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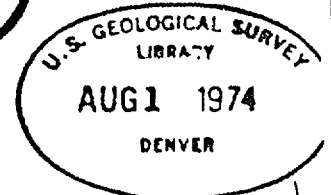
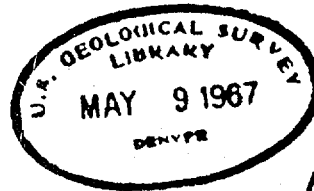
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DETERMINATION OF THE VERTICAL AND HORIZONTAL PERMEABILITIES OF FRACTURED WATER BEARING FORMATION

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Résumé

On peut considérer les formations aquifères fracturées comme des aquifères anisotropes. Lorsqu'il s'agit de fractures horizontales, la perméabilité horizontale correspond aux propriétés de transmission des fractures et la perméabilité verticale est celle de la roche mère. Au contraire, lorsqu'il s'agit de fractures verticales, la perméabilité verticale correspond aux propriétés de transmission des fractures et la perméabilité horizontale caractérise la roche mère. Autrement dit, on considère que la roche est isotrope. Le présent travail vise à déterminer les perméabilités verticale et horizontale des roches fracturées, aussi bien que leur capacité de réserve, par l'étude des données des essais de pompage procédé employé est fondé sur la méthode de la double pente (Saad *et al.*, 1964). De plus, le rapport des perméabilités verticale et horizontale peut fournir des indices sur les réseaux de fractures et il connaît la facilité de l'écoulement dans chaque direction.

ABSTRACT

Fractured water bearing formation may be considered as an anisotropic aquifer. That is the permeability through the fractures will be different from that of the original parent formation. In the present paper, the values of both the vertical and horizontal permeabilities as well as the storage coefficient have been determined through analysis of the pumping test data. The permeability in one of these two directions represents that of the fractures having the same direction, while the other characterizes the transmission property of the parent formation. The procedure of analysis is based on the double slope method (Saad *et al.* (1965) through analysis of the modified solution of the nonsteady flow toward a well partly penetrating the fractured water bearing formation. It can also be concluded that knowledge of the magnitude of the permeabilities in both directions may indicate the pattern and trend of the fractures.

INTRODUCTION

Anisotropic permeability of water bearing formation is a result of many reasons. Among them are the presence of fractures, with a certain pattern, in previous aquifer such as limestone. Consequently the permeability through the fractures is different and usually higher from that of the parent formation. The fractures may have either a horizontal or a vertical trend (Fig. 1) and thus they are represented hydraulically by the horizontal or the vertical permeability respectively. The permeability in the other direction represents the transmitting property of the parent water bearing formation.

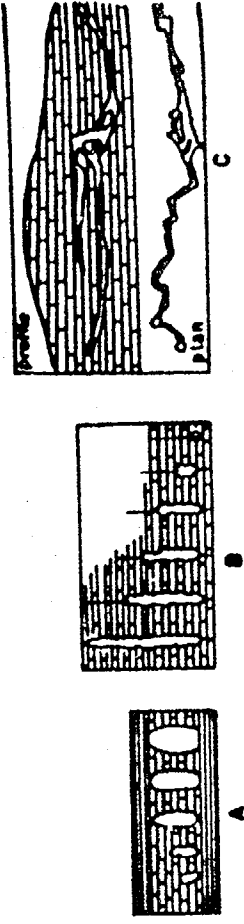


Fig. 1 — Cross section of fractured rock. A and B: Vertical fracture. C: Horizontal fracture.

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In practice, it frequently happens that the producing wells do not penetrate completely the fractured formation from which they are pumping. This is rather due to many technical reasons, among these are the large thickness of the fractured bed or the wide fracture openings. In both cases, drilling operation may be either expensive or impractical for large depths for excessive loss of mud circulation and other difficulties. For this, producing wells drilled in fractured water bearing formation, are usually partially penetrating the equifer.

The purpose of the present paper is to determine the permeability of both the fractures and the parent formation. These parameters as well as the storage coefficient can be determined through analysis of the pumping test data, records from a partially penetrating observation well, where the pumped well itself does not reach the bottom of the fractured formation. It is assumed that the fractures are either horizontal or vertical and that the storage coefficient remains constant in the whole region.

The procedure of analysis is based on the double slope method, Saad *et al.*, (1965), and on the modified solution of the nonsteady flow toward a partially penetrating well, Hantush, 1961, to account for anisotropy, Muskat, 1937.

The average drawdown (s) in an observation well, where the screens in both the pumped and observation wells extend for the whole length of each well, and that both are partially penetrating a water bearing formation, Figure 2, has been found by Hantush, (1961) as follows:

$$s = \frac{Q}{4\pi K_f b} [W(u) + (4b^2/\pi^2 d d') \sum_{n=1}^{\infty} (1/n^2) K_0(n\pi r/b) (\sin n\pi d/b) (\sin n\pi d'/b)] \quad (1)$$

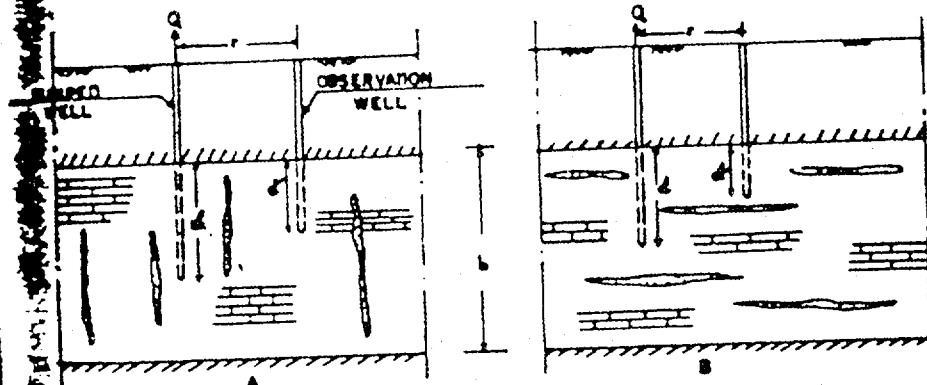


Fig. 2 — Diagrammatic representation of a well tapping fractured rock. A: Vertical fracture. B: Horizontal fracture.

Equation (1), can be modified, to account for anisotropic permeability resulting from the presence of fractures, by multiplying the term (r/b) by $(K_d/K_f)^{1/2}$, Muskat, (1937). Thus Equation (1), will reduce to:

$$s = \frac{Q}{4\pi K_f b} [W(u) + (4b^2/\pi^2 d d') \sum_{n=1}^{\infty} 1/n^2 K_0\{(n\pi r/b)(K_d/K_f)^{1/2}\} (\sin n\pi d/b) (\sin n\pi d'/b)] \quad (2)$$

Equation (2) shows that the rate of changes of the average drawdown behaves as of radial toward a well fully penetrating an aquifer.

For the determination of the hydraulic coefficients K_r and K_v , the double slope. Saad et al. (1965), can be used. The procedure of the mathematical analysis is outlined as follows:

i) differentiating (2) in Equation (2) with respect to $\log r$

$$ds/d(\ln r) = \frac{2.30 Q}{4\pi K_r b} e^{-u} = m \quad (1)$$

ii) differentiating (3) in Equation (3) with respect to $\ln r$

$$dm/d(\ln r) = \frac{(2.30)^2 Q}{4\pi K_r b} u e^{-u} = m' \quad (2)$$

iii) the double slope function $f(u) = m'/m$, can be found

$$f(u) = m'/m = 2.30u \quad (3)$$

Equations (5), (3), and the relationship $(u = r^2 S/4 K_r b t)$ and the data of pumping test, enable determining K_r and S , as will be shown later. The value of (K_v/K_r) , can also be found in terms in the summation form appearing in Equation (2), can be evaluated. For this, Equation (2) can be put in the following forms:

$$A = \sum_{n=1}^{\infty} \frac{(1/n^2) K_v \{(\pi r/b)(K_r/K_v)\}^{1/2}}{(\sin \pi x d/b)(\sin \pi x d'/b)} \quad (4)$$

Where

$$A = \left[\frac{s}{Q/4\pi K_r b} - W(u) \right] (\pi^2 d d'/4b^2) \quad (5)$$

Equation (6), can be further reduced to the form of Fourier series by multiplying both sides by $(\sin x) dx$, where $(x = \pi d'/b)$, and integrating between the limits 0 and π

$$\int_0^{\pi} A \sin(x) dx = \int_0^{\pi} \sum_{n=1}^{\infty} \frac{(1/n^2) K_v \{(\pi r/b)(K_r/K_v)\}^{1/2}}{(\sin \pi x d/b)(\sin \pi x d'/b)} (\sin \pi x) (\sin x) dx \quad (6)$$

Using the following two identities:

$$\int \sin(\pi x) \cdot \sin(m\pi x) dx = \frac{\sin(\pi - m)\pi x}{2(\pi - m)} - \frac{\sin(\pi + m)\pi x}{2(\pi + m)} \quad (7)$$

and

$$\int (\sin(\pi x))^2 dx = (x/2) - \frac{\sin(2\pi x)}{4\pi} \quad (8)$$

it can be shown easily that all the terms in the summation form in Equation (7) will tend to zero except when $n = 1$. Therefore Equation (7) will reduce finally to:

$$\int_0^{\pi} A \sin(x) dx = (\pi/2) K_v \{(\pi r/b)(K_r/K_v)\}^{1/2} \sin(\pi d/b) = 2A \quad (9)$$

- 1. Draw the curve).
- 2. Choose several points on the curve. This may be more accurate or in other words when plotting on the same semi-log plot.
- 3. Select few points on the $(m-1)$ curve, point, per cycle.
- 4. Knowing (m) and (m') at each point, find the correlation coefficient (r) .
- 5. Using Equation (3) with known (r) , find (K_v/K_r) .
- 6. Steps from v to vii, may be repeated for several values of (r) . However the computed value of (K_v/K_r) is a result of improper correlation.
- 7. Choose any point on the $(s-1)$ curve.
- 8. Compute the value of (s) at that point from tables of the well function $W(u)$.
- 9. Using Equation (11), the value of (K_r) , can be found.
- 10. Knowing (K_r) , determine the aquifer thickness (L) .
- 11. The depth of penetration of the zero-order modified Bessel function is $(2.405 L)$ or Dwight (1958).
- 12. The horizontal permeability ratio (K_v/K_r) .
- 13. The vertical permeability $(L^2 T)$.
- 14. The slope of the tangent at any point on the curve.
- 15. The slope of the tangent at any point on the curve.
- 16. The constant well discharge (Q) .
- 17. The distance between the pump and the observation point.
- 18. The average drawdown in an unconfined aquifer.
- 19. The storage coefficient or storage (S) .
- 20. The time since pumping started (t) .
- 21. The relation $r^2 S/4 K_r b t$.
- 22. The well function $W(u) = \int_0^{\infty} (e^{-u^2}/u) du$.

$$K_0[(\pi r/b)(K_1/K_2)^{1/2}] = (4A/\pi) \operatorname{cosec}(\pi d/b) \quad (11)$$

APPLICATION

For determining the hydraulic properties of both the fractured and parent formation, i.e., vertical and horizontal permeability and the storage coefficient from the data of the pumping test, the following procedure can be applied:

- i) Draw the drawdown-time curve on semi-log paper, with the time on the log scale ($S - \log t$ curve).
- ii) Choose several points on the curve and measure the slope (m), per cycle at each point. It may be more accurate if the chosen points comprise the latest portion of the original curve, or in other words when the time is large.
- iii) Plot on the same semi-log paper the measured slopes (m) versus $\log t$.
- iv) Select few points on the ($m - \log t$) curve and measure the slope (m') of the tangent at each point, per cycle.
- v) Knowing (m) and (m') at each time, find the double slope function $f(u)$, using Equation (5) at each point, and find the corresponding value of (u).
- vi) Using Equation (3) with known values of (u), (Q), (b) and (m), determine the value of K_1 .
- vii) Using the relationship ($u = r_s S/4 K_1 b t$), compute the value of S .
- viii) Steps from v to vii, may be repeated for other values of (m'), at different values of time. However the computed values of (K_1) and (S), should be the same at each point. Any deviation is a result of improper measurements of (m) or (m').
- ix) Choose any point on the ($s - \log t$) curve and record its coordinates.
- x) Compute the value of (u) at that particular point and find its corresponding value of $W(u)$, from tables of the well function, Wenzel, (1942).
- xi) Using Equation (11), the value of $K_0[(r/b)(K_1/K_2)^{1/2}]$, can be calculated. From which the value of (K_2/K_1), can be found, using the tables of the modified Bessel function $K_0(x)$.
- xii) Knowing (K_1), determine the value of (K_2).

APPENDIX - NOTATION

- L aquifer thickness (L).
- l depth of penetration of the pumped well (L).
- l' depth of penetration of the observation well (L).
- $K_0(x)$ the zero-order modified Bessel function of the second kind, tabulated, Watson (1944) or Dwight (1958).
- K_1 horizontal permeability radially from the well (L/T).
- K_2 vertical permeability (L^2/T).
- m slope of the tangent at any point on the drawdown-time curve ($s - \log t$) (L/cycle).
- m' slope of the tangent at any point on the drawdown-slope-time curve ($m - \log t$) (L/cycle).
- Q constant well discharge (L^3/T).
- r distance between the pumped and the observation wells (L).
- \bar{s} average drawdown in an observation well partially penetrating the aquifer (L).
- S storage coefficient or storativity.
- t time since pumping started (T).
- r_s relation $r_s S/4 K_1 b t$.
- $W(u) = \int_0^\infty (e^{-x^2/4u}) dx$ - well function of (u), tabulated by Wenzel (1942).

L'HYDRO DANS LES

Laboratoire d'Hy

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REMARQUES

Dans les dunes modernes du littoral méditerranéen, on trouve dans la partie inférieure, des réserves d'eau qui sont en fait des réserves d'eau. Elles sont en fait des réserves d'eau. Elles sont en fait des réserves d'eau. Elles sont en fait des réserves d'eau.

L'INTRODUCTION

Objet de l'étude

Dans les bas-fonds placés immédiatement proche de la surface, on trouve d'autres zones non encore explorées. Elles sont en fait des réserves d'eau. Elles sont en fait des réserves d'eau. Elles sont en fait des réserves d'eau.

Conclusions

L'ensemble des mesures et des sondages effectués dans le long du corridor de Força à 10 km au Nord du littoral méditerranéen, ont permis de constater que les réserves d'eau sont en fait des réserves d'eau. Elles sont en fait des réserves d'eau. Elles sont en fait des réserves d'eau.

Conditions écologiques générales

Les dunes sont de formation très récente. Elles sont en fait des réserves d'eau. Elles sont en fait des réserves d'eau. Elles sont en fait des réserves d'eau.

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