# Recent Advances in PTHA Methodology

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> NRC Workshop on Probabilistic Flood Hazard Assessment January 29–31, 2013

# **Crescent City PTHA Pilot Study**



Preliminary results will be shown from ongoing pilot study.

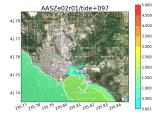
Supported by BakerAECOM, as part of a coastal modeling/mapping effort funded by the FEMA Region IX office as part of the new California Coastal Analysis and Mapping Project (CCAMP).

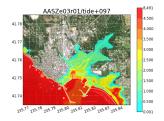
Simulations done with GeoClaw model (shallow water equations) www.clawpack.org/geoclaw

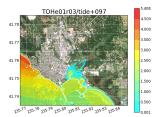
# Crescent City, CA

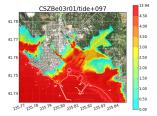


### Four sample event realizations









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- Cumulative probability distribution of tide stage.

#### Some recent advances

- Improved methodology for tidal uncertainty.
- Approaches to mapping probabilities in addition to depth.
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#### Some limitations:

- Proper probability density for slip distribution is hard to determine geophysics problem.
- How to add in the possibility of submarine landslides affecting tsunami size?
- Many other uncertainties, e.g. friction coefficient and other aspects of mathematical model / numerical method.

First define hazard curve: for each location (x, y):

 $P(\zeta) = P(\zeta; x, y) =$ Prob[inundation  $\geq \zeta$  in one year].

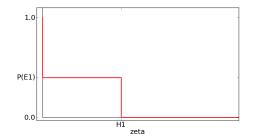
**Example**: If only one possible event  $E_1$  with recurrence time  $T_1$  (Poisson rate  $\nu_1 = 1/T_1$ ), that floods to level  $H_1(x, y)$ , then

$$P(\zeta) = \begin{cases} 1 & \text{if } \zeta = 0, \\ 1 - e^{-\nu_1} & \text{if } 0 < \zeta < H_1, \\ 0 & \text{if } \zeta > H_1. \end{cases}$$

### Hazard Curves

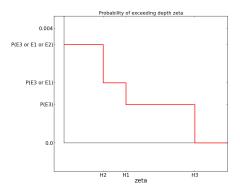
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## Hazard Curves

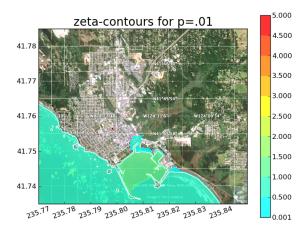
**Example:** Three possible events  $E_1$ ,  $E_2$ ,  $E_3$  with recurrence times  $T_1$ ,  $T_2$ ,  $T_3$ , that flood to levels  $H_1$ ,  $H_2$ ,  $H_3$ .



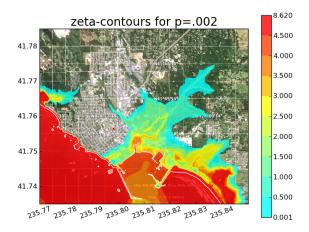
Where, for example,

 $P(E3 \text{ or } E1) = 1 - (1 - p_3)(1 - p_1) = 1 - e^{-(\nu_3 + \nu_1)}.$ 

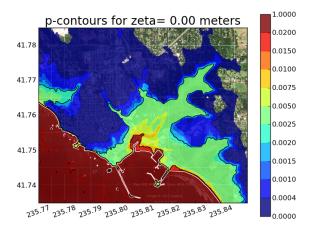
Standard view: Map of flooding depth for fixed probability



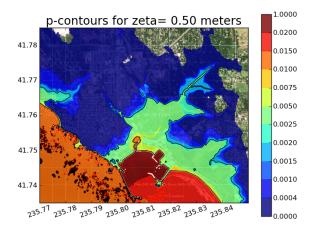
Standard view: Map of flooding depth for fixed probability



Alternative map: Probability of exceeding fixed depth

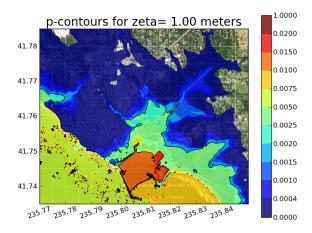


Alternative map: Probability of exceeding fixed depth

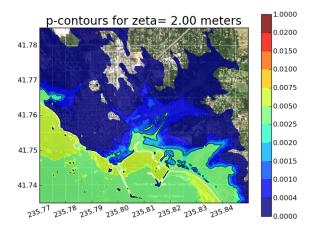


Preliminary Results Not for Use

Alternative map: Probability of exceeding fixed depth



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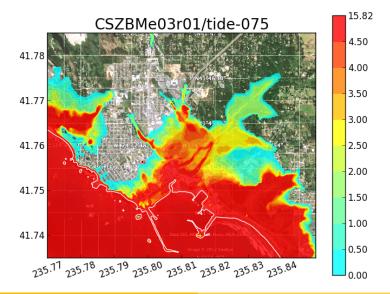
Preliminary Results Not for Use Bathymetry uses vertical datum MHW (Mean High Water).

Tidal range:

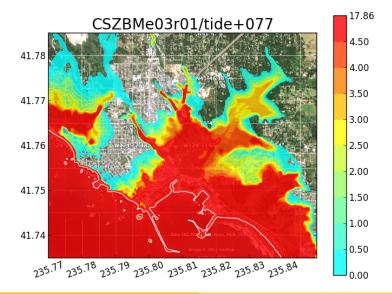
 $\begin{array}{l} \mbox{MHHW} \approx \mbox{MSL} \ + \ 97 \ \mbox{cm.} \\ \mbox{MHW} \ \approx \mbox{MSL} \ + \ 77 \ \mbox{cm.} \\ \mbox{MLW} \ \approx \mbox{MSL} \ - \ 75 \ \mbox{cm.} \\ \mbox{MLW} \ \approx \mbox{MSL} \ - \ 113 \ \mbox{cm.} \end{array}$ 

How much difference does modifying tide stage have on inundation?

## Sample inundation from CSZ M9 earthquake



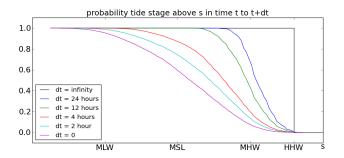
## Sample inundation from CSZ M9 earthquake



# **Tidal uncertainty**

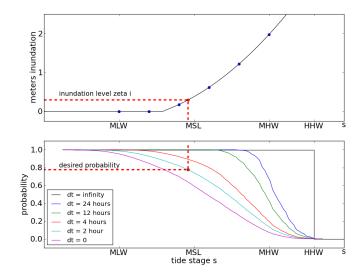
Suppose we determine for a given event that largest waves are all seen within a period of  $\Delta t$  hours after arrival of first wave.

Then we can use tide record to determine the cumulative probability that the tide stage will be above *s* at *some* time between *t* and  $t + \Delta t$ . (Where *t* is assumed to be a random time in tide cycle when first wave hits.)

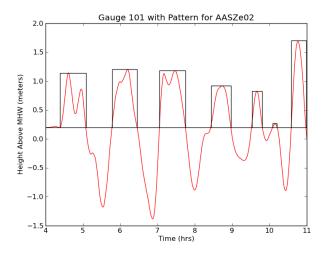


## **Tidal uncertainty**

For example, if  $\Delta t = 2$  hours is appropriate for this event:

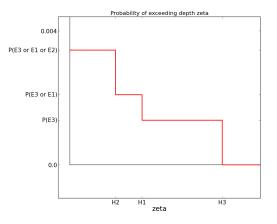


# The pattern method



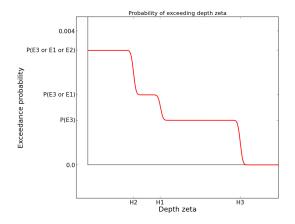
### Effect of tidal uncertainty on hazard curves

Example: Three events with no tidal uncertainty:

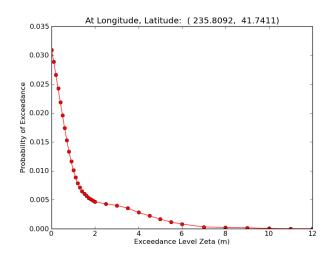


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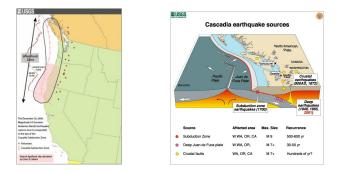


## Hazard curve with many events + tides



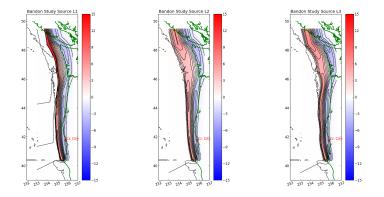
Preliminary Results — Not for Use

# Cascadia Subduction Zone (CSZ)



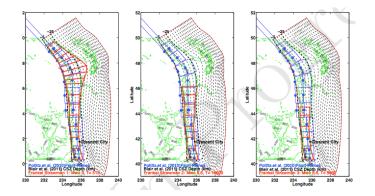
- 1200 km long off-shore fault stretching from northern California to southern Canada.
- Last major rupture: magnitude 9.0 earthquake on January 26, 1700.
- Tsunami recorded in Japan with run-up of up to 5 meters.
- Historically there appear to be magnitude 8 or larger quakes every 500 years on average.

# Sample seafloor deformations



R. C. Witter and Y. Zhang and K. Wang and G. R. Priest and C. Goldfinger and L. L. Stimely and J. T. English and P. A. Ferro, *Simulating tsunami inundation at Bandon, Coos County, Oregon, using hypothetical Cascadia and Alaska earthquake scenarios*, Oregon Department of Geology and Mineral Industries Special Paper 43, 2011. (15 realizations)

# CSZ fault geometry and possible slip regions



3 Possible slip regions (Art Frankel, USGS) and fault geometry of

Pollitz, McCrory, Wilson, Svarc, Puskas, Smith, http://doi.wiley.com/10.1111/j.1365-246X.2010.04546.x

and Blair, McCrory, Oppenheimer, http://pubs.usgs.gov/ds/633/

For a given event (e.g. CSZ  $M_w$  9.1) we need to use more than one possible slip distribution.

To simplify, consider an event with a single fault plane for which we know strike, dip, rake, and the integral of slip (from  $M_w$ ).

Assuming uniform slip gives one possible realization but not realistic.

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Can split plane into m pieces and distribute slip.

Let  $d_i = \text{slip on } i\text{th subfault, for } i = 1, 2, \ldots, m$ .

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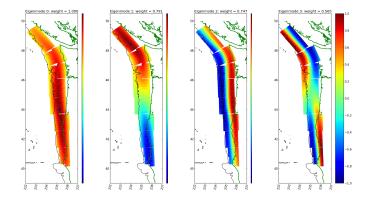
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#### How to generate many realistic slips $d_i$ ?

# Karhunen-Loève Modes of CSZ



Eigenfunctions of Covariance assuming correlation lengths of 40% along-strike and down-dip.

P. M. Mai and G. C. Beroza, *A spatial random field model to characterize complexity in earthquake slip*, J. Geophys. Res. 107(2002).

R. J. LeVeque, University of Washington NRC PFHA Workshop, January, 2013

Let  $(\lambda_k, v_k)$  be eigenvalue/vectors of C, for k = 1, 2, ..., m. Order so that  $\lambda_1 \ge \lambda_2 \ge \cdots \ge \lambda_m$ .

Eigenvectors for larger k are more oscillatory.

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Then realizations  $d = [d_1, \ldots, d_m]$  can be generated by

$$d = \mu + \sum_{k=1}^{m} z_k \sqrt{\lambda_k} v_k$$

where the  $z_k$  are chosen to be independent from N(0,1) (normally distributed with mean 0 and variance 1).

# Karhunen-Loève expansion

1

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#### Advantages:

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Moreover, Once we apply Okada model to *d*, oscillatory components get further smoothed.

Even lower dimensional space of seafloor deformations?

# Uncertainty quantification techniques

Why is dimension reduction important?

May be able to do better than brute-force Monte-Carlo by doing smaller number of simulations at carefully chosen points in stochastic space, represented by vectors

 $[z_1, z_2, \ldots, z_r]$  (r =number of stochastic dimensions)

and building up polynomial approximation to response surface (stochastic collocation).

Then use known probability distribution of the  $z_k$  to compute probability of response.

Hard to do in large number of dimensions, but can explore techniques such as sparse grids, importance sampling, MCMC, multi-level MC.

#### Aleatoric uncertainties: (Mathematical problem)

Given the correct probability density of slip distribution, there are techniques to reduce dimension and sample efficiently.

Epistemic uncertainties: (Geophysical problem)

- What is correct distribution/recurrence of different magnitude quakes?
- What is correct fault geometry?
- Is there tapering of slip?
- Is K-L expansion with Gaussian weights correct?
- What is correct correlation function?

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