

# Department of Energy

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John J. Linehan, Director Repository Licensing and Quality Assurance Project Directorate Division of High Level Waste Management U.S. Nuclear Regulatory Commission 4-H-3 Washington, DC 20555

TRANSMITTAL OF REQUESTED REPORT

Reference: Letter, Linehan to Stein, dtd. 5/14/90

Enclosed are three copies of the report you requested entitled: "Mineral Evaluation of the Yucca Mountain Addition, Nye County, Nevada." This report, dated December 27, 1989, was prepared by the Nevada Bureau of Mines and Geology in support of the U.S. Department of Energy's request of the U.S. Bureau of Land Management (BLM) for administrative land withdrawal of that portion of BLM lands falling within the Yucca Mountain site.

The enclosed report contains color photocopies in an attempt to fulfill your request for actual photographs of the report figures. The Nevada Bureau of Mines and Geology holds the original photographs and transparencies and these would have to be requested from them.

Additional copies of the subject report are being prepared by the Yucca Mountain Project Office and should be available within a few weeks. If you have any questions, please contact Ardyth M. Simmons of my staff at (702) 794-7998 or FTS 544-7998.

Carl P. Gertz, Project Manager YMP:AMS-4088 Yucca Mountain Project Office

Enclosure: "Mineral Evaluation of the Yucca Mountain Addition, Nye County, Nevada"



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# MINERAL EVALUATION OF THE YUCCA MOUNTAIN ADDITION, NYE COUNTY, NEVADA

Prepared for:

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by

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December 27, 1989

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## ILLUSTRATIONS

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#### PURPOSE

Early in January 1989, the U.S. Department of Energy filed an application with the U.S. Bureau of Land Management for an administrative land withdrawal of 4255.5 acres bordering the western edge of the Nevada Test Site (NTS) and the southern edge of the Nellis Air Force Range. Notice of the withdrawal application was published in the Federal Register (January 13, 1989; vol. 54, no. 9, p. 1452) by the Bureau of Land Management. This land is herein referred to as the Yucca Mountain Addition. Approximately 400 acres in the northeastern part of the Yucca Mountain Addition is within a 1,500-acre area that includes NTS and Nellis Air Force Range land being considered as a potential repository for high-level nuclear waste. BLM land including the Yucca Mountain Addition, along with adjacent parts of Nellis Air Force Range and the NTS, have been the focus of extensive study during the past 10 years. In requesting the withdrawal of this acreage, the Department of Energy stated that as preparations begin for more specific studies on the geology of Yucca Mountain it is essential that the subsurface area be secured from any interference that could hinder research to determine the suitability of Yucca Mountain for a high-level waste repository (DOE News, January 13, 1989).

The Federal Land Policy and Management Act of 1976 (FLPMA) requires that when a request for land withdrawal such as this is made, the requesting agency furnish to Congress "...a report prepared by a qualified mining engineer, engineering geologist, or geologist which shall include but not be limited to information on: general geology, known mineral deposits, past and present mineral production, mining claims, mineral leases, evaluation of future mineral potential, present and potential market demands." The Nevada Bureau of Mines and Geology (NBMG) has performed this mineral evaluation to meet these requirements.

#### METHODS OF STUDY

The NBMG mineral evaluation was accomplished over a fourmonth period by three senior scientists, a mineral technician, an analytical technician, a graduate geology student, an office manager, and a secretary-typist. Initial work consisted of a literature search, project planning, and obtaining NTS access.

Remote sensing data were collected and analyzed to determine surface alteration and linear structural patterns for the Yucca Mountain Addition and known precious-metal districts in the region. Ground control data were collected in both the Yucca Mountain Addition and precious-metal mining districts.

Field examination and sample collection in the Yucca Mountain Addition were accomplished over a one month period, and a week was

devoted to similar work in the mining districts. Samples were analyzed for selected elements by a commercial laboratory and by the NBMG laboratory. Mineralogy and petrology were determined by X-ray diffraction analyses and thin section examinations. The final stages of the work included computer processing of geochemical data and report writing which together took one month.

#### ACKNOWLEDGMENTS

D. A. Davis performed field work and prepared thin-sections, and P. M. Goldstrand also performed field work. L. J. Garside reviewed parts of the report. Other NBMG personnel who helped on this project are M. Desilets, D. Meeuwig, and L. E. Jacox.

#### **LOCATION**

The proposed Yucca Mountain Addition land withdrawal is located about 30 km southeast of the town of Beatty in southern Nye County, Nevada (Figure 1). The small community of Amargosa Valley lies about 20 km south of Yucca Mountain, and Las Vegas, the closest large city, is about 150 km to the southeast.

The land proposed for closure to public access and withdrawal from mineral entry consists of 4255.5 acres of public land specifically described as all of sections 7, 8, 9; the W/2 of section 10; the W/2 of section 15; all of sections 16, 17; the NE/4 of section 20; the N/2 and the N/2 of the S/2 of section 21; and the NW/4 and the N/2 of the SW/4 of section 22, Township 13 South, Range 49 East, Nye County, Nevada. This proposed withdrawal covers the southern part of Yucca Mountain and extends from the mouth of Windy Wash on the west, across the southern end of Solitario Canyon and the crest of Yucca Mountain, to the lower flanks of the mountain on the east. The northern boundary of this area is the southern boundary of the Nellis Air Force Range; the eastern boundary is the western boundary of the NTS.

#### **HISTORY AND PRODUCTION**

#### MINING DISTRICTS SURROUNDING THE YUCCA MOUNTAIN ADDITION

Recorded information on the history of the Yucca Mountain Addition and its immediate vicinity is sparse. Cane Springs, about 30 km east of Yucca Mountain, was a watering stop on the Death Valley Emigrant Trail in 1849. The trail descended down Fortymile Wash, passed the south end of Yucca Mountain, and continued into Death Valley southwest of the present-day town of

Amargosa Valley. The earliest mining activity in the area surrounding Yucca Mountain is associated, by popular accounts, with travel along the Emigrant Trail. An article in the Tonopah Daily Times of February 14, 1928 states that "...the old Hornsilver mine (northwest of Cane Springs) had been worked by Mormons in 1853." Ball (1907) visited the Hornsilver mine during a reconnaissance trip through the area in 1905. He made no comments on activity at the mine but, by noting it on his maps, documented that it was known at the time of his visit. Old prospect pits and shafts in the Calico Hills, about 13 km

northeast of Yucca Mountain, may also date from the 1905 period of activity. The Bullfrog district, located west of the town of Beatty and about 30 km west of southern Yucca Mountain, is a large district with recorded precious-metals production. The Original district with recorded precious-metals production. Bullfrog mine in the west part of the district was discovered in 1904; production from the district is recorded from 1907 to 1940 and totals about 3 million dollars (Couch and Carpenter, 1943).





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The earliest recorded mining activity in the districts immediately surrounding Yucca Mountain began about 1905 with discoveries of gold ore on the east slope of Bare Mountain. The camp of Telluride sprang up at this site, about 20 km northwest of the south end of Yucca Mountain, and gold ore was produced from 1913 to 1915 (Lincoln, 1923). Mercury was discovered in this same area in 1908 and small amounts of mercury were produced from the district intermittently through 1953 (Cornwall, 1972).

Fluorite was discovered during exploration for precious metals in the northern part of the Bare Mountain mining district in about 1906. Fluorspar has been mined in the district more-orless continuously since 1919, mainly from the Daisy mine (Papke, 1979), and the total production of fluorspar ore from the district is in excess of 300,000 tons.

In 1928 new discoveries of high-grade silver-gold ore near the Hornsilver mine resulted in a second rush to that area and the boom-camp of Wahmonie sprang up in the flats southeast of the old mine. Claim staking covered an area of several square km and a new, 150-m, shaft was sunk by the Wingfield interests (Goldfield Consolidated Mines Co. of Goldfield, Nevada) (Tonopah Daily Times, February 22, 1928). The boom collapsed quickly, however, and no<br>recorded production resulted from this activity. In 1940, the recorded production resulted from this activity. Wahmonie district and the adjacent Calico Hills area were included within the Tonopah Bombing and Gunnery Range (now known as the Nevada Test Site and Nellis Ranges) and withdrawn from mineral entry. No mineral activity has occurred in the Calico Hills or at Wahmonie since that time.

In the districts west of the proposed Yucca Mountain Addition, precious-metals mining activity revived in 1980 with the discovery of a disseminated gold orebody at the Sterling mine. The Sterling, located across Crater Flat about 11 km due west of Yucca Mountain began gold production in 1984 and has produced from 9000 to 7000 oz gold per year since that time (Nevada Bureau of Mines and Geology, 1984, 1988). In 1988, GEXA Gold Corp. announced discovery of significant reserves of gold ore on its Mother Lode property at the north end of Bare Mountain, about 8 km north of the Sterling mine. Announcements in early 1989 described reserves at the Mother Lode of about 4.4 million tons of ore averaging 0.054 oz/ton gold in a deposit hidden under shallow gravel cover near the old camp of Telluride.

Major new mining activity is also underway in the Bullfrog district. In 1982, St. Joe American began evaluation of the district and by 1985 had developed reserves of mineable ore at the old Montgomery-Shoshone mine northeast of the old camp of Rhyolite. Continued exploration in the district resulted in the discovery of an entirely new orebody near Ladd Mountain 1 km southeast of Rhyolite. Announced reserves are 3,088,000 tons grading 0.072 oz gold per ton at Montgomery-Shoshone and 14,300,000 tons grading 0.110 oz gold per ton at the new Bullfrog deposit (Jorgensen et al., 1989). Production from these two mines will soon eclipse all historic production from the entire district. In the western part of the Bullfrog district, north of

the Original Bullfrog mine, exploration by other companies has outlined 1.23 million tons of gold ore at the Gold Bar mine; this deposit, which was being evaluated in 1987 (Nevada Bureau of Mines and Geology, 1988), is now being mined.

In addition to precious-metals exploration and mining on Bare Mountain, the only other active mining operation in the immediate area of Yucca Mountain is recovery of volcanic cinder from cinder cones on the south end of Crater Flat. The deposit currently being mined is adjacent to U.S. Highway 95 about 13 km southwest of Yucca Mountain. Other cinder cones dot this portion of Crater Flat, however, and mining could extend to them at some time in the future.

White montmorillonite clay has been mined since the early 1950's from the New Discovery mine about a mile south of Beatty. In addition, minor clay production has come from claims located just northeast of Beatty. An unknown amount of ceramic silica was mined from from the Silicon mine at the northwest end of Yucca Mountain about 20 km northwest of the Yucca Mountain Addition.

#### YUCCA MOUNTAIN ADDITION

There is no record or evidence of any historic mining activity within the boundaries of the Yucca Mountain Addition. With the exception of the road in Solitario Canyon, there were probably no roads or trails traversing the area prior to the present period of DOE-related activity. No mining or prospecting excavations of any type were noted during the present study.

On June 29, 1987, Anthony J. Perchetti of Tonopah, Nevada staked a block of 10 lode mining claims (Yucca 1-10) in sections 9, 16, and 21, T13S, R49E. These claims were located end-to-end and generally follow the crest of southern Yucca Mountain within the area now included in the proposed withdrawal. Perchetti located 17 additional lode claims in a block west of his earlier claims 1-3 on December 1, 1988. On December 17, 1988, four lode mining claims, the Lucky 1-4, were located north of the Perchetti claims in section 9, T13S, R49E, by Robert F. Fowler of Lakewood City, California. It is believed that Mr. Fowler is an associate of Mr. Perchetti. In a letter to Mr. Carl Gertz, Waste Management Project Office, U.S. Department of Energy, Las Vegas, Nevada, dated January 6, 1989, Mr. Perchetti stated his belief that his claim block on Yucca Mountain had a likelihood of containing gold deposits similar to those found on GEXA's Mother Lode holdings on nearby Bare Mountain. Other than sampling, apparently no work was done on these claims by Perchetti. We have not seen the Perchetti sample results. In mid-1989 Mr. Perchetti sold his interests to In mid-1989 Mr. Perchetti sold his interests to the U. S. Department of Energy, and the U. S. Bureau of Land Management closed the case on all 27 Yucca claims on March 13, 1989. In addition, the U. S. Department of Energy purchased Mr. Fowler's Lucky claims in 1989, but as of October 17, 1989 no notice of closure was reported in the U. S. Bureau of Land Management's index of active mining claims in Nevada.

### REGIONAL LITHOLOGY

The geology of the area around the Yucca Mountain Addition is dominated by the rocks of the southwestern Nevada volcanic field as defined by Byers et al.(1976). Ash-flow tuffs and related rocks from at least four major Miocene calderas comprise this volcanic field, which once covered an area of more than  $10,000$  km<sup>2</sup> (Byers et. al, 1989). The Tertiary volcanic rocks overlie metamorphic and sedimentary rocks ranging in age from Precambrian to Mississippian.

The earliest extrusions of the southeastern Nevada volcanic field are 16-Ma (million year old) tuffs erupted from the Sleeping Butte caldera about 35 km northwest of the Yucca Mountain Addition (Christiansen et al., 1977). About 13.5 Ma (million years ago), ash flows were erupted that are considered by Carr et al. (1986a) to be from the Crater Flat - Prospector Pass Caldera complex, whose eastern wall is thought to lie beneath the Solitario Canyon area within the Yucca Mountain Addition. At about the same time, lavas and tuffs were deposited in the Calico Hills, 10 km northeast of the Yucca Mountain Addition.

The Paintbrush Tuff, an extensive formation consisting of ash flows and associated air fall units, was erupted approximately 13 Ma from the Claim Canyon caldera about 5 km north of the Yucca Mountain Addition (Byers et al., 1976 and 1989). The upper parts of the Paintbrush Tuff may have originated from the Oasis Valley caldera further to the northwest (Christiansen et al., 1977).

At Wahmonie, about 30 km east of the Yucca Mountain Addition, flows and tuffs with associated intrusives were deposited during the same time period as the Paintbrush Tuff. The andesitic to rhyodacitic Wahmonie rocks differ in composition from most other rocks of the southwestern volcanic field which are predominantly rhyolitic to latitic.

The Timber Mountain Tuff, the most voluminous formation in the southwestern Nevada volcanic field, was erupted about 11 Ma from the Timber Mountain caldera, a large, well-preserved feature centered about 25 km north of the Yucca Mountain Addition (Byers et al., 1976). Post-collapse volcanic activity in the Timber Mountain caldera is thought to have continued until about 10 Ma. Rhyolitic lavas and tuffs in the Bullfrog Hills about 30 km west of the Yucca Mountain Addition have also yielded ages as young as 10 Ma (Jackson, 1988).

Basalt yielding 10 and 4 Ma dates has been drilled beneath Crater Flat 5 to 10 km west of the Yucca Mountain Addition, and basalt flows and cinder cones about 1 Ma occur at the surface on Crater Flat (Carr, 1988a). Lava flows at the Lathrop Wells volcanic center 13 km south of the Yucca Mountain Addition are thought to be less than 100 Ka, and cinder cone deposition at the same site is estimated at less than 15 Ka (Crowe et al., 1988).

#### REGIONAL STRUCTURE

According to most recent work, structures within the rocks of the southwestern Nevada volcanic field mainly result from two interactive late Cenozoic tectonic processes: extension of the southern Basin and Range province, and caldera subsidence. Numerous west-dipping normal faults, low-angle detachment surfaces, and northwesterly strike-slip faults were formed, as well as fault patterns related to cauldron subsidence and resurgent doming. The southwestern Nevada volcanic field lies within the Walker Lane belt, a 100-km-wide northwesterly zone dominated structurally by lateral shear. Paleozoic and Mesozoic structures have been modified both by the Tertiary calderas and by crustal extension during middle to late Cenozoic time..

Steeply westward-dipping normal faults are abundant in the Yucca Mountain area (Scott and Bonk, 1984), and may represent the breakaway zone for detachment faulting bounding Bare Mountain 10 km east of the Yucca Mountain Addition. According to Hamilton (1988), normal faults at Yucca Mountain occurred above an 11 Ma hinge line between more-or-less flat-lying detachment to the east and moderately westward dipping detachment to the west. Hamilton believes that this detachment surface, which separates Tertiary volcanics from pre-Tertiary rocks, now extends from east of Yucca Mountain in the subsurface, through surface exposures on Bare Mountain and the Funeral Mountains, into Death Valley. On the other hand, Carr (1988a) relates Yucca Mountain normal faulting to reactivation along subsidence-bounding fractures of the proposed 14 Ma Crater Flat-Prospector Pass caldera. Carr (1988b) further speculates that the southwestern Nevada volcanic field calderas were emplaced along a north-south rift in the Walker Lane belt separating detachment faulted terrain to the west from gravitationally caused extensional faulting to the east. Rangebounding faults around Bare Mountain are considered by Hamilton to be part of a low to moderately dipping detachment surface, whereas Carr believes that the fault bounding the east side of Bare Mountain is a steeply dipping feature related mainly to subsidence of the Crater Flat-Prospector Pass caldera.

Geophysical data suggest that the faults bounding the Bare Mountain Precambrian and Paleozoic block dip at **300** or less (Ackermann et al., 1988). The Fluorspar Canyon fault, which forms the northern boundary of Paleozoic rocks in the Bare Mountain area is thought to be an extension of the nearly flat-lying Bullfrog detachment fault between Tertiary and pre-Tertiary rocks at the Original Bullfrog mine 30 km west of the Yucca Mountain Addition. Although surface outcrops along the Fluorspar Canyon fault suggest northerly dips of 60° or more, drilling at the Mother Lode gold deposit near the east end of the fault suggests a northerly dip of about 30° (M. R. Mapa, personal communication, 1989).

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#### LITHOLOGY

Within the Yucca Mountain Addition, pre-Quaternary surface lithology consists of only three members of two ash-flow formations with attendant bedded air-fall sequences. The 13.3 Ma Topopah Spring Member of the Paintbrush Tuff (ptu, Plate 1) is the oldest rock unit exposed in the area. It is characterized by the presence of pink to gray lithophysal ash-flow tuff with sparse phenocrysts (generally under 10 percent) overlain by a few meters of red-brown devitrified caprock containing about 15 percent phenocrysts, including distinctive large amber-colored biotite books. Above this is a thin black to reddish-brown vitrophyre with relatively abundant phenocrysts which is included in the Topopah Spring Member caprock by Scott and Bonk (1984).

Within the Yucca Mountain Addition, the Topopah Spring Member is overlain by a sequence composed mainly of nonresistant white to orange or light-purple, nonwelded, pumiceous bedded air-fall tuffs (bt, Plate 1). In the southwestern part of the Yucca Mountain Addition, this unit appears to contain a glassy purplish-gray welded tuff (sample SC 93). Above the nonwelded bedded tuff sequence is a distinctive glassy orange welded tuff (locally gray or bicolored orange and gray) which appears to grade downward into nonwelded orange bedded tuff.

Locally, above the orange welded tuff, is a black to darkgreenish-gray crystal-poor vitrophyre that grades up into glassy to devitrified ash-flow tuff typified by columnar jointing. According to Scott and Bonk (1984), this is the basal unit of the Tiva Canyon Member (pcu, Plate 1) of the Paintbrush Tuff. Most of the overlying Tiva Canyon Member is comprised of flaggy, crystalpoor (averaging less than 10 percent phenocrysts) devitrified gray ash-flow tuff much of which rings when struck (clinkstone zones of Scott and Bonk, 1984). The Tiva Canyon caprock is cliff-forming devitrifed to glassy, gray ash-flow tuff with moderately abundant phenocrysts.

The Rainier Mesa Member of the Timber Mountain Tuff is exposed in the' southwestern to central part of the Yucca Mountain Addition, mostly on Plug Hill (Tmr, Plate 1). It is a light-gray to light-pink devitrified ash flow containing relatively abundant quartz and feldspar phenocrysts. It is underlain by nonresistant white to light-pink airfall exposed in places along the wash west of Plug Hill (bt, Plate 1).

Quaternary gravels (QTac, Plate 1) consisting mainly of fanglomerate and alluvium underlie washes and canyons in the Yucca Mountain Addition, and locally occur low on the flanks of some ridges. Caliche cement is common in these gravels, particularly along and adjacent to faults. Some exhumed gravels may be Tertiary.

#### **STRUCTURE**

Structure within the Yucca Mountain Addition is almost totally dominated by normal faults that strike north-south to north-northeast and dip steeply west (Scott and Bonk, 1984). Most of the major faults belong to this set (e.g.: the Solitario Canyon fault, a fault along the north-northeast-trending arm of Abandoned<br>Wash: and a fault along the west side of Boomerang Point). A Wash; and a fault along the west side of Boomerang Point). minor subset of north-northwest to northwest faults dipping steeply southwest with apparent normal displacements is also present, particularly in the south central part of the Yucca Mountain Addition (Plate 1). Other fault orientations are rare.

Many of the faults in the Yucca Mountain Addition are marked by resistant silicified breccia (Figure 2). Such faults are especially common west of Yucca Crest near the Solitario Canyon fault, and on the west flank of Boomerang Point. Calcretecemented breccia is also very common along faults, and serves as<br>an indicator of poorly exposed faults in some cases. In many an indicator of poorly exposed faults in some cases. instances, fault breccia is cemented by silica and calcrete (Figure 3). Faults with clayey and(or) sandy gouge (Figure 4) are rare in the Yucca Mountain Addition.

### **ECONOMIC GEOLOGY EVALUATION**

This study was designed to test the potential for economic minerals within the Yucca Mountain Addition only. Extension of the findings of this study into other areas encompassed in the Yucca Mountain project and outside the Yucca Mountain Addition are not intended. In addition, because access to subsurface samples was not granted by the DOE due to quality assurance concerns, determinations of economic potential to depths greater than a few hundred meters were not possible.

Because of the presence of the Perchetti claims, and allegations of economic precious-metal potential made publically and privately by Mr. Perchetti, this economic evaluation is focused on a determination of precious-metal potential within the Yucca Mountain Addition. The field and laboratory methods used were mainly directed toward determination of precious-metal potential. However, other commodities were considered that could occur in economic amounts based on the geologic setting.

#### FIELD EXAMINATION

Fieldwork for the economic minerals evaluation in the Yucca Mountain Addition was begun on May 18, 1989, and completed on June 28, 1989. Outcrops were examined and samples collected along foot traverses made over the entire Yucca Mountain Addition (Plate 2).

Figure 2. Silica-cemented fault breccia near sample location YMSC 51, view to north.



Figure 3. Silica- and calcrete-cemented fault breccia near sample location YMSC 69. Located on the west side of Boomerang Point, view to north.



Figure 4. Gouge zone along Abandoned Wash fault. Fault downdrops devitrified Tiva Canyon Member ash flow on west (back-ground) against Paintbrush Tuff bedded tuff on east (fore-ground). Sample location YMSC 13.



Sample locations were marked on copies of the 1:12,000 scale geologic map of Scott and Bonk (1984) and on 1:12,000 scale enlargements of 7.5 minute quadrangle maps. During sample collection, visual descriptions of mineralogy, lithology, and structures encountered were recorded (Appendix A). Samples were collected mainly from veins, fracture coatings, fault zones, bodies of tectonic breccia, and areas of altered rock. Initially, 200 samples were collected from the Yucca Mountain Addition study area.

Samples were submitted to Geochemical Services Inc. (GSI), of Rocklin, California, for 15-element inductively-coupled plasma (ICP) emission spectroscopic analyses, and gold analysis by graphite furnace atomic absorption . Sample crushing and pulverization were done using either NBMG bucking facilities, or those of GSI located in Sparks, Nevada. Ten blank quartz samples were'submitted initially along with Yucca Mountain samples to monitor possible contamination during sample preparation.

Following receipt of analytical results for the initial samples, 20 samples- were submitted for corroborative analyses to the NBMG geochemical laboratory. In addition, 38 sample sites, including all sites that yielded samples with elevated trace

element contents, were revisited, resampled, and marked with aluminum tags and red flags (Figure 5). During this resampling work, 22 samples were collected from new sites. Altogether, 260 samples were collected from the Yucca Mountain Addition and subjected to multi-element analyses (Appendix B).

In addition to field work on the Yucca Mountain Addition, field data and samples were collected from two new gold and silver mines in the Rhyolite-Bullfrog area, a recently discovered gold deposit in the Bare Mountain area, and four abandoned mining areas known to have had past production of gold and silver. All of the current, potential, or past producers of precious metals examined have mineralized volcanic rock that is contemporaneous, or nearly so, with Yucca Mountain Addition rocks. A total of 122 samples (including 7 resubmitted blind for control) were analyzed from these areas by GSI with the same techniques used for the Yucca Mountain Addition samples.

Mineralogic and petrographic work was done by NBMG and Desert Research Institute personnel on selected samples from the Yucca Mountain Addition and the precious-metal districts examined. Xray diffraction analyses were performed on 54 samples from the Yucca Mountain Addition and 29 samples from three of the four precious-metal districts examined (Appendix C). In addition, 8 thin sections of representative veins, breccias, and altered rocks from the Yucca Mountain Addition were analyzed petrographically.



Figure 5. Sample location YMSC 52. A narrow zone of hematitic silicified air-fall tuff (beneath aluminum tag).

#### GEOCHEMICAL RESULTS

Several samples from the Yucca Mountain Addition were found to contain slightly anomalous amounts of silver, arsenic, bismuth, lead, and(or) thallium; however, the highest contents of most analyzed elements in samples collected from the Yucca Mountain Addition during this study are near or below background values in the earth's crust (Levinson, 1974). Analyses for three elements: palladium, selenium, and tellurium, all fell below detection levels (see Table 1). Most of the anomalous Yucca Mountain Addition samples were taken from the Paintbrush bedded tuff unit (pbt, Plate 1), and these are mainly silicified rock, but one is unaltered glassy air-fall tuff, and one is from a glassy tuff dike.

Perhaps the most interesting sample, in terms of its geochemistry, is YMSC 52, a sample of bright red silicified airfall tuff (Figure 2) from a northerly vein on the west flank of Yucca Crest. The vein is probably less than 20 cm wide and could only be followed on the surface for a few meters. This sample contains 4 to 8 ppm bismuth and 109 to 145 ppm lead. A single analysis of sample YMSC 92, brown silicified air-fall tuff from a zone approximately parallel to the nearby YMSC 52 vein, yielded a value of 1 ppm bismuth.

Sample YMSC 22, containing 3 to 4 ppm bismuth, 64 to 97 ppm zinc, and 2 ppm thallium, was collected from a 3-cm-wide dike of friable, glassy, pink and white air-fall tuff which dips steeply southwest. This tuff dike cuts an irregular 5- to 10-cm-wide northeasterly zone of gray opalized tuff (sample YMSC 22S) carrying 0.6 to 1.8 ppm bismuth. The host rock at this locality is gray glassy welded tuff which appears to be in the Paintbrush bedded tuff sequence.

Sample YMSC 31, a composite of purplish-gray silicified airfall tuff and hematitic gouge from a fault dipping steeply west, contains 0.6 to 1.5 ppm bismuth and 60 to 68 ppm zinc. Sample YMSC 31B, hematitic air-fall tuff taken from the hanging wall of this fault, contains 27 ppm arsenic and 0.3 ppm bismuth.

Other anomalous samples within the Paintbrush bedded tuff sequence are YMSC 14C, a bed of fine, well-sorted, and apparently unaltered lapilli tuff containing 0.4 to 0.5 ppm silver; and YMSC 88 from a near vertical northerly calcrete-silica vein system with 1 ppm bismuth.

Only two geochemically anomalous samples were collected from outside the Paintbrush bedded tuff sequence. Sample YMDD 36A, which contains 32 ppm arsenic and 1 ppm bismuth, is purplish-gray silicified breccia with some irregular veins of white opal collected from a fault dipping steeply southwest. Sample YMSC 45, with 0.1 ppm silver, is of calcrete vein material from a poor exposure. The host rock for both occurrences is devitrified ashflow tuff of the Topopah Spring Member of the Paintbrush Tuff.

None of the samples collected from the Yucca Mountain Addition can be said to have highly anomalous gold contents. The highest gold value for any Yucca Mountain Addition sample reported by GSI is 0.026 ppm (YMSC 66); however, analysis of a blind resubmitted sample resulted in a 0.003 ppm gold value, and gold was not detected in rock obtained by resampling the same outcrop (sample YMSC 66A). The highest gold value for all other samples from the Yucca Mountain Addition analyzed by GSI is 0.009 ppm, equivalent to 0.0003 ounces per ton.

Table 1. Elements analyzed, detection limits, and median values by area. Analyses by Geochemical Services Inc. All values in ppm.



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Analyses performed by the NBMG geochemical laboratory on Yucca Mountain Addition samples yielded higher gold numbers than those obtained by GSI (Table 2). The NBMG analyses were performed by fire assay with atomic absorption finish, a technique which can be expected to yield higher results than that used by GSI, which may not extract all of the gold in the sample during dissolution (P. Lechler, personal communication, 1989). The highest gold value obtained by NBMG on a Yucca Mountain Addition sample is 0.023 ppm, or about 0.0007 troy ounces per ton. A simple regression curve fitted to a plot of NBMG gold analyses versus GSI gold analyses for the same samples shows that the projected maximum NBMG value is 0.0245 ppm, about 0.0007 troy ounces per ton (Figure 6).

### STATISTICAL EVALUATION OF GEOCHEMICAL SAMPLING

Statistical calculations of sampling results from the Yucca Mountain Addition evaluation were made using statistical software developed by Koch (1987). In addition to calculations on data from the Yucca Mountain withdrawal, statistical calculations were performed on data from the Mother Lode property, from Rhyolite in the Bullfrog district, and from the Wahmonie district.

Geochemical values were analyzed by district and the results are presented in Tables 3 through 6. Table 1 provides a list of elements analyzed, shows detection limits, and shows median values of each element by area. Tables 3 through 6 show correlation coefficients for elements in samples from each of the four areas.

Correlation coefficients vary from +1 to -1. Correlation coefficient values quantify adequacy of fit to a linear regression curve (values of ±1 correspond to a perfect fit, and those of 0 to no linear fit), and signs indicate if the correlation is positive or negative. The significance of each coefficient was determined using a table for testing the null hypothesis  $p = 0$  (Snedecor and Cochran, 1967, p. 557). As the number of data pairs increases, the correlation coefficient will be significant at progressively lower values. Correlation coefficients significant at the 95 and 99 percent levels are indicated on Tables 3 through 6. Testing by the above methods may not be valid for data with non-normal distributions (e. g., mercury in Yucca Mountain samples).

#### Yucca Mountain Addition

Results from 196 samples collected in the Yucca Mountain Addition study area were used in statistical calculations (all samples taken during initial fieldwork with the exception of 4 samples for which analytical data was obtained too late for inclusion). In this group, gold was reported present above the limit of detection in 101 samples. Mercury was found above the limit of detection in only 14 samples, bismuth was found in only 8, and palladium, selenium, and tellurium were not found in any of the samples at the limit of detection of the analytical methods

Table 2. Comparison of GSI and NBMG analytical results. Twenty samples were analyzed for gold, arsenic, antimony, bismuth, and mercury at both labs. All values in ppm.

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Figure 6. Simple regression plot for gold analyses. NBMG versus GSI values for 20 samples. GSI values reported at <0.0005 ppm plotted at 0.00025 ppm.



Au ppm **(GSI)**

used (Table 1). All analyzed elements are present in very low concentrations (or are below detection limits) in all samples collected within the Yucca Mountain Addition. Correlation coefficients between elements (Table 3) do not show strong groupings of correlated elements. Arsenic, gold, and cadmium show weak correlations and there is a general grouping of base metals (copper, molybdenum, lead, antimony, and zinc along with bismuth and gallium) that display weak to moderate correlations.

#### Wahmonie District

Sampling at Wahmonie was less extensive than at the other three areas studied; only 12 samples are available for statistical evaluation. Median values from Wahmonie show definite enrichment of silver and tellurium and lower, but anomalous, concentrations of arsenic. Element correlations (Table 4) are limited to strong correlations between silver and mercury, silver and tellurium,

As **Au** ai I& MD Pb Sb Zn Bi Cd Ga To  $10$  .07 .10 .07 .10 .12 .10 .09 .09 .07 .09 .07 .10 .17^ .05 Ag  $.46*$   $-.16*$   $-.10$   $-.36*$   $-.03$   $-.07$   $-.25*$   $-.06$   $.26*$   $-.2$ As  $-22* -11$  .29\* .07 .06 .26\* .05 .24\* .19 **Au**  $1.09$   $.35*$   $.05$   $.26*$   $.11$   $-.04$   $-.16*$   $.10$  $\boldsymbol{\omega}$ k.04 .06 .05 -. 07 .04 f.07 .04  $H<sub>2</sub>$  $.25*$  .27\* .45\* .07 .30\* .27\* MD Pb  $04$  .34\* .74\*  $\pm 04$  .13 Sb  $.02$   $.12$   $.07$   $.32*$ Zn  $.25*$ 1-.09  $.15<sup>^</sup>$ Bi  $.08$  .  $21*$ Cd .03 Ga

Table 3. Geochemistry correlation coefficients, Yucca Mountain Addition. No coefficients are shown for Te because Te is below the detection limit in all samples.

\* Correlation coefficient significant at 99% level

A Correlation coefficient significant at 95% level

arsenic and cadmium, gold and tellurium, and mercury and tellurium. Based on our sample results, there does not seem to be a base-metals association at Wahmonie.

#### Mother Lode Deposit Area

Thirty-seven samples taken from the Mother Lode Mine area were used for statistical calculations. Gold is present above the detection limit in all 37 samples; silver is present in 34. The mean gold value of our-samples is 0.570 ppm; this is equivalent to about 0.02 oz per ton gold and compares favorably with the announced grade of the Mother Lode orebody (0.054 oz/ton) because many of our samples are of unmineralized rock collected from

|                | As     | Au           | $\alpha$ | <b>Hg</b> | Ю                     | Fb    | S <sub>b</sub> | 2n            | Bi     | $\alpha$                 | Ga   | <b>Te</b>     |
|----------------|--------|--------------|----------|-----------|-----------------------|-------|----------------|---------------|--------|--------------------------|------|---------------|
| Ag             | $-.18$ | .39          | $-.27$   | $.97*$    | .08                   | −.32  | -. 29          | .43           | $-.18$ | ⊧.œ                      | ├.39 | $.84*$        |
| As             |        | $\vdash .11$ | .03      | $-.12$    | $\mathord{\vdash}.11$ | -.03  | ⊢.04           | 10.⊾          | .04    | $.93*$                   | .43  | $\vdash$ . 10 |
| $\mathbf{A}$   |        |              | -.30     | .48       | ⊢.œ                   | ⊢.40  | $-.22$         | ⊢.02          | .01    | .03                      | - 45 | $.70*$        |
| $\alpha$       |        |              |          | −. 15     | -.03                  | .25   | $.62^{\circ}$  | . 46          | .34    | ⊢.07                     | .28  | $-.14$        |
| <b>Hg</b>      |        |              |          |           | .07                   | $-32$ | $-.11$         | .54           | $-.22$ | .01                      | ├.35 | $.86*$        |
| <b>Mo</b>      |        |              |          |           |                       | .30   | .12            | .43           | -.30   | -.20                     | .32  | .02           |
| P <sub>b</sub> |        |              |          |           |                       |       | -. 12          | .00           | .09    | $\mathsf{\mathsf{-.02}}$ | .10  | $-36$         |
| S <sub>b</sub> |        |              |          |           |                       |       |                | $.55^{\circ}$ | 1-.06  | -. 18                    | .50  | $\vdash$ . 24 |
| ${\bf Zn}$     |        |              |          |           |                       |       |                |               | $-.32$ | $-.12$                   | .32  | .28           |
| Bi             |        |              |          |           |                       |       |                |               |        | .03                      | -.39 | .19           |
| $_{\text{cd}}$ |        |              |          |           |                       |       |                |               |        |                          | .17  | ן0.⊣          |
| Ga             |        |              |          |           |                       |       |                |               |        |                          |      | $-.48$        |

Table 4. Geochemistry correlation coefficients, Wahmonie district.

\* Correlation coefficient significant at 99% level

\* Correlation coefficient significant at 95% level

outside the orebody. Median values for arsenic, antimony, and mercury are all high indicating that these elements are enriched in the area sampled. Correlation coefficients calculated from Mother Lode data (Table 5) show a precious-metals grouping (strong silver-arsenic-gold-copper-tellurium correlations) and a basemetals grouping which includes moderate to strong silver-antimonylead-zinc correlations, a weak zinc-cadmium correlation, and moderate lead-antimony-zinc-tellurium correlations. Mercury, known to be present in the district from its production history, was found to be present above detection limits in 21 of the 37 samples from the district. Mercury, however, does not correlate well with any of the elements in the precious-metals group and correlates only moderately with copper, molybdenum, and lead in the base-metals grouping.

|                       | As     | Au     | $\alpha$ | 比             | No.    | Pb     | S <sub>b</sub> | 2 <sub>c</sub> | Bi                       | $\mathbf{c}$ | Ga                    | Te               |
|-----------------------|--------|--------|----------|---------------|--------|--------|----------------|----------------|--------------------------|--------------|-----------------------|------------------|
| Ag                    | $.67*$ | $.82*$ | $.57*$   | .21           | .01    | $.58*$ | $.49*$         | .44*           | $-.09$                   | .31          | 13. –                 | .90 <sub>k</sub> |
| As                    |        | $.87*$ | $.81*$   | $.36^$        | .25    | .20    | .03            | .44            | 1-.06                    | .27          | $\mathord{\vdash}.11$ | $.78*$           |
| Au                    |        |        | $.76*$   | .25           | .20    | .26    | .13            | .29            | $-.06$                   | .21          | 1-.09                 | $.90*$           |
| $\boldsymbol{\alpha}$ |        |        |          | $.34^{\circ}$ | $.39^$ | $.40*$ | .30            | .27            | $\mathsf{L}.\mathsf{07}$ | .30          | -. 21                 | $.75*$           |
| <b>Hg</b>             |        |        |          |               | .03    | .05    | .02            | $.34^$         | .01                      | .15          | $\vdash$ .21          | .18              |
| Mo                    |        |        |          |               |        | -.04   | -.03           | )-.02          | .02                      | .12          | $-.19$                | .15              |
| Pb                    |        |        |          |               |        |        | $.95*$         | .23            | .05                      | .26          | -.09                  | $.57*$           |
| S <sub>b</sub>        |        |        |          |               |        |        |                | .17            | ⊦.04                     | .26          | -. 11                 | $.45*$           |
| 2n                    |        |        |          |               |        |        |                |                | $-.10$                   | $.52*$       | $-14$                 | $.41*$           |
| Bi                    |        |        |          |               |        |        |                |                |                          | $-.10$       | ⊢.01                  | -.07             |
| ${\bf c}$             |        |        |          |               |        |        |                |                |                          | ٠            | ⊢.04                  | $.38^{\circ}$    |
| Ga                    |        |        |          |               |        |        |                |                |                          | þ            |                       | -.08             |

Table 5. Geochemistry correlation coefficients, Mother Lod property.

\* Correlation coefficient significant at 99% level

A Correlation coefficient significant at 95% level

### Rhyolite Area

Samples from mines and mineralized outcrops in the area around Rhyolite in the Bullfrog district show high median values for gold and silver. Base-metals values in this area are very low. Examination of the correlation coefficients between elements (Table 6) shows a precious-metals grouping with strong silver-gold correlation and moderate silver-gold-antimony-bismuth-correlations. A broad base-metals association includes moderate copper-antimonybismuth, molybdenum-lead-cadmium, and zinc-cadmium correlations.

|                | As  | Au                  | $\alpha$ | Hg     | Mo     | Pb                    | 5 <sub>b</sub> | 2n    | Bi     | $\alpha$              | Ga                    | Te            |
|----------------|-----|---------------------|----------|--------|--------|-----------------------|----------------|-------|--------|-----------------------|-----------------------|---------------|
| Ag             | .15 | .94*                | .04      | .03    | ⊢.02   | $\mathord{\vdash}.11$ | .03            | 1-.06 | .01    | 108.⊣                 | $\infty$              | .02           |
| As             |     | .11                 | .03      | .18    | $.47*$ | $.40*$                | $.45*$         | .23   | ├.15   | .20                   | .19                   | -. 13         |
| Au             |     |                     | .00      | .03    | -.04   | -.05                  | .02            | -.08  | .12    | $\mathord{\vdash}.11$ | -.05                  | .29           |
| $\alpha$       |     |                     |          | $-.14$ | .00    | $\mathsf{-.07}$       | $.66*$         | .12   | $.41*$ | $.32^{\circ}$         | .25                   | ⊢.04          |
| <b>Hg</b>      |     |                     |          |        | -.07   | 10.⊣                  | -.07           | .15   | 1-.06  | .05                   | .03                   | .05           |
| Mo             |     |                     |          |        |        | $.64*$                | .26            | .14   | -.04   | $.43*$                | $-.12$                | .03           |
| P <sub>b</sub> |     |                     |          |        |        |                       | .16            | .14   | .11    | .29                   | 10.⊣                  | .35           |
| S <sub>b</sub> |     | $\hat{\phantom{a}}$ |          |        |        |                       |                | -.06  | $.51*$ | .09                   | $\mathord{\vdash}.11$ | ├. 13         |
| 2n             |     |                     |          |        |        |                       |                |       | -.20   | .52★                  | $.38^*$               | $\vdash$ . 12 |
| Bi             |     |                     |          |        |        |                       |                |       |        | $\vdash$ . 10         | -.20                  | $.51*$        |
| $C_{d}$        |     |                     |          |        |        |                       |                |       |        |                       | .20                   | -.05          |
| Ga             |     |                     |          |        |        |                       |                |       |        |                       |                       | $-.11$        |

Table 6. Geochemistry correlation coefficients, Rhyolite area.

\* Correlation coefficient significant at 99Z level

**<sup>A</sup>**Correlation coefficient significant at 95% level

#### Summary and Comparison of Areas

A comparison of element median values shows very substantial enrichments in both gold and silver in the established preciousmetal districts relative to the Yucca Mountain Addition (Table 1). Median precious-metals values in samples from Wahmonie, Mother Lode, and Rhyolite are 5 to 100 times the median values found in samples taken within the Yucca Mountain Addition. Tellurium, present in anomalous amounts in two of the three mining districts, was not found to be present above the detection limit in Yucca Mountain samples. With the exception of a moderate arsenic-gold correlation, correlated elements in samples from Yucca Mountain fall into only one general category - a base-metals association. Two groupings of elements are present in samples from the Mother Lode deposit area and from the Rhyolite district. Correlated

elements at Wahmonie, however, were found only in the precious-metals grouping (silver-tellurium, gold-tellurium, and mercury-tellurium).

#### MINERALOGIC AND LITHOLOGIC RESULTS

Although most of the samples considered to be geochemically anomalous are in the Paintbrush bedded tuff sequence, these samples are mineralogically variable. Silicified samples, such as YMSC 22S, YMSC 31, YMSC 52, YMSC 88, and YMSC 92 contain opal and opal-CT as shown by X-ray diffraction (Appendix C). Opal-CT was defined by Jones and Segnit (1971) as opal that yields X-ray diffraction patterns containing some cristobalite and tridymite peaks in addition to the pattern for amorphous silica. Unaltered phenocrysts of plagioclase, potash feldspar, biotite, and clinopyroxene are present, but pumice lapilli and shards are thoroughly replaced by isotropic to faintly birefringent silica (Figure 7).

The tuff dike (YMSC 22), on the other hand, is composed of largely unaltered glass lapilli and shards with unaltered phenocrysts of feldspar, biotite, and clinopyroxene. Minor amounts of finely disseminated montmorillonite and carbonate are present (Figure 8). Sample YMSC 14C is almost completely composed of volcanic glass with little or no calcite, opal, or clay.

Sample YMSC 45A is of calcrete breccia in ash-flow tuff with little or no secondary silica. The calcrete matrix consists of finely divided calcite with rounded granules of dark brown material (Figure 9) and very minor amounts of late opal-CT. Cavities in the matrix contain acicular carbonate (possibly calcite pseudomorphs after aragonite). Similar mineralogies and textures were observed in calcrete samples not considered. to have anomalous chemistries (Figures 10 and 11).

Sample YMDD 36A is silicified breccia from a fault cutting ash-flow tuff. It consists mainly of opal-CT and tridymite based on XRD analysis. It also contains late veinlets of white opaline silica containing tridymite and chalcedony probably similar to that shown in Figure 7.

Sample YMDD 36A is macroscopically similar to many other silicified breccia samples taken along faults in the Yucca Mountain Addition (Figure 2 shows a good example of an outcrop of such material). Except for sample YMDD 36A, none of these samples were found to have anomalous chemistries. Descriptions of two thin sections of this type of rock serve to illustrate the<br>mineralogical and textural features of these breccias. Sample mineralogical and textural features of these breccias. YMSC 29 consists of clasts of devitrified ash-flow set in siliceous matrix with local calcite. In thin section, the matrix is seen to consist of wedges of tridymite coated with brown opal-CT (Figure 12) and later local calcite, chalcedony, and possible cristobalite. Sample YMSC 66 is similar, containing a matrix of brown opal-CT coated by chalcedony with local late calcite infilling (Figure 13).

Figure 7. Photomicrograph of sample YMSC 22S. Silicified air-fall tuff with shards and lapilli almost completely replaced by isotropic silica (on bottom) and cut by white tridymite-rich opaline silica vein (top) containing discontinuous layers of birefringent quartz. Field of view about 3 mm x 2 mm. Cross-polarized light.



Figure 8. Photomicrograph of sample YMSC 22. Pumice lapilli in glass shard matrix with some fine birefringent montmorillonite. Field of view about 3 mm<br>x 2 mm. Cross-polarized light. Cross-polarized light.



Figure 9. Photomicrograph of sample YMSC 45. Breccia with calcrete matrix and devitrified ash-flow tuff clasts. Note rounded dark granules. Field of view 3 mm x 2 mm. Plane-polarized light.



Figure 10. Photomicrograph of sample YMSC 14. Calcrete layer in bedded tuff containing pumice fragments and a partial spherulite in fine calcite with dark granules. Field of view is 3 mm x 2 mm. Plane-polarized light.



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Figure 11. Photomicrograph of sample YMSC 14. Cavity in calcrete containing calcite after aragonite(?). Horizontal field approx. 750 microns. Cross-polarized light.



### SUBSURFACE DATA

Although geologic and petrographic data are available for three drill holes in the Yucca Mountain Addition, very little trace element data have been published. In addition, most of the data available are not applicable to precious- and(or) base-metal exploration because elements specific for the present work are not included, or because detection limits are too high.

A relatively detailed lithologic log, as well as data on fracture fillings, for drill core to a depth of 1533 m from holes USW G-3 and USW GU-3 are available (Scott and Costellanos, 1984). Based on this log, alteration of units intercepted in these drill holes consists of zeolitic alteration of bedded tuffs at depths of 600 m or more, minor clay alteration below 770 m, more extensive clay alteration below 1100 m, and local disseminated sulfides below 1170 m. Fracture fillings consist of clay, silica, zeolite, carbonate minerals, iron oxides and hydroxides, and fluorite, but zones of severe fracturing have not been noted.

Within the Yucca Mountain Addition, we are aware of only two gold analyses, both below a 0.02 ppm detection limit, on two samples from drill hole USW G-3 (Broxton et al., 1986). A single antimony analysis was below the 0.5 ppm detection limit. The analyses of these samples do not include other precious- or basemetals, or pathfinder elements associated with them. Results of XRD analyses of fracture fillings indicate that most are predominantly composed of silica (quartz, cristobalite, and
Figure 12a. Photomicrograph of sample YMSC 29. Silicified breccia matrix. Wedges of tridymite with brown opal-CT coating and late chalcedony. Field of view abount 1 mm x 1.5 mm. Plane-polarized light.



Figure 12b. Photomicrograph of sample YMSC 29. Same view as in Figure 12a, but with cross-polarized light.



Figure 13a. Photomicrogaph of sample YMSC 66. Matrix of silicified breccia. Brown opal-CT coated with grey birefringent chalcedony. Late carbonate filling. Field of view 1.5 mm x 1 mm. Cross-polarized light.



Figure 13b. Photomicrograph of sample YMSC 66. Planepolarized view of Figure 13a.



tridymite) and(or) calcite. However, several fractures containing up to an estimated 90 percent fluorite were identified at depths greater than 249 m (Scott and Costellanos, 1984).

More than 80 samples from drill hole USW G-1, about 3 km north of the Yucca Mountain Addition, have been analyzed for gold. All were below detection limits; however, for these analyses the gold detection limit was 0.12 ppm or higher. Antimony analyses of these samples were also all below a 1.5 ppm or higher detection limit. The data include a few zinc values of up to 235 ppm (Broxton et al., 1986).

Core from drill hole USW G-2, about 5 km north of the Yucca Mountain Addition, contains minerals characteristic of hydrothermal alteration, such as clay minerals, chlorite, and pyrite at depths of 1000 m or more, as well as fluorite veins and a single thin barite-calcite-chlorite vein (Caporuscio et al.,<br>1982). In addition, gold analyses of samples from this hole In addition, gold analyses of samples from this hole include a value of 0.06 gold in zeolitized tuff from a depth of 515 m (Broxton et al., 1986). Two samples from shallower depths have no detectable gold (at a detection limit of 0.02 ppm).

Uraniferous opal fracture fillings up to 1 cm thick were noted in drillcore from hole USW G-3 within the Yucca Mountain Addition, but the highest uranium content measured is 35 ppm (Szabo and Kyser, 1985). Gamma log results indicate that drill holes within the Yucca Mountain Addition did not encounter anomalously radioactive rock (Muller and Kibler, 1985).

## COMPARISON WITH SURROUNDING PRECIOUS-METAL DISTRICTS

Areas with precious-metal deposits which were examined and sampled for comparative purposes are the Wahmonie district; the Mother Lode deposit; and the Bullfrog district, a widespread group of past and present producing mines including the Rhyolite area, the Original Bullfrog mine, and the Gold Bar mine (Figure 1). All of these areas contain precious-metal mineralization in southwestern Nevada volcanic field rocks that are contemporaneous, or nearly so, with rocks underlying the Yucca Mountain Addition.

### Wahmonie District

precious-metal mineralization in the Wahmonie district occurs in a system of N30°E veins within a similarly oriented 8 km by 4 km elliptical alteration halo containing strongly oxidized argillized and silicified rock. Mineralization ages determined on adularia from altered rock in the Wahmonie district are 12.6 and 12.9 Ma (Jackson, 1988). Vein samples collected from the Vein samples collected from the Wingfield shaft dump contain up to 50 oz silver and 0.67 oz gold per ton. They consist mainly of macrocrystalline quartz with alunite and gypsum; and carry free gold, cerargyrite, argentite, hessite, iron and manganese oxides, and sulfides (Quade and Tingley, 1984). Samples of altered rock and veins collected during this study within a 400-m radius of the Wingfield shaft (samples W-1 through W-12, Appendix 4) contain up to 46 ppm silver

and 0.44 ppm gold along with anomalous arsenic, bismuth, mercury, antimony, and tellurium. Gold content ranges between 0.001 and 0.202 ppm. The silver to gold ratio ranges from 5 to more than 200, and averages about 75.

#### Mother Lode Deposit Area

The Mother Lode deposit is located at the base of the northeast flank of Bare Mountain about 15 km northwest of the Yucca Mountain Addition (Figure 1) near the eastern end of the Fluorite Canyon fault. The deposit, which was located by drilling along an inferred eastward extension of this fault under cover, has not been put into production, but announced reserves are 3.7 million tons at 0.05 oz gold per ton. An additional 0.7 million tons of similar ore about 1000 meters southeast of the Mother Lode deposit were recently announced by GEXA. The deposit does not crop out, but mineralized rock occurs within 1 m of the present surface in a large cross-shaped trench excavated at the locus of the most near-surface part of the deposit determined by drilling.

Two types of mineralized rock, separated by a near vertical northeasterly fault, are exposed in the trench. Altered pumiceous rhyolitic lapilli air-fall tuff occurs west of the fault, and moreor-less altered fine-grained volcanic sandstone with siltstone and limy siltstone interbeds are found east of the fault. The tuff, which is correlative with less altered rock on a hill just west of the deposit, is probably part of a bedded tuff sequence that overlies the Paintbrush Tuff. Examination of the trench walls showed that mineralized rock was originally overlain by 25 cm to 2 m of gravel and(or) caliche.

Alteration mineralogy in both rock types exposed in the trench is dominated by quartz and illite. Jarosite is locally abundant, particularly in the tuff. Sparse, coarsely crystalline, drusy irregular quartz and manganese oxide veins are present, particularly in the siltstone. Manganese oxide and quartz are also present along some faults in the trench.

The trench at the Mother Lode deposit is in the hanging wall of the north-dipping Fluorspar Canyon fault. Samples of limonitized Paleozoic rock and Tertiary dike rock were collected from dumps at old workings south of the fault in the Telluride mining district. The fault is the locus of chalcedonic silica The fault is the locus of chalcedonic silica deposition in the Mother Lode area, and is also marked by silicification 3 km to the southwest in Fluorite Canyon. Silicified and alunitized samples from the fault and the Paleozoic rocks south of the fault were dated at 12.2 Ma and 11.2 Ma, respectively (Jackson, 1988).

Thirty-five samples of mineralized and unmineralized rock were collected in the vicinity of the Mother Lode deposit (GEXA 1 through 35), and one sample of silicified breccia (FC-1) was taken from Fluorite Canyon. Gold content of these samples ranges from 0.001 ppm to more than 7 ppm; and elevated silver, arsenic, mercury, antimony, and tellurium contents are present in samples with high gold. The average silver to gold ratio is 0.8.

## Rhyglite Area

The Rhyolite area, which lies east of Beatty in the Bullfrog district about 30 km northwest of the Yucca Mountain Addition (Figure 1), contains the largest and highest grade known gold reserves of any area in the southwestern Nevada volcanic field. The bulk of the reserves are in Bond Gold Inc.'s Bond-Bullfrog mine, where a moderately westward-dipping orebody has an overall grade of 0.11 oz gold and 0.24 oz silver per ton. The orebody consists of a central core of silica-flooded ore overlain and underlain by quartz vein stockwork ore (Jorgensen et al., 1989). The ore occurs in rhyolitic tuffs probably equivalent to the Timber Mountain Tuff (Cornwall and Kleinhampl, 1964; and Byers et al., 1976). Vein gangue minerals are cryptocrystalline to coarsely crystalline quartz, calcite, and adularia. Limonite, fluorite, barite, pyrite, argentite, and base-metal sulfides are also present (Jorgensen et al., 1989). Production from the Bond-Bullfrog pit is slated to begin in September 1989 at a planned rate of 200,000 oz gold per year.

Bond Gold Inc. also holds reserves of 3.1 million tons of ore at the Montgomery-Shoshone mine, which produced most of the gold in the Bullfrog district in the past, The Montgomery-Shoshone ore is similar to that at Bond-Bullfrog, which lies about 1.5 km to the southwest, and occurs in the same lithologic units (Jorgensen et al., 1989). Adularia from the Montgomery-Shoshone Mine was dated at 9.5 Ma (Morton et al., 1977).

Although we were unable to collect samples from the Bond-Bullfrog pit, samples were taken on Ladd Mountain west of the pit (BH 21 through BH 25, Appendix 1) and from a roadcut northeast of the pit (BH 13 through BH 14, Appendix 1). In addition, samples were collected from a glory hole at the Montgomery-Shoshone mine (MS-1 through MS-7, Appendix 1), from the National Bank mine area 1 km northwest of the Bond-Bullfrog pit (BH 2 through BH 8, Appendix 1); and from the Tramps mine area 2 km northwest of the Bond-Bullfrog pit (BH 9 through 12, and BH 26 through 29V, Appendix 1). All samples were collected from rocks considered correlative with the Paintbrush and Timber Mountain Tuffs (Cornwall and Kleinhampl, 1964; and Byers et al., 1976). The vein samples consist mostly of chalcedonic to coarsely crystalline drusy quartz with dark gray leached calcite and clear to white adularia. Iron and manganese oxides occur in many samples, and free gold was noted at the Tramps Mine. Several samples of silicified breccia were collected. Country rock adjacent to veins is generally silicified and bleached ash-flow tuff. Basalt dikes and(or) flows occur at the Bond-Bullfrog, National Bank, and Montgomery-Shoshone deposits, and are mineralized and altered locally. Montmorillonite-illite was identified in a mass of argillized rock in the Montgomery-Shoshone glory hole.

Analyses of 34 samples collected from the Rhyolite area yielded gold values of up to 13.9 ppm, along with high silver. Many samples are slightly anomalous in molybdenum, particularly those from the Montgomery-Shoshone Mine. A few samples have anomalous arsenic, mercury, and antimony, but Rhyolite area

samples are relatively low in these elements compared with those from Wahmonie and the Mother Lode deposit. Only a single sample from the Rhyolite area contains high tellurium.

## Original Bullfrog Mine

The Original Bullfrog mine is in the southwest corner of the Bullfrog district about 30 km west-northwest of the Yucca Mountain Addition (Figure 1). The ore consists mostly of a mass of nearly solid quartz some 20 m thick dipping about 20° north, overlain by rhyolite with abundant quartz stringers and underlain by sheared shaly Paleozoic rock (Ransome et al., 1910). Quartz veins do not extend into the underlying shale or associated limestone. A nearly horizontal zone of bleached and intensely sheared rock occurs along a roadcut below the main lode and appears to separate quartz-veined rhyolite in the hanging wall from dark-green shale in the footwall. The quartz vein material consists mostly of coarsely crystalline quartz, with some calcite, malachite, and chrysocolla. Native gold occurs as visible particles with limonite associated with chrysocolla.

Adularia from ash-flow tuff altered to alkali feldspar and montmorillonite and cut by quartz-adularia veinlets was dated at 8.7 Ma (Jackson, 1988). The age of the ash-flow tuff that hosts the Original Bullfrog lode is not known, but it underlies the Bullfrog Member of the Crater Flat Tuff which is considered to be 14.0 Ma (Byers et al.,1976).

Although only five samples were collected from the Original Bullfrog mine (BH 15 through BH 20), results of chemical analyses are sufficiently different from the rest of the Bullfrog district to show that this deposit merits separate discussion. Gold values in Original Bullfrog samples range between 0.02 and 117 ppm, and the silver to gold ratio averages about 12. Other anomalous trace elements are bismuth, copper, and antimony. Molybdenum is anomalous, although at low levels, and single samples were anomalous in cadmium, lead, tellurium, or thallium.

## Gold Bar Mine

The Gold Bar mine is a producing silver-gold deposit located about 30 km northwest of the Yucca Mountain Addition in the northwest corner of the Bullfrog district (Figure 1). The orebody consists of a northeasterly zone of brecciated rhyolite and basalt about 300 m by 30 m in plan which dips  $65^{\circ}$ NW (Ransome et al., 1910). The age of the mineralization and host rock is not known.

Five samples were collected from a muck pile in the Gold Bar pit (GB 1 through GB 5, Appendix 1), but none were found to be of ore-grade material. Gold contents range between 0.006 ppm in unmineralized ash-flow tuff and 0.08 ppm in basalt cut by calcitequartz veins. The average silver to gold ratio is 11, and other trace element contents are low, with the exception of slightly anomalous antimony. Veins consist of macrocrystalline calcite and drusy quartz. Electrum and gold-bearing pyrite are present (Ransome et al., 1910).

# Comparison of the Yucca Mountain Addition with Precious-Metal Districts

Based on ages and correlations discussed above, lithologic units exposed in the Yucca Mountain Addition are older than mineralization in all of the surrounding precious-metal districts, and rocks correlative with Yucca Mountain Addition units contain silver-gold deposits in the Rhyolite area. Precious-metal mineralization in the Yucca Mountain Addition is, therefore, -permissive within the constraints of timing and host-rock lithology.

Jackson (1988) suggested that silver-gold mineralization in the southwestern Nevada volcanic field is related to a single episode of widespread hydrothermal activity associated with postcollapse volcanic activity following eruptions from the Timber Mountain caldera. The Yucca Mountain Addition falls within this area of hydrothermal activity. However, based on geochemistry and fluid inclusion data, the Bare Mountain district deposits, including the Mother Lode deposit, are thought to be a near surface expression of a porphyry molybdenum system not clearly related to caldera activity (Noble and Weiss, 1989).

According to Jorgensen et al. (1989), geologic data suggest that silver-gold mineralization in the Bullfrog district formed from hydrothermal solutions migrating along the Bullfrog detachment fault. Several writers (including Hamilton, 1988) have suggested that faults bounding Bare Mountain, including the Fluorite Canyon fault, are the eastward extension of the Bullfrog detachment. The Mother Lode deposit and other gold-silver mineralization in the Bare Mountain area could also be considered to be related to fluid migration along a detachment fault. If, as some believe (e.g., Hamilton,1988), the Bullfrog-Bare Mountain detachment surface extends eastward under Yucca Mountain, precious-metal mineralization could also occur there at depth.

The eastern caldera wall fractures of the proposed Crater Flat-Prospector Pass caldera system may lie beneath Yucca Mountain (Carr, 1984). If this is so, precious-metal mineralization related to post-collapse hydrothermal activity may occur at depth in the Yucca Mountain Addition.

Results of surface examination and sampling in the Yucca Mountain Addition have delineated some important differences between it and the surrounding precious-metal districts. Based on the general absence of hydrothermal alteration, we believe that the data from drill holes G-3 and GU-3 indicate that geology similar to that exposed on the surface extends to a depth of at least 600 m beneath the crest of Yucca Mountain.

The Yucca Mountain Addition is underlain by rock types that are mineralized in the surrounding precious-metal districts. However, exposures of Tertiary intrusive igneous rocks and pre-Tertiary rocks that occur in the Wahmonie and Bullfrog districts, and in the Mother Lode deposit area, do not occur in the Yucca Mountain Addition.

Direct surface observations indicate that no areas of hydrothermal alteration similar to those in the Wahmonie and Bullfrog districts, or at the Mother Lode deposit, occur within the Yucca Mountain Addition. No significant areas of strongly bleached, limonitized, argillized, or silicified rock are present in the Yucca Mountain Addition.

The silicified and(or) calcrete fault breccias, silica and(or) carbonate veins, and small amounts of altered tuff encountered in the Yucca Mountain Addition do not carry anomalous gold, and only a few samples contain slightly anomalous silver, lead, and bismuth. In comparison, veins, breccias, and altered rock from areas with known precious-metal mineralization have moderately to highly anomalous gold and silver, and generally contain anomalous amounts of associated trace metals such as arsenic, mercury, and antimony. In addition, correlations among elements in samples from the Yucca Mountain Addition show little resemblance to elemental correlations from precious-metal districts.

The silica and(or) carbonate veins and breccias which commonly occur along faults in the Yucca Mountain Addition are mineralogically distinct from veins and breccias in the preciousmetal districts examined. The siliceous component in these rocks in the Yucca Mountain Addition is mainly opal-CT, tridymite and(or) cristobalite with only minor chalcedonic quartz. Vein and breccia silica in the precious-metal districts is chalcedonic to coarsely crystalline quartz with little or no opal. Carbonate in Yucca Mountain Addition veins and breccias is very fine-grained calcrete, whereas that in the precious-metal districts is coarsely crystalline. Silica veins and breccias in the Yucca Mountain Addition do not contain abundant manganese oxide or limonite, but some contain abundant hematite. By contrast, limonite and manganese oxide are abundant in veins in the precious-metal districts.

### REMOTE SENSING ANALYSIS

Remote sensing methods were used to compare fault patterns, lineament patterns, and alteration in the Yucca Mountain Addition to those in the Wahmonie mining district, the Calico Hills area, the Mother Lode deposit area, and the Rhyolite area.

### REMOTE SENSING METHODOLOGY

#### Geologic Structure and tineaments

Six visible and near infrared bands of Landsat Thematic Mapper (TM) digital imagery and one band of SPOT panchromatic satellite imagery, with resolutions of 30 m and 10 m, respectively, were analyzed on a Terramar MicroImage computer image processing system. Structural features and lineaments were

enhanced on a single band of TM imagery and on the SPOT imagery by applying a high-pass filter. The high pass filter performs an edge enhancement and emphasizes linear features by exaggerating differences in brightness values between adjacent picture elements. Directional filters were also applied to both TM and SPOT imagery to emphasize lineaments with north-south, east-west, northeast-southwest, and northwest-southeast orientations. Lineaments were drawn manually on the edge-enhanced images. Because the number of lineaments drawn in each area was too small to show a statistically significant pattern, rose diagrams were not constructed.

The lineaments enhanced and recognized on SPOT and Landsat TM imagery are linear topographic and tonal features and are not unequivocally related to geologic structure. Linear cultural features such as roads and vegetation boundaries are also enhanced by applying filters to satellite imagery. Linear features should be interpreted with care since linear artifacts parallel to the sensor scan line direction may be emphasized, directional exaggerations may be introduced, and spurious linears generated from electronic "ringing" (a shadow or halo around sharp edges) may be produced. None of the linear features recognized on the satellite imagery during this investigation have been checked in the field. Until field checking or drilling data confirm the presence of faults, these lineaments cannot be assumed to be geologic structures.

The direction and length of faults of more than 0.3 km in length were measured on geologic maps of the Yucca Mountain Addition and in surrounding precious-metal mining districts and mineralized areas. For each area, the results were tabulated and plotted on rose diagrams to determine the most prominent fault orientations, and the circular variance was calculated. Lineaments detected on the TM and SPOT imagery were then compared with measured fault orientations in the Yucca Mountain Addition and the surrounding mining districts.

## Rock Alteration

Landsat Thematic Mapper imagery of the Yucca Mountain Addition and surrounding mineralized areas was processed to locate and enhance rock alteration which may be associated with mineralization. Distinctive spectral features which have been observed in the visible and near-infrared part of the spectrum on laboratory reflectance curves of pure minerals and which can be identified using Landsat TM bandpasses consist of ferrous and ferric iron, the hydroxyl ion, bound water, and the carbonate and sulfate ions. Even in low concentrations, iron and hydroxyl or water absorption features often dominate rock spectra, since many rock-forming minerals (e.g., quartz and feldspar) do not have distinctive spectra in the visible and near infrared part of the spectrum.

Three-band color ratio composite images of the Yucca Mountain Addition and surrounding mining districts were produced to show areas of argillic and iron oxide alteration. The ratios make use

of spectral absorption features characteristic of iron minerals near 450 nm and between 750 and 900 nm and absorption features near 2200 nm which are characteristic of clay minerals. The distinctive spectral characteristics of clay minerals causes areas with clay to appear bright in a TM band 5/TM band 7 ratio image while iron oxide minerals appear bright in a TM band 3/TM band 1 ratio image. Because of spectral features associated with vegetation, vegetation appears bright on a TM band 4/TM band 3 image.

The vegetation, iron oxide, and clay ratio images were combined to form a color composite image where TM4/TM3 was encoded blue, TM3/TM1 was encoded green, and TM5/TM7 was encoded red. Areas with iron oxide alteration appear green; areas with argillic alteration appear red; and areas containing both argillic and iron oxide alteration appear yellow on the imagery.

After the Landsat TM imagery was processed, selected areas that showed evidence of argillic and iron oxide alteration were checked in the field and samples were collected for X-ray diffraction analysis. Reflectance spectra in the visible and near infrared were also recorded in the field and in the laboratory with a Geophysical Environmental Research, Inc. field spectroradiometer.

# RESULTS FOR YUCCA MOUNTAIN ADDITION

### Faults and Lineaments

The lengths and orientations of faults longer than 0.3 km mapped by Lipman and McKay (1965) in the Yucca Mountain Addition were measured on the 1:24,000-scale map, tabulated, and plotted in **100** increments on rose diagrams as a function of fault frequency (Figure 14) and cumulative length of faults (Figure 15). Similar plots (Figures 16 and 17) were constructed for faults mapped by Scott and Bonk (1984) in the Yucca Mountain Addition (1:12,000 scale map). The orientations of mineralized veins or fractures from which samples were collected and analyzed as part of this investigation, where recorded, were also plotted (Figure 18). For each of the rose diagrams, the total number (n) or cumulative length in kilometers (1) of faults, veins, and fractures is given as well as the circular variance (cv). The circular variance is a measure of the dispersion of the data and varies between 0 and 1. Small values indicate that fault orientations are tightly grouped, and large values indicate a dispersed group of orientations.

In the Yucca Mountain Addition, the maximum number of faults or cumulative fault length, in all five cases, was between northsouth and N10°E or north-south and N10°W on the rose diagrams. The orientations with the greatest number of faults also approximately coincide with the orientations of the greatest cumulative length. Mean directions were between  $NS^{o}W$  and  $NI^{o}W$ , or at  $NS^{o}E$ . Circular variance ranged from 0.18 to 0.34. This at N3°E. Circular variance ranged from 0.18 to 0.34. indicates that many of the faults at the surface in the Yucca Mountain Addition are subparallel and fault intersections would be Figure 14. Yucca Mountain Addition normalized fault orientation frequency, as measured on the geologic map by Lipman and McKay (1965). The total number of faults measured (0.3 km or longer) is indicated by "n"; "cv" is the circular variance which can vary between 0 and 1.

N

 $n = 48$  $cv = 0.22$  Figure 15. Yucca Mountain Addition orientation of normalized cumulative fault length frequency for faults longer than 0.3 km, as measured on the geologic map by Lipman and McKay (1965). Total fault length measured in kilometers is indicated by "1"; "cv" is the circular variance which can vary between 0 and 1.



 $1 = 49.6$  $cv = 0.18$ 

Figure 16. Yucca Mountain Addition normalized fault orientation frequency for faults longer than 0.3 km, as measured on the geologic map by Scott and Bonk (1984). The total number of faults measured is indicated by "n"; "cv" is the circular variance which can vary between 0 and 1.

 $n = 69$ <br>cv = 0.19



Figure 17. Yucca Mountain Addition orientation of normalized cumulative fault length frequency for faults longer than 0.3 km, as measured on the geologic map by Scott and Show that the secretary and the secretary of the bond fault length measured in kilometers is indicated by "l"; "cv" is the circular variance which can vary between 0 and 1.

 $1 = 43.2$  $cv = 0.19$ 



Figure 18. Orientation of normalized frequency of mineralized veins and faults sampled in this investigation in the Yucca Mountain Addition. The total number of veins and faults of known orientation is indicated by "n"; "cv" is the circular variance which can vary between 0 and 1.

N

 $n = 69$  $cv = 0.34$  less likely to occur than in areas with a greater range of fault orientations.

SPOT digital data of the Yucca Mountain Addition were processed with a high-pass filter (Figure 19) and lineaments were drawn on the resulting image. Directionally filtered images were also produced and lineaments recognized on these (i.e., Figure 20) were transferred to a non-directional biased filtered image with a percentage of the original image added back to the filtered image. Figure 20 was enhanced for northwest-southeast lineaments, but shows lineaments in many other orientations as well. Lineaments drawn on Figure 19 are indicated on Figure 20 with arrows.

Many of the lineaments coincide with mapped faults, and some extend the faults into the alluvium. The northern part of the Solitario Canyon fault does not show up well, but a significant<br>lineament follows the wash in Solitario Canyon. Three east-west lineament follows the wash in Solitario Canyon. lineaments that do not coincide with mapped faults have been recognized, the southernmost one being the most prominent (Figures 19 and 20). If these relatively long lineaments are faults, the If these relatively long lineaments are faults, the number of fault intersections would be higher in the Yucca Mountain Addition than is indicated by geologic mapping. However, none of the lineaments drawn on the images were checked in the field to establish the presence or absence of faults.

### Alteration

A Landsat TM color ratio composite was produced from TM4/TM3, TM3/TM1, and TM5/TM7 as blue, green, and red, respectively (Figure 21). The ratio image produced from the Landsat. TM imagery is at a smaller scale and lower resolution than the SPOT lineament images. Areas in the Yucca Mountain Addition that are altered show up as yellow, indicating high clay and iron oxide content, and-consist primarily of bedded tuffs in or stratigraphically above the Paintbrush Tuff (bt and pbt on Plate 1).

These bedded tuffs were sampled in the field and reflectance spectra were recorded with a field spectroradiometer. The bedded tuffs are iron stained and silicified in places and, based on Xray diffraction analysis (Appendix C), contain primarily volcanic glass, with some calcite, opal, cristobalite, and montmorillonite. Spectral curves recorded with the spectroradiometer indicated clay absorption features in the bedded tuffs which were absent in the underlying upper unit of the Topopah Spring Member (Figures 22 and  $23$ .

# RESULTS FOR PRECIOUS-METAL DISTRICTS AND MINERALIZED AREAS

### Wahmonie Distriet

## Faults and Lineaments

Mapped fault frequency, length, and orientation were measured on the 1:24,000 geologic map of the Skull Mountain quadrangle (Ekren and Sargent, 1965). The rose diagram of fault frequency (Figure 24) shows that the greatest number of faults are oriented

Figure 19. Lineaments drawn on an enhanced SPOT panchromatic image of the Yucca Mountain Addition. The Yucca Mountain Addition is outlined in red. Lineaments have not been checked in the field to establish the presence or absence of faulting.



Figure 20. Lineaments shown in Figure 19 are indicated with arrows on this SPOT panchromatic image enhanced for northwest lineaments. The Yucca Mountain Addition is outlined in red. This image covers the same area as Figure 19. Lineaments have not been checked in th field to establish the presence or absence of faulting.



Figure 21. Landsat TM color ratio composite of the Yucca Mountain Addition. The Yucca Mountain Addition is outlined in red. Red on the image indicates areas with argillic alteration, green indicates areas with iron oxide alteration, and yellow areas on the image have been affected by both clay and iron oxide alteration. Yellow areas with altered bedded tuff are outlined in black.



Figure 22. Spectral curve of altered bedded tuff, sample YMSF3 in the Paintbrush Tuff, showing a clay absorption feature near 2200 nm.



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Figure 23. Relatively featureless spectral curve of the unaltered top of the Topopah Spring Member of the Paintbrush Tuff, sample YMSFll.



Figure 24. Normalized fault orientation frequency for faults Le 24: Normanized radio offendation rrequency for radio Ekren and Sargent (1965) in the Wahmonie minin district. The total number of faults measured i indicated by "n"; "cv" is the circular variance which can vary between 0 and 1.



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between north-south and N10°W. A secondary frequency maximum is located between N30°E and N50°E. The Wahmonie district contains a large range of fault orientations in comparison to the Yucca Mountain Addition.

A rose diagram of orientations of cumulative mapped fault length for the Wahmonie district shows that the orientation having the greatest cumulative fault length is between N30°E and N40°E and the circular variance is 0.64 (Figure 25). This coincides with the secondary maximum for fault frequency on Figure 24. Quade and Tingley (1984) noted that the mineralized vein system in the Wahmonie mining district follows a structure that is oriented  $N30$ <sup>°</sup>E.

Since the eastern border of the SPOT satellite image is slightly west of Wahmonie, Landsat TM imagery of the Wahmonie district was processed to enhance lineaments (Figures 26 and 27). The Landsat TM image of the Wahmonie area is at a smaller scale than lineament images of the other mineralized areas and the Yucca Mountain Addition. Within the area of alteration mapped by Ekren and Sargent, which has been outlined on Figure 26, lineaments of diverse orientations were noted. Three long lineaments are the most prominent. The first lineament is oriented north-northeast and curves northwest at its north extent; it coincides with the eastern limit of Tertiary intrusive rocks in the Wahmonie area. Most of the shafts, adits, and prospect pits in the district are located near the southern extent of this lineament. The second lineament is oriented north-south, extends across most of the Landsat TM image, and intersects the first lineament in the area of abundant mine workings. The third lineament extends northwest across the area of alteration and intersect the other two. It coincides with mapped faults which have not been joined together. Other near east-west lineaments have also been noted (Figures 26 and 27). The Wahmonie lineaments have not been checked in the field for correlation with geologic structure.

## Alteration

The Landsat TM color ratio composite of the Wahmonie mining district shows an east-northeast-trending zone of argillic and iron oxide alteration (Figure 28). Locations with high concentrations of iron oxide are displayed in green; areas with argillic alteration are shown in red; and areas with both iron oxide and argillic alteration are colored yellow. The alteration shown in Figure 28 includes the alteration mapped by Ekren and Sargent (1965), but provides more information on the specific alteration types.

Illite and montmorillonite were found during this investigation in altered volcanic rocks near the Wingfield shaft by X-ray diffraction analysis (Appendix C). Surface alteration at Wahmonie is not confined to one stratigraphic unit as it is in the Yucca Mountain Addition.

Hydrothermally altered rock units include andesites, dacites, latites, and tuffs of Wahmonie Flats. Granodiorite, andesite, and rhyolite intrude the tuffs and flow rocks.

Figure 25. Orientation of normalized cumulative fault length frequency for faults longer than 0.3 km, as measured on the geologic map by Ekren and Sargent (1965) in the Wahmonie mining district. The total fault length measured in kilometers is indicated by "1"; "cv" is the circular variance which can vary between 0 and 1.

 $1 = 108.8$ .  $cv = 0.61$ 



Figure 26. Lineaments drawn on the enhanced Landsat TM filtered image of the Wahmonie mining district. The altered area has been outlined in red. This lineament image is at a smaller scale and lower resolution than those for other areas. Lineaments have not been checked in the field to establish the presence or absence of faulting.



Figure 27. Lineaments shown in Figure 26 are indicated with arrows on this Landsat TM image of the Wahmonie mining district which has been enhanced for northwest lineaments. The altered area has been outlined in red. This lineament image is at a smaller scale and lower resolution than those for other areas. Lineaments have not been checked in the field to establish the presence or absence of faulting.



Figure 28. Landsat TM color ratio composite of the Wahmonie mining district. Red on the image indicates areas with argillic alteration, green indicates areas with iron oxide alteration, and yellow areas on the image have been affected by both clay and iron oxide alteration.



# Calico Hills

## Faults and Lineaments

Fault orientation and length in the altered rocks of the Calico Hills were measured on the 1:24,000 geologic maps of the Topopah Spring and Jackass Flats quadrangles (Orkild and O'Connor, 1970; McKay and Williams, 1964). The fault-frequency maximum on the rose diagram occurs between N20°W and N30°W. Large numbers of faults occurs within 20° on both sides of this maximum and a secondary maximum is located at between N30°E and N40°W. The circular variance is 0.77. The wide distribution of faults of all orientations is apparent on Figure 29.

The greatest cumulative fault length is oriented between north-south and N10 $\textdegree{W}$  (Figure 30). This is 20 $\textdegree{}$  east of the fault frequency maximum. A secondary maximum is located between N40°E and N50\*E. The circular variance is 0.80 and there is a wide range of fault directions as there was in Figure 29.

SPOT imagery was used to enhance lineament patterns in the Calico Hills area. The lineaments drawn on the filtered SPOT images were primarily northeast lineaments, which coincided with mapped faults along part of their length (Figures 31 and 32). The lineaments in the Calico Hills area have not been field checked to confirm the presence or absence of faults.

## Alteration

The Calico Hills, which lie east of Yucca Mountain, are the closest exposed area of widespread hydrothermal alteration to the Yucca Mountain Addition. Altered rocks include the Topopah Spring Member of what McKay and Williams (1964) mapped as the Piapi~ Canyon Formation and of what Orkild and O'Connor (1970) mapped as the Paintbrush Tuff, as well as older rhyolite flows, tuffaceous beds, and intrusions of the Calico Hills. The area of altered tuffs partially surrounds Carboniferous Eleana Formation and Devonian limestone and dolomite intruded by rhyolite.

The Landsat TM color ratio composite of the Calico Hills area (Figure 33) shows a semicircle of alteration around the east, west, and south margins of the Paleozoic exposures. Argillic alteration (red), is interspersed with argillic and iron oxide alteration. (yellow); some patchy iron oxide alteration (green) is located primarily east and west of the Paleozoic rocks. As in the Wahmonie mining district, and in contrast to the Yucca Mountain Addition, the alteration is not confined to bedded tuff or to a single geologic unit.

#### Mother Lode Deposit Area

Faults and Lineaments

GEXA's Mother Lode gold deposit is located beneath less than 1 m of alluvium at the north end of Bare Mountain. The host rocks are an unnamed Tertiary tuff and a sedimentary unit of questionable age.

Figure 29. Normalized fault orientation frequency for faults longer than 0.3 km, as measured on the geologic map by Orkild and O'Connor (1970) and McKay and Williams (1964) in the Calico Hills. The total number of faults measured is indicated by "n"; "cv" is the circular variance which can vary between 0 and 1.



Figure 30. Orientation of normalized cumulative fault length frequency for faults longer than 0.3 km, as measured on geologic map by Orkild and O'Connor (1970) and McKay and Williams (1964) in the Calico Hills. The total fault length measured in kilometers is indicated by "1"; "cv" is the circular variance which can vary between 0 and 1.

 $1 = 152.0$  $cv = 0.80$ 



Figure 31. Lineaments drawn on an enhanced SPOT panchromatic image of the Calico Hills. Lineaments have not been checked in the field to establish the presence or absence of faulting.



Figure 32. Lineaments shown in Figure 31 are indicated with arrows on this SPOT panchromatic image of the Calico Hills which has been enhanced for northeast lineaments. This image covers the same area as Figure 31. Lineaments have not been checked in the field to establish the presence or absence of faulting.



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Figure 33. Landsat TM color ratio composite of the Calico Hills. Red on the image indicates areas with argillic alteration, green indicates areas with iron oxide alteration, and yellow areas on the image have been affected by both clay and iron oxide alteration. The area covered by the SPOT lineament images is outlined in red.



The most detailed published map of the area is at a scale of 1:62,500 (Cornwall and Kleinhampl, 1961). Mapped faults in the vicinity of the deposit are either confined to Paleozoic rocks or are concentrated west and north of the deposit. Because few faults were mapped in the vicinity of the Mother Lode deposit, no analysis of fault orientations was performed.

On the SPOT image enhanced for lineaments, the most prominent lineament extends northeast across the image (Figures 34 and 35). This lineament coincides, in part, with a 4-km-long fault mapped by Cornwall and Kleinhampl (1961) in volcanic rocks northwest of the Mother Lode deposit. The Mother Lode deposit area has not been field checked to establish if there is a correlation between the lineaments shown on Figures 34 and 35 and geologic structure.

#### Alteration

Hydrothermal alteration is not evident on the Landsat TM color ratio composite in the vicinity of the Mother Lode deposit (Figure 36). This is expected since the deposit is covered by alluvium. However, a bright spot appears just west of the deposit, suggesting the presence of altered tuff. Cornwall and Kleinhampl (1961) have mapped a northeast-trending dike in the Paleozoic rocks, which if extended into the alluvium, would pass through or near the Mother Lode deposit.

Mineralized samples collected from a trench in the deposit. contain quartz, opal, illite, and montmorillonite. Unmineralized altered tuff samples from a hill just west of the deposit contain alunite, kaolinite, and montmorillonite (Appendix 4).

## Rhyolite Area. Bullfrog Mining District

### Faults and Lineaments

Fault orientations and lengths were measured on the 1:48,000 map of the Bullfrog quadrangle produced by Cornwall and Kleinhampl (1964). Rose diagrams of both fault frequency and cumulative fault length showed a maxima at  $N10^{\circ}E$  to  $N20^{\circ}E$  and a large angular distribution with circular variances of 0.79 and 0.76, respectively (Figures 37 and 38).

Prominent lineaments on a filtered SPOT image of the Rhyolite area are oriented primarily northeast and northwest (Figures 39 and 40). The Bullfrog detachment fault, an east-west-trending feature cited by Jorgensen et al. (1989) and previous workers, extends across the bottom of the image beneath the alluvium. It does not appear as a lineament on the SPOT imagery. The lineaments in the Rhyolite area have not been checked in the field for correlation with geologic structure.

### Alteration

According to Jorgensen et al. (1989), mineralization in the Bullfrog district is confined to a 3-km-wide zone along and north of the Bullfrog detachment fault. This mineralized zone is included in the lower half of the Landsat TM alteration image in Figure 41. Mineralization in the district occurs in fault-veins,

Figure 34. Lineaments drawn on the enhanced SPOT panchromatic image in the vicinity of the Mother Lode deposit. Lineaments have not been checked in the field to establish the presence or absence of faulting.



Figure 35. Lineaments shown in Figure 34 are indicated with arrows on this SPOT panchromatic image of the Mother Lode deposit enhanced for northwest lineaments. Lineaments have not been checked in the field to establish the presence or absence of faulting.


Figure 36. Landsat TM color ratio composite in the area of the Mother Lode deposit. Red on the image indicates areas with argillic alteration, green indicates areas with iron oxide alteration, and yellow areas on the<br>image have been affected by both clay and iron oxide alteration. The area covered by the SPOT lineament images is outlined in red.



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Figure 37. Normalized fault orientation frequency for faults longer than 0.3 km, as measured on the geologic map by Cornwall and Kleinhampl (1964) in the Bullfrog mining district. The total number of faults measured is indicated by "n"; "cv" is the circular variance which can vary between 0 and 1.



 $n = 262$  $cv = 0.79$ 

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Figure 38. Orientation of normalized cumulative fault length frequency for faults longer than 0.3 km, as measured on the geologic map by Cornwall and Kleinhampl (1964) in the Bullfrog mining district. The total fault length measured in kilometers is indicated by "1"; "cv" is the circular variance which can vary between 0 and 1.

 $1 = 206.4$  $cv = 0.76$ 



Figure 39. Lineaments drawn on the enhanced SPOT panchromatic image of the Bullfrog mining district. Lineaments have not been checked in the field to establish the presence or absence of faulting.



Figure 40. Lineaments shown in Figure 39 are indicated with arrows on this SPOT panchromatic image of the Bullfrog mining district enhanced for northeast lineaments. This image covers the same area as Figure 30. Lineaments. have not been checked in the field to establish the presence or absence of faulting.



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Figure 41. Landsat TM color ratio composite of the Bull mining district. Red on the image indicate  $\frac{1}{2}$  and  $\frac{1}{2}$  areas with  $\frac$ argillic alteration, green indicates areas argillic alteration, green indicates areas with iron oxide alteration, and yellow areas on the image have been affected by both clay and iron oxide alteration.<br>The area covered by the SPOT lineament images is<br>outlined in red. The area covered by the SPOT lineament images is



veinlets, and stockworks (Jorgensen et al., 1989). Shallow intrusives have been mapped in the' area by Ransome et al. (1910) and Cornwall and Kleinhampl (1961).

On the Landsat TM alteration image (Figure 41), argillic and iron oxide alteration do not appear widespread in the Rhyolite area, although the rocks are light-colored and bleached (Figure 39). The alteration zone and mineralization at Bond Gold Inc.'s Bullfrog mine was below the alluvium and did not crop out extensively'at the surface. Although argillic alteration is associated with precious-metal mineralization in the Rhyolite district, silicification and potassic alteration, which are not detectable by the remote sensing techniques used during this investigation, are probably more important (Jorgensen et al.,1989.; and Jackson, 1988).

## SUMMARY OF REMOTE SENSING ANALYSIS

Rose diagrams and circular variance calculations for fault orientation and cumulative fault length in the Yucca Mountain Addition showed significant differences from those for surrounding precious-metal districts. Faults in the Yucca Mountain Addition are tightly grouped around a near north-south orientation and the circular variance is low (between 0.18 and 0.34). This indicates that many of the faults exposed at the surface in the Yucca Mountain Addition are subparallel and, therefore, fault intersections are less likely than in areas with greater circular variance. Fault intersections relate to structural preparation, or the lack or it, for mineralization.

In the surrounding precious-metal districts, circular variance of fault orientations ranged between 0.61 and 0.80. This is much greater variance that in the Yucca Mountain Addition, and suggests greater opportunity for the existence of fault intersections and structural preparation favorable for mineralization.

The lineament images, which were prepared from filtered digital SPOT and Landsat TM imagery, showed extensions of previously mapped faults and indicated additional lineament trends that may have structural significance. Structural significance of lineaments, however, needs to be established from surface field checking or drill hole data.

Alteration images prepared from Landsat TM satellite data showed rock alteration patterns in the Yucca Mountain Addition and in the nearby mining districts and mineralized areas. In the Yucca Mountain Addition, alteration is confined to two bedded tuff units. In the mining districts, alteration crosses many lithologic units.

## APPRAISAL OF MINERAL RESOURCES

## BASE-METALS RESOURCES

The volcanic rock section that underlies Yucca Mountain is unfavorable for the development of base-metals resources. Our surface geochemical sampling of faults and fracture zones does, however, show weak correlations between several base-metal elements and we could be sampling a very distal geochemical halo around a concentration of base metals in the deep subsurface. The closest deposits of base metals, at Bare Mountain about 15 km to the west, and at Mine Mountain about 40 km to the northeast, are found in Paleozoic carbonate rocks. If these rocks are present at depth beneath Tertiary volcanic rocks at Yucca Mountain, they are probably at depths in excess of 1,500 m. Base metal deposits of the types expected to be present in Paleozoic carbonate rocks include polymetallic replacement deposits or, if intrusive rocks are also present, tungsten, copper, and lead-zinc skarn deposits (Cox and others, 1989; Cox and Singer, 1986). Under present economic conditions, these types of deposits, if present at depths around the minimum estimate of 1,500 m, would not be mineable using standard, accepted mining technology. Base-metals potential within the Yucca Mountain Addition is rated very low.

The types of volcanic rocks underlying the Yucca Mountain Addition are favorable for mercury deposits. However, our sampling indicates that mercury is not present in the area at levels above the detection level of our analytical procedures.

## PRECIOUS-METALS RESOURCES

The absence of subsurface trace element geochemical data applicable to precious-metals exploration renders unequivocal determinations of mineral potential within the Yucca Mountain Addition impossible. However, we believe that the geochemical, lithologic, and mineralogic data collected from the surface of the Yucca Mountain Addition, in conjunction with remote sensing data, preclude the presence of surface or shallow deposits of precious metals mineable under current economic conditions, and utilizing presently available technology.

The highest gold content of samples collected during this study is 0.023 ppm, or 0.0007 troy ounces per short ton (\$0.28 per short ton at \$400 per ounce). Silver contents are similarly low, with the highest value at about 0.5 ppm (about 0.015 troy ounces per short ton, or \$0.15 per short ton at \$10 per ounce).

Gold and silver deposits that are currently producing, have been exploited in the past, or have established future potential do occur in rocks correlative with Yucca Mountain area units to the west, northwest, and east of the Yucca Mountain Addition within a 35 km radius. However, samples of silica-carbonate veins and breccias collected from in or around these deposits have

different mineralogies from carbonate, silica, and silicacarbonate veins or breccias collected from the Yucca Mountain Addition. In addition, rock samples from within or near the precious-metal deposits yielded gold, silver, and pathfinder element values at much higher over-all levels than samples from the Yucca Mountain Addition.

## INDUSTRIAL MINERALS AND MATERIALS

There are no identified industrial minerals and materials resources within the Yucca Mountain Addition (this includes salable, stakable, and leasable solid or fluid minerals). Within the restricted area of the proposed withdrawal, there are only two general geologic environments with possibilities for the occurrence-of industrial minerals and materials.

Gravel-covered pediments have potential for sand and gravel deposits. Most sand and gravel produced in Nevada goes to the highway construction industry for portland and bituminous concrete aggregate, base, or fill material, and to the building industry for construction aggregate. As in the past, sand and gravel operations in Nevada will continue to be developed as close to consuming areas as possible. Because of their low unit value, sand and gravel deposits will not permit much transportation of any kind. Sand and gravel deposits, while possibly present within the Yucca Mountain Addition, do not have any unique value over similar material occurring in other areas in southern Nevada and their potential for development is rated very low.

Rhyolitic volcanic flows and tuffs are potential hosts for a number of industrial minerals including zeolites, montmorillonitic clays, perlite, and pumice. Deposits of perlite or lightweight aggregate are unlikely because specifically favorable rock types have not been identified. Because of the remoteness of the area from population centers, only high-value commodities would have much probability of economic production. This would, in effect, limit possible commodities to zeolites and montmorillonite. Both of these high-value materials are commercially produced in Nevada. Within the area proposed for withdrawal, however, both montmorillonitic clays and zeolites are present only in minor amounts as alteration and secondary minerals at the surface or near the surface. Drill hole data indicate that zeolites of unknown quality occur about 600 m below the crest of Yucca Mountain in the Yucca Mountain Addition, and smectite clay occurs in significant amounts below 1,100 m. Because the total relief within the Yucca Mountain Addition is less than 400 m, it is highly unlikely that either of these materials is economically exploitable.

## ENERGY RESOURCES

#### Oil and gas Resourees

Nevada's petroleum potential can be predicted in a very general fashion on the basis of known production, shows of oil and gas, and proximity to areas of potential source and reservoir rocks. Areas of medium to high potential are located in the eastern part of the state, where most source rocks are found, and where these rocks have not been heated beyond the petroleum generation "window" to temperatures at which hydrocarbons have<br>been destroyed. Although the Yucca Mountain Addition is in an Although the Yucca Mountain Addition is in an area that has prospective potential for oil and gas (Smith and Gere, 1960), it lies within areas considered to have low oil and gas potential according to the most thorough regional studies available (Sandberg, 1983; and Garside et al., 1988). Therefore, the Yucca Mountain Addition is considered to have low potential for oil and gas resources.

Mississipian shales, which are thought to be the source rocks for most of the producing oilfields in Nevada (Poole and Claypool, 1984), are exposed 10 km west and 10 km or more northeast of the Yucca Mountain Addition. Mississipian rocks about 30 km northeast of the Yucca Mountain Addition have marginal source rock geochemistry (Center for Neotectonic Studies, 1989). These rocks yielded conodont color-alteration index (CAI) values of 1.5 to 2 (Harris et al., 1979; and S. P. Nitchman, personal communication, 1989), which are considered to be within the oil generation window (Poole and Claypool, 1984). Samples of Ordovician to Mississippian rocks within 15 km of the Yucca Mountain Addition to the east, northeast, and northwest, including borehole samples from depths of 1,300 m to 1,800 m, have yielded CAI values ranging between 3 and 6 (Harris et al., 1979; Carr et al.; 1986b, and Center for Neotectonic Studies, 1989). CAI values greater than 2 indicate temperatures above the oil-generation window, and values of 3 or higher indicate temperatures above the limit of oil preservation (Poole and Claypool, 1984). Nitchman (personal communication, 1989), believes that Mississippian source rocks on Bare Mountain yield high CAI values due to a heating episode related to overthrusting, and speculates that mature Mississippian source rocks comprise part of a parautochthonous upper Paleozoic sequence beneath the volcanic section in the Yucca Mountain area. Chamberlain (1989) believes that such rocks may have had the same depositional and thermal histories as Mississippian rocks in the productive Railroad Valley and Pine Valley oilfields about 150 km to the northeast of the Yucca Mountain Addition.

The geologic history of the Yucca Mountain Addition is largely unfavorable for the preservation of large hydrocarbon accumulations that may have been generated from Paleozoic source rocks during the Mesozoic (as postulated by Chamberlain). Extensive Miocene calderas within 5 km of the Yucca Mountain Addition were probably related to large magma chambers that would have created subsurface temperatures high enough to have destroyed any large oil accumulations that may have existed in the area.

This is supported by the CAI data for Paleozoic rocks surrounding the Yucca Mountain Addition which indicates that they were subjected to temperatures between 140'C (Carr et al., 1986b) and more than 300'C (Epstein et al., 1977).

In the absence of deep drill data beneath Yucca Mountain to test speculations regarding the presence of possible source rocks, estimates of hydrocarbon potential in the Yucca Mountain Addition must be based on surface and shallow drill data. Available data indicate hydrocarbon favorability is low, but not nil.

#### Geothermal Resources

Geothermal resources in Nevada are widespread and varied; the state has about 900 reported thermal springs and wells (Garside and Schilling, 1979). The higher-temperature resources are concentrated in northern Nevada, but many areas of the state have potential for low- to moderate-temperature resources. Data from thermal springs, water wells, and geothermal exploration wells listed in Garside and Schilling (1979), and Trexler and others (1983) have been used to define areas of the state that have potential for geothermal resources.

Warm springs in Oasis Valley about 20 km west of the Yucca Mountain Addition have surface water temperatures as high as 43 C (Garside and Schilling, 1979), and the Yucca Mountain Addition is within an area defined as prospectively valuable for geothermal resources (Godwin et al., 1967). However, based on temperatures measured in drill holes, thermal water (water with temperatures above those expected for a normal temperature gradient at the depth sampled) is not present in the Yucca Mountain Addition. Temperatures measured in drill holes below the static water level in the Yucca Mountain Addition range up to 54°C at a depth of 1,200 m (Sass et al., 1988).

Geothermal exploration based on temperature gradients in shallow drill holes is often considered sufficient to indicate the presence of a geothermal area, with gradients over most economically attractive areas in excess of 7 C per 100 m (Combs and Muffler, 1973). Temperature profiles from drill holes in the Yucca Mountain area range up to 6C per 100 m, and average about 3C (Sass et al., 1988). The average conductive heat flow in the Yucca Mountain area is anomalously low with respect to the regional heat flow, probably due to groundwater flow beneath the depth of exploration (Sass et al., 1988).

Based on the above information, the Yucca Mountain Addition does not appear to have potential for the discovery of geothermal resources.

## Uranium Potential

Low-grade uranium deposits occurring in volcanic rocks are associated with rhyolite intrusions or occur in the ring fracture zone or moat areas of calderas. Such features are not known to occur in the Yucca Mountain Addition. Radioactive rock was not found during radiometric analysis of drill holes, and fracture fillings in drill hole USW G-3 contain very low uranium contents

(up to 35 ppm) in comparison to ore from economic uranium deposits<br>which contains several hundred ppm or more uranium. The which contains several hundred ppm or more uranium. potential for surface or shallow economic uranium deposits is low in the Yucca Mountain Addition.

## SUMMARY OF MINERAL POTENTIAL

There are no identified mineral resources within the Yucca Mountain Addition. The potential for mineral deposits for which rock types underlying the area are favorable is rated very low.

The structural preparation of the Yucca Mountain Addition for mineralization, when compared with that for nearby mining districts, was rated as low from an analysis of the orientations of faults exposed at the surface. In the Yucca Mountain Addition, alteration is confined to two thin glassy bedded tuffs, whereas in the mining districts it crosses formation boundaries. Shallow intrusive rocks and pre-Tertiary rocks are exposed in the mining districts, but not in the Yucca Mountain Addition.

Base metal deposits that could be present in the deep subsurface (at depths of 1,500 meters or more) would not be mineable using standard, accepted mining technology. Surface exposures have very low mercury contents, indicating that economic deposits of this element, which are found in volcanic rocks, are not present.

Although the southwestern Nevada volcanic field does host economic near-surface disseminated gold-silver deposits, our work indicates that such deposits are not present within the Yucca Mountain Addition. Potential for precious-metal deposits mineable under current economic conditions, using presently available technology, is very low.

Potential for economic deposits of industrial minerals is also very low. Economic deposits of zeolites or clay are considered unlikely, although both have been encountered by drilling. Sand and gravel deposits present within the Yucca Mountain Addition have no unique value.

Based on presently available data, the Yucca Mountain Addition has low potential for energy resources including oil and gas, geothermal power, and uranium.

Determinations of potential for deeply buried mineral deposits under the Yucca Mountain Addition which could be mineable under future economic conditions, or using new technologies, are not within the scope of this investigation.

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# APPENDIX A. SAMPLE DESCRIPTIONS FROM FIELD NOTES, YUCCA MOUNTAIN ADDITION AND SURROUNDING MINING DISTRICTS



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### APPENDIX B. CHEMICAL ANALYSES, IN PPM, OF SAMPLES FROM THE YUCCA MOUNTAIN ADDITION AND SURROUNDING MINING DISTRICTS

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\* reanalysis of original sample

t analysis of resampled material

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\*\* reanalysis of hand sample

 $8 - g$ 



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\*\* reanalysis of hand sample

 $6 - B$ 

### APPENDIX C. MINERALOGIC RESULTS FROM X-RAY DIFFRACTION, YUCCA MOUNTAIN ADDITION AND SURROUNDING MINING DISTRICTS

#### YUCCA MOUNTAIN ADDITION



\* A = alunite, B = biotite, Ca = calcite, Cl = clinoptilolite, Cr = cristobalite,  $H =$  hematite,  $I =$  illite,  $I-M =$  interstratified illite-montmorillonite,  $J = j$ arosite, KF = potash feldspar, M-I = montmorillonite-illite,  $M =$  montmorillonite,  $0 =$  opal,  $0$ -CT = opal-CT, PF = plagioclase feldspar,  $Q =$  quartz,  $T =$  tridymite, and VG = volcanic glass.

 $C-1$ 

# YUCCA MOUNTAIN ADDITION (CONT.)

-



 $C-2$ 

## WAHMONIE MINING DISTRICT



MOTHER LODE DEPOSIT AREA



## ORIGINAL BULLFROG MINE



 $C-3$ 

RHYOLITE AREA



\* A = alunite,  $3$  = biotite,  $Ca =$  calcite,  $Cl =$  clinoptilolite, Cr = cristobalite,  $H =$  hematite,  $I =$  illite,  $I-M =$  interstratified illite-montmorillonite, J = jarosite, KF = potash feldspar, M-I montmorillonite-illite, M = montmorillonite, 0 = opal, 0-CT = opal-CT, PF = plagioclase feldspar,  $Q =$  quartz, T = tridymite, and VG = volcanic glass.