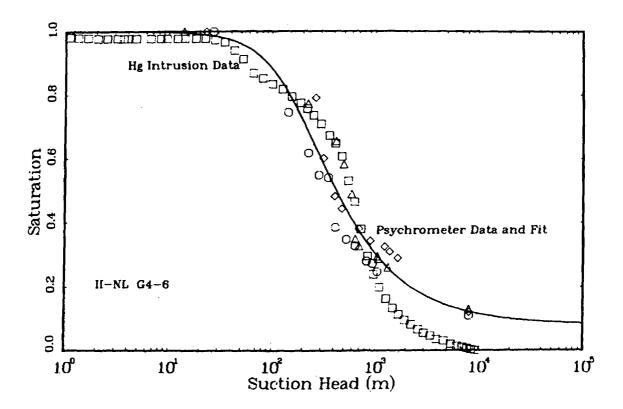


Figure C.6

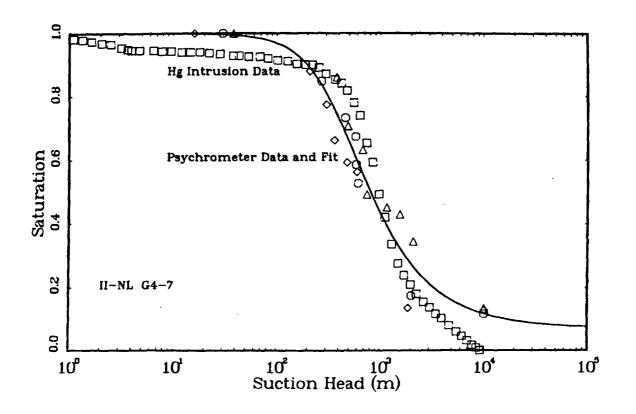


Figure C.7

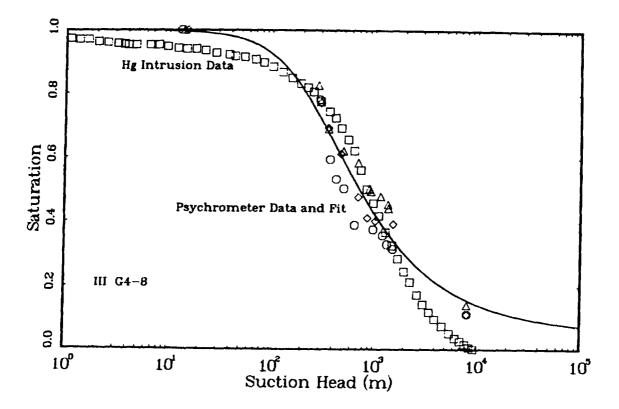
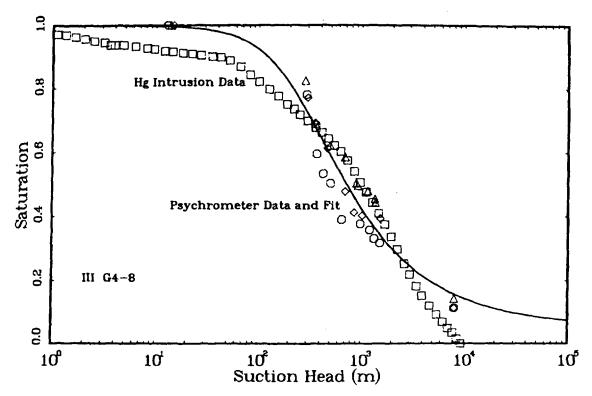


Figure C.8

(Note: This is the result of the first of two Hg Intrusion tests on III G4-8)



(Note: This is the result of the second of two Hg Intrusion tests on III G4-8) Figure C.9

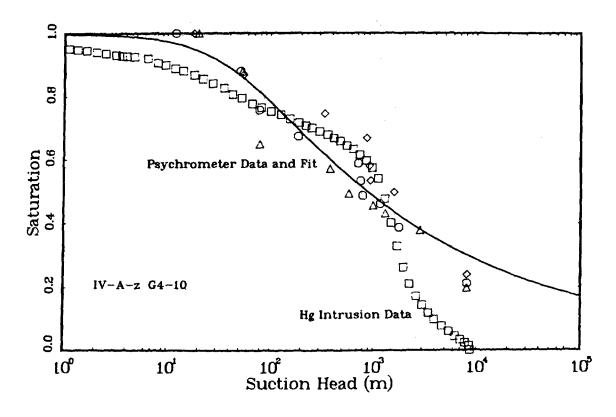


Figure C.10

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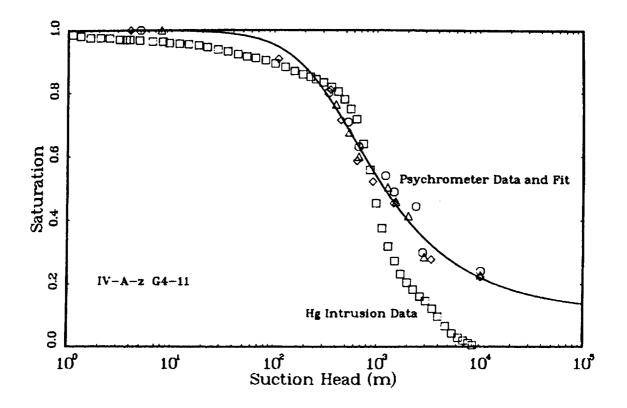


Figure C.11

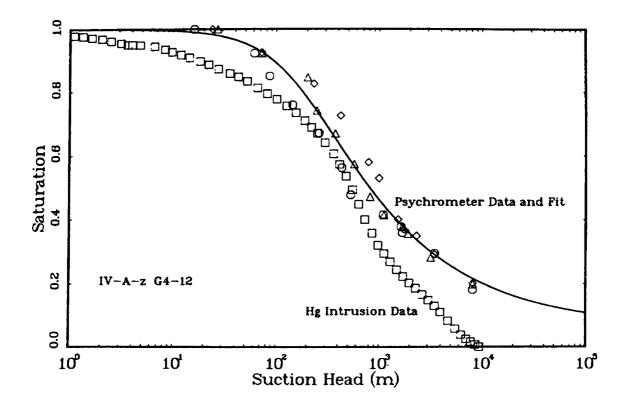


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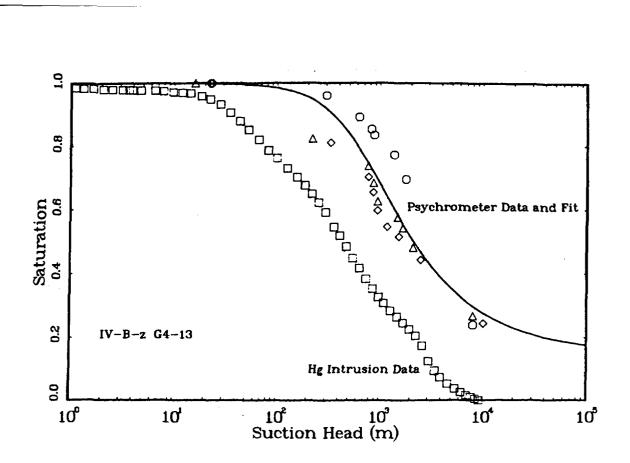


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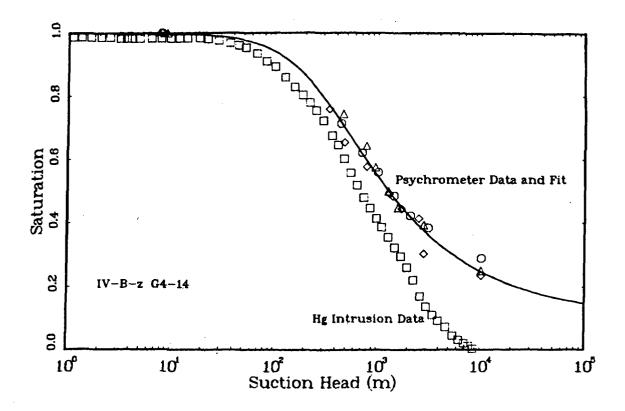
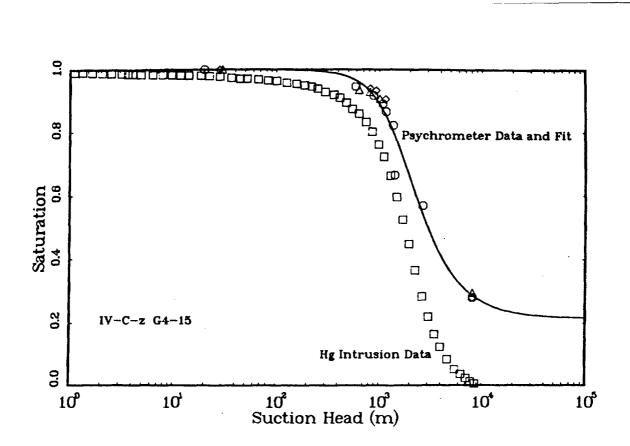


Figure C.14





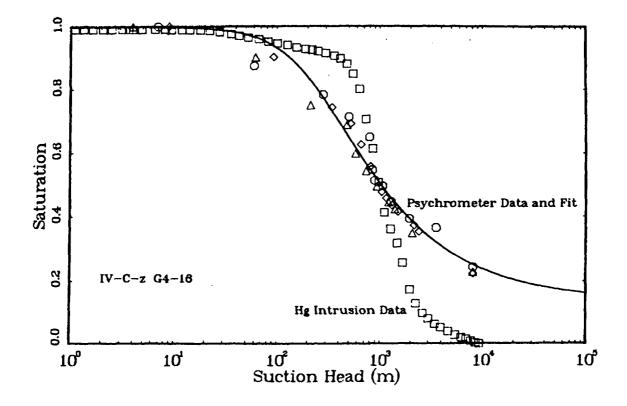


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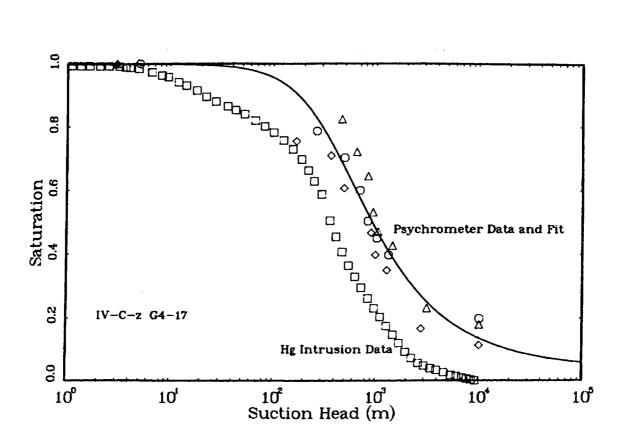
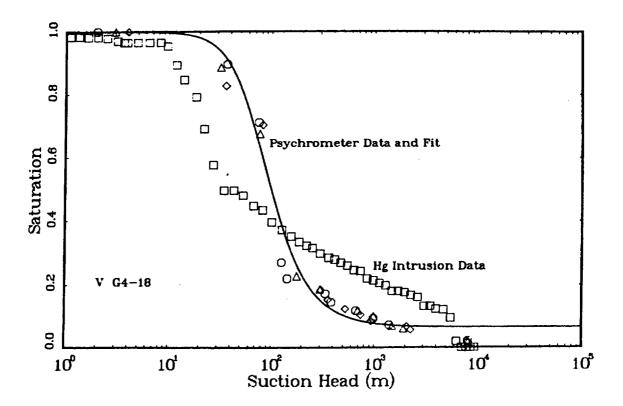
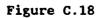
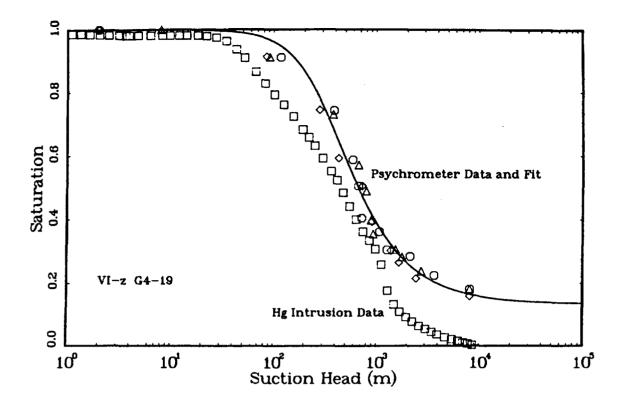
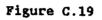


Figure C.17









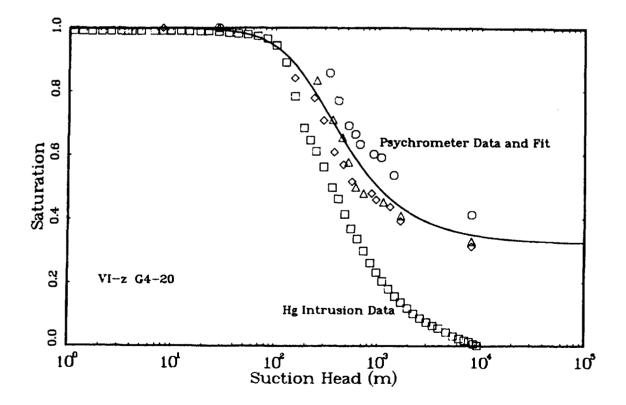


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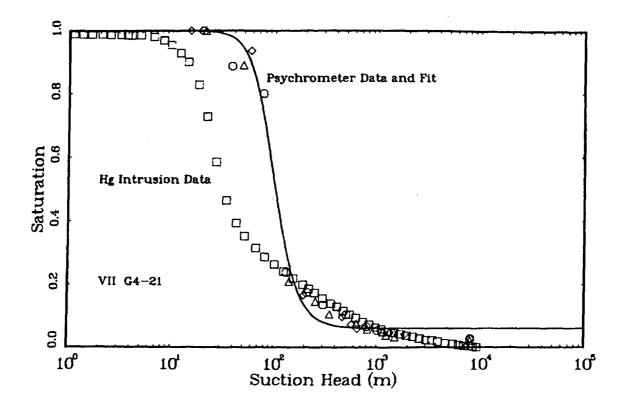
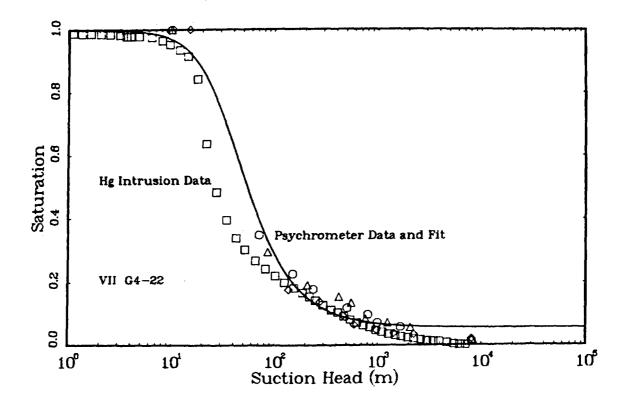


Figure C.21





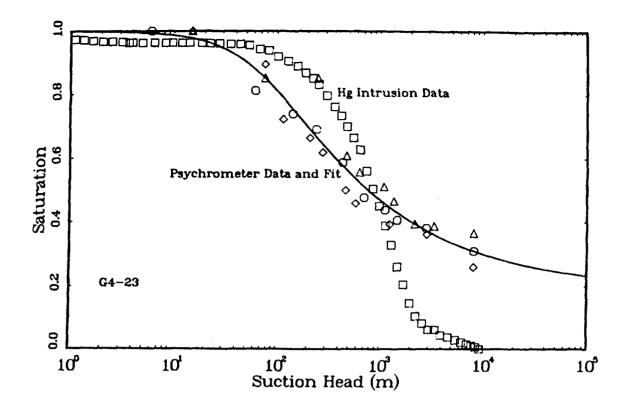


Figure C.23

APPENDIX D

COMPARISON OF VAN GENUCHTEN AND HAVERKAMP SATURATION CURVE FITS AND CONDUCTIVITY CURVES

The Van Genuchten and Haverkamp curve fits of saturation data were compared for 14 different samples. The form of the equations is listed below. In the following figures, the Van Genuchten and Haverkamp curve fits and the saturation data are plotted for each sample investigated along with the relative conductivity curves calculated from the curve fits of the saturation data by the method of Mualem (1976). In general, when the shape of the curve was well defined by the data points (e.g., G4-1), the Haverkamp and Van Genuchten curve fits of the data were nearly identical, and the relative conductivity curve calculated from them was also nearly identical. The root-mean-square (RMS) errors for the two curve fits were also similar. The RMS error values are contained in Table D.1. The major difference between the two curve fits is in the estimation of the conductivity from them. If the Haverkamp saturation curve is used, the relative conductivity equation based on the work of Mualem (1976) must be numerically integrated to obtain the relative conductivity curve with the expression being undefined at the limits of integration. If the Van Genuchten curve is used, the relative conductivity equation can be analytically integrated. Relative conductivity curves using both saturation curve fits and the method of Mualem are included in the plots which follow.

The Van Genuchten curve was chosen as the standard method to fit the saturation data because it gave as good a fit (i.e., as low an RMS value) as other methods, and it yielded an analytical expression for the hydraulic conductivity when calculated using the method of Mualem.

D-1

The defining equations for the two curve fits follow.

Van Genuchten

$$S = (S_{s} - S_{r}) \left[\frac{1}{1 + |\alpha h|^{B}} \right]^{\lambda} + S_{r}$$

Haverkamp

$$S = (S_{s} - S_{r}) \frac{\alpha}{\alpha + (h)^{\beta}} + S_{r}$$

h - Suction head

S - Saturation as a function of pressure

 S_{s} - Maximum saturation (~1)

 s_r - Residual saturation

$$\lambda = 1 = 1/B$$

a and B are curve-fit parameters

Sample	Unit	Van Genuchten	Haverkamp RMS Error (x10 ⁻²)
G4-1	I-A	6.7	6.3
G4-4	II-L	8.9	9.6
G4-7	II-NL	6.4	6.4
G4-9	III	9.4	9.0
G4-11	IV-A-z	3.1	3.1
G4-14	IV-B-z	2.5	2.4
G4-17	V	6.6	6.3
GU3-2	I-A	13.0	12.0
GU3-3	I-A	6.0	6.3
GU3-7	I-B	8.9	13.0
GU3-10	I-B	5.0	5.8
GU3-14	IV-A-v	4.5	5.1
GU3-16	I-B-v	6.1	10.0
GU3-18	v	9.8	9.9

TABLE D.1. RMS Error Values for Selected Haverkamp and Van Genuchten Curve Fits

*RMS: Root-mean-square error between the curves predicted value (Y_p) and the measured value (Y_M) . N is the number of data points.

RMS Error =
$$\left(\frac{\sum_{i=1}^{N} (Y_{Pi} - Y_{Mi})^{2}}{N}\right)^{1/2}$$

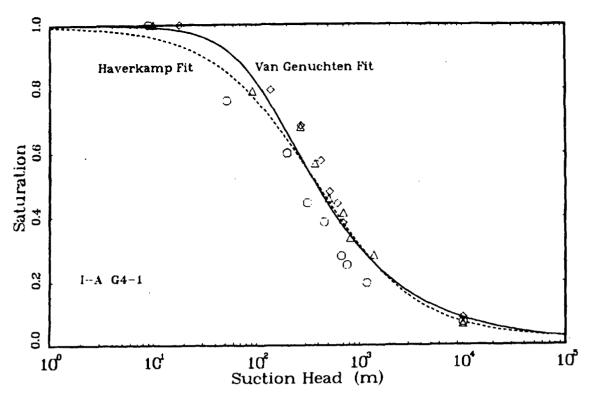


Figure D.1

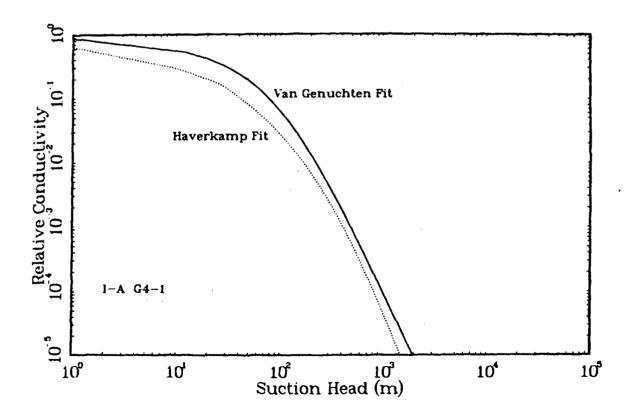


Figure D.2

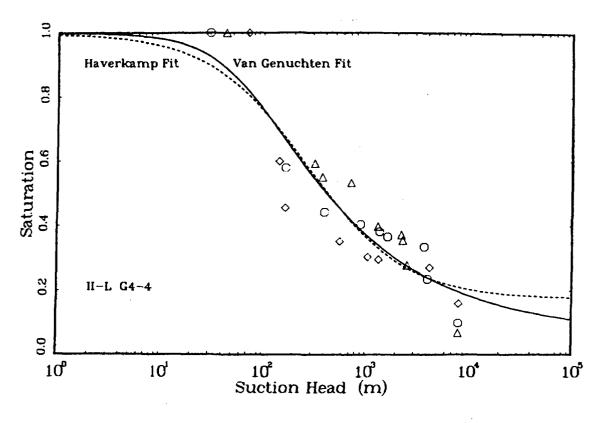


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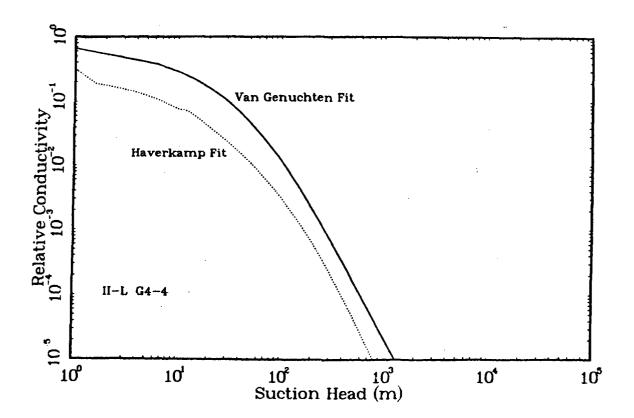
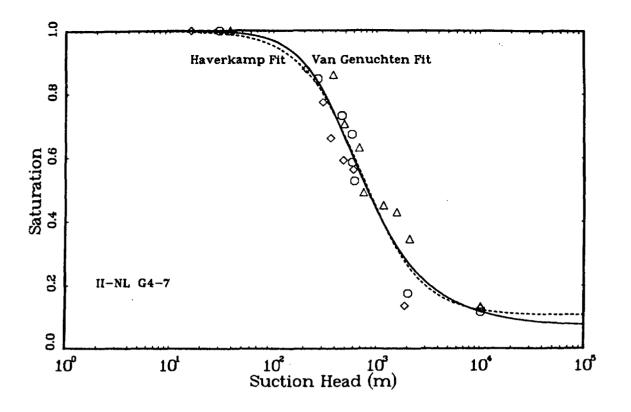


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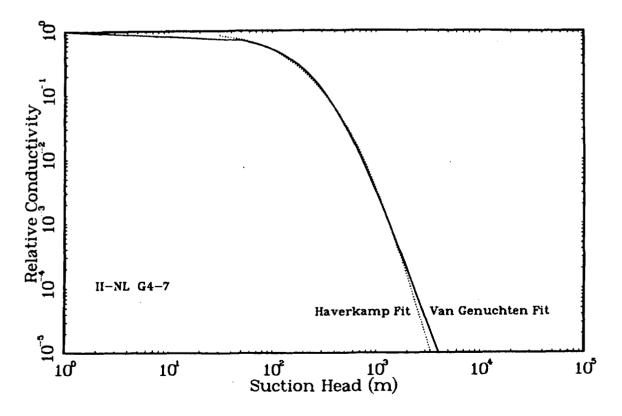


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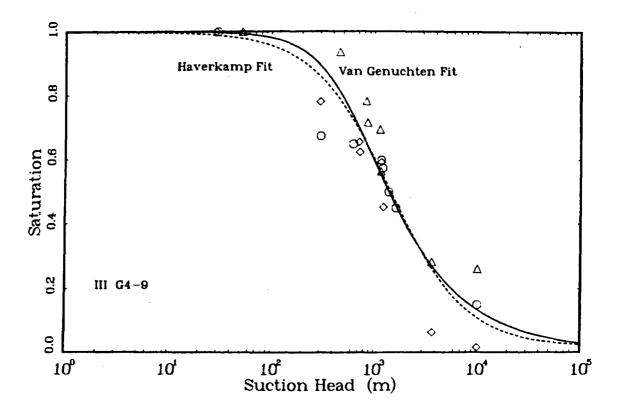
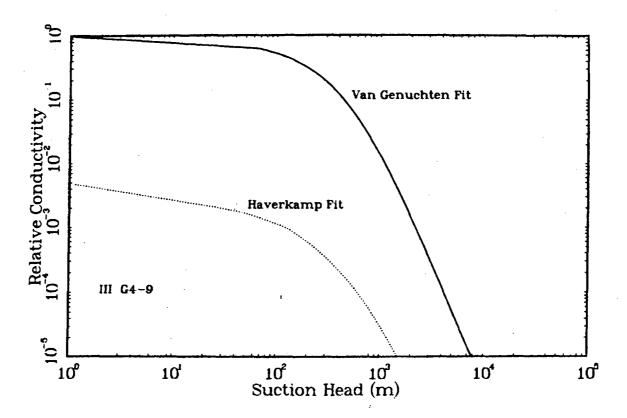


Figure D.7





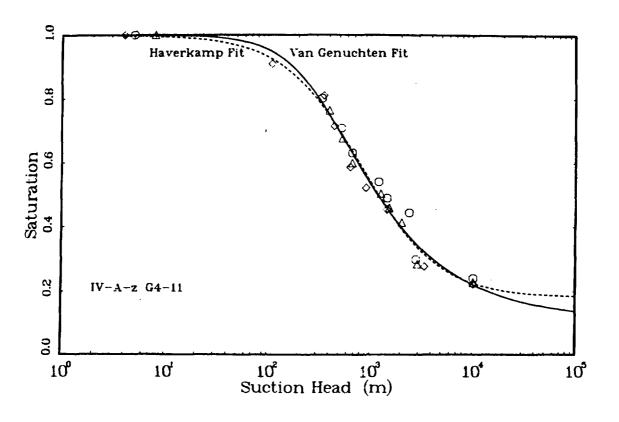


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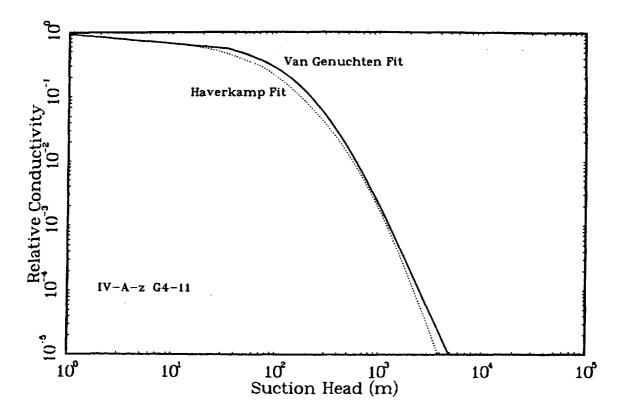


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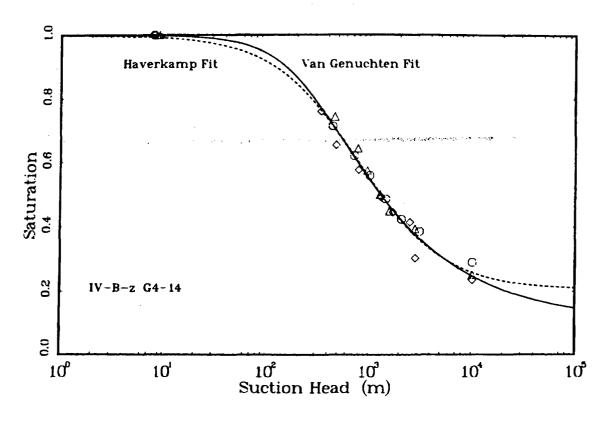
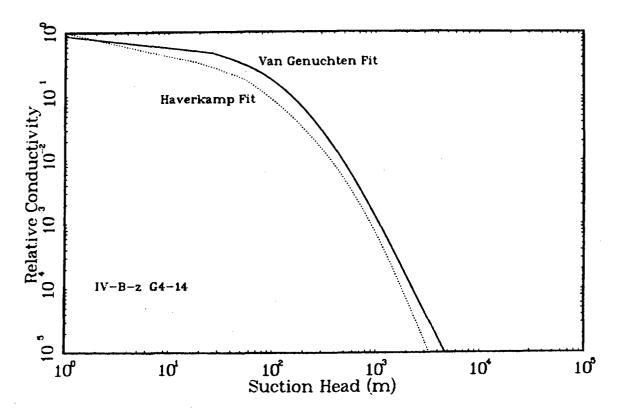


Figure D.11





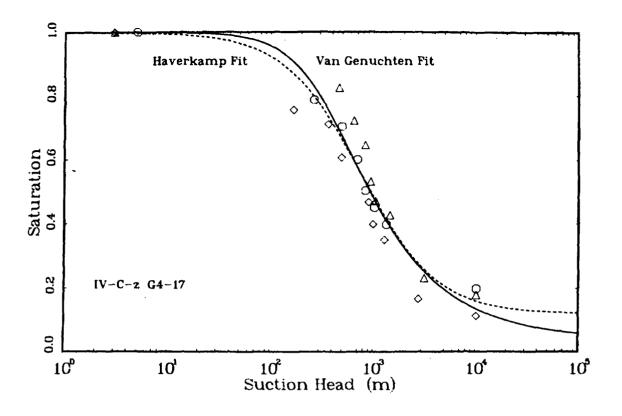


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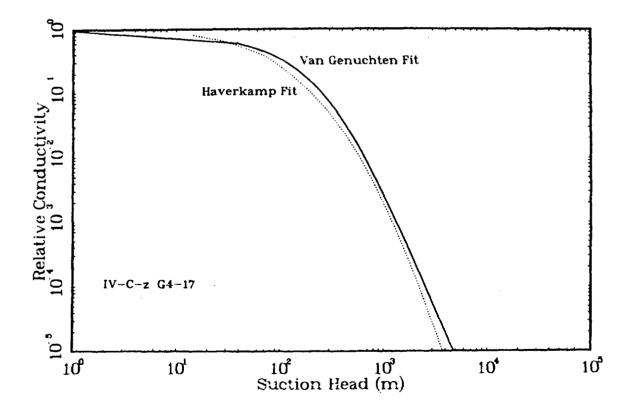


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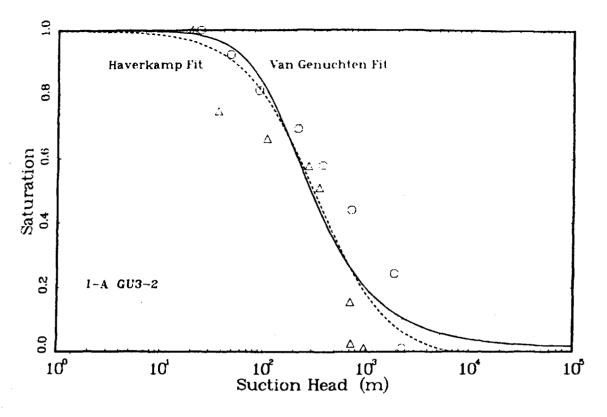


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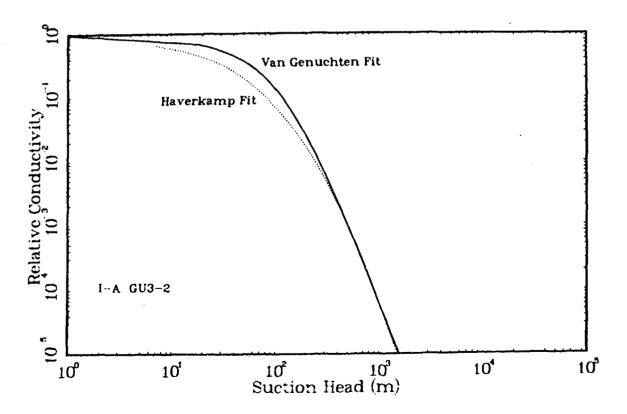


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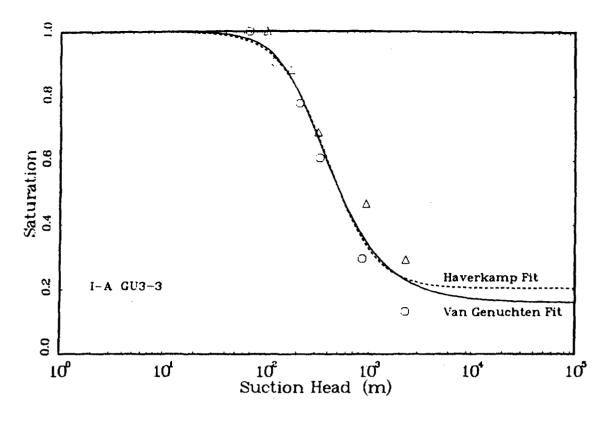


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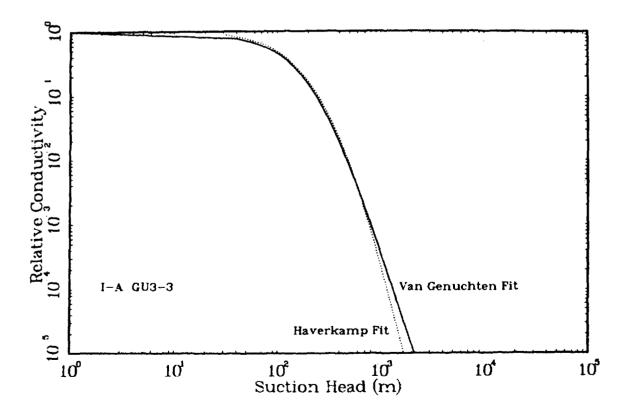


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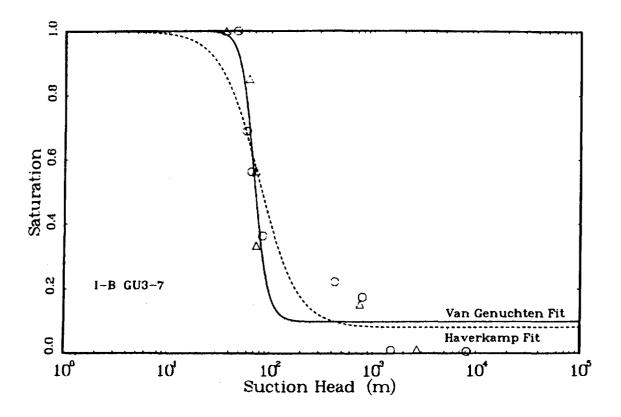


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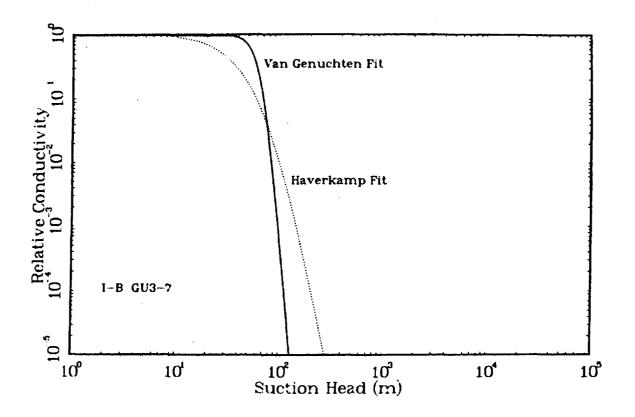


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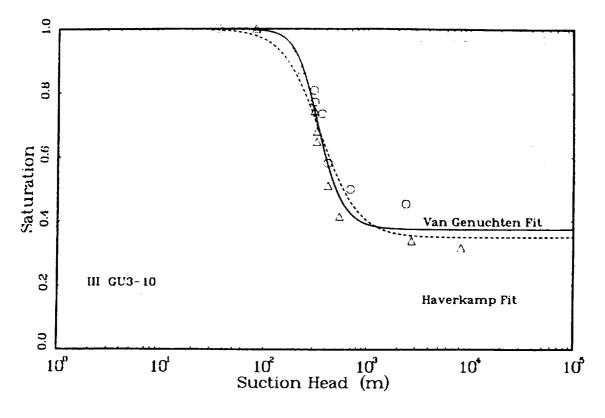


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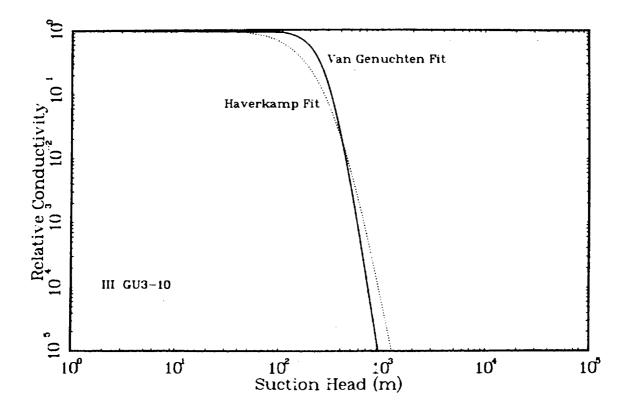


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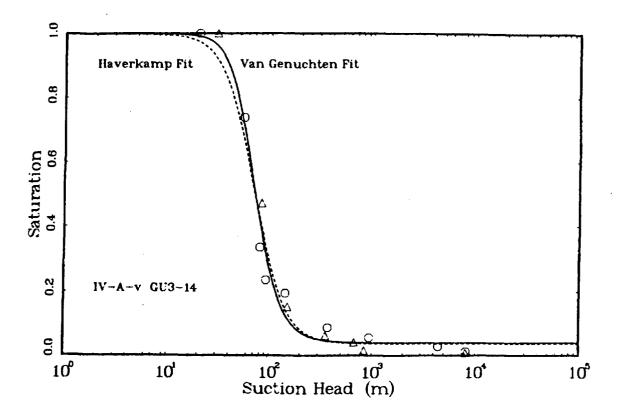


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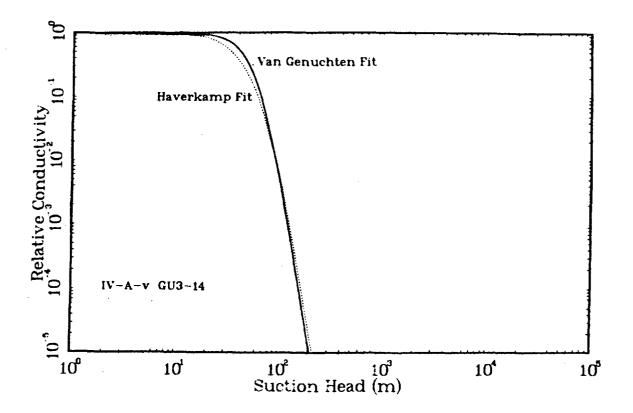


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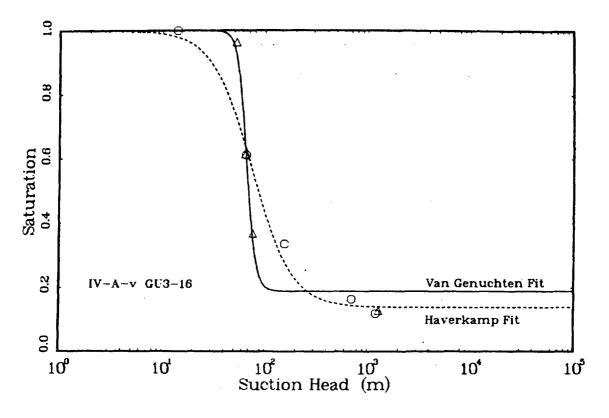


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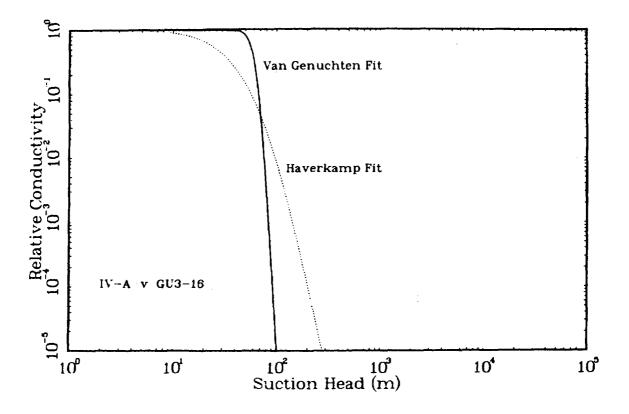


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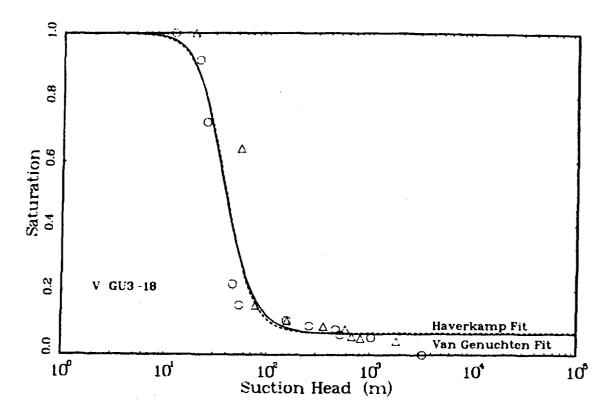


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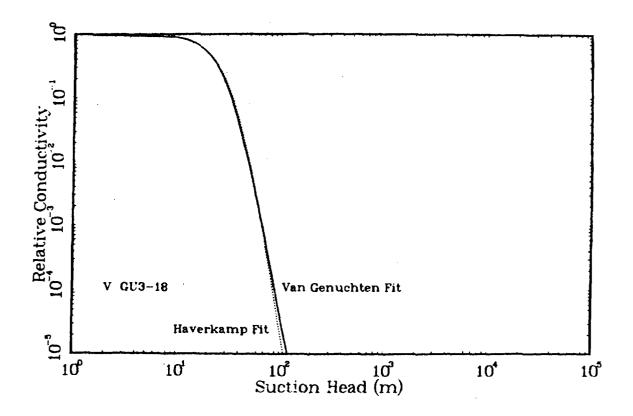


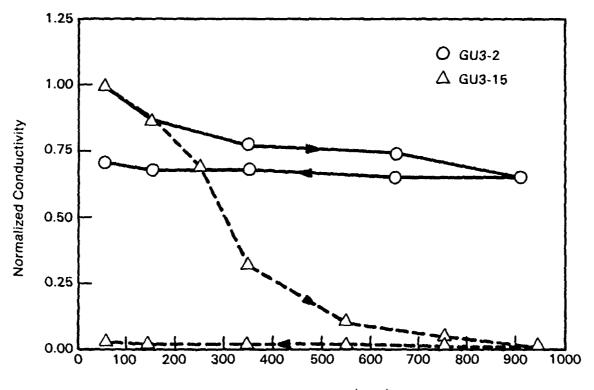
Figure D.28

D-17 - D-18

APPENDIX E

DATA PLOTS FOR CONFINED SATURATED CONDUCTIVITY TESTS

This appendix contains plots of the results of the confined, saturated conductivity tests. Figures E.1 thru E.3 plot the results of the matrix testing [see Tables A.5 to A.7]. Figures E.4 thru E.18 plot the results of the fracture testing [see Tables A.8 - A.12] and the curve fit of the data [see Table 8].

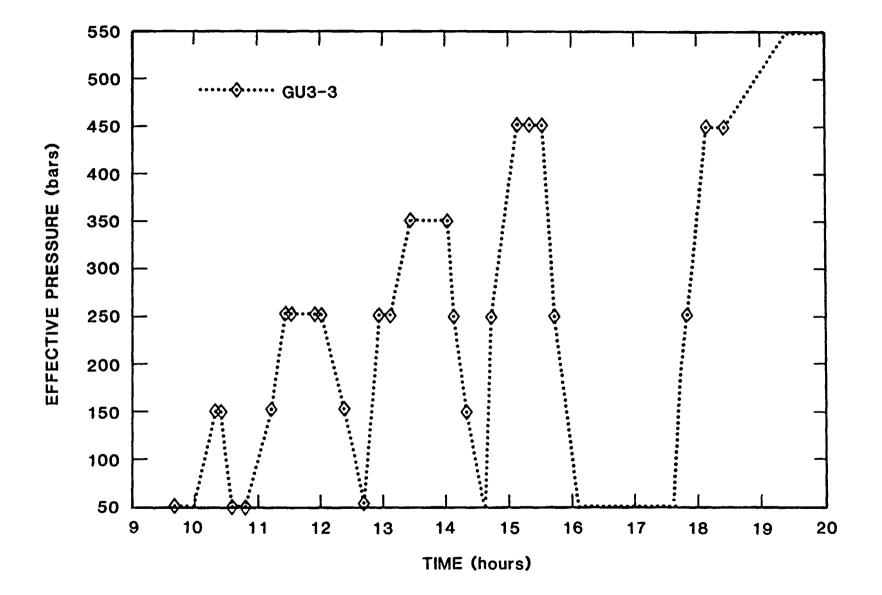


Effective Pressure (bars)

Figure B.1. Normalized Conductivity for Two Tuff Samples from Yucca Mountain as a Function of Increasing Effective Pressure

B-2

4



4

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Figure E.2. Effective Pressure Versus Time for Tuff Sample GU3-3 from Yucca Mountain

B-3

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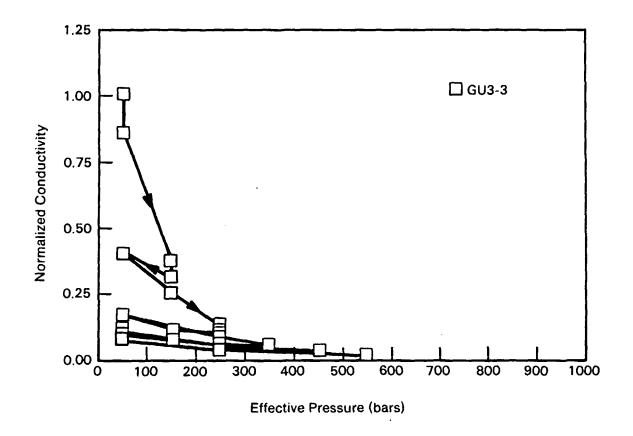


Figure B.3. Normalized Conductivity for GU3-3 from Yucca Mountain as a Function of Increasing Effective Pressure

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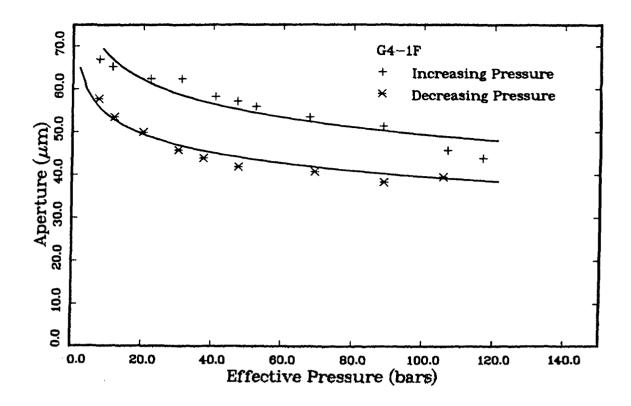


Figure E.4. Computed Aperture Versus Effective Pressure for Sample G4-1F

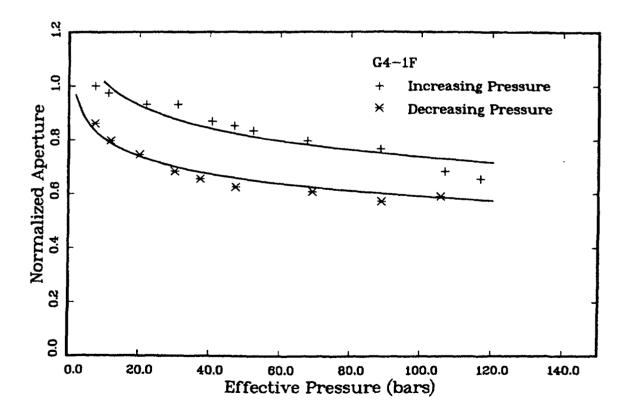


Figure E.5. Normalized Aperture Versus Effective Pressure for Sample G4-1F

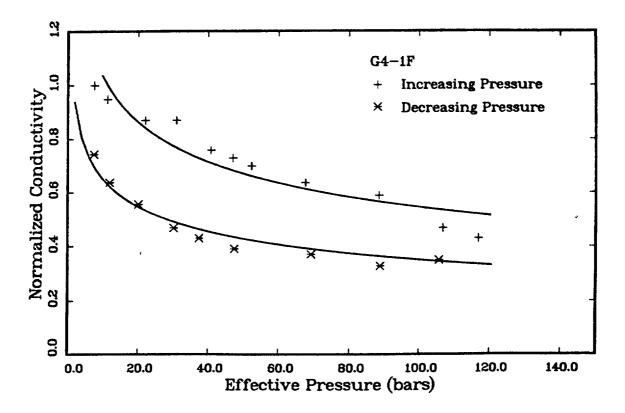


Figure B.6. Normalized Conductivity Versus Effective Pressure for Sample G4-1F

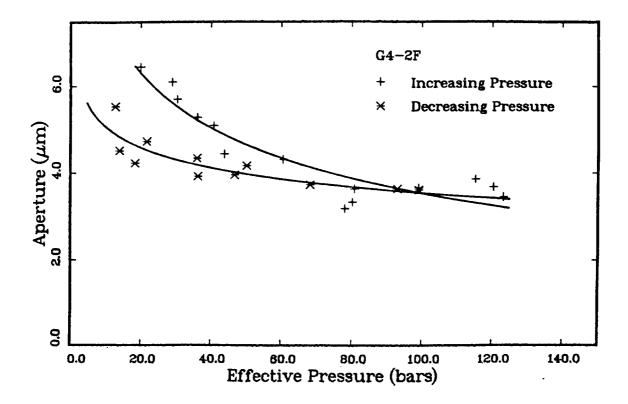


Figure E.7. Computed Aperture Versus Bffective Pressure for Sample G4-2F

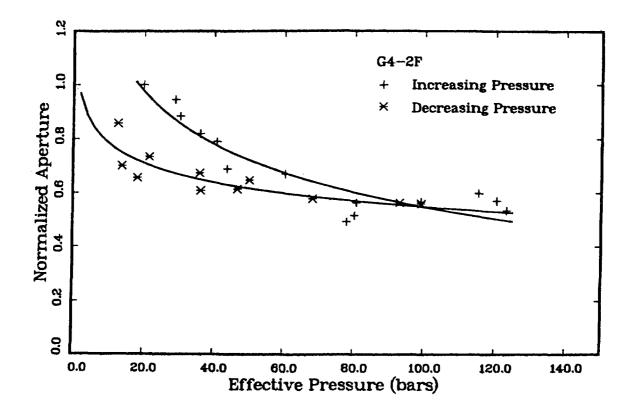


Figure E.8. Normalized Aperture Versus Effective Pressure for Sample G4-2F

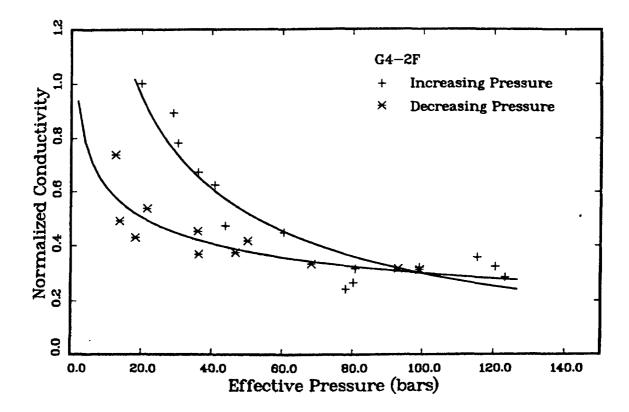


Figure E.9. Normalized Conductivity Versus Effective Pressure for Sample G4-2F

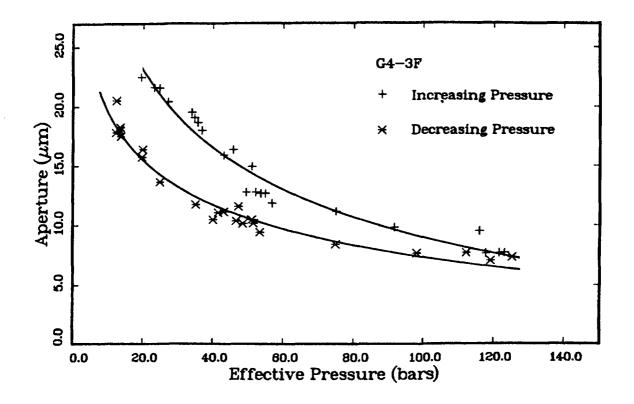


Figure E.10. Computed Aperture Versus Effective Pressure for Sample G4-3F

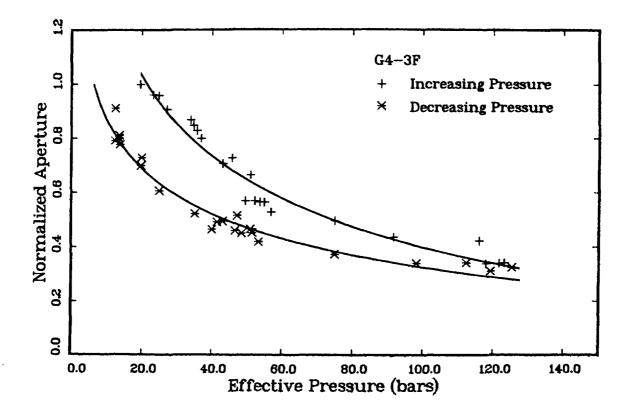


Figure B.11. Normalized Aperture Versus Effective Pressure for Sample G4-3F

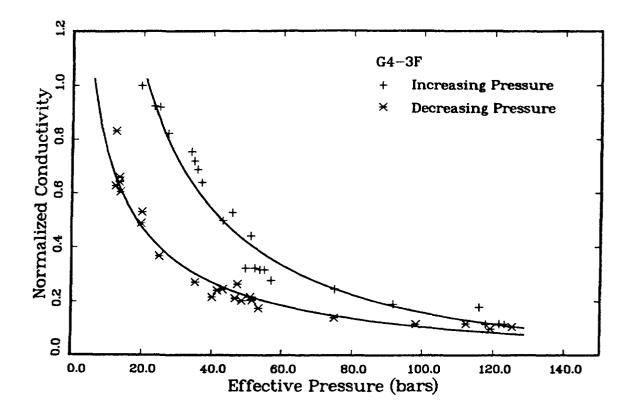


Figure E.12. Normalized Conductivity Versus Effective Pressure for Sample G4-3F

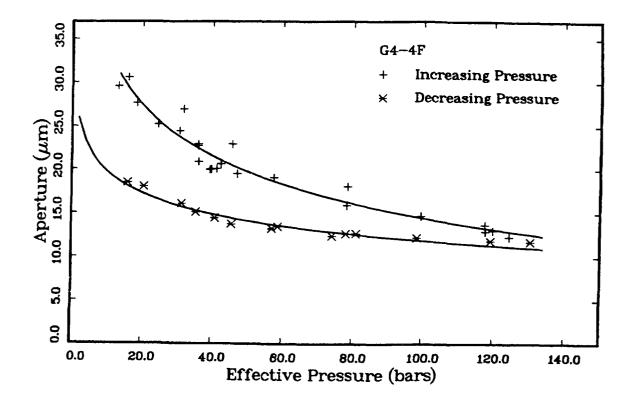


Figure E.13. Computed Aperture Versus Effective Pressure for Sample G4-4F

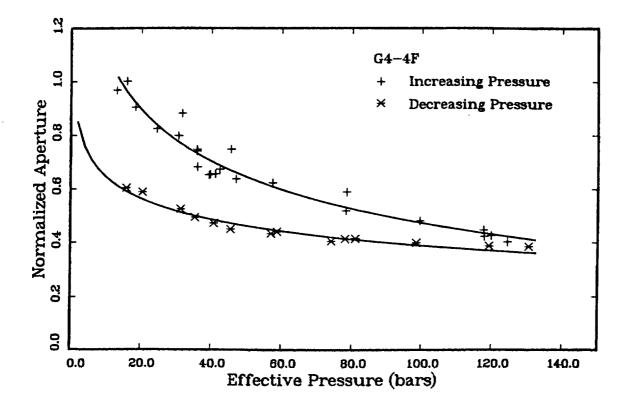


Figure E.14. Normalized Aperture Versus Effective Pressure for Sample G4-4F

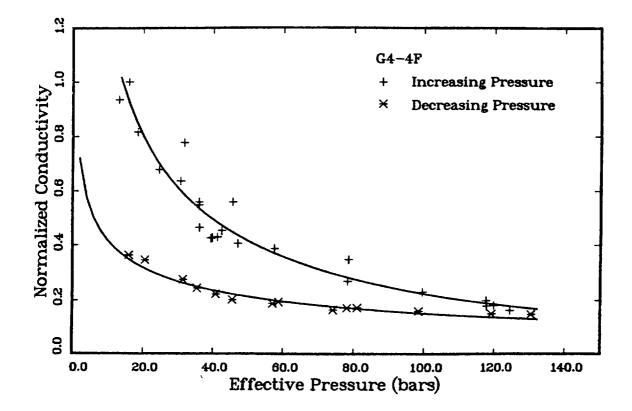


Figure B.15. Normalized Conductivity Versus Effective Pressure for Sample G4-4F

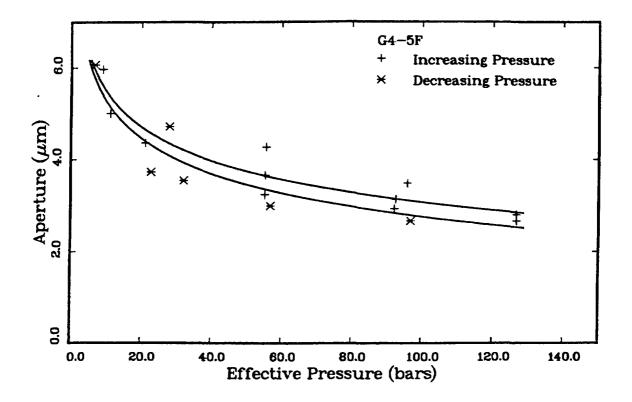
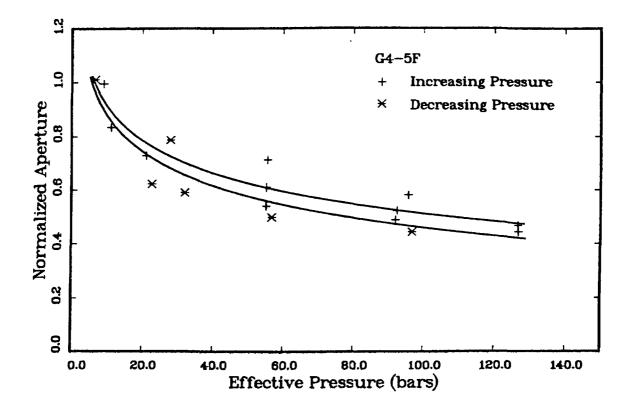
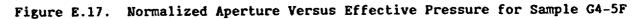


Figure E.16. Computed Aperture Versus Effective Pressure for Sample G4-5F





E-11

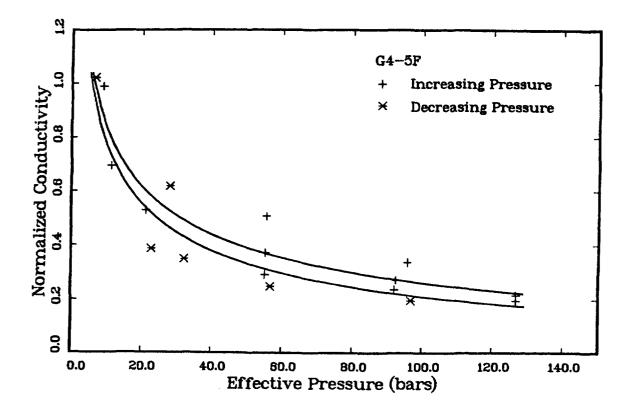


Figure E.18. Normalized Conductivity Versus Effective Pressure for Sample G4-5F

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