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Seth M. Coplan, Section Leader NTS Project Section Repository Projects Branch Division of Waste Management U.S. Nuclear Regulatory Commission Washington, D.C. 20555

OCT 1 0 1984

GEOCHEMISTRY MEETING SUMMARY

As requested in your letter of September 11, 1984, a signed copy of the summary for the July 10-12 Geochemistry Meeting held at Los Alamos, New Mexico is attached. Enclosures 1 through 4 to the summary are also included.

We agree with your suggestion that summaries for future technical meetings should be finalized and signed at the conclusion of the meeting. Agendas prepared for future technical meetings will provide for sufficient time to draft, type, review, and finalize the summaries.

Please contact J. S. Szymanski of my office if you have any questions regarding this matter.

Donald L. Vieth, Director Waste Management Project Office

WMPO:JSS:119

Enclosures: As stated

WM-11

cc w/encl: N. K. Stablein, NRC, Washington, D.C. P. T. Prestholt, NRC, Las Vegas, NV M. A. Glora, SAI, Las Vegas, NV D. T. Oakley, LANL, Los Alamos, NM G. L. DePoorter, LANL, Los Alamos, NM M. B. Blanchard, WMPO, DOE/NV

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NRC-DOE GEOCHEMISTRY MEETING SUMMARY JULY 10-12, 1984

LOS ALAMOS NATIONAL LABORATORY LOS ALAMOS, NEW MEXICO

Attendees

A list of attengees and their organizational affiliations is attached as Enclosure I.

Background/Facts

An agenda, and copies of viewgraphs used by the NRC and DOE speakers, are attached as Enclosures 2, 3, and 4 respectively.

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Observations

The NRC had the following observations:

- 1. The meeting was conducted in a professional manner and provided a useful exchange of information. The NNWSI speakers encouraged open, productive discussion.
- 2. The current and previous workshop agendas were structured along the technical disciplines that are being explored by NNWSI. This leads to presentations that appear to give only limited attention to potential licensing issues, and application of planned investigations, data, or information to resolution of those issues. Restructuring the agenda of future meetings by issue (for example, the NRC site issues presented at this meeting) would be one way of addressing this NNWSI shortcoming.
- 3. USGS has described alternative conceptual models regarding movement of water in unsaturated fractured tuff. The alternatives differ with respect to how much water moves through fractures as opposed to the rock matrix itself. Also, there is uncertainty as to how much recharge occurs at the site (between 0.1 and 8 mm per year). As a result, there is a wide range of possible groundwater residence times that are consistent with what is now known about the Yucca Mountain site. In planning investigations of geochemical processes at the site and in interpreting the results of investigations completed to date, the NNWSI have assumed the hydrologic conditions and models that lead to the longer, less conservative residence times. It is the NRC staff's view that these assumptions are still questionable. Accordingly, the staff considers that the full range of residence times should be considered by the NNWSI in planning and interpreting the geochemistry investigations.
- 4. Values for many geochemical parameters will be needed to support site performance assessment calculations of future behavior, as required by 10 CFR 60. "Accurate" values (i.e. values to be existent) for given parameters under future repository conditions may be difficult to establish by measurement or calculation. The use of bounding values and limiting geochemical conditions, which would support reasonably

conservative performance assessment calculations, could simplify the collection of needed information. The NNWSI Program seems to be giving little consideration to a reasonably conservative <u>vs</u> "accurate" performance analysis, i.e. the collection of limiting or bounding values. Greater consideration of a reasonably conservative approach by NNWSI could greatly facilitate licensing decisions.

5. The solubility or apparent concentration limit of radionuclides in site groundwater/rock systems is likely to be an important barrier to the migration of radioactivity from emplaced waste to the accessible environment. The rate of radionuclide release can be no greater than the product of the apparent concentration limit and the groundwater volumetric flux.

Several aspects of the NNWSI approach to radionuclide solubility may be subject to concern. These include:

- (1) A possible over reliance on solubility values based on geochemical models rather than experimental measurements. Solubility calculations must assume equilibrium which may not be appropriate, and rely on a thermodynamic data base, which may be inadequate for some elements, particularly actinides at elevated temperatures, as well as aluminosilicates.
- (2) Neither the solubility measurements nor the modeling calculations seem to be giving enough emphasis to the effects of altered groundwater composition on radionuclide solubility. The <u>in situ</u> groundwater composition will be altered by contact with waste package and engineered facility components at elevated temperature and radiation fields during migration. It would be desirable to consider the solubility of radionuclides in this altered groundwater. This solubility may be the most representative of the source terms for far-field analysis.
- 6. The mineralogy/petrology program may have much to offer the overall hydrology program. Information was presented regarding mineralization both in the rock matrix and in fractures above the water table. The origin of such mineralization could shed light on which of the alternative conceptual models for unsaturated zone flow is most nearly correct. Also, fracture density and mineralogy offers an opportunity to determine the paleo flow currents through the Topopah Springs. These data/observations may lead to useful information on mineral stabilities on a repository time scale and aid in extrapolating laboratory data to the long term and in determining paleo flow paths. The NRC staff considers that mineralogy studies should be pursued and factored into whatever interpretations are made regarding groundwater movement in the unsaturated zone.
- 7. Reaction path calculations pertaining to rock/water interactions have been done by the NNWSI using the assumption of a closed system. These calculations bear on determining the design bases for the engineered barriers. The NRC staff would encourage that open-system calculations be done as well.

- 8. The potential physiochemical effects of kinetics have not been adequately addressed in NNWSI studies of mineral stabilities. The NRC staff believes that certain features of Yucca Mountain petrology and mineralogy (e.g. variation of zeolite mineralogy with depth) may be more correctly explained if the potential role of kinetics is investigated fully.
- 9. In this meeting there was little discussion of groundwater flow paths and rates of flow in the saturated zone, and, likewise, little mention of geochemical conditions along potential radionuclide migration routes in this zone. Additional information on groundwater flow paths and geochemical conditions in the saturated zone beneath Yucca Mountain is required for a complete and accurate assessment of the potential radionuclide isolation performance of the Yucca Mountain site.
- 10. One phenomenon that may pertain to movement of water in the unsaturated zone is cycles of wetting and drying. Such cycles would affect many unsaturated zone geochemical processes in a variety of ways -- some of which are obvious and others more subtle. Some of the more subtle aspects do not appear to have received full consideration in NNWSI studies. For example, proposals to empirically determine unsaturated zone groundwater travel times by CI dating do not seem to have fully considered the implications of wet-dry cycles with respect to how CI samples might be emplaced and interpreted for dating.
- 11. Ionic species in the groundwater of the unsaturated zone may be concentrated due to evaporation and condensation in the near field. Upon cooling, potential flow of these fluids to the Calico Hills may adversely affect the radionuclide sorptive capacity of the tuffaceous host rocks. It is the view of the NRC staff that such a scenario be addressed.
- 12. The model of secondary mineral (zeolite) stability is important to future decisions about the location of the repository horizon and the choice of a suitable backfill. Two theoretical models have been proposed for zeolite diagenesis at Yucca Mountain. The more recent model suggests that the secondary mineral stability is controlled primarily by the activity of SiO, in solution and is not strongly affected by temperature. At present, this model is not well supported by available data and relies on a number of assumptions. In particular, the phase(s) controlling silica activity . and the mineral stability fields are not known. It is the less conservative of the two models since it implies that irreversible phase (mineralogical) changes will not occur in the vicinity of the waste at temperatures less than 200°C. An earlier model suggested that mineral stability at Yucca Mountain would be a strong function of temperature. This more conservative model should not be abandoned but should be considered an alternate working hypotheses until more field, experimental and theoretical data have been obtained.
- 13. Studies of clay mineral and zeolite dehydration <u>under vacuum</u> do not seem to provide any direct or indirect ("baseline") data that bear on investigations of the behavior of repository host rocks under thermal loading. A much more relevant approach is to investigate clay mineral and zeolite dehydration under prescribed and controlled temperature (T) and

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partial pressure of H_2^0 ($P_{H_2^0}$) conditions that mimic $T-P_{H_2^0}$ conditions that are expected to develop in the near field of an HLW repository in tuff beneath Yucca Mountain.

- 14. From the workshop presentations, it is not clear that a sound approach has been developed by NNWSI for determining that all significant species are included in the data base for geochemical modeling and for picking the controlling solid for modeling solubilities in specific groundwaters.
- 15. The <u>in situ</u> tests of fracture transport planned for the exploratory shaft may not have adequately addressed problems of interpretation of results due to (1) multiple fractures; (2) plume formation and position of sample borings within that plume; (3) extrapolation from 2 m to 2 km; and (4) the fact that only one result will be obtained, not allowing adequate estimation of uncertainty of that result.
- 16. Colloids have been suggested as a possible means for radionuclide transport; hence research on colloids and their properties may be warranted. However, the applicability of the colloid work, as described in the workshop, to repository performance assessment is not clear.
- 17. The sorption work should be guided more by consideration of key nuclides than it appeared to be from the presentations.
- 18. It is still not clear how results of experiments, particularly sorption work with high water/rock ratios (water-dominated system) are going to be applied to the unsaturated zone with very small water/rock ratios (rockdominated system).
- 19. Simple models may often be most useful for sensitivity analysis. Use of a model as complicated as TRACR3D may not be warranted. Analytical models may be more appropriate at this time.
- 20. There is a need for a conceptual geochemical model of mineral stability that integrates the field observations and the laboratory work; until this is done the experimental work, no matter how well conceived, will appear unrelated to repository performance.
- 21. The defense wastes at SRL likely will not be stored in South Carolina but will be sent to a national repository. The NNWSI geochemistry program apparently has not considered inclusion of defense wastes (in addition to spent fuel and reprocessing wastes) in their site analysis scheme. It would be desirable for NNWSI to do so.

The DOE had the following observations:

- 1. Workshop was generally useful; it acquainted us with NRC concerns, especially with respect to characterizing the unsaturated zone. The value of the workshop should be weighed against that of data reviews.
- 2. Speakers were required to skip back and forth, referring to other talks that were presented out of logical sequence due to the NRC-imposed agenda.

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- 3. DOE understands that it is NRC's view that in developing plans there is a lot of room for professional judgment at this time; DOE fully concurs.
- 4. Regulatory framework is still evolving and is ambiguous. Regulatory uncertainties that affect project direction due to delays in issuing of definitive positions exist for: 10 CFR 60 including the unsaturated zone amendment; 40 CFR 191, Draft 4 especially with respect to the definition for accessible environment; rewriting of Reg. Guide 4.17; and technical guidelines on solubility, which are being rewritten (communication at the workshop level would be more helpful as opposed to written interaction on proposed changes).
- 5. In lieu of these regulations, meaningful interaction between NNWSI and NRC is desirable. However, DOE is responsible for the planning and direction of the Project. NRC and NNWSI both expect sensitivity analysis to redirect emphasis on the Project. NRC should recognize that Project redirection takes time and proper planning as well as requiring sound reasons that are related to the overall issue of radiological safety of the public and environment.
- 6. NRC staff need to become familiar with simplifying assumptions that can be made in the TRACR3D radionuclide transport code.

The State of Nevada had the following observations:

- 1. Since the age of the waters in the various parts of the repository system is important, it is critical to consider all age dating techniques. We have heard exhaustive discussions on the use of Cl. Is there a program to look at other age-dating techniques such as He, H, Kr, and I? All of these methods have problems, but they may be complimentary. The stable isotopes such as D, O, and C should also be examined. These efforts should be coordinated with the USGS programs.
- 2. If Rainier Mesa is intended to be an analog to Yucca Mountain, then there needs to be a more comprehensive understanding of the hydrologic and geochemistry conditions. Our experience indicates Rainier Mesa may be a valid analog.
- 3. Since so many parameters are dependent upon knowing unsaturated zone water chemistry and actual in situ saturation, a program should be identified to obtain these data.
- 4. DOE studies have used hydrologic parameters to determine water flow rates in the unsaturated zone. This approach is based on numerous assumptions. There need to be examples of where these techniques have been used successfully.
- 5. What is the Los Alamos program to determine actual input water quality to the bedrock? This seems to be an important parameter to the current experiments and modeling activities.
- 6. We understand that both NRC and DOE have developed heat flow codes and are modeling the proposed Yucca Mountain Repository. We would request complete description and documentation of these activities.

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- 7. Discussions have identified that some very minor minerals may contribute ions to the complexing of radionuclides. We feel that it is important to have a program to identify the amounts and locations of these minerals.
- 8. A statement was made that manganese oxides coat some fracture surfaces. Since flow may occur in these fractures, these coatings should not be ignored. Particular attention should also be paid to sorption experiments on natural fracture surfaces both in the field and in the laboratory.
- 9. Geostatistics should be used to identify alternative flow paths to the accessible environment. Particular attention should be given to potential flow paths that bypass zeolites or other sorptive minerals, as a worst-case scenario.
- 10. If laboratory experiments and modeling continue to use reducing conditions as one possible scenario, then more information is needed on Redox conditions in the repository block in both the saturated and unsaturated zones. The mineralogy/petrology data should be integrated with the geochemical data in establishing a more complete understanding of redox conditions.
- 11. If the repository is to be located in the unsaturated zone, then vapor and aerosol transport should be considered. We heard little discussion on this subject at this workshop.
- 12. We feel that multiple working hypotheses must be considered in developing the geochemical and mineralogical/petrological research plans for NNWSI. Our perception from this workshop is that most of the research evolved from single hypotheses. We do not feel that this is the best approach to asses the characteristics of the Yucca Mountain site.

Agreements

- DOE and NRC agreed to conduct discussions concerning the results of numerical simulations of pore water movement under the influence of thermal fields.
- 2. DOE and NRC agreed to conduct discussions concerning a restructuring of the format and emphasis of future technical meetings. Specifically, a format is needed that addresses more narrowly defined issues that focus on specific phases of repository performance; for example, radionuclide transport in the unsaturated zone.

Open Items

As noted

DOE Requests of NRC

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- 1. Would like to see the QA Review Plan as soon as possible.
- 2. Would like to get the technical positions on Sorption, Solubility, and Mineralogy/Redox conditions as quickly as possible.

3. Would like NRC to review the Los Alamos Geochemistry Program Plan and provide feedback.

NRC Requests of DOE

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1. Correlation of the Los Alamos work plan with "issues" as presented.

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Seth M. Coplar

Division of Waste Management US NRC

Enclosure 1

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JULY 10-12, 1984

LOS ALAMOS NATIONAL LABORATORY

Name	Organization	FTS Phone Number
Aamodt, Paul L.	Los Alamos	843-7960
Beaton, John A.	Los Alamos	843-9805
Bermanis, Henry L.	Weston	202-963-6821
Bilhorn, Susan G.	NRC	
Bish, David L.	Los Alamos	843-4337
Blencoe, James G.	NRC/ORNL	624-7041
Brown, Lee F.	Los Alamos	843-8067
Broxton, David E.	Los Alamos	843-2492
Bryant, Ernest A.	Los Alamos	843-4498
Byers, Frank M.	Los Alamos	843-8753
Campbell, Katherine	Los Alamos	843-2799
Carlos, Barbara H.	Los Alamos	843-6879
Coplan, Seth M.	NRC	427-4177
Crowe, Bruce M.	Los Alamos	843-4299
DePoorter, Gerald L.	Los Alamos	843-1033
Duffy, Clarence J.	Los Alamos	843-5154
Elzeftawy, Atef	NRC	427-4675
Erdal, Bruce R.	Los Alamos	843-5338
Essington, Edward H.	Los Alamos	843-3057
Fiore, Joy H.	SAI/NV	575-1203
Fuentes, H.	Los Alamos	843-3004
Furman, Marvin J.	DOE/HQ	444-7062
Glora, Michael A.	SAI/NY	575-1463
Grisham, Genevieve F.	Los Alamos	843-5321
Hess, John W.	State of Nevada	702-798-8882
Jacobson, Roger L.	State of Nevada	702-673-7373
Johnson, Carl A.	State of Nevada	702-885-3744
Johnson, Mark	Los Alamos	843-3308

Name	Organization	FTS Phone Number
Kelmers, Donald A.	NRC/ORNL	624-6870
Kerrisk, Jerry F.	Los Alamos	843-3348
Knauss, Kevin	LLNL	532-1372
Kovach, Linda A.	NRC	427-4693
Levy, Schon S.	Los Alamos	843-9504
Meijer, Arend	Los Alamos	843-0831
Merson, Thomas J.	Los Alamos	843-5726
Michels, Ronald D.	Los Alamos	843-5816
Myers, C. Wes	Los Alamos	843-6722
Nitsche, Heino	LBL	451-4483
Norris, A. Edward	Los Alamos	843-5442
Oakley, Donald T.	Los Alamos	843-1310
Ogard, Allen E.	Los Alamos	843-6344
Ortiz, Terri	SNL	846-0267
Panno, Samuel Y.	Weston	202-963-6848
Planner, Harry N.	Los Alamos	843-1582
Polzer, Wilfred L.	Los Alamos	843-3073
Prestholt, Paul T.	NRC	598-6125
Rowley, John C.	Los Alamos	843-1378
Rundberg, Robert S.	Los Alamos	843-4559
Siegel, Malcolm D.	NRC/SNL	846-5448
Stablein, Newton K.	NRC	427-4611
Starmer, R. John	NRC	427-4541
Szymanski, Jerry S.	WMPO/NY	575-1503
Thomas, Kimberly W.	Los Alamos	843-4379
Travis, Bryan J.	Los Alamos	843-1254
Vaniman, David T.	Los Alamos	843-1165
Wolery, Thomas J.	LLNL	532-5789

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Enclosure 2

JULY 10-12, 1984

LOS ALAMOS NATIONAL LABORATORY

Tuesday - July 10, 1984

- 8:00 a.m. Registration
- 8:15 a.m. Welcome/Introductions D. T. Oakley
- 8:30 a.m. Technical Overview for Workshop DOE Representative
- 9:30 a.m. Break
- 9:45 a.m. LOS ALAMOS PRESENTATION Overview and Geochemistry Program Plan - G. L. DePoorter
- 10:15 a.m. NRC PRESENTATION Geochemical Issues Specific to NNWSI
- 10:45 a.m. Discussion Los Alamos and NRC
- 11:45 a.m. Lunch
- 12:45 p.m. SESSION I: UNSATURATED ZONE PROCESSES Overview - R. S. Rundberg (Los Alamos)
- 1:05 p.m. NRC Comments
- 1:15 p.m. Discussion Los Alamos and NRC Discussion Presentations Unsaturated Zone Transport (including Vapor) - B. J. Travis Unsaturated Zone Water Chemistry - A. E. Ogard Suggested Discussion Topic Relation between Saturated and Unsaturated Zone Geochemical Processes
- 2:45 p.m. Break
- 3:00 p.m. SESSION II: ROCK-WATER INTERACTIONS Presentations Unsaturated Zone Travel Time 36C1 Experiment for ES - A. E. Norris Rock-water Equilibrium - C. J. Duffy Reaction Path Calculations - J. F. Kerrisk
- 3:30 p.m. NRC Comments
- 3:40 p.m. Discussion Los Alamos and NRC Suggested Discussion Topic Equilibrium vs. Kinetic Effects
- 4:45 p.m. End of Session

Wednesday, July 11, 1984

- 8:00 a.m. SESSION III: 3-D MINERALOGY AND SORPTION STRATIGRAPHY Mineralogy/Petrology Update - D. T. Vaniman (Los Alamos)
- 8:20 a.m. NRC Comments
- 8:30 a.m. Discussion Los Alamos and NRC Discussion Presentation - Los Alamos Evaluation of Statistical Methods - K. Campbell Suggested Discussion Topics 3-D Mineralogy and Sorptive Mineral Stratigraphy Fracture Mineralogy Determining Paleo Flow Paths from Present Mineralogy
- 10:00 a.m. Break
- 10:15 a.m. SESSION IV: MINERAL STABILITY LOS ALAMOS PRESENTATION Mineral Stability

10:20 - 10:35 Hydrothermal Stability - C. J. Duffy 10:35 - 10:50 Thermal Stability - D. Bish

- 10:50 a.m. NRC Comments
- 11:00 a.m. Discussion Los Alamos and NRC
- 12:00 Noon Lunch
- 1:00 p.m. SESSION V: KEY RADIONUCLIDES AND SOLUBILITY Key Radionuclides - J. F. Kerrisk (Los Alamos)
- 1:10 p.m. NRC Comments
- 1:15 p.m. NRC PRESENTATION Solubility
- 1:35 p.m. Los Alamos Comments
- 1:45 p.m. EQ3/6 Geochemical Model T. Wolery (LLNL)
- 2:00 p.m. NRC Comments
- 2:10 p.m. LOS ALAMOS PRESENTATION Solubility and Speciation

2:10 - 2:40Solubility Calculations - J. F. Kerrisk2:40 - 3:00Solubility Measurements - Los Alamos3:00 - 3:20Solubility Measurements - LBL

Wednesday, July 11, 1984

- 3:20 p.m. Break
- 3:40 p.m. Discussion Los Alamos and NRC Suggested Discussion Topics Key Radionuclides Solubility and Speciation of Key Radionuclides
- 4:00 p.m. SESSION VI: REDOX CONDITIONS AND REDOX BUFFERING Los Alamos Presentation - A. E. Ogard
- 4:20 p.m. NRC Comments
- 4:30 p.m. Discussion Los Alamos and NRC Suggested Discussion Topics Redox Conditions in the Unsaturated Zone Redox Buffering by Fe-Ti Minerals
- 4:45 p.m. End of Session

Thursday, July 12, 1984

- 8:00 a.m. SESSION VII: TRANSPORT MODELING AND MATRIX DIFFUSION Transport Modeling - B. J. Travis (Los Alamos)
- 8:20 a.m. NRC Comments
- 8:30 a.m. Discussion Los Alamos and NRC Discussion Presentations - Los Alamos Fracture Flow Experiments (Laboratory and ES) Diffusion Experiments (Laboratory and ES) Colloid Size Determination Discussion Topics: Matrix Diffusion Vapor Transport Aerosol Transport of Radionuclides
- 9:45 a.m. Break
- 10:00 a.m. SESSION VIII: SORPTION Update on Sorption - K. W. Thomas (Los Alamos)
- 10:35 a.m. NRC Comments
- 10:45 a.m. Discussion Los Alamos and NRC Suggested Discussion Topics Sorption Measurements on Crushed vs Intact Samples Batch vs Flow Experiments Effects of Water Composition and Redox Conditions Sorption Kinetics Key Radionuclides
- 12:00 Noon Lunch
- 1:00 p.m. Preparation of Minutes and Closeout
- 3:30 p.m. Feedback and Discussion
- 4:45 p.m. End of Workshop

Enclosure 3

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NRC - Geochem 155 25 Handouts

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Technical Position

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It is the position of the NRC staff that based on our current level of knowledge of the NNWSI site, assessment of the Technical Criteria of 10CFR60 in terms of the performance elements of NUREG-0960 requires that, at a minimum, the following issues concerning geochemistry be addressed.

3.0 Geochemistry

- 3.1. What are the geochemical conditions preceding waste emplacement in the saturated/unsaturated zone?
 - 3.1.1 What is the mineralogy/petrology/chemistry of the backfill, disturbed zone/farfield host rock prior to waste emplacement?
 - 3.1.2 What is the mineralogy/petrology/chemistry along potential release pathways of the disturbed zone/farfield host rock prior to waste emplacement?
 - 3.1.3 What are the geochemical conditions of the groundwater in the disturbed zone/farfield in the saturated/unsaturated zone.
- 3.2 What are the geochemical conditions/processes following waste emplacement, in the saturated/unsaturated zone?
 - 3.2.1 How are the mineral stabilities of the Yucca Mountain tuffs affected by anticipated changes in the temperature and pressure?
 - 3.2.1.1 What is the rock/groundwater ratio in the backfill, disturbed zone/farfield in the saturated/unsaturated zone through time?
 - 3.2.1.2 Can chemical equilibrium be assumed in rock/water interactions under various flow regimes in the saturated/unsaturated zone?
 - 3.2.2 What is the anticipated spatial distribution of alteration products due to increased temperatures and altered fluid flow paths.

PRELIMINARY DRAFT FOR DISCUSSION AT NNWSI GEOCHEMISTRY WORKSHOP 84/07/06

- 3.2.3 What are the anticipated geochemical conditions of the groundwater due to increased temperature within the unsaturated/saturated zone?
- 3.3 What are the anticipated geochemical processes/conditions affecting release and transport of the radionuclides in the saturated/unsaturated zone?
 - 3.3.1 How does solubility/concentration of radionuclides under differing flow regimes affect transport?
 - 3.3.1.1 How does precipitation/co-precipitation affect radionuclide migration/retardation from the vicinity of the outermost packing material/rock/backfill interfaces to the accessible environment through time?
 - 3.3.1.2 How does speciation affect radionuclide solubility?
 - 3.3.1.3 How do colloids/particulates affect the solubility of radionuclides?
 - 3.3.2 What is the importance of reaction and sorption kinetics in radionuclide migration/retardation in the disturbed zone/farfield environment under various flow regimes?
 - 3.3.3 How do redox conditions of the groundwater and redox buffering potential of Fe-Ti oxides affect radionuclide speciation/retardation?
 - 3.3.3.1 What are the effects of gamma and alpha radiolysis products on backfill and disturbed zone/farfield host ` rock relevant to radionculide transport?
 - 3.3.4 How does backfill and disturbed zone/farfield mineralogy (under anticipated flow regimes) influence radionuclide migration through time?
 - 3.3.5 How do colloids/particulates/organics under differing flow regimes affect radionuclide migration/retardation in the disturbed zone/far field through time?

PRELIMINARY DRAFT FOR DISCUSSION AT NNWSI GEOCHEMISTRY WORKSHOP 84/07/06

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- 3.3.6 How does matrix diffusion affect radionuclide migration/retardation in the disturbed zone/farfield through time?
- 3.3.7 How does vapor transport affect radionuclide migration/retardation in the disturbed zone/farfield through time?

PRELIMINARY DRAFT FOR DISCUSSION AT NNWSI GEOCHEMISTRY WORKSHOP 84/07/06

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L. KOVACH MRC WORKSHOP 7-10-84

DEFINITIONS

SITE ISSUE A QUESTION ABOUT A SPECIFIC SITE THAT MUST BE ADDRESSED AND RESOLVED TO COMPLETE THE LICENSING ASSESSMENT OF SITE SUITABILITY AND/OR DESIGN SUITABILITY IN TERMS OF 10 CFR PART 60. (NEED NOT BE CONTROVERSIAL.)

PERFORMANCE ISSUEA BROAD QUESTION CONCERNING THE OPERATION OR
LONG-TERM PERFORMANCE OF THE 'ARIOUS COMPONENTS
OF THE REPOSITORY SYSTEM.



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2-D TEMPERATURE PROFILE

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GEOCHEMICAL ISSUES FOR NIWSI

- 1. WHAT ARE THE GEOCHEMICAL CONDITIONS <u>PRECEDING</u> WASTE EMPLACEMENT IN THE SATURATED/UNSATURATED ZONE?
- 2. WHAT ARE THE GEOCHEMICAL CONDITIONS/PROCESSES FOLLOWING WASTE EMPLACEMENT IN THE SATURATED/UNSATURATED ZONE?
- 3. WHAT ARE THE GEOCHEMICAL PROCESSES/CONDITIONS AFFECTING <u>RELEASE</u> AND <u>TRANSPORT</u> OF THE RADIONUCLIDES IN THE SATURATED/UNSATURATED ZONE?

GEOCHEMICAL CONDITIONS PRECEDING WASTE EMPLACEMENT

- 1. MINERALOGY/PETROLOGY/CHEMISTRY OF THE BACKFILL AND FAR FIELD HOST ROCK
- 2. MINERALOGY/PETROLOGY/CHEMISTRY ALONG POTENTIAL RELEASE PATHWAYS
- 3. GEOCHEMICAL CONDITIONS OF THE GROUNDWATER

GEOCHEMICAL CONDITIONS FOLLOWING WASTE EMPLACEMENT

1. MINERAL STABILITIES OF BACKFILL AND HOST ROCK

2. SPATIAL DISTRIBUTION OF ALTERATION PRODUCTS

3. GEOCHEMICAL CONDITIONS OF THE GROUNDWATER

SUMMARY OF GEOCHEMICAL ISSUES

INITIAL CONDITIONS

GEOCHEMICAL CONDITIONS PRECEDING WASTE EMPLACEMENT

MINERALOGY MINERALOGY ALONG RELEASE PATHWAYS GROUNDWATER GEOCHEMISTRY

CHANGING CONDITIONS

GEOCHEMICAL CONDITIONS FOLLOWING WASTE EMPLACEMENT MINERAL STABILITIES SPATIAL DISTRIBUTION OF ALTERATION GROUNDWATER CONDITIONS

RELEASE AND TRANSPORT

GEOCHEMICAL PROCESSES AFFECTING RELEASE AND TRANSPORT OF RADIONUCLIDES

SOLUBILITY/CONCENTRATION REACTION/SORPTION KINETICS REDOX MINERALOGY COLLOIDS/PARTICULATES/ORGANICS MATRIX DIFFUSION VAPOR TRANSPORT

GEOCHEMICAL PROCESSES AFFECTING RELEASE AND TRANSPORT OF RADIONUCLIDES

1. SOLUBILITY/CONCENTRATION OF RADIONUCLIDES

1

2. REACTION/SORPTION KINETICS

3. REDOX CONDITIONS

4. MINERALOGY

5. COLLOID/PARTICULATES/ORGANICS

6. MATRIX DIFFUSION

7. VAPOR TRANSPORT

BRIEFING ON THE NRC GENERIC TECHNICAL POSITION

ON

"SOLUBILITY"

R J STARMER

NRC/NNWSI GEOCHEMISTRY TECHNICAL MEETING LOS ALAMOS, NEW MEXICO JULY 11, 1984

BACKGROUND

SOLUBILITY IS A RELATIVELY SIMPLE CHEMICAL CONCEPT WHICH MAY BE DEFINED PRECISELY IN TERMS OF EQUILIBRIUM CHEMICAL THERMODYNAMICS. THE MEASURE OF SOLUBILITY IS MORE DIFFICULT, EVEN IN SIMPLE SYSTEMS. IN THE REPOSITORY ENVIRONMENT, SOLUBILITY CAN BE DEFINED ONLY IN AN OPERATIONAL SENSE. ANY ATTEMPT TO DEFINE OR USE THE TERM SOLUBILITY IS OPEN TO LENGTHY DISCUSSION. THIS REQUIRES A STATEMENT OF THE NRC STAFF POSITION ON THE SUBJECT. SOLUBILITY. A property of a substance by virtue of which it forms mixtures with other substances which are chemically and physically homogeneous throughout. The degree of solubility (often spoken of as "solubility") is the concentration of a solute in a saturated solution at any given temperature. The degree of solubility of most substances increases with rise in temperature, but there are cases (notably the organic salts of calcium) where a substance is more soluble in cold than in hot solvents.

The solubility of a given element is the sum of the stoichiometric concentrations of all dissolved species containing the element.

> solubility. The ability or tendency of one substance to blend uniformly with another, e.g., solid in liquid, liquid in liquid, gas in liquid, gas in gas. Solids vary from 0 to 100% in their degree of solubility in liquids, depending on the chemical nature of the substances; to the extent that they are soluble, they lose their crystalline form and become molecularly or ionically dispersed in the solvent to form a true solution.

solubility (Chem.). The mass of a dissolved substance which will saturate 100 g of a solvent under stated conditions.

solubility (PHYS CHEM) The ability of a substance to form a volution with another substance.

SOLUBILITY, REAL. In the case of salts which form a solution in a given solvent in which the salt exists partly as unionized molecules and partly as ions, the maximum concentration of unionized salt that may exist under equilibrium conditions is called the real solubility, and the maximum total concentration of ionized and unionized salt is called the apparent solubility. •

GARRELS AND CHRIST (1965) P. 55: "THE SOLUBILITY OF A GIVEN ELEMENT IS THE SUM OF THE STOICHIOMETRIC CONCENTRATIONS OF ALL DISSOLVED SPECIES CONTAINING THE ELEMENT."

SOLUBILITY OF SOMETHING

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USE OF "SOLUBILITY" ESTIMATES

1) SCREENING FOR KEY RADIONUCLIDES

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- 2) A WASTE PACKAGE PERFORMANCE MEASURE
- 1) A FAR-FIELD PERFORMANCE ASSESSMENT SOURCE TERM
STATEMENT OF POSITION

Each site which takes credit for solubility should design a minimum set of experiments to produce an information base adequate for licensing.

The set of experiments can be minimized if the following points are considered.

- 1. The range and proportions of solid compositions in the system that could react with ingressing or egressing waters;
- 2. The range of repository water compositions needed for determining radionuclide solution concentrations; and

3. The sources and effects of uncertainties.

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- 1) USE A MATRIX OF COMBINATIONS OF WATER AND MATERIALS WHICH IS RATIONAL AND SUPPORTABLE
- 2) DETERMINE ALTERED GROUNDWATER COMPOSITIONS
- 3) DETERMINE "SOLUBILITY" (MAXIMUM STEADY-STATE CONCENTRATION) LIMITS FOR THESE WATER COMPOSITIONS AT PRESSURES AND TEMPERATURES OF THE MATRIX SET UP IN (1) ABOVE <u>IN THE ABSENCE OF THE REACTING SOLIDS</u>

ASSUMPTIONS

- Knowledge of radionuclide solution concentration will be essential to near field and far field radionuclide transport modeling.
- It is feasible to determine maximum concentration of radionuclides in a representative site-sepcific hydrothermal system via laboratory experiments.
- A conservative assessment of radionuclide solution concentration should be based on approach to equilibrium from oversaturation.
- o Solubility measurements in the absence of other geochemical mechanisms give a conservative estimate of radionuclide concentration.

IF ALL DISSOLVED COMPONENTS ARE PRESENT IN THE PRECONDITIONED WATER <u>THEN</u>

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THE FIRST (AND ONLY POSSIBLE) PRECIPITATE OR SUITE OF PRECIPITATES WILL BE THE CONTROLLING SOLID(S)

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UNCERTAINTY

Water/Rock ratio

·Colloids/Particulates

Supersaturation

Kinetics -

Other suggested sources of uncertainty could include:

- o Use of defonized, distilled or synthetic ground waters,
- Proportions of mineral solids versus radionuclide solids, and
- The extrapolation of short-term results to predicted long-term results.

MODELING

Alone, models are not suitable for estimating source term concentrations. Source problems with calculations are:

equilibrium is assumed

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Unrealistic in view of changing T, mineralogy, pH, and redox conditions. Also, these calculations generate solubilities in most stable state, hence nonconservative.

Incomplete mass transfer

data base limitations

4. experimental verification

Enclosure 4

NNWSI/NRC GEOCHEMISTRY WORKSHOP

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July 10-12, 1984

LOS ALAMOS NATIONAL LABORATORY

VIEW GRAPHS AND BACKUP DATA

OVERVIEW AND GEOCHEMISTRY PROGRAM PLAN

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LOS ALAMOS NATIONAL LABORATORY CONTRIBUTION

WBS 2.3.1 GEOCHEMISTRY WBS 2.3.2 MINERALOGY/PETROLOGY



Los Alamos



WBS 2.3.1 GEOCHEMISTRY

- GROUNDWATER CHEMISTRY
- SOLUBILITY DETERMINATION
- SORPTION AND PRECIPITATION
- DYNAMIC TRANSPORT PROCESSES
- RETARDATION SENSITIVITY ANALYSIS
- APPLIED DIFFUSION
- HYDROTHERMAL GEOCHEMISTRY
- NATURAL ISOTOPE CHEMISTRY



- HOST ROCK STRATIGRAPHY AND MINERALOGY
- THREE-DIMENSIONAL DATA BASE FOR MINERALOGY/ PETROLOGY OF TRANSPORT PATHWAYS
- FRACTURE MINERALOGY
- MINERALOGIC CONTROLS ON RADIONUCLIDE MOBILITY (Eh AND pH)
- HEAT LOAD EFFECTS ON PHYSICAL AND SORPTIVE
 PROPERTIES OF MINERALS
- ALTERATION HISTORY OF YUCCA MOUNTAIN



2.3.1 GEOCHEMISTRY SUMMARY NETWORK

Los Alamos





CHEMICAL ANALYSIS DIVITRIFIED TOPOPAH

SHAFT TESTING

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GEOSTATISTICAL EVALUATION

OF TOPOPAH AND CALICO ES SAMPLES

EVALUATION OF X-RAY

DIFFRACTION DATA ES

SUMMARY REPORT ON

DIST & VARI OF TOPOPAN

PETROGRAPHIC STRATIGRAPHY

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BACKUP VIEWGRAPHS AND DATA

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6/84



2.3.1.4 SOLUBILITY **DETERMINATION NETWORK**

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ALTERED ROCK



2.3.1.6 DYNAMIC TRANSPORT PROCESS NETWORK

Los Alamos

1987 1986 1983 1984 1985 **CRUSHED TUFF** COLUMN **CONTINUE CRUSHED TUFF** EXPERIMENTS **COLUMN EXPERIMENTS** SUMMARY REPORT ON ASSESSMENT REPORT ON KINETICS OF SORPTION **KINETICS OF SORPTION** CONDUCT UNSATURATED DEVELOP FLOW COLUMN TECHNIQUES FOR UNSATURATED EXPERIMENTS FLOW COLUMN **COMPLETE UNSATURATED FLOW** EXPERIMENTS COLUMN EXPERIMENTS PRELIMINARY REPORT SUMMARY REPORT ON UNSATURATED **ON MATRIC POTENTIAL** FLOW COLUMN EXPERIMENTS **OF YM TUFFS** CONTINUE FRACTURE FLOW EXPERIMENTS COMPLETE **CONDUCT FRACTURE FRACTURE FLOW** FLOW EXPERIMENTS EXPERIMENTS PRELIMINARY REPORT ON **REPORT ON TRANSPORT DIFFUSION COEFFICIENTS** AND RETARDATION IN OF RN IN YM TUFF FRACTURED TUFF **REPORT ON TRANS** OF RN BY FRACTURE FLOW SATURATED CONDS COMPLETE SATURATED SOLID SATURATED SOLID TUFF **TUFF FLOW** FLOW EXPERIMENTS EXPERIMENTS PRELIMINARY PAPER ON COLLOID PROPERTIES **REPORT ON TRANS OF** PARTICULATES THROUGH RELATIVE TO TRANS & RETARDATION UNFRACTURED TUFF 6/84









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2.3.2.3 TRANSPORT PATHWAY NETWORK

Los Alamos







UNSATURATED ZONE PROCESSES

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FLOW CHARACTERISTICS
AFFECTED BY MATRIC POTENTIAL
PERMEABILITY DEPENDS ON SATURATION
DIFFUSIVITY DEPENDS ON SATURATION
DISPERSIVITY MAY DEPEND ON SATURATION
FRACTURE VS. POROUS FLOW DEPENDING ON INFILTRATION RATE





PRESENT KNOWLEDGE OF TRANSPORT IN THE YUCCA MOUNTAIN UNSATURATED ZONE MATRIC POTENTIAL / POROSIMETRY DATA SATURATION VS. DEPTH DIFFUSIVITY AT 100 % SATURATION MATRIX PERMEABILITY AT 100 % SATURATION TRACR3D CALCULATIONS OF FRACTURE FLOW SCENARIOS
















FIGURE 11.—Relative mole fractions of cations in the tuffaceous rocks, fracture waters, and experimental solutions.



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Fig. 3. Relative Na-K-Ca concentration in Yucca Mountain water.

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BACKUP VIEWGRAPHS AND DATA



REPRESENTATIVE GROUNDWATERS

USW H-3 BELOW REPOSITORY UE25p#1 PALEOZOIC AQUIFER J-13 SURROUNDING AQUIFER

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SNOW OR RAIN JUVENILE RECHARGE

Los Alamos





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TABLE I

ELEMENTAL CONCENTRATIONS OF GROUNDWATERS FROM THE VICINITY OF YUCCA MOUNTAIN

Concentration (mg/l) Field Well No. Ca Mg Na K Li Fe Mn A1 Si ____PH___ USW VH-1ª 7.5 10 1.5 80 1.9 0.090 23 USW H-6 7.4 5.5 0.22 74 2.1 0.10 0.12 0.04 0.12 20.0 9.4 USW H-3 0.8 0.01 1.5 0.22 0.13 0.01 0.51 16.9 124 USW H-5 7.1 1.1 0.03 2.3 0.04 0.01 0.17 17.4 54 N.D. USW G-4 7.1 9.2 0.15 2.5 0.08 56 0.04 0.02 0.02 19.6 USW H-1ª 7.5 6.2 <0.1 51 1.6 0.04 19 USW H-4 7.4 10.8 0.19 2.6 0.16 0.04 25.9 84 0.03 0.005 UE-256#1^b 7.7 19.7 0.68 3.3 0.28 0.04 0.004 0.03 31.5 56 UE-255#1^C 7.2 2.5 0.30 0.04 28.7 18.4 0.68 46 0.69 0.36 UE-255#1d 7.3 17.9 0.66 3.0 0.17 80.0 0.07 0.06 28.8 37 J-13 6.9 11.5 1.76 45 5.3 0.06 0.04 0.001 0.03 30.0 UE-29a#2 7.0 11.1 0.34 1.2 0.10 0.05 0.04 25.8 51 0.03 J-12^ª 7.1 2.1 38 5.1 25 14 UE-25p#1 87.8 31.9 13.4 0.32 0.1 30 6.7 171 <0.1 <0.1

^aData from L. Benson et al.¹

^bIntegral water sample.

^CBullfrog zone, 4th day.

^dBullfrog zone, 28th day.

Note: Ionic or molecular species are not listed; concentration is based on the element.

	ANION	CONCENT	ra'	TIONS	S AND	OTHE	ER 1	EASURI	EMENTS	
OF	GROUNI	WATERS	IN	THE	VICIN	ITY	OF	YUCCA	MOUNTAI	N

TABLE II

	Concentration (mg/l)									
Well No.	_ <u>F</u> _	<u>C1</u>	so ₄ ²⁻	HCO3	NO2	NO3		Ehf	Detergent	
USW VH-1ª	2.7	10	45	165						
USW H-6	4.1	7.7	27.5		N.D. ^e	5.3	5.6	395		
USW H-3	5.4	8.3	31.2		<0.10	0.2	<0.1	-143	<0.02	
USW H-5	1.3	5.7	14.6		N.D.	8.6	6.3	353	<0.005	
USW G-4	2.4	5.5	15.7		N.D.	5.5	6.4	402		
USW H-1 ^a	1.0	5.8	19	122						
USW H-4	4.5	6.2	23.9		N.D.	4.7	5.8	216	>2	
UE-256#1 ^b	1.2	7.1	20.6		N.D.	0.6	1.8	220		
UE-25b#1 ^C	1.5	9.8	21.0		0.5	2.2	<0.1	-18	2.7	
UE-255#1 ^d	1.2	6.6	20.3		N.D.	4.5	1.8	160	0.02	
J-13	2.1	6.4	18.1		N.D.	10.1	5.7		N.D.	
UE-29a#2	0.56	8.3	22.7		N.D.	18.7	5.7	305		
J-12 ⁴	2.1	7.3	22	119						
UE-25p#1	3.5	37	129		N.D.	<0.1		360	<0.2	

^aData from L. Benson et al.¹

b Integral water sample.

^CBullfrog zone, 4th day.

^dBullfrog zone, 28th day.

e_{N.D.} = not detected.

fmV vs H₂ electrode.

TABLE III

ELEMENTAL CONCENTRATIONS OF GROUNDWATER

FROM THE VICINITY OF YUCCA MOUNTAIN

	Field	Concentration (mmols/1)								
Well No.	_ <u>pH</u>	Ca	Mg	Na	K	Li	Fe	Mn		Si
USW VH-1ª	7.5	0.25	0.06	3.48	0.049	0.013				0.83
USW H-6	7.4	0.14	0.009	3.22	0.053	0.014	0.002	0.0007	0.004	0.71
USW H-3	9.4	0.02	0.0002	5.39	0.038	0.032	0.002	0.0002	0.019	0.60
USW H-5	7.1	0.03	0.001	2.35	0.059	0.006	0.0002	N.D.	0.006	0.62
USW G-4	7.1	0.23	0.007	2.43	0.064	0.012	0.0007	0.0004	0.0007	0.70
USW H-1 ^a	7.5	0.15	<0.004	2.22	0.041	0.006				0.71
USW H-4	7.4	0.27	0.008	3.65	0.066	0.023	0.0005	N.D.	0.001	0.92
UE-256#1 ^b	7.7	0.49	0.028	2.43	0.084	0.040	0.0007	N.D.	0.001	1.12
UE-255#1 ^C	.7.2	0.46	0.028	2.00	0.064	0.043	0.012	0.006	0.001	1.02
UE-256#1 ^d	7.3	0.45	0.027	1.61	0.077	0.024	0.001	0.001	0.002	1.03
J-13	6.9	0.29	0.072	1.96	0.136	0.009	8000.0	0.00002	0.0010	1.07
UE-29a#2	7.0	0.28	0.014	2.22	0.031	0.014	0.0009	0.0005	0.001	0.92
J-12 ⁸	7.1	0.35	0.086	1.65	0.130					0.90
UE-25p#1	6.7	2.19	1.312	7.43	0.343	0.046	<0.002	0.002	0.003	0.62

^aData from L. Benson et al.¹

^bIntegral water sample.

^CBullfrog zone, 4th day.

^dBullfrog zone, 28th day.

TABLE IV ANION CONCENTRATIONS OF GROUNDWATER IN THE VICINITY OF YUCCA MOUNTAIN

	Concentrations (mmols/L)									
Well No.	F	_ <u>C1</u> _	so ₄ ²⁻	Alkalinity	NO ₂	NO3				
USW VH-1 ^a	0.14	0.28	0.47							
USW H-6	0.22	0.22	0.29	2.75	N.D. ^e	0.09				
USW H-3	0.28	0.23	0.32	4.72	<0.002	0.003				
USW H-5	0.07	0.16	0.15	2.00	N.D.	0.14				
USW G-4	0.13	0.16	0.16	2.34	.N.D.	0.09				
USW H-1 [®]	0.05	0.16	0.20	•						
USW H-4	0.24	0.17	0.25	2.82	N.D.	80.0				
UE-256#1 ^b	0.06	0.20	0.21	2.20	N.D.	0.01				
UE-25b#1 ^C	80.0	0.28	0.22	2.41	0.011	0.03				
UE-256#1 ^d	0.06	0.19	0.21	2.13	N.D.	0.07				
J-13	0.11	0.18	0.19	2.34	N.D.	0.16				
UE-29a#2	0.03	0.23	0.24	1.77	N.D.	0.30				
J-12 ^a	0.13	0.21	0.23							
UE-25p#1	0.18	1.04	1.34	11.44	N.D.	<0.002				

^aData from L. Benson et al.¹ ^bIntegral water sample. ^cBullfrog zone, 4th day. ^dBullfrog zone, 28th day.

e_{N.D.} = not detected.

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Fig. 4. Relative Na-K-Mg concentration in Yucca Mountain water.

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Fig. 10. Relative fluoride content as a function of relative sodium content for waters from Yucca Mountain.



Fig. 11. Ratio of ion activity product to equilibrium constant for calcite as a function of the ratio for magnesite for waters from Yucca Mountain. The dolomite saturation line is also shown.

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Fig. 12. Ratio of ion activity product to equilibrium constant for calcite as a function of the ratio for fluorite for water from Yucca Mountain.



Fig. 13. Total sulfate content as a function of total chloride content for water from Yucca Mountain (see also Fig. 14).



Fig. 14. Total sulfate content as a function of total chloride content for water from Yucca Mountain-expanded scale (see also Fig. 13).

ROCK-WATER INTERACTIONS

³⁶Cl Experiment for the Exploratory Shaft

Rock-Water Equilibrium

Reaction Path Calculations

³⁶Cl Surface Experiment



10CFR60.113: PREEMPLACEMENT GROUNDWATER TRAVEL TIME TO ACCESSIBLE ENVIRONMENT AT LEAST 1000 YEARS OR AS SPECIFIED BY COMMISSION

10CFR60.112: EPA RELEASE STANDARDS

NNWSI INFORMATION NEED 1.2.3: *HYDROLOGIC TRAVEL TIME IN UNSATURATED ZONE*







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³⁶CI WATER MOVEMENT TRACER INACTIVE CHLORIDE INTERFERENCE

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- WELL J-13 WATER 15 mg/L CI⁻
- EFFECT MITIGATED BY

 DOWNHOLE WATER TRACER
 LARGE BLOCKS
- CHLORIDE IN TUFF
 - MEASURED
 - SHORT LEACHING TIME



WHY ³⁶CI?

- CHLORIDE MOVEMENT IS UPPER BOUND FOR $\rm T_cO_4^-$ AND I $^-$
- NON SORBING
- LOW RATE OF ISOTOPIC EXCHANGE
- NO VAPOR-PHASE TRANSPORT
- HALF-LIFE RANGE 5×10^4 TO 2×10^6 YEARS
- .: UNIQUE CHARACTERISTICS, BUT COMPLEMENTARY TO ¹⁴C DATING



- ³⁶CI/CI MEASURED IN ~30 SAMPLES AS EXPLORATORY SHAFT IS MINED
- UNSATURATED ZONE PORE WATER DATED BETWEEN 5×10^4 AND 2×10^6 YEARS
- BOMB PULSE WILL INDICATE FRACTURE FLOW
- DATA USED FOR
 - CORRECT MODELING OF UNSATURATED ZONE HY-DROLOGIC FLOW
 - 10CFR60.113 REQUIREMENT OF 1000-YEAR FLOW
 - SETTING LIMIT ON $T_cO_4^-$ AND I⁻ MOVEMENT
- COMPLEMENTS OTHER PORE WATER DATING TECHNIQUES



SOLID PHASES ARE NOT ALL IN EQUILIBRIUM FLUID COMPOSITION IS CONTROLLED BY ROCK THE SYSTEM IS NOT IN EQUILIBRIUM EQUILIBRIUM MODELING IS USEFUL



Na-Si-AL PHASE DIAGRAM AT 25℃







WHAT EQUILIBRIA ARE CLOSELY APPROACHED?

Los Alamos

SILICON AND ALUMINUM FORM SILICATE FRAMEWORKS ARE NOT EASILY EXCHANGED WITH WATER EQUILIBRATION MAY BE POOR

SODIUM, POTASSIUM, AND CALCIUM MORE EASILY EXCHANGED EQUILIBRIUM MORE CLOSELY APPROACHED



- * WATER COMPOSITION AND LOCAL MINERALS IN TUFFACEOUS AQUIFER RESULT FROM DISSOLUTION OF VOLCANIC GLASS.
- * EQ3/6 REACTION-PATH MODEL OF RAINIER MESA MINERAL FORMATION (LA-9912-MS).



- * GLASS DISSOLUTION IN WATER WITH DISSOLVED CO₂.
- * CATION LEACH RATES FROM GLASS A FUNCTION OF pH.
- * AQUEOUS SILICA ACTIVITY CONTROLLED BY SUPPRESSING PRECIPITATION OF SILICA MINERALS.



ESTIMATED ZEOLITE FORMATION CONSTANTS

- * CLINOPTILOLITE, MORDENITE, AND HEULANDITE.
- * METHOD DEVELOPED BY Chen USE DATA FOR OTHER SILICATES FROM EQ3/6 DATA BASE.



* CLINOPTILOLITE Na₂(Al₂Si₁₀O₂₄)·8H₂O Ca(Al₂Si₁₀O₂₄)·8H₂O

 $K_2(Al_2Si_{10}O_{24}) \cdot 8H_2O$ Mg(Al_2Si_{10}O_{24}) \cdot 8H_2O

* MORDENITE Na(AlSi₅O₁₂)·3H₂O

 $K(AISi_5O_{12})\cdot 3H_2O$

* HEULANDITE Ca(Al₂Si₇O₁₈)·6H₂O



RESULTS WITH CRISTOBALITE CONTROLLING AQUEOUS SILICA ACTIVITY

- * MINERAL PRECIPITATES ARE CRISTOBALITE, SMECTITE CLAYS, CLINOPTILOLITE, AND MORDENITE.
- * WATER IS PRIMARILY A SODIUM-BICARBONATE WATER.


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75 C SUPPRESS OTZ CHALC NONTRON









75 C SUPPRESS OTZ CHALC NONTRON



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RESULTS WITH QUARTZ CONTROLLING AQUEOUS SILICA ACTIVITY

- * MINERAL PRECIPITATES ARE QUARTZ AND VARIOUS CLAYS; ALBITE UNDER SOME CONDITIONS.
- * WATER IS PRIMARILY A SODIUM-BICARBONATE WATER; HIGHER pH.

Na-Si-AL PHASE DIAGRAM AT 25℃



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Ca-Si-AL PHASE DIAGRAM AT 25℃





CONCLUSIONS FROM THE MODEL

- * AQUEOUS SILICA ACTIVITY CONTROLS THE STABILITY OF ZEOLITES CLINOPTILOLITE AND MORDENITE.
- * ESTIMATED ZEOLITE FORMATION CONSTANTS ARE REASONABLE.



PRE-1953 1953-1964 POST-1964

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USEFUL FOR TRACING INFILTRATION RATE OVER LAST 25 YEARS.

RELATED STUDIES GREENLAND ICE CORE BORDEN LAND FILL SOCORRO SANDY LOAM



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pheric box model.

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NNWSI PROGRAM

BOMB PULSE 34CI INFILTRATION

Los Alamos

PICK SAMPLING SITES NEAR EXPLORATORY SHAFT FRACTURED LOCATION ON YUCCA MOUNTAIN CREST YUCCA WASH

DETERMINE CHLORIDE PROFILES

DETERMINE ³⁶CI/CI PROFILES

SUBCONTRACTOR - HYDRO GEO CHEM, INC.

3-D MINERALOGY AND SORPTION STRATIGRAPHY

Mineralogy/Petrology Update

Evaluation of Statistical Methods



MINERALOGY/PETROLOGY PLAN

LOS Alamos

TRANSPORT PATHWAY MINERALOGY

A) HOST ROCK

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B) AWAY FROM THE REPOSITORY

• ALTERATION MINERALS

- A) ALTERATION HISTORY
- **B) MINERALS IN FRACTURES**

O MINERAL STABILITY

- A) PAST HISTORY OF YUCCA MOUNTAIN
- **B) SMECTITE DEHYDRATION**
- C) ZEOLITE DEHYDRATION AND REACTION KINETICS

O EXPLORATORY SHAFT TESTING

- A) DEEP SAMPLES WITH MAPPED-OUT RELATIONS
- **B) EXPOSURE OF CONTACTS BETWEEN MAJOR**
 - PETROLOGIC INTERVALS



TRANSPORT PATHWAYS -

POSSIBILITY OF SLOW TRANSPORT IN THE UNSATURATED ZONE: SIGNIFICANT RETARDATION MAY OCCUR IN OR NEAR THE HOST ROCK

- KNOWLEDGE OF THE HOST ROCK -
 - A) FOR ACCURATE DETERMINATION OF POSITION DURING CONSTRUCTION (E.G. - CONSTRUCTION ACROSS FAULT OFFSETS)
 - B) FOR TRANSFER OF RESULTS FROM LOCAL STUDIES TO OTHER PARTS OF THE EXPLORATION BLOCK



USW G-4

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1577.0 1000 1000 177.2.2.1 14000 1000 1700 177.2.2.1 14006.8	1945.4 Ministration		8 32 0 2 2 2 0 32 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2	222 222 222 222 222 222 222 222			С. С. С. (ЗУВАНЧА П. С. С. (ЗУВАНЧА П. С. С. С. С. С. С. (З) (ЗУВАНЧА П. С. (З)	400 400.4 UNPER 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	EXPAND 265.6 CAMOCK	ED SCALE 228 WITHC 228 WITHC 260 CRYSTAL - 24 20	STRATIGRAPHIC SUBUNITS (CONTACTS FROM R. SPENGLER, USGS) LITHOLOGIC PATTERN
USW G-4	1931	1190-[<u>7////</u>	-900-	934 - LV///////////////////////////////////	817-	004 - <u>17 - 27 - 18 - 27 - 27 - 27 - 27 - 27 - 27 - 27 - 2</u>	514 - <u>[]</u> -156.7 656 - <u>]</u> -169.5	41 7	(EXPANDED SCALE IN UPPERMOST PART) - GO - GO - 100	241- - 75.9	SAMPLE FOOTAGE (DEPTH) SILICIC OUANTZ SILICIC OUANTZ SAMPLED DEPTH (METERS)

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DEPTH METERS)

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MAJOR MINERAL COMPONENTS IN THE DEVITRIFIED TOPOPAH SPRING MEMBER, PAINTBRUSH TUFF. 10-1 00-5 S 1 USW H-6] ¥5¥ N 1500 TWA, VUCCA AND PAH MEMBERS, PARTERUGH TUFF Upper Paintbrush Tuff UPPER TOPOPAH TRIDYMITE 1300 10-2615 0-10% -LOWER TOPOPAH BRINK, BELOW ZONE OF ABLINDANT LITHOPHYSAE 10 äs 1180 BASAL VITROPHYRE 800 1600 Upper Pantbrush Tuff CRISTOBALITE 1300 LEVEL (m) 6-80% -10% Ø-1100 SEA 800 ABOVE 1600 Upper Paintbrush Tuff ELEVATION 1500 **OUARTZ** -20-40% 10-20% ----Upper Paintbrush Tuff AUKAU FELDSPAR 1300 00-00% 1100 1 km

VERTICAL EXAGGERATION = 3.5x

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ROCK TYPES BEYOND THE REPOSITORY: YUCCA MOUNTAIN

CLASSIFICATION OF SIGNIFICANT LITHOLOGIC VARIANTS
 NEAR THE REPOSITORY

- A) DEVITRIFIED TUFF (ALKALI FELDSPARS & SILICA MINERALS)
- B) ZEOLITES AND CLAY (HEULANDITE, CLINOPTILOLITE & SMECTITE)
- C) VITROPHYRE (DENSE GLASS)

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- D) VITRIC TUFF (POORLY CONSOLIDATED GLASS)
- CLASSIFICATION OF SIGNIFICANT LITHOLOGIC VARIANTS
 AWAY FROM THE REPOSITORY
 - A) DEVITRIFIED TUFF (ALKALI FELDSPARS & SILICA MINERALS)
 - B) ZEOLITIZED TUFF (CLINOPTILOLITE & MORDENITE)
 - 1) FOUR MAPPABLE ZEOLITIZATION INTERVALS



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OXYGEN ISOTOPE GEOTHERMOMETRY

from quartz fracture filling in altered Tpt vitrophyre

USW VH-2 3565 ft (1086.6 m) depth







Data needed to calculate temperature:

- δ^{18} O of quartz -- determined in lab
- assumed or known δ^{18} O of water that deposited the quartz
- relationship between temperature and quartz-water oxygen isotope fractionation factor (experimentally determined — published)





Oxygen isotope fractionation factor:

$$\alpha = 1000 + \delta^{18}O_{\text{quartz}} / 1000 + \delta^{18}O_{\text{water}}$$

Temperature--fractionation factor relationship:
(Friedman and O'Neil, 1977)

$$T^{\circ}C = \frac{3.38 \times 10^{6}}{10^{3} \ln \alpha + 2.90} - 273$$

$$= 91 \text{ or } 95^{\circ}C \quad \text{VH-2 } 3565$$
Los Alamos



- Tectonic tilting recorded by geopetal structures in zeolitized tuffs probably occurred before 11.3 million years ago.
- The age of zeolitization at Yucca Mountain is also probably pre—11.3 million years ago, based on textural relations in zeolitized tuffs.





- THE GREATEST DIFFERENCES BETWEEN FRACTURE MINERALOGY AND ROCK-MATRIX MINERALOGY OCCUR IN DEVITRIFIED OR GLASSY TUFFS
- ZEOLITES IN FRACTURES OCCUR FAR ABOVE THEIR SHALLOWEST OCCURRENCES AS BULK-ROCK ALTERATION PRODUCTS
- HEULANDITE IN UPPER FRACTURES AND NEAR THE LOWER TOPOPAH SPRING VITROPHYRE MAY HAVE AN ORIGIN VERY DIFFERENT FROM CLINOPTILOLITE IN LOWER ZEOLITIZED INTERVALS

X-RAY DIFFRACTION ANALYSES OF FRACTURE FILLINGS USW G-4 UNSATURATED ZUNE

Unit	Sample	Estimated % of Fracture Surface Coated	Estimated Max. 7 Rock Contamination	Smectite	Cli cilolite/ Heulandite	Mordenite	Quartz	Cristobalite	Alkli Feldspar	Triaymite
الاستينية المتنابية عن المن اب	810	40	5	23 ± 5	-	59 ± 10	18 ± 3			
	887 '	· 90	30	2 ± 1			12 ± 2	5 ± 3 · · · ·	74 ± 10	7 ± 4
[a]	984	100	20				79 ± 5	-1	14 ± 5	6 ± 3
iKya	1001	10	40	~1			21 ± 4	5 ± 3	73 ± 10	
tho	1008	100	15	~1		11 ± 5	40 ± 5	10 ± 5	38 ± 10	
Ë	1038 X	100	20	15 ± 5			24 ± 8	5 ± 3	36 ± 10	20 ± 5
	1072 #1	100	30				65 ± 5	-1	22 ± 10	12 ± 4
٢	1072 #2	30	30	~1		7 ± 4	14 ± 1	8 ± 4	61 ± 10	9 ± 4
	1083 #3	90	10	6 ± 2			9 ± 2	5 ± 3	65 ± 10	14 ± 5
	1148	50	25	2 ± 1			22 ± 8	7 ± 4	69 ± 10	
	1160	50 .	10				32 ± 5	6 ± 4	62 ± 10	
	1173	25	10	3 ± 1		58 ± 15	23 ± 7	16 ± 7		
a c So	1201	95	50				43 ± 9	9 ± 4	48 ± 15	
T to	1244	60	15	~1			40 ± 5	8 ± 4	51 ± 10	
÷÷.	1254 #1	100	15		52 ± 10	13 ± 5	20 ± 2	15 ± 5		
	1254 #3	100	10		47 ± 10 H	11 ± 5	26 ± 5	16 ± 4		
nde	1258	100	5	7 ± 3	29 ± 5		12 ± 3	7 ± 4	45 ± 10	
	1309	90	5		100 H					
	1341 blue	100	50	2 ± 1	20 ± 10	12 ± 6	5 ± 3	6 ± 4	55 ± 10	
Vitrophyre	1341 cream	100	10	67 ± 10	28 ± 10		5±4			
	1341 betge	35	10	5 ± 2	80 ± 10 H	15 ± 5				
Vitric	1350	100	20	20 ± 10	80 ± 10					
	1362	20	>50	3 ± 1	42 ± 10	18 ± 5		6 ± 4		31 ± 10
Nomelded	1381	100	50	~1	67 ± 10 C			32 ± 5		
	1436	100	20	-1	99 C					
Bedded and	1513	100	20		55 ± 10 ·	22 ± 10	2 ± 2	15 ± 8		6 ± 2
Nonwe1ded	1542	100	15	2 ±1	74 ± 10	24 ± 10				
Zeolitic	1643	100	20		20 ± 10	80 ± 10				
	1694 Not roci	readily discernil matrix	ble from	2 ± 1	70 ± 10	28 ± 10			• • • • • • •	
Bedded	1707	70	10		4 ± 2	96 ± 4			A.F	. . .
Zeolitic	1716	30	50		25 ± 10	25 ± 10	18 ± 3	_	27 ± 10	5 ± Z
	1763 white	100	Can't tell as	88 ± 10	8 ± 3			4 ± 2		
Nonwe Ided SWL	1763 orange	100	rock grains	68 ± 10	23 ± 10	5 ± 3		4 ± 3		

Topopah Spring Member

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Prow Pass



SEQUENCE OF FRACTURE COATINGS IN THE UNSATURATED ZONE OF USW G-4



PALEO FLOW PATHS -

Sr-ISOTOPIC VARIATIONS BETWEEN ZEOLITE GROUPS (UPWARD

FLOW, DOWNWARD FLOW; DISTANT OR LOCAL TRANSPORT?)

PALEO WATER LEVELS -

- A) PRESERVATION OF UNPROTECTED, NONWELDED GLASS SHARDS
- **B) TEXTURAL TRANSFORMATIONS WITHIN CLINOPTILOLITE**

OCCURRENCES

Sr - ISOTOPIC DIFFERENCES BETWEEN CARBONATE AND TUFF SOURCES

	Source Rock	Rb (ppm)	Sr (ppm)	875r/865r	Comments
Ι.	Caliche				
	Carboniferous carbonates	1-40	200-1000	0.7077-0.7085	Reworked Paleozoic
	Devonian carbonates		**	0.7078-0.7084	may be the dominant
	Silurian carbonates			0.7088	components of the caliche layer at
	Ordovician carbonates	**	60	0.7077-0.7088	Yucca Mountain. Presented here are
	Cambrian carbonates	4	W	0.7089-0.7094	worldwide averages taken from Peterman et al., 1970 and Veizer, 1974.
11.	Paintbrush Tuffs				
,	Tiva Canyon Member	56-103	142-205	0.7075-0.7135	The 0.7075 87Sr/ 86Sr ratio of the
•	Pah Canyon Member	ND	ND	0.7101	Tiva Canyon Member is
	. Topopah Spring Member	125-195	22-110	0.7107-0.7179	latitic caprocks.
	All Tuffs	56-195	22-205	0.7075-0.7179	However, these cap- rocks have generally devitrified to stable mineral phases and are not subject to dissolution. Ratios taken from Noble and

Hedge, 1969.

POLAROID 1542 22m





1310 POLAROID



STATISTICAL METHODS FOR:

- A. EXTRAPOLATION OF LIMITED BOREHOLE DATA (MAPPING PROBLEM)
- B. DETERMINING THE VARIATION IN MINERALOGIC
 COMPOSITION OF THE REPOSITORY HORIZON
 AND ALONG POTENTIAL FLOW PATHS



- A. FOR MAPPING PETROLOGIC ZONES BELOW REPOSITORY HORIZONS:
 - UP TO FOURTEEN HOLES (FIVE CONTINUOUSLY CORED)
 - CONTACTS BETWEEN ZONES KNOWN TO WITHIN ONE TO FOUR METERS
- B. FOR DETERMINING MINERALOGIC VARIABILITY WITHIN THE

TOPOPAH SPRING MEMBER OF THE EXPLORATION BLOCK:

- UP TO FIVE HOLES (TWO CONTINUOUSLY CORED)
- X-RAY DIFFRACTION MEASUREMENTS OF CRISTOBALITE, TRIDYMITE, SMECTITE, QUARTZ AND ALKALI-FELDSPAR


- WHAT TECHNIQUES ARE AVAILABLE?
- WHAT IS THE MAGNITUDE OF ERROR ASSOCIATED WITH ESTIMATES OF THICKNESS AND DEPTH OF PETROLOGIC ZONES ACROSS THE EXPLORATION BLOCK?
- WHAT SAMPLING (BOREHOLE) DENSITY WOULD BE REQUIRED TO REFINE THESE ESTIMATES TO A SPECIFIED LEVEL OF PRECISION?



• TREND SURFACE FITTING WITH STANDARD PREDICTION ERRORS OR "JACKKNIFED" ERROR ESTIMATES

• KRIGING, WITH STANDARD KRIGING ERROR AND "JACKKNIFED" EVALUATION OF MODEL FIT



• SMALL DATA SET MEANS CHOICE OF SURFACE FORM LIMITED TO FORMS WITH FEW PARAMETERS (E.G., QUADRATIC)

• POOR FIT AND LARGE ERROR ESTIMATES

• POSSIBLY USEFUL AS PRELIMINARY TO KRIGING



• SMALL DATA SET MEANS CHOICE OF MODEL

IS SOMEWHAT ARBITRARY; VARIABILITY OF

RAW VARIOGRAM IS LARGE.

• REASONABLE ERROR ESTIMATES CAN BE OBTAINED.



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- HOW DOES THE VARIABILITY BETWEEN HOLES COMPARE TO THE VARIABILITY WITHIN HOLES?
- HOW DO X-RAY DIFFRACTION RESULTS COMPARE WITH PETROLOGIC ASSESSMENTS?
- CAN ESTIMATES OF VARIABILITY BE IMPROVED BY DRILLING MORE HOLES OR BY ANALYZING MORE SAMPLES FROM AVAILABLE HOLES?



• ANALYSIS OF VARIANCE (ANOVA) FOR

ZONE AND BOREHOLE EFFECTS

CANONICAL DISCRIMINATION FOR COMPARISON OF MINERALOGIC AND PETROLOGIC ASSESSMENT



• TWO-WAY ANOVA FOR INDIVIDUAL MINERALS SUGGESTS STRONG BOREHOLE EFFECTS AND BOREHOLE-ZONE INTERACTIONS, IMPLYING LACK OF HOMOGENEITY BETWEEN HOLES.

• SOME LINEAR COMBINATIONS (E.G., QUARTZ PLUS SMECTITE) EXHIBIT ONLY ZONE EFFECTS.



- GOOD MATCH BETWEEN PETROLOGIC AND MINERALOGIC ASSESSMENT FOR G-3 AND G-4.
- PREPARATION OF SAMPLES IS VERY IMPORTANT.
- ACCURACY OF X-RAY DIFFRACTION MEASUREMENTS ON CONSISTENTLY PREPARED SAMPLES IS SUFFICIENT TO BE USEFUL.











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MINERAL STABILITY

Hydrothermal Stability

Thermal Stability

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MINERAL STABILITY

WHY IS IT IMPORTANT?

Los Alamos

WILL SORPTIVE MINERALS REMAIN THROUGHOUT REPOSITORY LIFETIME? PREDICTION OF CHEMICAL CONDITIONS CONSTRAINTS ON PAST HYDROLOGIC CONDITIONS





THE MINERALOGY IN YUCCA IS EVOLVING TOWARD LOW SILICA ACTIVITIES

SIGNIFICANT CHANGE ON REPOSITORY TIMESCALE IS UNLIKELY

IF CLINOPTILOLITE BREAKS DOWN IT IS LIKELY TO PRODUCE CLAYS No-Si-AL PHASE DIAGRAM AT 25°C







MINERAL STABILITY IMPORTANT PARAMETERS

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ACTIVITY OF SiO₂ LITHOSTATIC PRESSURE WATER PRESSURE TEMPERATURE







IMPORTANT ADDITIONAL DATA NEEDED

Los Alamos

THERMODYNAMIC DATA ON CLAY AND ZEOLITE ENDMEMBERS

THERMODYNAMICS OF CLAY AND ZEOLITE SOLID SOLUTIONS

KNOWLEDGE OF CONTROLS ON SILICA POLYMORPH DISSOLUTION AND PRECIPITATION



DATA REFERENCE FREE ENERGY HEAT CAPACITY ENTROPY OF MIXING VOLUME

EXPERIMENTAL TECHNIQUES X-RAY DIFFRACTION MAGIC ANGLE NMR CALORIMETRY PHASE EQUILIBRIA DETERMINATIONS SOLUBILITY MEASUREMENTS



CRYSTAL STRUCTURE OF ANALCIME

Los Alamos



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ZEOLITE AND SMECTITE THERMAL STABILITY

Los Alamos

- IMPORTANT MINERAL REACTIONS CAN OCCUR IN A PARTIALLY SATURATED ENVIRONMENT AT TEMPERATURES BELOW 200°C.
 - SMECTITE DEHYDRATION IRREVERSIBLE?
 - CLINOPTILOLITE DEHYDRATION
 - GLASS DEHYDRATION
- THESE REACTIONS WILL ONLY BE IMPORTANT IN THE NEAR FIELD BUT POTENTIALLY CAN AFFECT RETARDATION BY SORPTION AND ROCK PERMEABILITY
- EXPERIMENTS CONDUCTED
 - EFFECTS OF HEATING ON CLINOPTILOLITE, HEULANDITE, AND MORDENITE VOLUME
 - DEHYDRATION BEHAVIOR OF CLINOPTILOLITES, HEULANDITE, MORDENITE, AND VOLCANIC GLASS.
 - KINETICS OF DEHYDRATION OF Ca-CLINOPTILOLITE
- EXPERIMENTS PLANNED
 - KINETICS OF DEHYDRATION OF CLINOPTILOLITES
 - EFFECT OF PH,O ON SMECTITE AND ZEOLITE DEHYDRATION
















KEY RADIONUCLIDES AND SOLUBILITY

Key Radionuclides

Solbuility Calculations

Solubility Measurements



- * REPOSITORY INVENTORY WHICH RADIONUCLIDES ARE PRESENT IN LARGE QUANTITIES?
- * EPA LIMITS HOW WELL MUST DIFFERENT RADIONUCLIDES BE CONTAINED?
- * RETARDATION PROCESSES WHICH PROCESSES AFFECT WHICH ELEMENTS?

REPOSITORY INVENTORY

(Ci/1000MTHM)

	100 yr	1000 yr	10,000 yr	100,000 yr
PWR				
spent fuel				
A. P.	3.21×10 ⁵	8.40×10 ³	6.40×10^{3}	2.47×10 ³
Act.	6.79×10 ⁶	1.72×10 ⁶	4.43×10 ⁵	3.87×10 ⁴
F. P.	3.39×10 ⁷	1.93×10 ⁴	1.86×10 ⁴	1.43×10 ⁴
Total	4.10×10 ⁷	1.75×10 ⁶	4.68×10^{5}	5.55×10 ⁴
PWR HLW				
A. P.	3.21×10 ⁵	8.40×10 ³	6.40×10 ³	2.47×10 ³
Act.	3.21×10^{5}	8.27×10 ⁴	2.05×10 ⁴	2.53×10 ³
F. P.	3.35×10 ⁷	1.92×10 ⁴	1.85×10 ⁴	1.42×10 ⁴
Total	3.41×10 ⁷	1.10×10 ⁵	4.54×10 ⁴	1.92×10 ⁴

REPOSITORY INVENTORY

	Primary	Radionucli	des (% total	l activity)
	100 yr	1000 yr	10,000 yr	100,000 yr
PWR				
spent fuel				
	¹³⁷ Cs 49% [*]	²⁴¹ Am 51%	²³⁹ Pu 51%	²³⁹ Pu 32%
	⁹⁰ Sr 34%	²⁴⁰ Pu 27%	²⁴⁰ Pu 39%	⁹⁹ Tc 17%
	²⁴¹ Am 9%	²³⁹ Pu 17%	⁹⁹ Tc 3%	⁹³ Zr 6% [*]
	²³⁸ Pu 3%	²³⁹ Np 1%	²⁴³ Am 1%	⁵⁹ Ni 4%
PWR HLW				
	¹³⁷ Cs 58%	²⁴¹ Am 38%	⁹⁹ Tc 28%	⁹⁹ Tc 49%
	⁹⁰ Sr 41%	²⁴³ Am 14%	²⁴³ Am 15%	⁹³ Zr 18% [*]
	²⁴¹ Am 0.5%	²³⁹ Np 14%	²³⁹ Np 15%	⁵⁹ Ni 11%
		⁹⁹ Tc 12%	⁵⁹ Ni 10%	²³⁹ Pu 3%

Activity of short-lived daughters included.

RADIONUCLIDES - EPA LIMITS

Cumulative release limits at 10000 yr (Ci/1000MTHM)

$^{241}Am = 100$	240 Pu = 100
243 Am = 100	242 Pu = 100
$^{14}C = 100$	226 Ra = 100
$^{135}Cs = 1000$	90 Sr = 1000
$^{137}Cs = 1000$	99 Tc = 10000
$^{237}Np = 100$	$^{126}Sn = 1000$
238 Pu = 100	other $\alpha = 100$
$^{239}Pu = 100$	other non $-\alpha = 1000$

RADIONUCLIDES - OLD EPA LIMITS

Cumulative release limits at 10000 yr (Ci/1000MTHM)

$^{239}Pu = 100$	other $\alpha = 10$ other non $-\alpha = 500$
$^{237}Np = 20$ $^{238}Du = 400$	$^{126}Sn = 80$
$^{137}Cs = 500$	99 Tc = 10000
$^{135}Cs = 2000$	90 Sr = 80
$^{14}C = 200$	226 Ra = 3
243 Am = 4	242 Pu = 100
$^{241}Am = 10$	240 Pu = 100

RADIONUCLIDES - PWR SPENT FUEL

(order based on curie content/EPA limit)

100 yr	1000 yr	10,000 yr	100,000 yr
²⁴¹ Am	²⁴¹ Am	²³⁹ Pu	²³⁹ Pu
¹³⁷ Cs	²⁴⁰ Pu	²⁴⁰ Pu	²³⁴ U
⁹⁰ Sr	²³⁹ Pu	²⁴³ Am	²⁴² Pu
²³⁸ Pu	²⁴³ Am	²³⁴ U	²³⁷ Np
²⁴⁰ Pu	²³⁴ U	²⁴² Pu	²²⁶ Ra
²³⁹ Pu	²⁴² Pu	²³⁷ Np	²³⁰ Th
²⁴¹ Pu	¹⁴ C	⁵⁹ Ni	²³³ U
²⁴⁴ Cm	²³⁷ Np	¹⁴ C	²³⁶ U
⁶³ Ni	²³⁸ Pu	²³⁶ U	²²⁹ Th
²⁴³ Am	⁵⁹ Ni	²³⁸ U	²³⁸ U

RADIONUCLIDES - PWR HLW

(order based on curie content/EPA limit)

1000 yr	10,000 yr	100,000 yr
²⁴¹ Am	²⁴³ Am	²³⁹ Pu
²⁴³ Am	²³⁹ Pu	²³⁷ Np
²⁴⁰ Pu	²⁴⁰ Pu	⁵⁹ Ni
²³⁹ Pu	⁵⁹ Ni	⁹³ Zr
¹⁴ C	²³⁷ Np	^{93m} Nb
⁵⁹ Ni	¹⁴ C	²³³ U
²³⁷ Np	⁹³ Zr	²²⁹ Th
⁹³ Zr	^{93m} Nb	⁹⁹ Tc
^{93m} Nb	⁹⁹ Tc	¹²⁸ Sn
⁹⁹ Tc	⁹⁴ Nb	²³⁴ U
	1000 yr ²⁴¹ Am ²⁴³ Am ²⁴⁰ Pu ²³⁹ Pu ¹⁴ C ⁵⁹ Ni ²³⁷ Np ⁹³ Zr ^{93m} Nb ⁹⁹ Tc	1000 yr10,000 yr241 Am243 Am243 Am239 Pu240 Pu240 Pu239 Pu59 Ni14 C237 Np59 Ni14 C237 Np93 Zr93 Zr93 m Nb93 m Nb99 Tc99 Tc94 Nb

RADIONUCLIDES

²²⁶Ra decay scheme



History

- First created in 1975 to model hydrothermal reactions between sea water and basalt
- Present development is focused on interactions in geologic disposal of nuclear waste

Calculate

- Chemical speciation (simple ions and complexes) in aqueous solutions
- Thermodynamic equilibrium in rock/water systems
- Progress of reacting rock/water systems (kinetic model)

Status

- Operational in current mode at LLNL, other national laboratories, universities, and industry
- We are continuing to improve these codes and extend their capabilities

• The importance of using geochemical codes to make long-term predictions of radionuclide migration is well established.

- EQ3/6 is the geochemical code now being upgraded for use by the Nevada Nuclear Waste Storage Investigations (NNWSI) and Office of Nuclear Waste Isolation (ONWI) Programs.
- Geochemical modeling task plans for both the tuff (NNWSI) and salt (ONWI) programs have been developed.
- The current geochemical modeling work includes code and data base development tasks, but does NOT include basic research.

EQ3/6 IS A CODE PACKAGE CONSISTING OF

- EQ3NR SPECIATION-SOLUBILITY CODE
- EQ6 REACTION-PATH CODE
- MCRT THERMODYNAMIC DATA PROCESSOR

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- DATA FILES
- BATA FILE PREPROCESSORS
- EQ6 GRAPHICS POSTPROCESSOR

STATUS OF EQ3/6 DOCUMENTATION

- EQ3NR USER'S GUIDE (1983)
- EQ6 USER'S GUIDE (IN REVISION, TBP 1984)
- EQ3/6 BRINE MODIFICATIONS (IN REVISION, TBP 1984)
- MCRT USER'S GUIDE (NO DRAFT YET, TBP 1984)
- EQ6 GRAPHICS GUIDE (1ST DRAFT, TBP 1984)
- EQ6 FIXED FUGACITY (IN REVISION, TBP 1984)
- EQ6 PPTN KINETICS (IN PREPARATION, TBP 1984)

EQ3/6 VERSION 3245 IS NOW UNDER DEVELOPMENT

IMPROVEMENTS INCLUDE

- BRINE MODELING CAPABILITY (PITZER'S EQUATIONS) ~
- "ECONOMY" MODE FOR SOME EQ6 CALCULATIONS
- PRECIPITATION KINETICS
 (SIMPLE TREATMENT AT PRESENT)
- FULLY DOCUMENTED THERMODYNAMIC DATA FILE

EQ3NR - A MULTI-OPTION SPECIATION-SOLUBILITY CODE

- MODELS THE THERMODYNAMIC STATE OF AN AQUEOUS SOLUTION
- TYPICAL INPUTS
 - PH. OTHER SPECIFIC ION ELECTRODE MEASUREMENTS
 - EH (NOT RECOMMENDED)
 - TOTAL CONCENTRATIONS
 - ELECTRICAL BALANCE (ADJUST SPECIFIED ION)
 - SOLUBILITY EQUILIBRIA (ASSUMPTIONS)
 - GAS EQUILIBRIA (ASSUMPTIONS)
- TYPICAL OUTPUTS
 - PH
 - **REDOX POTENTIALS**
 - FREE CONCENTRATIONS
 - TOTAL CONCENTRATIONS
 - DEGREE OF ELECTRICAL IMBALANCE
 - SATURATION INDICES
 - A MODEL OF THE WATER TO FEED TO EQ6

EQ6 - A MULTI-OPTION REACTION PATH CODE

- Models the Path of a Reacting Aqueous System (example: water + rock)
- NON-EQUILIBRIUM MASS TRANSFER IS TREATED BY EITHER
 - SIMPLE TITRATION (OLD)
 - **•** KINETIC RATE LAWS (NEW)

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- PREDICTS EVOLUTION OF FLUID COMPOSITION AND FORMATION OF SECONDARY SOLIDS
- TEMPERATURE CAN BE CONSTANT OR VARIABLE

EQ6 MODELS MORE THAN ONE KIND OF PHYSICAL SYSTEM

• ZERO-DIM: CLOSED SYSTEM, THERMODYNAMICALLY HOMOGENEOUS (GOOD MATCH TO ROCKING DICKSON AUTOCLAVES)

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- PSEUDO-ONE-DIM: FLOW-THROUGH OPEN SYSTEM, FOLLOWS A PACKET OF WATER
- PSEUDO-ONE-DIM: FLOW-THROUGH OPEN SYSTEM, FOCUSED ON A FIXED VOLUME (E.G., A FLOW-THROUGH LEACHING CELL) - PLANNED
- ONE-DIM: FEASIBLE BUT NOT PLANNED

EQ3/6 code development tasks planned for the NNWSI (tuff) and ONWI (salt) Programs, 1984 to 1988

- 1. Precipitation Kinetics
 - Complements the existing EQ6 capability to model dissolution kinetics. Will eventually be expanded to include second-order effects and nucleation phenomena.

2. Fixed Fugacity Option

Allows us to model systems open to gases; especially the unsaturated zone, which contains a large reservoir of CO_2 and O_2 .

3. Sorption

Our eventual goal is the capability to model complex sorption isotherms, sorption kinetics and non-equilibrium sorption processes which control radionuclide migration.

4. Redox Disequilibrium and Kinetics

Allows us to model systems where redox couples are not in equilibrium with each other. Redox kinetics will be incorporated into EQ6 for modeling corrosion experiments.

5. Glass/Water Interactions Model

To model waste glass leaching, code modifications will be needed to account for the effect of diffusion, the development of an alteration layer and other effects.

6. Radiolysis

Allows us to model the effect of radiolysis products on leaching, corrosion and other processes operating in the waste package environment.

Code development tasks (continued)

7. Site-Mixing Concept

Solid solutions are now represented in EQ3/6 as percentages of end-member components. To model reactions involving clays and zeolites, a more sophisticated approach such as the site-mixing concept is needed.

Other Tasks That May Be Included

- 1. Geochemical Flow Model May be needed to model flow-through leaching experiments.
- 2. Isotopic Fractionation

May be needed to model the preferential substitution of one isotope for another of the same element in a solid or aqueous phase.



- * SOLUBILITY WHICH ELEMENTS HAVE LOW SOLUBILITY? HOW WILL SOLUBILITY LIMIT TRANSPORT?
- * SORPTION WHICH ELEMENTS HAVE HIGH K_d?
- * PHYSICAL PROCESSES
- * RETARDATION OF GASES OR VAPORS ¹⁴C, ¹²⁹I, VOLATILE OXIDES



- * SORPTION U, Pu, Np, Am, Th, Ba, Ra, Cs, Sr, Tc, Sn, Se, Ce, Eu, I, Pr.
- * SOLUBILITY AND SPECIATION -Pu, Am, Np.



- * DETERMINATION OF SOLUBILITY -CALCULATIONS. -MEASUREMENTS.
- * USE OF SOLUBILITY DATA -RADIONUCLIDE DISSOLUTION MODELS.



- * DEVELOP A THERMODYNAMIC DATA BASE FOR WASTE ELEMENTS.
- * CALCULATE SOLUBILITY AND SPECIATION IN WATER AT YUCCA MOUNTAIN.
- * VALIDATE MODEL BY COMPARASON WITH EXPERIMENTS.
- * USE EMPIRICAL SOLUBILITY DATA WHERE NECESSARY.



SOLUBILITY AND SPECIATION CALCULATIONS WITH EQ3/6

- * THREE WATERS J-13, UE25p-1, AND H-3.
- * SIX ELEMENTS U, Pu, Am, Sr, Ra, AND Tc.

WATER COMPOSITIONS

.

	J-13	UE25p-1	H-3
Na ⁺ (mm/l)	1.96	7.43	5.39
K ⁺ (mm/l)	0.14	0.34	0.04
Ca ²⁺ (mm/l)	0.29	2.19	0.02
Mg ²⁺ (mm/l)	0.07	1.31	0.00
SiO ₂ (aq) (mm/l)	1.07	0.62	0.60
CO_3^{2-} (total) (mm/l)	2.81	16.1	4.04
Cl ⁻ (mm/l)	0.18	1.04	0.23
S04 ⁻ (mm/l)	0.19	1.34	0.32
F ⁻ (mm/l)	0.11	0.18	0.28
pH	7.0	6.7	9.4
Eh (mV)	700	360	-143

URANIUM

	J-13	UE25p-1	H-3
Solub. (m/l)	3.65×10 ⁻³	1.74×10^{-3}	4.05×10^{-8}
Solid	Schoepite	Rutherfordine	Uraninite
Primary Aqueous Species	(UO ₂) ₂ CO ₃ (OH) ₃ (98%) UO ₂ (CO ₃) ₂ ² (1%)	$UO_2(CO_3)_2^{2-}$ (54%) $(UO_2)_2CO_3(OH)_3^-$ (31%) $UO_2(CO_3)_3^{4-}$ (13%) UO_2CO_3 (2%)	$UO_2(CO_3)_3^{4-}$ (86%) $U(OH)_5^-$ (8%) $UO_2(CO_3)_2^{2-}$ (7%)

PLUTONIUM

	J-13	UE25p-1	H-3
Solub. (m/l)	1.79×10 ⁻⁶	3.11×10 ⁻⁸	1.33×10 ⁻⁵
Solid	Pu(OH) ₄	Pu(OH) ₄	Pu(OH)4
Primary Aqueous Species	$PuO_{2}^{+} (71\%)$ $PuO_{2}F_{3}^{-} (20\%)$ $Pu(OH)_{5}^{-} (3\%)$ $PuO_{2}(CO_{3})_{2}^{2-} (2\%)$ $PuO_{2}F_{4}^{2-} (2\%)$	Pu(OH) ₅ (94%) Pu(OH) ₄ (6%)	Pu(OH) ₅ (100%)

AMERICIUM

	J-13	UE25p-1	H-3
Solub. (m/l)	9.87×10 ⁻⁹	2.16×10 ⁻⁸	6.85×10^{-10}
Solid	Am(OH)CO ₃	Am(OH)CO ₃	Am(OH)CO ₃
Primary Aqueous Species	$AmCO_3^+$ (80%) $AmOH^{2+}$ (8%) AmF^{2+} (4%) Am^{3+} (3%) $Am(CO_3)_2^-$ (3%)	$AmCO_{3}^{+}$ (83%) $Am(CO_{3})_{2}^{-}$ (6%) AmF^{2+} (4%) $AmSO_{4}^{+}$ (2%) $AmOH^{2+}$ (2%) Am^{3+} (2%)	$Am(CO_3)_2^-$ (46%) $Am(OH)_3$ (36%) $Am(OH)_2^+$ (12%) $AmCO_3^+$ (5%)

STRONTIUM

	J-13	UE25p-1	H-3
Solub. (m/l)	8.04×10 ⁻⁴	5.27×10 ⁻⁴	3.28×10 ⁻⁶
Solid	Strontianite	Strontianite	Strontianite
Primary Aqueous Species	Sr ²⁺ (96%) SrSO ₄ (4%)	Sr ²⁺ (86%) SrSO ₄ (14%)	Sr ²⁺ (96%) SrSO ₄ (4%)

RADIUM

	J-13	UE25p-1	H-3
Solub. (m/l)	3.39×10 ⁻⁷	9.29×10 ⁻⁸	2.94×10 ⁻⁷
Solid	RaSO ₄	RaSO ₄	RaSO ₄
Primary Aqueous Species	Ra ²⁺ (99%)	Ra ²⁺ (99%)	Ra ²⁺ (99%)

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TECHNETIUM

	J-13	UE25p-1	Н-З
Solub. (m/l)	Large	Large	2.06×10 ⁻¹²
Solid			Tc ₃ O ₄
Primary Aqueous Species	TcO ₄ (100%)	TcO ₄ (100%)	TcO ₄ (91%) TcO(OH) ₂ (9%)



PROJECT

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Loe Alamos

- ASSESS THE IMPORTANCE OF RADIONUCLIDE DISSOLUTION SOLUBILITY IN LIMITING (LA-9995-MS). *
- TWO DISSOLUTION MODELS:
 - SATURATION LIMITED.
- DIFFUSION LIMITED.



SATURATION-LIMITED MODEL

- * WATER FLOW THROUGH THE REPOSITORY IS SATURATED WITH WASTE ELEMENTS.
- * SIMPLE.
- * CONSERVATIVE.


RADIONUCLIDE DISSOLUTION

Los Alamos

DIFFUSION-LIMITED MODEL

- * DIFFUSION OF WASTE ELEMENTS INTO WATER FLOWING PAST WASTE LIMITS DISSOLUTION.
- * WASTE ELEMENTS SATURATED AT WATER/WASTE INTERFACE.
- * COMPLEX.
- * REALISTIC.



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- * COMPARE SOLUBILITY-LIMITED DISSOLUTION WITH BULK WASTE-FORM DISSOLUTION.
- * ELEMENTS WITH A DISSOLUTION RATE LESS THAN THE BULK FRACTIONAL DISSOLUTION RATE ARE LIMITED BY SOLUBILITY.
- * ELEMENTS WITH A DISSOLUTION RATE GREATER THAN THE BULK FRACTIONAL DISSOLUTION RATE ARE LIMITED BY BULK WASTE DISSOLUTION.

DISSOLUTION RATE VS SOLUBILITY



RADIONUCLIDE DISSOLUTION

Waste element solubilities (m/l)(oxidizing conditions in Well J-13 water) $Am = 1.0 \times 10^{-8}$ Cs = large $Pu = 1.8 \times 10^{-6}$ Tc = large $U = 2.1 \times 10^{-4}$ $Np = 3.0 \times 10^{-3}$ $Sr = 9.4 \times 10^{-4}$ $Ra = 1.0 \times 10^{-7}$ $Sn = 1.0 \times 10^{-9}$ C = large

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RADIONUCLIDE DISSOLUTION

NOMINAL MODEL PARAMETERS

Recharge rate = 8 mm/yr. Repository area = 114 m²/MTHM. Water flow = 910 l/MTHM yr. Bulk dissolution rate = 1.0×10^{-4} /yr. Rock porosity = 10 %. Water velocity = 8.0×10^{-2} m/yr. Diffusion coef. = 1.0×10^{-10} m²/yr. L = 4.5 m (SF) or 3.0 m (HLW).

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SOLUBILITY-LIMITED DISSOLUTION

- * SPENT FUEL:
 SLM U, Pu, Am, Sn.
 DLM U, Pu, Am, Sn, Ra.
- * HIGH LEVEL WASTE: SLM - Am, Sn. DLM - U, Pu, Am, Sn.



- * OVERALL MEASURE OF THE EFFECT OF SOLUBILITY IS THE RATIO OF THE RELEASE RATE OF RADIOACTIVITY (CI/MTHM yr) TO THE REPOSITORY INVENTORY (CI/MTHM).
- * COMPARE SOLUBILITY-LIMITED DISSOLUTION WITH BULK WASTE DISSOLUTION.

SATURATION-LIMITED RADIOACTIVITY RELEASE



DIFFUSION-LIMITED RADIOACTIVITY RELEASE





EFFECT OF VARYING MODEL PARAMETERS

- * WATER RECHARGE RATE.
- * WASTE ELEMENT SOLUBILITIES.
- * BULK FRACTIONAL DISSOLUTION RATE.





RADIONUCLIDE DISSOLUTION

EFFECT OF VARYING PARAMETERS - AT 1000 yr.

(Release rate/inventory (1/yr))

	Nominal	Reduced Recharge ^a	Reduced Solubility ^b
Sat.—limited model Spent fuel HLW	2.9×10 ⁻⁵ 4.8×10 ⁻⁵	4.5×10 ⁻⁶ 3.7×10 ⁻⁵	2.1×10 ⁻⁶ 3.3×10 ⁻⁵
Diff.—limited model			{
Spent fuel	1.9×10 ⁻⁶	1.8×10 ⁻⁶	9.5×10 ⁻⁷
· HLW	2.9×10 ⁻⁵	2.8×10^{-5}	1.7×10^{-5}
^a factor of 10. ^b fa	ctor of 100.	Los	Alamos

RADIONUCLIDES - PWR SPENT FUEL DIFFUSION-LIMITED MODEL RESULTS

(order based on release rate/EPA limit)

100 yr	1000 yr	10,000 yr	100,000 yr
¹³⁷ Cs	¹⁴ C	²³⁷ Np	²³⁷ Np
⁹⁰ Sr	²³⁷ Np	¹⁴ C	²²⁶ Ra
¹⁴ C	²⁴⁰ Pu	²³⁹ Pu	²³⁹ Pu
²³⁸ Pu	²³⁹ Pu	²⁴⁰ Pu	⁹⁹ Tc
²⁴⁰ Pu	²⁴¹ Am	²²⁶ Ra	¹³⁵ Cs
²³⁷ Np	⁹⁹ Tc	⁹⁹ Tc	²⁴² Pu
²³⁹ Pu	¹³⁵ Cs	¹³⁵ Cs	²³⁴ U
²⁴¹ Am	²²⁶ Ra	²⁴³ Am	²⁴³ Am
⁹⁹ Tc	²⁴³ Am	²⁴² Pu	²³³ U

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RADIONUCLIDES - PWR HLW DIFFUSION-LIMITED MODEL RESULTS

(order based on release rate/EPA limit)

100 yr	1000 yr	10,000 yr	100,000 yr	
¹³⁷ Cs	¹⁴ C	¹⁴ C	²³⁹ Pu	
⁹⁰ Sr	²⁴⁰ Pu	²³⁹ Pu	²³⁷ Np	
²³⁸ Pu	²³⁷ Np	²³⁷ Np	⁹⁹ Tc	
¹⁴ C	²³⁹ Pu	²⁴⁰ Pu	¹³⁵ Cs	
²⁴⁰ Pu	⁹⁹ Tc	⁹⁹ Tc	²²⁶ Ra	
²³⁷ Np	¹³⁵ Cs	¹³⁵ Cs	²³³ U	
²³⁹ Pu	²⁴¹ Am	²⁴³ Am	²³⁴ U	
⁹⁹ Tc	²³⁸ Pu	²³⁴ U	²⁴² Pu	
²⁴¹ Pu	²³⁴ U	²²⁶ Ra	²⁴³ Am	

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COMPARISON OF RADIONUCLIDE IMPORTANCE CRITERIA-SPENT FUEL

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	INVENTORY	INVENTORY EPA LIMIT	RELEASE RATE EPA LIMIT		
1000 yr	²⁴¹ Am	²⁴¹ Am	¹⁴ C		
	²⁴⁰ Pu	²⁴⁰ Pu	²³⁷ Np		
	²³⁹ Pu	²³⁹ Pu	²⁴⁰ Pu		
	²⁴³ Am	²⁴³ Am	²³⁹ Pu		
10000 yr	²³⁹ Pu	²³⁹ Pu	²³⁷ Np		
	²⁴⁰ Pu	²⁴⁰ Pu	¹⁴ C		
	⁹⁹ Tc	²⁴³ Am	²³⁹ Pu		
	²⁴³ Am	²³⁴ U	²⁴⁰ Pu		
100000 yr	²³⁹ Pu	²³⁹ Pu	²³⁷ Np		
	⁹⁹ Tc	²³⁴ U	²²⁶ Ra		
	⁹³ Zr	²⁴² Pu	²³⁹ Pu		
	⁵⁹ Ni	²³⁷ Np	⁹⁹ Tc		
[*] Diffusion-limit	Diffusion-limited model				



COLLOIDAL ²³⁹PuO₂

SOLUBILITY DISSOLUTION RADIOLYTIC EFFECTS

PRELIMINARY RESULTS, RESEARCH IN PROGRESS.

D. E. HOBART, T. W. NEWTON, P. D. PALMER



REACTIONS

SOLUBILITY IN ACID

 $PuO_2 \cdot nH_2O + 4H^+ = Pu^{4+} + (2+n)H_2O$ (1)

DISSOLUTION WITH OXIDATION $PuO_2 \cdot nH_2O + Ox = PuO_2^+ + Rd + nH_2O$ (2)

SPECIAL CASE $PuO_2 \cdot nH_2O + PuO_2^{2+} = 2PuO_2^{+} + nH_2O$ (3) $Q_3 = (PuO_2^{+})^2 / (PuO_2^{2+})$



NET RADIATION EFFECTS

 $PuO_2^{2^+} + \frac{1}{2}H_2O \stackrel{a}{=} PuO_2^+ + H^+ + \frac{1}{4}O_2$ 1.2 - 1.5% PER DAY FOR ²³⁹Pu, pH ~ 3

PuO₂⁺ ^a NO REDOX OBSERVED, pH 2

$$2H^{+} + PuO_2 \cdot nH_2O \stackrel{\text{\tiny def}}{=} PuO_2^{2^{+}} + H_2 + nH_2O$$

pH 2 - 3

$$Pu^{3^{+}} + H^{+} \stackrel{a}{=} Pu^{4^{+}} + \frac{1}{2}H_{2}$$

 $Pu^{4^{+}} + \frac{1}{2}H_{2}O \stackrel{a}{=} Pu^{3^{+}} + H^{+} + \frac{1}{4}O_{2}$ (IN ACID SOLUTIONS)

(REACTIVE SPECIES: H₂O₂, H, OH, HO₂)

ANALYTICAL METHODS

 $Pu(VI), (Pu0_2^{2^+})$

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SPECTROPHOTOMETRY IN 0.6 M HC10₄ AT 830.3 nm, $\varepsilon \sim 525 \text{ M}^{-1} \text{cm}^{-1}$.

 $Pu(V), (PuO_2^+)$

SPECTROPHOTOMETRY AFTER ADDING Ce(IV) TO CONVERT Pu(V) TO Pu(VI). Ce(IV) OXIDIZES Pu(V) RAPIDLY BUT Pu(IV) OR COLLOIDAL Pu(IV) VERY SLOWLY.



EXPERIMENTS

- 1. PuO₂⁺ ALLOWED TO DISPROPORTIONATE, BOTH ²³⁹Pu AND ²⁴²Pu USED.
- 2. PURIFIED COLLOIDAL SUSPENSIONS ALLOWED TO "DISSOLVE".
- 3. MIXTURES OF COLLOIDAL Pu(IV) WITH PuO_2^{2+} AND/OR PuO_2^{+} ANALYZED AS A FUNCTION OF TIME.



PRELIMINARY RESULTS

1. NET OXIDATION AND NET REDUCTION WERE BOTH OBSERVED.

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PROJECT

- 2. Pu(V) UNDERGOES BOTH DISPROPORTIONATION AND REDUCTION.
- 3. COLLOIDAL Pu(IV) DISSOLVES BY OXIDATION, PRIMARILY TO Pu(VI).
- 4. AT CONSTANT INITIAL Pu(VI)/Pu(V) RATIOS, THE BEHAVIOR DEPENDS ON THE CONCENTRATION OF COLLOID.
- 5. THERE IS NO DIRECT EVIDENCE FOR REPROPORTIONATION OF Pu(V).



SOLUBILITY

$$PuO_{2(coll)} + PuO_{2}^{2^{+}} = 2PuO_{2}^{+}$$
 (3')

$$4H^{+} + 2PuO_{2}^{+} = Pu^{4^{+}} + PuO_{2}^{2^{+}} + 2H_{2}O$$
 (4)

 $PuO_{2(coll)} + 4H^{+} = Pu^{4^{+}} + 2H_{2}O$ (1)

UPPER LIMIT ON Q₃ FROM DISPROPORTIONATION RESULTS. OBSERVED LOG Q₃ VALUES: -3.05, -2.66 (²⁴²Pu) -2.83 (²³⁹Pu)

LOWER LIMIT ON Q3 IF. REPROPORTIONATION ACTUALLY OCCURRED.

OBSERVED LOG Q₃ VALUES: -3.13, -3.71, -3.83(LOG Q₃ = -3.0 CORRESPONDS TO LOG Q₁ = -1.1 OR LOG K_{SP} = -57.1)



TOTAL Pu

TOTAL Pu

 $Pu(VI) \stackrel{a}{=} Pu(V)$

Pu(IV)con = Pu(VI)

LAWRENCE BERKELEY LABORATORY

H. NITSCHE AND N. EDELSTEIN

SOLUBILITY OF SELECTED ACTINIDE IONS IN 0.1 M Nacio4 AT

PH 7 AND IN GROUNDWATER (J-13) OF THE

NEVADA STORAGE SITE

- experiment still in progress

COMPARISON OF EXPERIMENTAL CONDITIONS IN NaClo₄ Solution AND IN J-13 GROUNDWATER

- a) NaClO_{μ} solution
 - pH 7
 - 0.1 <u>M</u> ionic strength
 - non-complexing
 - actinide solubility only controlled by hydroxide formation
 - will give "simple" understanding of solubility (baseline)
 - b) J-13 groundwater
 - pH 7
 - 0.1 <u>M</u> ionic strength
 - complexing $(CO_2, SO_4, NO_3, Cl, F, PO_4)$
 - actinide solubility controlled by complexing ability of ligands
 - will give "complex overall picture"
 - results can be compared with values gained by modeling

GOAL OF EXPERIMENT

- determine actinide solubility
- identify oxidation state in solution
- characterize solid phase
- DETERMINE PERTURBATION OF SOLUBILITY BY LIGANDS

EXPERIMENTAL CONDITIONS

- actinide ions have specific oxidation state
- oxidation state was
 - adjusted by electrochemistry
 - verified by absorption spectrophotometry
- background electrolyte 0.1 <u>M</u> NaCl0₄
- pH 7.0 \pm 0.1, controlled by pH-stat, 25 \pm 1°C
- inert atmosphere
- samples were shaken continuously
- solution filtered through 2 nm filters

actinide ions used in experiment:

 $\begin{array}{rcrc} & 237_{\text{Np0}2}^{+} \\ & 237_{\text{Np0}2}^{2+} \\ & 243_{\text{Am}}^{3+} \\ & 242_{\text{Pu}}^{4+} \\ & 242_{\text{Pu0}2}^{4+} \\ & 242_{\text{Pu0}2}^{2+} \end{array}$



Initial NpO₂+

- c = 10^{-3.36±0.01} M

- Kraus and Nelson (48): amorphous NpO₂(OH), $c \le 10^{-2.3} M$ (pH 7)
- Np in solution is $Np0_2^+$ (V)



 $\frac{\text{Initial NpO_2}^{2+}}{- c = 10^{-3.46} \pm 0.02} \text{ M}$ - Kraus and Nelson (48):

amorphous NpO₂(OH)₂, c ≤ 10^{-7.4} M (pH 7)

- Np in solution is $Np0_2^+$ (V)

<u>Conclusion</u>

- solubility value for initially Np0 $_2^+$ and Np0 $_2^{2+}$ is close to each other
- oxidation state in solution is V for both cases



Initial Am³⁺

$$-$$
 c = 10^{-3.53} ± 0.02 _M

- Rai, et al. (82) amorphous ${}^{243}\text{Am(OH)}_3$, $c \approx 10^{3.5}$ M (I = 0.02 M)
- Silva (82) crystalline 243 Am(OH)₃, c = $10^{-4.5}$ <u>M</u> (I = 0.1 <u>M</u>)
- Am in solution is Am^{3+}

<u>Conclusion</u>

- Value agrees well with Rai's value for amorphous ²⁴³Am(OH)₃
- solid phase probably amorphous Am(OH)
- oxidation state in solution did not change



Initial Pu⁴⁺ - $c = 10^{-7.56 \pm 0.22} M$ - Rai and Ryan (82): Pu0₂ crystalline, $c = 10^{-8} M (pH 7)$ Pu0₂ x H₂0 amorphous, $c = 10^{-6.5} M (pH 7)$ Initial Pu0⁺₂ - $c = 10^{-8.51 \pm 0.14} M$ - Baes and Mesmer: amorphous Pu0₂(OH), $c = 10^{-1.6} M (pH 7)$

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Initial Pu02²⁺

$$-$$
 c = 10^{-6.93} ± 0.13 M

Baes and Mesmer: amorphous NpO₂(OH)₂, $c \leq 10^{-7.4} M$ (pH 7)

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Problem

Low Pu concentration in solution requires method other than absorption spectrophotometry for oxidation state determination.

	TTA, pH O	TTA, pH 4		LaF3		HEXONE	
IV	111,V,VI,P	III,IV	V,VI,P	III,IV,P	V,VI	IV,VI	III,V,P

D.,4+	0.8±0.2	99.2±0.2	2.0±1.3	98.0±1.3	2.6±1.6	96.8±0.6	29.9±7.4	70.1±7.4
	(95.14	±8,2)	(91.84	±11.0)	(86.8	±7.1)	(101,43	46.6)

200+	2.2±2.0 97.3±1.3	1.6±0.6 98.4±0.6	5.2±4.3 94.9±4.3	27.0±15.5 72.9±15.2	
<u>2</u>	(101.7±5.7) (96.0±1.6)		(97.3±1.1)	(95.8±2.9)	

- 24	1.4±0.8	98.7±0.8	0.6±0.1	99.4±0.1	2.9±1.6	97.1±1.6	15.8±15.8	84.2±15.8
Pu02'	(99.7±0.2) (98.4±0.5)		(93.5±0.4)		(85.4±5.2)			

All values in %

Values in parentheses are overall chemical yield

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CONCLUSION

Pu⁴⁺

- solubility value is within literature values for crystalline PuO_2 and amorphous PuO_2 x H_2O
- oxidation states in solution are

III + IV + V ≤ 5%

- V + VI ~97%
- V ~ 60%-75%

$Pu0^+_2$

- solubility value is much smaller than literature value for amorphous PuO₂(OH)

- oxidation states in solution are

III + IV + V $\leq 10.0\%$

V + VI ~ 90%-98%

V ~50%-70%

Pu022+

- solubility value agrees with literature value for amorphous $\text{NpO}_2(\text{OH})_2$ and $\text{UO}_2(\text{OH})_2$
- oxidation states in solution are

 $III + IV + V \leq 5\%$ $V + VI \qquad \sim 97\%$

V ~ 65%-80%

Work in Progress:

- a) Study of solids by X-rcy powder diffraction
- b) repetition of experiments with J-13 groundwater

SUMMARY

- a) Gave solubility values at pH 7 in 0.1 <u>M</u> non-complexing solution of initially
 - $Np0_2^+$
 - Np02²⁺
 - Am³⁺
 - Pu⁴⁺
 - Pu
 - Pu0⁺
 - $Pu0_2^{-2+}$
- b) Studied oxidation state distribution in solution phase by either absorption spectrophotometry or extraction/coprecipitation method

REDOX CONDITIONS AND REDOX BUFFERING


 FE²⁺ - TI OXIDES
 0.16 - 0.33 v/o

 DENSITY OF TUFF
 2g/cm³

 EQUIVALENTS FE²⁺
 20-40/m³

REPOSITORY LOADING

WASTE ELEMENT	EQUIVALENTS	
Np	4.19 x 10 ⁵	
Pu	4.36 x 10 ⁶	
Тс	1.63×10^{6}	
U	5.64 x 10 ⁸	
TOTAL	5.70 x 10 ⁸	

3-6m OF TUFF THE SIZE OF THE REPOSITORY CONTAINS THIS AMOUNT OF FE²⁺.

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MINERALS

- A) OXIDATION EXSOLUTION IN Fe, TI-OXIDE MINERALS: WIDELY
 - SCATTERED FERROUS REMNANTS
- **B) FERROUS IRON IN SULFIDE BEARING ROCKS AT DEPTH**

• ROCKS

RETENTION OF SOME FERROUS IRON IN ROCKS OF THE LOWER

TOPOPAH SPRING VITROPHYRE



WEIGHT - PERCENT FeO IN SAMPLES FROM USW G-4 (DETECTION LIMIT = 0.2% FeO). DEVITRIFIED TOPOPAH 0 % 1089 FT 1190 0 % FΤ ZEOLITE INTERVAL I 0 % 1314 FT LOWER TOPOPAH VIŤROPHYRE 1330 FT 0.3 % ZEOLITE INTERVAL II 1438 0 % FT 1544 FT 0 % CENTRAL PROW PASS MEMBER 0 % 1871 FT ZEOLITE INTERVAL III 2100 FT 0 % CENTRAL BULLFROG MEMBER 2516 0 % F'T . ZEOLITE INTERVAL IV 0 % 2716 FΤ 2823 0 % FT



REDOX BUFFERING 0.16-0.33% FE-T: OXIDES 2g/cm³ TUFF DENSITY OF 20-40/m³ EQUIVALENTS FE2+

REPOSITORY	LOADING
WASTE ELEMENT	EQUIVALEN
Ne	4.19×10
P	4.36×10
ja (1.63 × 10
le	5.64 ×108
U	570×108
TOTAL	0.10710
	· · · · · · · · · · · · · · · · · · ·

0 1	0E	TUFF	THE	SIZE
3-6m	Q E DOS	itory	CONTAI	NS.
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TRANSPORT MODELING AND MATRIX DIFFUSION

Fracture Flow Experiment

Diffusion Experiment



- How do fractures affect flow through unsaturated tuff?
- How well can the fractured and unfractured layers retard radionuclide transport?
- What is the effect of waste heat load on hydrology?



- characterization: (inverse problem) what must flow properties be to produce given observations?
- experimental design: special case of predictive mode





INGREDIENTS FOR BUILDING MODELS

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- conservation equations (mass, momentum, energy)
- equations of state
- conceptual model of microphysics which can be integrated to provide macroscopic constitutive relations,
 - e.g., $\kappa(\sigma, \varepsilon)$: permeability as fcn. of saturation and porosity

 ϵ (P,T,C): porosity as fcn. of pressure, temperature, concentrations.

Limitations of models due mainly to shortcomings of idealized microphysics and/or to inadequate equations of state.



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SIMPLE ANALYSIS OF TRANSPORT IN FRACTURED ROCK



- 1. One dimensional, steady state, two dissolved species.
- 2. Fixed concentration boundary conditions at z = 0.
- 3. Radioactive decay. x---c and y---d
- 4. Reversible adsorption of dissolved species on both the matrix and the colloids.
- 5. No sources except at z = 0.
- 6. Constant porous matrix properties, porosity, permeability, density, and chemical compositon.



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WATER FLOW IN UNSATURATED FRACTURED ROCK







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Coupling of RADIOCOLLOID Equations with Geo-Transport Equations

TRACR3D ENHANCEMENT



Naturally occurring (pseudo) colloids travel with the flowing groundwater into the repository where they contact and absorb a protion of the dissolved nuclides.





Decay of Radionuclides: x --> c y --> d

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LANL——Colliod Size Distributions and Concentrations PNL——Adsorption of Actinides on Iron Sillicates and Clay Georgia Tech——Column and Adsorption Studies using Kaolinite Sweden——Formation of Actinide Colloids Column Study Canada——Soil Column Studies with Plutonium Italy——Leached Borosillicate Glass followed by Column Tests Germany——Adsorption of Actinides on Natural Colloids

Field Sites

LANL--DP Waste Site

Canada—Low Level Waste Site (early release of cesium, etc.)



- 1. Treatment of both pseudo and true colloids. The code initially will consider only one chemical type of pseudo and one type of true colloid. This reduces the complexity and storage requirements of the code. If necessary, these restrictions could later be relaxed.
- 2. The size distribution of the pseudocolloids remains fixed. The true colloids can exhibit both birth and growth behavior. Colloids are assumed to be in stable dilute concentrations; thus no coalescence or agglomeration is considered.
- 3. Colloids can adsorb/desorb two species (nuclides x and y). (The population balance coupled to the mass transport equations for species x and y). One or both of the dissolved species may form true colloids.
- 4. Irreversible colloid capture on the matrix.

MODEL ASSUMPTIONS

COLLOIDS--POPULATION BALANCE

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5. Pseudocolloids neither grow nor dissolve. The amount of radionuclides adsorbed is assumed to be only a monolayer; thus, not effecting the size nor mass of the initial pseudo-colloid size distribution.

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PROFE

- 6. The colloid capture rate expression may be a function of particle size.
- 7. Initially the natural colloids are nonradioactive and of a given size distribution.
- 8. The concentration of pseudocolloids is changed only by capture. No additional sources or sinks.
- 9. One dimensional problem, constant fluid/colloid velocity, constant boundary colloid concentration and size distribution. (at z=0)
- 10. Radioactive decay of adsorbed species. No chain reactions.



GENERATION OF TWO PHASE FLOW BY WASTE HEAT LOAD











N W S I PROJECT

Temperature Contours 100 years After Emplacement of Heat Load

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-180 -30 40 50 -200 60 70 -220 80 X 90 H -240-100 -280 -300 10 20 30 0 RADIUS (M)










DIFFUSION TEST RATIONALE

- 10CFR60.112: EPA RELEASE STANDARDS
- NNWSI INFORMATION NEED 1.3.3: BOUNDS ON EFFECTIVE RETARDATION IN DISPERSION/DIFFUSION/ADVECTIVE TRANSPORT PROCESS





- DIFFUSIVITY OF AQUEOUS SOLUTES INTO WATER-FILLED
 PORES IN
 - TOPOPAH SPRING TUFF
 - CALICO HILLS TUFF
- MEASURE IN LABORATORY
- EXTRAPOLATE TO FIELD
 - LITHOSTATIC LOAD
 - SATURATION
- CHECK EXTRAPOLATION
- USE FIELD RESULTS FOR TRANSPORT CODE CALCULA-TIONS



- EXPLORATORY SHAFT
 - MAIN TEST FACILITY TOPOPAH SPRING TUFF
 - DRILL ROOM TUFFACEOUS BEDS OF CALICO HILLS
- NEVADA TEST SITE
 - PACKER TESTING
 - OVERCORING TECHNIQUE DEVELOPMENT
- LOS ALAMOS
 - TRACER SELECTION
 - LABORATORY-SCALE EXPERIMENTS
 - QUANTITATIVE ANALYSES OF OVERCORED MATERIAL
 - COMPUTATIONS







DIFFUSION TEST TECHNIQUE

- SELECT TRACERS
- VERIFY PACKER PERFORMANCE
- DEVELOP OVERCORE TECHNIQUE
- LABORATORY DIFFUSIVITY MEASUREMENT
- DRILL IN SITU HOLES ~5 cm DIAMETER, 10 m DEEP
- INTRODUCE TRACERS
- SEAL FOR 3 MONTHS TO 1 YEAR
- OVERCORE
- ANALYZE FOR SOLUTE CONCENTRATION WITH DISTANCE FROM EMPLACEMENT
- DERIVE DIFFUSIVITY, D



DIFFUSION TEST EXPECTED VALUE

- $D \sim 10^{-7} \text{ cm}^2/\text{s}$
- DIFFUSION DISTANCE \sim 7 cm/YEAR



- MEASURE D IN EXPLORATORY SHAFT IN
 - MAIN UNDERGROUND FACILITY (TOPOPAH SPRING)
 - DRILL ROOM AT BOTTOM (CALICO HILLS)
- COMPARE LABORATORY AND FIELD RESULTS
- SEARCH FOR INHOMOGENEOUS DIFFUSION TO MEA-SURE MATRIX FLOW
- USE FIELD D VALUES FOR TRANSPORT CODE CALCULA-TIONS OF RADIONUCLIDE RETARDATION



PROPOSED ES FRACTURE FLOW EXPERIMENTS PURPOSES OF PROPOSED EXPERIMENTS

- 1. Confirmation of codes
- 2. Determination of parameters in codes



PROPOSED ES FRACTURE FLOW EXPERIMENTS LOCATIONS OF PROPOSED EXPERIMENTS

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- 1. Welded portion of Topopah Springs member
- 2. Calico Hills tuff
- 3. Transition region between welded and nonwelded portions of Topopah Springs member



- 1. Predict transport of water and solutes in test location using numerical codes
- 2. Drill source hole
- 3. Inject water containing suite of solutes and some particulates —— injected flows greater than natural flows, so measurable transport can occur
- 4. Drill sampling holes where calculations predict flow
- 5. Analyze water from sampling holes and compare with predictions













SORPTION



UPDATE ON SORPTION AND PRECIPITATION

Kim Thomas INC-11, Los Alamos NRC Geochemistry Workshop July 10-13, 1984



SORPTION RATIOS FOR THORIUM IN J-13 GROUNDWATER

Los Alamos

SORPTION RATIOS (m1/g)°

CORE	TRACED FEED pH	TRACED FEED CONCENTRATION (M)	EXPERIM	ental value (ph)	AVERAGE
G1-1292	5.61	5.39 x 10 ⁻⁸	566 (7.52)	389 (7.52)	478(63)
G1-2233	5.75	1.07 x 10 ⁻⁷	266 (7.00)	421 (6.93)	344(55)
G1-2289	5.76	5.86 x 10 ⁻⁸	140 (6.89)	146 (6.86)	143(2)
G1-2363	5.60	6.52 x 10 ⁻⁸	940 (6.83)	1484 (6.82)	1213(193)
SORPTION TIM	E WAS 12 DAY	S			

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N N W S

PROJECT

ANALYSIS OF WELL J-13 WATER PRE-TREATED WITH YUCCA MOUNTAIN TUFF

Los Alamos

3.08

6.18

0.066

0.031

Mn Mg SI Fe Sr Ba SAMPLE NO. (ppm) (ppm) (ppm) (ppm) (ppm) (ppm) ORIGINAL GROUNDWATER 1.76 0.012 31.8 0.011 0.039 0.001 1.660 q1-2840 0.009 33.2 0.007 0.037 0.002 q4-1502 0.016 0.001 23.0 0.016 0.001 0.001 gu3-1301 1.040 0.026 25.7 0.050 0.073 0.003 Y Π U Ca K Ał SAMPLE NO. (ppm) (ppm) (ppm) (ppm) (ppm) (ppm) ORIGINAL GROUNDWATER 0.021 0.028 11.5 0.060 5.26 0.025 a1-2840 0.010 0.001 11.90 0.057 5.07 0.010

ORIGINAL GROUNDWATER0.0210.02811.50.060g1-28400.0100.00111.900.057g4-15020.0000.0070.240.050gu3-13010.0170.01811.100.078

ANALYSIS OF WELL J-13 WATER PRE-TREATED WITH YUCCA MOUNTAIN TUFF (cont)

	Na	F	CI	NO ₂	NO ₃	SO4		
SAMPLE NO.	(ppm)	(ppm)	(ppm)	(ppm)	(ppm)	(ppm)	рH	Alk
ORIGINAL GROUNDWATER	45.1	2.1	6.4	-	10.1	18.1	6.9	2.339
g1–2840	48	2.5	7.7	0.00	8.8	18.6	7.56	1.922
g4-1502	51	2.6	8.4	0.00	7.1	19.8	8.09	2.082
gu3-1301	46	2.7	8.8	0.20	8.1	19.0	8.52	2.001

N N W S PROJECT

ANALYSIS OF WELL UE25p#1 WATER PRE-TREATED WITH YUCCA MOUNTAIN TUFF

Los Alamos

Mg	Mn	SI	Fe	Sr	Ba
(ppm)	(ppm)	(ppm)	(ppm)	(ppm)	(ppm)
31.9	0.103	17.4	0.045	0.390	0.260
14.5	0.045	15.6	0.030	0.009	0.006
26.7	0.015	13.7	0.021	0.009	0.059
v	TI	Ca	LI	ĸ	AI
V	TI	Ca	LI .	K	AI
(ppm)	(ppm)	(ppm)	(ppm)	(ppm)	(ppm)
0.023	0.039	87.8	0.720	13.4	0.139
0.011	0.010	15.8	0.557	8.00	0.063
0.021	0.022	19.5	0.557	9.51	0.064
	Mg (ppm) 31.9 14.5 26.7 V (ppm) 0.023 0.011 0.021	Mg Mn (ppm) (ppm) 31.9 0.103 14.5 0.045 26.7 0.015 V Ti (ppm) (ppm) 0.023 0.039 0.011 0.010 0.021 0.022	MgMnSi (ppm) (ppm) (ppm) 31.9 0.103 17.4 14.5 0.045 15.6 26.7 0.015 13.7 VTiCa (ppm) (ppm) 0.023 0.039 87.8 0.011 0.010 15.8 0.021 0.022 19.5	MgMnSiFe (ppm) (ppm) (ppm) (ppm) 31.9 0.103 17.4 0.045 14.5 0.045 15.6 0.030 26.7 0.015 13.7 0.021 VTiCaLi (ppm) (ppm) (ppm) 0.023 0.039 87.8 0.720 0.011 0.010 15.8 0.557 0.021 0.022 19.5 0.557	MgMnSiFeSr (ppm) (ppm) (ppm) (ppm) (ppm) (ppm) $\overline{31.9}$ $\overline{0.103}$ $\overline{17.4}$ $\overline{0.045}$ $\overline{0.390}$ 14.5 0.045 15.6 0.030 0.009 26.7 0.015 13.7 0.021 0.009 26.7 0.015 13.7 0.021 0.009 0.023 0.039 87.8 0.720 13.4 0.011 0.010 15.8 0.557 8.00 0.021 0.022 19.5 0.557 9.51

ANALYSIS OF WELL UE25p#1 WATER PRE-TREATED WITH YUCCA MOUNTAIN TUFF (cont)

	Na	F	Cl	NO ₂	NO ₃	Br
SAMPLE NO.	(ppm)	(ppm)	(ppm)	(ppm)	(ppm)	(ppm)
ORIGINAL GROUNDWATER	110.	3.5	37.		<0.1	
G1-2233-2	192.	4.2	34.8	-	-	-
GU-3-1301-2	106.	3.9	34.2	-	-	-
	PO4	SO4		Aik		
SAMPLE NO.	(ppm)	(ppm)	pН	meq/l	•	
ORIGINAL GROUNDWATER		129.	6.7	11.448		
G1-2233-2	-	195.	8.97	10.848		
GU-3-1301-2	-	189.	8.91	6.771		

COMMENT			
	$HCO_3 = 698$		
$CO_3 = 22.8;$	$HCO_3 = 569$		
$CO_3 = 14.6;$	$HCO_3 = 354$		

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SORPTION RATIOS IN UE-25p#1, J-13, AND DEIONIZED WATERS

			SORPTION RAT	10S
	CORE	ELEMENT	UE-25p#1	J-13
G	U3-1301	Sr	10(2)	32(8)
		Cs	45(5)	160(35)
		Ba	82(18)	570(60)
		Eu	>17 000	75(12)
			SORPTION RATIOS (ml/g)	
CORE	ELEMENT	UE-25p#1	J-13	DEIONIZED
G1-2233	Sr	2 000(330)	48 000(3 000)	>56 000
	Cs	7 500(1 100)	13 500(800)	13 000(1 600)
	Ba	41 000(6 300)	250 000(30 000)	55 000(5 300)
	Eu	>5 600	900(200)	810(100)

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COMPARISON OF TIN SORPTION RATIOS IN J-13 AND UE-25p#1° GROUNDWATERS Los Alamos

CORE	SORPTIC (ml)N.RATIOS /g)	DESORPT	ION RATIOS
	J-13	UE-25#1	J-13	UE-25#1
G1-2840	283	20000	780	18400
GU3-1301	168	4000	1280	6750
G4-1502	215	800	500	300
G1-2901	22000	35800	38000	52500

"WATER FROM DEPTH 1298-1792m.



LONG-TERM TC AND Np SORPTION RATIOS

Los Alamos

CORE	ELEMENT	pH RANGE	TRACER CONC. (M)	CONTACT TIME: UP TO 9 MONTH
G41502	Tc	8.71-9.00	8×10-10	0-0.02
	Np	8.32-8.70	1×10 ⁻¹⁰	4.0-5.1
				·
GU3-1301	Тс	8.41-8.72	7×10-10	0-0.04
	Np	8.46-8.67	6×10 ⁻¹¹	1.7-2.2
GU3-916	Тс	8.51-8.83	6×10 ⁻¹⁰	0.13-0.81
	Np	8.08-8.76	6×10 ⁻¹¹	4.8-5.4
]

SORPTION RATIO (ml/a)

Am AND Pu	SHORT-TERM SC	DRPTION RATIOS
·	ON TUFF G4-15	02

N N W S I PROJECT

SORPTION TIME	SORPTION RATIO (ml/g)			
	Am	Pu		
1 h	490 (20)	19 (1)		
4 h	715 (5)	22 (1.2)		
1 d	1100 (60)	27 (1)		
2 d	1200 (5)	34 (2.3)		
3 d	1450 (50)	34 (2.9)		
1 wk	1900 (20)	36 (1)		
3 wk	1550 (10)	34 (1)		
6 wk	1400 (5)	43 (1)		

COMPARISON OF SORPTION RATIOS^{\circ} (R_d) MEASURED UNDER ATMOSPHERIC AND N₂-CONTROLLED-ATMOSPHERE CONDITIONS^b Los Alamos

ELEMENT	EFFECT ON Rd VALUE
Cs	NONE
Sr	NONE
Ba	NONE
Се	NONE
Eu	NONE
Na	NONE
Se	NONE
Mn	LOWER IN CA (FACTORS OF 3 TO 4)
Sn	LOWER IN CA (FACTOR OF 2)
Am	NONE
Pu	HIGHER IN CA (FACTOR OF 2)
Тс	HIGHER IN CA (FACTOR OF <u>></u> 10)
U	HIGHER IN CA° (FACTORS OF 2 TO 3)
Np	HIGHER IN CA (FACTOR OF 2)

"TUFF STUDIED: YM-22, YM-38, YM-54

N N

WS

^bCA = CONTROLLED ATMOSPHERE: NITROGEN, \leq 0.2 ppm OXYGEN, AND \leq 20 ppm CARBON DIOXIDE.

°YM-38 (ZEOLITIZED) TUFF ONLY; OTHERWISE NO EFFECT.



COMPARISON OF Np SORPTION RATIOS UNDER AMBIENT AND CO2-CONTROLLED ATMOSPHERES Los Alamos

CORE	ATMOSPHERE	SORPTION RATIO (ml/g)
GU3-1203	AMBIENT CO2	0.49 0.35
GU3—1301	AMBIENT CO2	1.7 2.1
G41608	AMBIENT CO2	6.3 5.4

N N N W S I PROJECT	Np ISOTHERM SORPTION RATIOS UNDER CO ₂ - CONTROLLED ATMOSPHERIC CONDITIONS Loe Alamos			
	Np CONC (M)	SORPTION RATIO (ml/g)		
	6 x 10 ⁻⁴	2.4		
	T 40- A	0.04		

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6×10^{-4}	2.4
3×10^{-4}	0.94
7×10^{-5}	3.4
3×10^{-5}	3.3
8 x 10 ⁻⁶	4.9
3×10^{-6}	5.0
7×10^{-7}	5.0
4×10^{-7}	5.2
2×10^{-11}	5.4



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COMPARISON OF AVERAGE SORPTION RATIOS FROM BATCH AND CIRCULATING-SYSTEM SORPTION MEASUREMENTS

BATCH R-TO-CIRCULATING SYSTEM R RATIO

	Sr	Çs	Ba	Am	Pu
ZEOLITIZED TUFF	S			· · · · ·	<u></u>
JA-37 YM-49	0.74(0.04)	0.34(0.03)	0.88(0.18)	8.2(3.3) 2.0(0.7)	1.4(0.3) 0.40(0.15)
DEVITRIFIED TUFF	<u>s</u>			· · ·	
YM-22 YM-54 G1-1883	2.0(0.2) 1.3(0.2)	0.59(0.09) 1.3(0.3)	7.5(0.7) 1.6(0.6)	1.4(0.1)	1.4(0.4)

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SUMMARY OF DIFFERENCES IN SORPTION RATIOS BETWEEN CRUSHED-ROCK COLUMN AND BATCH MEASUREMENTS Los Alamos

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N W S

CORE	⁸⁵ Sr	¹³⁷ Cs	¹³³ Ba
YM-54	0	0	0
YM-38	N.A.	C (1.45)	N.A.
YM-22	0	0`´	0
JA-37	B (2.08)	0	N.A.
JA-32	0`´	N.A.	N.A.
GI-3658	B (11.8)	B (1.18)	B (1.91)
GI-3116		_ ` '	0
GI-2363	0	—	-
GI-2233	0	C	0
GI-1982	0	0	B (1.18)
GI—1883	0	Ō	B (1.23)
GI-1292	0	C	~~~
N /			
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Ŵ	Rd COMPARISON: TUFF WAFERS VS CRUSHED	TUFF	
S VI		Los Alamos	

TUFF SAMPLE	ELEMENT	R_{d} (WAFER) (ml/g)	R_d (BAICH) (ml/g)
G-1-1883	Sr	27	22
	Cs	230	190
	Ba	210	180
G-1-1982	Sr	80	62
	Cs	1000	1200
	Ba	710	800
G-1-1436 (DESORPTION)	Sr	96500	87000
	Cs	14900	24000

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BACKUP VIEWGRAPHS AND DATA

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LONG-TERM TECHNETIUM SORPTION RATIOS

	`	Traced Feed		Sorption Ratios (m²/g)		
Core	Tracer Feed pH	Concentration (M)	Sorption <u>Time</u>	Experimen (p	tal Values H)	Average
G-4-1502	8.74	8×10^{-10}	6 weeks	0.042 (8.60)	-0.003 (8.60)	0.02(0.02)
			13 weeks	0.06 (9.00)	-0.038 (8.81)	0.01(0.05)
			6 months	-0.085 (8.65)	-0.09 (8.65)	0
			9 months	-0.06 (8.71)	-0.05 (8.75)	0
GU-3-1301	8.67	7×10^{-10}	6 weeks	0.058 (8.41)	0.022 (8.47)	0.04(0.02)
			13 weeks	0.065 (8.63)	0.002 (8.69)	0.03(0.03)
			6 months	-0.003 (8.60)	0.004 (8.61)	0
			9 months	0.013 (8.70)	0.031 (8.72)	0.02(0.01)
GU-3-916	8.57	6×10^{-10}	6 weeks	0.50 (8.51)	0.94 (8.57)	0.72(.2)
			13 weeks	0.33 (8.75)	1.3 (8.70)	0.81(0.5)
			6 months	0.81 (8.61)	0.44 (8.66)	0.62(0.2)
			9 months	0.14 (8.81)	0.11 (8.83)	0.13(0.02)

		Encod Food		S	orption Rat:	Los
Core	Tracer Feed	Concentration (M)	Sorption Time	Experimen (p	tal Values H)	Average
G-4-1502	8.72	1×10^{-10}	6 weeks	4.6 (8.33)	3.5 (8.32)	4.0(1)
			3 months	5.1 (8.66)	3.4 (8.70)	4.3(1)
			6 months	4.7 (8.44)	5.0 (8.49)	4.8(1)
			9 months	5.2 (8.37)	5.0 (8.53)	5.1(1)
GU-3-1301	8.4	2×10^{-11}	6 weeks	1.8 (8.53)	1.7 (8.54)	1.7(1)
	8.6	6×10^{-11}	3 months	2.3 (8.63)	2.1 (8.65)	2.2(1)
			3 months	2.1 (8.65)	1.9 (8.67)	2.0(1)
			6 months	2.2 (8.55)	2.3 (8.62)	2.2(1)
		• .	9 months	1.9 (8.46)	1.8 (8.47)	1.9(1)
GU-3-916	Not Available	e 6 x 10 ⁻¹¹	6 weeks	4.8 (8.76)	4.9 (8.65)	4.8(1)
			3 months	5.1 (8.69)	5.4 (8.68)	5.3(1)
			6 months	5.6 (8.48)	5.2 (8.54)	5.4(1)
			9 months	5.1 (8.24)	5.1 (8.08)	5.1(1)

LONG-TERM NEPTUNIUM SORPTION RATIOS

Traced Feed	Traced	Sorption Ratios (mL/g)			
Concentration (M)	Feed pH	Experimen (p	tal Values oH)	Average Value ^b	
6×10^{-4}	7.0	2.0 (6.4)	2.7 (6.3)	2.4(1)	
3×10^{-4}	7.2	1.3 (6.4)	0.57 (6.5)	0.94(1)	
7 x 10 ⁻⁵	7.3	3.5 (6.7)	3.2 (6.7)	3.4(1)	
3×10^{-5}	7.2	3.4 (6.7)	3.1 (6.7)	3.3(1)	
8 x 10 ⁻⁶	7.3	4.7 (6.7)	5.1 (6.7)	4.9(1)	
3×10^{-6}	7.3	4.5 (6.7)	5.4 (6.7)	5.0(1)	
7×10^{-7}	7.3	5.2 (6.7)	4.7 (6.7)	5.0(1)	
4×10^{-7}	7.3	4.4 (6.7)	5.9 (6.7)	5.2(1)	
2×10^{-11}	7.2	5.2 (6.7)	5.6 (6.8)	5.4(1)	

NEPTUNIUM ISOTHERM SORPTION RATIOS UNDER CO₂-CONTROLLED ATMOSPHERIC CONDITIONS[®]

^aCore was G-4-1608, 75-500 μm . Atmosphere was enriched in CO₂ to ~5%. Sorption was for 6 weeks.

b_{Numbers} in parentheses are standard deviations of the mean.

	Amb			Pu ^C							
Conntion	pH V	alues	R, (m	e/g)	pH V	alues	R, (1	w/g)			
<u>Time</u>	Feed	<u>Sample</u>	Individual	Average	Feed	Sample	Individual	Average			
l hour		8.3	470	(00(00))	• •	8.4	19	10(1)			
	8.5	8.3	510	490(20)	8.4	8.2	20	19(1)			
4 hour		8.4	720		~ /	8.2	24				
	8.5	8.4	710	715(5)	8.4	8.1	21	22(1.2)			
l day		8.3	1000			8.3	26	07/1)			
	8.5	8.4	1200	1100(60)	8.4	8.3	28	27(1)			
2 day		8.4	1200		. .	8.3	36	24(2,2)			
	8.5	8.4	1200	1200(5)	8.4	8.3	31	J4(2•J)			
3 day		8.3	1500		. <i>i</i>	8.4	31	2449.00			
	8.5	8.3	1400	1450(50)	8.4	8.4	38	34(2.9)			
l week		8.5	1800						8.3	36	
	8.5	8.4	2000	1900(20)	1900(20) 8.4	8.3	37	36(1)			
3 week		8.5	1600			8.5	34				
	8.5	8.5	1500	1550(10)	8.4	8.5	34	34(1)			
6 week		8.5	1400		•	8.6	44				
	8.5	8.5	1400	1400(5)	8.4	8.6	43	43(1)			

AMERICIUM AND PLUTONIUM SHORT-TERM SORPTION RATIOS ON USW G-4-1502 CRUSHED ROCK²

^aThe rock particle size was 75-500 μ m.

^bThe americium concentration was 1.3×10^{-8} M.

^cThe plutonium concentration was 3.1×10^{-8} M.

d Aliquots of the same americium or plutonium feed solutions were used for all of these experiments.

^eThe values in parentheses are the standard deviation of the means.

			Pu		Am		
		Sorption	<u> </u>	6-week Rd		6-week R	
Core	Core Sorption Category	Time (weeks)	$\frac{K_d}{(mt/g)}$	3-week R _d	Kd (ml/g)	3-week R _d	
JA-37	Devitrified.	3	~300		18 000		
	Smectite	6	~490	1.63	37 000	2.06	
YM-22	Devidend Stad	3	37	1 73	1500	0 72	
	Devitiiied	6	64	1.73	1100	0.75	
YM-38	714643	3	58	2 07	6100		
	Zeolitized	6	120	2.07	5200	0.05	
YM-49	Zeolite and	3	140	0 1.29 30	29 00	0 97	
•	Glass	6	~180		2800	V •77	
YM-54		3	52	1 66	150	1 07	
	Devitrified	6	81	1.50	160	1.0/	
G-1-1292		3	~160		N7 N7 -		
	GIASE	6	~530	3.31	NOT ME	BEUrea	
G-1-1883	-	3	~51		4200	1 07	
	Devitrified	6.	~72	1.41	4500	1.07	
G-4-1502	.	3	34(1)		1550(10)	0.00	
	Zeolitized	6	43(1)	1.26	1400(5)	0.90	

THE 3- AND 6-WEEK SORPTION MEASUREMENTS FOR PLUTONIUM AND AMERICIUM

COMPARISON OF NEPTUNIUM SORPTION RATIOS UNDER AMBIENT AND CO₂-CONTROLLED ATMOSPHERIC CONDITIONS

		Traced	Traced Feed	Sorption Ratio (mg/g)		
Core ^a	Atmospheric ^D Conditions	Feed PH	Concentration (M)	Experiment (pl	tal Value H)	Average Value
GU-3-1203	Ambient	8.6	3×10^{-11}	0.5 (8.4)	0.49 (8.5)	0.49(1)
	co ₂	7.1	6×10^{-11}	0.36 (6.8)	0.35 (6.8)	0.35(1)
GU-3-1301	Ambient	8.4	2×10^{-11}	1.8 (8.53)	1.7 (8.54)	1.7(1)
	co ₂		2.5×10^{-11}	2.1 (6.65)	2.1 (6.63)	2.1(1)
G-4-1608	Ambient	8.7	3.5×10^{-11}	6.1 (8.3)	6.6 (8.4)	6.3(1)
	co ₂	7.2	2×10^{-11}	5.2 (6.7)	5.6 (6.8)	5.4(1)

^aFraction size was 75-500 μ m.

 $^{b}CO_{2}$ -atmosphere was controlled to be enriched to ~5% in $^{CO}2$.

^CSorption times were 6 weeks.

^dNumbers in parentheses are standard deviations of the mean.

AVERAGE SORPTION RATIOS FROM BATCH AND CIRCULATING-SYSTEM SORPTION MEASUREMENTS

•		Rd	(ml/g)
Element	Tuff Core	Batch	Circulating System
Sr	YM-22	53(3)	27(2)
	YM-54	62(12)	45(3)
	YM-37	287(14)	390(10)
Cs	YM-22	290(30)	490(50)
	YM-54	180(40)	120(10)
	JA-37	610(40)	1800(80)
Ba	YM-22	900(30)	120(10)
	YM-54	400(150)	130(10)
	JA-37	760(150)	860(40)
Am	YM-49	4300(1400)	2200(300)
	JA-37	28 000(10000)	3400(600)
	G-1-1883	4700(300)	3300(100)
Pu	YM-49	230(50)	570(170)
	JA-37	400(70)	290(20)
	G-1-1883	77(11)	56(11)

SUMMARY OF DIFFERENCES IN SORPTION RATIOS BETWEEN CRUSHED-ROCK COLUMN AND BATCH MEASUREMENTS^a

Core	85 _{Sr}	137 _{Cs}	133 _{Ba}
YM-54	- 0	0	. 0
YM-38		C > B. (1.45)	
YM-22	$C < B^{b}$	0	с < в ^ъ
JA-37	C < B (2.08)	$C > B (1.36)^{b}$	
JA-32	0		
G-1-3658	C < B (11.8)	C < B (1.18)	C < B (1.91)
G-1-3116	С < В	С < В	0
G-1-2901	C < B	C < B	С < В
G-1-2476	С < В	C < B	C < B
G-1-2363	0	С < В	C < B
G-1-2233	Û	С > В	0
G-1-1982	0	0	C < B (1.18)
G-1-1883	0	0	C < B (1.23)
G-1-1292	0	C > B	C < B

^a"O" indicates the column R, value fell within the range of measured batch values or that the ranges overlapped (if more than one column measurement was made). C < B or C > B indicates the column R, value was less than or greater than the batch R, value. If a number follows in parentheses, it is the factor by which column and batch values differ. Numbers are not given when only batch data on unwashed fractions are unavailable for comparison. This is because the apparent differences between column and batch values may not be real.

^bC = B when results from the "circulating" type of batch measurements are included.

SORPTION OF WAFERS ON TUFF G-1-1883ª

D d	R_{d} (ml/g)					
(hours)		Cs	Ba			
2.55	0.85	11	7.5			
4.53	8.8	25	20			
6.00	11	32	26			
10.0	18	. 63	50			
15.0	22	75	61			
21.1	26	100	87			
24.0	27	100	93			
120	24	190	130			
144	25	210	140			
192	26	230	140			
305	24	230	150			
(504)	(22)	(190)	(180)			
869	27	230	210			
[Column]	[14;28]	[129;283]	[85;162]			

SORPTION ON WAFERS OF TUFF G-1-1982^a

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Dime	R _d (m ² /g)				
(hours)	Sr	Cs	Ba		
5.0	17	65	45		
7.0	30	110	78		
12.0	36	180	130		
16.0	45	190	170		
112	84	620	580		
136	86	670	560		
190	88	760	560		
303	84	750	650		
(336)	(53)	(1120)	(670)		
(504)	(62)	(1200)	(800)		
860	80	1000	710		
[Column]	[53]	[1350;1720]	[483]		

^aThe numbers in parentheses are the contact time and average R values from batch measurements with washed fractions >38 μ m.